

# Notes on SCVs statistics

April 10, 2020

## 1 Deep SCVs in the Atlantic

Figures 1 to 5 are showing horizontal sections of relative vorticity at various depths. A lot of SCVs, mostly anticyclones, are visible. Their spatial distribution seems quite different depending on the depths.

We need a good identification algorithm to be able to isolate coherent vortical structures. Questions are:

- Is the Okubo-Weiss number a good criteria (following ?). Or do we need to consider the angular momentum instead (following ?)?
- Can we adapt py-eddy-tracker to do the work?

Then once we have a working algorithm for detection, objectives will be to :

- compute the number of vortices (or the probability of presence) as a function of lon, lat, depth, isopycnal.
- compute SCV's parameters: radius, thickness, stratification (N), Rossby number, Richardson number, etc.
- Associate the vertical T/S profiles with the cyclones/anticyclones to check if we will be able to detect them using the argo dataset.
- Add the tracking to be able to understand the life-cycle of the SCVs.

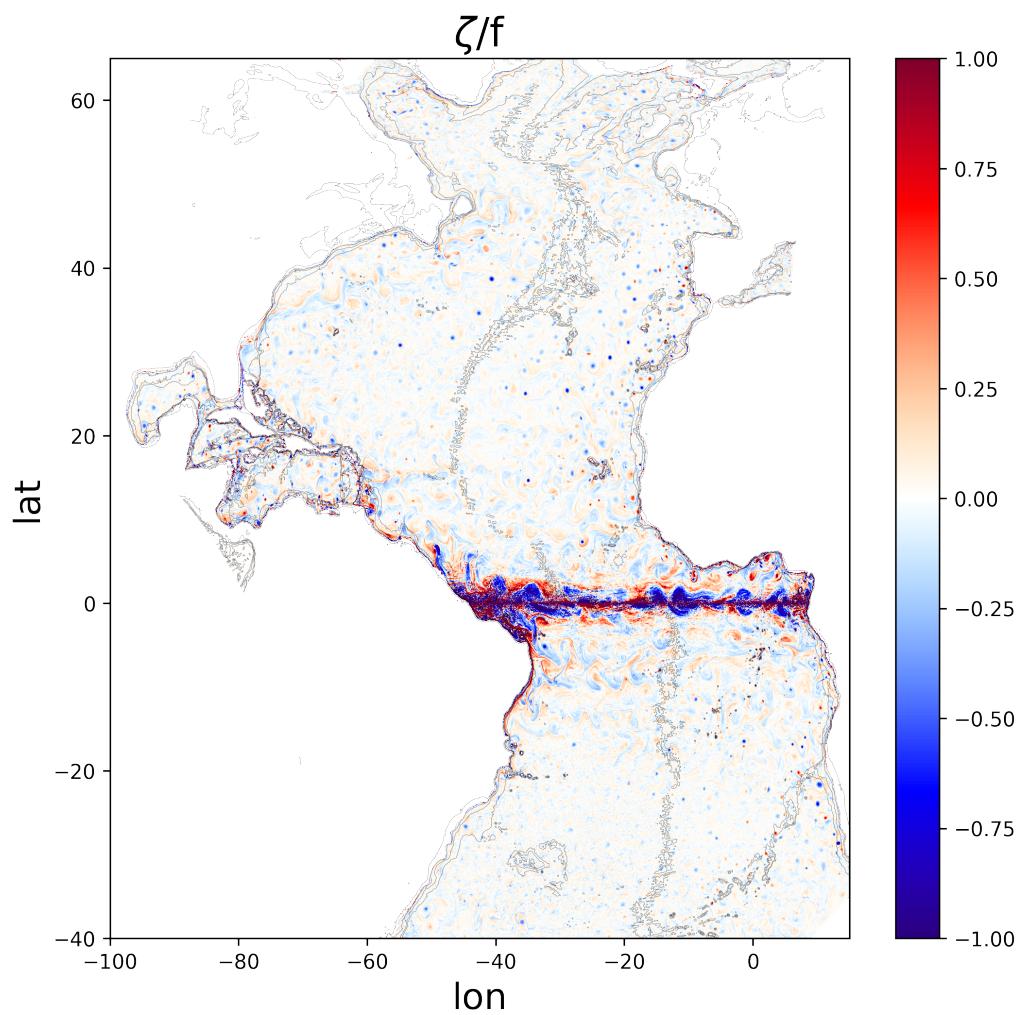


Figure 1: vorticity at 500 m depth

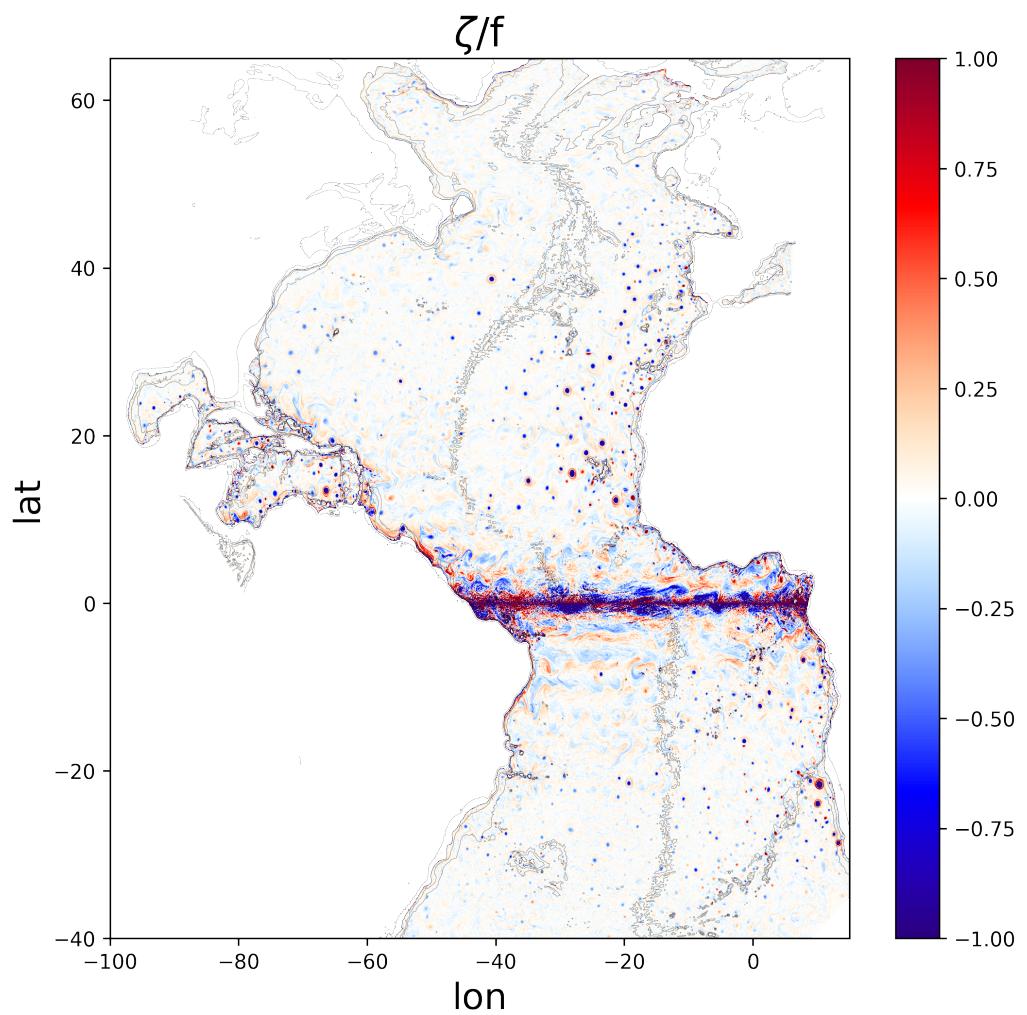


Figure 2: vorticity at 1000 m depth

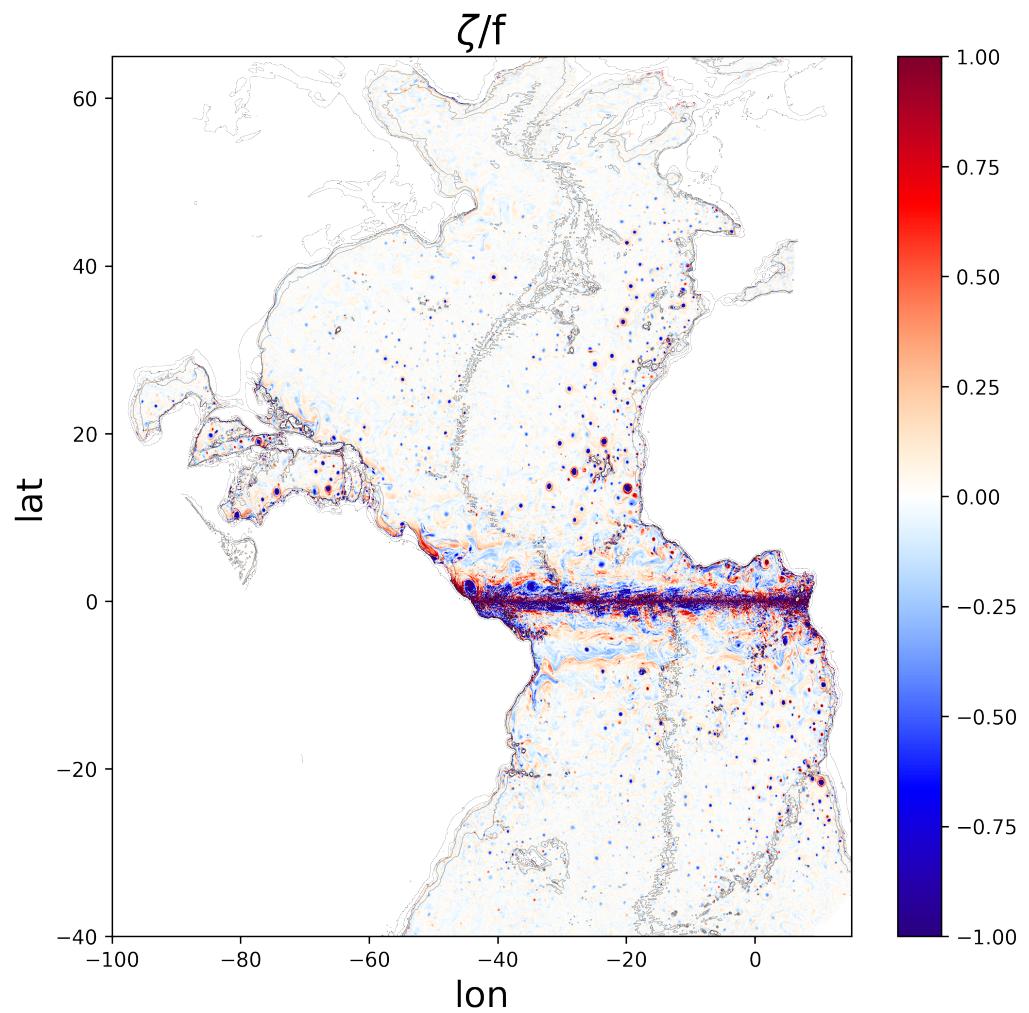


Figure 3: vorticity at 1500 m depth

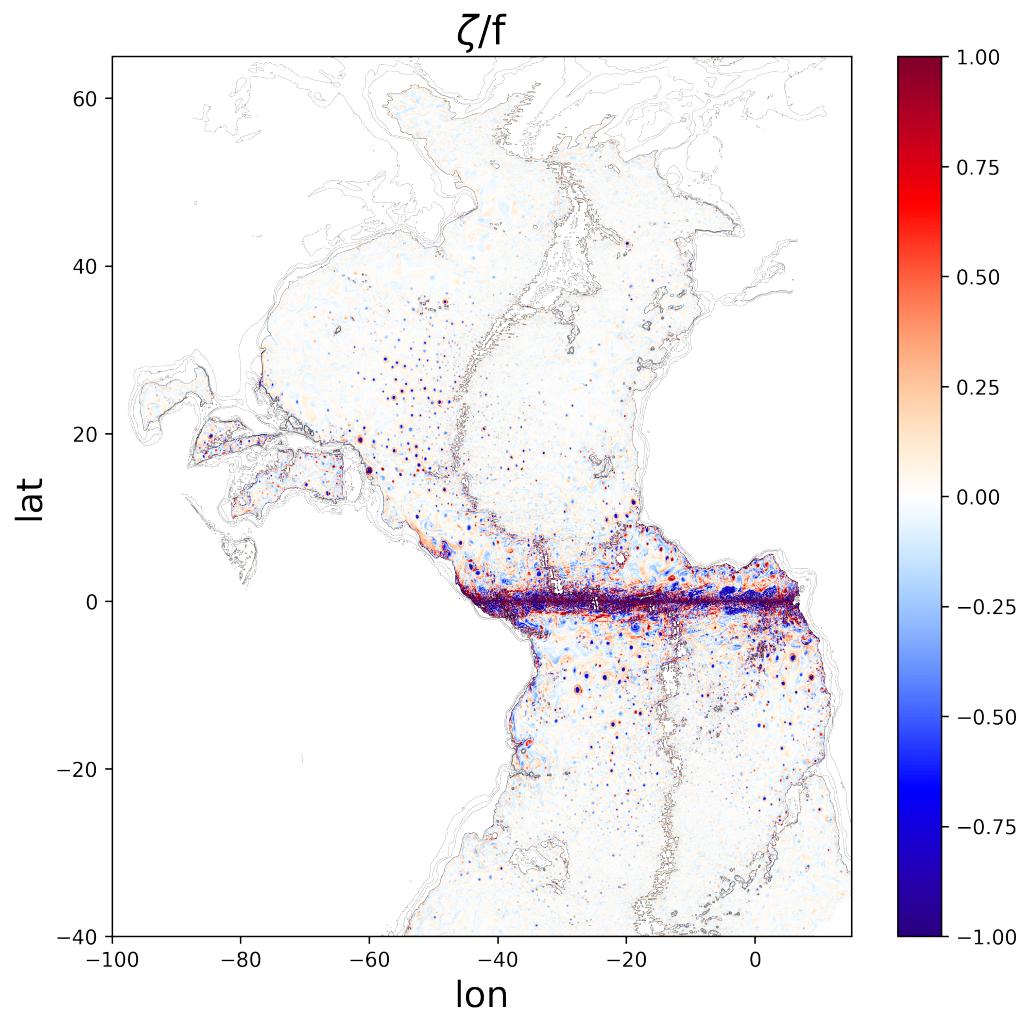


Figure 4: vorticity at 3000 m depth

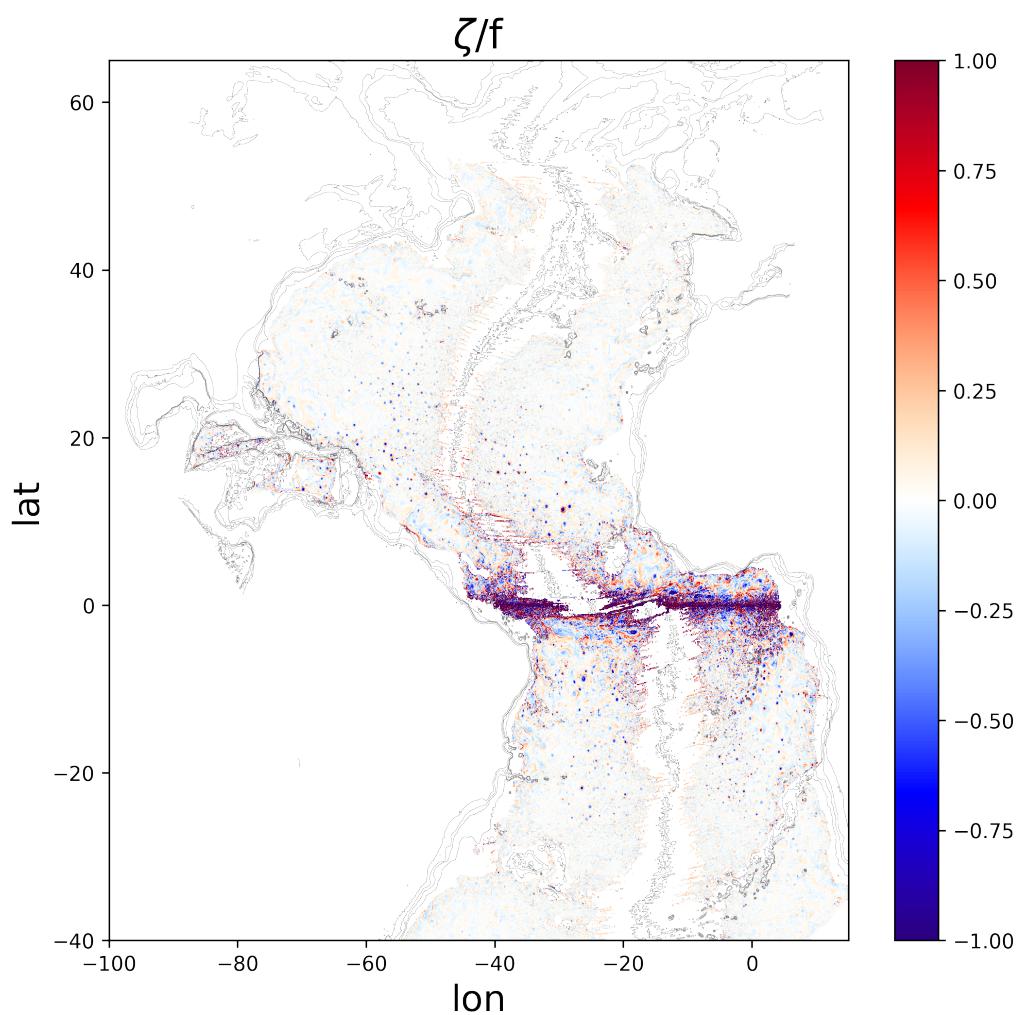


Figure 5: vorticity at 4000 m depth

## 2 Evaluating a few parameters for deep SCVs in the Atlantic

Important parameters to study the stability of the flow, and in particular vortices are:

- the gradient Rossby number:

$$Ro = \frac{\zeta}{f}$$

- the gradient Richardson number:

$$Ri = \frac{N^2}{|\partial_z u|^2}$$

with  $\zeta$  the vertical component of relative vorticity,  $f$  the Coriolis parameter,  $N^2 = \partial_z b$  the vertical stratification, and  $\partial_z u$  the vertical velocity shear.

The classical condition for instability for a front in thermal wind balance as derived by Hoskins (1974) can be written:

$$1 + Ro - Ri^{-1} < 0$$

Following Christian's work, in the presence of *curvature*, *i.e.*, when taking into account this criteria may be rewritten as:

$$1 + Ro - (1 + Cu)Ri^{-1} < 0$$

with the Curvature number  $Cu = \frac{2v}{fr}$ , where  $v$  is the azimuthal velocity and  $r$  the radius.

For a true axisymmetric vortex the criteria even becomes

$$(1 + Cu)(1 + Ro) - (1 + Cu)^2 Ri^{-1} < 0$$

### 2.1 Some examples

Figures 6 and 7 are showing the values of  $Ro$  and  $Ri$  as well as PV at 1000 m depth from a 3 km simulation over the western subtropical gyre and a zoom over one specific very intense anticyclone. We are not able to compute  $Cu$  yet as it requires knowledge of the vortex center (once the detection algorithm is working, that will be easy to do).

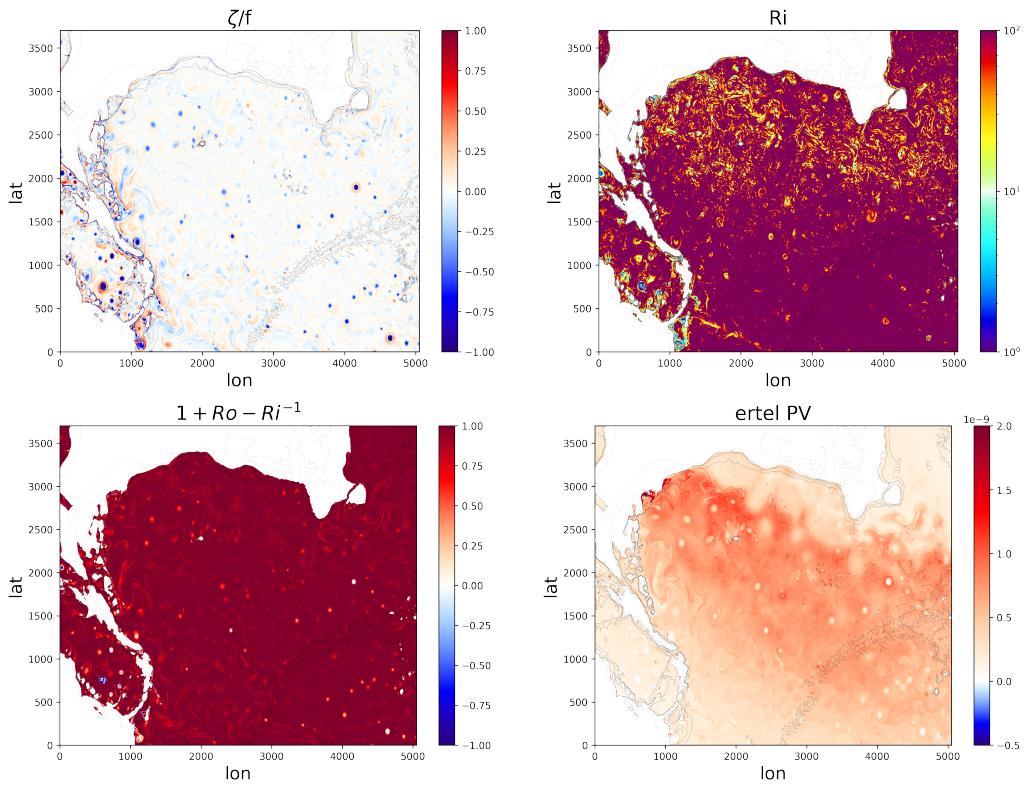


Figure 6: (a) Rossby Number, (b) Richardson Number, (c) Rayleigh's PV  $1 + Ro - Ri^{-1}$ , and full ertel PV at 1000 m depth from GIGATL3 zoomed over the western part of the subtropical gyre.

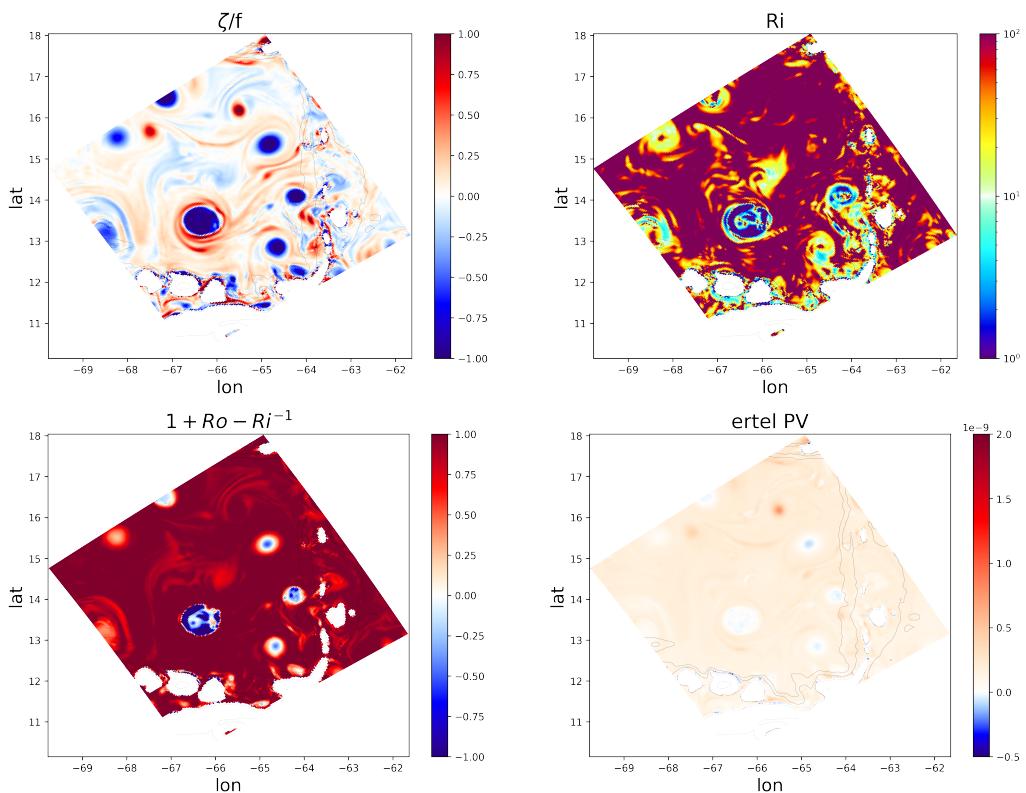


Figure 7: Zoom of Fig. 6.

## 2.2 Some statistics

We look globally at the statistics of some parameters ( $Ri, N$  for now) for cyclones and anticyclones.

Below are example of statistics at various depths for a snapshot of GIGATL3. They include horizontal sections of vorticity as well as distributions of  $Ri$ ,  $N/f$  and  $1 + Ro - Ri^{-1}$  for the full domain (minus the equatorial region), and only for “anticyclones” and “cyclones” here crudely defined as points with  $Ro < -0.5$  and  $Ro > 0.5$ , respectively.

## References

- CHELTON, DUDLEY B., SCHLAX, MICHAEL G., SAMELSON, ROGER M. & DE SZOKE, ROLAND A. 2007 Global observations of large oceanic eddies. *Geophys. Res. Lett.* **34** (15).
- HOSKINS, B. J. 1974 The role of potential vorticity in symmetric stability and instability. *Q.J.R. Meteorol. Soc.* **100**, 480–482.
- LE VU, BRIAC, STEGNER, ALEXANDRE & ARSOUZE, THOMAS 2018 Angular momentum eddy detection and tracking algorithm (ameda) and its application to coastal eddy formation. *Journal of Atmospheric and Oceanic Technology* **35** (4), 739–762.

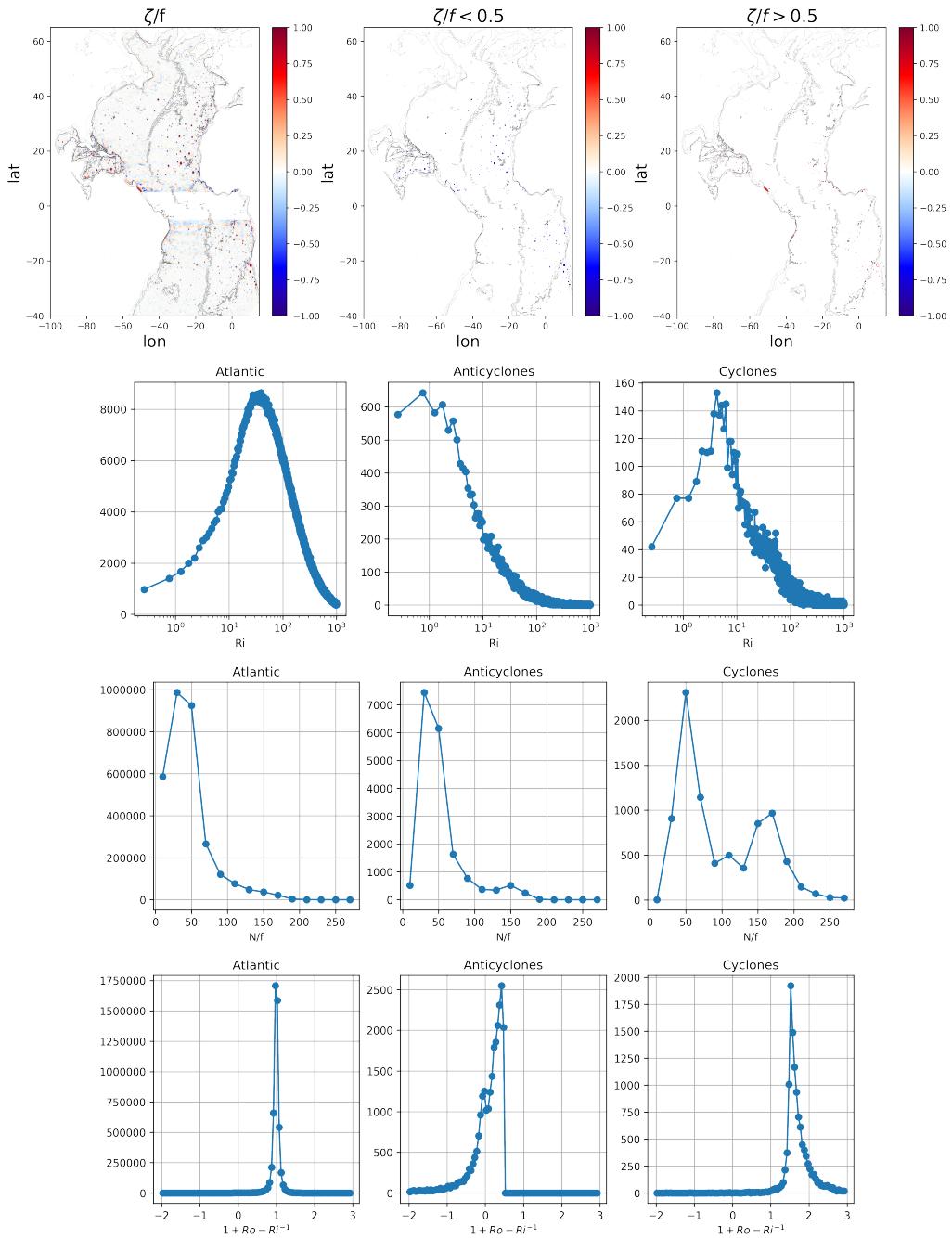


Figure 8: Statistics at 1000 m depth

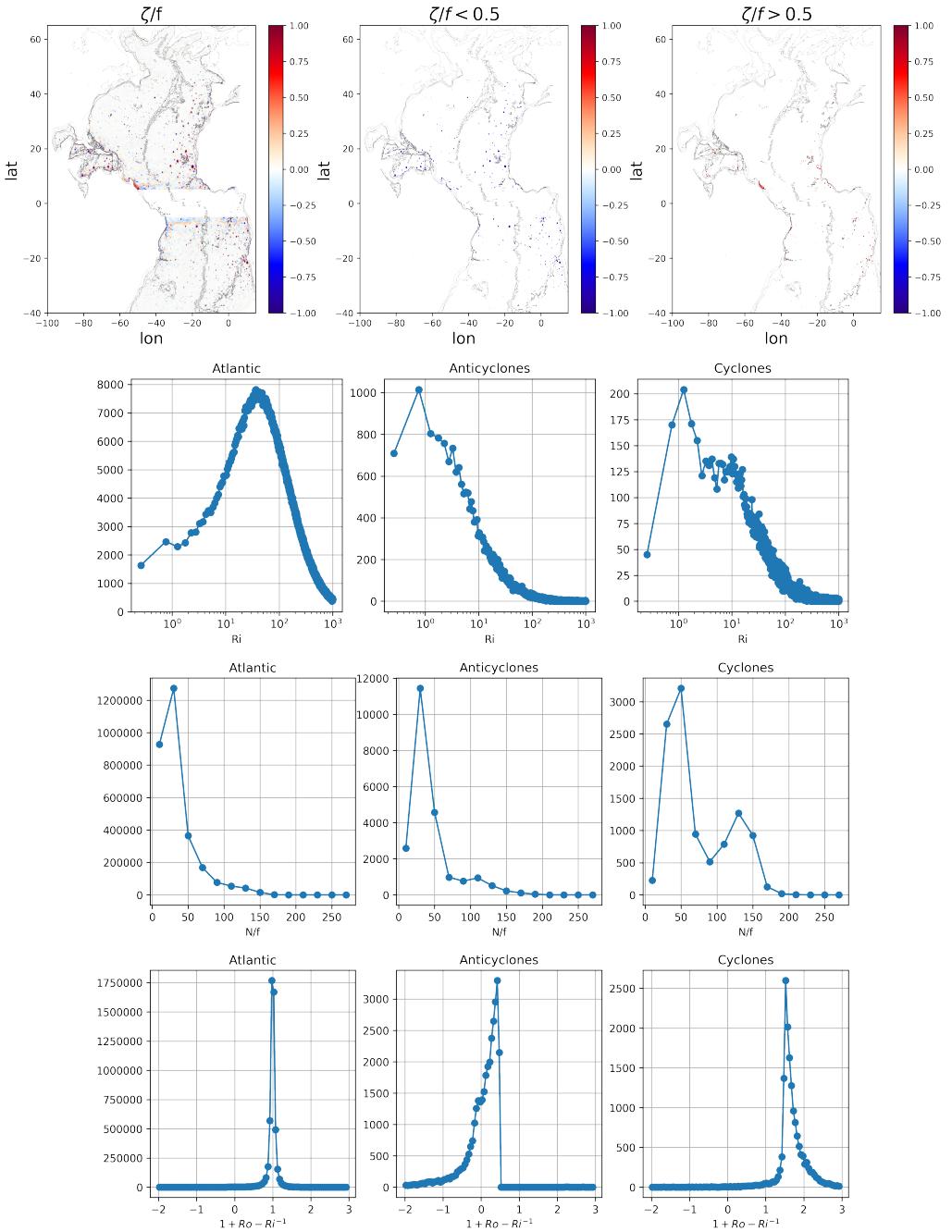


Figure 9: Statistics at 1500 m depth

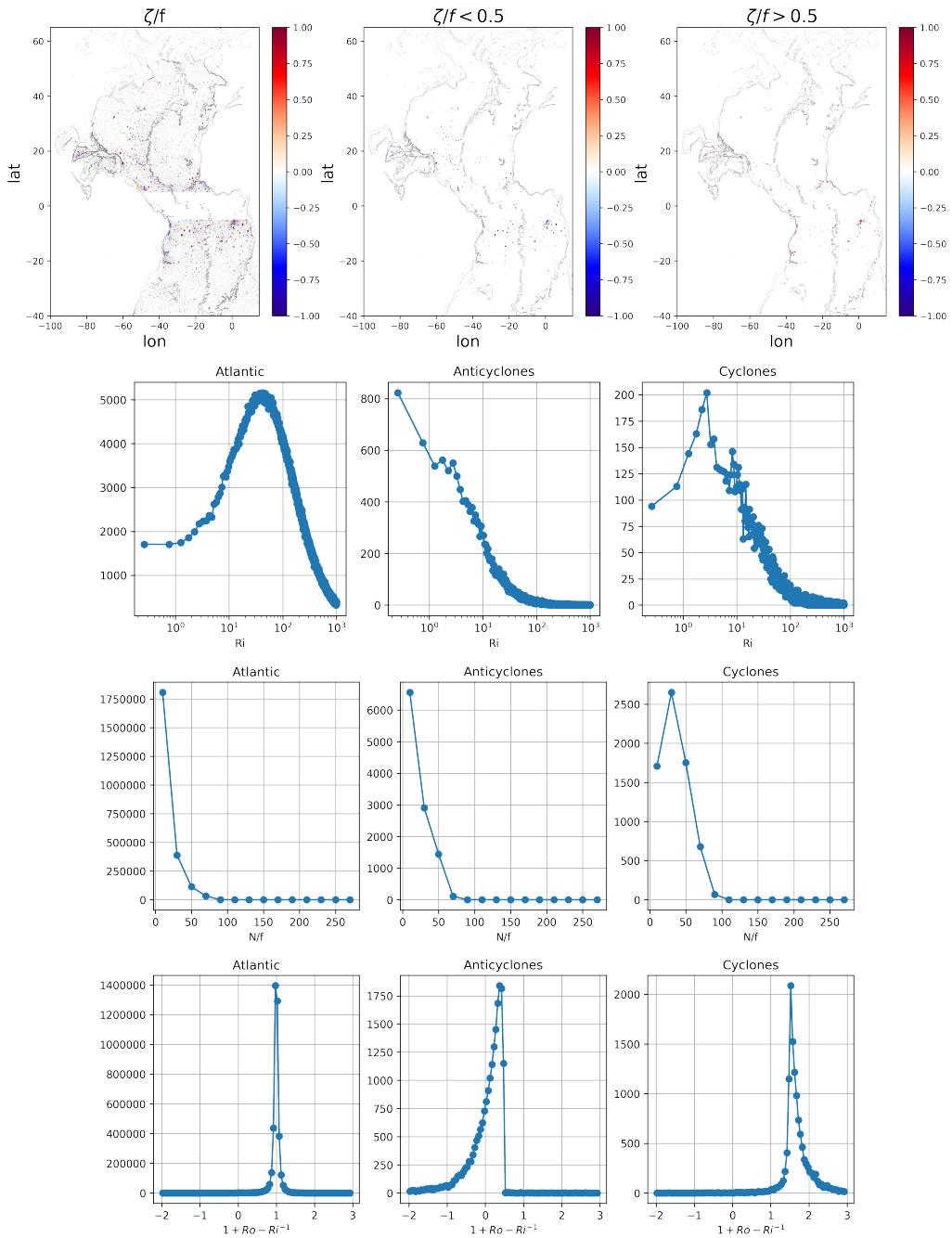


Figure 10: Statistics at 3000 m depth

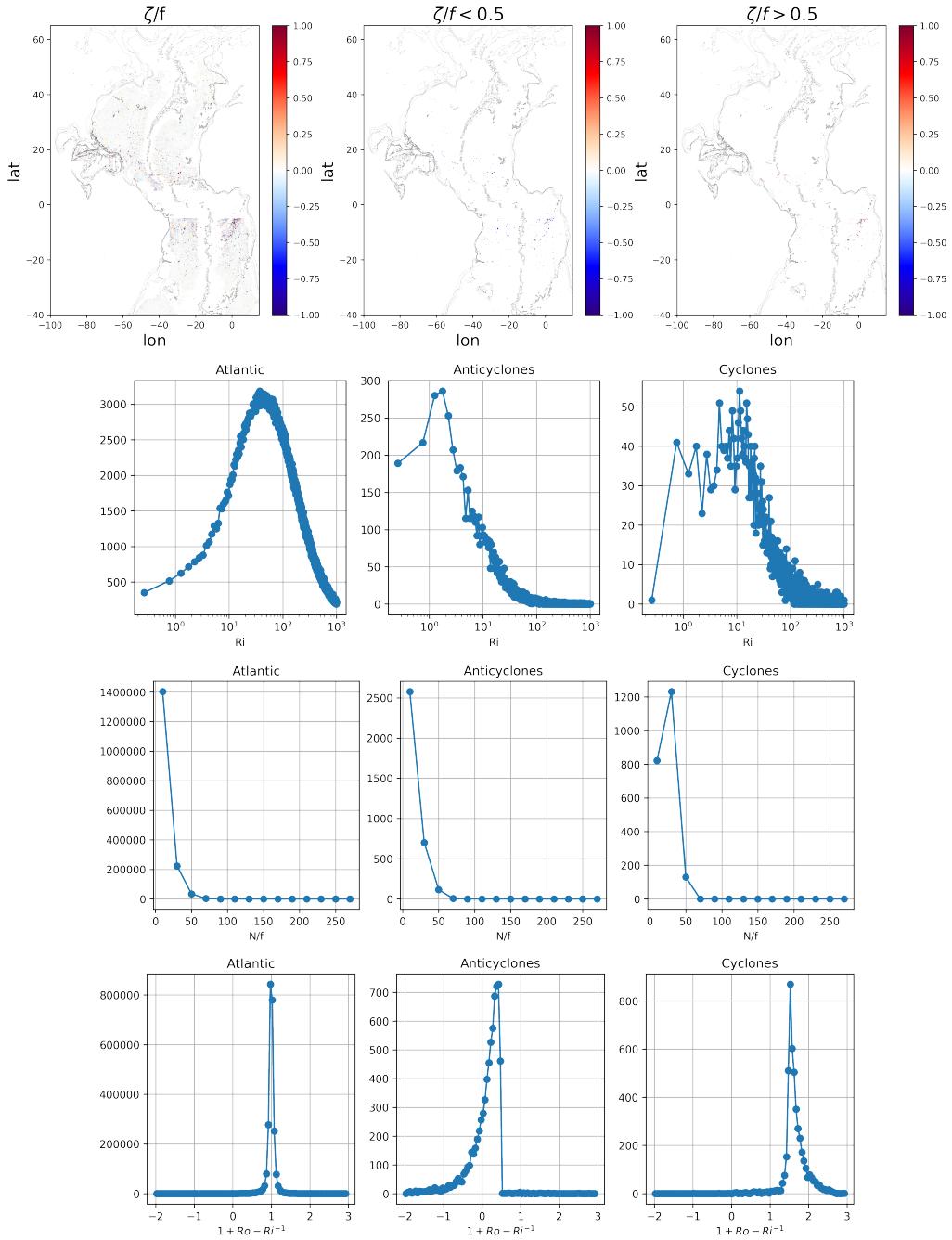


Figure 11: Statistics at 4000 m depth