

Geostrophic Speed

Contents

1 Key Characteristics

- **Force Balance:** Achieved when Coriolis force ($-f \times v$) balances pressure gradient force ($-\frac{1}{\rho}\nabla p$)
- **Flow Direction:** Parallel to isobars/isopycnals (anti-clockwise around lows in NH)
- **Mathematical Formulation:**

$$v_g = \frac{1}{f\rho} \frac{\partial p}{\partial n}$$

where:

- $f = 2\Omega \sin \phi$ (Coriolis parameter)
- ρ = fluid density
- $\frac{\partial p}{\partial n}$ = pressure gradient normal to flow

2 Key Relationships

2.1 Atmospheric Version

$$v_g = \frac{g}{f} \frac{\Delta Z}{\Delta n}$$

where g = gravitational acceleration (9.81 m/s²), ΔZ = geopotential height difference.

2.2 Oceanic Version

$$v_g = \frac{g}{f\rho_0} \frac{\partial \rho}{\partial x} \Delta z$$

Accounting for thermal wind adjustment in stratified flows.

3 Practical Calculation Example

Given parameters:

$$\begin{aligned}\Delta Z &= 60 \text{ m over } 300 \text{ km} \\ f_{45^\circ \text{N}} &= 10^{-4} \text{ s}^{-1} \\ g &= 9.8 \text{ m/s}^2\end{aligned}$$

Solution:

$$\begin{aligned}v_g &= \frac{9.8}{10^{-4}} \times \frac{60}{3 \times 10^5} \\ &= 98 \times 10^3 \times 2 \times 10^{-4} \\ &= 19.6 \text{ m/s} \quad (\sim 38 \text{ knots})\end{aligned}$$

4 Limitations and Extensions

- **Gradient Wind Approximation:** For curved flows:

$$\frac{v^2}{r} + fv = \frac{1}{\rho} \frac{\partial p}{\partial n}$$

- **Thermal Wind:** Vertical shear relationship:

$$\frac{\partial v_g}{\partial z} = \frac{g}{fT} \frac{\partial T}{\partial n}$$