

CMT3D_FLEXWIN Manual

Qinya Liu
Department of Physics
University of Toronto
January 30, 2012

CMT3D_FLEXWIN is re-written from my old CMT3D program to

- adapt to the output windows file from Alessia's FLEXWIN package, which runs through data and synthetics to pick time windows with decent match between data and synthetics.
- eliminate bandpass filter as it should have been applied before entering FLEXWIN.
- allow inversions to be done in global coordinates (`global_coord=.true.`), while by default, inversions are done in local RTP coordinates
- incorporate centroid time shift and half duration as inversion parameters (`npar=10` and `11`).
- reflect the gist of the program, which is to use Gaussian elimination to solve the inverse problem.

What still remains to be added is the possibility of CMT3D inversion for multiple subevents, which may be useful if N_{event} is relatively small, and may be handy for efficiency comparisons to CMT3D_CG.

1 Setup

1.1 Setup derivative cmt solutions

According to the theory in the Appendix, we need to compute the derivative synthetics

$$\frac{\partial s}{\partial m} \sim \frac{s(m + \Delta m) - s(m)}{\Delta m} \quad (1)$$

for all source parameters involved in the inversion. Depending on how many parameters are chosen (`npar`), we need to run at least 6 – 9 simulations to gather all the derivatives. Each SEM simulation requires a corresponding CMTSOLUTION, and the Δm should be chosen carefully to avoid either numerical inaccuracy or instability. For global code (SPECFEM3D_GLOBE), a choice may be

$$\Delta M = 1 \times 10^{24} \text{ dyne.cm}, \quad \Delta X = 2 \text{ km}, \quad \Delta \delta = 0.02 \quad (2)$$

and for basin code (SPECFEM3D), a choice may be

$$\Delta M = 1 \times 10^{22} \text{ dyne.cm}, \quad \Delta X = 0.5 \text{ km}, \quad \Delta \delta = 0.01 \quad (3)$$

where ΔX refers to depth perturbations and $\Delta \delta$ refers to perturbation in latitude and longitude.

Depending on what coordinate system you choose to calculate derivatives and run CMT3D inversions, different extensions need to be added to derivative CMTSOLUTIONs as well as the output synthetics (generated by `gen_der_cmt()` python script):

- local RTP (depth, lon, lat) coordinates (`global_coord=.false.`):
`ext=['Mrr','Mtt','Mpp','Mrt','Mrp','Mtp','dep','lon','lat']`
- global code, global XYZ coordinates (`global_coord=.true.`):
`ext=['Mxx','Myy','Mzz','Mxy','Mxz','Myz','xxx','yyy','zzz']`
 This coordinate system may not be necessary except to accommodate multiple event inversions in the future.

For derivations of rotation matrices, please refer to the Appendix.

1. Python script and modules are provided in `scripts` directory to facilitate this process, and allow full compatibility with CMT3D code:

```
gen_cmt_der('CMTSOLUTION',npar=9,dmoment=1.e24,
            ddepth=2,dlocation=2,utm=0)
```

where *ddepth* and *dlocation* are given in km ($\Delta\delta = \Delta X/111.12$), and *utm* = 0 refers to local coordinates and *utm* = -1 refers to global coordinates.

2. Also as indicated by the theory section, derivatives for centroid time and half duration simply depend on the synthetics, therefore no extra simulation is necessary.
3. After simulations for these derivative CMTSOLUTIONs finish, collect synthetic seismograms, convert them to sac files and add proper extension to the sac files according to the choice of coordinate system.
4. Note synthetics for the original CMTSOLUTION do not need to be obtained through an extra simulation, a simple linear combination of moment tensor derivatives can be realized by:

```
xadd_frechet_derivatives s syn_list cmt moment_scale(1e24)
```

1.2 Pre-processing data and synthetics

After gather raw data from IRIS or SCEDC, and collect synthetics as well derivative synthetics from SEM simulations you need to go through the pre-processing step to present properly cut and filtered data and synthetics to the FLEXWIN package. I have come a long way in writing such scripts, and right now I am settled on a recent one in python:

```
process_ds.py \
-d data,sac -s syn, -p CMTSOLUTION,STATIONS\
-t 6,30 -w -D [-r]
```

For detailed usage, simply type `process_ds.py`. The outcomes of this script are

- filtered data and (derivative) synthetics in `data_6_30` and `syn_6_30`.
- an input file to FLEXWIN named `INPUT_flexwin_6_30` is generated.

1.3 Executing FLEXWIN

After choosing the right parameters in `PAR_FILE` and user functions for flexwin package (please refer to FLEXWIN manual for details),

```
./flexwin < INPUT_flexwin_6_30 > OUTPUT_flexwin_6_30
```

After running flexwin through data and synthetics filtered at various period bands, all individual windows files in `MEASURE_t1_t2` directories can be cat into an input for `cmt3d_flexwin` package.

```
write_flexwin_out.py MEASURE_6_30 MEASURE_20_50 INPUT_cmt3d_flexwin_noweight
```

Now it is possible to pad time shift and weight information to the composite output windows file, either to facilitate a direct evaluation of misfit function, or to eliminate the weight calculations in CMT3D package, which can be rather cumbersome and inflexible.

```
write_flexwin_shift_weight.py -c 2/2/1 -d 1.15/0.55/0.78 -a 0.2 -r 2 -t \
-D data -s MEASURE INPUT_cmt3d_flexwin_noweight INPUT_cmt3d_flexwin_
```

Obviously you can not just blindly use this package without checking. Here are two auxiliary scripts:

```
plot_flexwin.py MEASURE
plot_data_and_syn.pl -d data_dir,data_ext -s syn_dir,syn_ext -m cmt \
-M OUTPUT_flexwin -A (Z/R/T) 0.05/0.05/0.08
```

The first script gives `flexwin_seismo.pdf` and the second one gives `ds_all.pdf`.

2 CMT3D

2.1 Parameter file

For structures of the main program and subroutine calls, refer to `call-graph`. The name of the parameter file is hard-wired in the main program as `INVERSION.PAR`.

```
CMTSOLUTION
CMTSOLUTION_NEW
6 .false. -- npar[, global_coord]
0.01 1.0 1.0e22 -- dlocation,ddepth,dmoment
flexwin.out
```

```
.true. .false.           -- weigh_data_files[, read_weight]
2 2 1 0 1.15 0.55 0.78  -- weights of data comp,az,dist
.true.                  -- station_correction
.true. .true. 0.0        -- zero_trace,double_couple,damping
.false.                 -- write_new_syn
```

The first line gives the old cmt solution file name. This is also the cmt solution you have used to generate derivative cmts and compute the synthetics from linear combinations. The second line gives name of the new cmt solution file written after the inversion is done.

File `flexwin.out` holds the start and end time of the windows outputted from FLEXWIN package, in the following format:

```
nfiles
data-file-1
syn-file-1
nwins-for-this-data-and-syn
tstart(1) tend(1) [tshift(1) weight(1)]
...
tstart(nwin) tend(nwin) [tshift(nwin) weight(nwin)]
data-file-2
syn-file-2
...
```

Only windows from this output file will be used to perform time integration in assembling the inversion matrices. Note the values in [] are recent additions where we allow weights to be specified in the input file together with an indicative logical `read_weight`. For now, ‘tshift’ are still dummy variables not used for aligning data and synthetics (a separate corr subroutine calculates it in the code, but future merge may be possible).

All other parameters are pretty self-explanatory:

- Number of parameters to be inverted can only be 6(moment-only), 7(moment+depth), 9(moment+locations). `global_coord` indicates if the input derivative synthetics are in terms of (X, Y, Z) or $(depth, lon, lat)$.
- The derivative synthetics should have been calculated based on the `ddepth`, `ddepth`, `dmoment` values. However, notice that strictly speaking, only the ones for moment are actual derivative synthetics; the ones for depth, longitude, and latitude (or R, S, E depending on coordinate system or utm) are actual synthetics, from which the code subtracts the old synthetics and divides it by `ddepth` or `ddepth` (refer to Appendix for more details).
- Weighting subroutine `compute_data_weights()` needs to be tinkered if you choose to `weigh_data_files`, and not importing weights directly from the input file by `read_weight`. The idea is that you utilize the `dist_km`, `dist_deg`, `azimuth`, and other file information from the sac headers to come up with a sensible weighting scheme for your particular scale/region.

- Station correction is used to align the data with the synthetics to accomodate 3D path effects that may bias your source inversion results.
- Different constraints on the moment-tensor are allowed, including zero-trace and double-couple. To improve the stability of the matrix inverse (when only very few windows are used), damping is also allowed before Gaussian elimination.
- Finally synthetics for the new cmt solution can be written, however, notice that these are obtained from ‘linearized’ approximation, not from an actual SEM simulation.

Some more words on weighting:

- If you want to invert either only surface waves or body waves, make sure you set the optimal parameters in flexwin.
- How do the weights work:

```
data_weights(nwint) = cmp_weight(nwint) &
    * (dist/REF_DIST) ** dist_exp_weight(nwint) &
    * corr(nwint) ** corr_exp_weight(nwint) &
    / ( naz(k) ** az_exp_weight)
```

where component weights are applied linearly, distance and azimuth weights are applied exponentially (such that 0 means no weight and 1 means linear weight). `naz(k)` is the number of traces available in a particular azimuth (+1). One can set `az_exp_weight` to be 0 to mute out the azimuthal effect, the same for distance weights.

This package has been fully tested for 6, 7, 9, 10 parameter inversions based on synthetic data.

A CMT3D Theory

A.1 Moment-tensor Inversions

We seek to minimize the difference between data $\{d_i(t), i = 1, \dots, N\}$ and corresponding synthetics $\{s_i(\mathbf{m}, t), i = 1, \dots, N\}$ computed for a given model parameters $\mathbf{m} = \{m_j, j = 1, \dots, M\}$ after applying some bandpass filter:

$$\phi(\mathbf{m}) = \frac{1}{2} \sum_{rq} \int W_{rq}(t) [s^b(\mathbf{m}, \mathbf{x}_r, t) - \mathbf{d}_i^b(\mathbf{x}_r, t)]^2 dt \quad (4)$$

where $W_{rq}(t)$ combines the taper function with weighting factor for the q 'th window at r 'th receiver. s^b and \mathbf{d}^b denote filtered data and synthetics.

If we take the first and second order derivatives of this misfit function at some given reference model \mathbf{m}^0 , we obtain

$$\frac{\partial \phi(\mathbf{m}^0)}{\partial m_j} = \sum_{i=rq} \int W_{rq}(t) [\mathbf{s}^b(\mathbf{m}^0, \mathbf{x}_r, t) - \mathbf{d}^b(\mathbf{x}_r, t)] \frac{\partial s_i(\mathbf{m}^0, \mathbf{x}_r, t)}{\partial m_j} dt \quad (5)$$

$$\frac{\partial^2 \phi(\mathbf{m}^0)}{\partial m_j \partial m_k} \sim \sum_{i=rq} \int W_{rq}(t) \frac{\partial \mathbf{s}^b(\mathbf{m}^0, \mathbf{x}_r, t)}{\partial m_j} \frac{\partial \mathbf{s}^b(\mathbf{m}^0, \mathbf{x}_r, t)}{\partial m_k} dt \quad (6)$$

where for the second-order derivatives, we have omitted the term related to second-order synthetics derivatives, i.e. approximate Hessian

By assuming that at the optimal model parameter \mathbf{m} , we have $\frac{\partial \phi}{\partial \mathbf{m}} = 0$, which translates to the quasi-Newton's method:

$$\frac{\partial \phi(\mathbf{m}^0)}{\partial m_j} + \frac{\partial^2 \phi(\mathbf{m}^0)}{\partial m_j \partial m_k} (m_k - m_k^0) = 0 \quad (7)$$

Therefore, we just need to compute the first and second-order derivatives according to equations ??, which further points to the computation of derivative synthetic seismograms $\frac{\partial s_i(\mathbf{m}^0)}{\partial m_j}$.

A.2 Derivative Synthetics

We know in general, the synthetics for model \mathbf{m} can be expanded with respect to a reference model \mathbf{m}^0 :

$$s(\mathbf{m}, t) = s(\mathbf{m}^0, t) + \sum_{j=1}^M \frac{\partial s_i(\mathbf{m}^0, t)}{\partial m_j} (m_j - m_j^0) \quad (8)$$

If source parameters only include moment-tensor elements, then the linear dependences can be written as:

$$s(\mathbf{m}, t) = \sum_{j=1}^M \frac{\partial s_i(\mathbf{m}^0, t)}{\partial m_j} m_j \quad (9)$$

Therefore derivatives for moment-tensor elements can be computed by

$$\frac{\partial s_i(\mathbf{m}^0, t)}{\partial M_j} = s(M_j^0, t) / M_j^0 \quad (10)$$

where $s(M_j^0, t)$ is the synthetics calculated by setting all the moment tensor elements other than the j th component to 0.

For other source parameters, including event location and depth, a finite-difference formula has

to be used:

$$\frac{\partial s_i(\mathbf{m}^0, t)}{\partial m_j} = \frac{s(m_j^0 + \Delta m_j^0, t) - s(m_j^0, t)}{\Delta m_j} \quad (11)$$

where $s(m_j^0 + \Delta m_j^0, t)$ refers to synthetics computed from only perturbing the j 'th component of the source parameters. Δm_j needs to be selected with care to avoid numerical noise. Typically 2 – 5 km for global earthquakes and 1 km for local earthquakes.

Derivatives for time shift and half duration can be derived from the original synthetics itself. Given a Gaussian source time function

$$g(t) = \frac{1}{\sqrt{\pi}t_h} e^{-\left(\frac{t-t_s}{t_h}\right)^2} \quad \text{or} \quad g(\omega) = e^{-\frac{\omega^2 t_h^2}{4}} e^{-i\omega t_s} \quad (12)$$

under the Fourier transformation pairs:

$$F(\omega) = \int_{-\infty}^{\infty} f(t) e^{-i\omega t} dt, \quad f(t+a) \Leftrightarrow F(\omega) e^{i\omega a}, \quad \dot{f}(t) \Leftrightarrow (i\omega) F(\omega) \quad (13)$$

Given Green's functions for a point source with delta moment function, we can express the synthetics for an earthquake as

$$\mathbf{s}(\mathbf{x}_r, \omega) = \frac{M}{i\omega} : \nabla G(\mathbf{x}_r, \mathbf{x}_s, \omega) g(\omega) \quad (14)$$

i.e., synthetics for error function as moment function. Therefore

$$\frac{\partial \mathbf{s}(\mathbf{x}_r, \omega)}{\partial t_{s,h}} = \frac{M}{i\omega} : \nabla G(\mathbf{x}_r, \mathbf{x}_s, \omega) \frac{\partial g(\omega)}{\partial t_{s,h}}. \quad (15)$$

Since

$$\begin{aligned} \frac{\partial g(\omega)}{\partial t_s} &= -(i\omega) g(\omega) \\ \frac{\partial g(\omega)}{\partial t_h} &= -\frac{1}{2} t_h \omega^2 g(\omega) = -\frac{1}{2} t_h (i\omega) \frac{\partial g(\omega)}{\partial t_s} \end{aligned} \quad (16)$$

we have

$$\begin{aligned} \frac{\partial s(\mathbf{m}_0, t)}{\partial t_s} &= -\dot{\mathbf{s}}(\mathbf{m}_0, t) \\ \frac{\partial s(\mathbf{m}_0, t)}{\partial t_h} &= -\frac{1}{2} t_h \partial_t \left[\frac{\partial s(\mathbf{m}_0, t)}{\partial t_s} \right] \end{aligned} \quad (17)$$

A.3 Rotation between Global and Local coordinates

A series rotation is needed to tranform local (r, θ, ϕ) coordinates (which is also what CMT moment tensor uses) into the global coordinates (X, Y, Z) , where X is pointing to 0° longitude, and vice

versa.

For a given point with $(\theta = 90 - \text{Lat}, \phi = \text{Long})$ on a **unit** sphere, the coordinates in the cartesian global coordinate system is given by

$$X = \sin \theta \cos \phi, \quad Y = \sin \theta \sin \phi, \quad Z = \cos \theta \quad (18)$$

Conversely, given the global coordinates (X, Y, Z) ,

$$\theta = \arccos(Z/r), \quad \phi = \arctan2(Y/\sin \theta, X/\sin \theta) \quad (19)$$

Note here if $\theta = 0$ (North pole), then ϕ is undetermined, and can be basically any number.

For rotation matrices, as an example, in 2-D, for the same point, the coordinates under new coordinate system is related to the old coordinates by

$$\begin{bmatrix} X_{new} \\ Y_{new} \end{bmatrix} = \begin{bmatrix} \cos \theta & \sin \theta \\ -\sin \theta & \cos \theta \end{bmatrix} \begin{bmatrix} X_{old} \\ Y_{old} \end{bmatrix} \quad (20)$$

After the first two simple rotations ($Z(\phi)$ and $Y(\theta)$) demonstrated by Figure 1, we have the relation

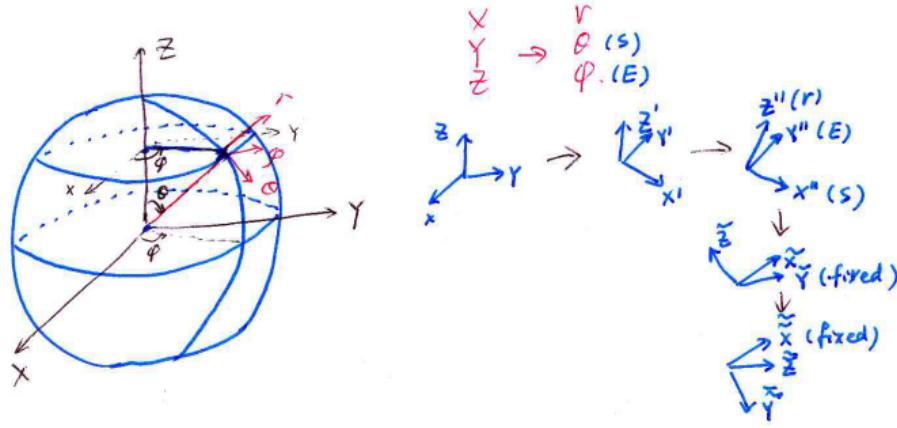


Figure 1: Rotation matrices

between (X, Y, Z) and (θ, ϕ, r) systems:

$$\begin{bmatrix} x_\theta \\ x_\phi \\ x_r \end{bmatrix} = \begin{bmatrix} \cos \theta & 0 & -\sin \theta \\ 0 & 1 & 0 \\ \sin \theta & 0 & \cos \theta \end{bmatrix} \begin{bmatrix} \cos \phi & \sin \phi & 0 \\ -\sin \phi & \cos \phi & 0 \\ 0 & 0 & 1 \end{bmatrix} \begin{bmatrix} X \\ Y \\ Z \end{bmatrix} \quad (21)$$

This is not the end of story, as the usual order of local coordinate system is (r, θ, ϕ) instead of (θ, ϕ, r) . Therefore two extra rotations ($Y(-90^\circ)$ and $X(-90^\circ)$) are needed, which give an extra

(left-applied) rotation matrix

$$\begin{bmatrix} 1 & 0 & 0 \\ 0 & \cos(-90^\circ) & \sin(-90^\circ) \\ 0 & -\sin(-90^\circ) & \cos(-90^\circ) \end{bmatrix} \begin{bmatrix} \cos(-90^\circ) & 0 & -\sin(-90^\circ) \\ 0 & 1 & 0 \\ \sin(-90^\circ) & 0 & \cos(-90^\circ) \end{bmatrix} \quad (22)$$

Multiply these two matrices together, we obtain the final rotation matrix from (X, Y, Z) to (r, θ, ϕ) :

$$R^T = \begin{bmatrix} \sin \theta \cos \phi & \sin \theta \sin \phi & \cos \theta \\ \cos \theta \cos \phi & \cos \theta \sin \phi & -\sin \theta \\ -\sin \phi & \cos \phi & 0 \end{bmatrix} \quad (23)$$

which can be verified by

$$\begin{bmatrix} 1 \\ 0 \\ 0 \end{bmatrix} = R^T \begin{bmatrix} \sin \theta \cos \phi \\ \sin \theta \sin \phi \\ \cos \theta \end{bmatrix} \quad (24)$$

The rotation matrix from local (r, θ, ϕ) to global (X, Y, Z) is simply:

$$R = \begin{bmatrix} \sin \theta \cos \phi & \cos \theta \cos \phi & -\sin \phi \\ \sin \theta \sin \phi & \cos \theta \sin \phi & \cos \phi \\ \cos \theta & -\sin \theta & 0 \end{bmatrix} \quad (25)$$

B Testing (further update awaits)

Two types of datasets are provided to accurately test this inversion package, one with station correction, and one without. Synthetics and derivative synthetics for corresponding CMTSOLUTION are always in `syn_T006_T030`. No pre-processing are needed for these dataset as they have gone through all the procedures. Although to achieve inversion result that is accurate to several digits after the decimal point, several tweaks with the usual pre-processing procedures have to be done.

- (obsolete, needs fix) we should not use `-h` in the synthetics processing to ensure that the synthetics data and the derivative synthetics go through the exact same set of processing procedures. we also need to take out the `rtrend` and `rmean` line in the filtering steps of `process_syn_new.pl` and `process_cal_data.pl` (?). Since the data and synthetics presented in `test_dataset` have gone through the script:

```
process_data_and_syn.pl -d data,sac -s syn, \
    -0 -m CMTSOLUTION -a station_file \
    -1 -b -26 -2 -t 6/30
```

which generated `process_ds.csh` script for all the commands, we can just tweak individual lines in `process_ds.csh` to achieve our test dataset. Also note that the `-i none`

option for data processing and `-h` option for synthetics processing have to be taken out from the script to make sure that synthetics and data are processed in exactly the same way.

- The synthetic data for the 7 parameter case are generated by the following command:

```
xadd_frechet_full s syn_file_list old_cmt true_cmt \
                  doment depth dx[dy]
```

which is a fancier version of the original

```
xadd_frechet_derivatives s syn_file_list \
                          cmt_file moment_scale
```

Of course we can't really use the true synthetics for the true cmt solution in a 7-par case, since we know that the synthetics is a non-linear function of depth, and the inversion won't accurately recover depth. Linearized 'pseudo-synthetics' will ensure accurate results, which is the purpose of testing.

- It is impossible to design test dataset for 7 parameter inversion with time shift, since time shifts interfere with the effect of waveform change caused by depth variation, therefore, may not give exact solution, and sometimes may even produce the incorrect solution. This is part of the reason why I also wrote the grid search package GRID3D_FLEXWIN.

C Auxilliary Scripts and Programs

A list of the scripts that are used before and after the package:

- `process_data_and_syn.pl`, which includes `process_data_new.pl` (global/regional) or `process_cal_data.pl` (socal), `process_syn_new.pl` (global/regional) or `process_trinet_syn_new.pl` (socal), `pad_zeros.pl` (include `pad_zero_to_syn`), `rotate.pl`, `sac1st`. Of course `sac` has to be present in the system.
- `prepare_meas_all.pl`
- `xadd_frechet_derivatives` and `xadd_frechet_full`