

The Aerosol Limb Imager: Acousto-Optic Measurements of Limb Scattered Sunlight for 2D Stratospheric Aerosol Profiles

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Abstract

The Atmospheric Limb Imager (ALI) is a prototype atmospheric instrument developed at the University of Saskatchewan with the long term goal satellite instrumentation to gather high resolution horizontal and vertical stratospheric aerosol profiles, including extinction and particle size. The immediate goal of the ALI prototype is to test and verify the use of an Acousto-Optical Tunable Filter (AOTF), the fundamental technology behind ALI, in a space environment and to gather atmospheric sulfate aerosol profiles with high spatial resolution. ALI will measure light from the atmosphere through the limb geometry which measures the radiance from scattered sunlight from sunlit atmosphere. ALI is a single channel instrument measuring radiance from the visible to the near infrared wavelengths (650-950 nm) and through successive images builds spectral information. The system uses a telescopic front end to pass collimated light through the AOTF for each line of sight which is then focused onto the the detector. The AOTF has the property of separating the incoming radiance into each of its linear polarized components and rotating the selected wavelength polarization by 90° allowing one to recover some polarization information of the incoming radiance. The ALI prototype was completed in August of 2014. A stratospheric balloon test flight from the Canadian Space Agency (CSA) balloon launch facility in Timmins, Ontario onboard the CNES CARMEN gondola platform occurred on September 20, 2014. The data for the flight undergoes a method to convert the data into relative radiances and are used in a MART retrieval method to determine aerosol extinction and size profiles.

1 Introduction

Aerosol has been monitored in several different geometries throughout history. Occultation was one of the first methods used on satellite instrumentation to measure atmospheric species, including ozone. Notable instruments included SAM II (*McCormick et al.*, 1979), SAGE II

(*McCormick*, 1987), and SAGE III (*Thomason and Taha*, 2003). Another geometry is the limb scatter geometry which measures radiance from the sunlit atmosphere as it scatters and interacts with molecules. However, the limb scatter geometry has a complexity that requires a forward model to be able to accurately compute the scattering interactions, both single and multiple, between different constituents. Limb scatter was first preformed on the Solar Mesosphere Explorer (SME) (*Barth et al.*, 1983) to measure ozone profiles in 1981. Two instruments from the previous generation of remote sensing instruments have successfully used limb scatter to determine aerosol atmospheric parameters, the Optical Spectrograph and InfraRed Imaging System (OSIRIS) a Canadian instrument onboard the Odin satellite (*Llewellyn et al.*, 2004) and SCanning Imaging Absorption spectroMeter for Atmospheric CHartographY (SCIAMACHY) onboard the ENVISAT (*Bovensmann et al.*, 1999) which are grating spectrometers that acquire data at a single altitude at a time so a series of exposures is required to create a vertical profile with approximately a 1.5 km and 3 km vertical resolution respectively. The Ozone Mapping Profiler Suite - Limb Profiler (OMPS-LP) (*Dittman et al.*, 2002) is a current generation instrument which images the atmosphere with spectral and vertical spatial information with 1 km vertical resolution. The new proposed method with ALI allows for two dimensional spatial images of a wavelength with polarized light scanned spectrally to be acquired giving both vertical and horizontal resolution of the environment through the use of the innovative AOTF technology. Aerosol is a broadband scattering feature with effects that can be observed from a wide range of wavelengths from 525 nm extending into the near infrared and does not require high spectral resolution in order to determine aerosol extinction profiles.

Sulfuric aerosol is a fundamental portion of the global climate balance. Stratospheric aerosol loading has been increasing at a rate of 4-7% per year from 2002 to 2009 due to moderate volcanic eruptions (*Vernier et al.*, 2011) which have lead to a negative radiative forcing effect that has partially offset the radiative forcing from CO₂ (*Solomon et al.*, 2011). Aerosol is a spherical partical that scatters incoming radiance away from the surface of the planet causing an overall cooling effect which is dependant on particle size distribution (*Kiehl and Briegleb*, 1993). This effect is fully defined by aerosol microphysics requiring accurate knowledge of concentration and particle size distributions. The method of aerosol transportation across the tropopause are not completely understood, recently after the 2011 eruption of Nabro it was determined that the asian monsoon has the ability to penetrate through the thermal boundary and ejected aerosol from the Nabro plume into the stratosphere (*Bourassa et al.*, 2012). Also anthropogenic sources of SO₂ can enter the stratosphere through the Asian tropopause aerosol layer that forms during the Asian monsoon (*Vernier et al.*, 2011; *Neely et al.*, 2014).

A lidar instrument, CALIPSO, is able to achieve aerosol profiles of 333 m horizontal resolution and 60 m, 180 m, and 300 m vertical resolution for the altitude ranges of 8.2-20.2 km, 20.2-30.1 km, and 30.1-40.0 km respectively (*Winker et al.*, 2003). Additional instrumentation with high resolutions on the order of 150 m vertically and horizontally will satisfy the needs for current scientific investigation and will allow the determination of aerosol loading from volcanic

and anthropogenic sources causing a surface cooling effect and as well account for losses from the stratosphere due to transportation that will counteract the previous cooling effect. ALI, a prototype for a next generation of instruments will address the much needed horizontal and vertical resolution previously stated required to help answer questions about global climate. The ALI prototype discussed in the work has been design with balloon geometries in mind, accounting for a larger field of view to be able to image the whole limb with a range of 35 km vertically. However, for satellite geometry with a low earth orbit (approximately 600 km) a considerable smaller field of view will be required to have the same 35x35km field of view.

In this paper the first section will outline the technology behind the AOTF used within the ALI system and then describe the optical design behind ALI including an ulterior optical presentation comparing the benefits and drawbacks between two optical layout which will be followed by a comparison with the Belgium instrument ALTIUS. An overview of ALI's maiden flight, including the conditions and trajectory of the flight, onboard the CNES CARMEN gondola from the CSA balloon launch facility in Timmins, Ontario will be presented. Post flight analysis of the data is presented including the measurements taken during the campaign, the conversion from raw measurements into calibrated data and, an retrieval algorithm for aerosol extinction and particle size parameters will be outlined and then preformed on the data from the campaign.

2 Instrument Design

2.1 Acousto-Optical Tunable Filter Characteristics

An AOTF uses an acoustic wave, set by a Radio Frewuency (RF), that is prorogated through the crystal and forms a standing wave to allow the diffraction of a specific wavelength forming a filtering effect. The use of an AOTF for an imaging system has several distinct advantages due to its low mass, fast stabilization times of a few microseconds, and no moving parts and has only recently become possible thanks to improvements to non-collinear acosto devices with the use of birefringent materials with large apertures (*Chang, 1974; Voloshinov et al., 2007*) which allowed for the ability to image with AOTF technology. In order for the AOTF to allow the diffraction of a specific wavelength a momentum matching criteria must be held where the wave vectors of the acoustic wave match the difference of the incoming and diffracted light wave vectors as seen in Figure 1. This condition is known as the Bragg matching criteria and is given by

$$\mathbf{k}_i = \boldsymbol{\kappa} + \mathbf{k}_d \quad (1)$$

where $|\mathbf{k}_i| = \frac{2\pi n_i}{\lambda}$ is the wave of the incoming light, $|\mathbf{k}_d| = \frac{2\pi n_d}{\lambda}$ is the momentum of the diffracted light, and $|\boldsymbol{\kappa}| = \frac{2\pi F}{\nu}$ is the momentum of the acousto wave. The parameters λ , F , and ν are the wavelength in vacuum, the frequency of the RF wave, and the phase velocity in the crystal. The condition given in Equation 1 and the wave vector diagram gives the following

relation for a birefringent material undergoing Bragg diffraction

$$\lambda = \frac{\Delta n \nu \sin^2(\theta_i + \alpha)}{F \sin \theta_i} \quad (2)$$

where Δn is the absolute difference between the ordinary and extraordinary indices of refraction, θ_i is the angle of incidence of the incoming light, and α is the cut angle which is the cut angle from the optical axis to the piezoelectric transducer. Also, through the described interaction the diffracted light goes through a 90° rotation in polarization (*Voloshinov*, 1996). Lastly, a wide aperture is required for an AOTF used for imaging purposes.

In order to be able to perform the required spectral imaging a 10x10 mm aperture imaging quality ATOF was acquired from Brimrose of America. It is optically tuned, meaning designed for a specific wavelength octave, for a range of 600 nm to 1200 nm corresponding to a RF range of 156 to 70 MHz and made from Tellurium Dioxide (TeO_2), a birefringent crystal with indices of refraction at 800 nm of 2.226 and 2.373 for the ordinary and extraordinary modes respectively (*Uchida*, 1971). The extraordinary light is diffracted at 2.7 degrees off of the optical axis of the device. In order to achieve a constant diffraction angle of the first order beam a wedge is attached to the exit aperture of the AOTF to compensate for the angular change that occurs from altering the wavelength. The ordinary light undergoes diffraction but at a nonconstant angle off of the optical axis with respect to wavelength due to the ordinary polarization being uncompensated. It is important to note that it is currently not possible to have both polarizations compensated and as such only one polarization will be captured by ALI.

Characterization of the AOTF is required for calibrated use in an instrument. The AOTF is surrounded by a telocentric front and back end since it will allow the best determination of the central wavelength of the device at a sacrifice of the optimal FWHM results. A 100 mm plano-convex lenses were chosen for both lens to best fill the AOTF aperture. The light source is a 100 W quartz-tungsten halogen bulb that is collimated and passed into the first telocentric lens. The light passes through the AOTF focused then is re-collimated through the second lens. The signal enters a HORIBA iHR320 spectrometer with a maximum spectral resolution of 1.175 nm. The spectral results were imaged on a Synapse 354308 front-illuminated CCD detector with 1024x256 pixels. Images were taken at a set of RFs spaced every 150 kHz from 75 MHz to 160 MHz. For each image the results are spectrally averaged across the rows and a typical spectral measurement result can be seen in the top left panel of Figure 2. The maximum value of each image is taken to be the central wavelength through the AOTF at each respective RF.

The maximum values from each of the images were determined as well as the corresponding wavelengths for these values. It was noted that the curve appears to follow a power function of the form

$$F = a\lambda^b. \quad (3)$$

A linear least squares fit was performed in log space finding the coefficients a and b . The fit

was performed and appeared to match the data quite well but a relative error analysis was performed and it was seen that there was only an agreement better than 0.6% within the testing range and would increase for wavelengths outside the testing region. A better fit was desired to characterize the AOTF's RF-spectral dependence so a modified power function was used in the form of

$$F = a\lambda^{b+c \log \lambda} \quad (4)$$

or a quadratic least squares fit. These results can be seen in the bottom half of Figure 2. The agreement of this form is less than 0.1% throughout the whole wavelength range and the determined RF and wavelength relation can be seen in

$$F = \exp(19.793)\lambda^{-3.381+0.168 \log \lambda} \quad (5)$$

where λ is in nanometers and F is in MHz. It should be noted that even though the AOTF optical range is 600 nm to 1200 nm this analysis only measures wavelengths from 600 nm to 1080 nm. The low quantum efficiency of the CCD and the low near IR emitted from the light source causes wavelengths past this point to be noisy.

The same set of data was used to determine the FWHM for each of the above determined wavelengths. The results of this study are shown in the top right of Figure 2. Wavelengths past approximately 1080 nm are too noisy to be able to determine the FWHM of the signal which is not an issue since ALI uses a CCD camera which also cannot measure radiances past 1080 nm. However, the AOTF spectral resolution is well within the limits that are required in order to determine aerosol extinction in the upper troposphere and lower stratosphere which is discussed under section 1.

The diffraction efficiency was determined using the same recorded data from the previous experiment combined with a second set of spectral measurements. The source used in the previous experiment had its radiance passed through the same experiment set up except with the AOTF and the horizontal polarizer removed. These spectral images contained only the initial light that would be incident to AOTF. By directly comparing the counts from the AOTF experiment to the measured incidence radiance the diffraction efficiency of the AOTF is found with a simple ratio. The determined diffraction efficiency is between 56% to 64% across the measured spectral range and should be noted that the diffraction efficiency changes slightly with respect to incoming angle.

2.2 Optical Design and Performance

The AOTF limits the optical system to only having two practical layouts since the incoming light must enter the device at less than the acceptance angle. This is the maximum angle light can enter the device and still undergo the diffraction interaction. These two layouts are a telecentric and a telescopic system. The telescope or afocal system causes a wavelength gradient to be formed across the field of view of the image whereas the telecentric design overcomes this problem but has a larger FWHM (*Suhre et al.*, 2004) and the differences will

be discussed in the following.

A telecentric layout in both image and object space has an advantage for the imaging quality of the AOTF system. The filtered image has a consistent central wavelength across the entire image with a larger spectral FWHM since the diffracted wavelength is dependent on incident angle as seen in Equation 2. This system does have two inherent issues. First, this method is sensitive to any surface defects of the crystal since the light enters the crystal in focused bundles.

Second, a blurring effect is added to the final image that is dependant on wavelength. The blurring effect is caused by the inherent change of optical path length. First, the optical path between the first two lenes is a fixed distance, however the AOTF is made of TeO_2 or paratellurite and has a high index of refraction. The crystal also has a high dispersive property, or Abbe number, so the index of refraction depends on the wavelength. The dispersive nature gives an apparent change in the optical path length. The change in path length, d , is given by

$$d = \frac{n(\lambda) - 1}{n(\lambda)} t \quad (6)$$

where $n(\lambda)$ is the index of refraction with a wavelength dependance and t is the thickness of the crystal. The AOTF crystal causes the optical path in air to be lengthened by d , and in order to compensate, the length, d , must be added to the path to account for the discrepancy, but this can only be accounted for a specific wavelength. Due to the high dispersive nature defocusing will occur at the image plane for all other wavelengths. In order to correct for this problem additional compensating optics would need to be added or the CCD would need to be actively moved as the wavelengths are being scanned.

In the telescopic system the AOTF has collimated light for each line of sight passing through the device and this has a few fundamental changes that alter the system's imaging quality. First, the light passing through the AOTF from a single line of sight enters the AOTF at the same angle, so the image will have a smaller FWHM than the telecentric counterpart. However, each line of sight will be diffracted with a different fundamental wavelength due to the angular dependance in the AOTF Bragg diffraction wavelength equation (Equation 2). The final image has a smaller spectral bandpass but there will be a wavelength gradient radiating out from the center of the image. Second, since light now passes through the AOTF collimated, the focal point of the image no longer changes with wavelength. Instead a lateral displacement of each line of sight occurs based on the angle of incidence and the diffracted wavelength which causes a slight magnification of the final image. The lateral displacement that occurs is given by the following relation

$$\delta = (n(\lambda) - 1) \frac{t\theta}{n(\lambda)} \quad (7)$$

where δ is the displacement from the original path, but this magnification is a negligible change overall amounting to a maximum 0.07 mm exiting the AOTF.

The telescopic system offers the ability of having an image plane location that is not dependant on the wavelength being imaged and was the ultimate optical choice for ALI.

A telecentric system would lead to an optical layout that would either require mechanical components to move the imaging plane, additional optical elements to counter act the change in the optical path, or have lower resolution spatially. Using mechanic components to move the camera would be an addition failure point and would have to be well calibrated across the wavelength range. For the second solution, an alternative custom lens or prism would need to be created in order to counteract the defocusing effect of the AOTF which would cause further reflections within the system increasing the possibility of significant stray light and well as decreasing the signal throughput and the overall cost. The telescopic design would also allow for a greater focus on spatial resolution that could be achieved which is necessary in order to be able to image the fine structure of aerosol extinction on the order of 100s of metres.

For ALI a simple three lens optical layout was chosen for the system. Two lens before the AOTF to form the telescopic front end and one lens after to focus the light for imaging. Consideration was given when picking the lenses to be able gather a large amount of light and pass it through the AOTF itself in order to reduce the required exposure times during the mission. In order to be able to achieve the increased light throughput a demagnification occurs in the first half of the optical chain resulting in a tightening of the collimated light bundles before entering the AOTF. A slit plate is added at the imaging point between the first two lenses to help reduce unwanted signal from continuing though the optical chain. Before the light enters the AOTF a linear polarizer is used to remove the incoming horizontal polarized light which is the polarization that is not imaged by ALI, removing the 0th and 1st order horizontal polarization. After the AOTF the diffracted wavelength has undergone a 90° rotation in polarization so a second linear polarizer orientated vertically is used to remove 0th order vertical polarization. The diffracted light is bent 2.7° from the optical axis and to compensate, the rest of the optical chain is bent and is passed into an imaging lens which images the signal on a QSI 616s CCD which is a 16 bit output with 1536×1024 $9 \mu\text{m}$ pixels with a 15 count error on the readout. As such ALI measures light that is polarized 90° from the sun or what is known as the vertical polarization. The vertical polarization only account for 10 to 35% of the total incoming radiance. A ray tracing diagram for ALI's optical system was created using the CODE V optical design software and can be seen in Figure 3.

Another minor consideration for the system is the narrower FWHM for each line of sight that passes through the ATOF giving finer spectral resolution however the draw back is that the central wavelength of each line of sight is dependant upon the angle of incident on the AOTF crystal therefore a wavelength gradient appears in the final image. The longest central wavelength occurs in the center of the image and radiates outward towards shorter wavelengths as apparent in Equation 2. Using the telescopic layout ALI would have a wavelength gradient of approximately 7 nm at 650 nm central wavelength and 11 nm at 950 nm, however for a space based instrument with a relatively smaller field of view the gradient could be reduced to as small as 2 nm across the whole image which, at worse, is slightly larger than the FWHM of the AOTF. This effect would be a significant problem for instruments measuring trace gases absorption lines, for example NO_2 , but aerosol is a broadband enhancing feature in which its

effects are smooth in atmospheric measurements from 500 nm well into the IR. The wavelength dependent magnification mentioned earlier only amounts to a change of a approximately 4 pixels or 0.4% in both directions from the inherent magnification from 650 nm to 950 nm and overall this change is considered negligible. The final optical specification for ALI can be found in Table 1.

Instruments that measure the limb have a concern about stray light entering the optical path from paths outside of the systems field of view and significantly contaminating the measurement. Even though the AOTF has good method of removing stray light; having a large amount of stray light would drop the signal to ratio and would not be preferable. A baffle was designed and built using the optimal baffle geometry in order to reduce the unwanted signal will maintaing a height to pitch ratio greater than 0.5 to keep stray light at a minimum (*Fischer et al.*, 2008).

A SolidWorks rendition of the completed version of ALI can be seen in Figure 4. ALI is tilted at 3° from the horizon so ALI's complete 6° vertical field of view would span from a tangent point that would approach the ground to the float altitude.

Effective focal length (mm)	74.3
Front optics magnification	0.67
Back optics magnification	1.27
Field of view (°)	6.0 x 6.0
F-number	7.5
Image Size (mm)	9x9
Image Size (pixels)	1000x1000
Spectral Range (nm)	650-950

Table 1: ALI Final System Optical Parameters.

ALTIUS, a Belgium instrument, uses similar technology as ALI except uses a telecentric optical layout and is designed to measure atmospheric trace gases (*Dekemper et al.*, 2012). Trace gases have narrow absorbtion and emission features that require specific wavelength knowledge. A telecentric layout will give a constant wavelength across the whole field of view to be able to resolve absorbtion features. The optical specifications are similar between the two instruments, however two key differences will be noted. First, the field of view of ALTIUS is smaller at 5.73x5.73° and if ALI would have used a telecentric design then it would have had an similar field of view due to the geometry and optical requirements of the telecentric layout. In a telecentric design the AOTF aperture directly limits your instruments field of view and both systems use an ATOF crystal with an optical aperture of 10x10 mm. Second, the f/# for ALTIUS is 14.32 compared to ALI's 7.5 which allows ALI to increase light throughput at the cost of slightly higher abberations in the final image which allowed for a reduction in ALI's exposure time without making a complicated prototype instrument. The visible channel of ALTIUS was breadboarded and tested by imaging a smoke stack in order determine NO₂ slant column density at 3.5 km away with a 10 second measurement times. It was noted that

improvements to the system would result from methods to reduce the amount the stray light that enters the system as it is non-negligible as well as a reduction in the time in between measurements.

3 Stratospheric Balloon Flight

3.1 Flight Day Conditions and Flight Path

The balloon launch base in Timmins, Ontario is located at 48.47°N 81.33°W and ALI arrived at the base on August 28, 2014 with a launch window from September 8 to 14, 2014. In between the arrival and launch ALI was integrated onto the CNES CARMEN gondola with CARMEN’s systems, including communications and power. ALI was orientated so it would be 90° from the direction of the sun in regard to the gondolas pointing system with an overall southern field of view during the mission. CARMEN is a gondola with pointing capabilities and is piloted from the ground station in Timmins during the flight. The pointing precision is then fine tuned to a pointing of better than 1° with the use of an onboard star tracker.

The flight of CARMEN was delayed due to poor weather conditions during the launch window. On September 20, 2014 at 05:35 UTC (01:35 local time) ALI was launched as the Nimbus 7 mission onboard the CNES CARMEN gondola from the CSA Timmins balloon launch facility. During the launch, the sky was clear with light winds allowing for an safe and uneventful launch. The ascent of CARMEN occurred in darkness and reached its flight altitude of 36.5 km at 8:17 UTC. First light was observed by ALI at 9:39 UTC and recorded measurements until 14:42 UTC when the primary aerosol mission was completed. ALI was powered off at 17:15 UTC during this time ALI recorded measurements for secondary goals, including an azimuth scan. A visualization of the flight path with all major landmarks noted can be found on Figure 5. Temperature profiles for the ambient atmosphere and instrument can be seen in Figure 6. The black curve is the ambient atmospheric temperature surrounding the gondola during the flight from the ECMWF (*Molteni et al.*, 1996). The blue, green, and red are from temperature sensors onboard ALI located on the baffle, camera, and RF driver respectively. The baffle temperature sensor was attached just on the inside of the ALI right by the entrance aperture for the system and monitors the temperature at the front of the system. The camera sensor is attached to the back of the CCD camera and the RF driver sensors measures the temperature of the RF driver. ALI was thermally insulated with styrofoam to keep the system warm, while at the same time direct heating from the sun was a concern and the system was covered in a reflecting material to reduce solar heating.

During the mission, ALI ran in two primary operational modes, a dark mode and a aerosol imaging mode. The first mode, the dark mode was primarily used during ascent and intermittently between aerosol modes. During this mode the filtering of the AOTF is disabled, meaning with no RF signal being applied to the crystal, and as such the wavelength has no dependance on the images. Eight exposures are taken at with 0.05, 0.1, 0.5, 1, 2, 3, 5, 10 second exposure times with the camera shutter operating. The second operational mode was

the aerosol mode and it records 13 measurements which each consisted of a pairs of images. A pair of images, one with the AOTF enabled and one disabled, were gathered for every 25 nm between 650 to 950 nm each measurement set took approximately 25 s to acquire. The exposure times were determined by making ground based measurements of all of the wavelengths in the aerosol mode at a variety of exposure times. This data was used to determine the value at which the well of the CCD would be three quarters full on the ground. Then using the known geometry from the ground and assumed geometry from the balloon in combination with the SASKTRAN model, which will be discussed in subsection 3.3, simulated radiance profiles for both geometries were determined. The determined ground based exposure times were scaled to balloon exposure times by the ratio of the balloon based modeled radiances over the ground based modeled radiances.

3.2 Measurements

The raw flight data (level 0) must be converted to relative radiances (level 1), which is the data normalized to a laboratory measurement value, before they can be used to retrieve aerosol extinction and particle size. The transformation includes removal of dark current, DC bias, stray light, and application of flat fielding calibration. Using the dark images from the ascent of the flight combined with laboratory dark images a table of values based on exposure time and CCD temperature was computed to be able to remove the DC bias and dark current from each raw image accurately. Once completed, the time independent offset caused by the DC bias allows for the rest of the analysis to be preformed in counts per second, simplifying comparisons between images.

Stray light removal has always been difficult in atmospheric instrumentation due to the difficulty in accurately discerning the signal in regards to the stray light contamination. Furthermore, ALI's optical system has the further addition of unwanted light internal to the instrument because of the rejection of one of the polarizations due to the nature of the AOTF. The signal enters the optical system unpolarized, but only one polarization can have a consistent output angle from the AOTF. The entering light is passed through a linear polarizer with an extinction ratio of at least 100,000:1 to remove the unwanted polarization however a small percentage is not absorbed. Furthermore, a second linear polarizer is used after the AOTF to reject all of the unwanted radiance that did not meet the Bragg criteria and once again a small percentage of this radiance is not absorbed. The active filtering of the AOTF allows for an image to be measured when the filtering device is disabled, which is with no applied RF, allowing only the stray light to be captured by the measurement known as a 'dark image', and during ALI's aerosol mode a 'dark image' was captured before every measurement image. By removing the 'dark images' from the signal-stray light contaminated images the end result is images that only contain the measured signal. The previous method was tested in the lab with a known source with even illumination across the field of view of the system, the resulting final image is left with a even decrease in intensity radiating from the center of the field of view as its expected with the known vignetting of ALI. The vignetting is caused by the aperture of

the AOTF itself, by using a simple optical layout as chosen for the prototype the larger the angle for the field of view the more light that get blocked by the AOTF's apperture causing a known vignetting for the images and will need to be calibrated out of any measurements.

To finalize the data into level 1 relative radiances a flat fielding calibration is preformed, which involves taking all the incoming fielding of views and normalize them to be unity. The flat fielding coefficients needed to be applied not only spatially across each image but spectrally across a series of images. To determine the flat fielding coefficients ALI was set up in the lab viewing the light emitted from a quartz-tungsten halogen bulb that is passed through a diffusing plate to give an evenly distributed signal across the entire field of view of ALI. Measurements were taken at varying integration times at every wavelengths used in ALI's aerosol mode. Since ALI is most sensitive at 750 nm it was chosen to be the reference point. The center pixels were used since they experience little vignetting, specifically the center 25 by 25 pixels. All pixels for every image were normalized to the mean of the reference point pixels. From these normalized images, the flat fielding coefficients were the values needed to multiply the normalized images to achieve unity. The coefficients were then averaged for each wavelength and applied to the flight data to yield the final relative radiances. In Figure 7 image number 212 at 750 nm is shown with a before and after comparison. The before image is the raw data directly after the mission with no calibrations preformed and the after image in the bottom panel is the same image after the all the calibrations have been applied.

To increase the precision of the measurements from the flight the images were averaged in cells of 25 horizontal and 10 vertical pixels decreasing the resolution but greatly reduces the noise on the radiances. Thus resulting in a reduction of the vertical resolution of ALI to 700 m at 15 km to 200 m at 30 km. Furthermore, a loss of resolution was speculated to occur in the flight data because of the drastic change in temperature of the optics during the flight which is a secondary reason for the pixel averaging. A large number of vertical profiles with a horizontal dependance are determined The final relative radiances can be seen in Figure 8 and Figure 9.

3.3 Retrievals

The modeled radiances were computed with the SASKTRAN radiative transfer engine (*Bourassa et al.*, 2008) for for High Resolution (SASKTRAN-HR) (*Zawada et al.*, 2015) measurements using the newly developed polarization module (*Dueck S. and Degenstein*). The model uses a given atmospheric state to solve the radiative transfer equation to determine the final radiance at the observer that follows

$$I(s_1) = I(s_0)e^{-\tau(s_0,s_1)} + \int_{s_0}^{s_1} k(s)J(s)e^{-\tau(s,s_1)}ds \quad (8)$$

where $I(s_1)$ is the radiance at the observer through a path from s_0 to s_1 , the first term is the contribution of light that is attenuated along the line of sight from the sun to the observer at s_1 . The second term takes the source term, $J(s)$, which is the radiance scattered into the

line of sight and integrates the path along line of sight with attenuation to determine the scattering contribution to the final radiance. The extinction, given by $k(s)$, is the sum of the number density, $n(s)$, multiplied by the scattering cross section, $\sigma(s)$, over all species and τ is the optical depth. The polarized output of SASKTRAN-HR gives the Stokes vectors for the radiance on its internal coordinate grid which can be retrieved from the model and then rotated to match ALI's coordinate system.

The relative radiance level 1 data from ALI are used create a measurement vectors, y , in the following form

$$\mathbf{y} = \log \left(\frac{\mathbf{I}(\mathbf{z}, \lambda)}{I(z_{ref}, \lambda)} \right) - \log \left(\frac{\mathbf{I}_{rayleigh}(\mathbf{z}, \lambda)}{I_{rayleigh}(z_{ref}, \lambda)} \right) \quad (9)$$

where $\mathbf{I}(\mathbf{z}, \lambda)$ is the measured radiance from ALI and $I(z_{ref}, \lambda)$ is the radiance at a reference altitude used to normalize the signal from a high altitude where there is little aerosol contribution, for ALI the highest possible altitude where the signal is above the noise threshold is around 25-30 km tangent height. The second term is modeled values from SASKTRAN-HR with only background neutral atmosphere to remove the rayleigh signal from the measured radiances which is done to improve the speed of the convergence of the retrieval. The ALI measurement vector is similar to the measurement vector used for the OSIRIS retrieval (*Bourassa et al.*, 2007; *Bourassa et al.*, 2011). A base aerosol state or a priori, \mathbf{x} , for aerosol extinction profile is used in the SASKTRAN-HR model. The forward model vector is constructed similarly to the measurement vector and follows.

$$\mathbf{F}(\mathbf{x}) = \log \left(\frac{\mathbf{I}_{mod}(\mathbf{z}, \lambda)}{I_{mod}(z_{ref}, \lambda)} \right) - \log \left(\frac{\mathbf{I}_{rayleigh}(\mathbf{z}, \lambda)}{I_{rayleigh}(z_{ref}, \lambda)} \right) \quad (10)$$

where $I_{mod}(z, \lambda)$ is the modeled radiance for the measurement and $I_{mod}(z_{ref}, \lambda)$ is the measurement at the same reference altitude for normalization. The forward model is used in combination with the measurement vector to update the extinction profile in the Multiplicative Algebraic Reconstruction Technique (MART) algorithm which was developed for use in the OSIRIS retrievals (*Bourassa et al.*, 2012) with the following iterative technique

$$x_{i+1} = x_i \sum_j \frac{y_j}{F(z_j)} W_{ij} \quad (11)$$

where x_i is the aerosol extinction at each shell altitude, i and j is the tangent point from the measurements. W_{ij} is the weighting matrix that relates the importance of each measurement vector to each shell altitude. This method was outlined by *Degenstein et al.* (2009) and allows for fast retrievals without calculating a Jacobian.

An error estimation was also needed to be able to fully classify the capabilities of ALI and was performed using a perturbation method. Once a retrieval has been completed for an image the result is used to estimate the error in the returned extinction. For each altitude, the measurement vector is perturbed by the error resulting from the level 0 to level 1 data conversion parameters used and from the readout electronics. The MART retrieval is rerun

and the change of the extinction is determined. These are compiled to form a gain matrix, \mathbf{G} , with size is n by m which are the shell altitudes and the tangent altitudes grids respectively, with elements g_{ij} . The error at each retired altitude is then given by

$$e_i = \left(\sum_j g_{ij} g_{ji} \right)^{0.5} \quad (12)$$

which gives the error at at retrieved altitude.

Once the retrieval has been preformed for a complete series of wavelengths a determination for the Angström exponent occurs which is preformed in a similar manor as outlined by *Rault and Loughman* (2013) for the OMPS aerosol particle size retrieval. The Angström exponent is an approximation to Mie scattering since the value of the Angström exponent, α , is an analogy to particle size. Lower Angström exponents correspond to larger particle sizes and vice versa for small particle sizes. Since the measurements observe relatively the same atmosphere over the time of one complete aerosol cycle; the differences between extinction ratios at the different wavelength can be used to gather a understanding of aerosol particle size in the form

$$\frac{n\sigma}{n_0\sigma_0} = \left(\frac{\lambda}{\lambda_0} \right)^{-\alpha} \quad (13)$$

where n is the aerosol concentration, and σ is the scattering cross section. Since the cross section of aerosol is dependent on not only the wavelength but also the particle size and distribution the Angström exponent give some information on particle size distribution. For the retrieval described here a single mode log-normal distribution is assumed with a mode radius, r_g , and the mode width, σ_g , which is the same as the OSIRIS version 5.07 aerosol product. At each retrieved altitude the extinction from each wavelength is used to fit the Angström exponent. Then the median value from all the altitudes is used as the new size parameters in next iteration of the retrieval. The mode is set at a constant 1.6 and the mode radius is allowed to vary to match the Angström exponent as retrieved. The mode radius is updated and the process is repeated, re-retrieving the extinction profiles, until the Angström exponent converges. At the end of the last iteration the Angström exponent profile is reported as the final particle size product.

3.4 Results

In order to be able to use the ALI data in the MART method certain quantities were needed for the model; albedo, ozone concentration and cross sections, and aerosol cross sections. The albedo is required since an absolute radiance calibration was never preformed with ALI and the albedo cannot be determined directly form ALI's measurements, which is important since the amount of ground scatter in SASKTRAN-HR must correspond to the ground scatter during ALI's flight. Second, the ozone absorbtion features from the Chappuis band appear in the ALI measurements from 650 to 750 nm. The absorbtion must be accounted for to not artificially

change the determined aerosol profiles. The ozone profiles were acquired from OSIRIS. Five scans that were within 48 hours of the balloon flight and within 500 km of the launch facility were averaged together to be the ozone profile used in the SASKTRAN-HR model, with cross sections from *Burrows et al.* (1999). The albedo is from the ADAM database which has monthly values for albedo over the surface on earth at a resolution of $0.1^\circ \times 0.1^\circ$ grid at 1 nm spectral resolution (*Muller et al.*, 2013). The aerosol cross sections come from the Mie scattering derivation that was originally proposed by Mie and was implemented efficiently by *Wiscombe* (1980). For the purpose of the retrieval an a priori was used with a mode radius of $0.08 \mu\text{m}$ and a mode width of 1.6 which is considered a standard size distribution for aerosol (*Deshler et al.*, 2003).

The complete mission consisted of 216 images that were recorded in illuminated conditions. The MART retrieval method was run on a select signal set for the purpose of the analysis, specifically the last complete set of images from 650 to 950 nm consisting of images 204-216. They were chosen due to being the last in the mission and the sun was the highest in the sky giving the brightest atmosphere leading to the best signal to noise during the mission. A sample of the retrieval set can be observed in Figure 10 which highlights the 725, 825, and 925 nm retrievals. The left panels shows the measurement vector from ALI in black with the forward model radiance profile from SASKTRAN-HR in blue. For each of the wavelengths, the algorithm determines altitudes where the value of the measurement vector is less than the known noise and does not allow aerosol to be retrieved in those regions. Instead the scaling factor, given by $\alpha = yF^{-1}$ is scaled to the aerosol profile above and below the last retrieved point to keep the aerosol profile smooth, as discontinuities are nonphysical and can lead to a convergence failure in the MART algorithm. The middle panel shows the convergence between the measurement vector and the forward model result. For the center wavelengths, being 700-875 nm, a difference of less than 2% is seen from 12 to 22 km with a few outliers.

The aerosol profile for the three wavelengths is shown in blue with the shading representing the error for the retrieval strictly based on measurement error and neglecting any model and atmospheric state errors. The green curve is the average 750 nm aerosol extinction profiles of the same five OSIRIS scans used for the ozone profile and red is the 750 nm aerosol extinction from SALOMON (*Berthet et al.*, 2002) which was launched from the Timmins balloon base as the Nimbus 5 mission on September 12, 2014. The aerosol extinction for ALI and OSIRIS are within 50% of each other for a majority of the 725 nm and 825 nm profile while the upper altitude of the 925 nm are considerable different, in some places over 100% difference occurs. However, agreement is best for 750 nm profile and displays a reasonable spectral dependance to also should be noted that the three instruments follow the same overall shape. First, a bend in the profile occurs at approximately 25 km, then increases approximately linearly until 15 km where aerosol extinction leaves the linear trend and forms the peak of the measurement. The agreement in shape is an excellent result for the ALI mission due to the good numerical comparison with OSIRIS as well as the overall profile shape agreement between all three instruments.

The particle size method was used as outlined in the previous section using the Angström exponent to retrieve one instance of size distribution during the ALI mission. The first panel of Figure 11 shows the median Angström exponent that was determined after each iteration and convergence can be seen after a couple iterations. The particle size determined for ALI in the last complete set of aerosol images can be seen in the second panel of Figure 11 which yields a final Angström exponent of between 2 at the high altitudes (20 km). Assuming a mode width of 1.6 yields a mode radius of 0.12 μm .

In order to determine the Angström exponent a least squares fit was used for all usable wavelengths at each altitude. A wavelength at a altitude was rejected if the forward model at that shell altitude was not within 3% of the measurement vector. In the case shown in Figure 11, the 20.5 km shell altitude, only 8 of the 13 possible wavelength contributed to the Angström exponent.

4 Conclusions and Future Prospects

A description of the prototype ALI using an AOTF for active filtering in the visible to near IR with a telescopic optical layout with the purpose to measure aerosol extinction from the upper troposphere and lower stratosphere with a high vertical and horizontal resolution with images was presented here. The AOTF has fast stabilization times and the ability to disable the filter gives an excellent method to remove stray light from the final measurements. ALI is able to measure aerosol microphysics in remote atmospheric sensing.

ALI was tested on board the CARMEN gondola from the balloon launch facility at Timmins, Ontario. Aerosol extinction profiles were determined and had good comparisons to ORISIS and SALOMON in profile shape. The absolute extinction values are different by a large amounts but could be attributed to the large amount of unaccounted systematics in the retrieval algorithm. Overall ALI performed admirably and verified the use of this technology for future atmospheric remote sensing missions.

Future upgrades to ALI would include the realignment of the optics for flight temperatures to allow for higher resolution measurements and retrievals. Secondly, replacing the CCD currently used on ALI with a camera with faster readout would allow for a higher quantity of data to be taken by reducing the approximate 25 s readout time down to a smaller value necessary for a satellite missions on the order of 1 s per wavelength.

This work would have not been possible without funding from the CSA to design and build ALI through the FAST program as well as the CSA building and managing the launch facility in Timmins, Ontario. Also, thanks to CNES for funding and overseeing the launches at Timmins in 2014. As well thanks Nick Lloyd for help in development of the flight code, without their efforts this work would have not been accomplished.

References

- Barth, C., D. Rusch, R. Thomas, G. Mount, G. Rottman, G. Thomas, R. Sanders, and G. Lawrence (1983), Solar mesosphere explorer: Scientific objectives and results, *Geophysical research letters*, *10*, 237–240.
- Berthet, G., J.-B. Renard, C. Brogniez, C. Robert, M. Chartier, and M. Pirre (2002), Optical and physical properties of stratospheric aerosols from balloon measurements in the visible and near-infrared domains. i. analysis of aerosol extinction spectra from the amon and salomon balloonborne spectrometers, *Applied optics*, *41*, 7522–7539.
- Bourassa, A. E., D. A. Degenstein, R. L. Gattinger, and E. J. Llewellyn (2007), Stratospheric aerosol retrieval with optical spectrograph and infrared imaging system limb scatter measurements, *Journal of Geophysical Research (Atmospheres)*, *112*, D10217, doi:10.1029/2006JD008079.
- Bourassa, A. E., D. A. Degenstein, and E. J. Llewellyn (2008), SASKTRAN: A spherical geometry radiative transfer code for efficient estimation of limb scattered sunlight, *Journal of Quantitative Spectroscopy and Radiative Transfer*, *109*, 52–73, doi:10.1016/j.jqsrt.2007.07.007.
- Bourassa, A. E., C. A. McLinden, C. E. Sioris, S. Brohede, A. F. Bathgate, E. J. Llewellyn, and D. A. Degenstein (2011), Fast no₂ retrievals from odin-osiris limb scatter measurements, *Atmospheric Measurement Techniques*, *4*, 965–972, doi:10.5194/amt-4-965-2011.
- Bourassa, A. E., L. A. Rieger, N. D. Lloyd, and D. A. Degenstein (2012), Odin-OSIRIS stratospheric aerosol data product and SAGE III intercomparison, *Atmospheric Chemistry & Physics*, *12*, 605–614, doi:10.5194/acp-12-605-2012.
- Bourassa, A. E., A. Robock, W. J. Randel, T. Deshler, L. A. Rieger, N. D. Lloyd, E. T. Llewellyn, and D. A. Degenstein (2012), Large volcanic aerosol load in the stratosphere linked to asian monsoon transport, *Science*, *337*, 78–81.
- Bovensmann, H., J. Burrows, M. Buchwitz, J. Frerick, S. Noël, V. Rozanov, K. Chance, and A. Goede (1999), Sciamachy: Mission objectives and measurement modes, *Journal of the Atmospheric Sciences*, *56*, 127–150.
- Burrows, J., A. Richter, A. Dehn, B. Deters, S. Himmelmann, S. Voigt, and J. Orphal (1999), Atmospheric remote-sensing reference data from gome2. temperature-dependent absorption cross sections of o₃ in the 231–794nm range, *Journal of Quantitative Spectroscopy and Radiative Transfer*, *61*, 509–517.
- Chang, I. C. (1974), Noncollinear acousto-optic filter with large angular aperture, *Applied Physics Letters*, *25*, 370, doi:10.1063/1.1655512.

- Degenstein, D., A. Bourassa, C. Roth, and E. Llewellyn (2009), Limb scatter ozone retrieval from 10 to 60 km using a multiplicative algebraic reconstruction technique, *Atmospheric Chemistry and Physics*, 9, 6521–6529.
- Dekemper, E., N. Loodts, B. V. Opstal, J. Maes, F. Vanhellemont, N. Mateshvili, G. Franssens, D. Pieroux, C. Bingen, C. Robert, L. D. Vos, L. Aballea, and D. Fussen (2012), Tunable acousto-optic spectral imager for atmospheric composition measurements in the visible spectral domain, *Applied Optics*, 51, 6259–6267, doi:10.1364/AO.51.006259.
- Deshler, T., M. Hervig, D. Hofmann, J. Rosen, and J. Liley (2003), Thirty years of in situ stratospheric aerosol size distribution measurements from laramie, wyoming (41 n), using balloon-borne instruments, *Journal of Geophysical Research: Atmospheres (1984–2012)*, 108.
- Dittman, M. G., J. W. Leitch, M. Chrisp, J. V. Rodriguez, A. L. Sparks, B. K. McComas, N. H. Zaun, D. Frazier, T. Dixon, R. H. Philbrick, et al. (2002), Limb broad-band imaging spectrometer for the npoess ozone mapping and profiler suite(omps), *Proceedings of SPIE*, vol. 4814, 120–130.
- Dueck S., A. E., Bourassa, and D. A. Degenstein (), SASKTRAN-HR Polarization Module, *In Preparations*.
- Fischer, H., M. Birk, C. Blom, B. Carli, M. Carlotti, T. v. Clarmann, L. Delbouille, A. Dudhia, D. Ehhalt, M. Endemann, et al. (2008), Mipass: an instrument for atmospheric and climate research, *Atmospheric Chemistry and Physics*, 8, 2151–2188.
- Kiehl, J. T., and B. P. Briegleb (1993), The relative roles of sulfate aerosols and greenhouse gases in climate forcing, *Science*, 260, 311–314, doi:10.1126/science.260.5106.311.
- Llewellyn, E., N. D. Lloyd, D. A. Degenstein, R. L. Gattinger, S. V. Petelina, A. E. Bourassa, J. T. Wiensz, E. V. Ivanov, I. C. McDade, B. H. Solheim, J. C. McConnell, C. S. Haley, C. von Savigny, C. E. Sioris, C. A. McLinden, E. Griffioen, J. Kaminski, W. F. J. Evans, E. Puckrin, K. Strong, V. Wehrle, R. H. Hum, D. J. W. Kendall, J. Matsushita, D. P. Murtagh, S. Brohede, J. Stegman, G. Witt, G. Barnes, W. F. Payne, L. Piche, K. Smith, G. Warshaw, D. L. Deslauniers, P. Marchand, E. H. Richardson, R. A. King, I. Wevers, W. McCreath, E. Kyrola, L. Oikarinen, G. W. Leppelmeier, H. Auvinen, G. Megie, A. Hauchecorne, F. Lefevre, J. de La Noe, P. Ricaud, U. Frisk, F. Sjöberg, F. von Scheele, and L. Nordh (2004), The OSIRIS instrument on the Odin spacecraft, *Canadian Journal of Physics*, 82, 411–422, doi:10.1139/p04-005.
- McCormick, M. (1987), Sage ii: an overview, *Advances in space research*, 7, 219–226.
- McCormick, M., P. Hamill, W. Chu, T. Swissler, L. McMaster, and T. Pepin (1979), Satellite studies of the stratospheric aerosol, *Bulletin of the American meteorological Society*, 60, 1038–1046.

- Molteni, F., R. Buizza, T. N. Palmer, and T. Petroliaigis (1996), The ecmwf ensemble prediction system: Methodology and validation, *Quarterly Journal of the Royal Meteorological Society*, *122*, 73–119.
- Muller, J.-P., P. Lewis, F.-M. Bréon, C. Bacour, I. Price, L. Chaumat, P. Prunet, L. Gonzales, C. Schlundt, M. Vountas, et al. (2013), A surface reflectance database for esas earth observation missions (adam), *Living Planet Symposium*.
- Neely, R. R., P. Yu, K. H. Rosenlof, O. B. Toon, J. S. Daniel, S. Solomon, and H. L. Miller (2014), The contribution of anthropogenic so₂ emissions to the asian tropopause aerosol layer, *Journal of Geophysical Research: Atmospheres*, *119*, 1571–1579, doi:10.1002/2013JD020578.
- Rault, D. F., and R. P. Loughman (2013), The omps limb profiler environmental data record algorithm theoretical basis document and expected performance, *Geoscience and Remote Sensing, IEEE Transactions on*, *51*, 2505–2527.
- Solomon, S., J. S. Daniel, R. R. Neely, J.-P. Vernier, E. G. Dutton, and L. W. Thomason (2011), The persistently variable background stratospheric aerosol layer and global climate change, *Science*, *333*, 866–870, doi:10.1126/science.1206027.
- Suhre, D. R., L. J. Denes, and N. Gupta (2004), Telecentric confocal optics for aberration correction of acousto-optic tunable filters, *Applied Optics*, *43*, 1255–1260, doi:10.1364/AO.43.001255.
- Thomason, L. W., and G. Taha (2003), Sage iii aerosol extinction measurements: Initial results, *Geophysical research letters*, *30*.
- Uchida, N. (1971), Optical properties of single-crystal paratellurite (teo₂), *Phys. Rev. B*, *4*, 3736–3745, doi:10.1103/PhysRevB.4.3736.
- Vernier, J.-P., L. W. Thomason, J.-P. Pommereau, A. Bourassa, J. Pelon, A. Garnier, A. Hauchecorne, L. Blannot, C. Trepte, D. Degenstein, and F. Vargas (2011), Major influence of tropical volcanic eruptions on the stratospheric aerosol layer during the last decade, *Geophysical Research Letters*, *38*, n/a–n/a, doi:10.1029/2011GL047563, 112807.
- Voloshinov, V. (1996), Spectral and polarization analysis of optical images by means of acousto-optics, *Optics Laser Technology*, *28*, 119–127, doi:10.1016/0030-3992(95)00079-8.
- Voloshinov, V. B., K. B. Yushkov, and B. B. J. Linde (2007), Improvement in performance of a TeO₂ acousto-optic imaging spectrometer, *Journal of Optics A: Pure and Applied Optics*, *9*, 341–347, doi:10.1088/1464-4258/9/4/006.
- Winker, D. M., J. R. Pelon, and M. P. McCormick (2003), The calipso mission: Spaceborne lidar for observation of aerosols and clouds, 1–11.

- Wiscombe, W. J. (1980), Improved mie scattering algorithms, *Applied optics*, *19*, 1505–1509.
- Zawada, D. J., S. R. Dueck, L. A. Rieger, A. E. Bourassa, N. D. Lloyd, and D. A. De-
genstein (2015), High resolution and monte carlo additions to the sasktran radiative
transfer model, *Atmospheric Measurement Techniques Discussions*, *8*, 3357–3397, doi:
10.5194/amtd-8-3357-2015.

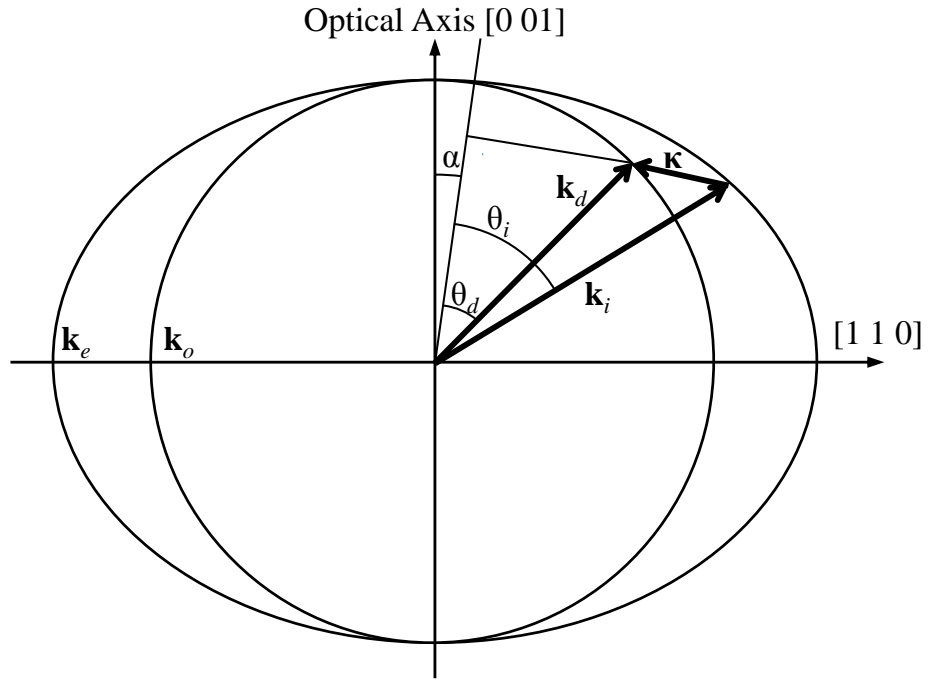


Figure 1: The wave vectors generated by the AOTF experiment. From Equation 1, the incident wave vector, diffracted wave vector, and acoustic wave vector are shown.

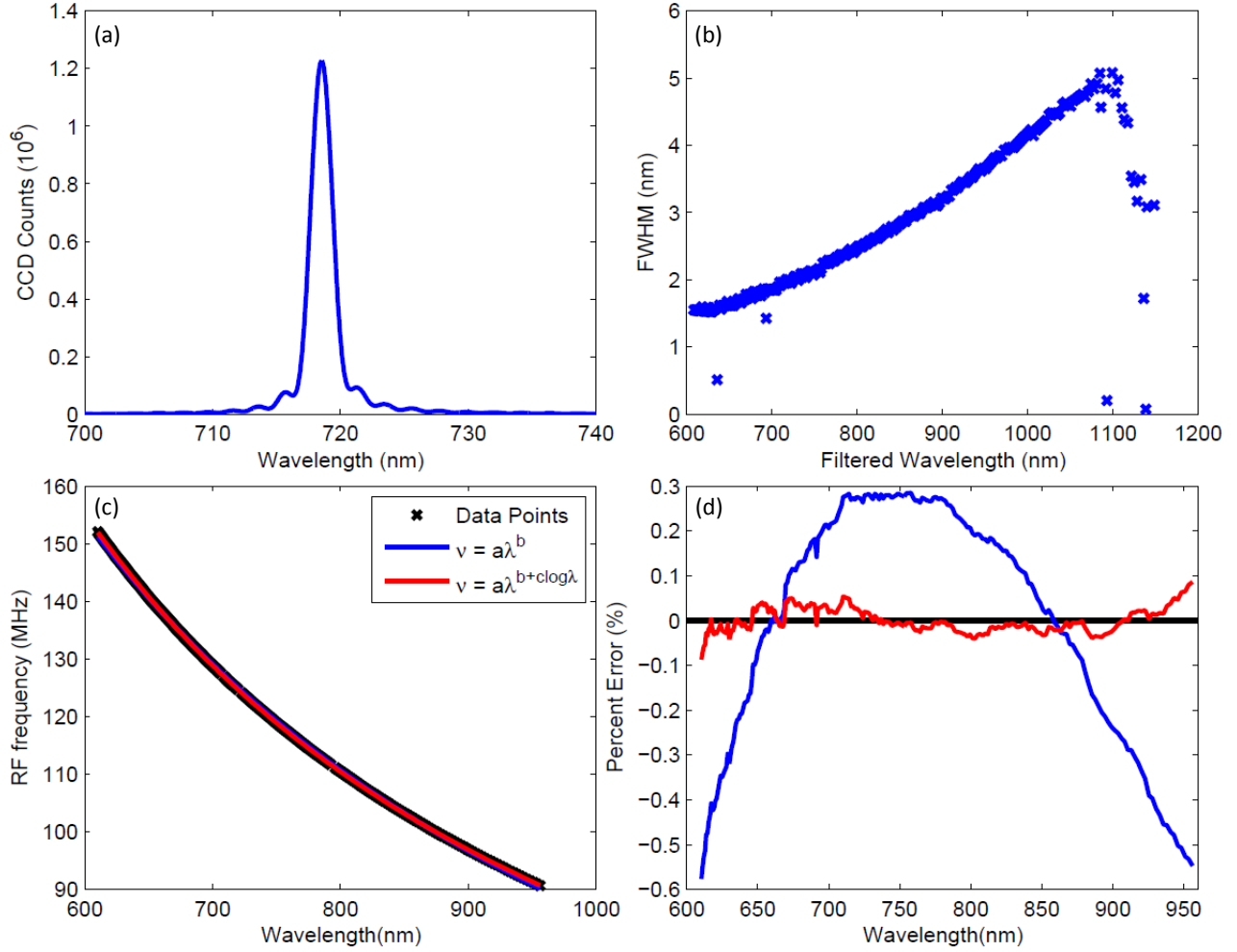


Figure 2: (a) A standard image taken from the AOTF calibration experiment when the tuning frequency of the AOTF was at 124.96 MHz. (b) The FWHM for each of the determined wavelengths for the AOTF. The FWHM at 600 nm is 1.5 and as the wavelengths get longer the FWHM increases to 4.9 at 1080 nm. (c) The calibration curves for the AOTF RF versus the diffracted wavelength which contains the data points recorded and two best fit curves. (d) The percent error with respect to the measured frequency for the two best fit curves in the previous panel.

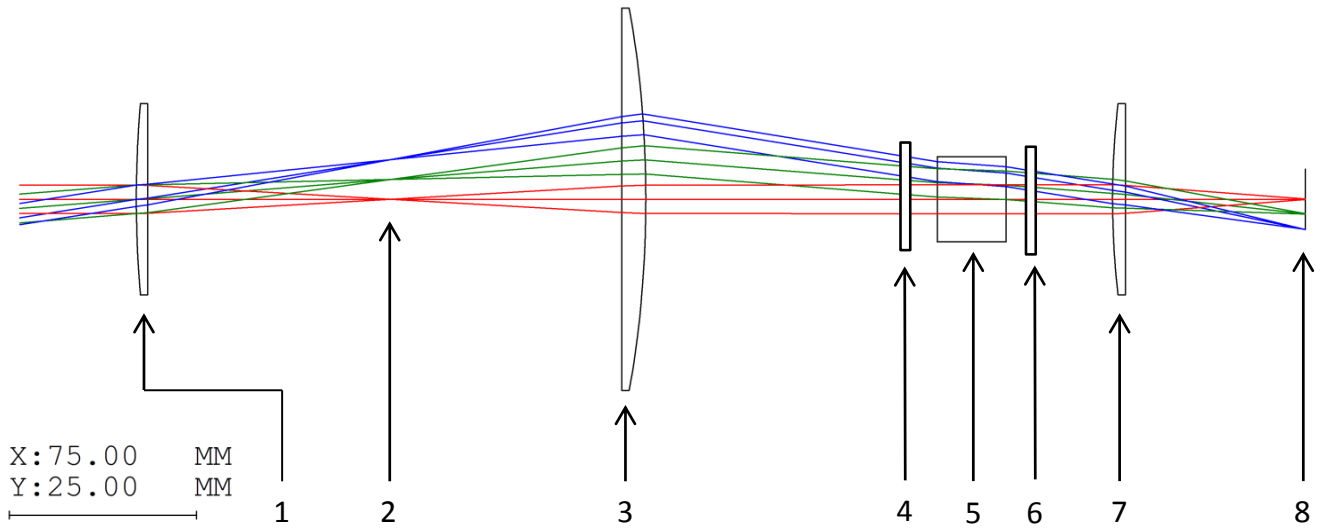


Figure 3: Ray Tracing diagram of the telescopic lens system for ALI simulated by Code V optical design software. The elements in the system are the following: (1) 150 mm focal length plano-convex lens. (2) Slit plate. (3) 100 mm focal length plano-convex lens. (4) Vertical linear polarizer. (5) Brimrose AOTF. (6) Horizontal linear polarizer. (7) 50.4 mm focal length plano-convex lens. (8) Imaging plane.

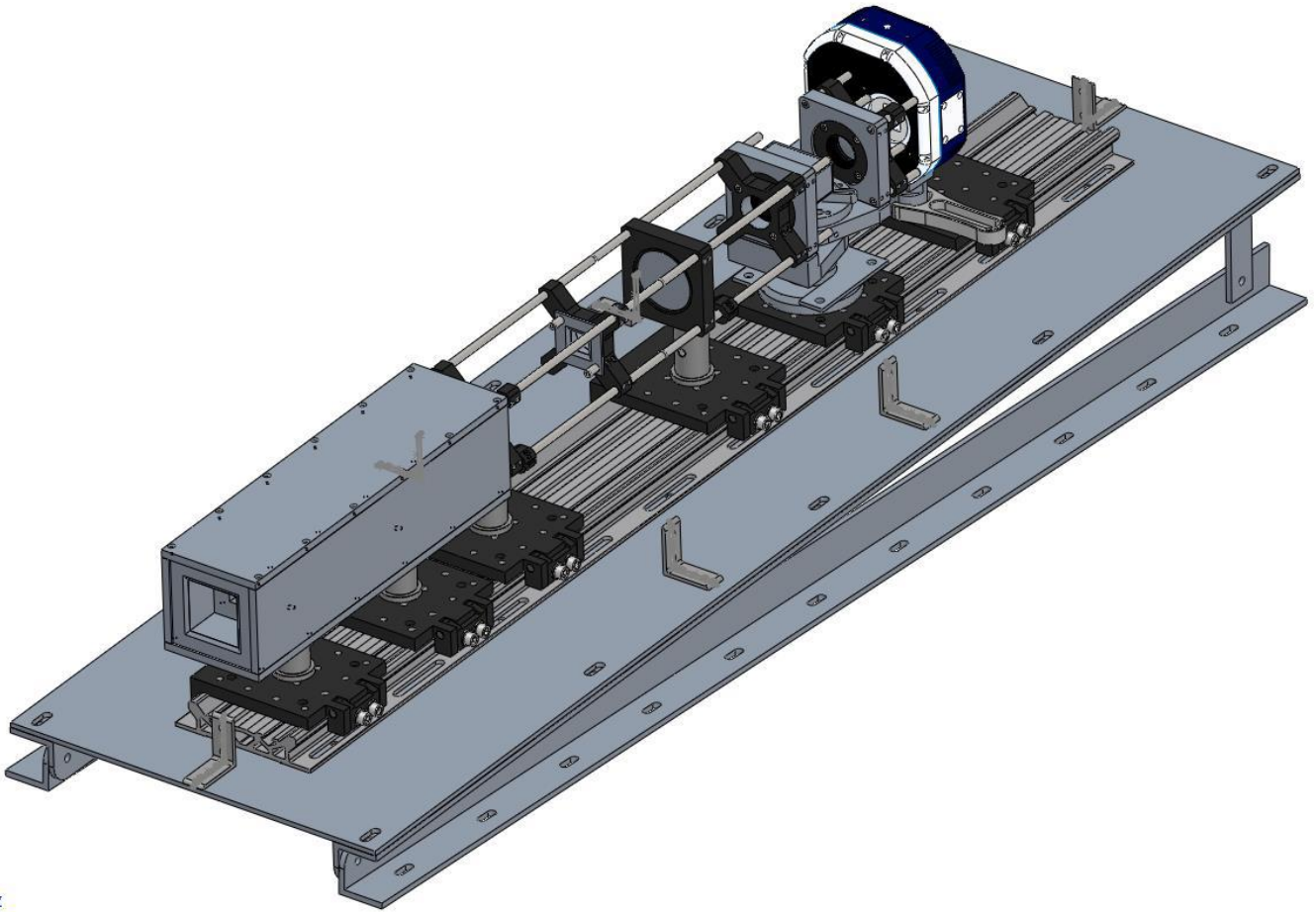


Figure 4: An isometric view of the complete ALI system with the baffle and field of view slant. Light tight case absent from diagram.



Figure 5: The GPS data from ALI during the Nimbus 7 mission generated via Google Earth. The colour of the line represents the absolute speed of the gondola during the mission. Important landmarks are noted on the image. The end of mission represent the end of the aerosol mission. No GPS data was collected from ALI after power down.

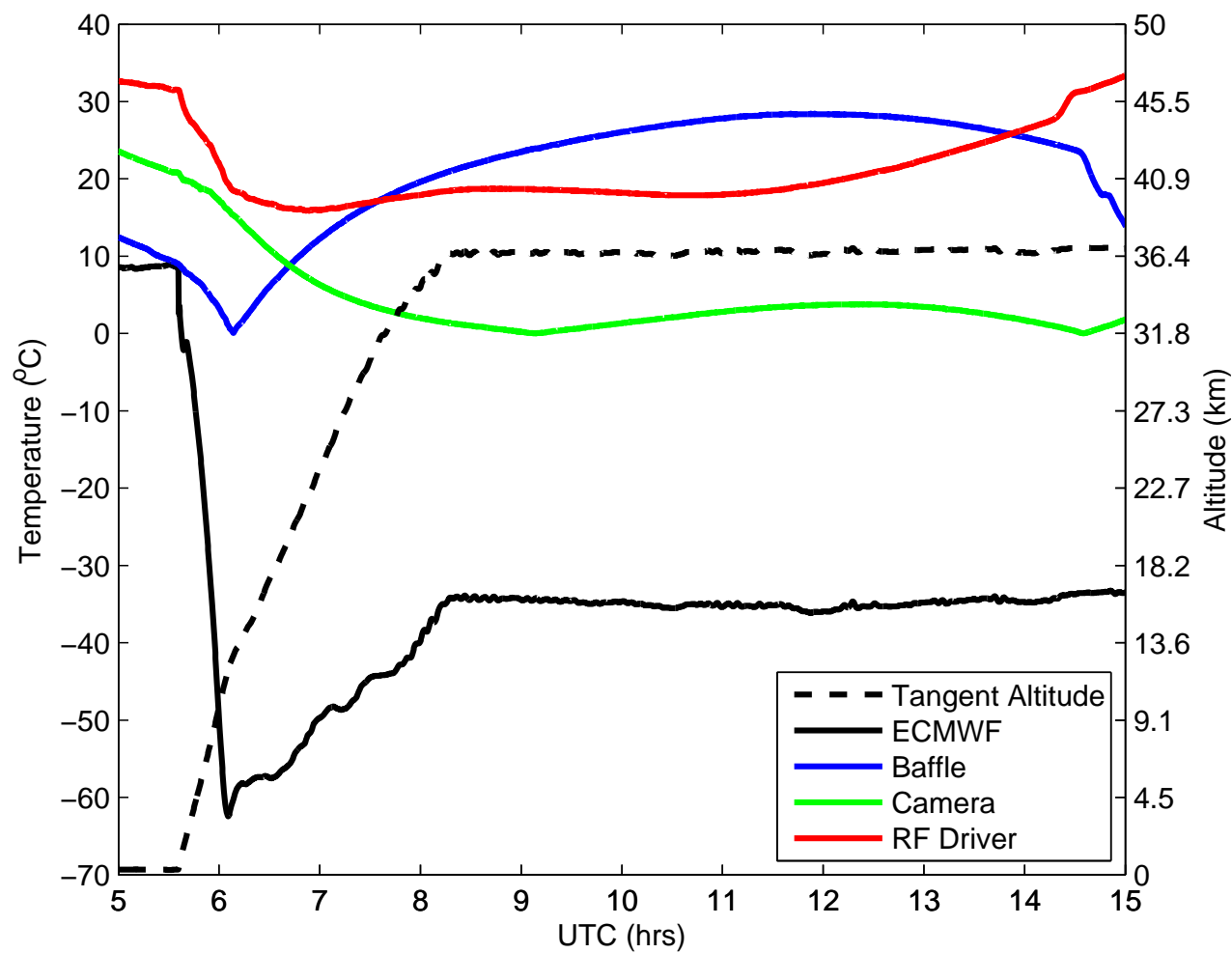


Figure 6: Temperature and altitude profiles from the NIMBUS 7 flight.

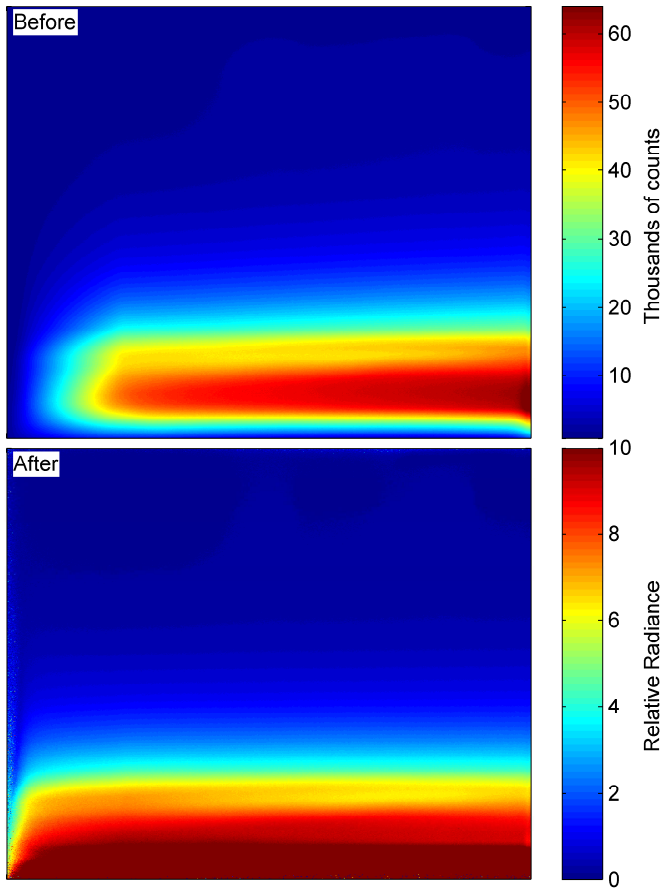


Figure 7: Comparison of the same image, image number 212, at 750 nm. The upper panel is the raw level 0 data and the lower panel is the relative radiance level 1 data.

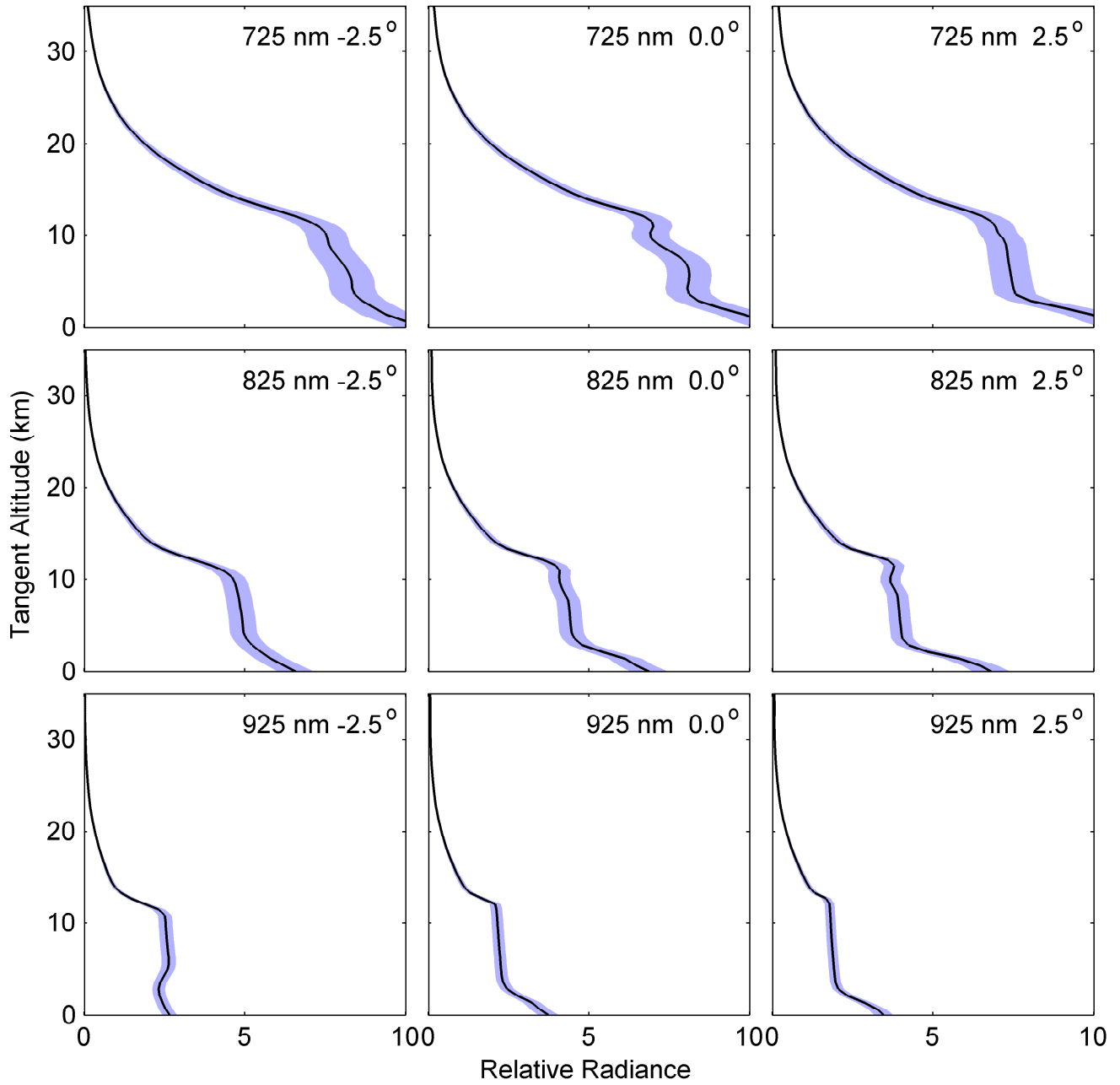


Figure 8: Level 1 relative radiances as measured from ALI at approximately 14:20 UTC (images number 207, 211, and 215) looking 90° from the sun facing southwards. The top middle, and bottom row are measurements taken at 725, 825, and 925 nm respectively. The center column is viewing the atmosphere directly in front of ALI, While the left column is looking to the left at -2.5° and the right at 2.5°.

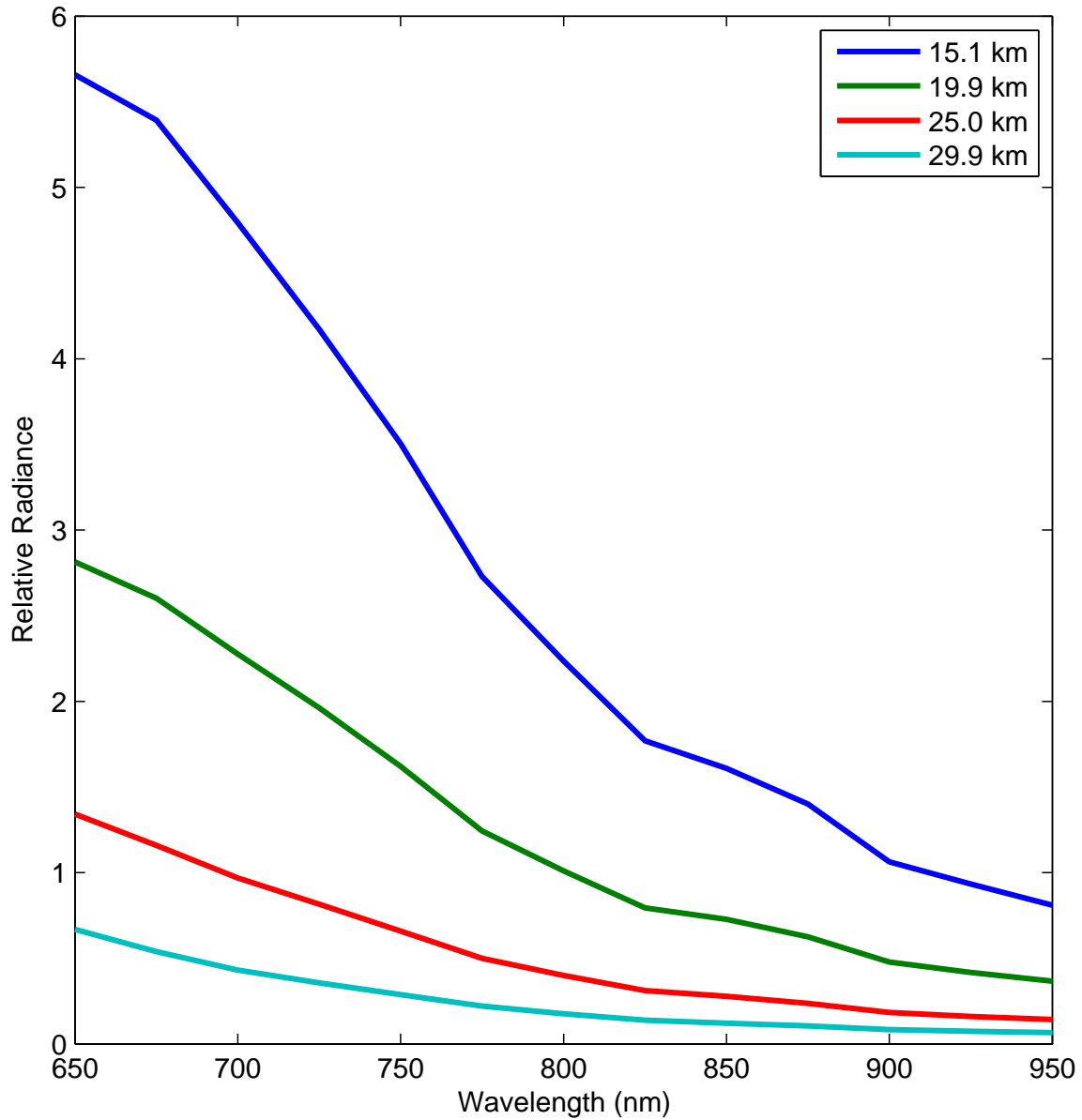


Figure 9: Level 1 relative radiances spectrally from 650 nm to 950 nm as measured from ALI at approximately 14:20 UTC consisting of images number 204 to 216 looking 90° from the sun facing southwards. These spectral profiles are presented at several tangent altitudes with a horizontal field of view of 0°.

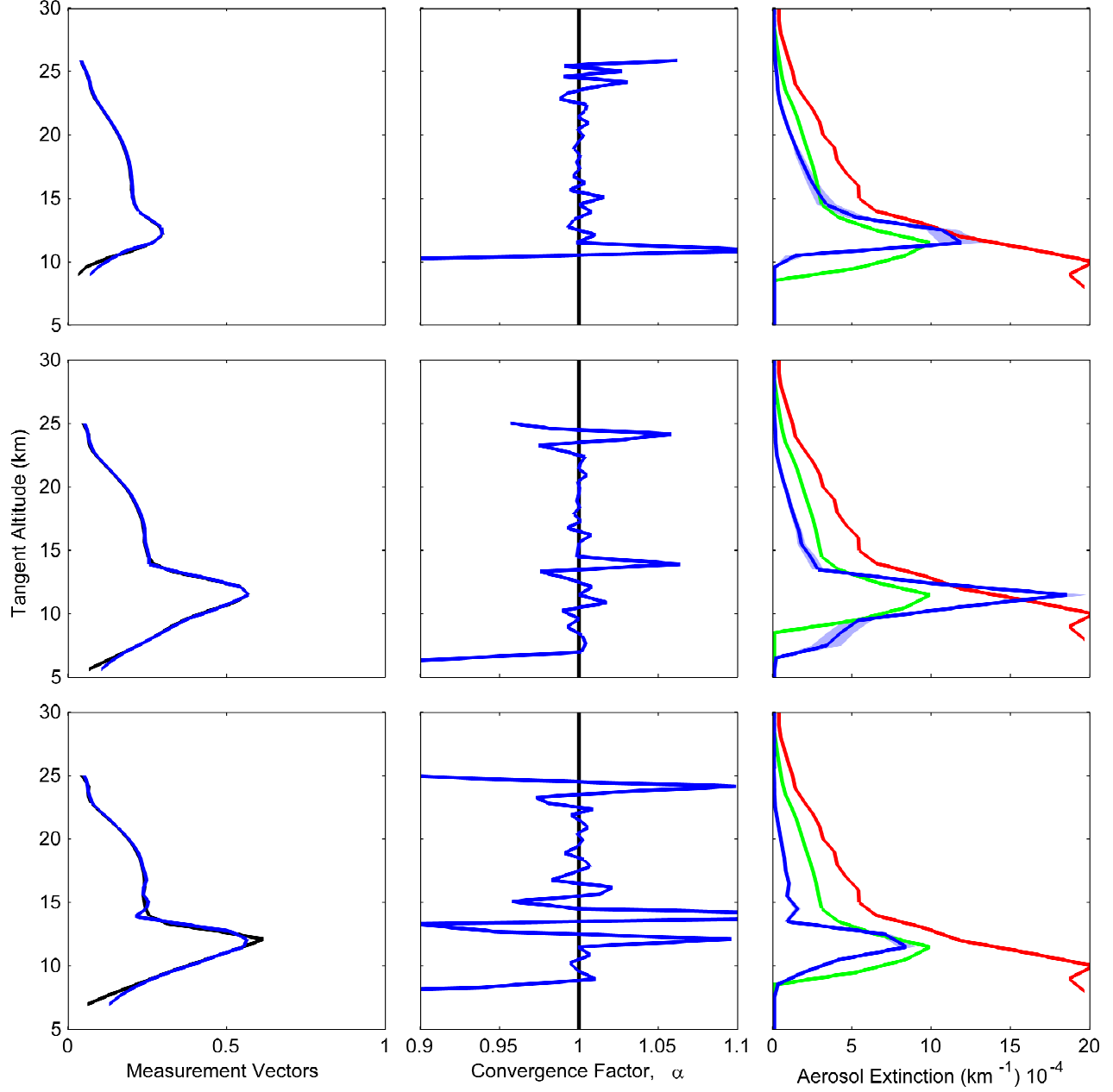


Figure 10: An example of three aerosol retrievals from images 207, 211, and 215, with center wavelengths of 725, 825, and 925 nm respectively and vertically displayed in the figure from top to bottom. The left column shows the measurement vector, y , in black with the retrieved forward model, F , in blue. The center column shows the ratio of the y over F known as α and is the convergence factor between the ALI measurement and the forward model. The final column is ALI aerosol extinction in blue with the associated error represented by the light blue shading. The green is 750 nm aerosol extinction measured by OSIRIS, and red is 750 nm extinction as measured by SALOMON.

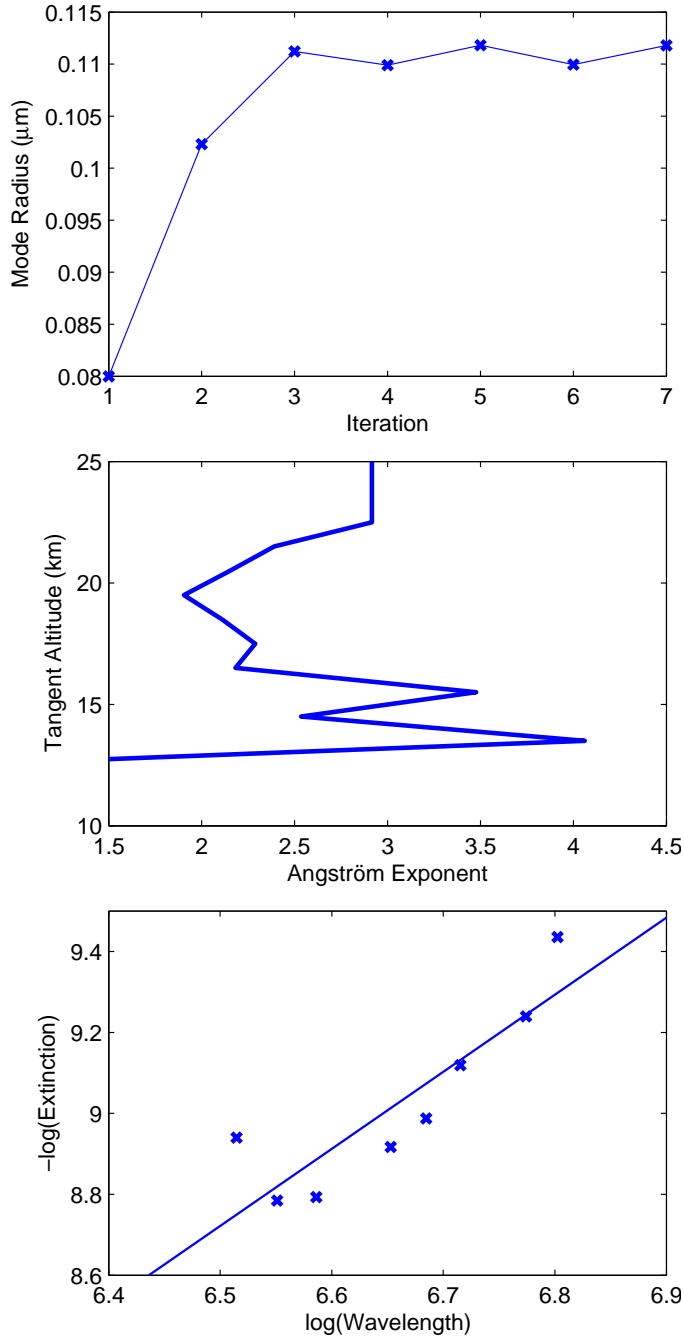


Figure 11: The top panel shows the convergence of the mode radius throughout the iterations in the retrieval. The second panel is the final Angström exponents determined for images 204-217 during the Timmins 2014 campaign. And the last panel demonstrate a least squares fit to determine the Angström exponent at 20.5 km shell altitude.