

OCEAN WAVES

INTRODUCTION

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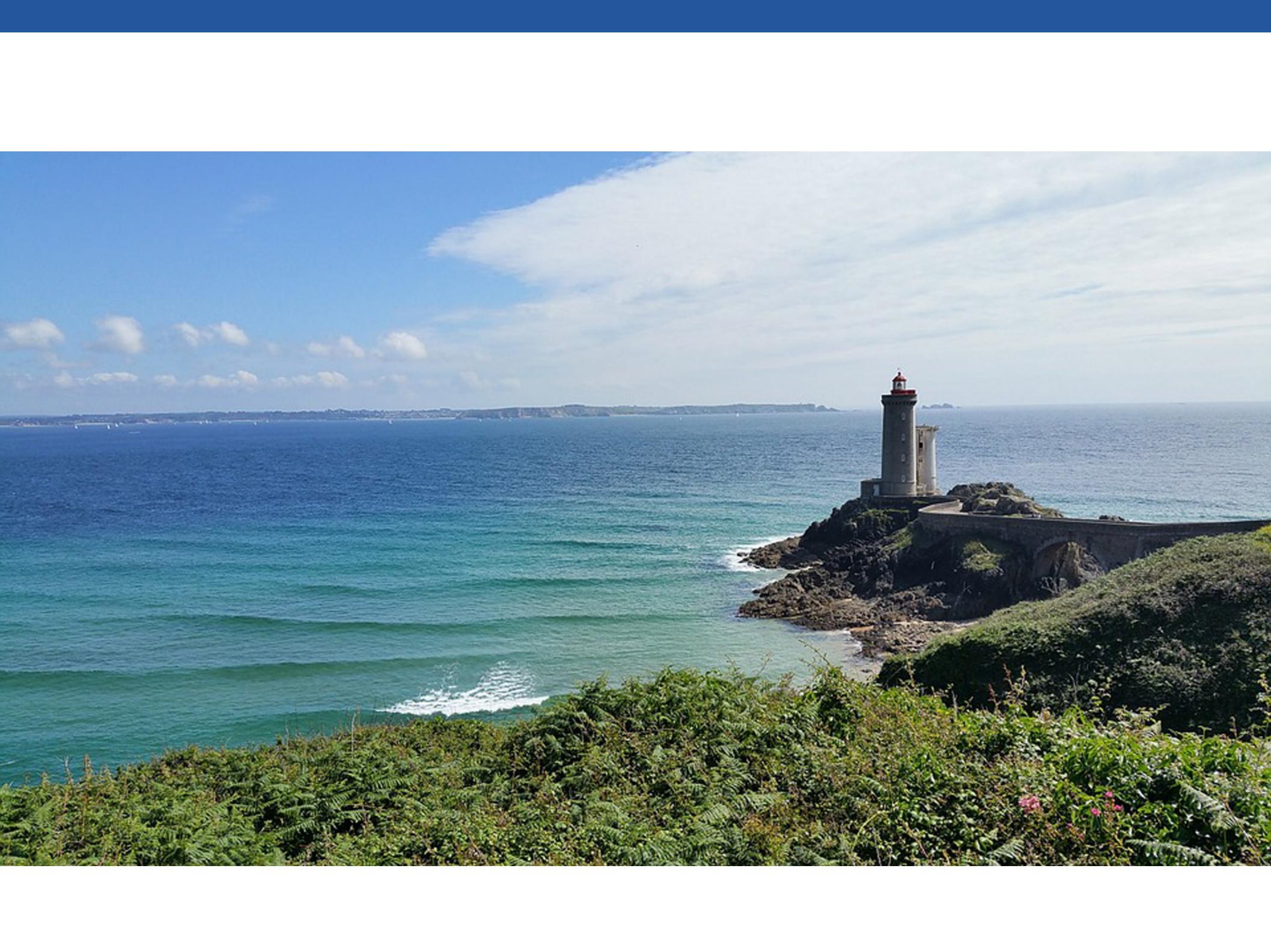
- **Lesson 1 :**
 - Introduction about Ocean waves
 - Surface waves
 - Internal Waves (Introduction)
 - Internal Waves in the 2-layer shallow-water model
- **Lesson 2 :**
 - Internal Waves with a continuous stratification
- **Lesson 3 :**
 - Generation of internal waves
- **Lesson 4 :**
 - Dissipation and interaction of internal waves
 - Activities: Numerical simulation of internal waves

Presentations and material will be available at :

jgula.fr/Ondes/

Useful references

- *The course is largely based on:*
 - Gerkema- Zimmerman (2008). *An introduction to internal waves* <https://www.jgula.fr/Ondes/gerkema.pdf>
- *Other useful references are:*
 - Leblond-Mysak (1977) : *Waves in the ocean*
 - Whitham (1974) : *Linear and nonlinear waves*
 - Gill (1982) : *Atmosphère-Ocean Dynamics*
 - Kundu-Cohen (1987). *Fluid Mechanics. Third edition*
 - Cushman-Roisin. *Introduction to geophysical fluid Dynamics*
- *Some important research articles and reviews are available here:* <https://www.jgula.fr/Ondes/>





Space and time scales ?

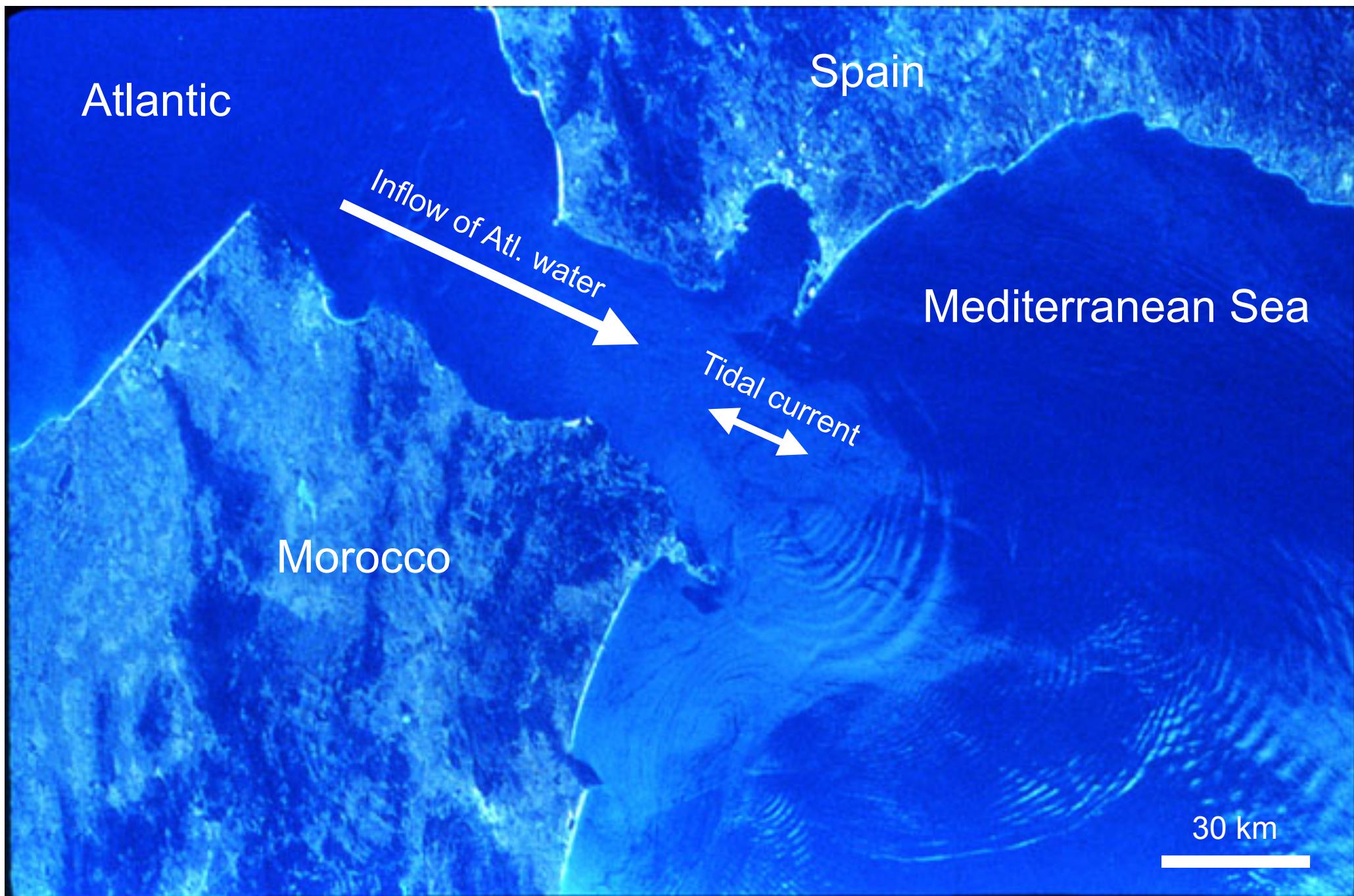


Surface capillary wave

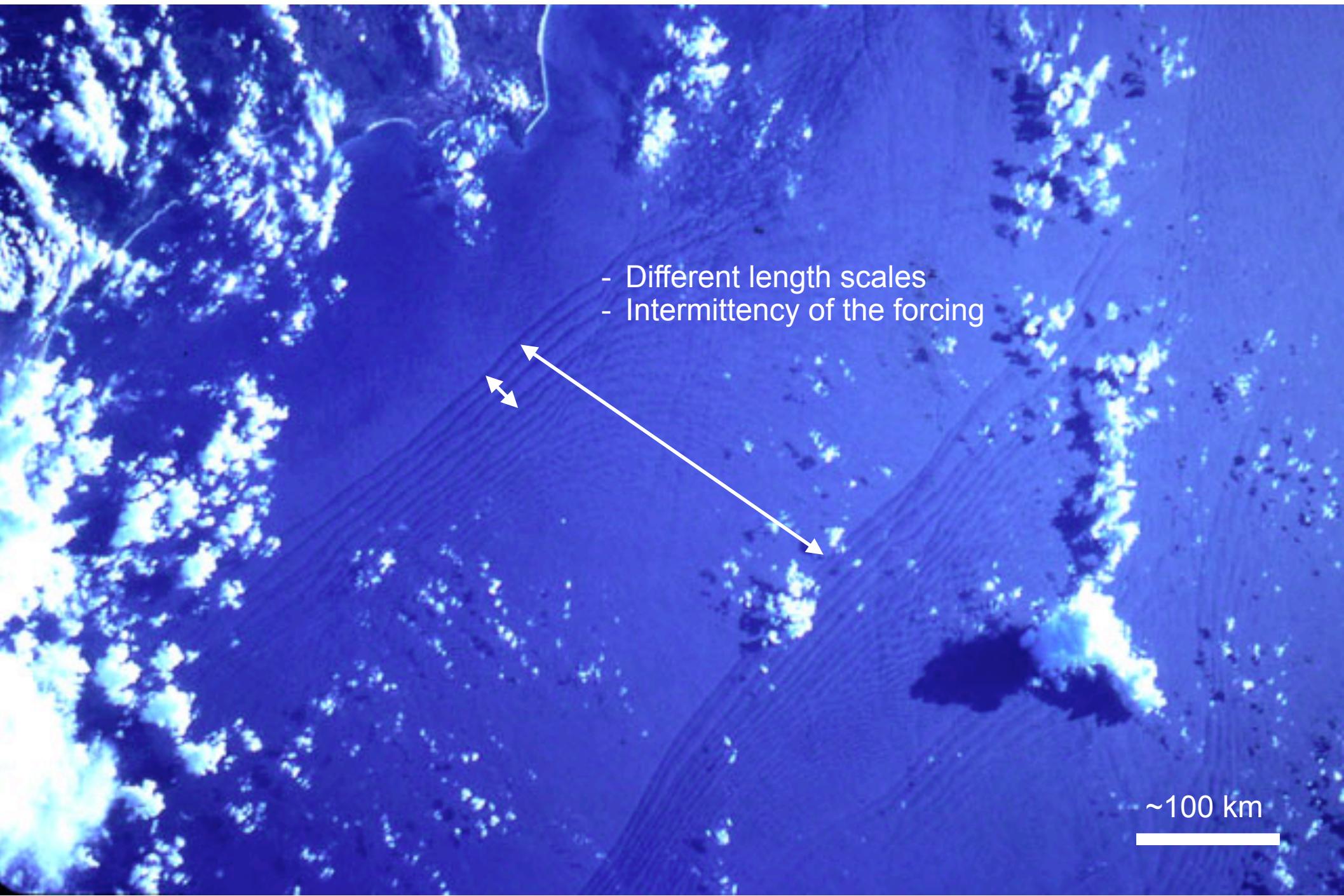
a few cm



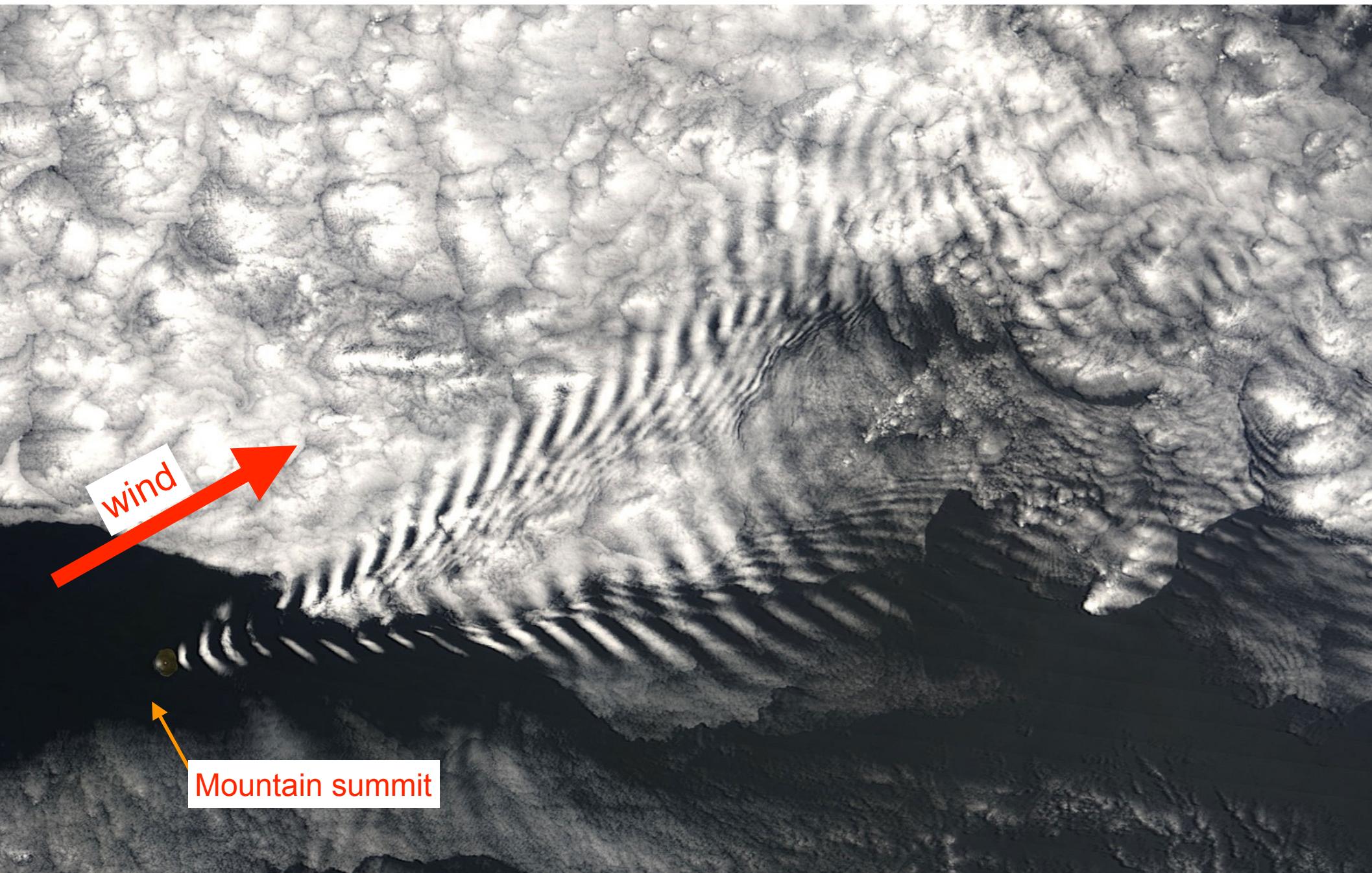
Soliton in the Strait of Gibraltar



Wave packets in the South China Sea



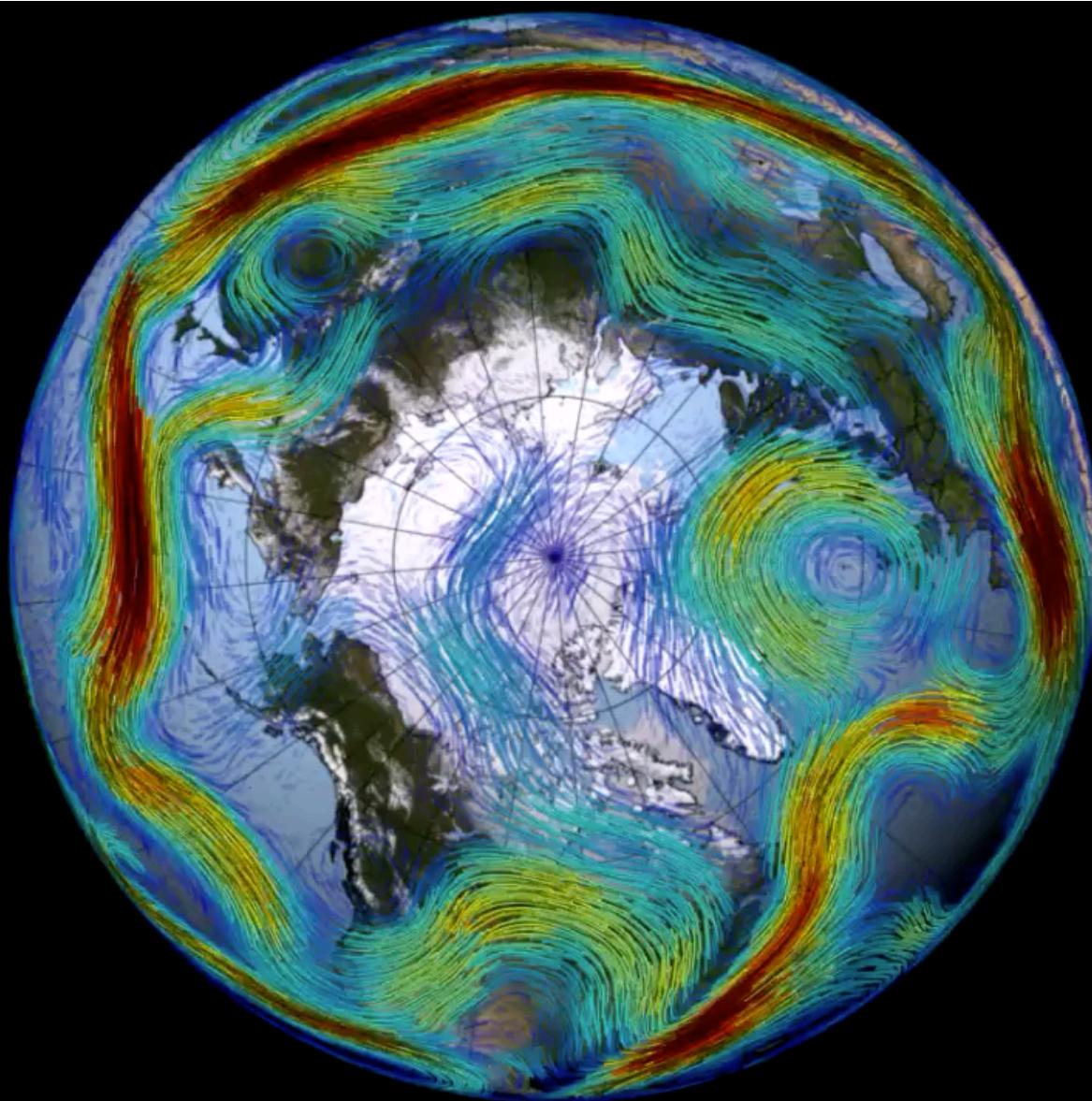
Standing waves generated leeward of a mountain summit



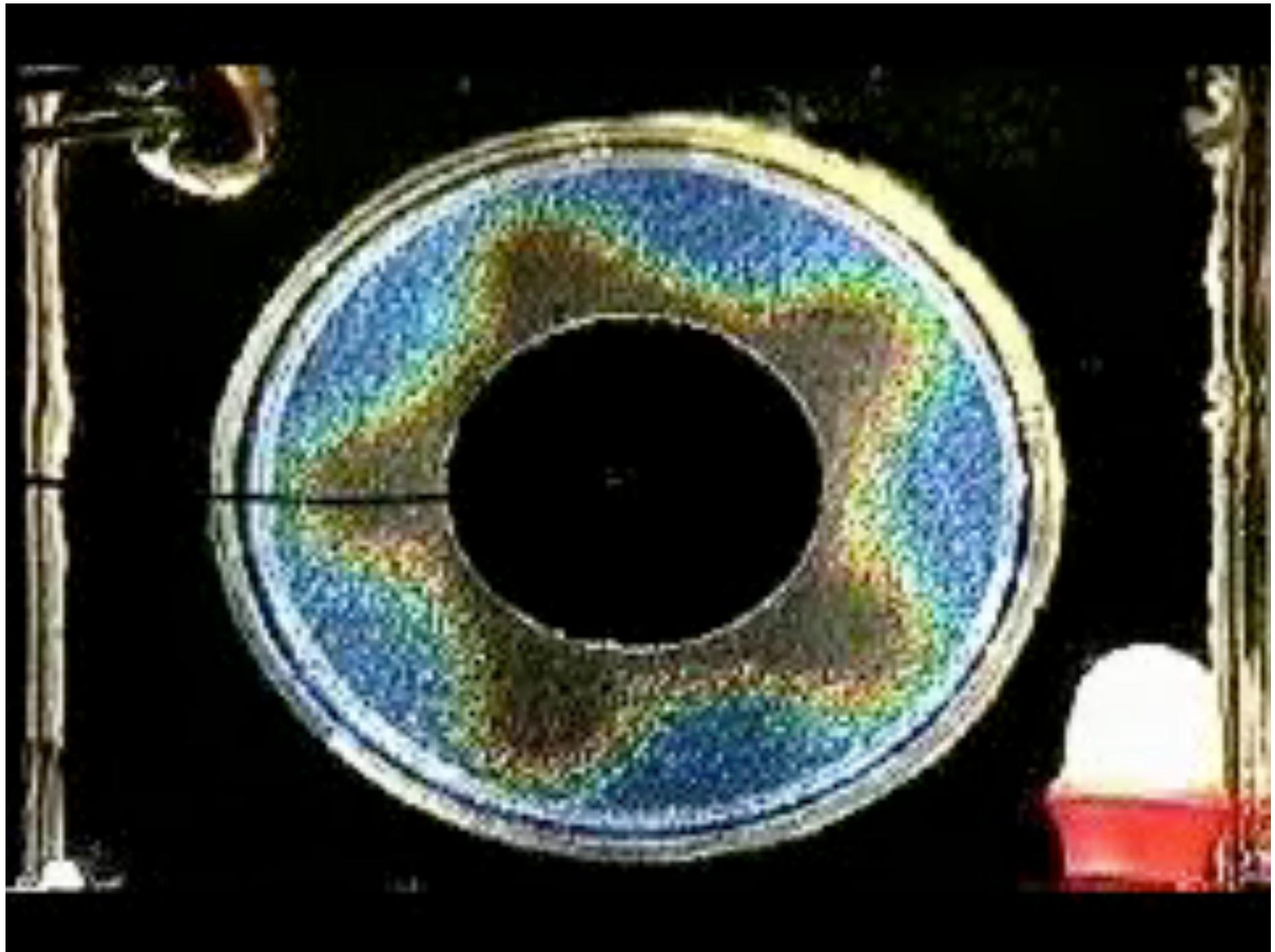


© marit kastberg

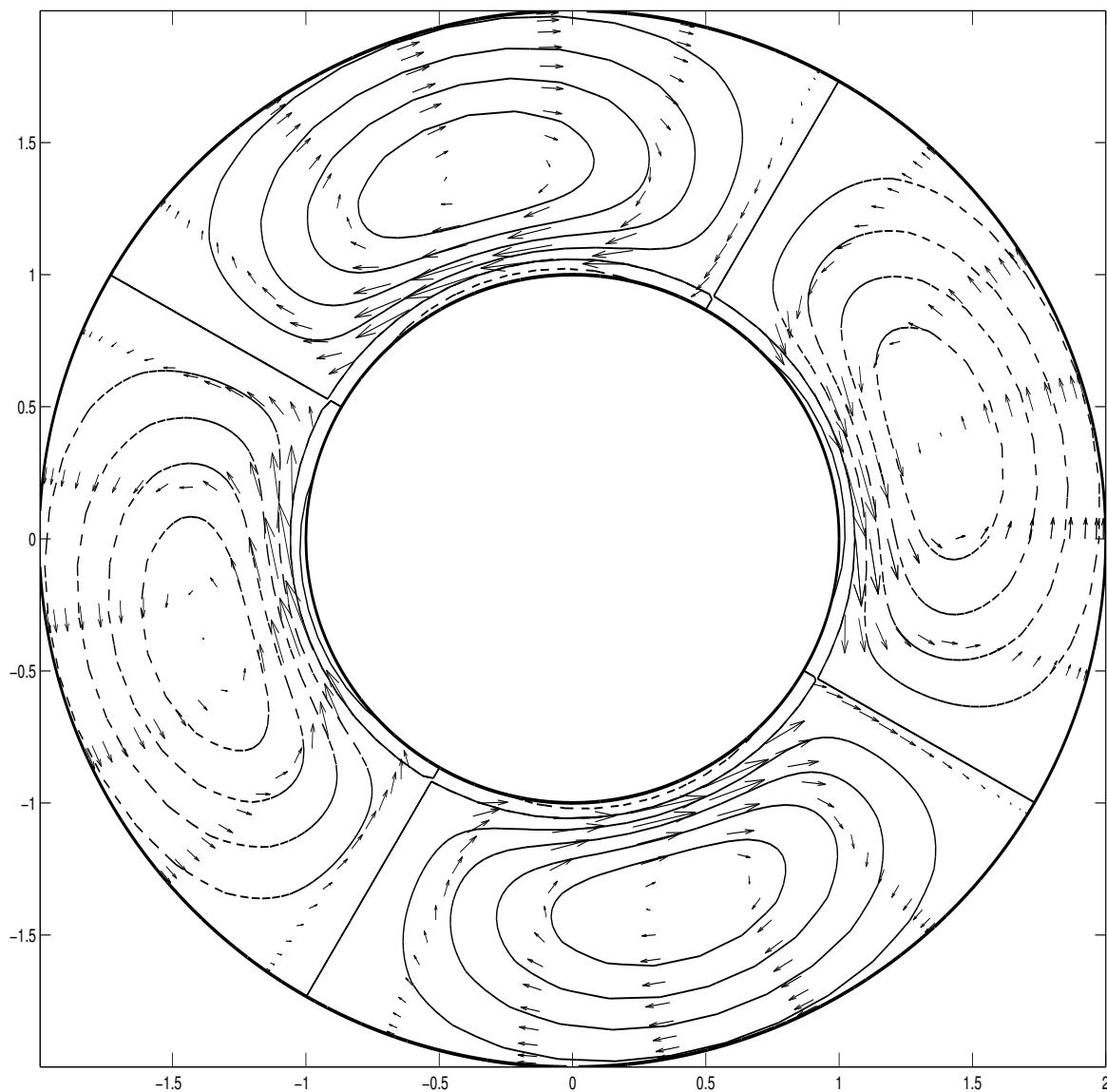
Planetary-scale waves



Replication of planetary waves in a rotating tank

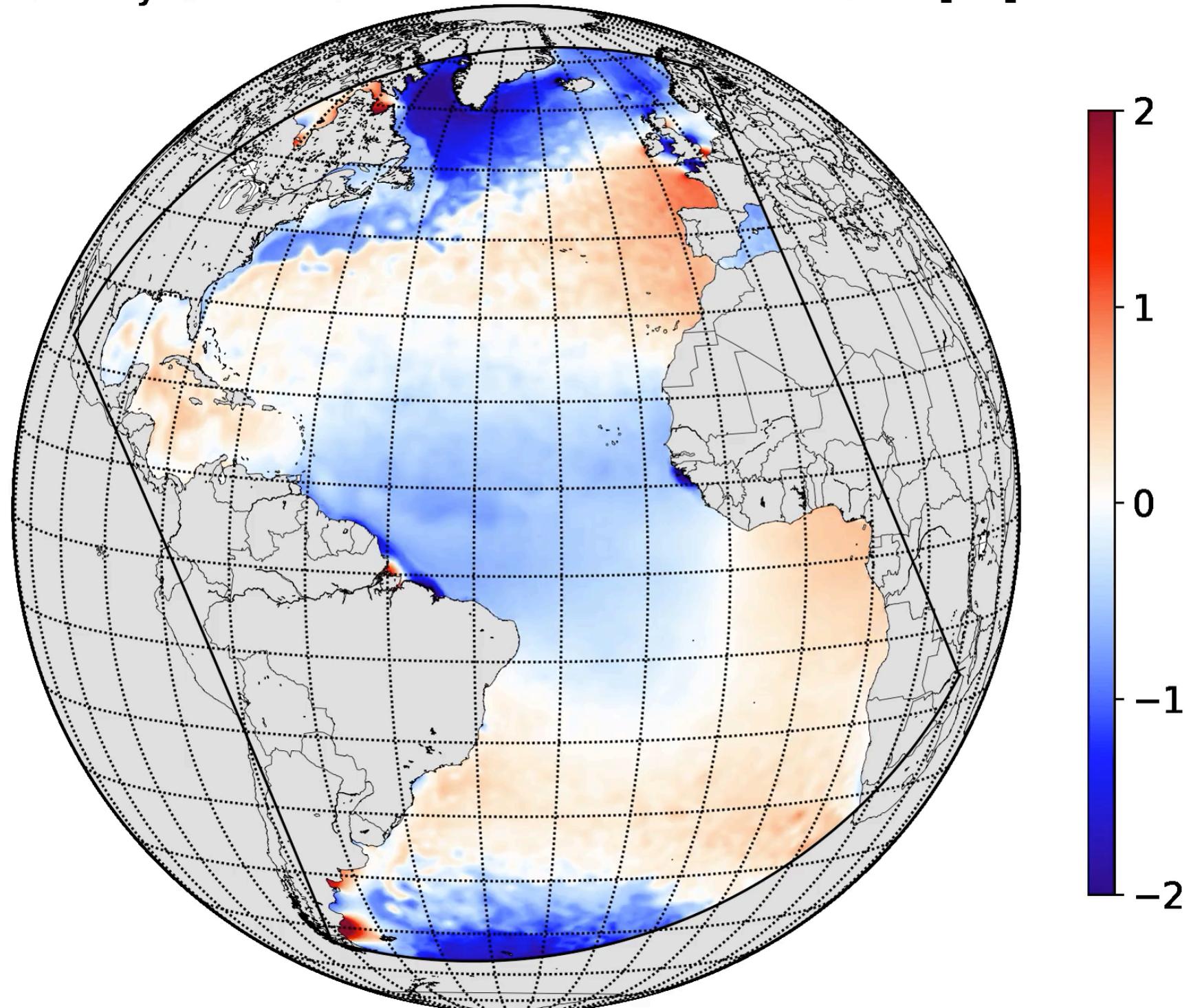


Theory of planetary waves

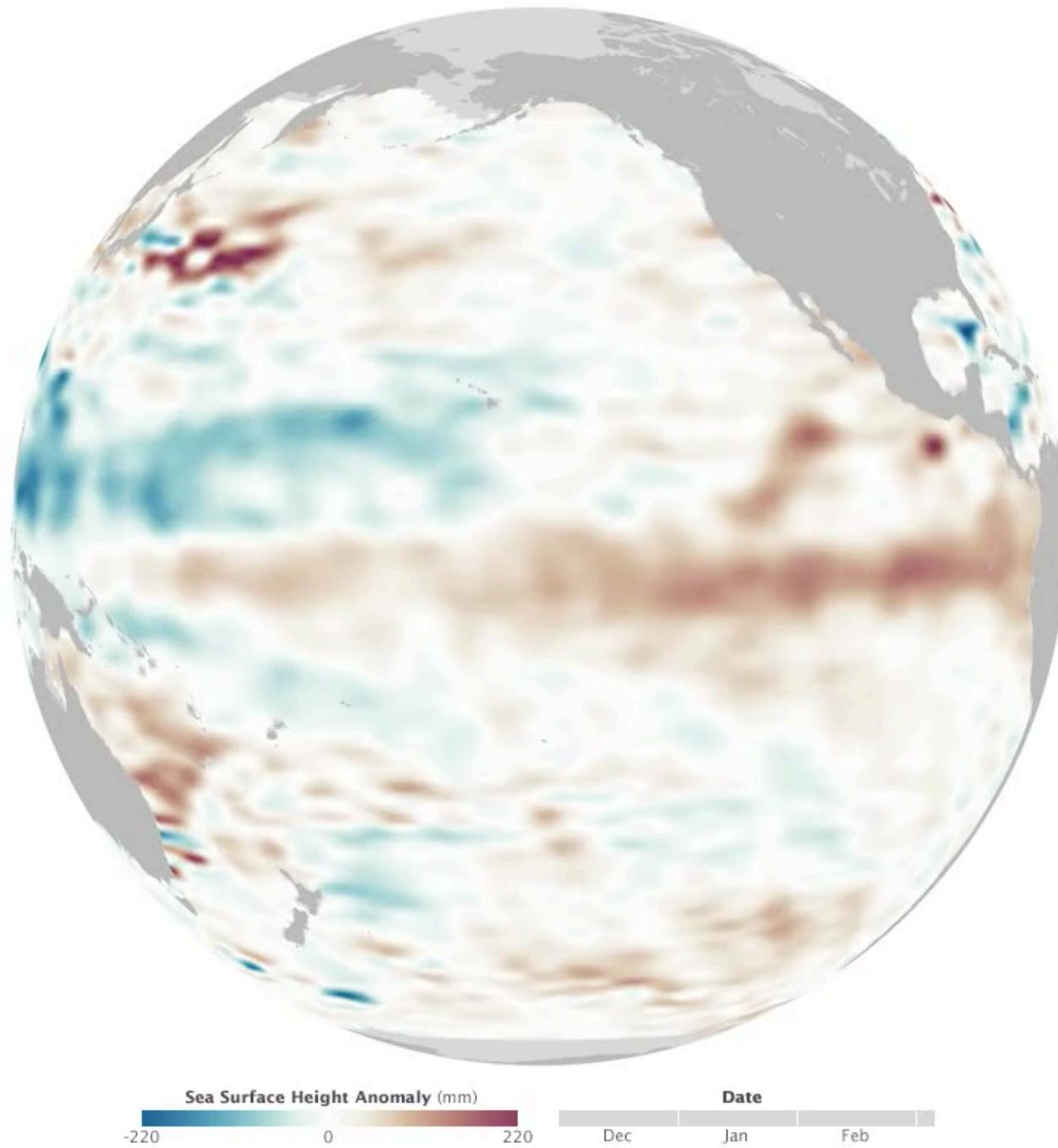


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SSH [m]



Equatorial Kelvin waves seen in sea surface height



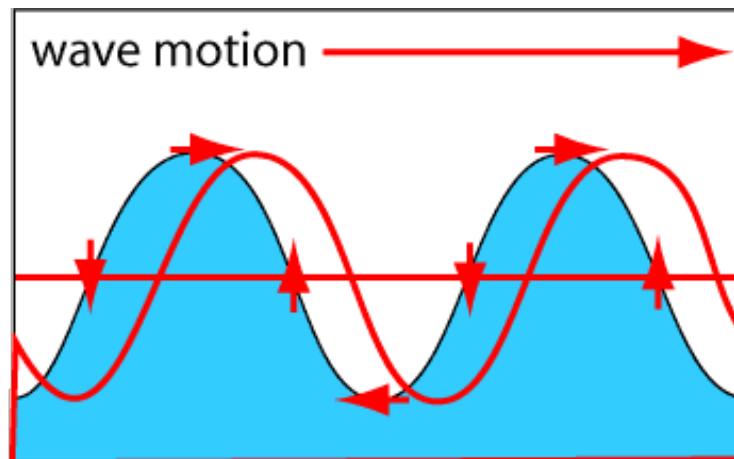
Ocean Waves

Definition of a wave:

- *A wave is a recognizable signal that is transferred from one part of the medium to another with a recognizable velocity of propagation. The signal may be any feature of the disturbance, such as a maximum or an abrupt change in some quantity, provided that it can be clearly recognized and its location at any time can be determined. [Whitham: « Linear and nonlinear waves »]*

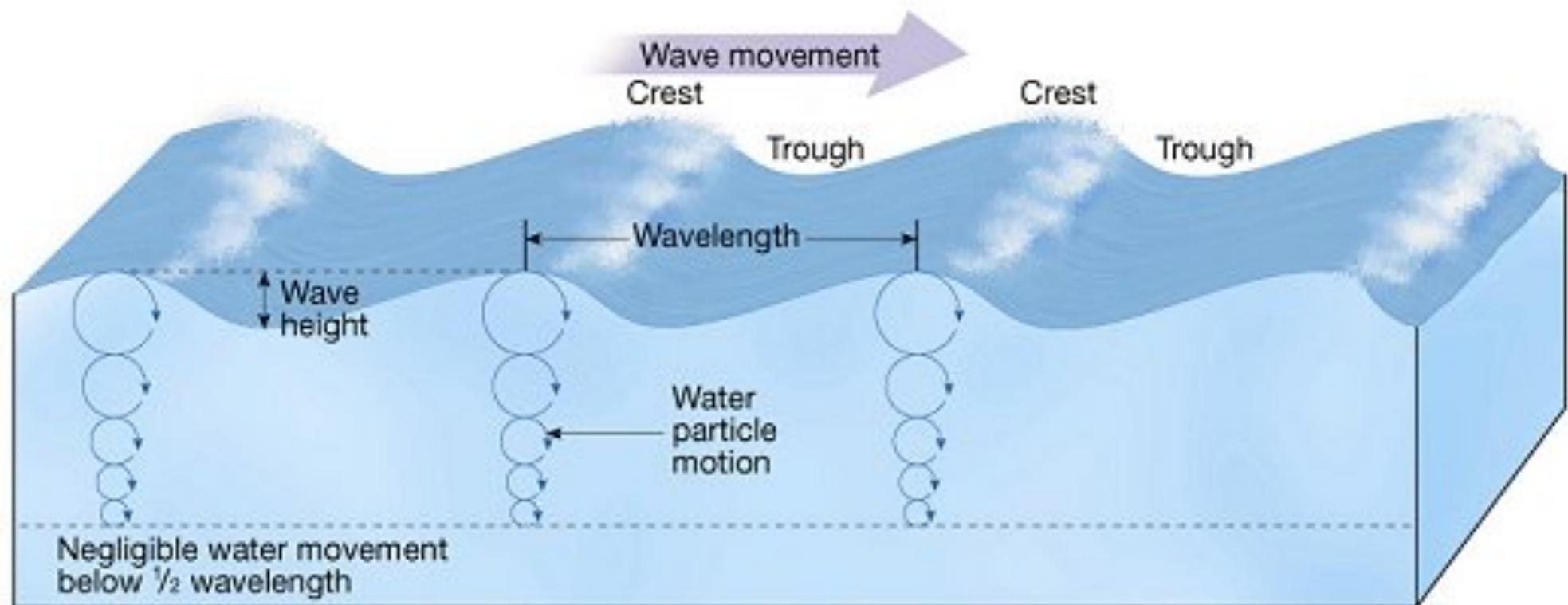
Ocean Waves

- A wave results when fluid is displaced from a position of equilibrium.
- The restoration of the fluid to equilibrium will produce a movement of the fluid back and forth, called a wave orbit.



Ocean Waves

- Waves propagate energy but do not transport water



Ocean Waves

The different type of waves can be classified on the basis of:

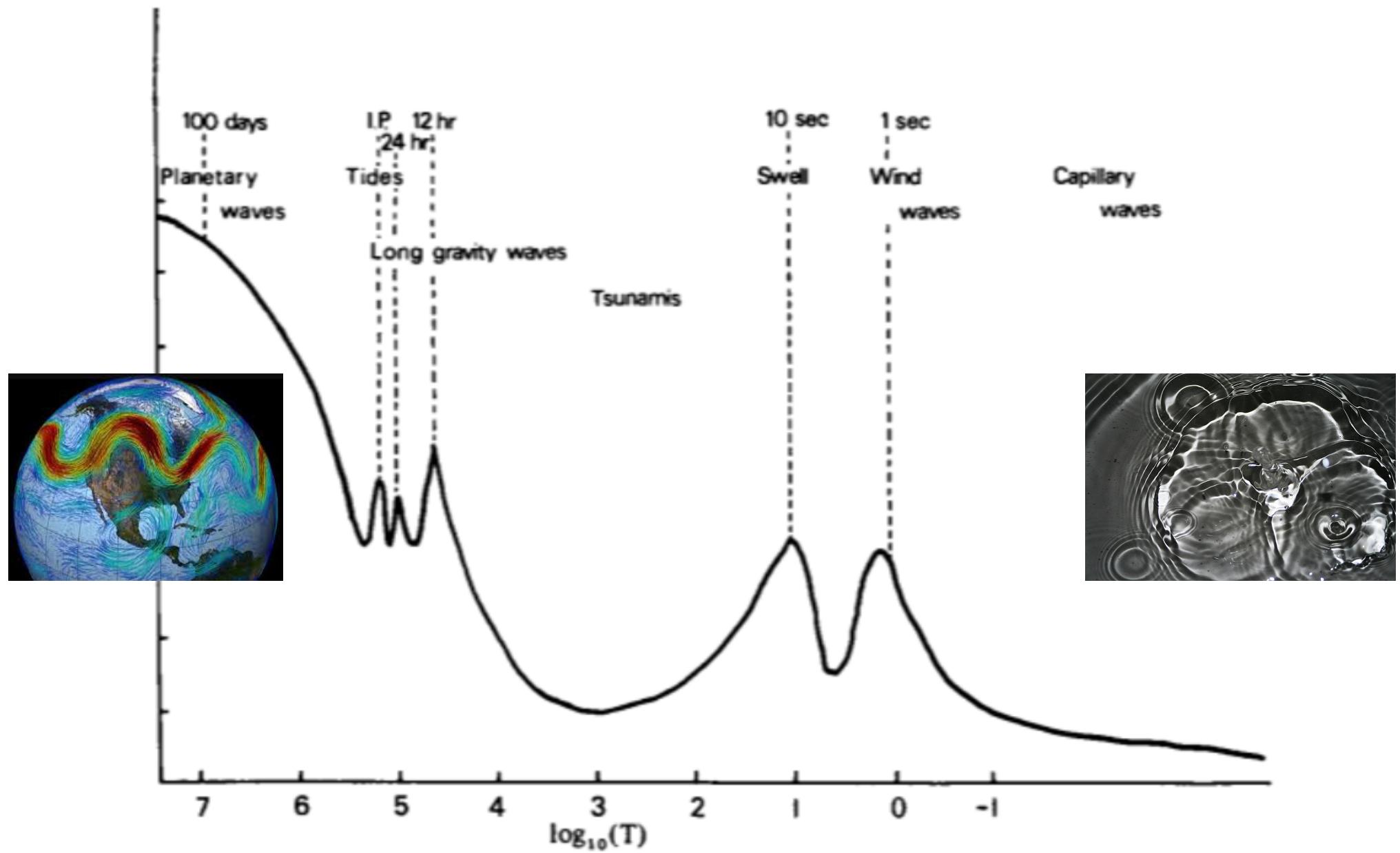
- Restoring force
- Disturbing force
- Wavelength
- Free wave Vs forced wave

Ocean Waves

Ocean Waves

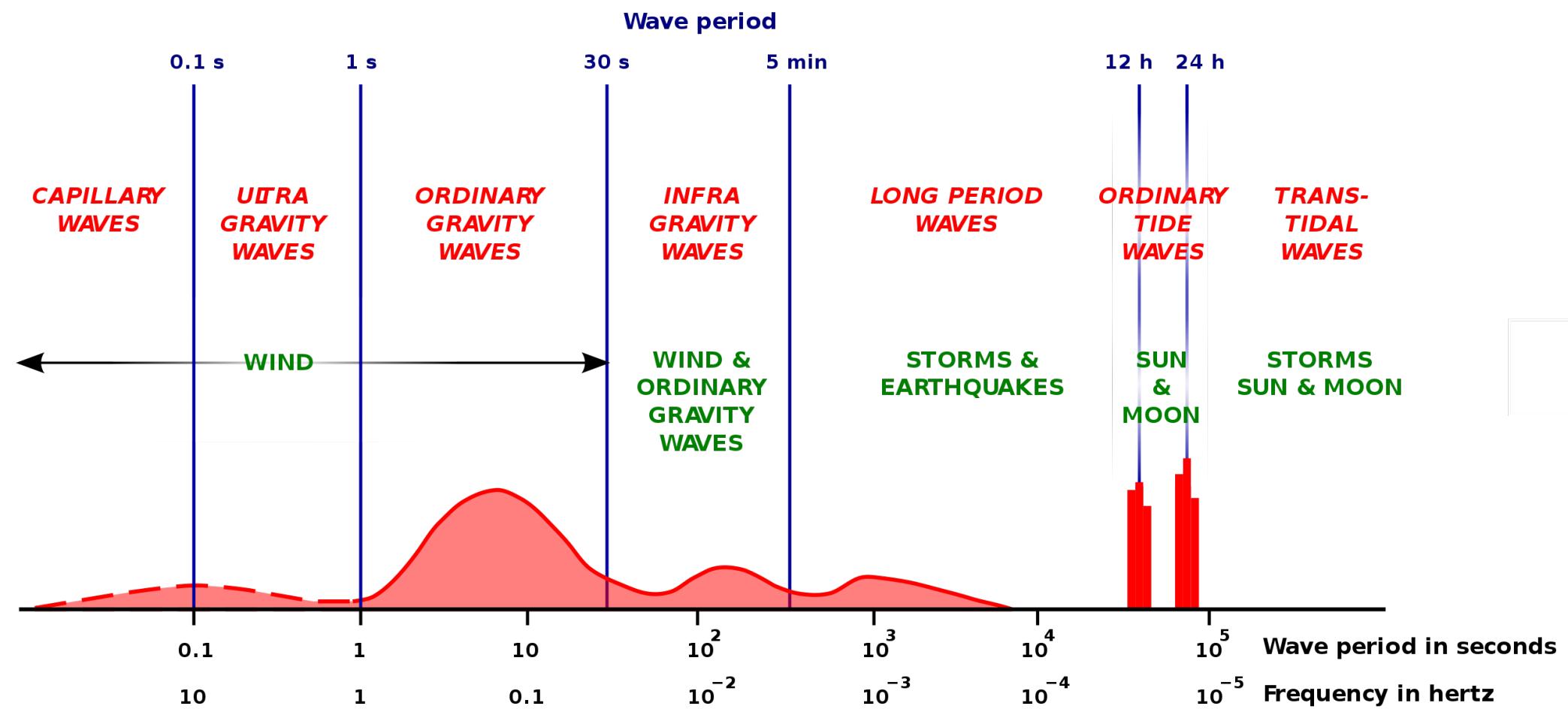
	Disturbing Force	Restoring Force	Wavelength	Periods
Acoustic waves		Pressure (compressibility)	1 mm - 10 km	< 1 s
Capillary waves		Surface tension	1 mm - 1 cm	< 1 s
Surface gravity waves	Wind, boat, earthquake, etc.	Gravity	1 cm - 100 km	1 s - 1 day
Internal waves	Tides, Wind, Topography, etc.	Gravity (stratification) + Coriolis	1 m - 100 km	1 s - 1 day
Rossby waves		PV (variation of Coriolis with latitude)	100 km - 1000 km	Days - Months
Kelvin wave		Pressure gradient + Coriolis	10 km - 100 km	Days - Months
Equatorial Waves		Pressure gradient + Coriolis	10 km - 1000 km	Days - Years

Ocean Waves

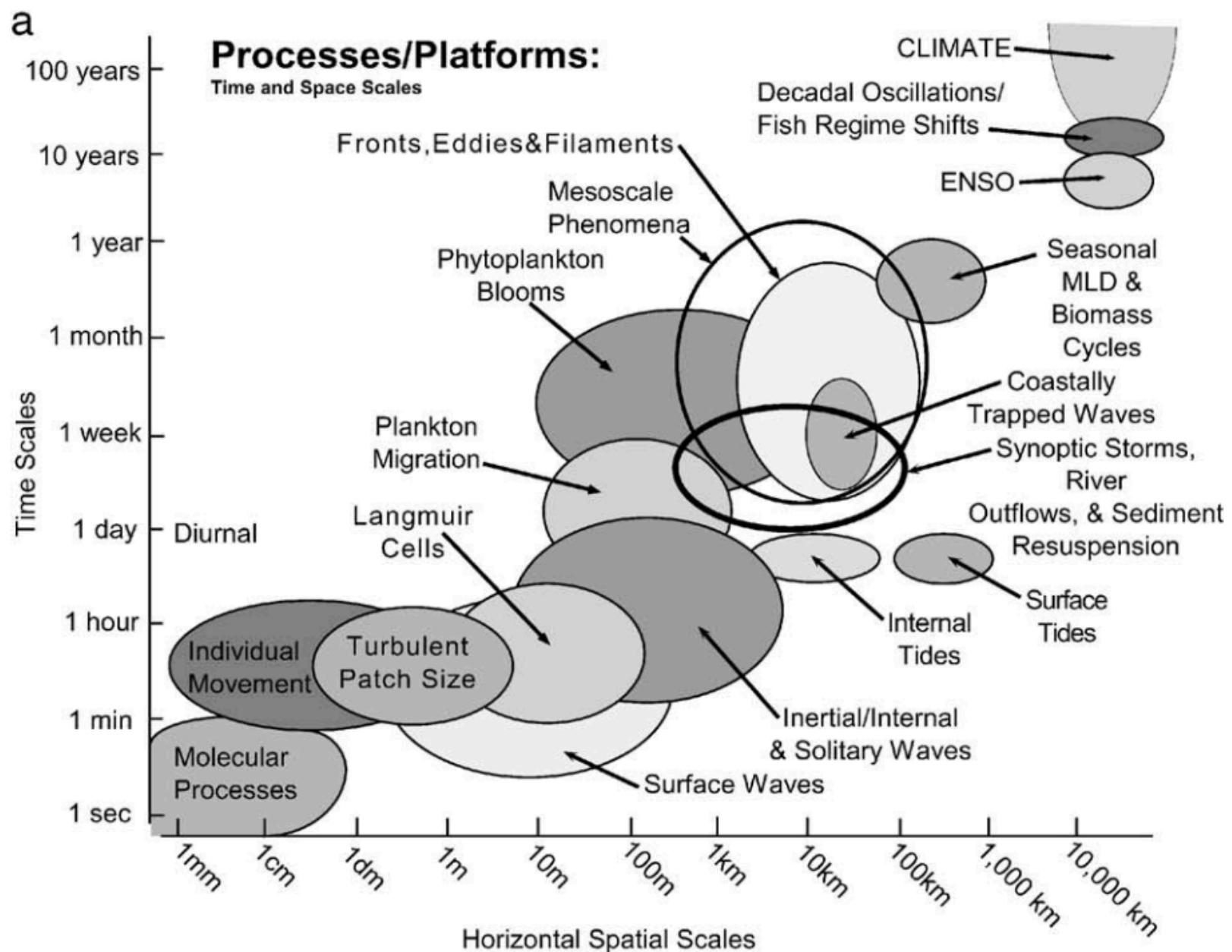


Schematic energy spectrum of ocean variability [Leblond & Mysak]

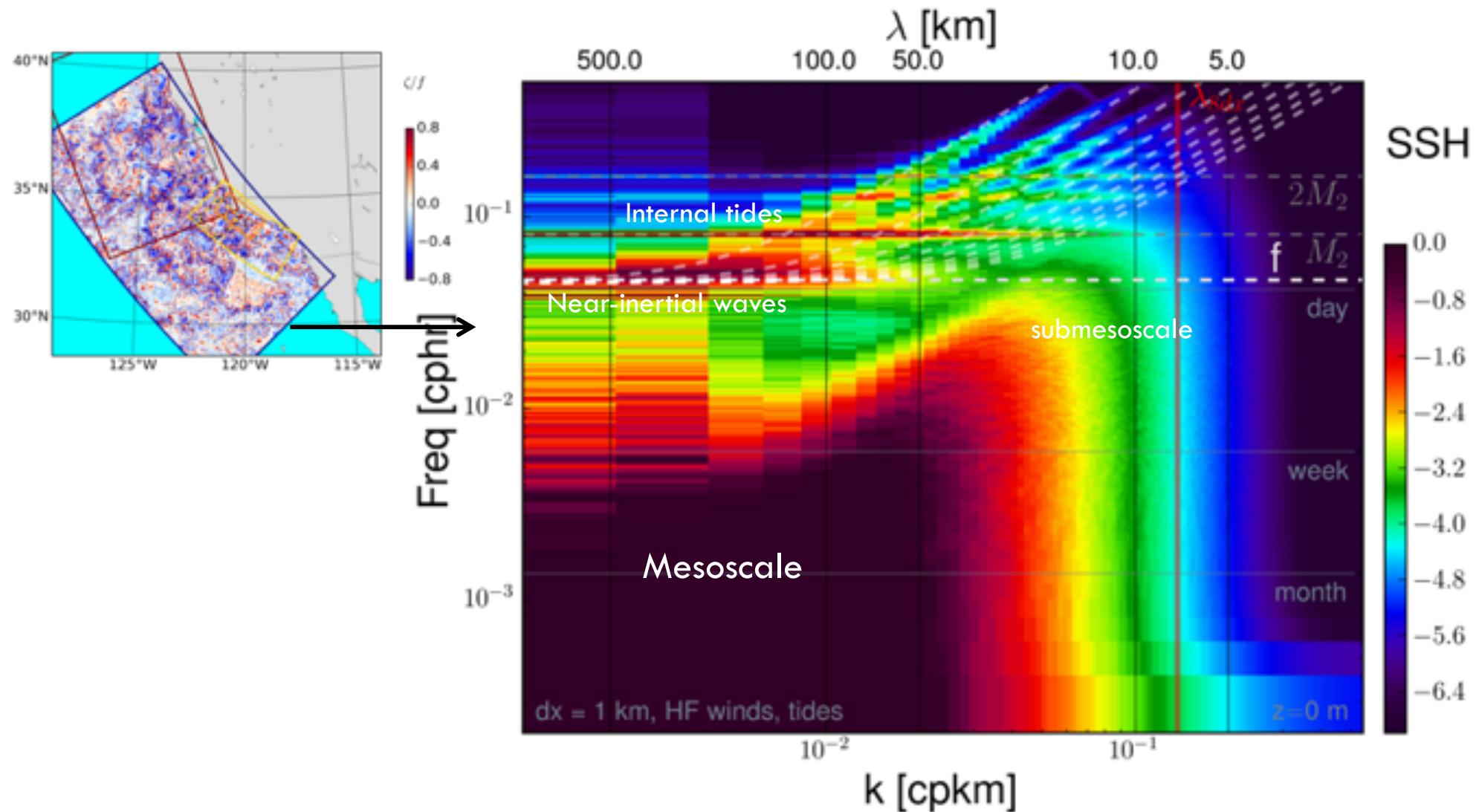
Surface gravity Waves



Ocean Waves



Ocean Waves



Azimuthally-averaged 2D frequency-wavenumber spectra for SSH in California Current

- Mathematically two main classes of waves:

Hyperbolic waves and dispersive waves

Hyperbolic waves

1. **Hyperbolic waves** are formulated in terms of hyperbolic partial differential equations, for example:

$$\eta_t + c\nabla\eta = 0$$

$$\eta_{tt} - c^2\nabla^2\eta = 0$$

Hyperbolic waves

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$$\eta_t + c\nabla\eta = 0$$

$$\eta_{tt} - c^2\nabla^2\eta = 0$$

- With general solutions in the form:

$$\eta = f(x - ct)$$

$$\eta = f(x - ct) + g(x + ct)$$

- Very frequent in acoustics, elasticity, electromagnetism, etc.

Hyperbolic waves

1. Examples of Hyperbolic waves

- Flood wave, tidal bores



- Shock wave



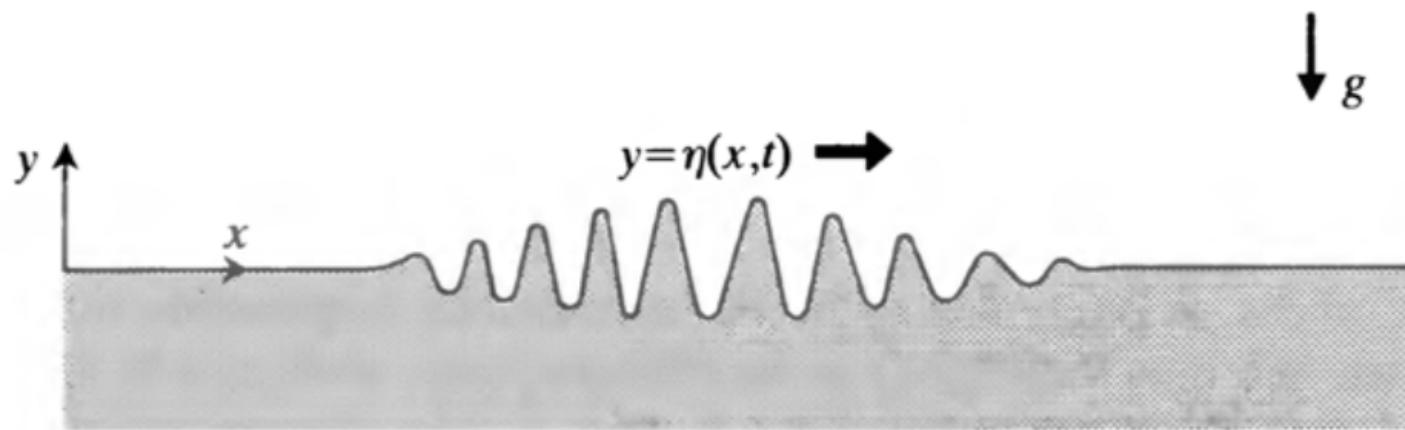
Dispersive waves

2. **Dispersive waves** come from a variety of partial differential equations, they are characterized principally by their dispersion relation:

$$\omega = f(k)$$

Connecting the frequency and the wave number.

They are visualized as a group of waves where the different Fourier components propagate at different speeds



Dispersive waves

2. Dispersive waves:

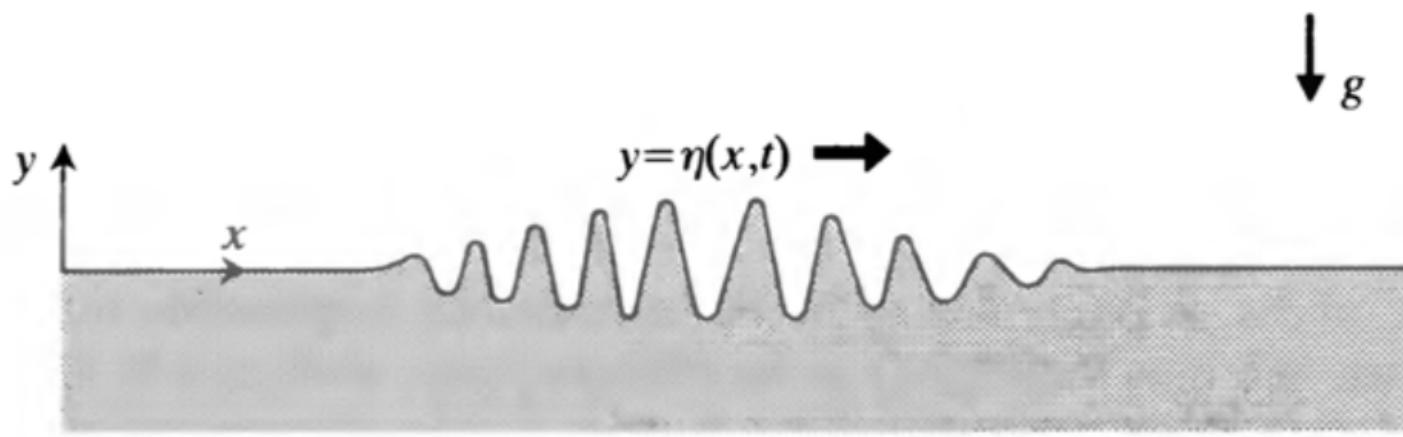
$$\omega = f(k)$$

The wave speed is

$$c = \frac{\omega}{k}$$

Energy propagates with the group velocity:

$$c_g = \frac{\partial \omega}{\partial k}$$

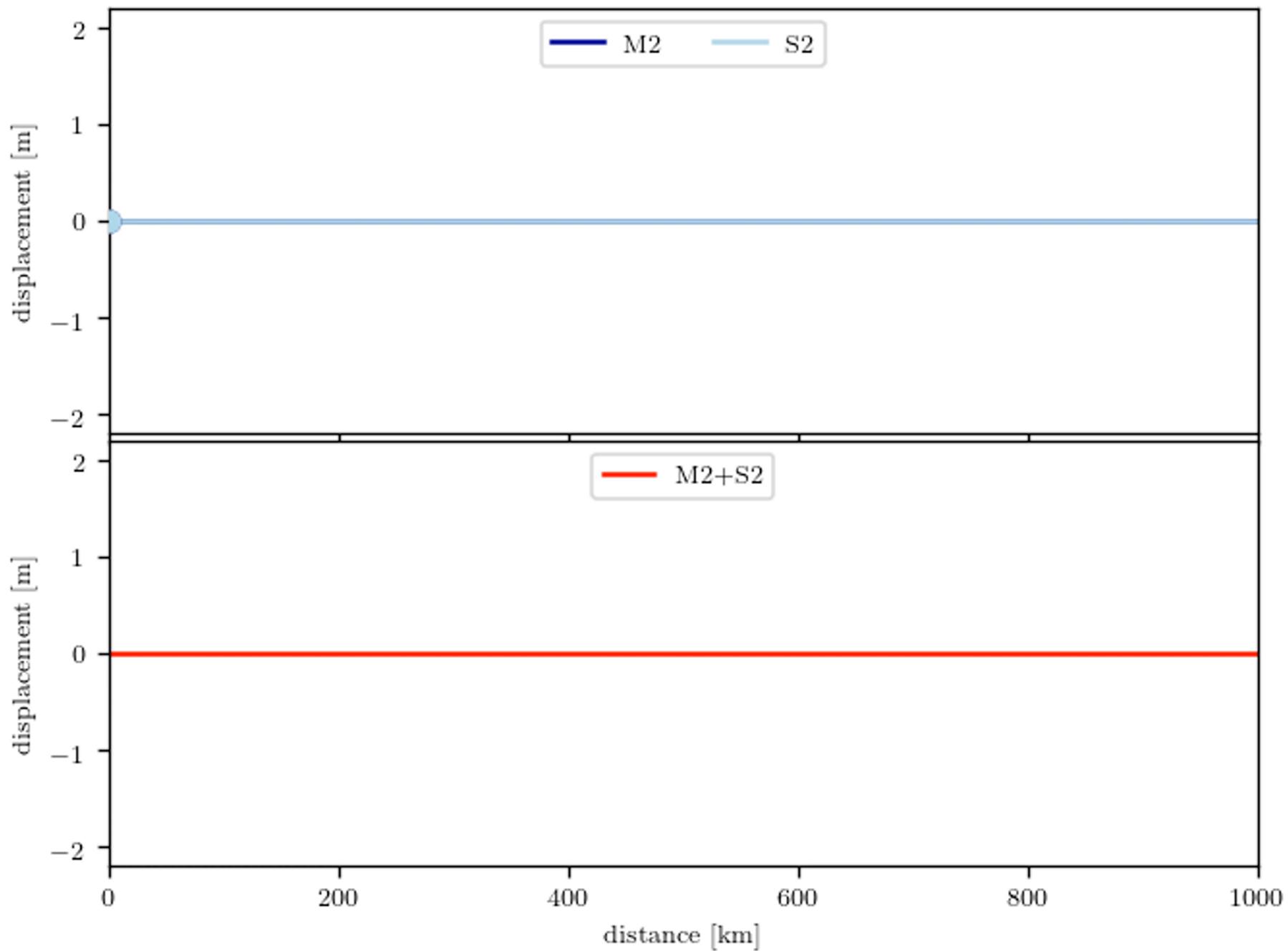


Dispersive waves

Example of the interaction of tidal waves :

- “M2 wave” (M=Moon) $f_{M2} = A_{M2} \exp(i(\omega_{M2}t - k_{M2}x))$
- “S2 wave” (S=Sun). $f_{S2} = A_{S2} \exp(i(\omega_{S2}t - k_{S2}x))$
- f can be the sea surface, current, displacement of isopycnals, etc.

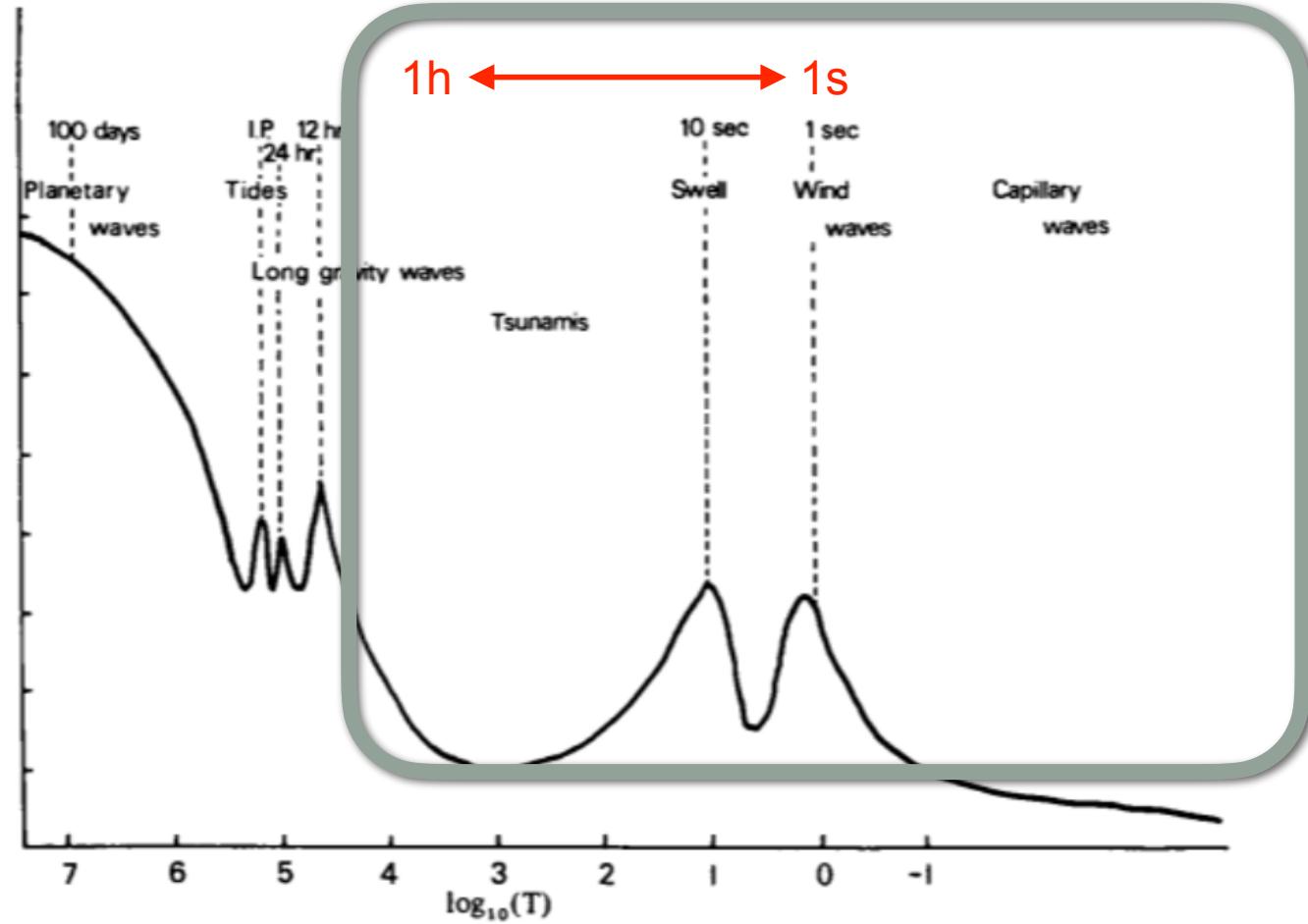
Dispersive waves



OCEAN WAVES

1. Surface waves [very brief reminder]
2. Internal waves
3. Long waves
 - a. Rossby waves [Rossby, Poincare, etc.]
 - b. Coastal trapped waves [Kelvin waves, etc.]
 - c. Equatorial waves [Rossby, Kelvin, Yanai]

1. Surface waves



Schematic energy spectrum of ocean variability [Leblond & Mysak]

1. Surface waves

Solving the linearized incompressible Euler equations:

Incompressible Euler equations (convective or Lagrangian form)

$$\begin{cases} \frac{D\rho}{Dt} = 0 \\ \frac{D\mathbf{u}}{Dt} = -\frac{\nabla p}{\rho} + \mathbf{g} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}$$

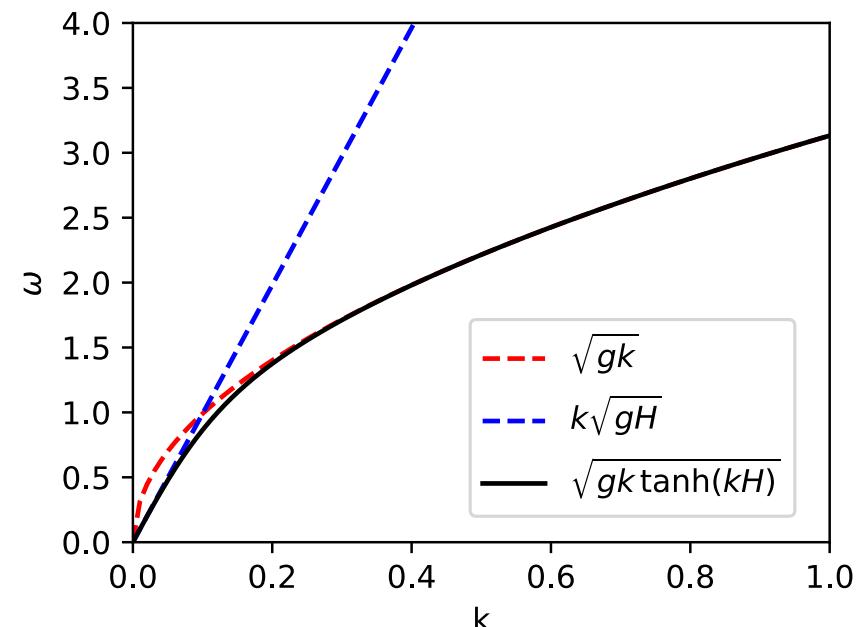
$$\begin{aligned} u_t &= -p_x \\ v_t &= -p_y \\ w_t &= -p_z - g \\ 0 &= u_x + v_y + w_z, \end{aligned}$$

Looking for a wave solution: $\Phi = F(z)e^{i(\vec{k} \cdot \vec{x} - \omega t)}$

We get the dispersion relation :

$$\omega = \sqrt{gk \tanh(kH)} \text{ (black line)}$$

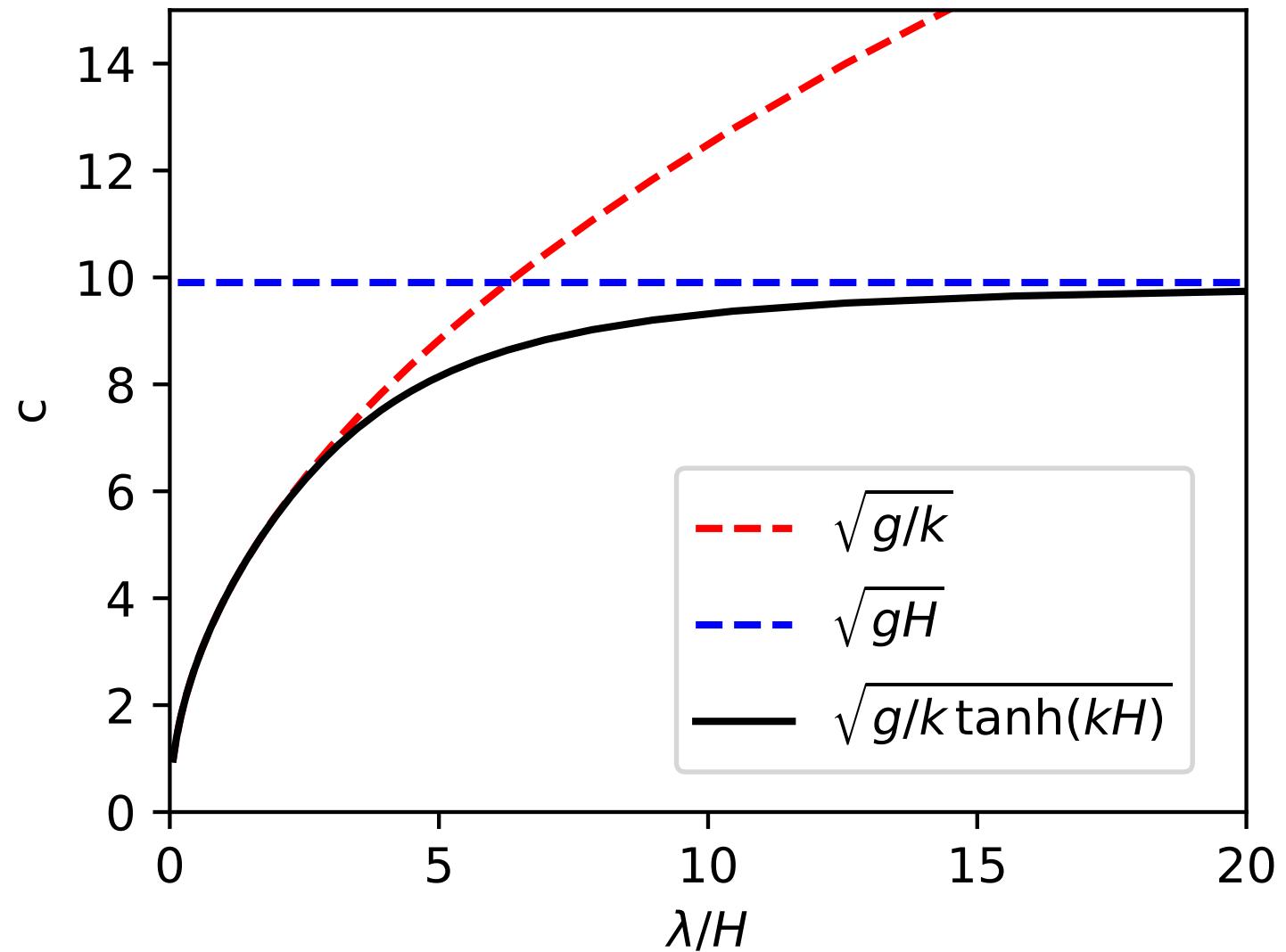
With H the water depth.



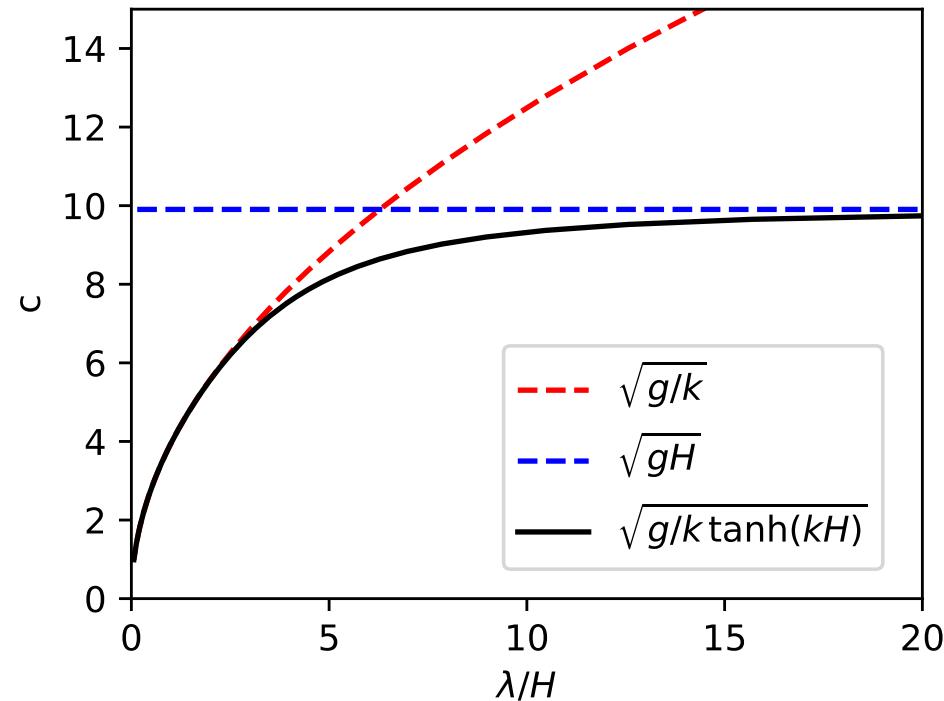
1. Surface waves

Phase speed:

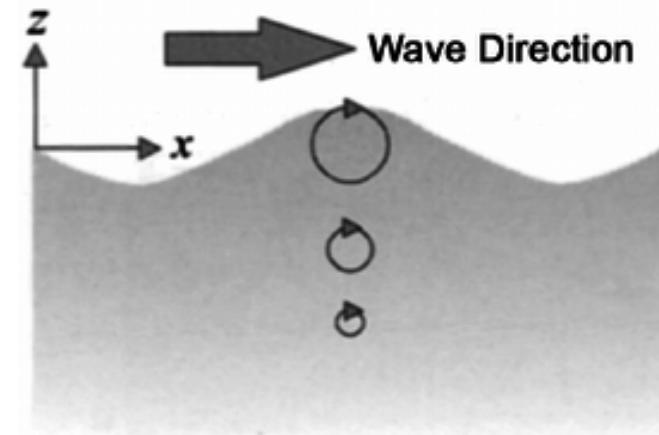
$$c = \frac{\omega}{k}$$



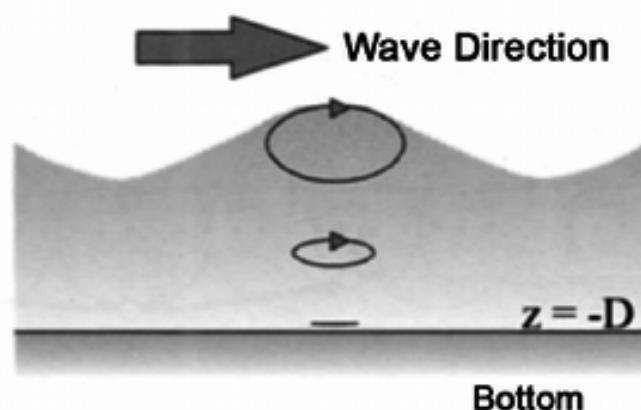
1. Surface waves



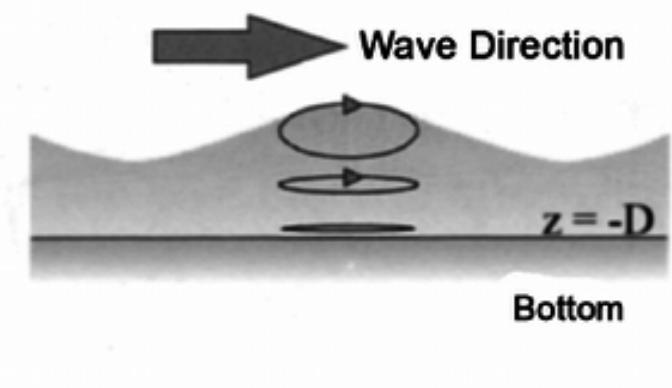
Deep Water



Intermediate Depth

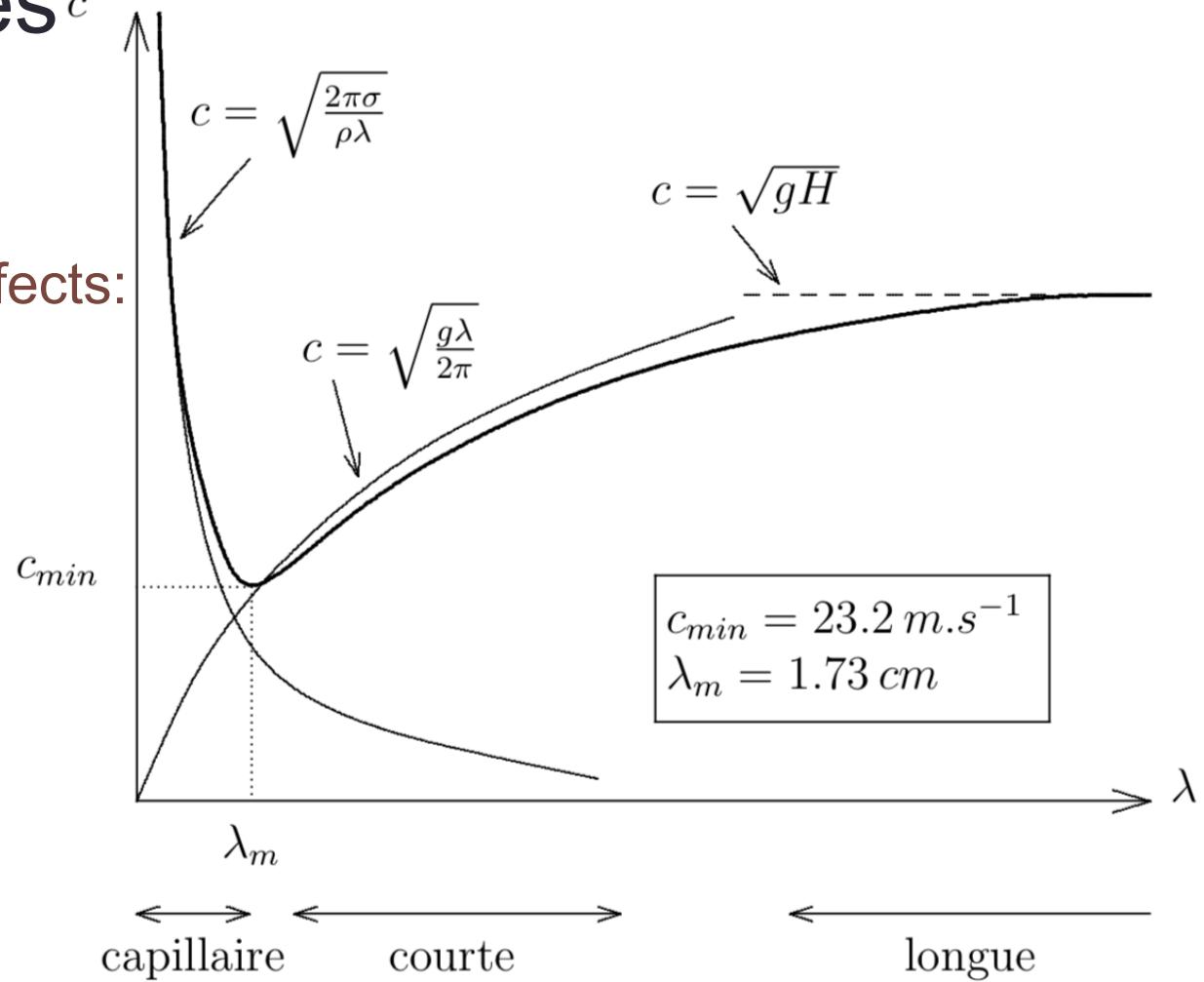


Very Shallow Water



1. Surface waves^c

With surface tension effects:



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