

Indian Institute of Technology, Kanpur Department of Earth Sciences

ES0213: Fundamentals of Earth Sciences

Lecture 04. Solar System: origin and characteristics

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Aims of this lecture



- Origin and history of Solar System
- Chemistry and structure of young Earth

Origin of Solar System and Planets

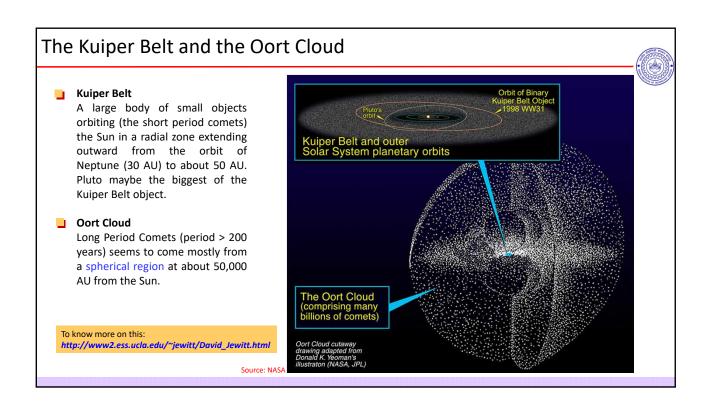


■ How was the Solar System Formed?

A viable theory for the formation of the solar system must be

- based on physical principles (conservation of energy, momentum, the law of gravity, the law of motions, etc.),
- able to explain all (at least most) the observable facts with reasonable accuracy, and
- able to explain other planetary systems.
- How do we go about finding the answers?
 - Observe: looking for clues
 - **Guess:** come up with some explanations
 - Test it: see if our guess explains everything (or most of it)
 - Try again: if it doesn't quite work, go back to step 2.

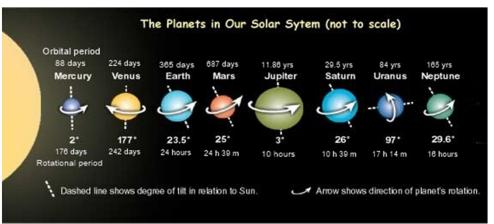
Solar System from space Sun, a star, at the center... Inner Planets (Mercury, Venus, Earth, Mars) ~ 1 AU; They are all rocky planets... Asteroid Belt, ~ 3 AU Outer Planets (Jupiter, Saturn, Neptune, Uranus), ~ 5-40 AU; They are all gaseous planets... Pluto: Odd one, a comet... not a planet Kuiper Belt ~ 30 to 50 AU Oort Cloud ~ 50,000 AU; Where comets come from... Source: NASA

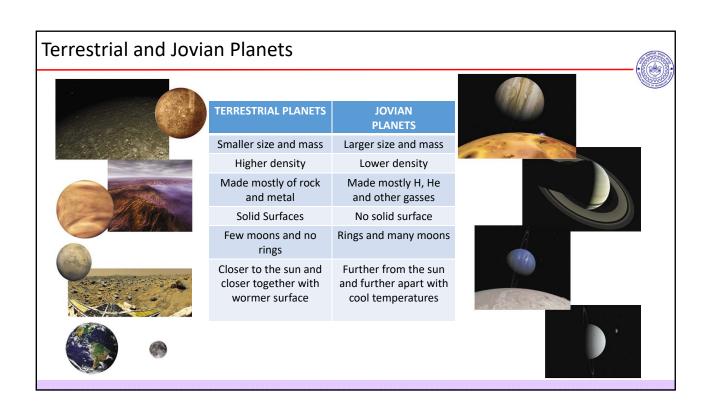


Orbital path and spin



- All the planets orbit the Sun in the same direction
- The rotation axes of most of the planets and the Sun are roughly aligned with the rotation axes of their orbits.
- Orientation of Venus, Uranus, and Pluto's spin axes are not similar to that of the Sun and other planets.





| Planet | Relative Size | Dist. Frm. Sun (AU) | Avg. Equt. rads (Km) | Mass (Earth=1) | Avg. Density (g/cm³) | Orbital Period | Rotation Period | Axis Tilt | Avg. Surf. T (K) | Composition | Known Moons (2004) | Rings |
|---------|---------------|------------------------|-------------------------|-------------------|----------------------------|-------------------|--------------------|-----------|---------------------|---------------------------|--------------------------|-------|
| Mercury | | 0.387 | 2.44 | 0.055 | 5.43 | 87.9 d. | 58.6 d. | 0.0° | 700 K (D) 100K (N) | Rocks, Metals | 0 | No |
| Venus | • | 0.723 | 6.051 | 0.82 | 5.24 | 225 d. | 243 d. | 177.3° | 740 K | Rocks, Metals | 0 | No |
| Earth | • | 1 | 6.378 | 1 | 5.52 | 1.00 yr. | 29.93 h. | 23.5° | 290 K | Rocks, Metals | 1 | No |
| Mars | | 1.52 | 3.397 | 0.11 | 3.93 | 1.88 yr. | 24.6 h. | 25.5° | 240 K | Rocks, Metals | 2 | No |
| Jupiter | | 5.2 | 71.492 | 318 | 1.33 | 11.9 yr. | 9.93 h. | 3.1° | 125 K | H, He, Hydrogen compounds | 79* | Yes |
| Saturn | | 9.54 | 60.268 | 95.2 | 0.7 | 29.4 yr. | 10.6 h. | 26.7° | 95 K | H, He, Hydrogen compounds | 31* | Yes |
| Uranus | • | 19.2 | 25.559 | 14.5 | 1.32 | 83.8 yr. | 17.2 h. | 97.9° | 60 K | H, He, Hydrogen compounds | 27 | Yes |
| Neptune | • | 30.1 | 24.764 | 17.1 | 1.64 | 165 yr. | 16.1 h. | 29.6° | 60 K | H, He, Hydrogen compounds | 14 | Yes |

Origin of Solar System and Planets



Historically, two hypothesis were put forward to explain the formation of the solar system....

Gravitational Collapse of Planetary Nebula (Latin for "cloud")

Solar system formed form gravitational collapse of an interstellar cloud or gas.

Close Encounter (of the Sun with another star)

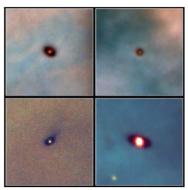
Planets are formed from debris pulled out of the Sun during a close encounter with another star

But, it cannot account for the angular momentum distribution in the solar system, Probability for such encounter is small in our neighborhood...

Origin of Solar System and Planets



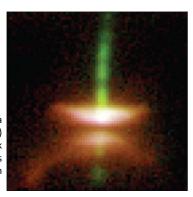
- Planet formation is a common outcome of star formation; but solar system is unique.
- The solar system may be unique, but to know the origin of Earth and other planets of the solar system observations of other planetary systems are necessary.
- Most developing stars, as viewed through telescopes, are approximately 99 percent gas and 1 percent dust; still they appear opaque at visible wavelengths.

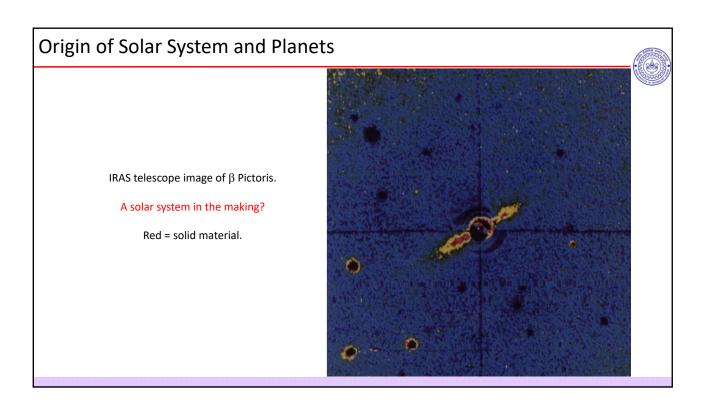


Hubble images of 4 protoplanetary disks around young stars in the Orion nebula, located 1,500 light-years from the Sun. The red glow in the center of each disk is a newly formed star approximately 1 million years old.

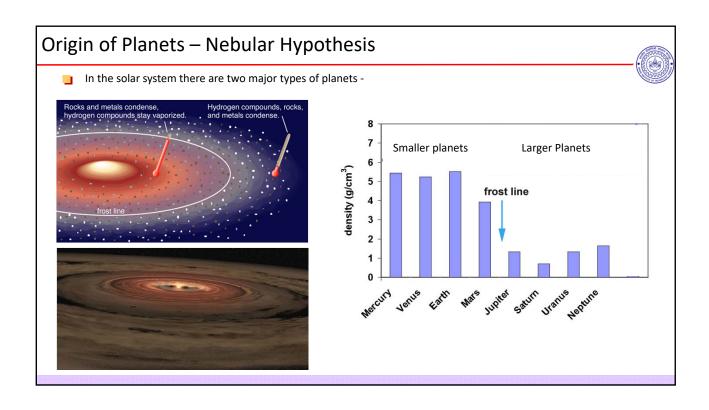
Image of Herbig-Haro 30, a prototype of a young (~1 my old) star surrounded by a thin, dark disk and emitting powerful bipolar jets of gas.. The gas jets, shown in green, are driven by accretion.

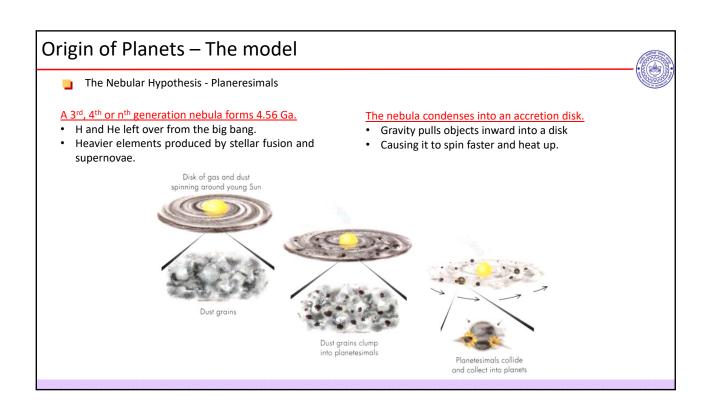
http://hubblesite.org/gallery/album/nebula_collection/





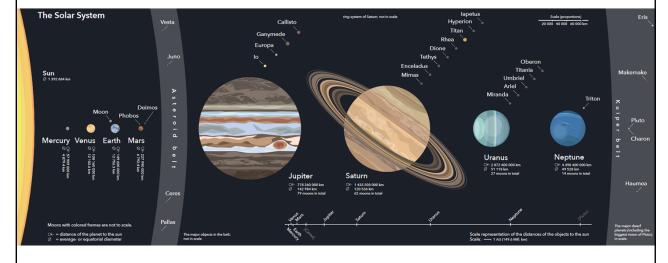
Gravitational Collapse — a scenario The solar nebula was initially somewhat spherical and a few light years in diameter very cold rotating slightly It was given a "push" by some event perhaps the shock wave from a nearby supernova As the nebula shrank, gravity increased, causing collapse As the nebula "falls" inward, gravitational potential energy is converted to heat. Conservation of Energy As the nebula's radius decreases, it rotates faster Conservation of Angular Momentum Finally it forms a disk like shape Orderly motion





The solar system





Origin of Planets – the solar system



- The primary difference between the inner and outer planets (rock versus gas and ice) is thought to reflect the temperature gradient in the solar nebula.
- Near the Sun, mainly silicates and metal would have condensed from the gas (so-called refractory materials), whereas beyond the asteroid belt, temperatures were low enough for ices (i.e., water, methane, ammonia) containing more volatile elements to have condensed, as well as solid silicates.
- Was the process that as the nebula cooled and solids formed unidirectional? NO. The solids typically were re-melted, re-evaporated, and re-condensed repeatedly as materials were circulated through different temperature regimes and variously affected by nebular shock waves and collisions between solid objects.

Origin of the larger planets – The models



The formation of giant planets starts with condensation and coalescence of rocky and icy material to form objects several times as massive as Earth. These solid bodies then attract and accumulate gas from the circumstellar disk (Pollack et al., 1996).

- Jupiter and Saturn, the two largest outermost planets fit well with this model as they have H and He, almost equal to solar proportions; but they have heavier elements (residing in the core) in greater than solar proportions.
- Uranus and Neptune have much less abundance of H and He compared Jupiter and Saturn; also have atmospheres consisting of solar ice.

An alternative to the standard model is that the rock and ice balls are not needed to induce the formation of gas-giant planets; they can form directly from the gas and dust in the disk, which can collapse under its own gravity like mini versions of the Sun (Boss, 2002).

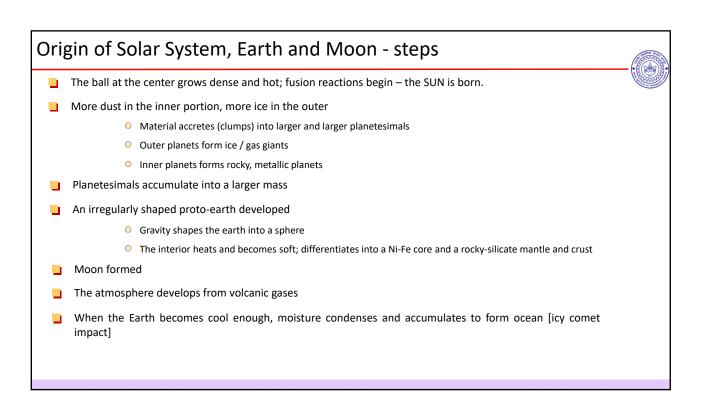
Origin of the inner planets – The models

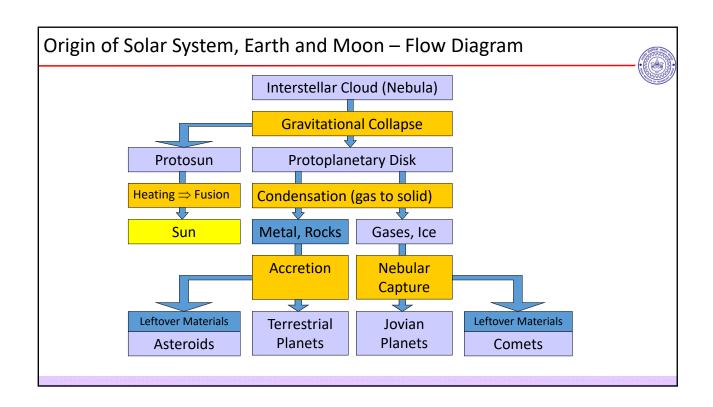


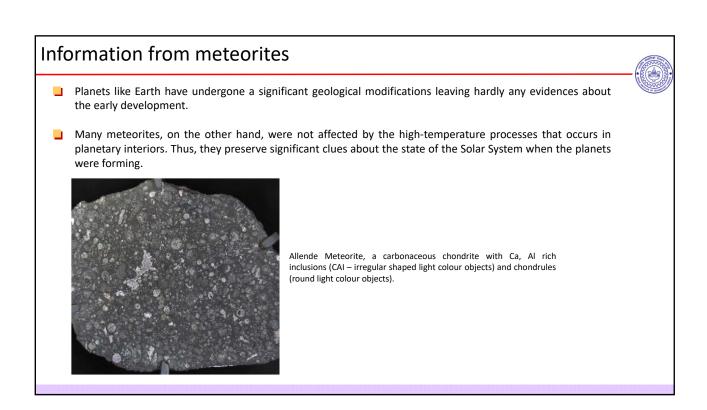
The model for the formation of inner planets is based on the information from meteorites and proposes two distinct stages of planer formation – early- and late-stage.

- The dust grains aggregated slowly at first, and growth accelerated along with object size as small objects were embedded into larger ones. The larger bodies are less affected by the presence of gas and their subsequent evolution is governed by mutual gravitational attractions. Gravitational interactions gave the largest planetesimals nearly circular and coplanar orbits—the most favorable conditions for sweeping up smaller objects.
- ☐ The later stages of planet formation took much longer, involved progressively fewer objects. The main phase of terrestrial planet formation probably took a few tens of millions of years. The final stages were marked by the occasional collision and merger of planetary embryos, which continued until the orbits of the resulting planets separated sufficiently to be protected from additional major collisions.

Moon rocks provide one of the most persuasive pieces of evidence that Earth and the Moon have a common origin. 1 During middle to late stages of Earth's accretion, short 5 billion years ago, a late impact quickly propelled a shower of distriction both the impact of the Earth into open and Earth into open



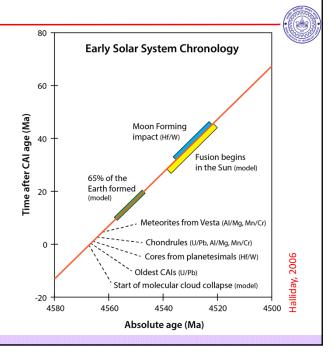




Information from meteorites

The most primitive chondritic meteorites contain inclusions made up of minerals that condense at high temperature from a gas of solar nebula composition. These objects, called calciumaluminum-rich inclusions (CAI), have been precisely dated using the decay of uranium to lead, where time is measured by the accumulation of the lead decay products formed at 4,567 (±1) Ma. This age is now generally accepted as "time zero" for the Solar System.

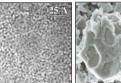
> Recent geochronological data and models for the sequence and timing of events in the early Solar System.



The primitive particles – Presolar Grains

- "Presolar grains" are stardusts manufactured by individual stars before the birth of our Solar System that are preserved in primitive meteorites.
- Each of these grains contains chemical elements that were made or reprocessed by an individual star and keeps information about the initial composition and chemical evolution.

| Туре | Size | Concentration in Meteorites | Sources |
|--|--------------------|-----------------------------|--|
| Diamond (C) | 1-5 nanometers | 1000 parts per million | Supernovae |
| Silicon carbide (SiC) | 0.1-10 micrometers | 10 parts per million | Carbon-rich giant stars, or supernovae |
| Graphite (C) | 1-10 micrometers | 2 parts per million | Supernovae and carbon-rich giant stars |
| Aluminum oxide (Al ₂ O ₃) | 1-5 micrometers | 0.1 parts per million | oxygen-rich giant stars |
| Spinel (MgAl ₂ O ₄) | 1 micrometer | 2 parts per billion | oxygen-rich giant stars |
| Silicon nitride (Si ₃ N ₄) | 1 micrometer | 2 parts per billion | Supernovae |







Nano-diamond

Graphite

Composition of primitive Earth



- At the pre-accretion stage, a very strong differentiation affected by the solar wind, light pressure, short-lived temperature rise, and magnetic separation GAS COMPONENTS MOSTLY REMOVED.
- The young Earth was a relatively cold celestial body. Nowhere within it had temperature exceeded the melting point of the Earth matter.
- The primordial Earth had a relatively uniform composition. Therefore, there was no Earth core at that time, and there was no chemical stratification into the mantle and crust.
- The source matter was hydro-silicates, carbonates, and sulfur as well as bases and some other easily fusible elements.

Composition of primitive Earth

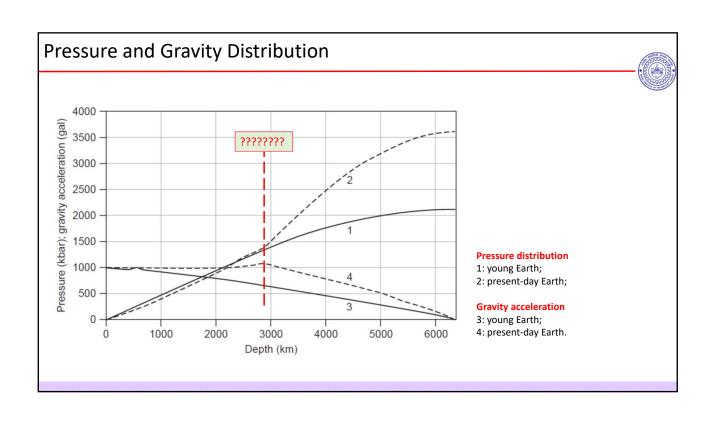


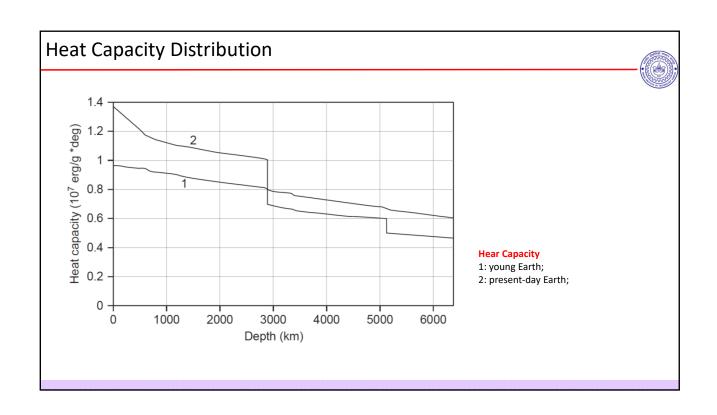
| | Avg. Comp. of coaly chondrites ^d | Avg. Comp. of chondrites ^c | Comp. of Primordial matter (calculated) | Comp. of Core | Comp. of Mantle ^b | Comp. of conti. Crust ^a | Oxides |
|----------------------------------|---|---------------------------------------|---|------------------|---------------------------------|------------------------------------|--------------------------------|
| | 33.000 | 38.040 | 30.710 | - | 45.400 | 59.3 | SiO ₂ |
| | 0.110 | 0.110 | 0.410 | - | 0.600 | 0.7 | TiO ₂ |
| | 2.530 | 2.500 | 2.540 | = | 3.700 | 15.0 | Al₂O₃ |
| | = | = | = | - | 1.970 | 2.4 | Fe₂O₃ |
| | 22.000 | 12.450 | 22.760 | 49.340 | 6.550 | 5.6 | FeO |
| | 0.240 | 0.250 | 0.090 | - | 0.130 | 0.1 | MnO |
| ^a Ronov and Yaroshev | 23.000 | 23.840 | 25.810 | = | 38.400 | 1.9 | MgO |
| (1978). | 2.320 | 1.950 | 1.570 | = | 2.300 | 7.2 | CaO |
| (1978). | 0.720 | 0.950 | 0.300 | - | 0.430 | 2.5 | Na₂O |
| | = | 0.170 | 0.016 | = | 0.012 | 2.1 | K ₂ O |
| b Dmitriyev (1973) an | 0.490 | 0.360 | 0.280 | - | 0.410 | = | Cr ₂ O ₃ |
| Ringwood (1966). | 0.380 | = | = | - | = | 0.2 | P_2O_5 |
| fillers and Cont. (405) | - | = | 0.070 | - | 0.100 | = | NiO |
| ^c Urey and Craig (195 | 13.600 | 5.760 | 2.170 | 6.690 | - | - | FeS |
| | - | 11.760 | 13.100 | 43.410 | - | = | Fe |
| d "Outlines of compar | | 1.340 | 0.180 | 0.560 | = | = | NiO |
| planetology." Nauka (| 98.390 | 99.480 | 100.000 | 100.000 | 100.000 | 100.0 | Total |

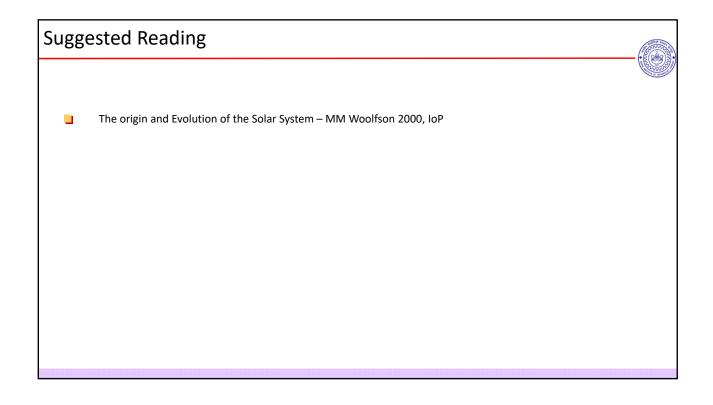
Structure of primitive Earth



| Depth (km) | Density (g/cm³) | Temperatu re (K) | Pressure (kbar) | g (cm/s²) | | epth km) | Density (g/cm³) | Temperature (K) | Pressure (kbar) | |
|---------------|--------------------|---------------------|--------------------|--------------|---|-------------|--------------------|--------------------|--------------------|--|
| 0 | 3.92 | 260 | 0 | 985 | 2 | 200 | 6.15 | 1379 | 1042 | |
| 200 | 4.15 | 1147 | 82 | 980 | 2 | 400 | 6.25 | 1378 | 1133 | |
| 400 | 4.38 | 1385 | 168 | 973 | 2 | 600 | 6.35 | 1377 | 1223 | |
| 400 | 4.5 | 1385 | 168 | 973 | 2 | 800 | 6.44 | 1376 | 1309 | |
| 600 | 4.76 | 1457 | 261 | 986 | 3 | 000 | 6.52 | 1375 | 1393 | |
| 670 | 4.85 | 1294 | 285 | 955 | 3 | 400 | 6.66 | 1373 | 1548 | |
| 670 | 5.02 | 1294 | 285 | 955 | 3 | 800 | 6.78 | 1371 | 1688 | |
| 800 | 5.16 | 1433 | 358 | 941 | 4 | 200 | 6.9 | 1369 | 1810 | |
| 1000 | 5.36 | 1411 | 456 | 921 | 2 | 600 | 6.99 | 1367 | 1912 | |
| 1200 | 5.53 | 1400 | 556 | 898 | 5 | 000 | 7.07 | 1365 | 1995 | |
| 1400 | 5.68 | 1393 | 656 | 874 | 5 | 400 | 7.11 | 1363 | 2057 | |
| 1600 | 5.81 | 1387 | 754 | 848 | 5 | 800 | 7.15 | 1361 | 2097 | |
| 1800 | 5.93 | 1384 | 852 | 821 | 6 | 200 | 7.18 | 1359 | 2116 | |
| 2000 | 6.04 | 1381 | 1381 | 793 | 6 | 360 | 7.18 | 1358 | 2116 | |







Next Lecture



We shall talk about the Earth as a System and learn some of its basic principles