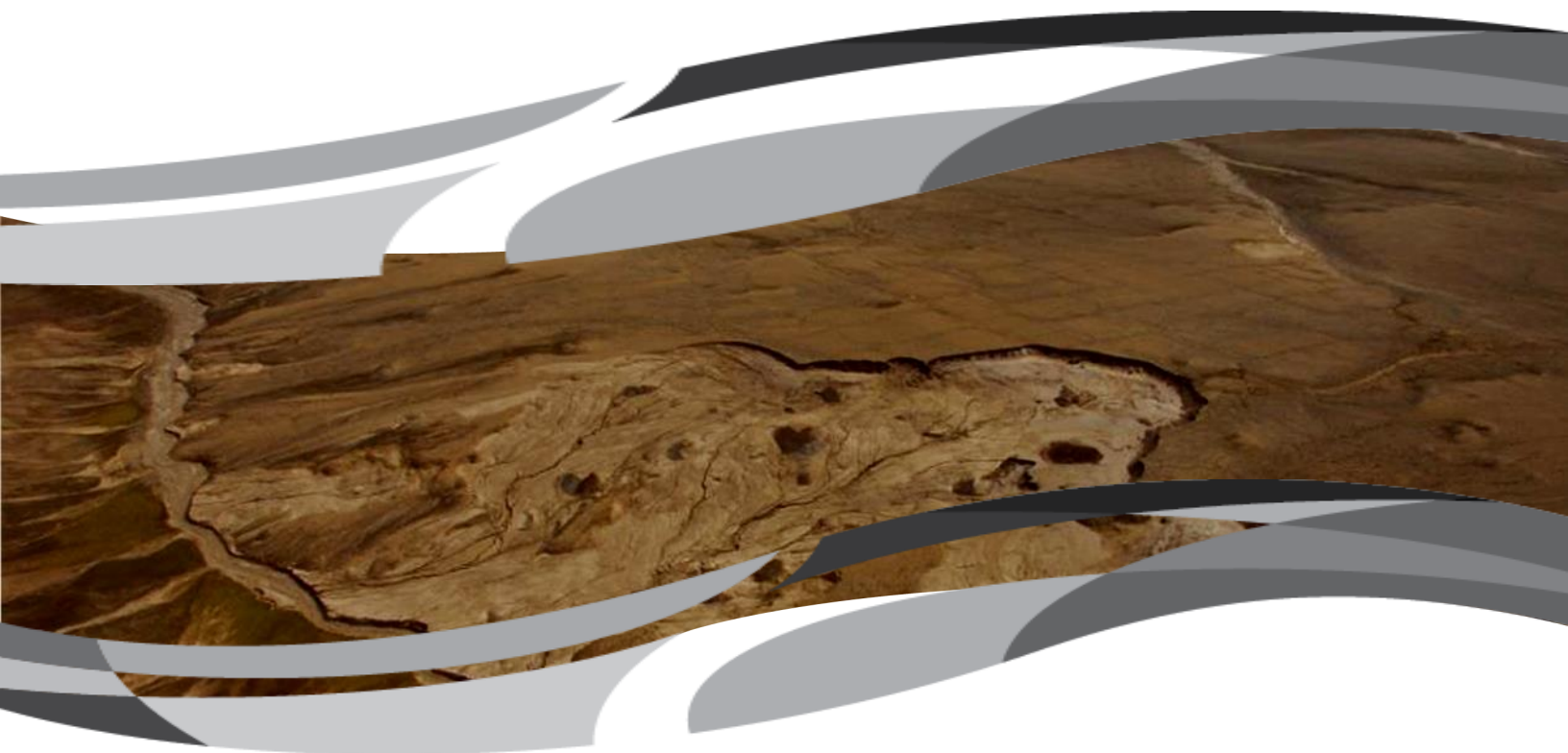




NWT Open Report 2015-021

**Inventory of active retrogressive thaw slumps on eastern
Banks Island, Northwest Territories**



R.A. Segal, T.C. Lantz, and S.V. Kokelj

**NORTHWEST TERRITORIES
GEOLOGICAL SURVEY**

Government of
Northwest Territories

Cover Image:

A retrogressive thaw slump on Banks Island, Northwest Territories. Credit: Ecosystem Classification Group (2012).

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Introduction

The purpose of this report is to describe the spatial data distribution of terrain affected by active retrogressive thaw slumping on eastern Banks Island, Northwest Territories (NWT). Thawing of ice-rich permafrost can lead to significant terrain modification, alter terrestrial and aquatic ecosystems, and pose a threat to infrastructure. This remote mapping project contributes data on the distribution and density of thaw slumps across a study region (approximately 60 km by 96 km) on eastern Banks Island.

Background

Retrogressive thaw slumps are a form of thermokarst that develop in sloping terrain underlain by ice-rich permafrost (Figures 1 and 2; Lewkowicz 1987; Lantuit and Pollard 2005; Lacelle et al. 2010). These long-lived geomorphic disturbances commonly occur along coastlines, lakeshores, and river valleys. Thaw slumps can pose a threat to northern infrastructure (Kokelj and Jorgenson 2013) and can impact freshwater, marine, and terrestrial ecosystems (Lantuit and Pollard 2008; Kokelj et al. 2009; Mesquita et al. 2010; Kokelj et al. 2013; Lantz et al. 2009).

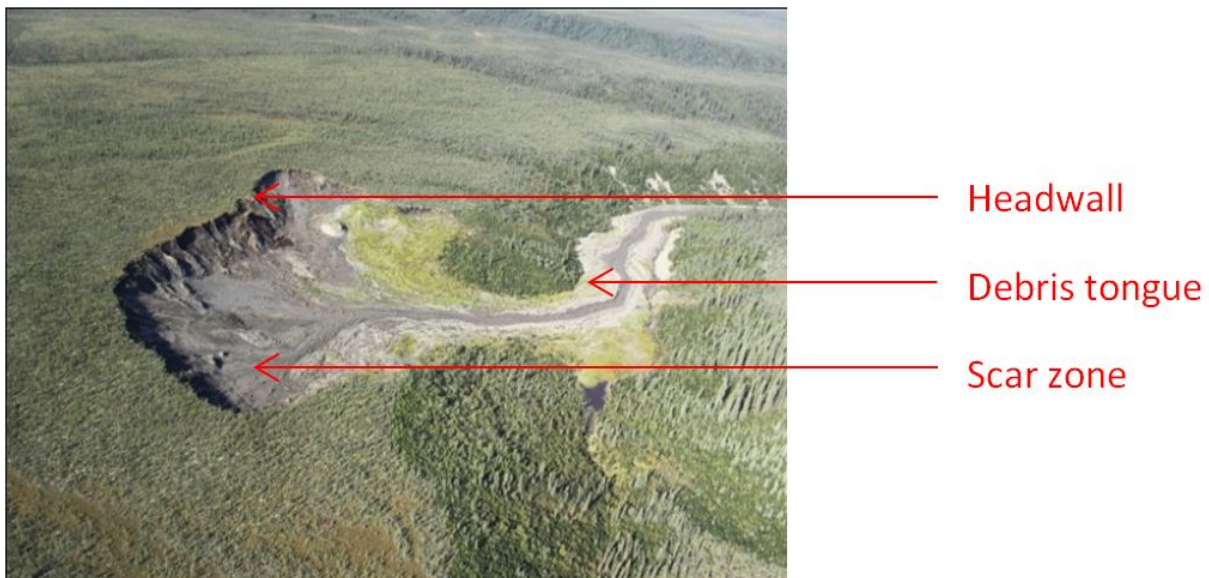


Figure 1. A retrogressive thaw slump in the Peel Plateau area (67.280°N, 135.162°W) in 2010. Arrows denote the features described in the text.



Figure 2. A retrogressive thaw slump on Banks Island. The white bar in the lower right-hand corner of the photo is ~45 m long. Image from the Ecosystem Classification Group (2012).

Active thaw slumps are characterised by an ice-rich headwall that defines the upslope boundary of the slump. Ablation of exposed ground-ice can cause retrogressive slumping and growth of the disturbance. The scar area of an active slump typically consists of saturated mineral soils and detritus, which contrasts significantly with the surrounding undisturbed forest or tundra. The headwall or upslope margins of a slump are typically crescent-shaped (Figure 1). The saturated materials in the scar zone can move downslope by a variety of processes, which can result in the development of a debris tongue (Burn and Lewkowicz 1990). Mass flows may cause the debris tongue to extend downslope and infill small valleys (Figure 1), or the materials may be transported into a lake or coastal zone. Slump stabilisation occurs when the headwall becomes covered and insulated by debris (Burn and Friele 1989; Burn and Lewkowicz 1990; Kokelj and Jorgenson 2013).

Thaw slumps are significant landscape features that impact infrastructure, and aquatic and terrestrial ecosystems (Kokelj et al. 2013; Mesquita et al. 2010; Lantuit and Pollard 2008; Kokelj et al. 2009). Sediment and solutes released by slumping can alter the chemistry of disturbed soils and surface waters. Mass movement of sediments and runoff can alter the chemistry and turbidity of adjacent rivers, lakes, and coastal environments (Kokelj et al. 2005; Lantuit and Pollard 2008; Kokelj et al. 2009; Mesquita et al. 2010; Vonk et al. 2012; Kokelj et al. 2013; Malone et al. 2013; Thienpont et al. 2013). Following slump stabilisation, vegetation communities and the underlying permafrost soils can take centuries to recover from the disturbance (Kokelj et al. 2009; Lantz et al. 2009).

Eastern Banks Island is a cold region located in the Canadian high Arctic (Fig. 3) that contains retrogressive thaw slumps. Mean annual air and ground temperatures are approximately -14 °C and -6 °C respectively (NASA GISS 2014; Global Terrestrial Network for Permafrost 2010). During the Quaternary, Banks Island is thought to have been completely covered by the Laurentide Ice Sheet, which left behind a sequence of morainal deposits. In the eastern Banks Island study region, the Jesse Moraine covers all but the northwestern portion of the study region, which is characterised as undifferentiated drift (Lakeman and England 2012). Eastern Banks Island has undergone significant increases in mean annual air temperature in the past 55 years (+2.3 °C; (NASA GISS 2014). The topography of this region is relatively subtle, ranging from 0 m above sea level (ASL) to about 250 m ASL (Government of Canada et al. 2000). The slumps identified in this ice-rich region were generally small and slow-growing.

This report provides spatial data on the size, location, and types of active retrogressive thaw slumps on eastern Banks Island, NWT. Appendix A presents the accompanying Shapefiles, Google Earth files and metadata.

Compilation Methods

All active retrogressive thaw slumps in the eastern Banks Island study region were digitized on-screen (Figure 3) by viewing georeferenced imagery in Google Earth, ESRI ArcMap (versions 10.0 and 10.1), and the NWT Spatial Data Warehouse Geospatial Portal (<http://www.geomatics.gov.nt.ca/sdw.aspx>, NWT Centre for Geomatics 2013; Latitude Geographics Group Ltd.). Slumps were digitised using 2004 Quickbird imagery (0.6 m resolution) where available, as well as SPOT 5 and SPOT 4 imagery (10 m resolution, from 2006, 2009, and 2010). Slumps were identified based on the presence of exposed sediments, poorly developed vegetation, and a well-defined headwall. Thaw slump outlines were consolidated into a single shapefile using ArcMap. Waterbodies were removed from the study area using CanVec hydro features (waterbody products; 1:50 000 resolution) (Government of Canada 2016) prior to the slump density calculations.

The results of this assessment are summarised in two figures that show: A) the distribution of retrogressive thaw slumps across the eastern Banks Island study region (Figure 3), and B) the distribution of thaw slump sizes (Figure 4).

Results and Summary

The study area with the waterbodies removed was 4701 km². There were 832 active retrogressive thaw slumps digitised in the eastern Banks Island study region (Figure 3). These slumps ranged in size from 299 m² to 224 121 m² and had an average area of 16 917 m² (1.69 ha). A histogram showing the size distribution of retrogressive thaw slump areas is shown in Figure 4. This plot shows that the distribution of active slumps is roughly log-normal. There was an average of 17.7 slumps / 100 km². However, slumps were distributed throughout deposits of the Jesse moraine complex, but absent from the northwestern part of the study region that is underlain by undifferentiated drift (Lakeman and England 2012).

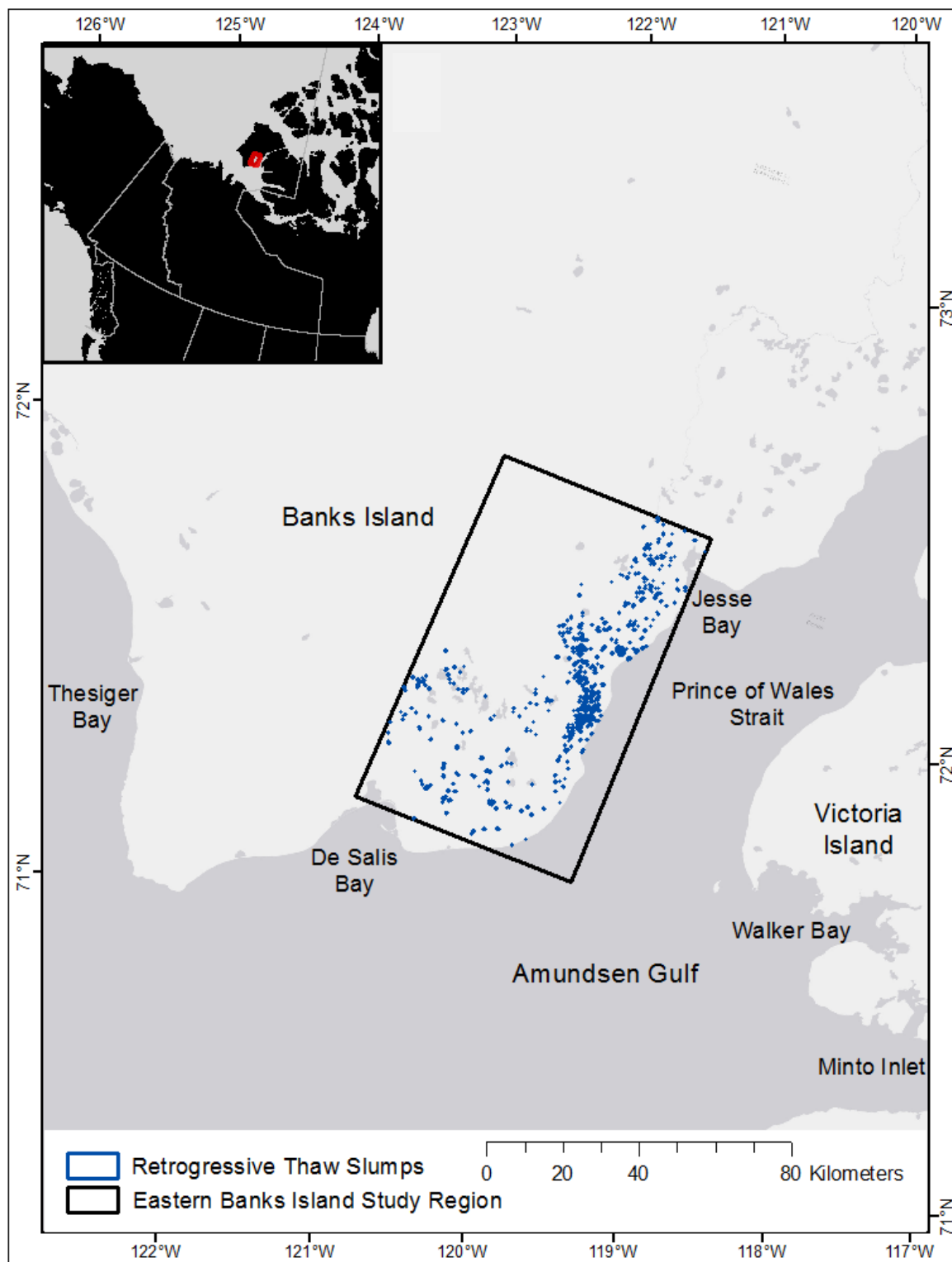


Figure 3. Map of the study region on eastern Banks Island (black outline) and active retrogressive thaw slumps (blue). The location map in the upper left corner shows the position of the study area in northwestern Canada.

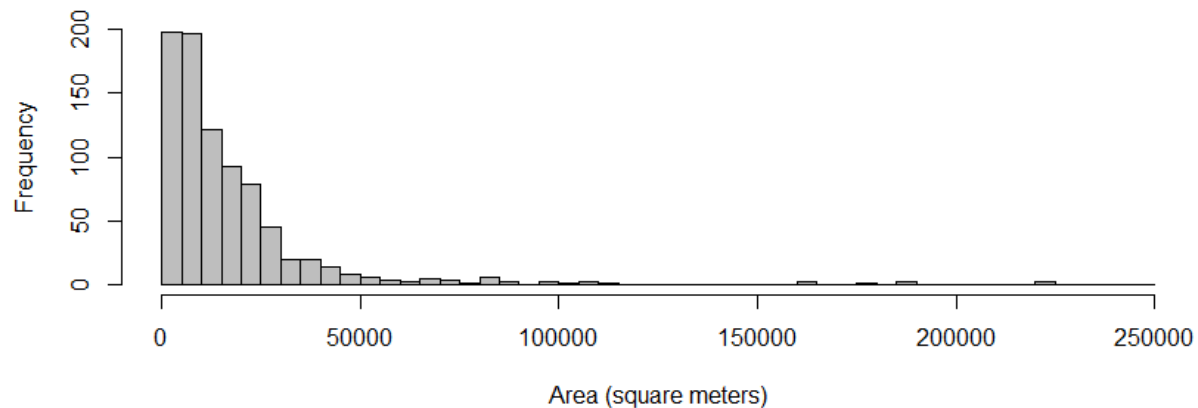


Figure 4. Histogram showing the size distribution of active retrogressive thaw slumps on eastern Banks Island.

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