

SANDIA REPORT

SAND2008-5621

Unlimited Release

Printed December 2008

Sandia National Laboratories Environmental Fluid Dynamics Code: Sediment Transport User Manual

Phi Hung X. Thanh, Matthew D. Grace, and Scott C. James

Prepared by
Sandia National Laboratories
Albuquerque, New Mexico 87185 and Livermore, California 94550

Sandia is a multiprogram laboratory operated by Sandia Corporation,
a Lockheed Martin Company, for the United States Department of Energy's
National Nuclear Security Administration under Contract DE-AC04-94AL85000.

Approved for public release; further dissemination unlimited.



Issued by Sandia National Laboratories, operated for the United States Department of Energy by Sandia Corporation.

NOTICE: This report was prepared as an account of work sponsored by an agency of the United States Government. Neither the United States Government, nor any agency thereof, nor any of their employees, nor any of their contractors, subcontractors, or their employees, make any warranty, express or implied, or assume any legal liability or responsibility for the accuracy, completeness, or usefulness of any information, apparatus, product, or process disclosed, or represent that its use would not infringe privately owned rights. Reference herein to any specific commercial product, process, or service by trade name, trademark, manufacturer, or otherwise, does not necessarily constitute or imply its endorsement, recommendation, or favoring by the United States Government, any agency thereof, or any of their contractors or subcontractors. The views and opinions expressed herein do not necessarily state or reflect those of the United States Government, any agency thereof, or any of their contractors.

Printed in the United States of America. This report has been reproduced directly from the best available copy.

Available to DOE and DOE contractors from

U.S. Department of Energy
Office of Scientific and Technical Information
P.O. Box 62
Oak Ridge, TN 37831

Telephone: (865) 576-8401
Facsimile: (865) 576-5728
E-Mail: reports@adonis.osti.gov
Online order: <http://www.doe.gov/bridge>

Available to the public from

U.S. Department of Commerce
National Technical Information Service
5285 Port Royal Rd
Springfield, VA 22161

Telephone: (800) 553-6847
Facsimile: (703) 605-6900
E-Mail: orders@ntis.fedworld.gov
Online order: <http://www.ntis.gov/help/ordermethods.asp?loc=7-4-0#online>



SAND2008-5621
Unlimited Release
Printed December 2008

Sandia National Laboratories Environmental Fluid Dynamics Code: Sediment Transport User Manual

Phi Hung X. Thanh
Sandia National Laboratories
Thermal/Fluids Science and Engineering Department
P.O. Box 969
Livermore, CA 94551-0969
phthanh@sandia.gov

Matthew D. Grace
Sandia National Laboratories
Thermal/Fluids Science and Engineering Department
P.O. Box 969
Livermore, CA 94551-0969
mgrace@sandia.gov

Scott C. James
Sandia National Laboratories
Thermal/Fluids Science and Engineering Department
P.O. Box 969
Livermore, CA 94551-0969
scjames@sandia.gov

Abstract

In conjunction with this study, the USACE is funding development of a combined hydrodynamic, sediment transport, and water quality model for the lake based on the U.S. Environmental Protection Agency sponsored Environmental Fluid Dynamics Code (EFDC). Sandia National Laboratories (SNL) has been commissioned to perform the study and has made several updates to the code which is now called SNL-EFDC. The newly incorporated SEDZLJ sediment transport algorithm, which allows direct incorporation of site-specific erosion data from the Sediment Erosion with Depth flume (SEDflume), facilitates accurate prediction of sediment behavior.

This document describes the sediment dynamics subroutines and input files for the Sandia National Laboratories Environmental Fluid Dynamics Code (SNL-EFDC), which is a combined hydrodynamic, sediment transport, and water quality model based on the U.S. Environmental Protection Agency sponsored Environmental Fluid Dynamics Code (EFDC) developed by John Hamrick (1992). SNL-EFDC is enhanced with the incorporation of the SEDZLJ sediment transport model developed by Ziegler, Lick, and Jones (Jones 2001; Jones and Lick 2001; Ziegler et al. 2000). Detailed descriptions of the input files containing data from Sediment Erosion with Depth flume (SEDflume) measurements by Roberts et al. (2003) are provided, along with a description of the SEDZLJ routines. This model is unique in that it directly incorporates SEDflume data. Both a theoretical description of sediment transport employed in SNL-EFDC and an explanation of the source code are provided. This user manual is meant to be used in conjunction with the original EFDC manual (Hamrick 1996).

Through this document, the authors provide information for users who wish to model sediment transport with EFDC and who also seek a clear understanding of the source code, which is available on Sourceforge under GNU's General Public License: <http://sourceforge.net/projects/snl-efdc/>.

Acknowledgement

The research described in this document was funded by the U.S. Army Corps of Engineers, under Section 206 of the Water Resources Development Act of 1996, and by Sandia National Laboratories' Enabling Predictive Simulation Research Institute. Sandia is a multiprogram laboratory operated by Sandia Corporation, a Lockheed Martin Company, for the United States Department of Energy's National Nuclear Security Administration under contract DE-AC04-94AL85000.

Table of Contents

<i>Abstract</i>	4
<i>Acknowledgement</i>	4
<i>Introduction</i>	7
<i>SEDZLJ Sediment Input Files</i>	8
ERATE.SDF	8
CORE_FIELD.SDF	9
BED.SDF	10
<i>SEDZLJ Sediment Source Codes</i>	13
<i>SCANSEDZLJ.f90</i>	13
S_SEDIC.f90	13
S_MAIN.f90	15
S_SHEAR.f90	16
S_BEDLOAD.f90	21
S_SEDZLJ.f90	25
S_MORPH.f90	37
<i>Variable Dictionary</i>	39
<i>Special Notes</i>	45
<i>References</i>	46

List of Figures

Figure 1 : The input file for erosion rate data from SEDflume: <code>erate.sdf</code>	9
Figure 2: The input file for sediment bed cores of each cell in the entire domain: <code>core_field.inp</code>	10
Figure 3: The input file for sediment bed layer information: <code>bed.sdf</code>	12

Introduction

This document describes the Sandia National Laboratories Environmental Fluid Dynamics Code (SNL-EFDC), which is a modified version of the EPA's public-domain surface-water flow, sediment transport, and water-quality model developed by John Hamrick while at the Virginia Institute of Marine Sciences (Hamrick 1992). EFDC simulates flow and transport of sediment (in bedload and in suspended load), algae (including growth kinetics), and toxic substance (kinetic reactions and transport). SNL-EFDC improves EFDC with updated sediment dynamics subroutines (e.g., sediment exchange with the bed) developed by Ziegler, Lick, and Jones (Jones 2001; Jones and Lick 2001; Ziegler et al. 2000). Sediment erosion is calculated using data collected with a Sediment Erosion with Depth flume (SEDflume). SEDflume measures erosion rates as a function of shear stress and depth from relatively undisturbed cores taken directly from the sediment bed below the water body of interest. The use of SEDflume data provides more accurate sediment erosion rates, which are directly input to the model.

The SNL-EFDC code includes the following subroutines: `s_main.f90`, `s_sedic.f90`, `s_shear.f90`, `s_bedload.f90`, `s_morph.f90`, `s_sedzlj.f90`, and `SCANSEDZLJ.f90`. The main subroutine (`s_main.f90`) calls the other subroutines, orchestrates the numerical sediment dynamics, and outputs sediment fluxes to the EFDC transport routines. The initializer subroutines (`s_sedic.f90` and `SCANSEDZLJ.f90`) read data from the input files and convert the SEDflume data into EFDC variables. The shear subroutine (`s_shear.f90`) calculates the shear stress from the velocity profile output of the hydrodynamics portion of EFDC and any wind or wave contributions. Shear-stress information is required to calculate sediment erosion and deposition processes. The bedload subroutine (`s_bedload.f90`) calculates the probability of sediment being transported as saltating bedload or suspended load. It also calculates the sediment bedload flux and bedload sediment concentration. The primary sediment dynamics subroutine is the SEDZLJ subroutine (`s_sedzlj.f90`), which calculates erosion and deposition and determines the change in the bed-layer thickness, using the data read by `s_sedic.f90`. Lastly, the morphology subroutine (`s_morph.f90`) allows for the incorporation of the changing hydrodynamic boundary conditions due to sediment erosion and deposition. These sediment transport subroutines use three input files (`erate.sdf`, `bed.sdf`, and `core_field.sdf`), which are read by `s_sedic.f90` and `SCANSEDZLJ.f90`. Detailed description and examples for these input files are given below. In addition, a variable dictionary is provided at the end of the report. The source code for SNL-EFDC is available at <http://sourceforge.net/projects/snl-efdc/>.

SEDZLJ Sediment Input Files

ERATE.SDF

The example input file `erate.sdf` shown in Figure 1 contains erosion rate data from SEDflume. First, data about the thickness of each of the bed layer are provided in centimeters. The first two bed layer thicknesses are always initially zero (0.0) and are placeholders for the active and deposited layers (see the section on `S_SEDZLJ.F90`). Sediment layer 1 is the active layer, where bed-armoring can occur, and layer 2 is the depositional layer. Layers 3 through KB are described with SEDflume data and they may only decrease in thickness due to erosion. Note that the minimum number of layers is 3, where layers 1 and 2 are the initially zero-thickness active and depositional layers. A dimensionless multiplier for the active layer thickness (TACTM) follows the layer thicknesses. In addition, `erate.sdf` can have multiple core samples (although data from only a single core appears in Figure 1). The sediment bed will have different sets of core data as specified in `core_field.sdf` with integer identification (1, 2, 3, etc.) corresponding to the order of each core's data set in `erate.sdf`. As shown in the `core_field.sdf` file in Figure 2, the numbers 1 and 2 correspond to the first and second data tables in `erate.sdf`, respectively (note that only the first line of core 2 is shown). Within each table, the critical shear stress values are listed for each of the bed layers (4.25 dynes/cm² in our example). The number of bed layers is specified in Card 36 of the main input file `efdc.inp`, as well as in `bed.sdf`. Critical shear stress for erosion (TAUTEMP; dynes/cm²) is the minimum shear stress necessary for erosion to occur. The next set of inputs describes the bed or sediment core density (BDEN; g/cm³), which accounts for both solid sediment density. Water density (WATERDENS; g/cm³) and the solid sediment density (SEDDENS; g/cm³) are listed next. In the example in Figure 1, the water density is reported at STP as 1.0 g/cm³ and the solid sediment density is for quartz at 2.6 g/cm³. The particle size distribution as a percentage is listed after the sediment density information. In the file `efdc.inp`, the number of particle sizes, NSED (also referred to as the number of particle size classes), is specified in both Card 22 (`efdc.inp`) and in `bed.sdf`. Next, erosion rate data are listed. The erosion rate data start with the value of the applied shear with the following line providing the erosion rate for each sediment layer at that shear stress. In Figure 1, the applied shear stresses are 0, 2, 4, 8, 10, and 20 Pa, and the erosion rates are listed in cm/s. For multiple core samples, critical shear stress, bulk density, particle size distribution, and initial bed erosion rates are reported in the same format below the first set of core sample data.


```

*****
# Layer Thicknesses (cm) TSED0S #
0.0 0.0 15.0 15.0 15.0
# Active Layer Thickness Multiplier TACTM #
5.0
# Critical Shear Stresses for Erosion (dynes/cm^2) TAUTEMP & TAUCOR
#Data for sediment 1
4.25 4.25 4.25 4.25 4.25
# Bed core Density (g/cm^3) BDEN #
1.9 1.9 1.9 1.9 1.9
# Water and Sediment Density (g/cm^3) WATERDENS and SEDDENS #
1.0 2.6
# Particle Size Distribution (%) PNEW #
2.955 6.649 12.664 13.493 14.679 27.094 9.223 10.231 3.184
2.955 6.649 12.664 13.493 14.679 27.094 9.223 10.231 3.184
2.955 6.649 12.664 13.493 14.679 27.094 9.223 10.231 3.184
2.955 6.649 12.664 13.493 14.679 27.094 9.223 10.231 3.184
2.955 6.649 12.664 13.493 14.679 27.094 9.223 10.231 3.184
# Initial Bed Erosion Rates (cm/s) EORATE & TAULOC #
0.0 !shear stress 1
0.000E+00 0.000E+00 0.000E+00 0.000E+00 0.000E+00 !erosion rate of
each layer
2.00000 !shear stress 2
0.000E+00 0.000E+00 0.000E+00 0.000E+00 0.000E+00
4.00000 !shear stress 3
8.000E-05 8.000E-05 8.000E-05 8.000E-05 8.000E-05
8.00000 !shear stress 4
7.000E-03 7.000E-03 7.000E-03 7.000E-03 7.000E-03
10.0000 !shear stress 5
1.140E-02 1.140E-02 1.140E-02 1.140E-02 1.140E-02
20.0000 !shear stress 6
0.649E-01 0.649E-01 0.649E-01 0.649E-01 0.649E-01
# Critical Shear Stresses for Erosion (dynes/cm^2) TAUTEMP & TAUCOR
#Data for sediment 2 (same structure as above)
*****

```

Figure 1 : The input file for erosion rate data from SEDflume: erate.sdf.

CORE_FIELD.SDF

The core_field.sdf file specifies which core data set listed in erate.sdf corresponds to which cells of the bed. The first entry in the file is the number of different data sets (or the number of different data cores collected). After the number of cores, a matrix with the core identification number for each cell (defined in the cell.inp and celllt.inp files). Note that core data exist for water cell (active cells) and non-water cells. In the example core_field.inp file in Figure 2, the model has 9 rows and 15 columns.

```

*****
2 !Number of cores
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
2 2 2 1 1 1 1 1 1 1 1 1 1 1 1
*****

```

**Figure 2: The input file for sediment bed cores of each cell in the entire domain:
core_field.inp.**

BED.SDF

The input file `bed.sdf` contains information about the sediment bed layers. The first line contains names of the model parameter flags, whose values appear in the second line. When `VAR_BED = 1`, the variable sediment bed option is activated, resulting in the usage of data from `core_field.sdf`, allowing the model to incorporate spatially varying sediment characteristics. The `BEDLOAD` flag for bedload transport appears next (bedload is typically only active for river systems). `NEQUIL` is the flag that activates adsorption of toxics onto sediments (currently not available in SNL-EFDC). `KB` is the maximum number of sediment layers, and should coincide with the `KB` value in Card 36 of the `EFDC.INP` input file. `ISEDTIME` is the number of time steps before sediment transport is activated. This option allows the velocity profile to develop before sediment transport occurs. The `ISMORPH` flag activates morphology calculations for cases when erosion or deposition of sediment on the bed significantly changes the topology of the bed. `ISMORPH` incorporates the change in topology into the hydrodynamic model. `IFWAVE` is the flag that turns on wind-driven waves that introduce additional shear at the sediment bed. `MAXDEPLIMIT` is the maximum fraction of sediment in the bottom water layer allowed to deposit onto the sediment bed during a single time step (it should usually be 1.0 unless instabilities are observed and no less than 0.5). `ITBM` and `NSICM` are array parameters read by `SCANSEDZLJ.f90`. `ITBM` is the number of different `SEDflume` shear-stress interpolants used to calculate erosion of newly deposited sediments. These values are listed `erate.sdf` and in the example from Figure 3, they have values 0, 2, 4, 8, 10, and 20 Pa. `NSICM` is the number of size classes of interpolants used to estimate erosion rates of newly deposited sediment (described below). `ZBSKIN` (μm) is the roughness parameter, which accounts for both grain-size roughness and bed-form roughness (topographical variations from dunes and ripples at the bottom of the water body). If `ZBSKIN` is zero, then each model cell's roughness is the local average particle size (μm), the average particle diameter in the cell. `TAUCONST` (dynes/cm^2) is an option that, when greater than zero, allows the user to set the shear stress from the fluid flow to a constant value (mainly used for debugging). If `TAUCONST` is zero, the stresses are calculated from the velocity profile resulting from the hydrodynamic model. The next line of data contains the discretized sediment size classes (μm). In the example file in Figure 3, there are nine different size classes used to approximate the sediment distribution for the model of interest. For each of these size classes, the discretized diameter (`D50`; μm) is listed, and values of critical shear for

erosion (TCRE; dynes/cm²) and critical shear for suspension (TCRSUS; dynes/cm²) are reported in the following lines. The critical shear for erosion is calculated according to Soulsby (1997) and the critical shear for suspension is calculated according to equations (8) and (9) of van Rijn (1984a). These two values are used to calculate the fraction of sediment in suspended load versus bedload. The rest of the `bed.sdf` input file contains information for newly deposited sediments from SEDflume or some other data source. Newly deposited sediments are defined as those that have undergone erosion (and perhaps suspension) and have settled on the bed again. The next line of data provides the NSICM interpolant sediment size classes, SCLOC(1:NSICM) (μm). For each interpolant size class (SCLOC), the critical shear stress for erosion (TAUCRITE; dynes/cm²) is provided. These values are either site-specific data from SEDflume or can be calculated according to equations (6) or (9) and (10) from Roberts et al. (1998). Lastly, a table of erosion rate data for the newly deposited sediment is listed (ENRATE; cm/s). ENRATE is calculated for each SCLOC at the shear values listed in `erate.sdf` (0, 2, 4, 8, 10, and 20 Pa). These are typically from site-specific data or Table 1 of Robert et al. (1998).

```

*****
# VAR_BED Bedload Nequil KB ISEDTIME IMORPH IFWAVE MAXDEPLIMIT #
      1      1      0      5      100      0      0      1.0
# ITBM NSICM Array Parameters #
      6      8
# ZBSKIN (um), > 0 to set constant TAUCONST (dynes/cm^2), > 0 to set
constant #
      1500.0      10.0
# Discretized Size Classes (um) #
237.7 427.0 603.5 853.5 1070.5 1570.5 2415.0 3415.0 5450.1
# Critical Shear for Erosion (dynes/cm2) TCRE #
1.5 2.4 3.3 4.25 7.6 9.5 10.8 16. 24.8 !Soulsby 1997
# Critical Shear for Suspension (dynes/cm2) TCRSUS #
1.5 2.6 4.5 21.2 53.6 67.3 87.9 122.6 192.0 !van Rijn 1984b
Eqs. (8) and (9)
# Sediment Bed Sizes (um) SCLOC #
125.0 222.0 432.0 1020.0 2000.0 2400.0 3000.0 6000.00
# Critical Shear for Erosion of Bed Surface (dynes/cm^2) TAUCRITE
#!Robert et al. 1998, then 4.14d
1.20 2.27 2.96 4.17 5.46 5.88 6.42 8.48
# Erosion Rates for Active and Deposited Layers (cm/s) # !Robert et
al. 1998, Table 1
1E-9 6.60E-5 4.66E-4 3.29E-3 6.17E-3 4.36E-2 !125 um !These six
values are for the 6 shear stresses specified in ITBM and defined in
erate.sdf as the initial bed erosion rates.
1E-9 5.97E-5 5.97E-4 5.96E-3 1.25E-2 1.25E-1 !222 um
1E-9 3.65E-4 2.16E-3 1.27E-2 2.25E-2 1.33E-1 !432 um
1E-9 2.01E-4 1.14E-3 6.51E-3 1.14E-2 6.49E-2 !1020 um
1E-9 2.25E-5 1.63E-4 1.19E-3 2.25E-3 1.63E-2 !2000 um
1E-9 1.12E-5 8.40E-5 6.28E-4 1.20E-3 8.98E-3 !2400 um
1E-9 3.96E-6 3.07E-5 2.28E-4 4.59E-4 3.56E-3 !3000 um
1E-9 2.03E-8 1.67E-7 1.44E-6 2.88E-6 2.49E-5 !6000 um
*****

```

Figure 3: The input file for sediment bed layer information: bed.sdf.

SEDZLJ Sediment Source Codes

SCANSEDZLJ.f90

SCANSEDZLJ.f90 loads data from `bed.sdf` and evaluates its compatibility with parameters such as number of sediment layers, particle size class, etc. from `EFDC.INP`. This routine identifies some errors in the input files and also helps with dynamic memory allocation.

S SEDIC.f90

This routine is called from the EFDC routine `INPUT.for` and reads the three SEDZLJ input files (`erate.sdf`, `core_field.sdf`, and `bed.sdf`). Flag option variables in `bed.sdf` are read first. After these flags, the roughness and shear stress are loaded. After the shear stress, the discretized particle size classes (D50) are loaded. The critical shear stresses for erosion (TCRE) and critical shear stresses for suspension (TCRSUS) for each sediment size class (D50) are read next. The file `bed.sdf` will also be read later for newly deposited sediment erosion information. Note that while the number of sediment size classes is defined in `EFDC.INP` (as the number of cohesive sediments), NSICM is defined here as the number of interpolant size classes defined.

After obtaining the first portion of data in `bed.sdf`, `s_sedic.f90` then reads `core_field.sdf`. The number of different core samples and the core sample corresponding to each cell in the domain is read. Next, if `VAR_BED` equals 1 (multiple SEDflume cores are read), the routine reads `bed` and core sample data from `erate.sdf` for erosion calculations before any deposition or sorting occur. `s_sedic.f90` first reads layer thicknesses (TSEDOS; g/cm^2), which are in units of cm, from the `erate.sdf` file. Note that these thicknesses are converted to g/cm^2 by dividing by the bulk density for the bed layer (BDEN) to make subsequent mass flux calculations consistent. `TAUTEMP` (dynes/cm^2) contains the critical shear stresses necessary to erode each sediment layer. `BDEN` (g/cm^3) is the bulk density for the bed layers; this is used in conjunction with bed thickness to define bed layer thicknesses (TSED, TACT, etc., all in units of g/cm^2). `WATERDENS` and `SEDDENS` are the water density and solid sediment density (from `erate.sdf`). `PNEW` is the mass percentage of each sediment size class in each layer, expressed as a percentage, but the form of `PNEW` in the code is the fraction `PER` (`PER` = `PNEW`/100.0). Lastly, `TAULOC` (dynes/cm^2) and `EORATE` (cm/s) are loaded, with `TAULOC` being the applied shear and `EORATE` the erosion rate of each sediment layer subject to the applied `TAULOC`. `EORATE` is the erosion rate of the site-specific (original bed) data while `ENRATE` from `bed.sdf` is the erosion rate of newly deposited sediment.

`s_sedic.f90` then sets the EFDC variables equal to the data loaded from the input files. The critical shear stress of each core, `TAUCOR` (dynes/cm^2), and `ERATE` (equivalent of `EORATE`; cm/s) are initially equated to the corresponding values from `erate.sdf` with only a change in units (cgs to mks) to correspond with the convention of EFDC. The weight fraction, `PER`, is calculated by dividing the particle's mass percentage `PNEW` by 100, converting mass percentage to mass fraction. Also, the bulk density (`BULKDENS`; g/cm^3) is the mass of solids per unit volume:

$$\chi_s \rho_s = \frac{\rho_s (\rho - 1)}{\rho_s - \rho_w}, \quad (1)$$

where χ_s is the mass fraction (PER) of the sediment in the bed, ρ_s is the density of the sediment (SEDDENS), the product ($\chi_s \rho_s$) is the mass of solids per unit volume (BULKDENS), ρ_w is the density of water (WATERDENS), and ρ is the density of the sediment bed (BDEN). The sediment bed density (BDEN) is defined as:

$$\rho = \chi_s \rho_s + \chi_w \rho_w, \quad (2)$$

where the mass fractions χ_s and χ_w sum to 1. The bed layer thicknesses (i.e., BEDLINIT, which contains the initial bed layer thickness, and HBED; m) must be non-zero and are passed to the EFDC variable HBED. The initial bed thickness BEDLINIT is reported in units of meters, hence the thickness values from the input file are multiplied by 0.01. `s_sedic.f90` activates all layers (i.e., sets LAYER equal to 1) and assigns each layer's thickness (TSED and TSED0; g/cm²) as BEDLINIT (cm) multiplied by the mass per unit volume (BULKDENS; g/cm³).

Total sediment bed thickness (HBEDA; cm) is read in next. The bed elevation (ZELBEDA; cm) is calculated with the bed elevation (BELV; cm) and the total sediment thickness (HBEDA; cm). `s_sedic.f90` then reads `bed.sdf` to obtain information about newly deposited sediments. SCLOC is a vector of the sediment interpolant size classes used to interpolate erosion rates (ENRATE; cm/s) of newly deposited sediments. The critical shear stress for each sediment interpolant size class (SCLOC) is read as TAUCRITE. The erosion rate of newly deposited sediment (ENRATE) for each sediment interpolant size class (SCLOC) at different shear stresses (TAULOC) are read next.

The settling velocity (cm/s) is (Cheng 1997)

$$w_s = \frac{v}{d} \left(\sqrt{25 + 1.2d^{*2}} - 5 \right)^{1.5}, \quad (3)$$

where w_s is the settling velocity, v is the kinematic fluid viscosity (cm²/s), d is the particle diameter (cm), and d^* is the non-dimensional particle diameter defined as

$$d^* = d \left[\frac{(s_s - 1)g}{v^2} \right]^{1/3}, \quad (4)$$

where s_s is the sediment specific gravity (ρ_s/ρ_w), and g is gravity (cm/s²). All of these values are calculated in cgs units (van Rijn 1984b). In the code, DWS is the settling speed (cm/s) and the non-dimensional diameter (d^*) is DISTAR. Because d (D50) is in units of μm , it is divided by 10,000 to convert it to centimeters. The sediment specific gravity, s_s , is read in from `ERATE.SDF`, gravity is hard wired to 980.0 cm/s², and the kinematic viscosity of water is also hard wired to 0.01 cm²/s.

`s_sedic.f90` also calls on two input files (`fetch.inp` and `stwave.inp`) as necessary to read information about wind fetch length and the corresponding wave field.

The routine ends after printing relevant information to the screen and closing the input files.

S MAIN.f90

This routine calculates sediment dispersion through the various water layers. It is called from `CALSED.for` where EFDC previously calculated cohesive sediment transport. `s_main.f90` replaces the current (and separate) cohesive and non-cohesive sediment transport routines in EFDC. The first step in `s_main.f90` involves calling the `s_shear.f90` subroutine to calculate SEDZLJ shear stresses.

If the morphology feature is activated, then `s_main.f90` calculates sediment bed height (HBED; m) for each layer in each cell. The layer thickness, TSED (g/cm^2), is divided by the solid mass per volume (BULKDENS) to obtain HBED in units of meter. The suspended sediment concentration (SED; kg/m^3) is also saved in another variable (SEDS; kg/m^3) for later calculations for vertical turbulent dispersion and settling. Next, `s_main.f90` checks if bedload transport (NCALC_BL) is active. The velocity profile in the model is allowed to develop for ISEDTIME time steps before the sediment transport is activated. If bedload sediment feature is activated (i.e., NCALC_BL = 1), `s_main.f90` calls the bedload subroutine `s_bedload.f90`. `s_bedload.f90` also evaluates the bedload concentrations and bedload fluxes to and from the bed.

`s_sedzlj.f90` is activated in `EFDC.inp` with an IWRSP value of 98 (simply a unique integer). Here, both NSEDFLUME and ISEDTIME are checked by `s_main.f90` before sediment transport calculations are performed.

`s_main.f90` calculates the vertical dispersion of suspended sediment using the same formulation as Hamrick (1992). Based on the number of layers, `s_main.f90` evaluates vertical dispersion. “IF” conditional statements are added to account for special cases with one or two layers. The sediment flux due to vertical dispersion is a function of vertical velocity (WVEL; m/s), settling velocity (WSETA; m/s), and suspended sediment concentration (SED). Once the vertical dispersion of suspended sediments has been calculated, `s_main.f90` calls `SEDZLJ` to calculate sediment exchange between the sediment bed and water column.

Once sediment transport is complete, the total sediment flux (QSBDBTOP; $\text{kg/m}^2\text{s}$) is calculated as a function of the inverse specific gravity (SSGI) and the sediment flux (SEDF; $\text{kg/m}^2\text{s}$), initial void ratio in each layer of the bed (VDRBED), and sediment void ratio (VDRDEPO). Next, `s_main.f90` goes through a series of calculations that estimate the sediment flux between water layers. Vertical sediment transport is due to turbulent dispersion and settling. Horizontal transport is due to advection and dispersion as calculated by the standard EFDC routines.

Once `s_main.f90` re-calculates the sediment concentration profile due to turbulent dispersion, the sediment flux (SEDF; $\text{kg/m}^2\text{s}$) due to settling is calculated by starting with the top layer of water and then calculating fluxes to the deeper water layers. SEDF is calculated as the product of

the water layer's height ($HP \times DZC$; m), change in sediment concentration over the last time step ($SED - SEDS$), and the inverse time step, $DELTI$ (1/s). For water layers below the top layer, the sediment flux ($SEDF$) is the current layer's concentration plus the flux from the layer directly above it. The sediment flux algorithm depends on the number of water layers and `s_main.f90` accounts for this with four distinct cases. The first case is for situations with more than three water layers, the second is for problems with exactly three water layers, the third case is for problems with two water layers, and the final case is for problems with only one water layer (for one- and two-dimensional representations).

Once the suspended sediment flux is calculated, `s_main.f90` calls the morphology routine `s_morph.f90` to update the grid for the ensuing hydrodynamics.

S SHEAR.f90

The shear subroutine uses the hydrodynamic results from EFDC to calculate the total shear stress at the sediment bed, which is then used in sediment erosion calculations. The shear stress and friction factor in `s_shear.f90` are calculated using Christofferson and Jonsson (CJ) log-law (1985), which is a function of the wind, wave, and convective information from the hydrodynamics of EFDC.

`s_shear.f90` first checks if the constant shear assumption is activated in the `bed.sdf` file.

```
IF (TAUCONST==0) THEN
```

If the TAUCONST flag is not activated (i.e., not equal to zero), then `s_shear.f90` calculates the shear stress based on the local velocities of each cell, accounting for shear stress from the velocity gradient as well as wind- and wave-induced shear stresses. First, wave characteristics are calculated (i.e., wavelength, orbital velocity, height, and frequency) according to the lines of code below.

```
IF (IFWAVE==1) THEN
!Wind wave fetch
!Convert wind input into current wind info for wind-driven wave calcs
  IF (ISDYNSTP==0) THEN
    WFTIM=DT*FLOAT(N)/TCWSER(1)+TBEGIN*(TCON/TCWSER(1))
  ELSE
    WFTIM=TIMESEC/TCWSER(1)
  ENDIF
  M1=MWTLAST(1)
  M2=M1+1
  DO WHILE (TWSER(M2,1)<WFTIM)
    M1=M2
    M2=M1+1
  END DO
  MWTLAST(1)=M1
  TDIFF=TWSER(M2,1)-TWSER(M1,1)
  WTM1=(TWSER(M2,1)-WFTIM)/TDIFF
  WTM2=(WFTIM-TWSER(M1,1))/TDIFF
```



```

FWINDS=WTM1*WINDS(M1,1)+WTM2*WINDS(M2,1)
IF(FWINDS>1.0) THEN
  IF(ABS(WINDD(M1,1)-WINDD(M2,1))<180.0) THEN
    FWINDD=WTM1*WINDD(M1,1)+WTM2*WINDD(M2,1)
  ELSE
    IF(WINDD(M1,1).GT.WINDD(M2,1)) THEN
      FWINDD=WTM1*WINDD(M1,1)+WTM2*(WINDD(M2,1)+360.0)
    ELSE
      FWINDD=WTM1*(WINDD(M1,1)+360.0)+WTM2*WINDD(M2,1)
    ENDIF
    IF(FWINDD>=360.0)FWINDD=FWINDD-360.0
    ENDIF
    !Convert wind into direction it is blowing "from"
    IF(FWINDD<=180.0) THEN
      FWINDD=FWINDD+180.0
      IF(FWINDD==360.0)FWINDD=0.0
    ELSE
      FWINDD=FWINDD-180.0
      IF(FWINDD==360.0)FWINDD=0.0
    ENDIF
    !Calculate which of the 8 wind zones (FWZONE) the wind is coming
from
    !Also the Waveangle CCW from East
    IF(FWINDD>=337.5.OR.FWINDD<22.5) THEN
      FZONE=1
      WVANGLE=4.712
    ELSEIF(FWINDD>=22.5.AND.FWINDD<67.5) THEN
      FZONE=2
      WVANGLE=3.927
    ELSEIF(FWINDD>=67.5.AND.FWINDD<112.5) THEN
      FZONE=3
      WVANGLE=3.142
    ELSEIF(FWINDD>=112.5.AND.FWINDD<157.5) THEN
      FZONE=4
      WVANGLE=2.356
    ELSEIF(FWINDD>=157.5.AND.FWINDD<202.5) THEN
      FZONE=5
      WVANGLE=1.571
    ELSEIF(FWINDD>=202.5.AND.FWINDD<247.5) THEN
      FZONE=6
      WVANGLE=0.7854
    ELSEIF(FWINDD>=247.5.AND.FWINDD<292.5) THEN
      FZONE=7
      WVANGLE=0.
    ELSEIF(FWINDD>=292.5.AND.FWINDD<337.5) THEN
      FZONE=8
      WVANGLE=5.4978
    ENDIF
    !Calculate Domain Average Depth
    !Needs to be calculated along each fetch
    !This is sufficient for small systems
    AVGDEPTH=SUM(HP(2:LA))/ FLOAT(LA-1)

```

```

FWINDSQ=FWINDS*FWINDS
!Calculate wave height, period, orbital velocity, and length
DO L=2,LA
  FC1=(FWINDSQ/9.8)*0.283*TANH(0.530*(9.8*AVGDEPTH/FWINDSQ)**0.75)
  FC2=TANH(0.0125*(9.8*FWDIR(L,FZONE)/FWINDSQ)**0.42/
    TANH(0.530*(9.8*AVGDEPTH/FWINDSQ)**0.75))
  FWVHT(L)=MIN(HP(L),FC1*FC2)
  FC1=(FWINDS/9.8)*7.54*TANH(0.833*(9.8*AVGDEPTH/FWINDSQ)**0.375)
  FC2=TANH(0.077*(9.8*FWDIR(L,FZONE)/FWINDSQ)**0.25/
    TANH(0.833*(9.8*AVGDEPTH/FWINDSQ)**0.375))
  FWVTP(L)=FC1*FC2
  WVFREQ(L)=2.0*PI/FWVTP(L)
  FC1=(2.0*PI/FWVTP(L))**2*HP(L)/9.8
  FC2=FC1+1.0/(1.0+0.6522*(FC1)+0.4622*(FC1)**2+
    0.0864*(FC1)**4+0.0675*(FC1)**5)
  WVLENGTH=FWVTP(L)*SQRT(9.8*HP(L)/FC2)

WVORBIT(L)=MAX(0.01,PI*FWVHT(L)/(FWVTP(L)*SINH(HP(L)*2.0*PI/WVLENGTH))
)
ENDDO
!*****
****
!Read EFDC STWAVE Wave Field
ELSEIF(IFWAVE==2) THEN
  NWVCOUNT=NWVCOUNT+1
IF(NWVCOUNT==STWVTIM) THEN
  NWVCOUNT=0
  STINC=STINC+1
DO L=2,LA
  IF(STINC>STWVNUM)EXIT
  IF(STWVTP(L,STINC)>0.0) THEN
    WVFREQ(L)=2.0*PI/STWVTP(L,STINC)
    FWVHT(L)=MAX(HP(L),FWVHT(L))
    FC1=(2.0*PI/STWVTP(L,STINC))**2*HP(L)/9.8
    FC2=FC1+1.0/(1.0+0.6522*(FC1)+0.4622*(FC1)**2+0.0864*(FC1)**4+
      0.0675*(FC1)**5)
    WVLENGTH=STWVTP(L,STINC)*SQRT(9.8*HP(L)/FC2)

WVORBIT(L)=MAX(0.01,PI*STWVHT(L,STINC)/(STWVTP(L,STINC)*
SINH(HP(L)*2.0*PI/WVLENGTH))
  WVANG(L)=STWVDR(L,STINC)
  ELSE
    WVFREQ(L)=0.0
    WVORBIT(L)=0.0
    WVANG(L)=0.0
  ENDIF
ENDDO
ENDIF
ENDIF
ENDIF

```

After the wave characteristics are calculated, `s_shear.f90` calculates shear stresses. Bed-form roughness is determined in three distinct ways. If the bed-form roughness is not provided from the SEDZLJ subroutines (i.e., `ZBSKIN = 0` and `NSEDFLUME = 0`), it is set equal to the EFDC roughness (`ZBR`, as loaded from `EFDC.INP`). If `NSEDFLUME` is greater than 0 but no bed-form roughness (`ZBSKIN`) is specified, then it is set equal to the average grain roughness (`D50AVG`). Otherwise, the bed-form roughness is equal to the value specified in `bed.sdf` (`ZBSKIN`).

```
DO L=2,LA
  ZBTEMP(L)=ZBR(L)
  IF (ZBSKIN.EQ.0.AND.NSEDFLUME.GT.0) THEN
    IF (D50AVG(L).LT.D50(1)) THEN
      ZBTEMP(L)=D50(1)/1.e6
    ELSE
      ZBTEMP(L)=D50AVG(L)/1.e6
    ENDIF
  ELSEIF (ZBSKIN.GT.0.AND.NSEDFLUME.GT.0) THEN
    ZBTEMP(L)=ZBSKIN/1.e6
  ENDIF
  KN(L)=30.0*ZBTEMP(L)
ENDDO
```

To obtain the shear stress, `s_shear.f90` calculates the average velocity magnitude.

```
DO L=2,LA
!Calculate Average Velocity Magnitude in cm/s
  UTMP=100.0*STCUV(L)*(UHE(L+1)+UHE(L))/(HU(L+1)+HU(L))+1.0E-12
  VTMP=100.0*STCUV(L)*(VHE(LNC(L))+VHE(L))/(HV(LNC(L))+HV(L))
  VELMAG=SQRT(UTMP**2+VTMP**2)
```

The friction factor (`FC`) is then calculated with the CJ log-law using the water height and bed-form roughness (`ZBTEMP`).

```
FC(L)=(0.42/LOG(HP(L)/(2.0*ZBTEMP(L))))**2
```

If there are no wave or wind inputs, then the shear stress can be calculated directly with only the velocity magnitude and the friction factor using the CJ log-law. Otherwise, a more complex method is used to incorporate the wind- and wave-induced shear stresses into the shear stress (`TAU`; dynes/cm²) calculation.

```
IF (IFWAVE==0.AND.UWVSQ(L)==0.0.OR.WVORBIT(L)==0.0) THEN
  TAU(L)=FC(L)*VELMAG**2
ELSE !Calculate Current Angle CCW From X axis
  IF (UTMP>0.0.AND.VTMP>0.0) THEN
    VELANG=ATAN(VTMP/UTMP)
  ELSEIF (UTMP<0.0) THEN
    VELANG=ATAN(VTMP/UTMP)+PI
  ELSEIF (UTMP>0.0.AND.VTMP<0.0) THEN
    VELANG=2*PI+ATAN(VTMP/UTMP)
  ELSEIF (UTMP==0.0) THEN
```

```

        VELANG=SIGN(0.5*PI,VTMP)
    ENDIF
    IF(IFWAVE==0) THEN !Set Orbital velocity in m/s and waveangle and
frequency
        WVORBIT(L)=SQRT(UWVSQ(L))
        WVFREQ(L)=WVFRQ
        WVANG(L)=WACCWE(L)
    ELSEIF(IFWAVE==1) THEN
        WVANG(L)=WVANGLE
    ENDIF
    !Calculate wave friction factor
    FWW(L)=2.0*(0.0747*(KN(L)*WVFREQ(L)/WVORBIT(L))**0.66
    SIGMAWV=FC(L)/FWW(L)*(VELMAG/(WVORBIT(L)*100.0))**2
    MMW=SQRT(1.0+SIGMAWV**2+2.0*SIGMAWV*ABS(COS(VELANG-WVANG(L))))
    JJW=WVORBIT(L)/(KN(L)*WVFREQ(L))*SQRT(MMW*FWW(L)/2.0)
    FWW(L)=MMW*0.15/JJW
    !Calculate wave boundary layer info
    DELW=KN(L)*0.273*SQRT(JJW)
    APROUGH=30.0*DELW*EXP(-5.62*DELW/KN(L)*SQRT(SIGMAWV/MMW))
    !Calculate new current friction factor
    FC(L)=2.0*(1.0/(2.38*LOG(30.0*HP(L)/(2.718*KN(L))))-
2.38*LOG(APROUGH/KN(L))**2
    !Iterate once more
    SIGMAWV=FC(L)/FWW(L)*(VELMAG/(WVORBIT(L)*100.0))**2
    MMW=SQRT(1.0+SIGMAWV**2+2.0*SIGMAWV*ABS(COS(VELANG-WVANG(L))))
    JJW=WVORBIT(L)/(KN(L)*WVFREQ(L))*SQRT(MMW*FWW(L)/2.0)
    FWW(L)=MMW*0.15/JJW

```

The total shear stress (TAU) is calculated in cgs units (dynes/cm²), and is converted to Pascals and renamed TAUB. From TAUB, the shear velocity u^* (USTAR; m/s) is calculated as:

$$u^* = \sqrt{\frac{\tau}{\rho_w}}, \quad (5)$$

where τ (TAUB; Pa) is the shear stress and ρ_w is the density of water (Jones 2001, page 14). Shear velocity is used in many of the subsequent calculations for sediment transport.

```

!Calculate total wave and current shear stress (dynes/cm^2)
    TAU(L)=0.5*FWW(L)*WVORBIT(L)**2*10000.0*MMW
    TAUB(L)=0.1*TAU(L)
    USTAR(L)=SQRT(TAUB(L)/RHO)
ENDIF
ENDDO

```

If the shear stress is defined as constant in bed.sdf, then s_shear.f90 skips the calculation of friction- and wave-induced shear and shear is simply set equal to the constant value specified in the input file.

```

ELSE
  DO L=2, LA
    TAU(L)=TAUCONST
    TAUB(L)=0.1*TAU(L)
    USTAR(L)=SQRT(TAUB(L)/1000.0)
  ENDDO
ENDIF

```

S BEDLOAD.f90

The `s_bedload.f90` routine calculates the bedload transport dynamics

`s_bedload.f90` first determines the ratio (USW) of shear velocity to settling velocity, which is a function of applied shear stress (TAU), density of water (WATERDENS), and settling velocity (DWS).

```
USW(L,K)=SQRT(TAU(L)/WATERDENS)/DWS(K)
```

Next, the subroutine calculates the fraction of sediment transported in bedload versus suspended load using both flow and sediment properties (i.e., sediment characteristics, shear stress, and shear velocity at each cell). Using the empiricism of Guy et al. (1966), `s_bedload.f90` assumes that all particles with a diameter of less than 200 μm are transported as suspended load (unless there is no flow). Particles with a diameter of greater than 200 μm and a ratio of shear velocity (USTAR) to settling velocity (DWS) of less than 4 are partially transported as bedload. The ratio of suspended to total sediment transported is calculated as (Jones 2001, eq. 3.7):

$$\frac{q_s}{q_t} = \begin{cases} 0 & \text{for } \tau < \tau_{cs} \\ \frac{\ln\left(\frac{u^*}{w_s}\right) - \ln\left(\frac{\sqrt{\tau_{cs}/\rho_w}}{w_s}\right)}{\ln(4) - \ln\left(\frac{\sqrt{\tau_{cs}/\rho_w}}{w_s}\right)} & \text{for } \tau > \tau_{cs} \text{ and } \frac{u^*}{w_s} < 4, \\ 1 & \text{for } \frac{u^*}{w_s} > 4 \end{cases} \quad (6)$$

where q_s/q_t is the ratio of suspended sediment flux to total sediment flux, u^* is defined in Eq. (5), w_s is the settling velocity (cm/s) defined in Eq. (3), and τ_{cs} (dynes/cm²) is the critical shear stress for suspension read from `bed.sdf`. The corresponding ratio of bedload sediment flux to total sediment flux is simply $1 - q_s/q_t$.

```

LCM_LOOP_1:DO L=2,LA
WHERE (D50(1:NSCM)>=200.0.AND.USW(L,1:NSCM)<4.0)
  WHERE (TAU(L)<=TCRE(1:NSCM).OR.USW(L,1:NSCM)<=0.0)
    BLFLAG(L,1:NSCM)=0
    PSUS(L,1:NSCM)=0.0
  ELSEWHERE
    BLFLAG(L,1:NSCM)=1
    PSUS(L,1:NSCM)=MAX((LOG(USW(L,1:NSCM))-
      LOG(SQRT(TCRSUS(1:NSCM)/DWS(1:NSCM)))/(LOG(4.0)-
      LOG(SQRT(TCRSUS(1:NSCM)/DWS(1:NSCM))),0.0)
  ENDWHERE
ELSEWHERE
  BLFLAG(L,1:NSCM)=0
  PSUS(L,1:NSCM)=1.0
ENDWHERE
ENDDO LCM_LOOP_1

```

After evaluating the fraction of sediment in bedload and suspended load, `s_bedload.f90` calculates the bedload velocity (BLVEL; cm/s). First, the velocity magnitude (VELMAG; m/s) is calculated in MKS units. For each sediment size class, a dimensionless transport parameter (TRANS), is defined as a function of the shear stress (TAU; dynes/cm²) and the critical shear stress (TCRE) provided in the `bed.sdf` input file, is (van Rijn 1984b):

$$\tau_* = \frac{\tau - \tau_{ce}}{\tau_{ce}}, \quad (7)$$

where τ_* (TRANS) is the transport parameter, τ (TAU; dynes/cm²) is shear stress at bed-fluid interface, and τ_{ce} (TCRE; dynes/cm²) is critical shear stress for erosion.

Next, `s_bedload.f90` calculates the saltation height, DZBL (cm), which is a function of transport parameter (TRANS), particle size class (D50), and non-dimensional particle diameter (DSTAR) calculated in `s_sedic.f90` (van Rijn 1984b):

$$h_b = 0.3dd^{*0.7}\tau_*^{0.5}, \quad (8)$$

where h_b (DZBL; cm) is the bedload or saltation height.

The bedload velocity u_b (BLVEL; cm/s) is calculated using the transport parameter (TRANS) and the size class (D50):

$$u_b = 1.5\tau_*^{3/5} \left[\left(\frac{\rho_s}{\rho_w} - 1 \right) g d \right]^{1/2}. \quad (9)$$

```

LCM_LOOP_2:DO L=2,LA
  VELMAG(L) =SQRT(U(L,1)**2+V(L,1)**2)
  TRANS(L,K)=MAX((TAU(L)-TCRE(K))/TCRE(K),0.0)
  DZBL(L,K)=MIN(100.*HP(L),
                 D50(K)/10000.0*0.3*DSTAR(K)**0.7*SQRT(TRANS(L,K)))
  BLVEL(L,K)=1.5*TRANS(L,K)**0.6*SQRT((SSG-1)*980.0*D50(K)/10000.0)
ENDDO LCM_LOOP_2

```

After the bedload velocity magnitude (BLVEL; cm/s) is calculated, `s_bedload.f90` converts it into two orthogonal components. The equation below involves the ratios of U and VELMAG and V and VELMAG, (the *x*- and *y*-components of the hydrodynamic velocity and the velocity magnitude; m/s), respectively. Therefore, the particular units of VELMAG, U, and V are irrelevant. Ultimately, the *x*- and *y*-components of bedload velocity (UBL and VBL; m/s) are in mks units.

```

WHERE(VELMAG(2:LA)>0.0)
  UBL(2:LA,K)=BLVEL(2:LA,K)*U(2:LA,1)/VELMAG(2:LA)
  VBL(2:LA,K)=BLVEL(2:LA,K)*V(2:LA,1)/VELMAG(2:LA)
ELSEWHERE
  UBL(2:LA,K)=0.0
  VBL(2:LA,K)=0.0
ENDWHERE

```

Next, `s_bedload.f90` calculates the horizontal bedload sediment flux. Three loops in `s_bedload.f90` solve the following differential equation (Jones 2001, eq. 3.23):

$$\frac{\partial(h_b C_b)}{\partial t} = \frac{\partial q_{b,x}}{\partial x} + \frac{\partial q_{b,y}}{\partial y} + Q_b, \quad (10)$$

where C_b (g/cm³) is the bedload concentration, $q_{b,x}$ (g/cms) is the bedload sediment line flux in the *x* direction, $q_{b,y}$ (g/cms) is the bedload sediment line flux in the *y* direction, and Q_b (g/cm²s) is the sediment flux between the sediment bed and the bedload. For example, $q_{b,x} = u_{b,x} h_b C_b$, where $u_{b,x}$ is the *x*-component (analogously for *y*) of the bedload velocity from Eq. (9). The component ratio of $u_{b,x}$ to $u_{b,y}$ is equal to the ratio of U to V from EFDC hydrodynamics.

The first loop in the code solves for $q_{b,x} dt$ (XBLFLUX; g/cm³) in discretized form.

```

LCM_LOOP_3:DO L=2,LA
I=IL(L)
J=JL(L)
IF(LIJ(I-1,J)>0) THEN
  IF(UBL(L,K).GT.0) THEN
    XBLFLUX(L,K)=DT/(DXP(L)*100.0)*CBL(1,LIJ(I-1,J),K)*UBL(LIJ(I-
1,J),K)*
                                DZBL(LIJ(I-1,J),K)
  ELSE
    XBLFLUX(L,K)=DT/(DXP(L)*100.0)*CBL(1,LIJ(I,J),K)*UBL(LIJ(I,J),K)*
                                DZBL(LIJ(I,J),K)
  ENDIF
ELSE
  XBLFLUX(L,K)=0.0
ENDIF
ENDDO LCM_LOOP_3

```

The second loop solves for $q_{b,y}dt$ (YBLFLUX; g/cm^3) in discretized form.

```

LCM_LOOP_5:DO L=2,LA
I=IL(L)
J=JL(L)
IF(VBL(L,K).GE.0) THEN
  YBLFLUX(L,K)=DT/(DYP(L)*100.0)*CBL(1,LSC(L),K)*VBL(LSC(L),K)*
                                DZBL(LSC(L),K)
ELSE
  YBLFLUX(L,K)=DT/(DYP(L)*100.0)*CBL(1,LIJ(I,J),K)*VBL(LIJ(I,J),K)*
                                DZBL(LIJ(I,J),K)
ENDIF
ENDDO LCM_LOOP_5

```

where the multiplicative factor of 100 converts cm to m. After the first two terms on the right hand side of the Eq. (10) are evaluated, the bedload sediment concentration (CBL; g/cm^3) is calculated.

```

LCM_LOOP_4:DO L=2,LA
I=IL(L)
J=JL(L)
IF(BLFLAG(L,K)==1) THEN
  IF(LIJ(I+1,J)>0) THEN
    CBL(2,L,K)=CBL(1,L,K)+(XBLFLUX(L,K)-
XBLFLUX(LIJ(I+1,J),K)+YBLFLUX(L,K)-
                                YBLFLUX(LNC(L),K)+QBSSED(L,K))/DZBL(L,K)
  ELSE
    CBL(2,L,K)=CBL(1,L,K)+(YBLFLUX(L,K)-
                                YBLFLUX(LNC(L),K)+QBSSED(L,K))/DZBL(L,K)
  ENDIF
ELSE
  CBL(2,L,K)=0.0
ENDIF
IF(DZBL(L,K)==0.OR.CBL(2,L,K)<0)CBL(2,L,K)=0.0

```



```

ENDDO LCM_LOOP_4
ENDDO SED_LOOP
CBL(1,2:LA,1:NSCM)=CBL(2,2:LA,1:NSCM)

```

S SEDZLJ.f90

The `s_sedzlj.f90` routine calculates erosion and deposition from the top-most layer of the sediment bed for all cells. First, the following variables are initialized: total erosion from each cell (ETOTO; g/cm²), total deposition to each cell (DEP; g/cm²), water column depth (HP; m) is converted to cm and renamed (HT; cm), suspended sediment deposition (D; g/cm²), bedload deposition (DBL; g/cm²), total erosion of each cell (E; g/cm²), total erosion of the top layer (ELAY; g/cm²), and sediment mass available for deposition (SMASS; g/cm²).

```

ETOTO(L)=0.0
DEP(L)=0.0
HT(L)=HP(L)*100.0
D(1:NSCM,L)=0.0
DBL(1:NSCM,L)=0.0
E(1:NSCM,L)=0.0
ELAY(1:NSCM)=0.0
SMASS(1:NSCM)=0.0

```

After initializing erosion and deposition variables, `s_sedzlj.f90` calculates the near-bed concentration (CTB; g/cm³). If there is only one layer in the water column, the near-bed concentration is calculated with the approximation that the suspended sediment concentration has the following profile (Jones 2001, eq. 3.12):

$$C_s(z) = C_0 e^{-\frac{w_s z}{D_v}}, \quad (11)$$

where $C_s(z)$ (g/cm³) is the suspended concentration as a function of water depth, C_0 (CTB; g/cm³) is the near-bed concentration, w_s is the settling speed, and D_v is the vertical dispersivity. For single water layer systems, the vertical dispersivity (D_v ; cm²/s) is calculated from an empirical relation for rivers (Fischer et al. 1979) as:

$$D_v = 0.067 h u^*, \quad (12)$$

where h (cm) is the water depth. In the code below, $VZDIF$ (cm²/s) is the vertical dispersivity, HT (cm) is h , and $USTAR$ is u^* . A physically reasonable lower limit of 20.0 cm²/s is hardwired for the vertical dispersivity. The average suspended sediment concentration (C_{avg} or SED; g/cm³) is assumed to be the vertically-integrated average concentration. Based on the vertical diffusion, `s_sedzlj.f90` calculates the near-bed concentration C_0 , which is solved according to

$$C_{avg} = \int_0^h C_0 e^{-\frac{w_s z}{D_v}} dz \Rightarrow C_0 = C_{avg} \left(-\frac{w_s h}{D_v} \right) \left(\frac{1}{1 - e^{-\frac{w_s h}{D_v}}} \right). \quad (13)$$

If there is more than one water layer in the water column, the near-bed concentration is assumed to be the cell-centered average sediment concentration of the first water layer (SED(L,1,N); g/cm³).

```

IF (KC == 1) THEN
  DO K=1,NSCM
    VZDIF(L)=MAX(20.0,0.067*HT(L)*USTAR(L)*100.)
    TEMP2=HT(L)*DWS(K)/VZDIF(L)
    CTB(K,L)=SED(L,1,K)*TEMP2*(1.0/(1.0-EXP(-TEMP2)))*1.0E-06
  ENDDO
ELSE
  CTB(1:NSCM,L)=SED(L,1,1:NSCM)*1.0E-06
ENDIF

```

The temporary thickness (TTEMP; g/cm²) of each sediment size class in the top layer is calculated as the product of the percentage of each sediment size class in the top or active layer (PER) and the thickness of the top bed layer (TSED).

```
TTEMP(1:NSCM,L)=PER(1:NSCM,1,L)*TSED(1,L)
```

With the temporary thickness, `s_sedzlj.f90` calculates deposition from suspended load, first checking if the near-bed concentration for a particular size class is zero, i.e., $CTB(K,L) < 1.0E-10$. The probability of deposition from suspended load is a function of shear stress (TAU), critical shear stress for suspension (TCRSUS) from the SEDflume input files, and probability calculations from Gessler (1967) and Krone (1962).

The probability of deposition (P) of cohesive, sub-200- μm , suspended sediment particles is (Krone 1962)

$$P = \begin{cases} 0 & \text{for } \tau > \tau_{cs} \\ 1 - \frac{\tau}{\tau_{cs}} & \text{for } \tau \leq \tau_{cs} \end{cases} \quad (14)$$

For sediment particles with a diameter greater than 200 μm (non-cohesive), the probability of deposition is assumed to follow a Gaussian probability density function (Gessler 1967)

$$P(Y) = \text{erf}\left(\frac{Y}{2}\right), \quad (15)$$

where

$$Y = \frac{1}{\sigma} \left(\frac{\tau_{cs}}{\tau} - 1 \right). \quad (16)$$

The error function is approximated as (Abramowitz and Stegun 1972):

$$P \approx 1 - \exp\left[-(Y/2)^2\right] (0.34802X - 0.09587X^2 + 0.74785X^3), \quad (17)$$

where $X = 1/(1+0.235235Y)$. When $Y < 0$,

$$P(Y) = 1 - P(|Y|). \quad (18)$$

In the lines of code below, P is PROB, and the index K refers to the size class.

```

DO K=1,NSCM
  IF(CTB(K,L)<1.0E-10)CYCLE
  IF(D50(K)>=200.0) THEN
    PY(K)=ABS(1.7544*(TCRSUS(K)/TAU(L)-1.0))
    PFY(K)=EXP(-0.25*PY(K)*PY(K))
    PX(K)=1.0/(1.0+0.235235*PY(K))
    PROB(K)=PFY(K)*(0.34802*PX(K)-0.09587*PX(K)*PX(K) +
0.74785*PX(K)**3)
    IF(TAU(L)<=TCRSUS(K))PROB(K)=1.0-PROB(K)
  ELSEIF(TAU(L)<=TCRSUS(K)) THEN
    PROB(K)=1.0-TAU(L)/(TCRSUS(K))
  ELSE
    PROB(K)=0.0
  ENDIF

```

All sediment is deposited from the bottom water layer. DZC is the water-layer thickness weight-factor. The product of DZC and HT is the thickness of the each water layer in cm. This product is multiplied with the sediment concentration in the bottom water layer (SED(K,1,L)) to find the sediment mass available in each cell per unit area (SMASS). MAXDEPLIMIT is the maximum fraction of sediment from the bottom water layer allowed to deposit onto the bed per time step. This fraction is specified in `bed.sdf` and is usually set to 1.0 unless instability results (i.e., negative sediment concentrations).

```
SMASS(K)=SED(K,1,L)*DZC(1)*HT(L)*MAXDEPLIMIT
```

Suspended sediment deposition, D, is calculated as the product of probability of deposition (PROB), near-bed concentration (CTB), settling speed (DWS) and the time step (DT; s). Once the suspended sediment deposition is calculated, `s_sedzlj.f90` ensures that the deposited material is less than or equal to the total mass in the first water layer.

```

D(K,L)=PROB(K)*(DWS(K)*DT)*CTB(K,L)
ENDDO
D(1:NSCM,L)=MIN(MAX(D(1:NSCM,L),0.0),SMASS(1:NSCM))

```

Next, `s_sedzlj.f90` calculates bedload deposition fraction. Bedload deposition occurs when:

1. Bedload concentration (CBL) exists for a particular size class,
2. Bedload velocity (BLVEL) is greater than zero, and
3. Bedload transport flag (BLFLAG) is turned on.

The bedload equilibrium concentration is calculated according to van Rijn (1984b, eqn. 21):

$$C_{eq} = 0.18 \frac{C_0 \tau_*}{d_*}, \quad (19)$$

where C_{eq} is the equilibrium concentration (CSEDVR; g/cm³), C_0 is the maximum concentration (empirically determined to be 0.65), τ_* is the transport parameter (TRANS) from Eq. (7), and d_* is the dimensionless particle diameter from Eq. (4). In SEDZLJ, deposition is the product of probability of deposition, the settling velocity, and the concentration at the bed:

$$D = Pw_s C, \quad (20)$$

where D (g/cm²s) is the deposition rate and P is the probability of deposition (in the case of dynamic equilibrium, P is also the probability of erosion).

In dynamic equilibrium where erosion and deposition rates are the same,

$$E = Pw_s C_{eq}, \quad (21)$$

where E (g/cm²s) is the erosion rate.

From the equilibrium concentration, the probability of deposition is calculated as:

$$P = \frac{E_b}{w_s C_{eq}}, \quad (22)$$

where P is the van Rijn probability of deposition (PROBVR), E_b is erosion into bedload (EBL; g/cm²), w_s is the settling velocity (DWS), and C_{eq} is the equilibrium bedload concentration (CSEDVR). Note that in the equation below for the amount of sediment eroded into bedload (CSEDSS; g/cm³), DWS is multiplied by the time step DT to calculate the amount of sediment eroded at each time step. The value of EBL is the total mass eroded for that particular time step. Hence, to obtain the erosion rate E_b , EBL must be divided by DT.

If there is no equilibrium bedload concentration ($CSEDVR \leq 0.0$), then all sediment in bedload deposits onto the bed (PROBVR = 1.0). If an equilibrium concentration exists ($CSEDVR > 0$), then the probability of deposition is the fraction of sediment eroded into bedload over the equilibrium concentration of bedload sediment (CSEDSS/CSEDVR).

DO K=1, NSCM

```
IF (CBL(1,L,K) > 1.0E-10 .AND. BLVEL(L,K) > 0.0 .AND. BLFLAG(L,K) == 1) THEN
  CSEDSS(K) = EBL(L) / (DWS(K) * DT)
  CSEDVR(K,L) = (0.18 * 0.65 * TRANS(L,K) / DISTAR(K))
  IF (CSEDVR(K,L) <= 0.0) THEN
    PROBVR(K,L) = 1.0
  ELSE
    PROBVR(K,L) = MIN(CSEDSS(K) / CSEDVR(K,L), 1.0)
ENDIF
```

In the case when the sediment eroded into bedload is less than zero ($CSEDSS < 0.0$), probability of deposition reverts back to Gessler's formulation (1967) (PROBVR = PROB). This simulates the case when there is only deposition (no erosion).

```
IF (CSEDSS(K) <= 0.0) PROBVR(K,L) = PROB(K)
```

Once the probability of bedload deposition is calculated, `s_sedzlj.f90` calculates the local bedload mass available for deposition, SMASS, which is a function of bedload concentration (CBL) and saltation height (DZBL; cm). Bedload sediment deposition (DBL) is a function of probability of bedload deposition (PROBVR), bedload concentration (CBL), settling speed

(DWS), and time step (DT). The routine ensures that bedload sediment deposited onto the bed does not exceed total mass in the bedload (SMASS).

```
SMASS(K)=CBL(1,L,K)*DZBL(L,K)
DBL(K,L)=PROBVR(K,L)*CBL(1,L,K)*(DWS(K)*DT)
```

ENDIF

ENDDO

```
DBL(1:NSCM,L)=MIN(MAX(DBL(1:NSCM,L),0.0),SMASS(1:NSCM))
```

Total sediment deposited is the sum of deposition from the bedload (DBL) and suspended load (D). The deposition rate (DEPO; g/cm²s) is the total sediment deposited (DEP; g/cm²) divided by the time step (DT). Once deposition is calculated for each time step, erosion is calculated. First, local deposition is added to the active layer (TSED(1,L)) is modified according to the amount of deposited materials. Mass fractions of each sediment size class are calculated for the active layer after deposition from suspended load or from suspended load and bedload.

```
DEP(L)=SUM(DBL(1:NSCM,L),BLVEL(L,1:NSCM)>0.0)+SUM(D(1:NSCM,L))
DEPO(L)=DEP(L)/(DT)
```

Once deposition is calculated for each time step, erosion is calculated. First, local deposition is checked (DEP(L) > 0.0). Then, top bed sediment layer thickness (TSED(1,L)) is modified according to the amount of deposited materials.

```
IF(DEP(L)>0.0) THEN
  LAYER(1,L)=1
  TSED(1,L)=TSED(1,L)+DEP(L)
```

If there is bedload transport (i.e., BLVEL > 0), the mass fraction (PER) is updated by dividing the sum of original mass in the first layer (TTEMP), mass deposited from the bedload (DBL), and suspended load (D) by the newly calculated top bed layer thickness (TSED). If there is no bedload transport, only the original mass and suspended load deposition are summed.

```
WHERE(BLVEL(L,1:NSCM)>0.0)
  PER(1:NSCM,1,L)=(TTEMP(1:NSCM,L)+DBL(1:NSCM,L)
    +D(1:NSCM,L))/(TSED(1,L))
ELSEWHERE
  PER(1:NSCM,1,L)=(TTEMP(1:NSCM,L)+D(1:NSCM,L))/(TSED(1,L))
ENDWHERE
ENDIF
```

In `s_sedic.f90`, at least three bed layers must exist to account for deposition and erosion calculations. Because one bed layer may erode entirely, the following code determines which bed layer is the active layer and which layer is below the active layer. Finding the bed layer directly below the active layer is necessary because there are cases where sediment bed layers completely erode. If the top layer from the last time step is not the active layer (it has eroded away), then `s_sedzlj.f90` finds the active layer (assigning `SLLN = LL`) and the layer below it. For example, if there are five bed layers, then the top two layers initially have zero thicknesses and they are always active (provided there is deposition or armoring), otherwise their thickness remains zero. The third, fourth, and fifth layers will have initial thicknesses determined by data for the original sediment bed from `erate.sdf`. When there is deposition, the second layer will exist (i.e., `LAYER(2,L)=1`). If there is no deposition (i.e., `DO L=3,KB`), the *erodible* layer is

the third layer. If all the sediment in the third layer has eroded away in a cell, the code will re-assign the fourth layer as the erodible layer in that cell.

```

IF (LAYER( 2 , L ) == 1 ) THEN
    SLLN( L ) = 2
ELSE
    DO LL = 3 , KB
        IF ( LAYER( LL , L ) == 1 .AND. LAYER( LL-1 , L ) /= 1 ) THEN
            SLLN( L ) = LL
            EXIT
        ENDIF
    ENDDO
ENDIF

```

Next, s_sedzlj.f90 assigns the active layer (SURFACE).

```

IF (LAYER( 1 , L ) == 1 ) THEN !is the top layer present?
    SURFACE = 1 !surface variable established
ELSE
    SURFACE = SLLN( L ) !otherwise the surface variable is SLLN
ENDIF

```

As a result of sediment deposition and erosion in the active layer, the average particle size (D50AVG; μm) of the top bed layer is updated using the updated mass fractions (PER). The equivalent variable in EFDC, SEDDIA50, is also updated.

```

D50AVG( L ) = SUM( PER( 1 : NSCM , SURFACE , L ) * D50( 1 : NSCM ) )
FORALL ( LL = 1 : KB ) SEDDIA50( L , LL ) = SUM( PER( 1 : NSCM , LL , L ) * D50( 1 : NSCM ) )

```

If the active layer (SURFACE = 1) has a non-zero thickness, then the critical shear stress of the bed is interpolated from the interpolant size classes (SCLOC) and the critical shear for the corresponding interpolant size class (TAUCRITE). For example, if the average diameter (D50AVG) falls between the two interpolant size classes SCLOC(1) and SCLOC(2), where TAUCRITE(2) > TAUCRITE(1), then the critical shear stress TAUCRIT for the updated bed with an average particle size D50AVG is linearly interpolated using TAUCRITE(1) and TAUCRITE(2). In the case where the top layer is not layer 1 or 2 (the cell is purely erosive), critical shear is set equal to the critical shear in of the original bed (TAUCOR). There is no need for interpolation here because the critical shear stress data for the original bed (TAUCOR) is used (i.e., data from SEDflume).

```

DO K = 1 , NSICM-1
    IF ( D50AVG( L ) >= SCLOC( K ) .AND. D50AVG( L ) < SCLOC( K+1 ) ) THEN
        NSCD( 1 ) = SCLOC( K )
        NSCD( 2 ) = SCLOC( K+1 )
        NSC0 = K
        NSC1 = K+1
        EXIT
    END IF
ENDDO
IF ( SURFACE == 1 .OR. SURFACE == 2 ) THEN

```

```

    TAUCRIT(L)=TAUCRITE(NSC0)+(TAUCRITE(NSC1)-TAUCRITE(NSC0))/
    (NSCD(2)-NSCD(1))*(D50AVG(L)-NSCD(1))
ELSE
    TAUCRIT(L)=TAUCOR(SURFACE,L)
ENDIF

```

The active layer thickness is defined as the depth into the sediment bed (depth below water-sediment interface) that is influenced by the water shear stress (penetration depth). In the case of bed armoring, the larger sediment do not move while smaller particles within the interstices of the larger sediment are sheared away into the water column down to a certain thickness (the active layer thickness). The active layer is re-calculated according to the bed armoring formula (Van Niekerk et al. 1992):

$$T_a = \alpha d_{50} \frac{\tau}{\tau_{ce}}, \quad (23)$$

where T_a (cm) is the active layer's thickness (TACT), $1.0 < \alpha < 10.0$ is the active layer thickness multiplier (TACTM), d_{50} (μm) is the active layer averaged particle diameter (D50AVG), τ (dynes/cm^2) is the shear stress at the bed (TAU), and τ_{ce} (dynes/cm^2) is the critical shear stress (TAUCRIT). Note: the units of TACT in the code is (g/cm^2) instead of cm like T_a . If the shear stress (TAU) is less than the critical shear stress (TAUCRIT), the active layer thickness (TACT) is proportional to the ratio of shear stress to critical shear stress (i.e., TAU/TAUCRIT). However, if the shear stress is greater than the critical shear stress, the active layer thickness (TACT) is only a function of average particle diameter (D50AVG). Note that the density (BULKDENS) appears in the equation for TACT because TACT has units of g/cm^2 .

```

IF (TAU(L)/TAUCRIT(L) < 1.0) THEN
    TACT(L) = 2.0 * D50AVG(L) * (BULKDENS(1,L) / 10000.0)
ELSE
    TACT(L) = 2.0 * D50AVG(L) * (TAU(L)/TAUCRIT(L)) * (BULKDENS(1,L) / 10000.0)
ENDIF

```

s_sedzlj.f90 checks for the existence of an active layer, which happens when there is armoring or deposition, but not when the model cell is purely erosive.

```

NACTLAY=0
DO LL=1,KB
    IF (LAYER(LL,L) > 0) THEN
DO K=1,NSCM
    IF (PER(K,LL,L) > 0.0 .AND. TAU(L) < TCRE(K) .AND. TAU(L) > TAUCRIT(L)) THEN
        NACTLAY=1
        LAYER(1,L)=1
    ENDIF
ENDDO
EXIT
ENDIF
ENDDO

```

When material is deposited onto (or eroded away from) the top layer, the thickness of the top layer is adjusted back to the active layer thickness, defined by Eq. (23), according to the following three scenarios:

1. If there is a net deposition and the top layer thickness TSED(1,L) is greater than the calculated active layer TACT(L), then mass from the top layer is transferred to the layer below it (i.e., the depositional layer). The amount of mass transferred is TSED(1,L) – TACT(L). Once this is done, the mass fraction (PER) for both layers is re-calculated.

```

IF (TSED(1,L)>0.0.OR.NACTLAY/=0) THEN
  IF (LAYER(1,L)==1) THEN
    IF (TSED(1,L)>TACT(L)) THEN
      FORALL (K=1:NSCM) PER(K,2,L)=(PER(K,2,L)*TSED(2,L)
                                     +PER(K,1,L)*(TSED(1,L)
                                     -TACT(L)))/(TSED(2,L)
                                     +(TSED(1,L)-TACT(L)))
      LAYER(2,L)=1
      SLLN(L)=2
      TSED(2,L)=TSED(2,L)+TSED(1,L)-TACT(L)
      TSED(1,L)=TACT(L)

```

2. When deposition occurs, but the net sediment flux is erosional, the new active layer thickness TACT(L) will be greater than the active layer thickness TSED(1,L). In this situation, mass from the layer below the active layer will be added to the top layer, and the new thickness TSED(1,L) will be TSED(1,L) + TACT(L) – TSED(1,L). The layer below is either from the newly deposited layer or from deeper layers with properties defined in `erate.sdf`. Mass fractions (PER) of both layers are then re-calculated.

```

ELSEIF (TSED(1,L)<TACT(L).AND.TSED(1,L)+TSED(SLLN(L),L)>TACT(L)
        .AND. TAU(L)>TAUCOR(SLLN(L),L)) THEN
  FORALL (K=1:NSCM) PER(K,1,L)=(PER(K,1,L)*TSED(1,L)+PER(K,SLLN(L),L)
    *(TACT(L)-TSED(1,L)))/TACT(L)
    TSED(SLLN(L),L)=TSED(SLLN(L),L)-(TACT(L)-TSED(1,L))
    TSED(1,L)=TACT(L)

```

3. If there is net erosion but the layer immediately below the top layer does not have enough mass to contribute to the active layer, `s_sedzlj.f90` will incorporate all of the material in the present erodible layer to the active layer in this time step and re-assign the erodible layer. However, the actual active layer thickness might not be the same as the active layer thickness requested by Eq. (23).

```

ELSEIF (TSED(1,L)<TACT(L).AND.TSED(1,L)+TSED(SLLN(L),L)<=TACT(L).AND.
        TAU(L)>TAUCOR(SLLN(L),L)) THEN
  FORALL (K=1:NSCM)
    PER(K,1,L)=(PER(K,1,L)*TSED(1,L)+PER(K,SLLN(L),L)*(TSED(SLLN(L),L)))/
      (TSED(1,L)+TSED(SLLN(L),L))
    PER(K,SLLN(L),L)=0.0
  ENDFORALL
  TSED(1,L)=TSED(1,L)+TSED(SLLN(L),L) !add available sediment to
layer
  TSED(SLLN(L),L)=0.0 !no more sediment available below this

```



```

        LAYER(SLLN(L),L)=0 !layer has been eliminated in the logical
variable
        SLLN(L)=SLLN(L)+1 !set next layer lower
        IF (SLLN(L)+1>KB) SLLN(L)=KB !do not allow specification of the
next lower layer to be below the bottom sediment layer
    ENDIF
ENDIF
ENDIF

```

After the sorting process is completed, s_sedzlj.f90 calculates erosion. If a particular layer has eroded or if the cell is in a state of deposition only, s_sedzlj.f90 identifies which sediment layers are intact before re-formulating bed properties.

```

ALL_LAYERS:DO LL=1,KB !go through all sediment layers so that they are
properly eroded and sorted
IF(LAYER(LL,L)==0)CYCLE !if the layer is gone don't consider it
IF(LAYER(1,L)==1.AND.LL/=1)EXIT !if it is depositional, there is no
need to consider erosion
IF(LL/=1) THEN
    IF(LAYER(LL-1,L)==1)EXIT
ENDIF

```

Before erosion calculations, the active (top) layer is re-characterized because the sediment in the active layer may have been adjusted. Average particle size (D50AVG; μm) of the layer is calculated using size classes (D50; μm) and weight fractions (PER). From D50AVG, critical shear stress (TAUCRIT) for the active layer are linearly interpolated from interpolant size class (SCLOC) and the corresponding critical shear for each interpolant (TAUCRITE). In the event that all the sediment bed layers are eroded away, the critical shear stress is set to 10^6 dynes/cm², so that there is no further erosion.

```

D50AVG(L)=SUM(PER(1:NSCM,LL,L)*D50(1:NSCM))
DO K=1,NSICM-1
    IF(D50AVG(L)>=SCLOC(K).AND.D50AVG(L)<SCLOC(K+1)) THEN
        NSCD(1)=SCLOC(K)
        NSCD(2)=SCLOC(K+1)
        NSC0=K
        NSC1=K+1
        EXIT
    ENDIF
ENDDO
IF(LL==1.OR.LL==2) THEN !for active layer
    TAUCRIT(L)=TAUCRITE(NSC0)+(TAUCRITE(NSC1)-TAUCRITE(NSC0))/(NSCD(2)
        -NSCD(1))*(D50AVG(L)-NSCD(1)) !interpolation
    TAUCOR(LL,L)=TAUCRIT(L)
ELSE
    SN01=TSED(LL,L)/TSED0(LL,L) !weighting factor 1 for interpolation
    SN11=(TSED0(LL,L)-TSED(LL,L))/TSED0(LL,L) !weighting factor 2
    IF(LL+1<=KB) THEN !Avoid array exceedance when LL=KB and used in
TAUCOR(LL+1,L)
        TAUCRIT(L)=SN01*TAUCOR(LL,L)+SN11*TAUCOR(LL+1,L) !
    ELSE
        TAUCRIT(L)=1.0e6 !set arbitrarily high
    ENDIF
ENDIF

```

ENDIF
ENDIF

The total erosion rate is interpolated across layer thicknesses and shear stresses. Linear interpolation is used to calculate the erosion rate at a specified shear stress (Jones 2001, eq. 3.3):

$$E(\tau) = \left(\frac{\tau_{i+1} - \tau}{\tau_{i+1} - \tau_m} \right) E_m + \left(\frac{\tau - \tau_i}{\tau_{i+1} - \tau_i} \right) E_{i+1}. \quad (24)$$

The subscripts i indicate data for a shear stress less than τ and $(i + 1)$ indicates a shear stress larger than τ . For example, consider a newly deposited sediment layer having a D50AVG of 150 μm . If the model calculates a shear stress (TAU) of 30.0 dynes/cm² (3.0 Pa), then the erosion rate (ERATEMOD) must be somewhere between the appropriate four values in bed.sdf (Figure 3: 6.60E-5, 4.66E-4, 5.97E-5, and 5.97E-4 cm/s).

Because erosion rate varies rapidly with flow condition and bed properties, it is logarithmically interpolated from SEDflume data as a function of depth (Jones 2001, eq. 3.4):

$$\ln[E(T)] = \left(\frac{T_0 - T}{T_0} \right) \ln(E^j) + \frac{T}{T_0} \ln(E^{j+1}). \quad (25)$$

where T (m) is the actual layer thickness, T_0 (m) is the initial layer thickness, the superscripts j and $j+1$ denote data for the interface at the top and the bottom of the appropriate layer, respectively.

There are two kinds of erosion rate (ERATEMOD) calculations: (i) the initial bed erosion rate and (ii) the erosion rate of newly deposited materials. For the calculation of initial bed erosion rate (not newly deposited), s_sedzlj.f90 first identifies the shear stress with respect to the erosion rate data table from erate.sdf in Figure 1, by comparing the actual shear stress (TAU) with the shear stresses (TAULOC) at which erosion rate data (ERATE) are measured. When the closest value of TAULOC to TAU is found, s_sedzlj.f90 calculates the modified erosion rate (ERATEMOD; cm/s) using the log-linear interpolation scheme described above. In cases where all of the sediment in the last layer is eroded away, ERATE is set to 10⁻⁹ cm/s and ERATEMOD = 0.

```
IF (TAU(L) < TAUCRIT(L)) EXIT
!Find the upper and lower limits of the Shear Stress for the
interpolation
DO K=1, ITBM-1
  IF (TAU(L) >= TAULOC(K) .AND. TAU(L) < TAULOC(K+1)) THEN
    TAUDD(1) = TAULOC(K)
    TAUDD(2) = TAULOC(K+1)
    NTAU0 = K
    NTAU1 = K+1
  EXIT
ENDIF
ENDDO
!Interpolate the erosion rates for shear stress and depth.
!This utilizes normal SEDflume data for deeper layers.
IF (LL > 2) THEN !calculate erosion rates of deeper layers
  SN00 = (TAUDD(2) - TAU(L)) / (TAUDD(2) - TAUDD(1)) !weighting factor 1 for
```

```

interpolation
  SN10=(TAUDD(1)-TAU(L))/(TAUDD(1)-TAUDD(2)) !weighting factor 2
  SN01=TSDD(LL,L)/TSDD(LL,L) !weighting factor 3
  SN11=(TSDD(LL,L)-TSDD(LL,L))/TSDD(LL,L) !weighting factor 4
  IF(LL+1<=KB) THEN !modeled erosion rate

ERATEMOD(L)=(SN00*EXP(SN11*LOG(ERATE(LL+1,L,NTAU0)))+SN01*LOG(ERATE(LL,
L,NTAU0)))
  &
+SN10*EXP(SN11*LOG(ERATE(LL+1,L,NTAU1)))+SN01*LOG(ERATE(LL,L,NTAU1))) *
BULKDENS(LL,L)
  ELSE !do not allow erosion through the bottom layer
    ERATEMOD(L)=(SN00*EXP(SN11*LOG(1e-9))+SN01*LOG(ERATE(LL,L,NTAU0)))
    & +SN10*EXP(SN11*LOG(1e-
9))+SN01*LOG(ERATE(LL,L,NTAU1))) *BULKDENS(LL,L)
  ENDIF

```

For erosion of newly deposited sediments (with layers 1 and 2 being the active depositional layers, respectively), erosion rate (ENRATE) from bed.sdf in Figure 3 are interpolated to calculate the erosion rate (ERATEMOD).

```

ELSE
  NSCTOT=NSCD(2)-NSCD(1)
  D50TMPP=D50AVG(L)-NSCD(1)
  SN00=(TAUDD(2)-TAU(L))/(TAUDD(2)-TAUDD(1))
  SN10=(TAUDD(1)-TAU(L))/(TAUDD(1)-TAUDD(2))
  SN01=D50TMPP/NSCTOT
  SN11=(NSCTOT-D50TMPP)/NSCTOT
  ERATEMOD(L)=(SN00*EXP(SN11*LOG(ENRATE(NSC0,NTAU0))
+SN01*LOG(ENRATE(NSC1,NTAU0)))
+SN10*EXP(SN11*LOG(ENRATE(NSC0,NTAU1))
+SN01*LOG(ENRATE(NSC1,NTAU1))) *BULKDENS(LL,L)
ENDIF

```

Once the erosion rate (ERATEMOD) is interpolated based on the layer-averaged particle size (D50AVG) in the cell, the amount of sediment eroded per time step (EBD; g/cm²) is calculated by multiplying the erosion rate (ERATEMOD) by the time step (DT).

```
EBD(LL,L)=ERATEMOD(L)*DT
```

The following variables are calculated from the total sediment eroded per time step (EBD): total erosion at each cell for each size class (E) and total erosion per layer for each size class (ELAY). Top layer thickness (TTEMP) is also re-calculated. If shear stress (TAU) is less than critical shear (TCRE) then all erosion values (E and ELAY) are set to zero and top layer thickness (TTEMP) is calculated without erosion (ELAY).

```

WHERE(TAU(L)>=TCRE(1:NSCM))
  E(1:NSCM,L)=E(1:NSCM,L)+PER(1:NSCM,LL,L)*EBD(LL,L)
  ELAY(1:NSCM)=PER(1:NSCM,LL,L)*EBD(LL,L)
  TTEMP(1:NSCM,L)=PER(1:NSCM,LL,L)*TSDD(LL,L)-ELAY(1:NSCM)
ELSEWHERE
  E(1:NSCM,L)=0.0
  ELAY(1:NSCM)=0.0

```

```
TTEMP(1:NSCM,L)=PER(1:NSCM,LL,L)*TSED(LL,L)
```

```
ENDWHERE
```

If amount of sediment eroded is greater than thickness of the bed layer (i.e., $TTEMP < 0$ and a bed layer has been completely eroded during this time step), thickness of the top layer (TTEMP) is set to zero and erosion at each cell for each size class (E) is set equal to the amount of sediment available in the eroding bed layer. Furthermore, sediment eroded from the layer (ELAY) is set equal the remaining layer thickness ($PER * TSED$). The total sediments eroded from the active (top) layer (ESED) and per cell (ETOTO) are calculated in units of g/cm^2 .

```
WHERE (TTEMP(1:NSCM,L) < 0.0)
```

```
TTEMP(1:NSCM,L)=0.0
```

```
E(1:NSCM,L)=E(1:NSCM,L)-PER(1:NSCM,LL,L)*EBD(LL,L)
+PER(1:NSCM,LL,L)*TSED(LL,L)
```

```
ELAY(1:NSCM)=PER(1:NSCM,LL,L)*TSED(LL,L)
```

```
ENDWHERE
```

The total sediments eroded from the top (active) layer (ESED) and per cell (ETOTO) are calculated in units of g/cm^2 .

```
ESED=SUM(ELAY(1:NSCM))
```

```
ETOTO(L)=SUM(E(1:NSCM,L))
```

Once erosion is calculated, thickness of the top layer (TSED) and mass fractions (PER) is updated. If active layer thickness is less than 10^{-3} cm, TSED becomes zero and the top layer is completely eroded (i.e., $LAYER(LL,L) = 0$). Otherwise, top layer thickness is set equal to calculated layer thickness after erosion (TEMP; cm), and the mass fraction (PER) is also updated accordingly.

```
TEMP=TSED(LL,L)-ESED
```

```
IF (TEMP <= 1.0E-3) THEN
```

```
TSED(LL,L)=0.0
```

```
LAYER(LL,L)=0
```

```
FORALL (K=1:NSCM) PER(K,LL,L)=0.0
```

```
ELSE
```

```
TSED(LL,L)=TEMP
```

```
FORALL (K=1:NSCM) PER(K,LL,L)=TTEMP(K,L)/TSED(LL,L)
```

```
ENDIF
```

If a change in morphology is considered (i.e., $IMORPH = 0$), then height of the bed (HBED; m) is calculated using the updated layer thicknesses (TSED).

```
IF (IMORPH==0) HBED(L,LL)=0.01*TSED(LL,L)/BULKDENS(LL,L)
```

```
ENDDO ALL_LAYERS
```

Once the total erosion is calculated, sediment bed fluxes (QBSSED and BED_SED_FLX; g/cm^2) are divided into bedload (EBL) and suspended load (ESUS; g/cm^2) using the probability for suspension (PSUS). When bedload transport is inactive ($BLFLAG = 0$), only suspended sediment flux (BED_SED_FLX) is calculated.

```
EBL(L)=0.0
```

```
ESUS(L)=0.0
```

```
DO K=1,NSCM
```

```
IF (BLFLAG(L,K)==1) THEN
```

```
QBSSED(L,K)=(1.0-PSUS(L,K))*E(K,L)-DBL(K,L)
```

```

BED_SED_FLX(L,K)=PSUS(L,K)*E(K,L)-D(K,L)
ELSE
BED_SED_FLX(L,K)=E(K,L)-D(K,L)
ENDIF
ENDDO
EBL(L)=SUM((1.0-PSUS(L,1:NSCM))*E(1:NSCM,L),BLVEL(L,1:NSCM)>0.0)
ESUS(L)=SUM(PSUS(L,1:NSCM)*E(1:NSCM,L),BLVEL(L,1:NSCM)>0)

```

After calculating erosion information, `s_sedzlj.f90` assigns EFDC variables. This is for backward compatibility with EFDC and the visualization program EFDC Explorer (Craig 2008). ETOTO is the total amount of eroded sediment. To convert this quantity into an erosion rate, `s_sedzlj.f90` divides ETOTO by the time step (DT).

```

!Flux calculations for EFDC
!Convert BED_SED_FLX from g/cm^2 to g/m^2*s
WDTDZ=DT*HPI(L)*DZIC(1)
SEDF(L,0,1:NSCM)=BED_SED_FLX(L,1:NSCM)*10000.0/DT
SED(L,1,1:NSCM)=SEDS(L,1,1:NSCM)+(SEDF(L,0,1:NSCM)-
SEDF(L,1,1:NSCM))*WDTDZ
ETOTO(L)=ETOTO(L)/(DT)

```

S MORPH.f90

Erosion and deposition may change the geometry of the sediment bed, which may in turn affect the hydrodynamics. Layer thickness (TSED), bulk density (BULKDENS), and bed height from the previous time step (HBEDA; m) are used to calculate change in bed height (DELBED; m).

BELV (m) defines the position of the sediment-bed/water-column interface for the current time step, HP is the water-column height for the current time step, P (m^2/s^2) is the product of water surface elevation and acceleration due to gravity, and HBEDA is the total bed height. The position of the sediment-bed/water-column interface (BELV) and the water column depth (HP) vary according to the change in bed height (DELBED). Note that the corresponding variables with values from the previous time step are BELV1, H1P, and P1.

```

DO L=2,LA
DELBED(L)=0.01*SUM(TSED(1:KB,L)/BULKDENS(1:KB,L))-HBEDA(L)
BELV1(L)=BELV(L)
HTMP(L)=HP(L)
H1P(L)=HP(L)
P1(L)=P(L)
HBEDA(L)=0.01*SUM(TSED(1:KB,L)/BULKDENS(1:KB,L))
HBED(L,1:KB)=0.01*TSED(1:KB,L)/BULKDENS(1:KB,L)
BELV(L)=ZELBEDA(L)+HBEDA(L)
HP(L)=HP(L)+DELBED(L)
ENDDO

```

The inverse of the water depth (HPI; 1/m) and the water flux (QMORPH; kg/m^2s) from the eroded sediment bed to the water column are also calculated using the current column height (HP) and previous water column heights (H1P).

```

DO L=2,LA
  HPI(L)=1.0/HP(L)
  QMORPH(L)=DELTI*DXYP(L)*(HP(L)-H1P(L))
ENDDO

```

s_morph.f90 also writes an output file for drying cells and error checking.

```

ITMP=0
DO L=2,LA
  IF(HP(L)<0.0) THEN
    IF(ABS(H1P(L))>=HWET) THEN
      ITMP=1
      WRITE(8, "('NEG DEPTH DUE TO MORPH
CHANGE',2I5,12F12.5)") IL(L), JL(L),
        HBED1(L,KBT(L)), HBED(L,KBT(L)), BELV1(L), BELV(L), DELT,
        QSBDBTOP(L), QWBDBTOP(L), HBEDA(L)
      WRITE(8, "('NEG DEPTH DUE TO MORPH CHANGE',
2I5,12F12.5)") L, KBT(L), (HBED(L,K), K=1, KBT(L))
    ELSE
      HP(L)=0.9*HDRV
    ENDIF
  ENDIF
ENDDO
IF(ITMP==1) THEN
  CALL RESTOUT(1)
  IF(NDRYSTP<0) THEN
    OPEN(1, FILE='DRYLOSS.OUT')
    CLOSE(1, STATUS='DELETE')
    OPEN(1, FILE='DRYLOSS.OUT')
    DO L=2,LA
      IF(VDWASTE(L)>0.0) THEN
        TMPVAL=VDWASTE(L)/DXYP(L)

WRITE(1, '(2I6,4E14.6)') IL(L), JL(L), VDWASTE(L), TMPVAL, QDWASTE(L)
      ENDIF
    ENDDO
    CLOSE(1)
  ENDIF
STOP
ENDIF

```

Changes in morphology may cause change in water column depth. Hence, the concentrations of all transported materials must be adjusted.

```

IF(ISTRAN(1)>0) FORALL(K=1:KC, L=2:LA) SAL(L,K)=HTMP(L)*SAL(L,K)/HP(L)
IF(ISTRAN(3)>0) FORALL(K=1:KC, L=2:LA) DYE(L,K)=HTMP(L)*DYE(L,K)/HP(L)
IF(ISTRAN(4)>0) FORALL(K=1:KC, L=2:LA) SFL(L,K)=HTMP(L)*SFL(L,K)/HP(L)
IF(ISTRAN(5)>0) FORALL(NT=1:NTOX, K=1:KC, L=2:LA) TOX(L,K,NT)=HTMP(L)*TOX(
L,K,NT)/HP(L)
IF(ISTRAN(6)>0) FORALL(NS=1:NSCM, K=1:KC, L=2:LA) SED(L,K,NS)=HTMP(L)*SED(
L,K,NS)/HP(L)
IF(ISTRAN(7)>0) FORALL(NS=1:NSND, K=1:KC, L=2:LA) SND(L,K,NS)=HTMP(L)*SND(
L,K,NS)/HP(L)

```

Variable Dictionary

The table below lists the SNL-EFDC variables in alphabetical order with a brief description. Variables from EFDC are denoted as such in the description.

Variable name	Description
BEDLOAD	Flag for bedload transport (1 = yes)
BDEN	Bulk density for each sediment layer (g/cm^3)
BED_SED_FLX	Suspended sediment flux (g/cm^2)
BEDLINIT	Initial bed thickness (m)
BELV	Position of the sediment-bed water column interface for the current time step (m)
BLVEL	Bedload velocity (cm/s)
BULKDENS	Mass of solids per unit volume (g/cm^3)
CBL	Bedload concentration (g/cm^3)
CSEDVR	Van Rijn's equilibrium bedload concentration (g/cm^3)
CSEDSS	Sediment eroded into bedload (g/cm^3)
CTB	Concentration at sediment bed (g/cm^3)
D	Deposition from suspended load (g/cm^2)
D50	Particle diameter for each size class (μm)
D50AVG	Average particle size (μm)
DBL	Deposition from bedload (g/cm^2)
DELBED	Change in the sediment bed height in one time step, from EFDC (m)
DELTI	Inverse time step ($1/\text{s}$)
DEP	Total deposition of each cell (g/cm^2)
DISTAR	Non-dimensional particle diameter for each size class

Variable name	Description
DT	Time step (s)
DWS	Settling speed (cm/s)
DZBL	Bedload or saltation height (cm)
DZC	Water-layer thickness factor
E	Total erosion of each cell in one time step (g/cm^2)
EORATE	Initial erosion rate for each layer subject to TAULOC (cm/s)
EBD	Total sediment eroded in one time step (g/cm^2)
EBL	Total erosion into bedload in one time step (g/cm^2)
ELAY	Total erosion of the top layer in one time step (g/cm^2)
ENRATE	Erosion rate for newly deposited bed (cm/s)
ERATE	EORATE (cm/s)
ERATEMOD	Erosion rate for each layer ($\text{g}/\text{cm}^2\text{s}$)
ESED	Total erosion from each layer in one time step (g/cm^2)
ESUS	Total erosion into suspended load in one time step (g/cm^2)
ETOTO	Total erosion of all size classes (g/cm^2)
FC	Friction factor used to calculate the applied shear stress
HBED	Bed thickness per layer, from EFDC (m)
HBEDA	Total sediment bed thickness (cm)
HP	Water column depth (m)
HT	Water column depth (cm)
IFWAVE	Flag that activates wind-driven surface waves (1 = yes)
IMORPH	Flag that activates bed morphology calculations (1 = yes)
ISEDTIME	Number of time steps before sediment transport is activated

Variable name	Description
ISMORPH	Flag that activates bed morphology calculations (1 = yes)
ITBM	Number of different SEDflume shear-stress interpolants to calculate erosion of newly deposited sediments
IWRSP	EFDC flag that activates SEDZLJ (98 = yes)
KB	Maximum number of sediment layers
LAYER	Logical array that is 1 if a layer is present at cell, otherwise it is 0
MAXDEPLIMIT	Maximum fraction of sediment that can deposit in one time step
NCALC_BL	Flag that activates bedload sediment transport (1 = yes)
NEQUIL	Flag to activate adsorption of toxics onto sediments (1 = yes)
NSED	Number of sediment particle size classes, from EFDC
NSEDFLUME	Flag to activate SEDZLJ routines (1 or 2 = yes), set by the EFDC variable IWRSP = 98 or 99
NSCM	Number of sediment size classes, from EFDC
NSICM	Number of size classes of interpolants to estimate erosion rates of newly deposited sediments
P	Product of water surface elevation and acceleration due to gravity (m^2/s^2)
PER	P/100.0
PNEW	Mass percentage of each size class in layer
PROB	Probability of suspended load deposition
PROBVR	Probability of deposition from bedload
PSUS	Percentage of total erosion into suspended load
QBSER	Net flux into bedload in one time step (g/cm^2)
QSBDBTOP	Total volumetric sediment flux, from EFDC ($\text{kg}/\text{m}^2\text{s}$)
SCLOC	Particle size for newly deposited bed erosion data (μm)

Variable name	Description
SED	Suspended sediment concentration (kg/m^3)
SEDF	Sediment flux ($\text{kg/m}^2\text{s}$)
SEDS	SED, used for vertical turbulent dispersion and settling calculations (kg/m^3)
SEDDENS	Sediment density (g/cm^3)
SLLN	Contains the index for the active layer of each cell
SMASS	Sediment mass available for deposition (g/cm^2)
SSGI	Inverse specific gravity
SURFACE	Sediment layer that corresponds to the active layer
TACT	Active layer thickness (g/cm^2)
TACTM	Active layer thickness multiplier
TAU	Total shear stress (dynes/cm^2)
TAUB	Total shear stress, for EFDC bed stress (Pa)
TAUCONST	Parameter that, when greater than zero, sets the shear stress to a specified constant value in the flow (dynes/cm^2)
TAUCOR	Critical shear stress for each core at each sediment layer (dynes/cm^2)
TAUCR	Critical shear stress for erosion for each size class (dynes/cm^2)
TAUCRIT	Interpolated critical shear stress for erosion per size class (dynes/cm^2)
TAUCRITE	Critical shear stress for erosion of newly deposited bed per size class (dynes/cm^2)
TAULOC	Applied shear stress from erosion data (dynes/cm^2)
TAUTEMP	Critical shear stress for the erosion of each layer (dynes/cm^2)
TCRE	Critical shear stress for erosion (dynes/cm^2)
TCSRUS	Critical shear stress for suspension (dynes/cm^2)
TEMP	Calculated layer thickness after erosion (cm)

Variable name	Description
TRANS	Dimensionless transport parameter for bedload calculations (van Rijn 1984b).
TSED	Current mass in each layer (g/cm^2)
TSED0	Initial “thickness” of each sediment layer converts to (g/cm^2)
TSED0S	TSED0 (g/cm^2)
TTEMP	Temporary thickness of a given layer (g/cm^2)
U	x -component of the hydrodynamic velocity (m/s)
UBL	x -component of the bedload velocity (cm/s)
USTAR	Shear velocity (m/s)
USW	Ratio of shear velocity to settling velocity
V	y -component of the hydrodynamic velocity (m/s)
VAR_BED	Flag that activates the variable sediment bed option (1 = yes)
VBL	y -component of the bedload velocity (cm/s)
VDRBED	Initial void ratio in each bed layer
VDRBEDO	Sediment ratio in each bed layer
VELMAG	Velocity magnitude (m/s)
VZDIF	Vertical diffusion for suspended sediments (cm^2/s)
WATERDENS	Density of water (g/cm^3)
WSETA	Sediment settling velocity, from EFDC (m/s)
WVEL	Sediment vertical velocity, from EFDC (m/s)
XBLFLUX	x -component of the horizontal bedload flux in one time step (g/cm^3)
YBLFLUX	y -component of the horizontal bedload flux in one time step (g/cm^3)
ZBSKIN	Parameter that, when greater than zero, sets the sediment bed roughness to a specified value (μm)

Variable name	Description
ZBTEMP	Variable that stores the calculated roughness parameter value (μm)
ZELBEDA	Average bed elevation (cm)

Special Notes

Currently, `s_sedzlj.f90` does not incorporate bedload sediment flux through boundaries of the domain. This feature will soon be added to SNL-EFDC. Bed consolidation is also an expected upgrade.

There are a few issues in EFDC not mentioned in the user's manual that we document in this section to assist users of SNL-EFDC.

Currently the EFDC routine `s_ser.inp` divides the salinity value by the mass of each water layer. This is somewhat confusing because salinity is an intrinsic variable.

Also, the variable `QSSER` from `efdc.inp` is added to flows specified in `QSER.inp` and it is easy to neglect the fact that both flow inputs may be active in a model.

References

- Abramowitz, M., and Stegun, I. (1972). *Handbook of Mathematical Functions*, Washington, D.C.
- Cheng, N. S. (1997). "Simplified settling velocity formula for sediment particle." *Journal of Hydraulic Engineering-ASCE*, 123(2), 149-152.
- Christoffersen, J., and Jonsson, I. (1985). "Bed friction and dissipation in a combined current and wave motion." *Ocean Engineering*, 17(4), 479-494.
- Craig, P. M. (2008). "EFDC Explorer."
- Fischer, H., List, E., Koh, R., Imberger, J., and Brooks, N. (1979). *Mixing in Inland and Coastal Waters*, Academic Press, New York, NY.
- Gessler, J. (1967). "The beginning of bedload movement of mixtures investigated as natural armoring in channels."
- Guy, H. P., Simons, D. B., and Richardson, E. V. (1966). "Summary of alluvial channel data from flume experiments." *Geological Survey Professional Paper*, 462(I).
- Hamrick, J. M. (1992). "A Three-Dimensional Environmental Fluid Dynamics Computer Code: Theoretical and Computational Aspects." The College of William and Mary.
- Hamrick, J. M. (1996). "User's Manual for the Environmental Fluid Dynamics Computer Code." Virginia Institute of Marine Sciences, Gloucester Point, Virginia.
- Jones, C. (2001). "SEDZLJ: A Sediment Transport Model," University of California, Santa Barbara.
- Jones, C., and Lick, W. (Year). "Sediment erosion rates: Their measurement and use in modeling." *The Texas A&M Dredging Seminar*, College Station, Texas, 1-15.
- Krone, R. B. (1962). "Technical Report DA-04-203 CIVENG-59-2." US Army Corps of Engineers, San Francisco, CA.
- Roberts, J. D., Jepsen, R. A., and James, S. C. (2003). "Measurements of sediment erosion and transport with the Adjustable Shear Stress Erosion and Transport Flume." *Journal of Hydraulic Engineering-Asce*, 129(11), 862-871.
- Roberts, J. D., Jepsen, R. A., and Lick, W. (1998). "Effects of particle size and bulk density on erosion of quartz particles." *Journal of Hydraulic Engineering-ASCE*, 124(12), 1261-1267.
- Soulsby, R. L. (1997). *Dynamics of Marine Sands*, Thomas Telford, London, UK.
- Van Niekerk, A., Vogel, K., Slingerland, R., and Bridge, J. (1992). "Routing of heterogeneous sediments over movable bed: model development." *J. Hydr. Engr.*, 118(2), 246-263.
- van Rijn, L. C. (1984a). "Sediment transport, II: Suspended load transport." *Journal of Hydraulic Engineering-ASCE*, 110(11), 1613-1641.
- van Rijn, L. C. (1984b). "Sediment transport, Part I: Bed load transport." *Journal of Hydraulic Engineering-ASCE*, 110(11), 1613-1641.

Ziegler, C. K., Israelsson, P. H., and Connolly, J. P. (2000). "Modeling sediment transport dynamics in Thompson Island Pool, Upper Hudson River." *Water Quality and Ecosystem Modeling*, 1, 193-222.