

Frehg User Manual

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Introduction

Fine Resolution Environmental Hydrodynamic and Groundwater model (Frehg)

Existing capabilities:

- Simulate 2D depth-integrated shallow water flow with wetting/drying and complex topography
- Simulate 3D variably-saturated subsurface flow
- Simulate coupled 2D surface and 3D subsurface flow
- Simulate simple advective-diffusive transport of passive scalars

Pre-requisite

- **laspack** for solving linear systems
(<http://www.mgnet.org/mgnet/Codes/laspack/html/node2.html>)
- **gcc** or other C-compiler
- **mpich** for message passing

Start a simulation

In the `frehg` directory:

Create an input folder to put model input data (e.g. `ex1_input`)

Create an output folder for saving model outputs (e.g. `ex1_output`)

Customize user setting file (named **input** by default), where:

- **sim_id** = Simulation ID. User can define case-specific settings in the source code by referring to certain `sim_id`
- **finput** = Directory of input data (e.g. `ex1_input/`)
- **foutput** = Directory of output data (e.g. `ex1_output/`)

Compile the code by typing **make** in terminal

Execute simulation by typing **./frehg** in terminal

If in parallel mode, execute by typing **mpiexec -np N ./frehg** (N = total number of sub-processes)

Surface domain

Set `sim_shallowwater = 1` to activate surface flow

Set `difuwave = 1` to use the diffusive wave approximation

Surface domain is characterized by:

- `NX` = number of grids in x direction (i-index)
- `NY` = number of grids in y direction (j-index)
- `dx` = grid resolution in x direction [m]
- `dy` = grid resolution in y direction [m]

Use the 2D domain on the right side for illustration:

- `NX = 3`, `NY = 4`, ghost cells are added along the boundaries
- Variables are stored in allocated 1D arrays
 - Numbers in the cells represent 1D array indices
 - A map is built for conversion between 1D/2D indices (named **smap** in the source code)
 - The map also stores the connections between grid cells

Set `bath_file = 1` to read bathymetry from a **bath** file, otherwise bottom elevation is zero all over the domain

29	22	23	24	25	26	i = 0
15	0	3	6	9	12	1
16	1	4	7	10	13	2
17	2	5	8	11	14	3
28	18	19	20	21	27	
j = 0	1	2	3	4	5	

Subsurface domain

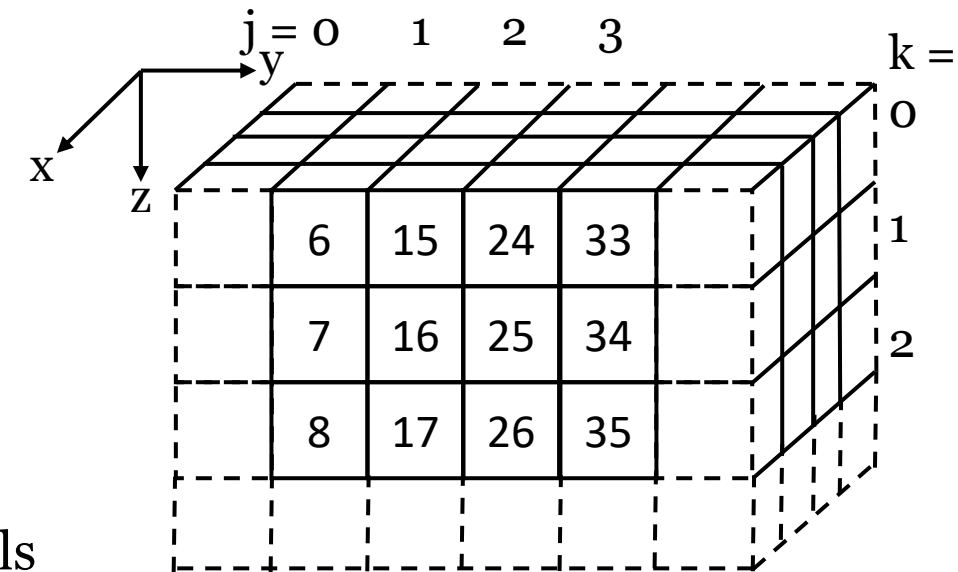
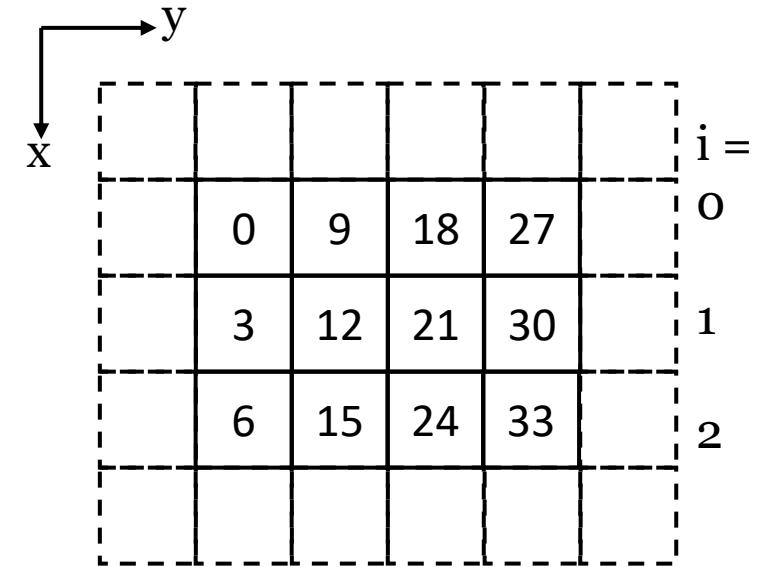
Set `sim_groundwater = 1` to activate subsurface flow

Subsurface domain is characterized by:

- `botZ` = bottom elevation of the domain [m]
- `dz` = grid resolution in z direction [m]
- `dz_incre` = used to create non-uniform z-discretization (not implemented for the current version)

Use the 3D domain on the right side for illustration:

- $NX = 3$, $NY = 4$, $NZ = 3$
- Ghost cells are added (not indexed) along the boundaries
- Variables are stored in allocated 1D arrays
 - Numbers in the cells represent 1D array indices
 - A map is built for conversion between 1D/3D indices (named **gmap** in the source code)
 - The map also stores the connections between grid cells



Simulation control

Parallelization (set `use_mpi = 1`):

- `mpi_nx` = number of threads in x direction (must be a factor of `NX`)
- `mpi_ny` = number of threads in y direction (must be a factor of `NY`)

Time control:

- `dt` = time step size [sec]
- `Tend` = end time [sec]
- `dt_out` = output frequency [sec]

If `sim_groundwater = 1`, variable dt is allowed by setting `dt_adjust = 1`. Simulation will start from `dt` and self-adjust between `dt_min` and `dt_max`:

- `dt_max` = maximum dt [sec]
- `dt_min` = minimum dt [sec]
- `Co_max` = maximum Courant number for subsurface solver

Then dt is adjusted using the water content + Courant number criteria as described in [3]

Surface domain – IC

Initial conditions for surface flow includes:

- **init_eta** = initial surface elevation [m]
- **init_tide** = initial tidal elevation at the tidal boundary [m]. This is used when surface elevation at a boundary is a constant.

Spatially-variable initial conditions can be read from file:

- **eta_file = 1** reads file **surf_ic** from **finput** folder, which contains cell-by-cell initial surface elevation values in the order of the 1D array
- **uv_file = 1** reads file **uu_ic** and **vv_ic** from **finput** folder, which contains cell-by-cell initial velocities in the order of the 1D array
 - If **uv_file = 0**, initial velocities are zero
- When restarting a simulation, the user should copy simulated surface elevation and velocities from **foutput** into **finput**, rename them to **surf_ic**, **uu_ic** and **vv_ic**, then set **eta_file** and **uv_file** to **1**.

Surface domain – inflow

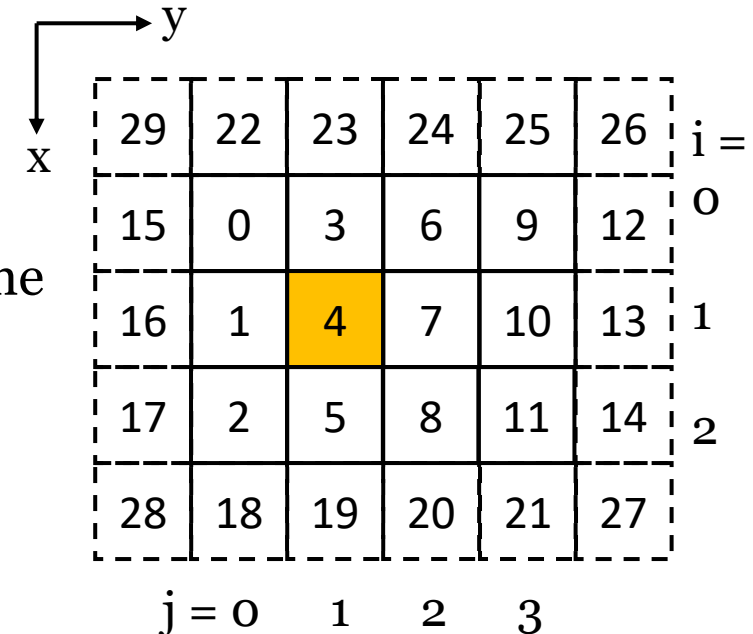
Set inflow boundary condition (or source/sink terms in general):

- **inflow_locX/locY** = x-start, x-end, y-start, y-end indices where inflow BC is applied, which is similar to how tide location is defined.
- **n_inflow** = total number of inflow BC to be applied
- **inflow_file** = 1 if time—variable inflow data is to be read
- **init_inflow** = constant inflow rate [m^3/s], which is used when **inflow_file** = 0

For example, if inflow BC is enforced for cell 4 in the right figure:

- **inflow_locX** = 1,1
- **Inflow_locY** = 1,1

If **inflow_loc** contains **N** grid cells, inflow rate for each cell is (assume constant) **init_inflow/N**



Surface domain – rainfall

Set evaporation/rainfall boundary conditions:

- **evap_file = 1** if time—variable evaporation data is to be read
- **q_evap** = constant evaporation rate [m/s], which is used when **evap_file = 0**
- **rain_file = 1** if time—variable rainfall data is to be read
- **q_rain** = constant rainfall rate [m/s], which is used when **rain_file = 0**

Evaporation from the subsurface domain:

- **evap_model = 1** if aerodynamic evaporation model is applied to estimate moisture-dependent evaporation rate on dry surface. Otherwise evaporation only occurs on wet surface.

In the current version, evaporation/rainfall have to be applied to the entire domain. Both evaporation and rainfall rates should be positive here.

Surface domain – wind

Set wind stress over wet surface:

- `sim_wind = 1` if wind effect is modeled
- `wind_file = 1` if time-variable wind speed and direction are read from data files (this function is not implemented yet, so set `wind_file = 0` for now)
- `init_windspeed` = wind speed [m/s]
- `init_winddir` = direction of wind [degrees] measured clockwise from the “north” direction
- `north_angle` = direction of “north” direction measured clockwise from the $-x$ direction
- `Cw` = wind drag coefficient
- `CwT` = decay rate of wind stress within the thin-layer [1]

Surface domain – parameters

Set model parameters for the surface domain:

- **grav** = gravitational constant [m/s^2]
- **viscx, viscy** = eddy viscosity [m^2/s] (currently version only supports constant eddy viscosity)
- **min_dept** = minimum depth below which a grid cell is considered as dry [m]
- **manning** = Manning's roughness coefficient
- **wtfh** = minimum depth to trigger the waterfall model [1]. In the current version, the waterfall model is disabled
- **hD** = depth below which drag coefficient is increased to avoid instability [1]
- **rhoa** = air density [kg/m^3]
- **rhow** = water density [kg/m^3]

Subsurface domain – IC

Initial conditions for subsurface flow includes:

- **init_wc** = initial water content. If **init_wc** is between saturated and residual water content, the domain will be initialized with constant **init_wc**
- **init_h** = initial pressure head [m]. If **init_wc** is out of range and **init_h** is not positive, the domain will be initialized with constant **init_h**
- **init_wt_rel** = initial water table relative to the topography [m] (must be positive). If neither **init_wc** nor **init_h** is used, a water table **init_wt_rel** m below topography will be set. Regions below the water table is fully saturated.
- **init_wt_abs** = initial water table [m]. If none of the above 3 settings is invoked, initial water table is located at **init_wt_abs**, which means a non-terrain-following water table is used.
- **h_file** = 1 if element-by-element initial head values is to be read from a file named **head_ic**.
- **wc_file** = 1 if element-by-element initial water content values is to be read from a file named **moisture_ic**.

Subsurface domain – BC

Boundary conditions for subsurface flow is set by:

- **bctype_GW** = type of boundary conditions used for x-minus, x-plus, y-minus, y-plus, z-minus(bottom), z-plus(top) boundaries, possible values are:
 - 0 = no flow
 - 1 = constant head (Dirichlet type)
 - 2 = constant flux (Neumann type)
 - 3 = free drainage (only used for bottom boundary)
 - When simulating surface-subsurface exchange, the top boundary cells with positive surface depth will automatically switch to Dirichlet condition.
- **qtop** = flux at top boundary [m/s]
- **qbot** = flux at bottom boundary [m/s]
- **htop** = head at top boundary [m]
- **hbot** = head at bottom boundary [m]

Note that **qtop** and **qbot** are positive if upward, negative otherwise. If **bctype_GW[5] = 2** and net infiltration exists (evaporation+rainfall is downward), infiltration becomes surface ponding and BC is Dirichlet type. If net exfiltration exists, exfiltration rate is used as qtop.

Subsurface settings

- `use_corrector = 1` activates the explicit corrector step [3], otherwise only the head form of the Richards equation is solved, which is non-conservative in the unsaturated zone
- `post_allocate = 1` activates the post-allocation step [3]
- `use_full3d = 1` activates 3D post-allocation, otherwise post-allocation is only 1D in the vertical direction. For large domain where horizontal mesh size \gg vertical mesh size, set `use_full3d = 0` is sufficient.
- `use_mvg = 1` activates the modified Mualem-van Genuchten model [3], which allows larger dt than the original van Genuchten model
 - `aev` = air entry value when using the MVG model [m]

Subsurface parameters

- K_{sx} , K_{sy} , K_{sz} = saturated hydraulic conductivities in x, y, z directions [m/s]
- S_s = specific storage [1/m]
- $soil_a$ = alpha parameter in the van Genuchten model
- $soil_n$ = n parameter in the van Genuchten model
- wcs = saturated water content
- wcr = residual water content

Scalar transport

The current version of frehg only simulates advective-diffusive transport of passive scalars.

- `n_scalar` = total number of scalars to be modeled
- `difux/difuy/difuz` = molecular diffusivities of the scalars.
- `disp_lon/disp_lat` = longitudinal/transverse dispersivity in the subsurface domain
- `scalar_surf_file = 1` if initial surface scalar concentration is to be read from a file named **scalar_surf_icN**, where N represents the index of the scalar.
- `init_s_surf = initial` scalar concentration for the surface domain, which is used if `scalar_surf_file = 0`
- `scalar_subs_file = 1` if initial subsurface scalar concentration is to be read from a file named **scalar_subs_icN**, where N represents the index of the scalar.
- `init_s_subs = initial` scalar concentration for the subsurface domain, which is used if `scalar_subs_file = 0`

Scalar transport

The current version of frehg only simulates advective-diffusive transport of passive scalars.

- `scalar_tide_file = 1` if time-variable scalar concentration for tide is to be read from a file named **scalarN_tideM**, where N is the index of scalar and M is the index of the tide.
- `scalar_tide_datlen` = number of time-variable scalar data points to be read from the **scalarN_tideM** file.
- `s_tide` = constant scalar concentration for tide, which is used when `scalar_tide_file = 0`
- `scalar_inflow_file = 1` if time-variable scalar concentration for inflow is to be read from a file named **scalarN_inflowM**, where N is the index of scalar and M is the index of the tide.
- `scalar_inflow_datlen` = number of time-variable scalar data points to be read from the **scalarN_inflowM** file.
- `s_inflow` = constant scalar concentration for tide, which is used when `scalar_inflow_file = 0`

References

- [1] **Li, Z.** and Hodges, B.R., 2019, Model instability and channel connectivity for 2D coastal marsh simulations, *Environ Fluid Mech*, (19): 1309,
<https://doi.org/10.1007/s10652-018-9623-7>
- [2] **Li, Z.** and Hodges, B.R., 2019, Modeling subgrid-scale topographic effects on shallow marsh hydrodynamics and salinity transport, *Advances in Water Resources*, (129) 1-15,
<https://doi.org/10.1016/j.advwatres.2019.05.004>
- [3] **Li, Z.**, Ozgen, I. and Maina, F.Z., 2020, A mass-conservative predictor-corrector solution to the 1D Richards equation with adaptive time control, *Journal of Hydrology*, (592)125809,
<https://doi.org/10.1016/j.jhydrol.2020.125809>