Lecture 6.

Absorption by atmospheric gases in the IR, visible and UV spectral regions.

Objectives:

- 1. Gaseous absorption in thermal IR.
- 2. Gaseous absorption in the visible and near infrared.
- 3. Gaseous absorption in UV.
- 4. Spectroscopic databases: HITRAN

Required reading:

L02: 3.2, 4.2.1

Additional reading:

Rothman L.S. et al., The HITRAN 2008 molecular spectroscopic database. Journal of Quantitative Spectroscopy and Radiative Transfer, V100, 533-572, 2009.

http://cfa-www.harvard.edu/hitran//

High resolution spectral modeling http://spectralcalc.com/

1. Gaseous absorption in the thermal IR.

Main atmospheric gases absorbing/emitting in the IR: CO₂, H₂O, O₃, CH₄, N₂O, CFCs.

- ✓ Each atmospheric gas has a specific absorption/emission spectrum its own spectral signature.
- ✓ Position of absorption line centers differs for isotopes of the same molecule.

Table 6.1 Main vibrational transition of water vapor isotopes.

| Gas | v ₁ , cm ⁻¹ | v ₂ , cm ⁻¹ | v ₃ , cm ⁻¹ | |
|-------------|-----------------------------------|-----------------------------------|-----------------------------------|--|
| $H_2^{16}O$ | 3657.05 | 1594.75 | 3755.93 | |
| $H_2^{17}O$ | 3653.15 | 1591.32 | 3748.32 | |
| $H_2^{18}O$ | 3649.69 | 1588.26 | 3741.57 | |
| $HD^{16}O$ | 2723.68 | 1403.48 | 3707.47 | |
| $D_2^{16}O$ | 2669.40 | 1178.38 | 2787.92 | |

- ✓ Continuum absorption by water vapor is defined as any observed absorption by water vapor not attributable to the Lorentz line contribution within 25 cm⁻¹ of each line. It has been suggested that it results from the accumulated absorption of the distant wings of lines in the far infrared. This absorption is caused by collision broadening between H₂O molecules (called **self-broadening**) and between H₂O and non-absorbing molecules (N₂) (called **foreign broadening**). The most recent work suggested that the large portion of the continuum might be due to collision-induced transitions and does not relate to the line wings.
- ✓ Atmospheric pressure strongly affects the absorption spectra of gases (through pressure broadening). This poses a major problem in computing the transfer of IR radiation through the atmosphere with varying pressure, temperature, and amount of gases.

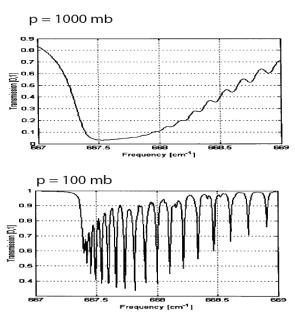


Figure 6.1 Example of *high spectral* resolution transmission spectra of a one-meter path with typical CO2 concentration at 1000 mb and 100 mb.

Table 6.2 The most important vibrational and rotational transitions for H_20 , CO_2 , O_3 , CH_4 , N_2O , and CFCs.

| Gas | Center | Transition | Band interval |
|------------------|--------------------------------|---|---------------------|
| | $v (cm^{-1}) (\lambda(\mu m))$ | | (cm ⁻¹) |
| H ₂ O | - | pure rotational | 0-1000 |
| | 1594.8 (6.3) | ν ₂ ; P, R | 640-2800 |
| | continuum* | far wings of the strong | 200-1200 |
| | | lines; water vapor | |
| | | dimmers (H ₂ O) ₂ | |
| CO ₂ | 667 (15) | ν ₂ ; P, R, Q | 540-800 |
| | 961 (10.4) | overtone and | 850-1250 |
| | 1063.8 (9.4) | combination | |
| | 2349 (4.3) | ν ₃ ; P, R | 2100-2400 |
| | | overtone and | |
| | | combination | |
| O ₃ | 1110 (9.01) | ν ₁ ; P, R | 950-1200 |
| | 1043 (9.59) | ν ₃ ; P, R | 600-800 |
| | 705 (14.2) | ν ₂ ; P, R | 600-800 |
| CH ₄ | 1306.2 (7.6) | V_4 | 950-1650 |
| N ₂ O | 1285.6 (7.9) | ν_1 | 1200-1350 |
| | 588.8 (17.0) | v_2 | 520-660 |
| | 2223.5 (4.5) | ν ₃ | 2120-2270 |
| CFCs | | | 700-1300 |

NOTE: Chlorofluorocarbons (**CFCs**) are a **family** of chemical compounds.

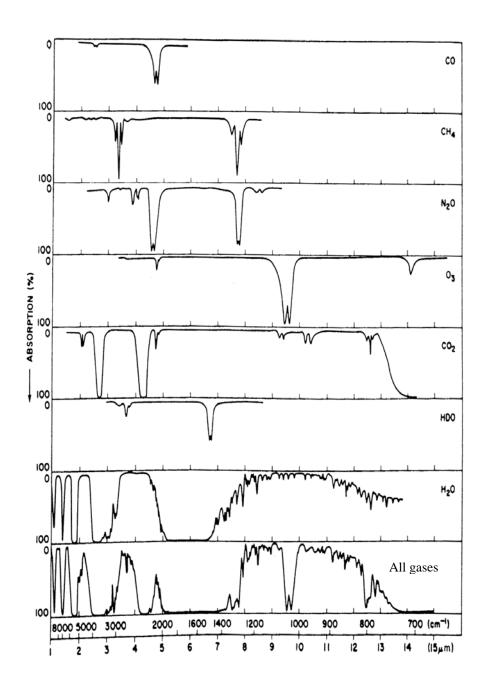


Figure 6.2 *Low-resolution* infrared absorption spectra of the major atmospheric gases. (compare to Figure 6.3 that shows transmission with higher spectral resolution)

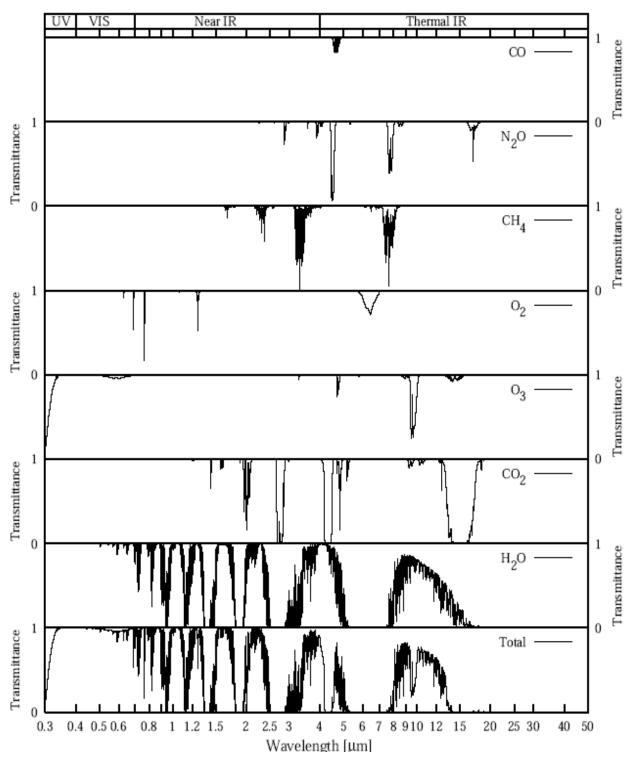


Figure 6.3 Transmission spectra of the major atmospheric gases.

2. Gaseous absorption in the visible and near-IR.

✓ Absorption of visible and near IR radiation in the gaseous atmosphere is primarily due to H_2O , O_3 , and CO_2 .

Table 6.3 Main Visible and near-IR absorption bands of atmospheric gases

| Gas | Center | Band interval |
|------------------|--------------------------------|---------------------|
| | $v (cm^{-1}) (\lambda(\mu m))$ | (cm ⁻¹) |
| H ₂ O | 3703 (2.7) | 2500-4500 |
| | 5348 (1.87) | 4800-6200 |
| | 7246 (1.38) | 6400-7600 |
| | 9090 (1.1) | 8200-9400 |
| | 10638 (0.94) | 10100-11300 |
| | 12195 (0.82) | 11700-12700 |
| | 13888 (0.72) | 13400-14600 |
| | visible | 15000-22600 |
| CO_2 | 2526 (4.3) | 2000-2400 |
| | 3703 (2.7) | 3400-3850 |
| | 5000 (2.0) | 4700-5200 |
| | 6250 (1.6) | 6100-6450 |
| | 7143 (1.4) | 6850-7000 |
| O_3 | 2110 (4.74) | 2000-2300 |
| | 3030 (3.3) | 3000-3100 |
| | visible | 10600-22600 |
| O_2 | 6329 (1.58) | 6300-6350 |
| | 7874 (1.27) | 7700-8050 |
| | 9433 (1.06) | 9350-9400 |
| | 13158 (0.76) | 12850-13200 |
| | 14493 (0.69) | 14300-14600 |
| | 15873 (0.63) | 14750-15900 |
| N_2O | 2222 (4.5) | 2100-2300 |
| | 2463 (4.06) | 2100-2800 |
| | 3484 (2.87) | 3300-3500 |
| $\mathbf{CH_4}$ | 3030 (3.3) | 2500-3200 |
| | 4420 (2.20) | 4000-4600 |
| | 6005 (1.66) | 5850-6100 |
| CO | 2141 (4.67) | 2000-2300 |
| | 4273 (2.34) | 4150-4350 |
| NO_2 | visible | 14400-50000 |

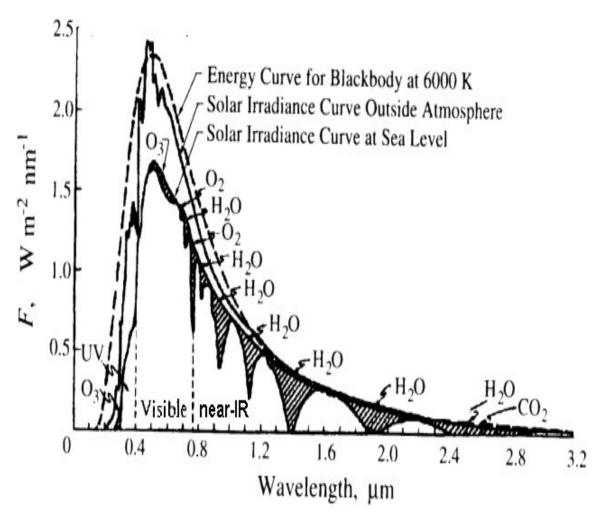


Figure 6.4 Solar spectral irradiance (flux) at the top of the atmosphere and at the surface.

NOTE: Atmospheric gases absorb only a small fraction of visible radiation.

3. Gaseous absorption in the UV.

Table 6.4 Wavelengths of absorption in the solar spectrum (UV + visible) by several atmospheric gases

| Gas | Absorption wavelengths (µm) | |
|---|-----------------------------|--|
| N_2 | < 0.1 | |
| O_2 | < 0.245 | |
| O ₃ | 0.17-0.35 | |
| | 0.45-0.75 | |
| H ₂ O | < 0.21 | |
| | 0.6-0.72 | |
| H ₂ O ₂ hydrogen peroxide | < 0.35 | |
| NO ₂ nitrogen oxide | < 0.6* | |
| N ₂ O | < 0.24 | |
| NO ₃ nitrate radical | 0.41-0.67 | |
| HONO nitrous acid | < 0.4 | |
| HNO ₃ nitric acid | < 0.33 | |
| CH ₃ Br methyl bromide | < 0.26 | |
| CFCl ₃ (CFC11) | < 0.23 | |
| HCHO formaldehyde | 0.25-0.36 | |

^{*} NO_2 absorb at λ < 0.6 μ m, but photodissociate at λ < 0.4 μ m

NOTE: To avoid very complicated calculations of electronic transitions, numerous measurements of the **absorption cross-sections** of the atmospheric atoms and molecules absorbing in the UV and visible have been performed in laboratory experiments. In general, the absorption cross section varies with temperature.

✓ Absorption of UV radiation in the gaseous atmosphere is primarily due molecular oxygen O₂ and ozone O₃.

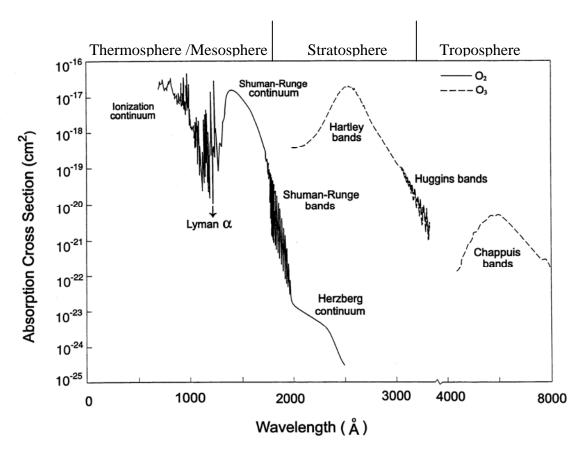


Figure 6.5 Spectral absorption cross-sections of O2 and O3

NOTE:

- a) Bands of O_2 and O_3 at wavelengths < 1 μm are electronic transitions.
- b) These absorption bands are relatively uncomplicated continua because practically all absorption results in dissociation of the molecule (so the upper state is not quantized);
- c) Despite the small amount of O_3 , no solar radiation penetrates to the lower atmosphere at wavelengths < 310 nm (because of large absorption cross-sections of O_3);

4. Spectroscopic database HITRAN

(HIgh-resolution TRANsmission molecular absorption database)

http://cfa-www.harvard.edu/hitran//

- ✓ The database is a long-running project started by the Air Force Cambridge Research Laboratories (AFCRL) in the late 1960's in response to the need for detailed knowledge of the infrared properties of the atmosphere.
- ✓ Free to users, need to order and install a free software package HAWKS (HITRAN Atmospheric Workstation)
- ✓ The current HITRAN'2008 Database (Version 13.0) contains 2,713,968 spectral lines for 39 different molecules.
- ✓ In addition to the line HITRAN2008 database, there are directories containing files of aerosol indices of refraction, UV line-by-line and absorption cross-section parameters, and more extensive IR absorption cross-sections

HITRAN Tree structure

