Ground water modeling

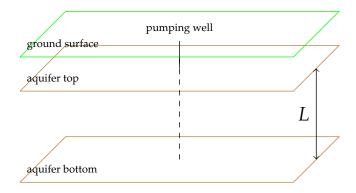
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Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 0/28

Thiem's law - Assumptions

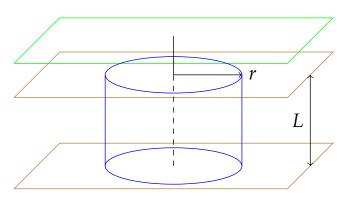
confined aquifer



- **⊙** homogeneous, with *hydraulic conductivity K*,
- **⊙** constant thickness, *L*,
- Θ IC: immobile water, $h = h_0 = \text{const}$,
- with a *fully penetrating* pumping well,
- Θ extracting a constant water flux Q,

Thiem's law - Cylindrical symmetry

confined aquifer



⇒ cylindrical symmetry:

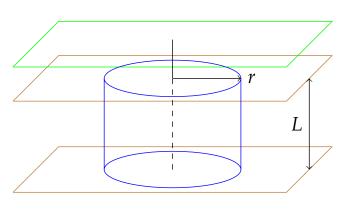
the hydraulic head h under stationary conditions is function of r only, the distance of a point from the pumping well.

$$h = h(r)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 2/28

Thiem's law - Water through the cylinder with radius r

confined aquifer



$$Q = -L 2\pi r q_r = L 2\pi r K \frac{\mathrm{d}h}{\mathrm{d}r}$$

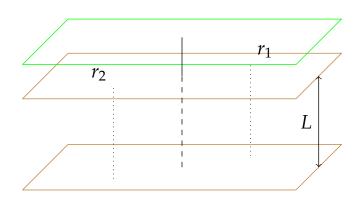
(Darcy's law)

$$\frac{\mathrm{d}h}{\mathrm{d}r} = \frac{Q}{2\pi r \, KL}$$

$$h(r) = \frac{Q}{2\pi KL} \ln(r) + c$$

(integrating...)

Thiem's law (try to avoid the constant c)

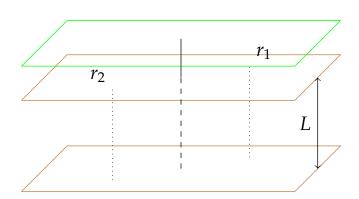


$$h(r_1) = \frac{Q}{2\pi KL} \ln(r_1) + c; \qquad h(r_2) = \frac{Q}{2\pi KL} \ln(r_2) + c$$

$$\Rightarrow h(r_2) - h(r_1) = \frac{Q}{2\pi KL} \ln(r_2) - \frac{Q}{2\pi KL} \ln(r_1) = \frac{Q}{2\pi KL} \ln\left(\frac{r_2}{r_1}\right)$$

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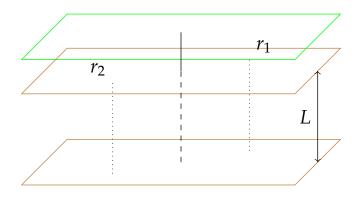
Thiem's law, in terms of drawdown



In terms of $drawdown \ s(r) = h(r) - h_0$ caused by the pumping well: $(h_2 - b)^2 - (h_1 - h_0)^2 = dfracQ2\pi \ KL \ln\left(\frac{r_2}{r_1}\right) \Rightarrow$

$$\Rightarrow$$
 $KL = T = \frac{Q}{2\pi(s_2 - s_1)} \ln\left(\frac{r_2}{r_1}\right)$

Thiem's law - Final results



In terms of *drawdown* $s(r) = h(r) - h_0$ caused by the pumping well:

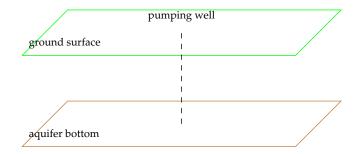
> Thiem's law (confined aquifer)

$$T = \frac{Q}{2\pi(s_2 - s_1)} \ln\left(\frac{r_2}{r_1}\right)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 6/28

Dupuit's law - Assumptions

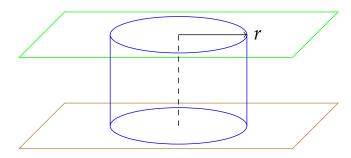
phreatic aquifer



- Dupuit's approximation,
- without recharge,
- homogeneous, with conductivity K,
- horizontal impermeable bottom at height b,
- IC: immobile water, $h = h_0 = \text{const}$,
- **•** with a *fully penetrating* pumping well,
- extracting a constant water flux Q,
- **⊙** same flux per unit length through the whole aquifer thickness.

Dupuit's law - cylindrical symmetry

phreatic aquifer



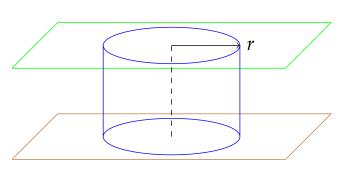
\Rightarrow cylindrical symmetry:

the hydraulic head h under stationary conditions is function of r only, the distance of a point from the pumping well.

$$h = h(r)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 8/28

Dupuit's law - water through the cylinder with radius r



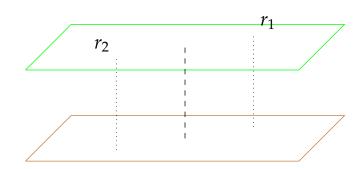
$$Q = -(h - b)2\pi r q_r$$

$$= (h - b)2\pi r K \frac{dh}{dr}$$

$$\Rightarrow \frac{d}{dr} \frac{(h - b)^2}{2} = \frac{Q}{2\pi K r} \Rightarrow$$

$$(h(r) - b)^2 = \frac{Q}{\pi K} \ln(r) + c$$

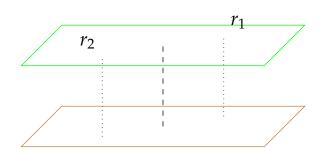
Dupuit's law - application



$$(h(r_1) - b)^2 = \frac{Q}{\pi K} \ln(r_1) + c; \qquad (h(r_2) - b)^2 = \frac{Q}{\pi K} \ln(r_2) + c \Rightarrow$$
$$(h(r_2) - b)^2 - (h(r_1) - b)^2 = \frac{Q}{\pi K} \ln(r_2) - \frac{Q}{\pi K} \ln(r_1) = \frac{Q}{\pi K} \ln\left(\frac{r_2}{r_1}\right)$$

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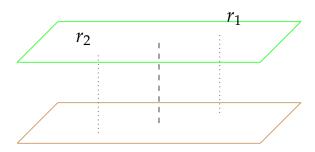
Dupuit's law - in practice



$$(h_2 - b)^2 - (h_1 - b)^2 = \frac{Q}{\pi K} \ln \left(\frac{r_2}{r_1}\right)$$

$$\Rightarrow K = \frac{Q}{\pi [(h_2 - b)^2 - (h_1 - b)^2]} \ln \left(\frac{r_2}{r_1}\right)$$

Dupuit's law - in practice, in terms of drawdown



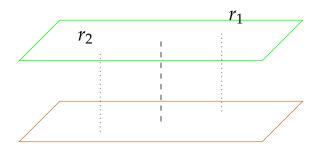
In terms of *drawdown*, $s(r) = h(r) - h_0$), and phratic aquifer thickness, L = h - b:

$$K = \frac{Q}{\pi[(h_2 - b)^2 - (h_1 - b)^2]} \ln\left(\frac{r_2}{r_1}\right)$$

$$\Rightarrow K = \frac{Q}{\pi[(s_2 - s_1)(L_2 + L_1)]} \ln\left(\frac{r_2}{r_1}\right)$$

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Dupuit's law - in practice

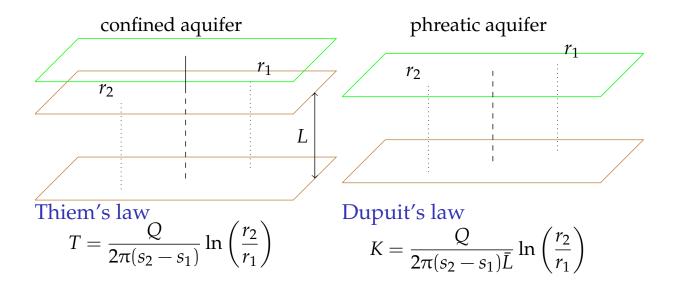


With $L_2 + L_1 = 2\bar{L}$, where \bar{L} is the average aquifer thickness

Dupuit's law (phreatic aquifer)

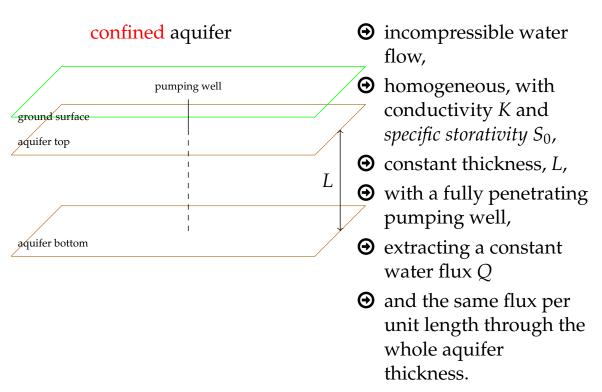
$$K = \frac{Q}{2\pi(s_2 - s_1)\bar{L}} \ln\left(\frac{r_2}{r_1}\right)$$

Thiem's vs. Dupuit's law

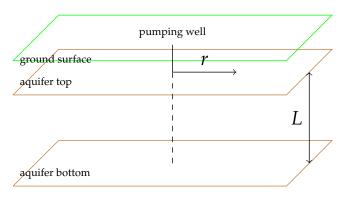


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Theis' law - Assumptions



Theis' law - assumptions, cylindrical symmetry



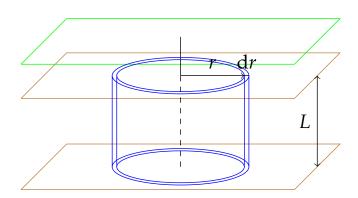
 \Rightarrow cylindrical symmetry:

the hydraulic head h is function of *r*, the distance of a point from the pumping well, and *t* only.

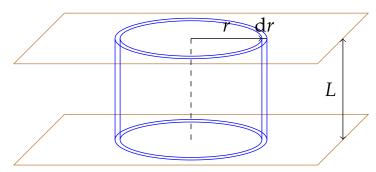
$$h=h(r,t)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 16/28

Theis' law - flux through cylinders with radius r and r + dr



$$Q(r) = -L 2\pi r q_r(r);$$
 $Q(r + dr) = -L 2\pi (r + dr) q_r(r + dr)$



Water balance in the volume between the two cylinders:

$$S_0 L2\pi r dr \frac{\partial h}{\partial t} = -L2\pi (r + dr)q_r(r + dr) + L2\pi r q_r(r) \Rightarrow$$

$$S_0 L2\pi r dr \frac{\partial h}{\partial t} = -L2\pi \left[(r + dr)q_r(r + dr) - r q_r(r) \right]$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 18/28

Theis' law

$$\begin{split} S_0 L 2\pi r \mathrm{d}r \frac{\partial h}{\partial t} &= -L 2\pi \left[(r + \mathrm{d}r) q_r (r + \mathrm{d}r) - r q_r (r) \right] \ \Rightarrow \\ S_0 L r \frac{\partial h}{\partial t} &= -L \frac{(r + \mathrm{d}r) q_r (r + \mathrm{d}r) - r q_r (r)}{\mathrm{d}r} \ \Rightarrow \\ S_0 L r \frac{\partial h}{\partial t} &= -L \frac{\partial}{\partial r} \left(r q_r (r) \right) \ \Rightarrow \\ S_0 L \frac{\partial h}{\partial t} &= K L \frac{1}{r} \frac{\partial}{\partial r} \left(r \frac{\partial h}{\partial r} \right) \ \Rightarrow \\ S \frac{\partial h}{\partial t} &= T \frac{1}{r} \frac{\partial}{\partial r} \left(r \frac{\partial h}{\partial r} \right). \end{split}$$

The problem is to solve the following equation

$$S\frac{\partial h}{\partial t} = T \left[\frac{\partial}{\partial x} \left(\frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left(\frac{\partial h}{\partial y} \right) \right],$$

with $h = h_0 + s$, where

- Θ $h_0(x, y, t)$ is the response to the boundary conditions and to the source terms, but the pumping well;
- \odot s(r,t) is the effect (drawdown) of the well.

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 20/28

Theis' law

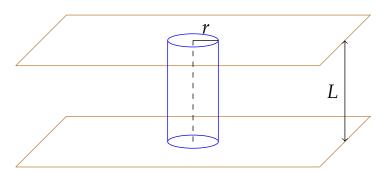
s(r, t) is solution of

$$S\frac{\partial s}{\partial t} = T\frac{1}{r}\frac{\partial}{\partial r}\left(r\frac{\partial s}{\partial r}\right),\,$$

with the following initial & boundary conditions:

$$\Theta$$
 $s(r,t) \to 0$ for $r \to +\infty$.

The last equation required to pose the problem is the boundary condition at the well... next slide.



Water balance for a cylinder with small radius *r* around the well:

$$S_0 L \pi r^2 \frac{\partial h}{\partial t} = -L2 \pi r q_r(r) - Q$$

(LHS negligile because r^2 smaller than r for $r \to 0$; Darcy's law; $h = h_0 + s$, but around the well contributions of h_0 negligible...)

$$\Rightarrow LK2\pi r \frac{\partial s}{\partial r} \rightarrow Q, \ r \rightarrow 0 \ \Rightarrow \lim_{r \rightarrow 0} r \frac{\partial s}{\partial r} = \frac{Q}{2\pi T}.$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 22/28

Theis' law

s(r, t) is solution of

$$\begin{cases} S\frac{\partial s}{\partial t} = T\frac{1}{r}\frac{\partial}{\partial r}\left(r\frac{\partial s}{\partial r}\right) \\ s(r,0) = 0 \\ \lim_{r \to +\infty} s(r,t) = 0 \\ \lim_{r \to 0} r\frac{\partial s}{\partial r} = \frac{Q}{2\pi T} \end{cases}$$

We search a solution which depends on r and t through

$$\xi = \frac{S}{4T} \cdot \frac{r^2}{t}.$$

We look for
$$s(r,t) = \tilde{s}(\xi(r,t))$$
, with $\xi = \frac{S}{4T} \cdot \frac{r^2}{t}$.

$$\frac{\partial s}{\partial t} = \frac{d\tilde{s}}{d\xi} \cdot \frac{\partial \xi}{\partial t} = -\frac{S}{4T} \cdot \frac{r^2}{t^2} \frac{d\tilde{s}}{d\xi} = -\frac{\xi}{t} \frac{d\tilde{s}}{d\xi};$$

$$\frac{\partial s}{\partial r} = \frac{d\tilde{s}}{d\xi} \cdot \frac{\partial \xi}{\partial r} = \frac{S}{4T} \cdot \frac{2r}{t} \frac{d\tilde{s}}{d\xi} = \frac{2\xi}{r} \frac{d\tilde{s}}{d\xi}.$$

$$r \frac{\partial s}{\partial r} = 2\xi \frac{d\tilde{s}}{d\xi},$$

$$S \frac{\partial s}{\partial t} = T \frac{1}{r} \frac{\partial}{\partial r} \left(r \frac{\partial s}{\partial r} \right) \Rightarrow -\frac{S\xi}{t} \frac{d\tilde{s}}{d\xi} = \frac{2T}{r} \frac{\partial}{\partial r} \left(\xi \frac{d\tilde{s}}{d\xi} \right)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 24/28

Theis' law

We look for
$$s(r,t) = \tilde{s}(\xi(r,t))$$
, with $\xi = \frac{S}{4T} \cdot \frac{r^2}{t}$.
$$r\frac{\partial s}{\partial r} = r\frac{2\xi}{r}\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = 2\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi},$$

$$\frac{1}{r}\frac{\partial}{\partial r}\left(r\frac{\partial s}{\partial r}\right) = \frac{2}{r}\frac{\partial}{\partial r}\left(\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}\right) = \frac{2}{r}\frac{\partial\xi}{\partial r}\cdot\frac{\mathrm{d}}{\mathrm{d}\xi}\left(\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}\right) = \frac{4\xi}{r^2}\cdot\frac{\mathrm{d}}{\mathrm{d}\xi}\left(\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}\right);$$

$$-\frac{S\xi}{t}\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = T\frac{4\xi}{r^2}\cdot\frac{\mathrm{d}}{\mathrm{d}\xi}\left(\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}\right) \Rightarrow -\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = \frac{\mathrm{d}}{\mathrm{d}\xi}\left(\xi\frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}\right).$$

We look for
$$s(r,t) = \tilde{s}(\xi(r,t))$$
, with $\xi = \frac{S}{4T} \cdot \frac{r^2}{t}$.

in terms of *s*:

in terms of \tilde{s} :

$$\begin{cases} S \frac{\partial s}{\partial t} = T \frac{1}{r} \frac{\partial}{\partial r} \left(r \frac{\partial s}{\partial r} \right) \\ s(r,0) = 0 \\ \lim_{r \to +\infty} s(r,t) = 0 \\ \lim_{r \to 0} r \frac{\partial s}{\partial r} = \frac{Q}{2\pi T} \end{cases} \begin{cases} -\xi \frac{d\tilde{s}}{d\xi} = \frac{d}{d\xi} \left(\xi \frac{d\tilde{s}}{d\xi} \right) \star \\ \tilde{s}(r,0) = 0 \\ \lim_{\xi \to +\infty} \tilde{s}(\xi) = 0 \\ \lim_{\xi \to +\infty} (\xi \frac{d\tilde{s}}{d\xi}(\xi)) = \frac{Q}{4\pi T} \end{cases}$$

$$\star \, \xi \frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = c e^{-\xi}, \qquad \lim_{\xi \to 0} \left(\xi \frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi}(\xi) \right) = \frac{Q}{4\pi T} \, \Rightarrow c = \frac{Q}{4\pi T}.$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling -

Theis' law

We look for
$$s(r,t) = \tilde{s}(\xi(r,t))$$
, with $\xi = \frac{S}{4T} \cdot \frac{r^2}{t}$.
$$\xi \frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = \frac{Q}{4\pi T} e^{-\xi} \ \Rightarrow \ \frac{\mathrm{d}\tilde{s}}{\mathrm{d}\xi} = \frac{Q}{4\pi T} \cdot \frac{\exp(-\xi)}{\xi},$$

$$\lim_{\xi \to +\infty} \tilde{s}(\xi) = 0 \ \Rightarrow -\tilde{s}(\xi) = \frac{Q}{4\pi T} \int_{\xi}^{+\infty} \frac{\exp(-\xi')}{\xi'} \mathrm{d}\xi'.$$

Theis' law (confined aquifer)

$$W(\xi) = \int_{\xi}^{+\infty} \frac{\exp(-\xi')}{\xi'} d\xi'.$$

$$s(r,t) = -\frac{Q}{4\pi T}W\left(\frac{S}{4T}\cdot\frac{r^2}{t}\right),$$

Summary

Thiem's law (confined aquifer, steady-state)

$$T = \frac{Q}{2\pi(s_2 - s_1)} \ln\left(\frac{r_2}{r_1}\right)$$

Dupuit's law (phreatic aquifer, steady-state)

$$K = \frac{Q}{2\pi(s_2 - s_1)\bar{L}} \ln\left(\frac{r_2}{r_1}\right)$$

Theis' law (confined aquifer, transient-flow regime)

$$s(r,t) = -\frac{Q}{4\pi T}W\left(\frac{S}{4T} \cdot \frac{r^2}{t}\right)$$

Physics of the Hydrosphere and the Cryosphere - Ground water modeling - slide 28/28