Global plate model choice impacts reconstructions of the latitudinal biodiversity gradient

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Lewis A. Jones1, Bethany J. Allen2,3, William Gearty4, and Lucas Buffan4

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1Grupo de Ecoloxía Animal, Departamento de Ecoloxía e Bioloxía Animal, Universidade de Vigo, 36310 Vigo, Spain.

2Department of Biosystems Science and Engineering, ETH Zürich, 4058 Basel, Switzerland.

3Computational Evolution Group, Swiss Institute of Bioinformatics, 1015 Lausanne, Switzerland.

4Division of Paleontology, American Museum of Natural History, New York, NY, 10024 USA.

5Département de Biologie, École Normale Supérieure de Lyon, Université Claude Bernard Lyon 1, 69342 Lyon Cedex 07, France.

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**All authors contributed equally to this work**

**Corresponding author:** LewisAlan.Jones@uvigo.es

¶ Abstract

Here goes the abstract…

# Keywords

Latitudinal biodiversity gradient, marine invertebrates, macroecology, global plate model, palaeogeographic uncertainty

# Introduction (600 words)

Today, species richness decreases from the tropics to the poles. This phenomenon is known as the latitudinal biodiversity gradient (LBG), and is one of Earth’s longest recognised macroecological patterns [1–4]. While observed across numerous taxonomic groups in the terrestrial [2,3] and marine realm today [5,6], the fossil record suggests this broadly unimodal gradient was not always present, with flattened and bimodal gradients observed across a range of taxonomic groups at various points in Earth’s geological history [e.g. 7,8–15]. Several studies have demonstrated that tropical peaks and poleward declines in taxonomic richness are restricted to the last 30 million years (Myr), and intervals of the Palaeozoic, when cool icehouse climatic regimes persisted [e.g. 4,15–18]. Conversely, during intervals of warmer climatic conditions (i.e. greenhouse and interglacial periods), various taxonomic groups have exhibited flattened LBGs, or even temperate peaks in biodiversity [e.g. 8,9,11–13,19,20]. However, recent work suggests that the recognition of different types of LBG in deep time is limited by incomplete and heterogeneous spatial sampling [21].

When examining the LBG in the present day, neontologists use the direct geographic distribution of organisms. However, palaeobiologists must contend with the shift of the continents over geological timescales; the geographic location of a fossil occurrence on the Earth’s surface today does not necessarily represent its location *in vivo*. Consequently, reconstructing the past geographic distribution of fossil occurrences is fundamental to reconstructing the LBG in deep time. To do so, palaeobiologists routinely use what is known as Global Plate Models–sometimes also called ‘palaeorotation models’ or ‘plate rotation models’–[14,18,20,22–24]. These models aim to reconstruct the tectonic evolution of the Earth by modelling the motion of the continental–and sometimes marine–plates across its surface through geological time. Many of these Global Plate Models exist [e.g. 25,26–28], varying in the way they define the geological boundaries of continents and how they rotate them through time, having consequence for how fossil occurrences might be palaeogeographically reconstructed [29]. However–to date–few palaeobiological studies (though see ref. [30,31]) have considered how different Global Plate Models might influence reconstructions of the latitudinal distribution of fossil occurrence.

Here, we test whether Global Plate Model choice influences the recognition of ‘unimodal-type’ latitudinal biodiversity gradients throughout the Phanerozoic (the last 540 million years). To do so, we reconstruct the palaeogeographic distribution of fossil occurrences for five major marine invertebrate groups using data from The Paleobiology Database and three Global Plate Models. Subsequently, we reconstruct the latitudinal biodiversity gradient using coverage-based rarefaction–a common sampling-standardisation approach–and quantify the strength of the gradient through time and the variability between Global Plate Models. We hypothesis that reconstructions of the latitudinal biodiversity gradient are more sensitive to plate rotation model choice with age of rotation.

# Materials and Methods (600 words)

## Occurrence data

We downloaded Fortunian–Piacenzian (541–0 Ma) fossil occurrence data from The Paleobiology Database (PBDB; <https://paleobiodb.org/>) for five major marine invertebrate groups (Bivalvia, Brachiopoda, Cephalopoda, Gastropoda, Trilobita) on mars 06 2023. Fossil occurrence data were downloaded using the PBDB API service and were restricted to marine environments, valid taxa and regular preservation (i.e. excluding form taxa and ichnotaxa). Occurrence data were subsequently binned into stratigraphic time bins following the Geological Timescale 2020 [32]. To do so, we used the bin\_time() function from the palaeoverse R package ver. 1.1.1 using the ‘majority’ approach [33]. Subsequently, we removed all occurrences with less than 95% of their age range covered by their assigned temporal bin. After data preparation, the occurrence dataset contained 414,366 occurrences from 76,998 collections.

## Palaeogeographic reconstruction and binning

To reconstruct the palaeogeographic distribution of fossil occurrences, we used occurrences’ present-day coordinates and midpoint age from assigned temporal bins with three Global Plate Models: PALEOMAP [27], GOLONKA [26], and MERDITH2021 [28]. Palaeogeographic reconstructions were generated using the GPlates Web Service (<https://gwsdoc.gplates.org>) via the palaeorotate() function in palaeoverse ver. 1.1.1 [33]. Subsequently, for each Global Plate Model, fossil occurrences were binned into one of twelve equal-area latitudinal bins (assuming a regular spheroid Earth model with a radius of ~6,371 km), using the estimated palaeolatitudes ([Table 1](#tbl-bins)).

Table 1: Equal-area latitudinal bins used in this study. Bins are generated assumming a regular spheroid Earth model with a mean radius of ~6,371 km.

| Bin | Maximum | Midpoint | Minimum | Area (m2) | Proportion of Area |
| --- | --- | --- | --- | --- | --- |
| 1 | 90.00 | 73.235 | 56.47 | 4.24e+13 | 0.083 |
| 2 | 56.47 | 49.150 | 41.83 | 4.25e+13 | 0.083 |
| 3 | 41.83 | 35.920 | 30.01 | 4.25e+13 | 0.083 |
| 4 | 30.01 | 24.745 | 19.48 | 4.25e+13 | 0.083 |
| 5 | 19.48 | 14.540 | 9.60 | 4.25e+13 | 0.083 |
| 6 | 9.60 | 4.800 | 0.00 | 4.25e+13 | 0.083 |
| 7 | 0.00 | -4.800 | -9.60 | 4.25e+13 | 0.083 |
| 8 | -9.60 | -14.540 | -19.48 | 4.25e+13 | 0.083 |
| 9 | -19.48 | -24.745 | -30.01 | 4.25e+13 | 0.083 |
| 10 | -30.01 | -35.920 | -41.83 | 4.25e+13 | 0.083 |
| 11 | -41.83 | -49.150 | -56.47 | 4.25e+13 | 0.083 |
| 12 | -56.47 | -73.235 | -90.00 | 4.24e+13 | 0.083 |

## Quantifying the latitudinal biodiveristy gradient

* Metrics used to quantify the gradient

# Results (600 words)

* Summary of reconstructions (could all points be reconstructed for each model?)
* Summary of results from metrics, do different gradients emerge?

# Discussion (700 words)

* Recap on importance of GPMs for deep time macroecology?
* What have we shown?
* Are some times or areas more problematic than others?
* Importance for other fields beyond palaeobiology?
* Consider importance of GPM choice in future work… or not?

# Data accessibility

# Authors’ contributions

# Funding

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