

shallow_water01.f90 – semi-implicit one-dimensional shallow water equations

The model uses Arakawa C-grid staggering. Linearized equation of motion. Bottom friction.

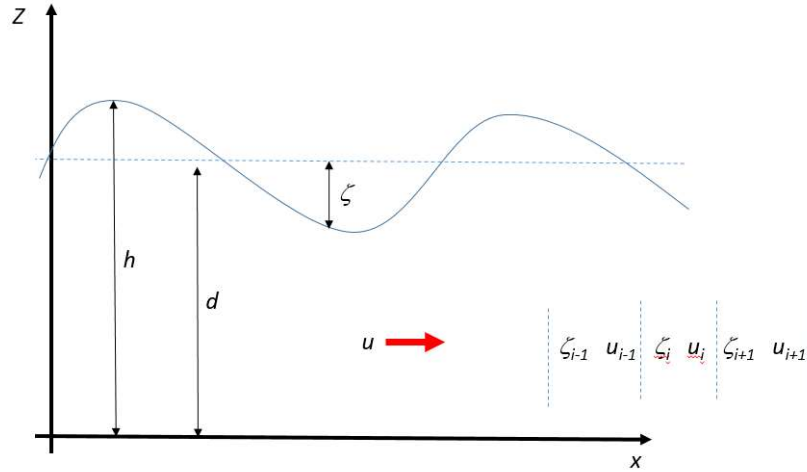


Fig 1: Geometry of the 1-D shallow water problem. Disturbed water depth h , un-disturbed water depth d , sea surface elevation ζ , horizontal velocity u . The lower right corner shows the C-grid staggering.

The momentum equation is formulated on the old (n) and new time level ($n+1$). The weight between both levels is w_{imp} .

$$u^{n+1} = u^n - \Delta t c_D \frac{1}{h} u^n |u^n| - \Delta t g (1 - w_{imp}) \frac{\partial \zeta^n}{\partial x} - \Delta t g w_{imp} \frac{\partial \zeta^{n+1}}{\partial x} \quad (1)$$

The same is valid for the equation of continuity:

$$\zeta^{n+1} = \zeta^n - \Delta t (1 - w_{imp}) \frac{\partial (h u)^n}{\partial x} - \Delta t w_{imp} \frac{\partial (h u)^{n+1}}{\partial x} \quad (2)$$

with

$$h = d + \zeta \quad (3)$$

Our strategy is to form a system of equations for all elements of ζ^{n+1} , solve this system iteratively and finally compute the velocity at the new time step u^{n+1} .

Inserting equation (1) into (2) yields:

$$\zeta^{n+1} = \zeta^n - \Delta t (1 - w_{imp}) \frac{\partial(h u)^n}{\partial x} - \Delta t w_{imp} \frac{\partial(h^n u^*)}{\partial x} + \Delta t^2 w_{imp}^2 g h \frac{\partial^2 \zeta^{n+1}}{\partial x^2} \quad (4)$$

with the interim solution

$$u^* = u^n - \Delta t c_D \frac{1}{h} u^n |u^n| - \Delta t g (1 - w_{imp}) \frac{\partial \zeta^n}{\partial x} \quad (5)$$

If we define

$$div = -\Delta t (1 - w_{imp}) \frac{\partial(h u)^n}{\partial x} - \Delta t w_{imp} \frac{\partial(h^n u^*)}{\partial x} \quad (6)$$

then it follows

$$\zeta_i^{n+1} = \frac{1}{1+c_E+c_W} [\zeta^n + div + c_E \zeta_{i+1}^{n+1} + c_W \zeta_{i-1}^{n+1}] \quad (7)$$

with

$$c_E = \begin{cases} \frac{\Delta t^2 w_{imp}^2 g h_E}{\Delta x^2} & \text{if } h_{i+1} > 0 \\ 0. & \text{otherwise} \end{cases} \quad (8)$$

$$c_W = \begin{cases} \frac{\Delta t^2 w_{imp}^2 g h_W}{\Delta x^2} & \text{if } h_{i-1} > 0 \\ 0. & \text{otherwise} \end{cases} \quad (9)$$

with

$$h_E = 0.5 (h_i + h_{i+1}) \quad (10)$$

$$h_W = 0.5 (h_i + h_{i-1}) \quad (11)$$

i.e. the water depth at the velocity point u_i and u_{i-1} .

After the system of equations (6) is solved iteratively with the Gauss-Seidel-Method, the velocity at the new time level is computed with

$$u^{n+1} = u^* - \Delta t \, g \, w_{imp} \frac{\partial \zeta^{n+1}}{\partial x} \quad (12)$$

Reference:

Backhaus JO (1983) A semi-implicit scheme for the shallow water equations for application to shelf sea modelling. Continental Shelf Research, Vol. 2, No. 4, pp. 243 - 254. [https://doi.org/10.1016/0278-4343\(82\)90020-6](https://doi.org/10.1016/0278-4343(82)90020-6)