## Un-weighted least-squares fitting in HFRPtoolbox

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## 1 Method

This is the part of this manuscript (([Kim et al.(2007)Kim, Terrill, and Cornuelle])). The weighted least-squares fitting (WLS, [Wunch(1996)]) method to estimate the vector current ( $\mathbf{u}$ ,  $2 \times 1$  elements) at a vector grid point ( $x_k, y_k$ ) assumes that each radial velocity is a linear combination of the vector current components. The radial velocities ( $\mathbf{r}$ ,  $L \times 1$  elements) within the search radius ( $d_0$ ) of the vector grid point and the projection matrix ( $\mathbf{G_a}$ ,  $L \times 2$  elements) are:

$$\mathbf{r} = \mathbf{G_a} \mathbf{u} + \mathbf{n_a} \,, \tag{1.1}$$

where

$$\mathbf{G_{a}} = \begin{bmatrix} \mathbf{g}_{1}^{\mathrm{T}} \\ \mathbf{g}_{2}^{\mathrm{T}} \\ \vdots \\ \mathbf{g}_{L}^{\mathrm{T}} \end{bmatrix} = \begin{bmatrix} \cos \theta_{1} & \sin \theta_{1} \\ \cos \theta_{2} & \sin \theta_{2} \\ \vdots & \vdots \\ \cos \theta_{L} & \sin \theta_{L} \end{bmatrix}, \mathbf{r} = \begin{bmatrix} r_{1} \\ r_{2} \\ \vdots \\ r_{L} \end{bmatrix}, \tag{1.2}$$

 $\mathbf{g}_i = [\cos \theta_i \sin \theta_i]^{\mathrm{T}}$ ,  $\theta_i$  denotes the bearing angle at the *i*th radar grid point  $(x_i, y_i)$  measured counter-clockwise from East, and  $\mathbf{n}_a$  is the residual when the radial velocities are fit by a single vector current (<sup>T</sup> denotes the matrix transpose). The least-squares estimate for the vector current is

$$\hat{\mathbf{u}} = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}} \mathbf{R}^{-1} \mathbf{G}_{\mathbf{a}} + \mathbf{P}^{-1}\right)^{-1} \mathbf{G}_{\mathbf{a}}^{\mathrm{T}} \mathbf{R}^{-1} \mathbf{r}. \tag{1.3}$$

The model covariance matrix,  $\mathbf{P} \equiv \langle \mathbf{u}\mathbf{u}^T \rangle$ , is the expected covariance of the unknown velocity, where  $\langle \cdot \rangle$  denotes the expected value. The measurement uncertainty covariance matrix is  $\mathbf{R} \equiv \langle \mathbf{n_a} \mathbf{n_a}^T \rangle$ . In simple least-squares fitting, both  $\mathbf{P}$  and  $\mathbf{R}$  are assumed to be a scaled identity matrix:

$$\mathbf{P} = \sigma_s^2 \mathbf{I} \,, \ \mathbf{R} = \sigma_r^2 \mathbf{I}. \tag{1.4}$$

With the assumption of infinite signal variance ( $\sigma_s^2 = \infty$ ) and unit error variance ( $\sigma_r^2 = 1$ ), the WLS method matches the un-weighted least-squares fitting (UWLS) method, which is the standard method in the estimate of vector currents ([Lipa and Barrick(1983)], [Gurgel(1994)], [Graber et al.(1997)Graber, Haus, Chapman, and Shay]). The UWLS method also assumes constant vector currents within the search radius ( $d_0$ ) of the vector grid point, and the estimate is

$$\hat{\mathbf{u}} = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}} \mathbf{G}_{\mathbf{a}}\right)^{-1} \mathbf{G}_{\mathbf{a}}^{\mathrm{T}} \mathbf{r} , \qquad (1.5)$$

where  $\mathbf{G}_{\mathbf{a}}^{\mathrm{T}}\mathbf{G}_{\mathbf{a}}$  is called the geometric covariance matrix (2×2 elements), of which inverse matrix is geometric dilution of precision (GDOP).

The a posterior uncertainty matrices in the WLS and UWLS methods are

$$\hat{\mathbf{P}} = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}}\mathbf{R}^{-1}\mathbf{G}_{\mathbf{a}} + \mathbf{P}^{-1}\right)^{-1}, \qquad (1.6)$$

$$\hat{\mathbf{P}} = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}}\mathbf{G}_{\mathbf{a}}\right)^{-1}, \tag{1.7}$$

respectively.

The normalized uncertainty matrix  $(\varphi)$  is introduced in order to compare the uncertainty in the estimate between the UWLS and WLS methods. The corresponding normalized uncertainty matrix in the OI method is the *a posteriori* uncertainty matrix (Equation 1.6) divided by the observational error variance  $(\sigma_x^2)$ :

$$\varphi = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}}\mathbf{G}_{\mathbf{a}} + \mathbf{P}^{-1}\sigma_{r}^{2}\right)^{-1}, \tag{1.8}$$

recall that  $\mathbf{R} = \sigma_r^2 \mathbf{I}$ . On the other hand, the uncertainty matrix in the UWLS method is already scaled with the observational error variance to the implied error variance 1 cm<sup>2</sup>s<sup>-2</sup>:

$$\varphi = \left(\mathbf{G}_{\mathbf{a}}^{\mathrm{T}}\mathbf{G}_{\mathbf{a}}\right)^{-1}.\tag{1.9}$$

## References

- [Graber et al.(1997) Graber, Haus, Chapman, and Shay] Graber, H. C., B. K. Haus, R. D. Chapman, and L. K. Shay (1997), HF radar comparisons with moored estimates of current speed and direction: Expected differences and implications, J. Geophys. Res., 102(C8), 18,749–18,766.
- [Gurgel(1994)] Gurgel, K.-W. (1994), Shipborne measurements of surface current fields by HF radar, L'Onde Electrique, 74(5).
- [Kim et al.(2007)Kim, Terrill, and Cornuelle] Kim, S. Y., E. Terrill, and B. Cornuelle (2007), Mapping surface currents from hf radar radial velocity measurements using optimal interpolation, J. Geophys. Res., submitted.
- [Lipa and Barrick(1983)] Lipa, B. J., and D. E. Barrick (1983), Least-squares methods for the extraction of surface currents from CODAR crossed-loop data: Application at ARSLOE, IEEE J. Oceanic Eng., 13(2), 507–513.
- [Wunch(1996)] Wunch, C. (1996), The ocean circulation inverse problem, Cambridge University Press.

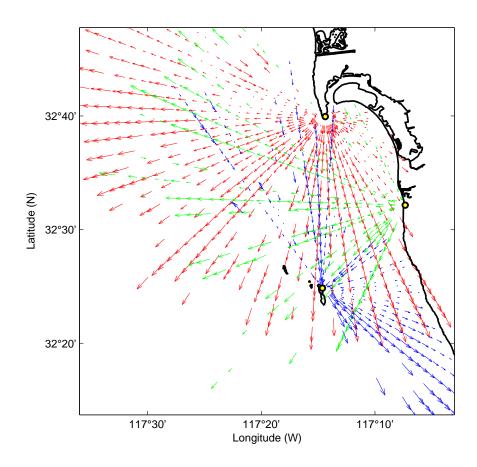


Figure 1: An example of the radial velocity.

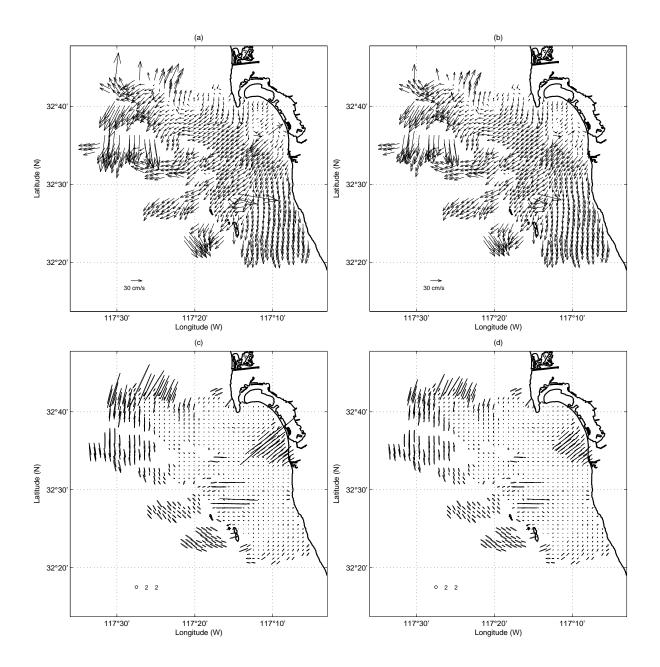


Figure 2: An example of the surface current vector and a posteriori uncertainty normalized by the observational error variance (a) UWLS (b) WLS (c) UWLS ( $\varphi$ ) (d) WLS ( $\varphi$ )