# 2.4 RayDec

The name RayDec stands for "Rayleigh wave ellipticity estimation using the random decrement technique". Figure 2.3 gives an overview of the principles of the method. The technique was published in *Geophysical Review Letters*, the following article ("Single Station Determination of Rayleigh Wave Ellipticity by Using the Random Decrement Technique (RayDec)" by Hobiger, Bard, Cornou, and Le Bihan, 2009a) describes the method in detail.

### 2.4.1 Abstract

Rayleigh wave ellipticity as a function of frequency is closely linked to underground structure, i.e. shear wave velocity profile and sediment thickness. The possibility to calculate these underground properties by inverting ellipticity curves has recently been shown. We propose a new technique enabling the Rayleigh wave ellipticity to be recovered over a wide frequency range by using ambient noise recordings. Based on the random decrement technique commonly used to characterize dynamic parameters of buildings, this method eliminates all wave types except Rayleigh waves. We apply the method to noise synthetics simulated for different underground structures and show its applicability to real seismic noise data.

### 2.4.2 Introduction

In order to properly evaluate local site amplification it is important to have a good knowledge of the local soil properties, *i.e.* shear-wave velocities and sediment thickness. These soil properties can be obtained by inverting dispersion curves. One way is to use the dispersion curves of surface waves, as derived from active or passive seismic array measurements. However, it usually provides reliable information at frequencies higher than the resonance frequency and the inverted velocity profiles do not carry information on the deeper structures. Additional information around the resonance frequency can be added by using H/V spectra. Recent studies have shown the possibility to evaluate the S-wave velocity profile by inversion of H/V spectra [Fäh et al., 2003, Arai and Tokimatsu, 2004] or joint inversion of H/V spectra and dispersion curves [Arai and Tokimatsu, 2005, Picozzi et al., 2005]. The classical H/V technique [Nakamura, 1989] calculates the ratio between the horizontal and the vertical spectrum of ambient vibrations recorded on a single three-component seismic sensor by  $H/V(f) = \frac{\sqrt{|E(f)|^2 + |N(f)|^2}}{\sqrt{2+|E(f)|^2}}$  where E(f), N(f)

single three-component seismic sensor by  $H/V(f) = \frac{\sqrt{|E(f)|^2 + |N(f)|^2}}{\sqrt{2}|Z(f)|}$  where E(f), N(f) and Z(f) are the spectra of the east-west, north-south and vertical displacements of the sensor. Both Rayleigh and Love waves contribute to the H/V spectrum [Bonnefoy-Claudet et al., 2008] and the presence of both surface wave types has to be accounted for in inversion algorithms. Therefore, the proportion between Rayleigh and Love waves in the noise wave field has to be inferred or assumed prior to the inversion [Scherbaum et al., 2003, Arai and Tokimatsu, 2004, Köhler et al., 2006]. Extraction of Rayleigh wave ellipticity and its subsequent inversion [Boore and Toksöz, 1969, Fäh et al., 2001, Malischewsky and Scherbaum, 2004] would avoid making such assumptions.

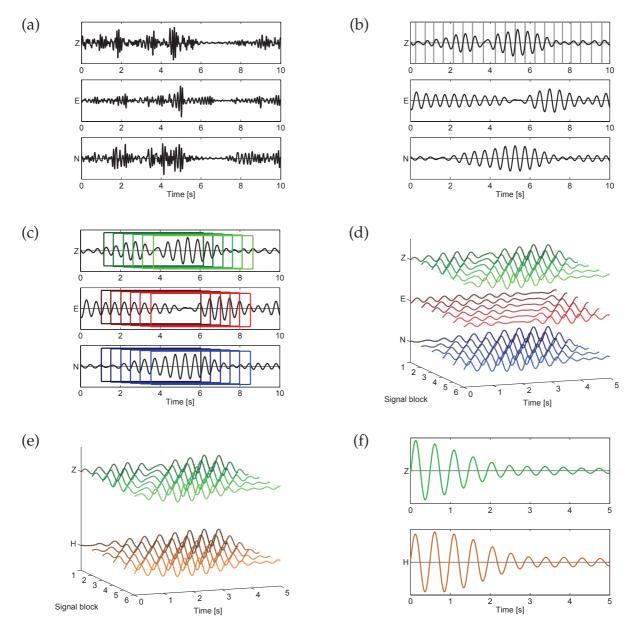


Figure 2.3: Graphic outline of the RayDec method: (a) A single three-component seismic sensor records a time series of data. (b) On the vertical component of the filtered signal, all zero crossings from negative to positive are searched for. (c) For every identified zero crossing, signals of length  $\Delta$  are extracted on all three components, delaying the horizontal by  $\frac{1}{4f}$ , which corresponds to the 90° phase shift to the vertical component. (d) These signal blocks are stacked. (e) Both horizontal signals are projected in the direction which maximizes the correlation to the vertical signal. (f) All these signals are summed, which yields a single vertical and horizontal signal. The ellipticity at frequency f is then estimated by comparing the respective energies of these summed signals.

We propose a new method (RayDec) to estimate the Rayleigh wave ellipticity spectrum by using the random decrement technique [Asmussen, 1997]. A usual application of the random decrement technique is the measurement of resonance frequencies and damping parameters of buildings [Dunand, 2005, Michel et al., 2008]. We propose to apply the random decrement technique to three-component seismometer records of seismic noise. By using the vertical component as a master trigger and stacking a large number of horizontal and vertical signals, Rayleigh waves will be emphasized with respect to Love and body waves. Therefore, the obtained ellipticity curve will be closer to the true curve than the H/V curve.

In this paper, we will present the method, apply it to synthetic seismic noise and finally show its applicability to real seismic noise records with one example.

## 2.4.3 Methodology

Assume that a three-component sensor record consists of three time-series of seismic noise (e(t) in east-west direction, n(t) in north-south direction and v(t) in vertical direction). Each of these time-series has a time length T and consists of N data points. In order to extract the ellipticity of Rayleigh waves we suggest the following method, based on the random decrement technique [Asmussen, 1997]. The principles of the method are to emphasize Rayleigh waves with respect to other wave types by summing a large number of specially tuned signal windows, calculate the energy on the vertical and horizontal summed signals and henceforward estimate Rayleigh wave ellipticity. The first step consists of filtering the signals by using a narrow-band Chebyshev filter of order 4 and bandwidth df centered on frequency f. We use the Chebyshev filter because of its sharpness. Filtering was performed on a linear frequency scale. The resulting signals are called  $e_f(t)$ ,  $n_f(t)$  and  $v_f(t)$ .

Then we search for all times  $\tau_i$  where the signal on the vertical component changes its sign from negative to positive ( $v_f(\tau_i) \leq 0$ ,  $v_f(\tau_i + dt) > 0$ ). Every time this condition is satisfied, we store signals of length  $\Delta$  on all three components, leading to buffered signals  $v_{f,b,i}(t)$ ,  $e_{f,b,i}(t)$  and  $n_{f,b,i}(t)$  defined in terms of the original signals as (for  $0 \leq t \leq \Delta$ ):

$$v_{f,b,i}(t) = v_f(\tau_i + t),$$

$$e_{f,b,i}(t) = e_f(\tau_i - \frac{1}{4f} + t),$$

$$n_{f,b,i}(t) = n_f(\tau_i - \frac{1}{4f} + t).$$
(2.42)

The shift of  $\frac{1}{4f}$  on the eastern and northern components accounts for the 90° phase shift between the vertical and horizontal components of Rayleigh waves. Then we build a horizontal signal  $h_{f,b,i}(t)$  by projecting the east-west and north-south components onto an axis forming an azimuth angle  $\theta_i$  with the northern direction by

$$h_{f,b,i}(t) = \sin(\vartheta_i) \cdot e_{f,b,i}(t) + \cos(\vartheta_i) \cdot n_{f,b,i}(t). \tag{2.43}$$

The angle  $\theta_i$  is chosen in such a way that it maximizes the correlation between the vertical signal  $v_{f,b,i}(t)$  and the horizontal signal  $h_{f,b,i}(t)$  given by:

$$C_{f,b,i}(\vartheta_i) = \int_0^\Delta v_{f,b,i}(t) \cdot h_{f,b,i}(t) dt$$

$$C_{f,b,i}(\vartheta_i) = \sin(\vartheta_i) \int_0^\Delta v_{f,b,i}(t) \cdot e_{f,b,i}(t) dt + \cos(\vartheta_i) \int_0^\Delta v_{f,b,i}(t) \cdot n_{f,b,i}(t) dt. \tag{2.44}$$

The correlation is maximized by

$$\vartheta_{i} = \tan^{-1} \left( \frac{\int_{0}^{\Delta} v_{f,b,i}(t) \cdot e_{f,b,i}(t) dt}{\int_{0}^{\Delta} v_{f,b,i}(t) \cdot n_{f,b,i}(t) dt} \right). \tag{2.45}$$

There are two possible  $\vartheta_i \in [0^\circ, 360^\circ]$  satisfying equation (2.45), with  $\vartheta_{i,2} = \vartheta_{i,1} + 180^\circ$ , one of them leading to a negative, the other to a positive correlation. The  $\vartheta_i$  leading to a positive correlation is chosen. The method does not distinguish between prograde and retrograde particle motion as prograde Rayleigh waves are treated as retrograde waves arriving from the opposite azimuth.

Although the chosen angle  $\vartheta_i$  maximizes the correlation between  $v_{f,b,i}(t)$  and  $h_{f,b,i}(t)$ , the correlation can be low. Thus, we make use of the correlation factor defined as

$$c_{f,b,i} = \frac{\int_0^\Delta v_{f,b,i}(t) \cdot h_{f,b,i}(t)dt}{\sqrt{\int_0^\Delta v_{f,b,i}^2(t)dt \cdot \int_0^\Delta h_{f,b,i}^2(t)dt}}.$$
 (2.46)

In the case of perfect correlation between  $v_{f,b,i}(t)$  and  $h_{f,b,i}(t)$  this factor equals 1. On the other side, it will vanish when the vertical and horizontal signals are not correlated at all. The signals  $v_{f,b,i}(t)$  and  $h_{f,b,i}(t)$  are weighted by  $c_{f,b,i}^2$  and summed for all times  $\tau_i$  to obtain two buffered signal windows of length  $\Delta$ :

$$v_{f,S}(t) = \sum_{i} c_{f,b,i}^{2} \cdot v_{f,b,i}(t)$$
 (2.47)

$$h_{f,S}(t) = \sum_{i} c_{f,b,i}^2 \cdot h_{f,b,i}(t).$$
 (2.48)

Finally, the ellipticity  $\epsilon$  is calculated as the square root of the ratio of the energies in the buffered signal windows  $h_{f,S}(t)$  and  $v_{f,S}(t)$ :

$$\epsilon = \sqrt{\frac{\int_0^{\Delta} (h_{f,S}(t))^2 dt}{\int_0^{\Delta} (v_{f,S}(t))^2 dt}}.$$
(2.49)

By repeated filtering of the signal over the entire frequency range and applying the presented method the ellipticity of Rayleigh waves as a function of frequency is determined.

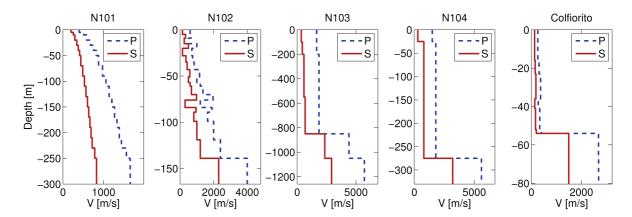


Figure 2.4: S- and P-wave velocity profiles for the models N101, N102, N103 and N104 [Cornou et al., 2009]. The original figure has been modified to improve the readability.

The free parameters of the method are the width of the frequency filter df and the length of the buffered signal  $\Delta$ .

It could be supposed that Love and body waves influence the results. In an assumed isotropic medium, Love waves are not present on the vertical component and their horizontal movements are thus uncorrelated with the vertical signal. Although incidental correlations of Love and Rayleigh waves arriving at the same time cannot be excluded, such effects will rapidly lose importance because of stacking and can thus be neglected. Arriving P- and S-waves can be present on the vertical and horizontal components, but as their movements are not phase delayed, they will also disappear due to averaging over a sufficiently large number of windows.

# 2.4.4 Application to synthetic noise

In order to test the method, we used a set of synthetic noise simulated for the models N101, N102, N103 and N104 of the ESG 2006 benchmark [Cornou et al., 2009]. The compressional (P) and shear wave (S) velocity profiles of these models are displayed in figure 2.4.

The noise synthetics were obtained in the following way: Noise sources were approximated by forces of random orientation and amplitude located 0.5 m below the surface [Moczo and Kristek, 2002]. The forces are randomly distributed in time and space. The time function of each source is a delta-like signal with a flat Fourier frequency-amplitude spectrum between 0.1 and  $20\,Hz$ . In this frequency range, the associated wave field has been computed by using the wavenumber-based technique for horizontally layered structures proposed by Hisada [1994, 1995]. All types of surface and body waves are included in the simulations.

### 2.4.4.1 Influence of the parameters of the method

As mentioned above, there are two free parameters that can be adjusted to improve the performance of the method: df, the width of the frequency filter, and  $\Delta$ , the length of the buffered signal. To investigate the optimal parameters, we used 30 minutes of noise simulated for the model N102. The noise sources are located in the far field as well as in the local field.

Figure 2.5 (a) shows the calculated ellipticity for different bandwidths df of the frequency filter compared to the theoretical ellipticity of the model. The used bandwidths are proportional to the central frequencies. Too weak filtering (df = 0.5f) as well as too strong filtering (df = 0.1f) do not give satisfying results. An intermediate setting of 0.2f reproduces the real curve the best. Therefore, we fixed the bandwidth to 0.2f in the following.

Figure 2.5 (b) shows the resulting ellipticity for different lengths  $\Delta$  of the buffered signal windows in comparison to the theoretical curve. We fix  $\Delta$  as a function of frequency, so that the signals consist of the same number of periods at all frequencies. As can be seen in the figure, at least ten cycles are necessary for a good coverage of the signal whereas further increasing of the window length does not improve the results. Therefore,  $\Delta = 10/f$  is used in the following.

## 2.4.4.2 Minimum required signal length and temporal stability of the results

The required signal length is illustrated in figure 2.5 (c). The same signals as in subsubsection 2.4.4.1 were used and the first 1800, 600, 300 and 120 seconds of signal were analyzed. The results of the analysis for signals of 1800, 600 and 300 seconds only slightly differ. However, the curve obtained for 120 seconds of signal is less smooth than the other curves. At low frequencies, the signal statistics deteriorate due to the smaller number of buffered windows. It can be deduced that - for sites with resonance frequencies around 1 Hz - a signal length of five minutes can be sufficient to measure an accurate ellipticity curve, as long as enough Rayleigh waves are present in the signal. For smaller underground resonance frequencies, the measurement has to be longer. Indeed, longer measurements improve the statistical features and the reliability of the results at all frequencies.

The temporal stability of the obtained ellipticity curves is shown in figure 2.5 (d). Five different synthetic signals of ten minutes each were simulated and the ellipticity curves were calculated for each window. The overall behavior of the curves is identical. Only the fourth ten-minute-window shows some instabilities between 0.1 and 0.2 Hz emphasizing the need for longer signal lengths at low frequencies. For the other frequencies, the five ellipticity curves are very close. This illustrates the temporal stability of the results.

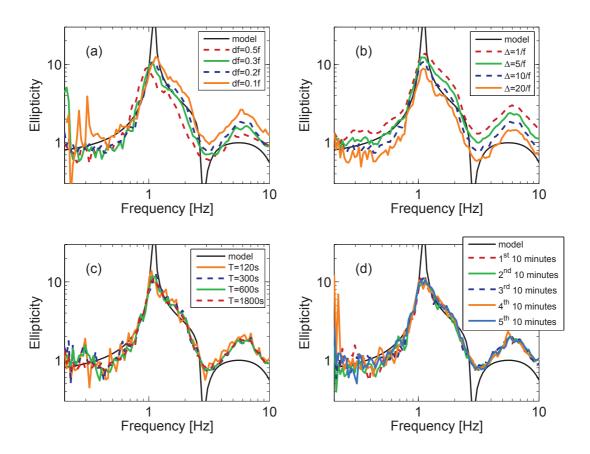


Figure 2.5: Sensitivity of the results with respect to the parameters of the method demonstrated on simulated seimic noise for the N102 model. The real ellipticity curve of the model is included for comparison. Curves (a) and (b) were calculated for 30 minutes of simulated seismic noise. (a) Effect of the filter bandwidth df on the resulting ellipticity curve ( $\Delta = 10/f$ ) for df = 0.5f, df = 0.3f, df = 0.2f and df = 0.1f. (b) Effect of the length  $\Delta$  of the buffered signal (df = 0.2f) for  $\Delta = 1/f$ ,  $\Delta = 5/f$ ,  $\Delta = 10/f$  and  $\Delta = 20/f$ . (c) Effect of the length of the signals (time length T, df = 0.2f,  $\Delta = 10/f$ ) for T = 1800s, T = 600s, T = 300s and T = 120s. (d) Analysis of five different signals of ten minutes each (df = 0.2f,  $\Delta = 10/f$ ). The original figure has been modified to improve the readability.

### 2.4.4.3 Results for the models N101, N103 and N104

RayDec has been applied to synthetic noise simulated for different models. The results of these simulations are shown in figure 2.6. For model N101 (figure 2.6 (a)), the curve is close to the real curve over the whole frequency range, although the peak ellipticity is overestimated. For this model layout, reflected body waves can deteriorate the results. Therefore, more signal would be needed to obtain a more robust ellipticity curve. For model N103 (figure 2.6 (b)), there is a good agreement between the simulated and the theoretical curves around the resonance frequency and at higher frequencies, with some minor deviations between 0.3 and 2 Hz.

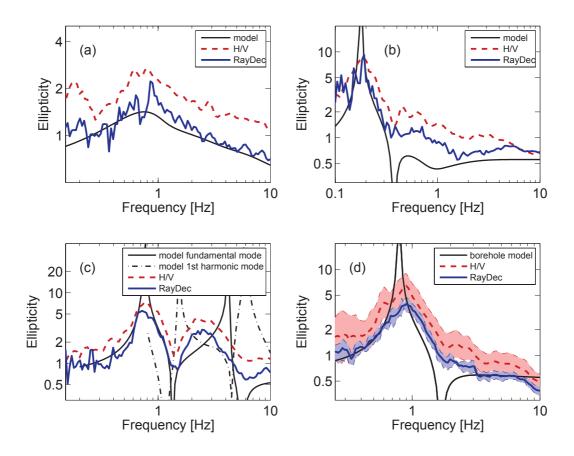


Figure 2.6: Ellipticity curves obtained for ten minutes of synthetic noise simulated for three different underground models (N101(a), N103 (b) and N104 (c)) and for one hour of real noise from Colfiorito basin (d). The used parameters are df = 0.2f and  $\Delta = 10/f$ . In the four graphs, the results are compared to the real ellipticity curves of the fundamental mode and the first harmonic mode (only in (c)) and to the results obtained by the H/V technique. Curve (d) also contains the error ranges of the H/V and RayDec curves. The original figure has been modified to improve the readability.

Below 0.15 Hz, the curve differs considerably from the theoretical curve. Model N104 (figure 2.6 (c)) exhibits some unfavorable behavior, as the fundamental mode and the first harmonic mode of Rayleigh wave motion overlap above 1.4 Hz. RayDec cannot distinguish between two modes. Therefore, its results are a mixture of the two modes in this frequency range. For lower frequencies, RayDec reproduces well the theoretical curve around the first peak of the fundamental mode. The H/V curve follows the same trend, but the absolute values of the RayDec curve better reproduce the right flank of the fundamental peak.

It can be stated that, in general, the results of RayDec are much closer to the theoretical curves than the results obtained with the H/V method. However, RayDec regularly fails in reproducing very small ellipticity values. At these points, the horizontal components of the Rayleigh wave vanish.

Therefore, it is problematic to find horizontal signals correlated to the vertical signal. The same reasoning is valid for the peak frequency ellipticity in the case of a theoretically vanishing vertical component. This explains why the peak ellipticity cannot be satisfactorily estimated. On the other hand, small ellipticities can also be hidden by the presence of higher Rayleigh wave modes. In cases where higher Rayleigh wave modes are close to the fundamental modes, RayDec only indicates an apparent ellipticity which cannot be simply related with the true ellipticity of each mode.

# 2.4.5 Application to real noise data

To test the RayDec method on real data, we applied it to one hour of seismic noise recorded in the center of Colfiorito basin in central Italy. Colfiorito basin is a well-known site which has been extensively investigated by different geophysical methods [Di Giulio et al., 2006]. The theoretical ellipticity curve is calculated on the basis of S-and P-wave velocity profiles obtained by a borehole measurement (see figure 2.4). The parameters used were df = 0.2f and  $\Delta = 10/f$ . The hour of signal was cut into six ten-minute signals which were analyzed separately. The resulting curves were then averaged. When comparing with the theoretical ellipticity curve, one should keep in mind that a borehole measurement is made at a selected point, whereas a seismic noise record is influenced by the whole surrounding underground structure. The RayDec curve is in better agreement with the theoretical curve than the H/V curve, especially on the slopes of the fundamental resonance frequency peak (figure 2.6 (d)). However, both the H/V and RayDec curve's peak frequencies are shifted with respect to the borehole model's peak frequency.

#### 2.4.6 Conclusion

RayDec is a fast and simple method to determine the ellipticity of Rayleigh waves. As it suppresses Love and body waves efficiently, its results are closer to theory than H/V results. The necessary signal length depends on the resonance frequency of the underground structure, but some minutes of signal can be sufficient to get satisfying results, as long as enough Rayleigh waves are present in the wave field, although longer signal lengths are preferable. However, as it is a one-sensor technique, RayDec gives two possible azimuths and cannot distinguish between prograde and retrograde particle motion. By coupling RayDec with array measurements the correct azimuth and the sense of particle motion could be retrieved. Besides that, it could be envisaged to estimate the proportion of Rayleigh waves in the wave field by comparing RayDec results including only Rayleigh waves with H/V measurements including all wave types. The future work consists of trying to improve the inversion of S-wave velocity profiles by using RayDec's ellipticity estimations. This is a non-trivial endeavor which will be addressed in a future paper.

## 2.5 Conclusion

One of the main fields of interest of this thesis is the measurement and use of the ellipticity of Rayleigh waves. The rectilinearity and planarity filters indicate which wave type is predominant at a given time and can be used to discriminate against undesired wave types and enhance the desired signal part. However, these methods do not allow the direct retrieval of ellipticity and will not be considered in the following chapters. Nevertheless, they could be used as a preprocessing step for real data applications or for studies on specific seismic phases (P, S, surface waves).

The new methods DELFI and RayDec, as well as H/V, will be tested on synthetic seismic signals in chapter 5. In chapter 6, the potentials and limitations of the inversion of ellipticity curves will be investigated. These inversions will subsequently be applied to real data measurements collected during the NERIES project for 14 European sites in chapter 7. The ellipticity curves used for these inversions will be measured by RayDec. A concise overview of DELFI and RayDec was presented at the European Signal Processing Conference (EUSIPCO) 2009 in Glasgow (UK). The associated paper ("Rayleigh wave ellipticity estimation from ambient seismic noise using single and multiple vector-sensor techniques" by Hobiger, Le Bihan, Cornou, and Bard, 2009b) which includes a comparison with the MUSIQUE technique (see section 4.6) can be found in appendix A.