

# Turbulent seismoacoustic signals from a hurricane landfall

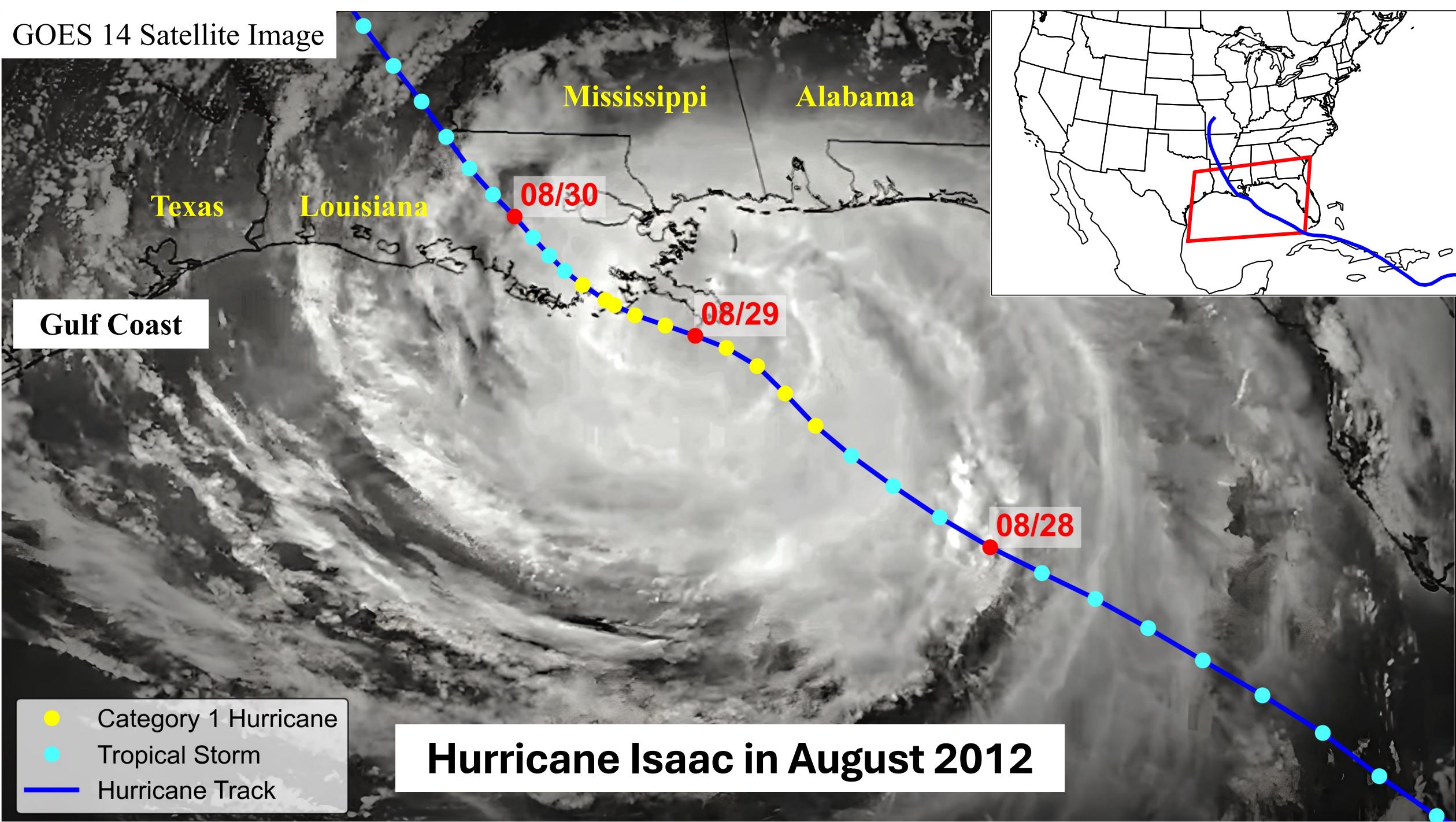
Qing Ji<sup>1</sup>    Ipsita Dey<sup>2,3</sup>    Eric Dunham<sup>1</sup>

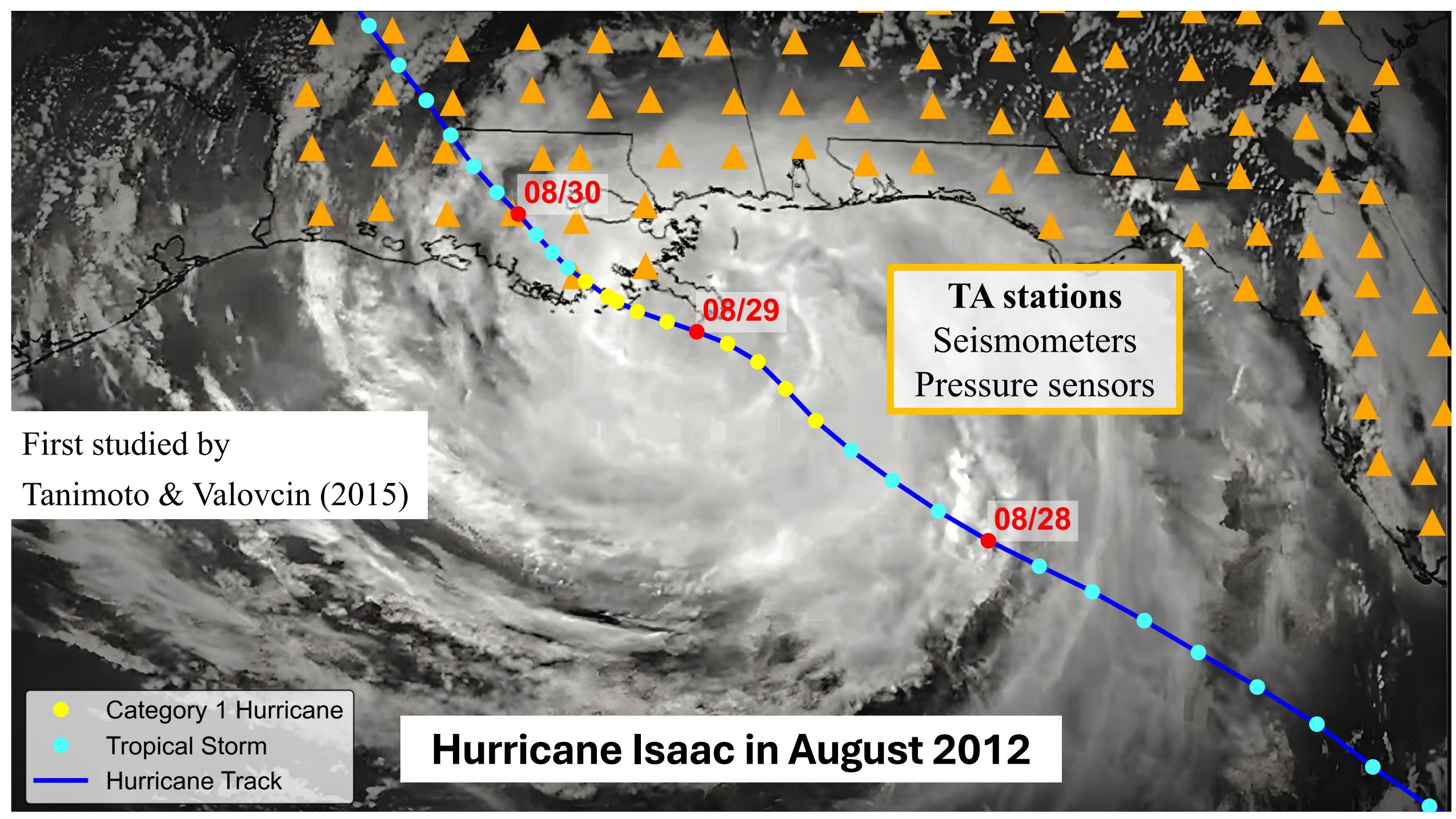
1 Department of Geophysics, Stanford University

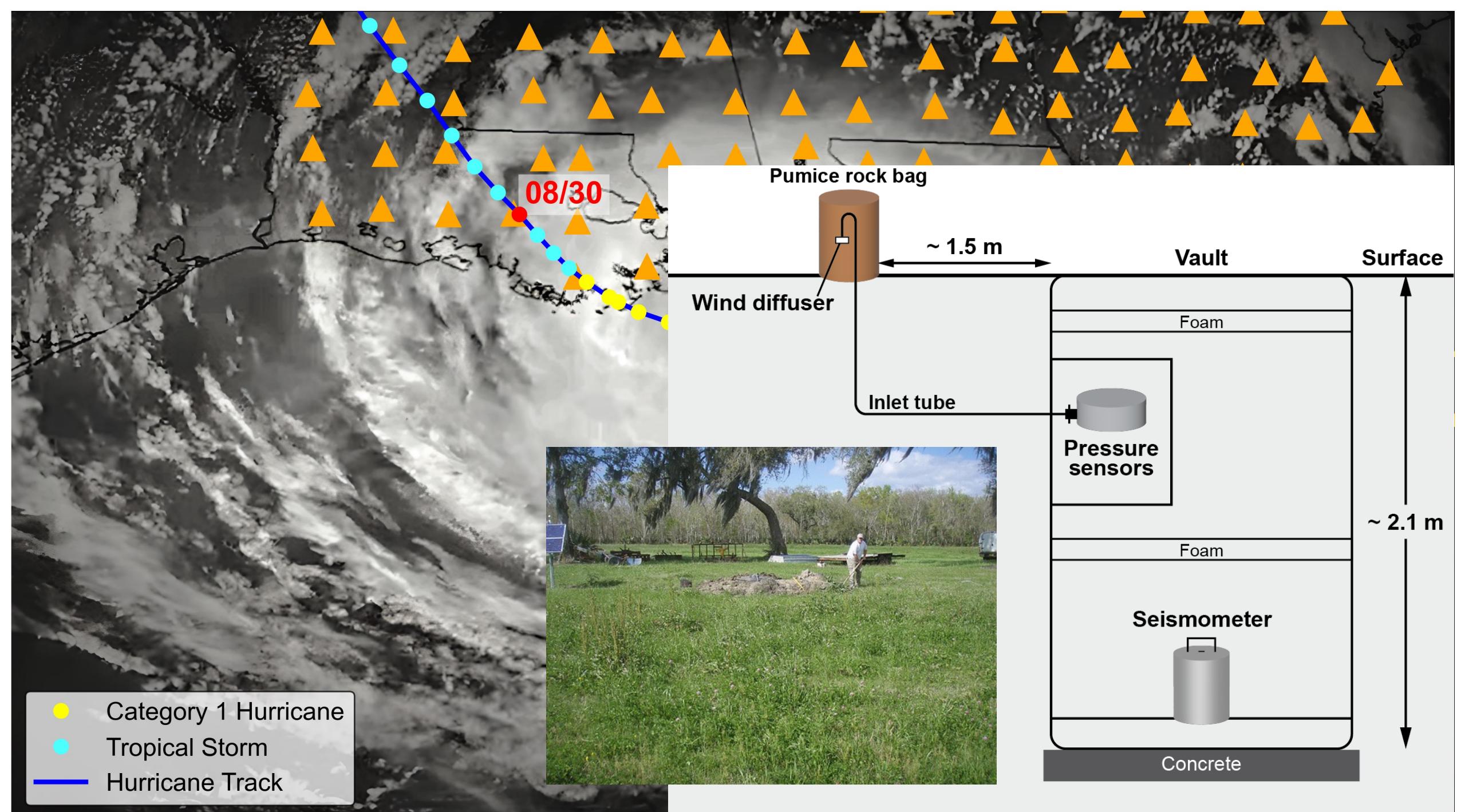
2 Department of Earth System Science, Stanford University

3 National Centre for Atmospheric Science and Department of Meteorology, University of Reading, UK

GOES 14 Satellite Image

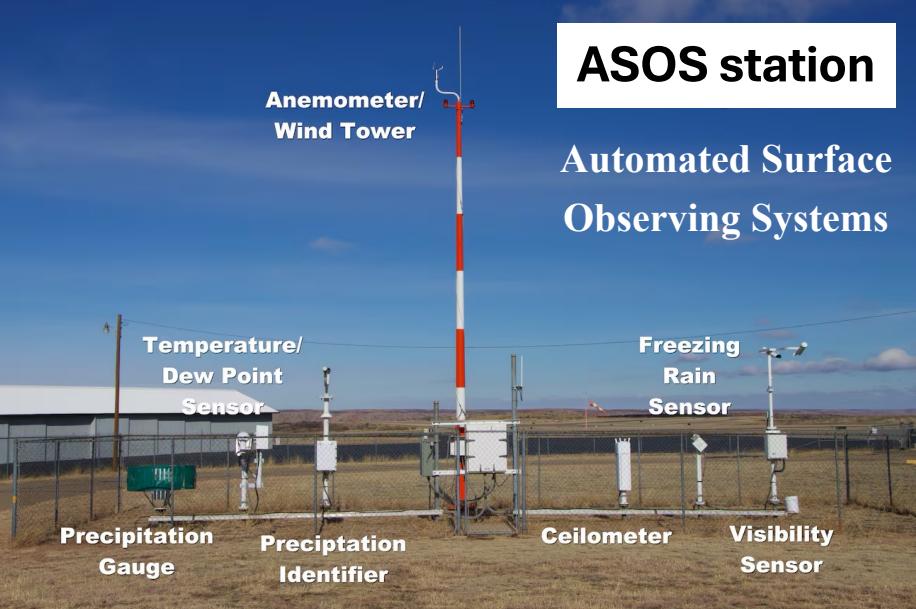




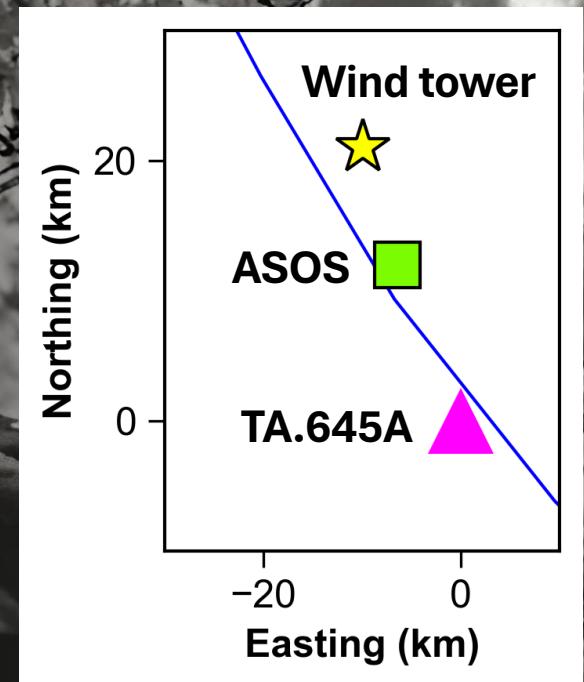


## ASOS station

Automated Surface  
Observing Systems



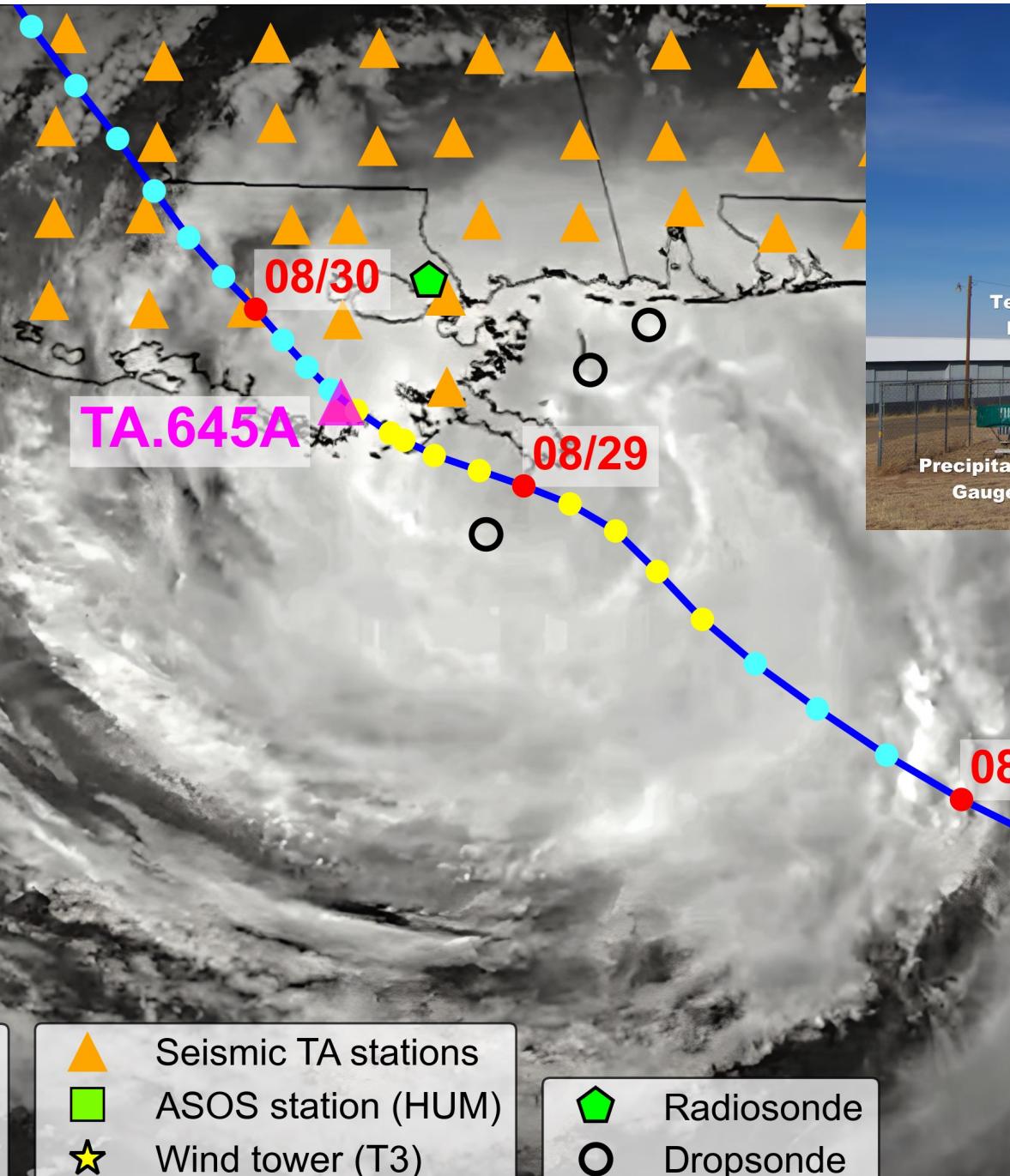
## Portable wind tower



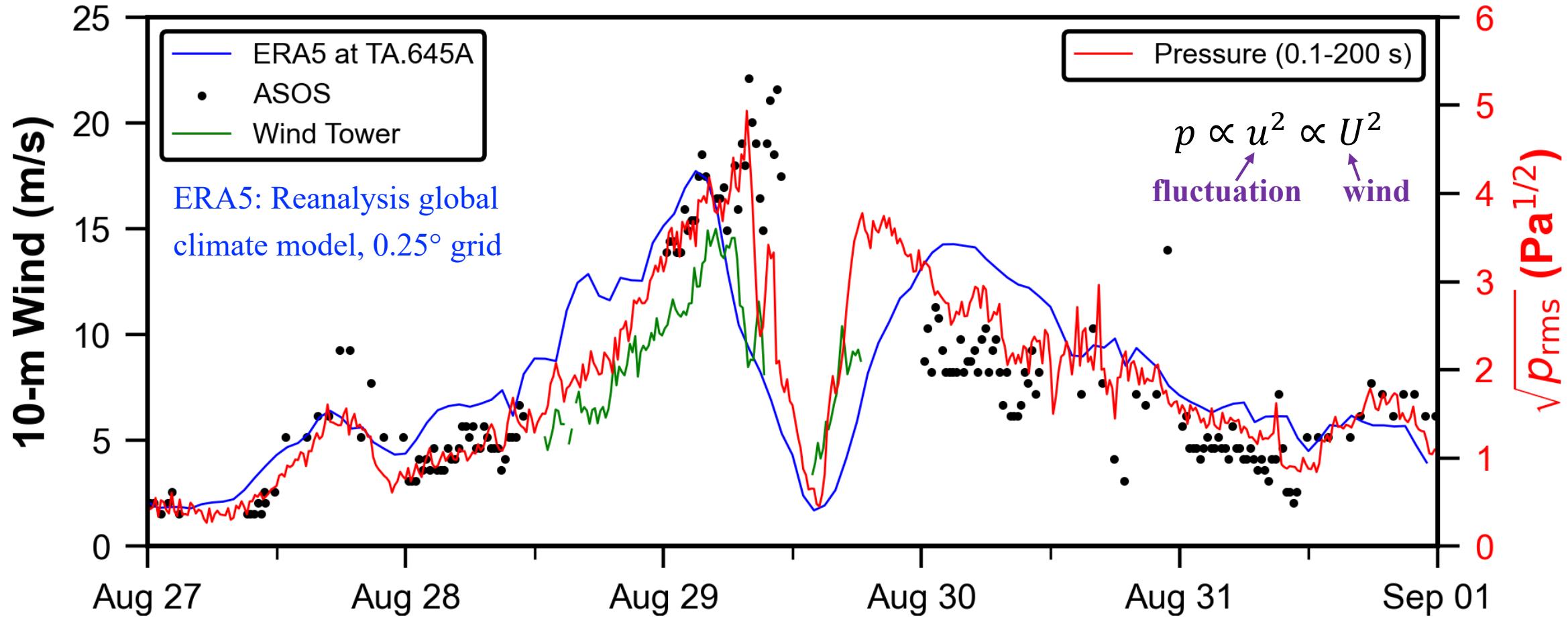
- Category 1 Hurricane
- Tropical Storm
- Hurricane Track

- ▲ Seismic TA stations
- ASOS station (HUM)
- ★ Wind tower (T3)

- ◆ Radiosonde
- Dropsonde

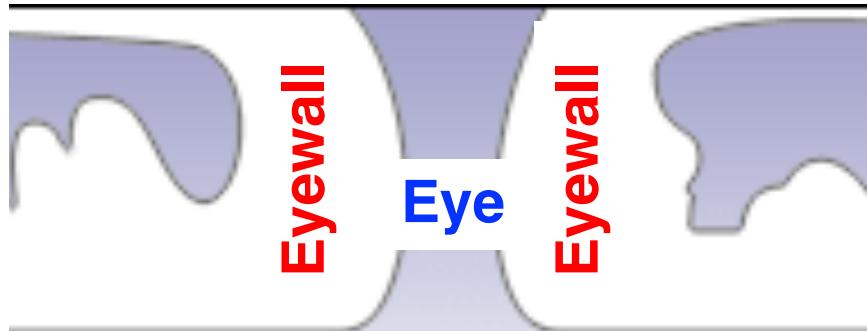


# Infrasound data for hurricane study



ASOS: Real-time monitoring of the environment sampling quantities  
(e.g., wind speed, turbulence analysis, ...)

# As hurricane passes the station .....



TA Station

L: Long period channel

D: Pressure

LDO

Barometric Pressure

08/29

Low Pressure Center  
(968 mb)

08/30

(mbar)

LDF

Infrasound

Eyewall Eye Eyewall

Period: 2 – 20 s

60  
980  
1000

0  
-60

LHZ

Microseism

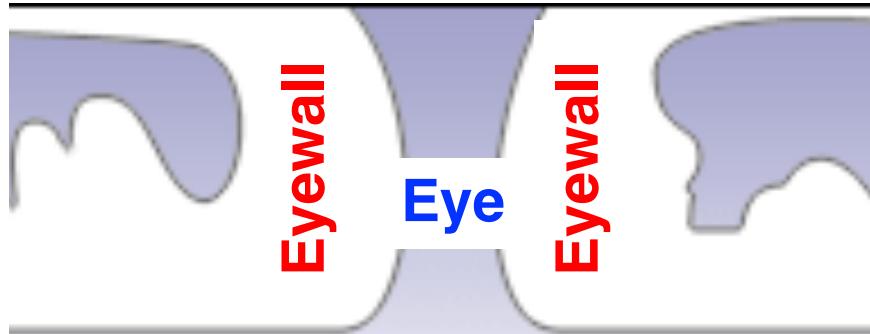
Teleseismic EQ

Teleseismic EQ

30  
0  
-30

( $\mu\text{m}$ )

# As hurricane passes the station .....



TA Station

L: Long period (1 Hz sampling)

D: Pressure

LDO

Barometric Pressure

08/29

08/30

Low Pressure Center  
(968 mb)

(mbar)

1000  
980

LDF

Infrasound

Period: 20 – 100 s

40  
0

-40

LHZ

Vertical Displ.

Eyewall Eye Eyewall

15  
0

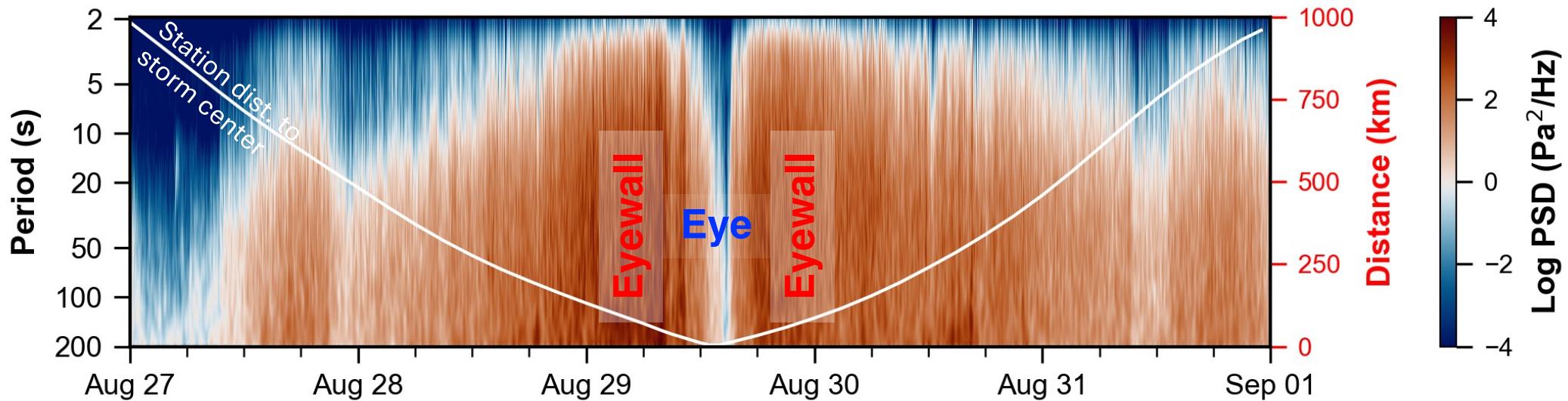
-15

Teleseismic EQ

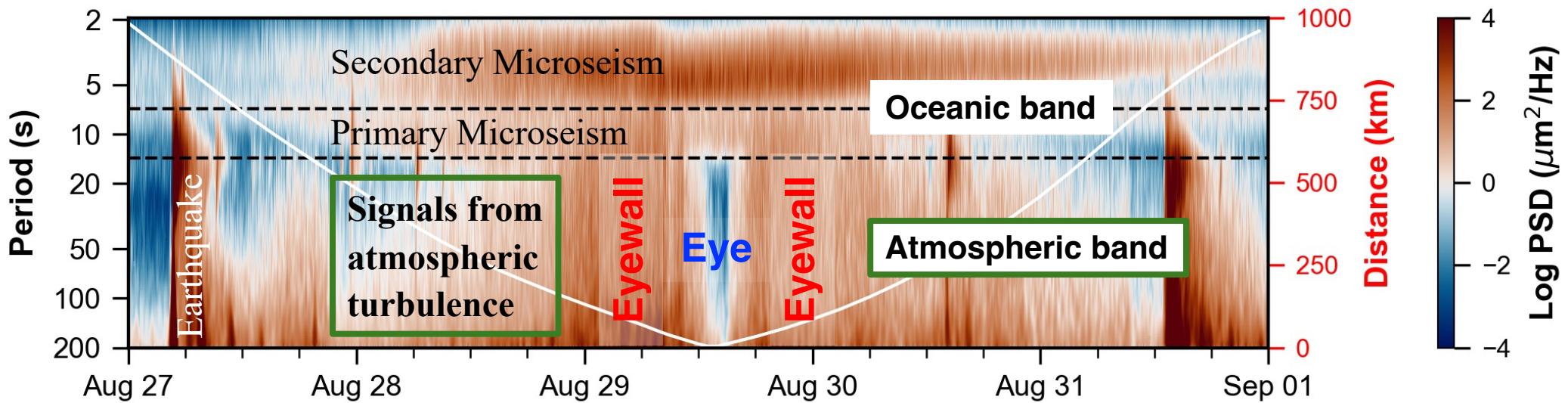
Teleseismic EQ

# Wavelet spectrograms of infrasound & seismic data

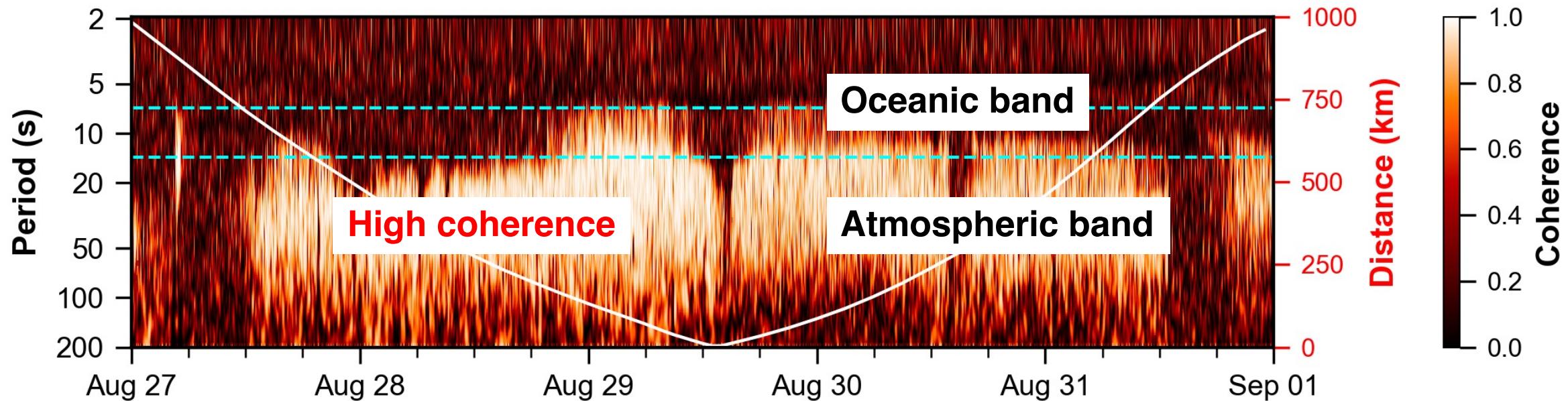
**Surface Pressure**



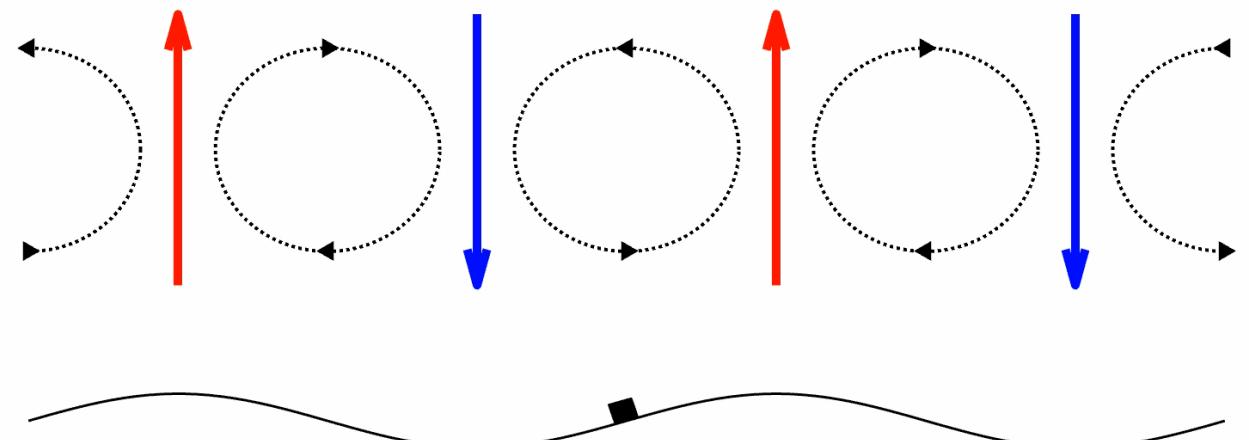
**Seismic Vertical Displacement**



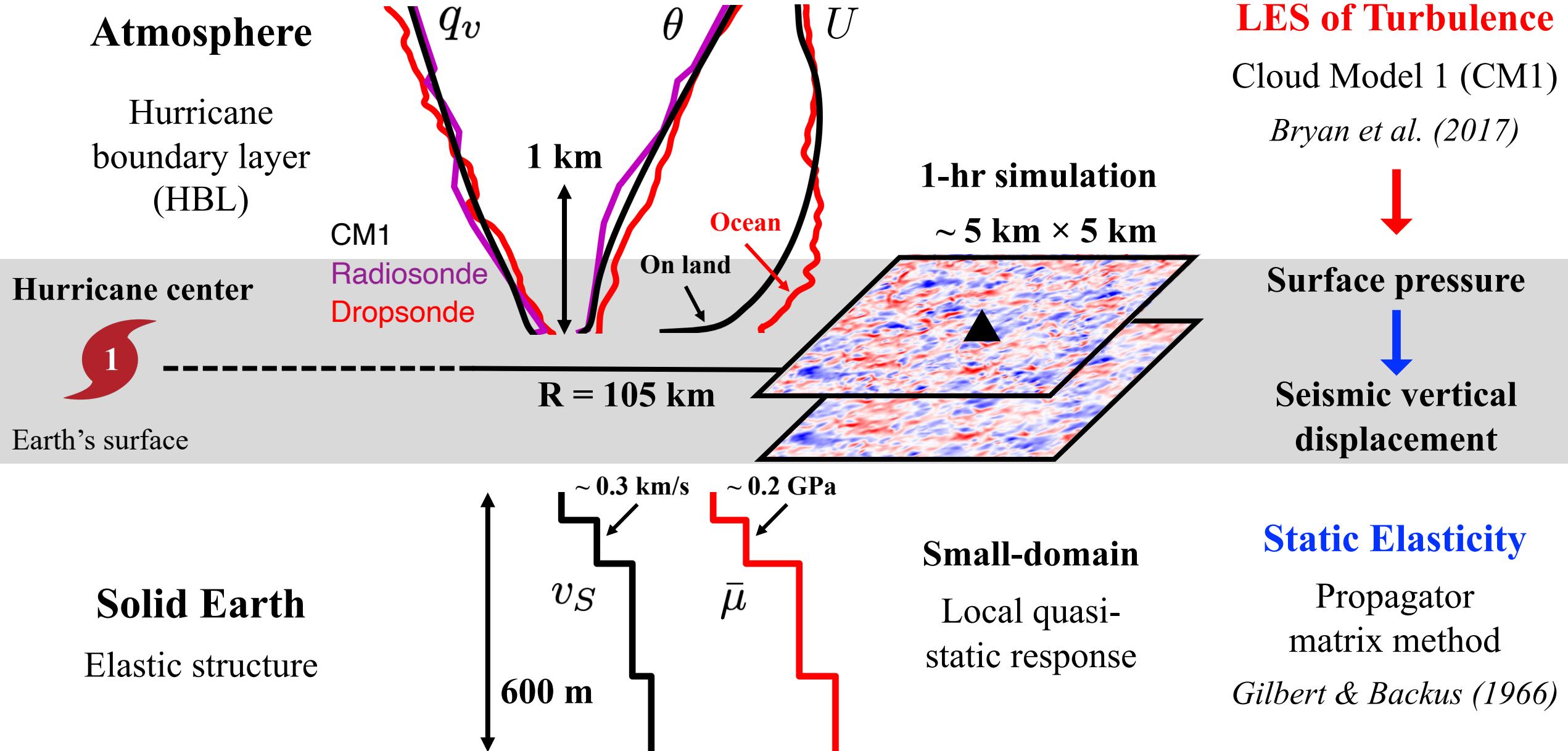
# High coherence indicates local quasi-static response



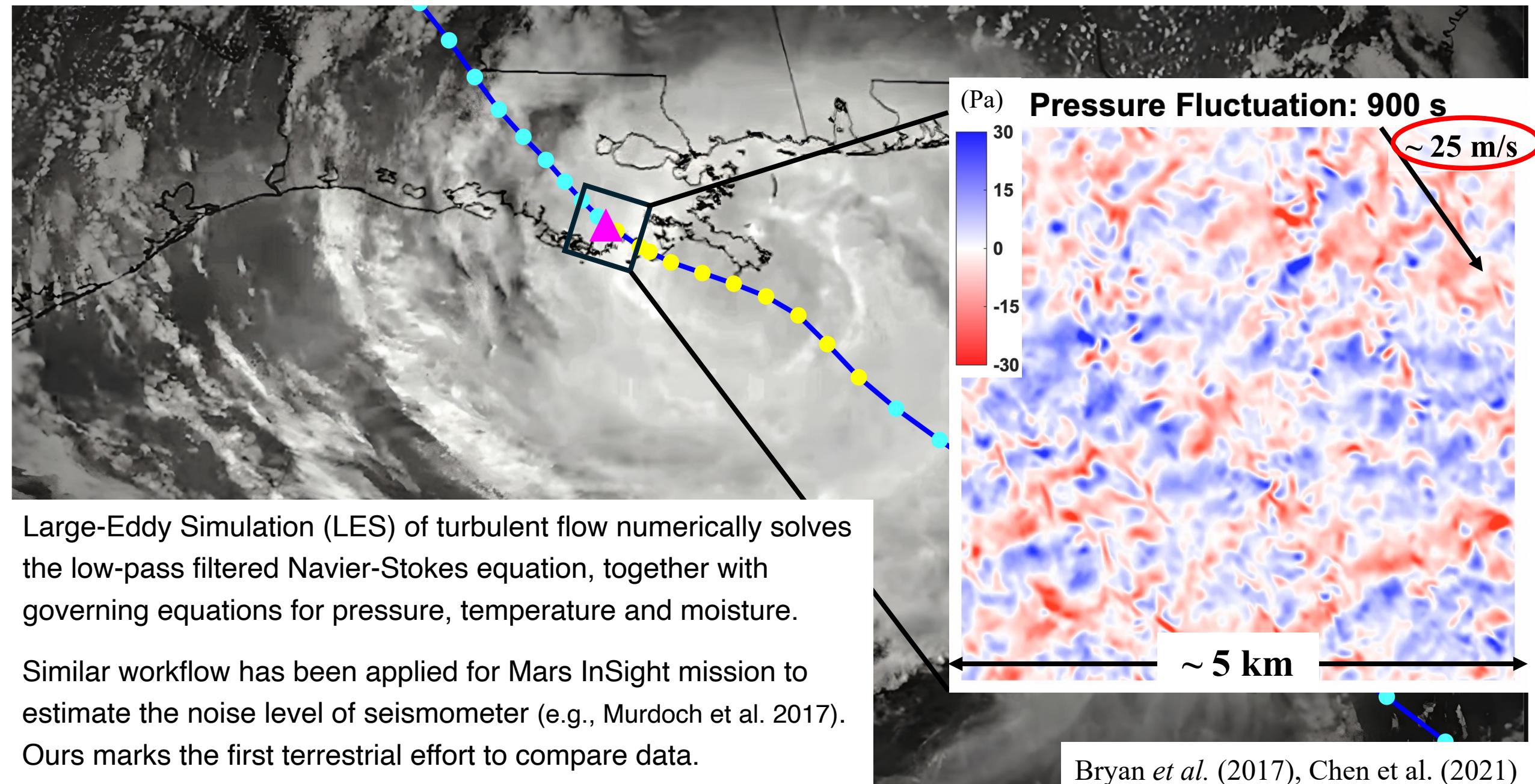
Simplified illustration  
Sorrells (1971) theory  
Pressure wave model



# Interdisciplinary modeling



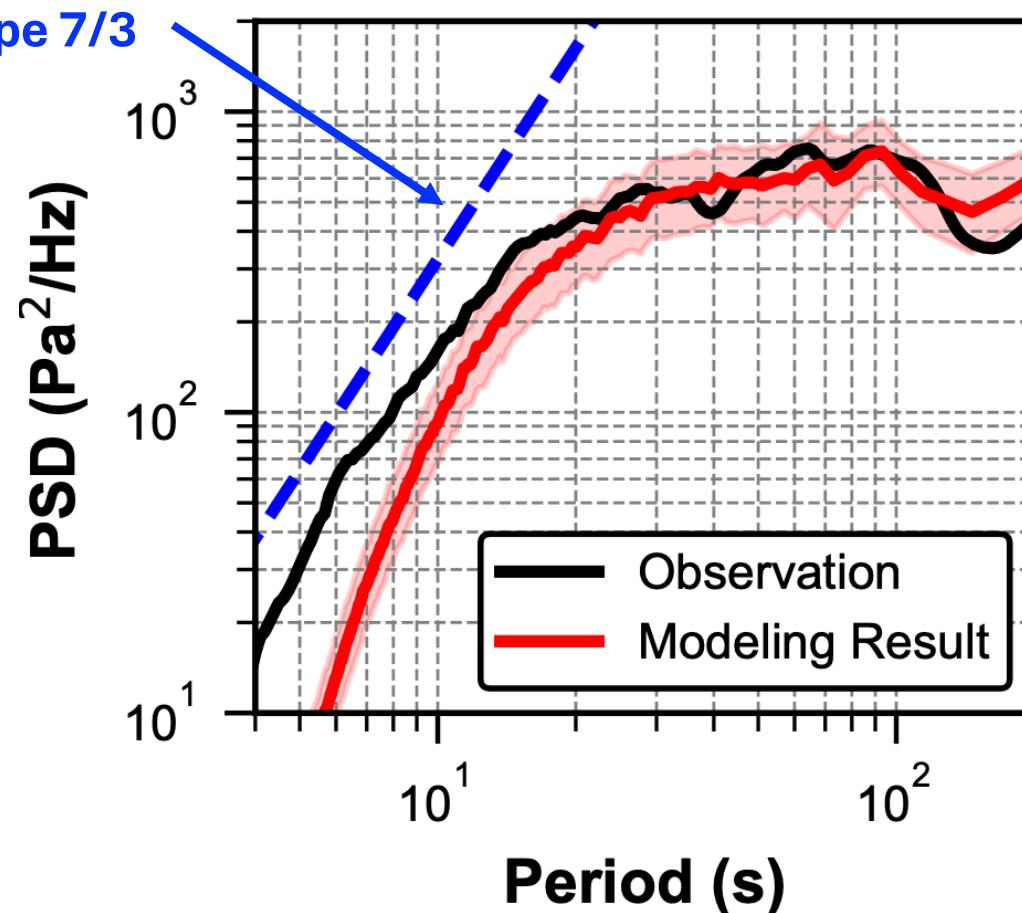
# LES of Hurricane Boundary Layer (HBL) on land



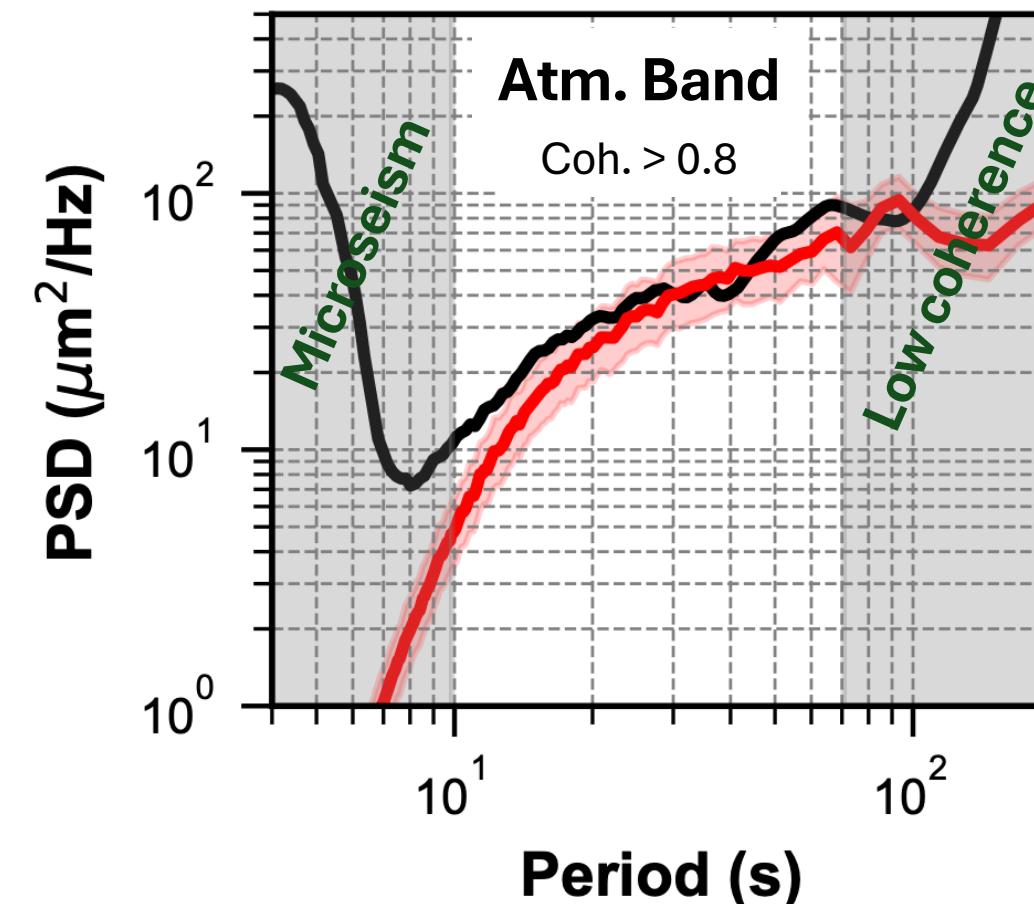
# Turbulent origin of signals

Inertial subrange

## Surface Pressure



## Vertical Seismic Displ.

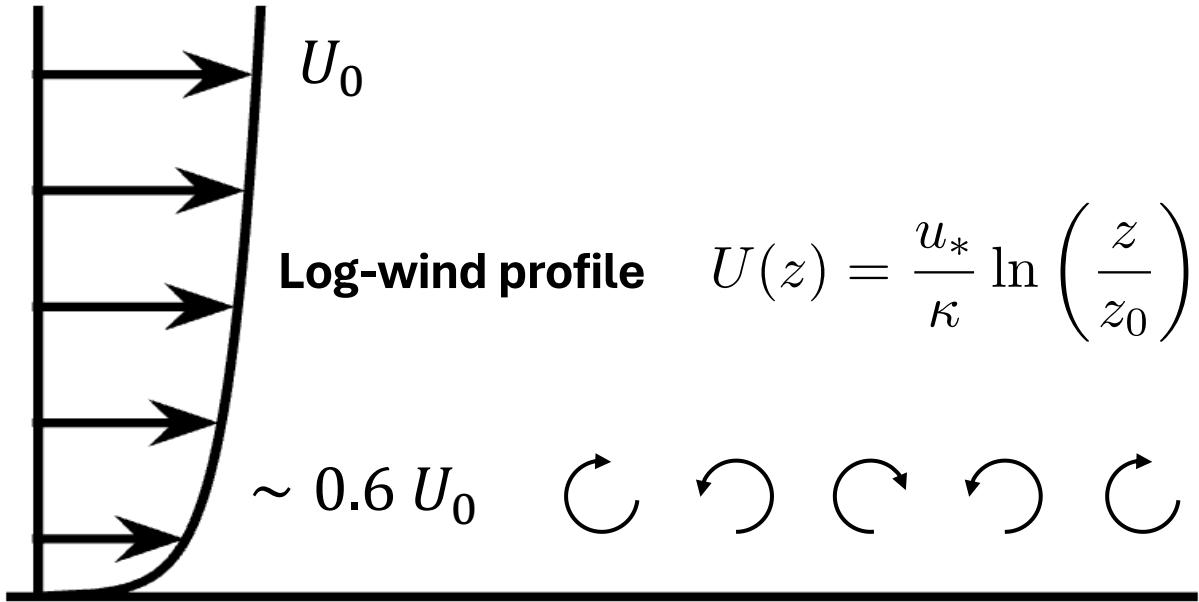


1. Insights into Sorrells theory

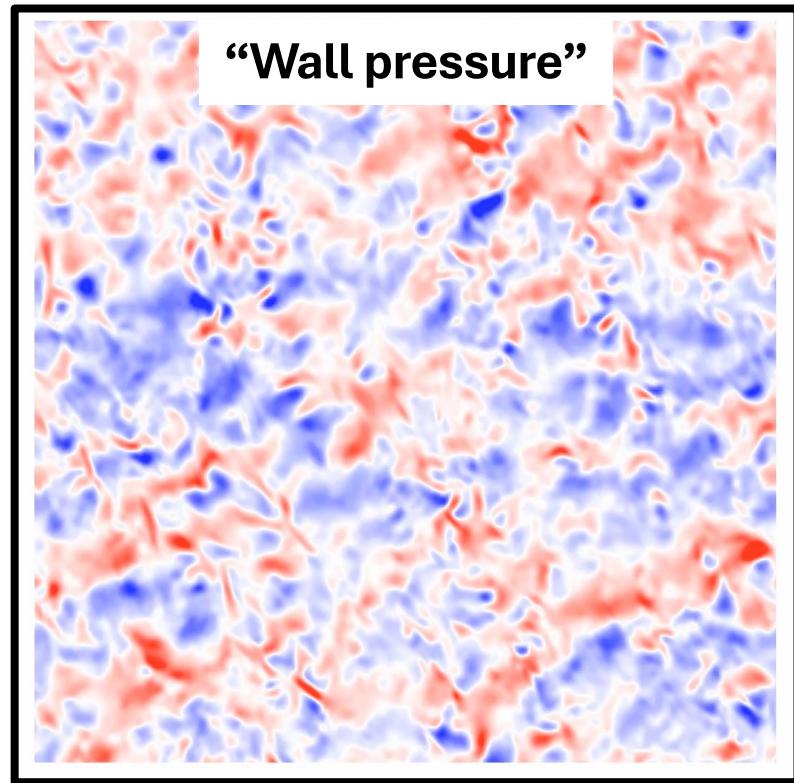
2. Infrasound data for turbulence analysis

# Convection velocity of a turbulent field

Turbulent boundary layer



$$\text{Log-wind profile} \quad U(z) = \frac{u_*}{\kappa} \ln \left( \frac{z}{z_0} \right)$$



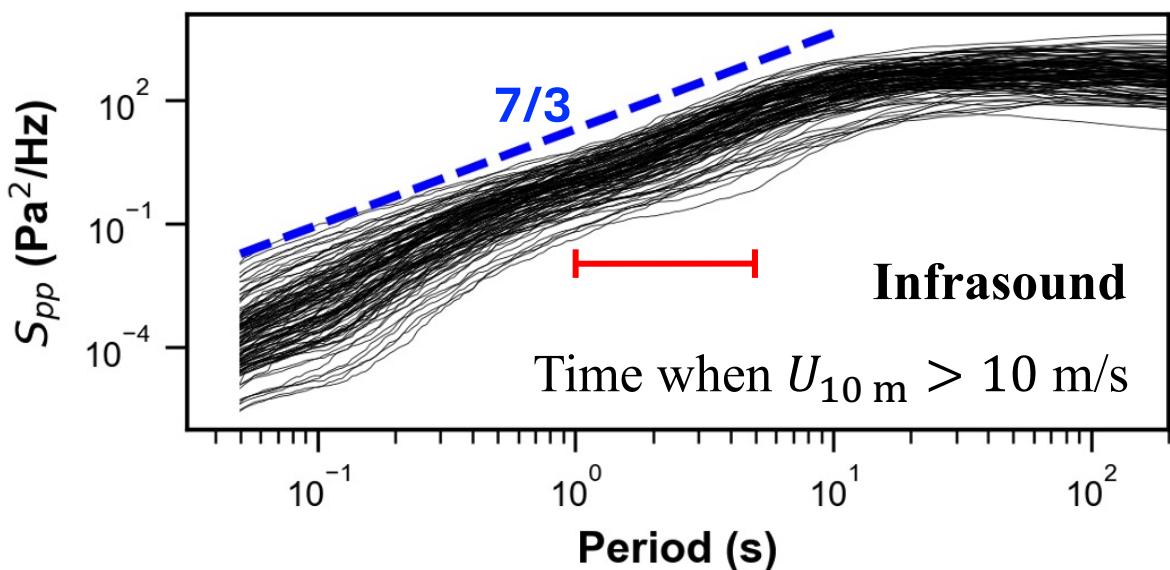
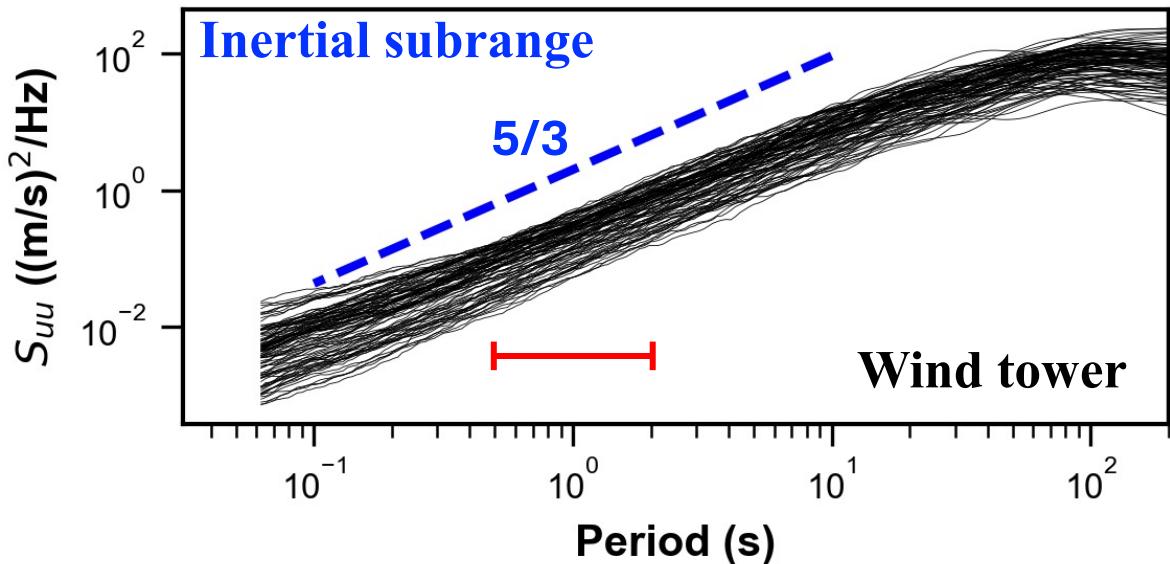
Velocity is zero at the wall, but wall stress / pressure advects with  $U_c$  (convection velocity)

"Shadow" of eddies at height where wind speed  $U = U_c \approx 0.6 U_0$

Sorrells theory: "Pressure wave speed"  $c \approx U_c$

Similar in Mars application, where  $c \approx 0.5 U_0$  work best for Sorrells results (Murdoch et al. 2017)

# Turbulent spectra from wind tower and infrasound



## Streamwise velocity PSD

$$S_{uu}(f) = \alpha_u \left( \frac{\varepsilon U}{2\pi} \right)^{2/3} f^{-5/3}$$

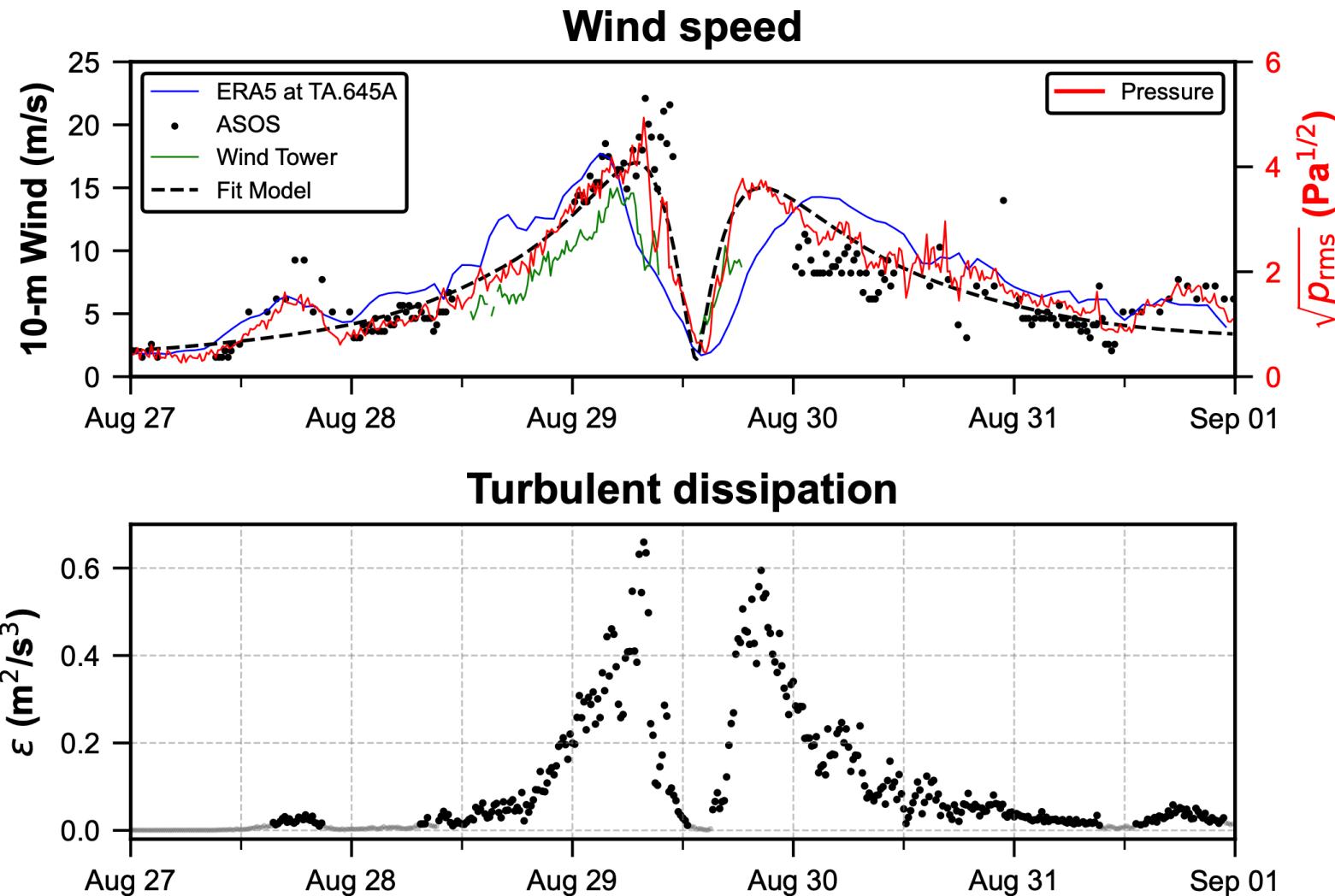
*Dissipation rate  $\varepsilon$  is a key parameter describing the turbulence statistic, and contributes to an important energy source for hurricanes. (Bister & Emanuel, 1998)*

## Pressure PSD

$$\frac{1}{\rho^2} S_{pp}(f) = \alpha_p \left( \frac{\varepsilon U_c}{2\pi} \right)^{4/3} f^{-7/3}$$

Kolmogorov (1941), Obukhov (1949), George et al. (1984) ...

# Continuous monitoring of hurricane landfall



Fluctuation:  $p_{rms} \propto u_{rms}^2$

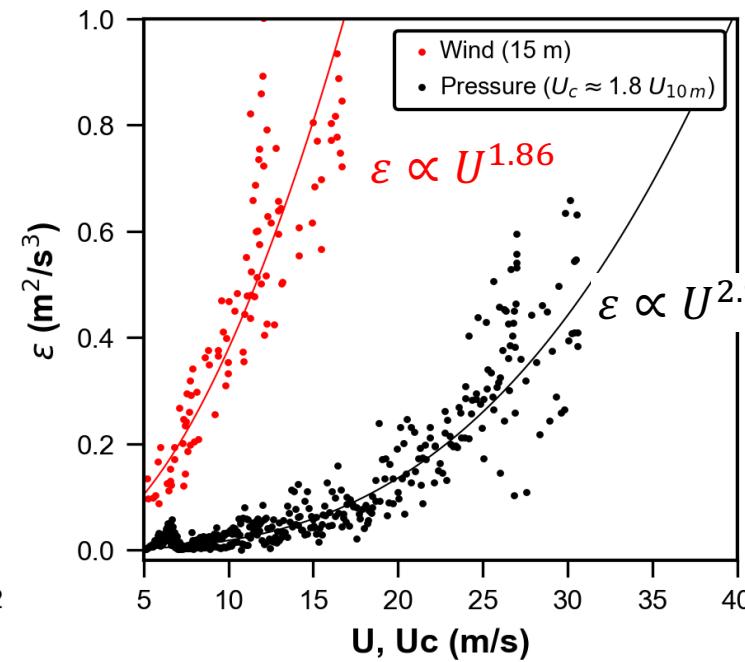
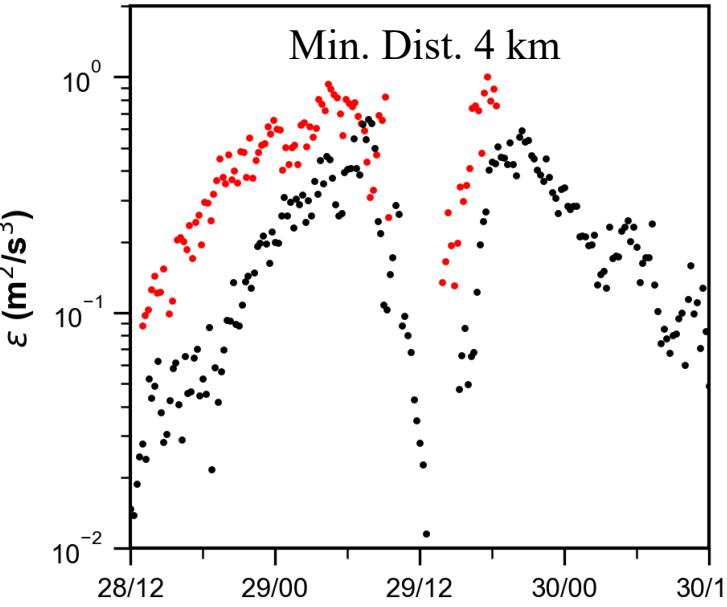
Turbulent intensity:  $u_{rms} \propto U_{10\text{ m}}$

Fit wind model:  $U_{10\text{ m}} = U(r)$

ERA5 reanalysis has a larger radius of maximum wind.

$O(0.1 \text{ m}^2/\text{s}^3)$  or  $O(0.1 \text{ W/kg})$  dissipation rate

Interpreted as  $\epsilon$  for height  $\sim 100$  m where the wind speed equal to  $U_c$

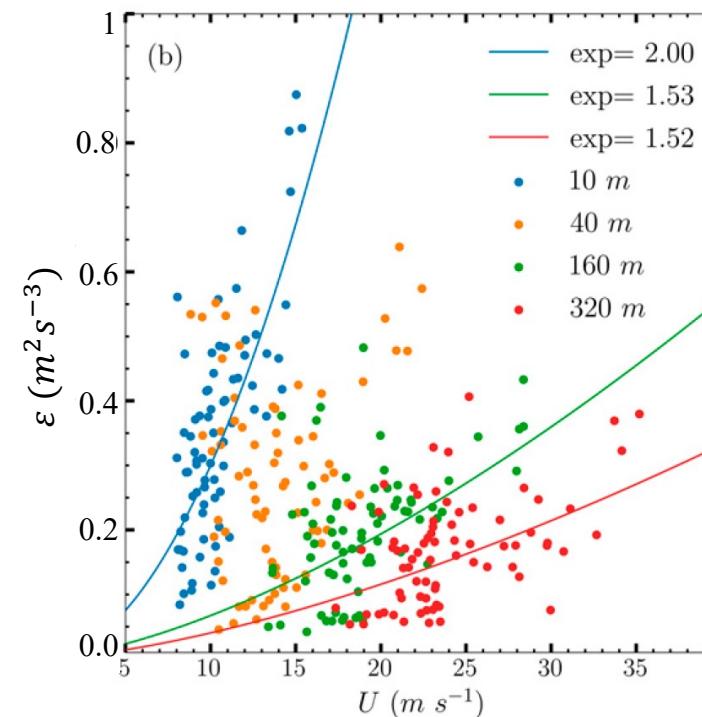
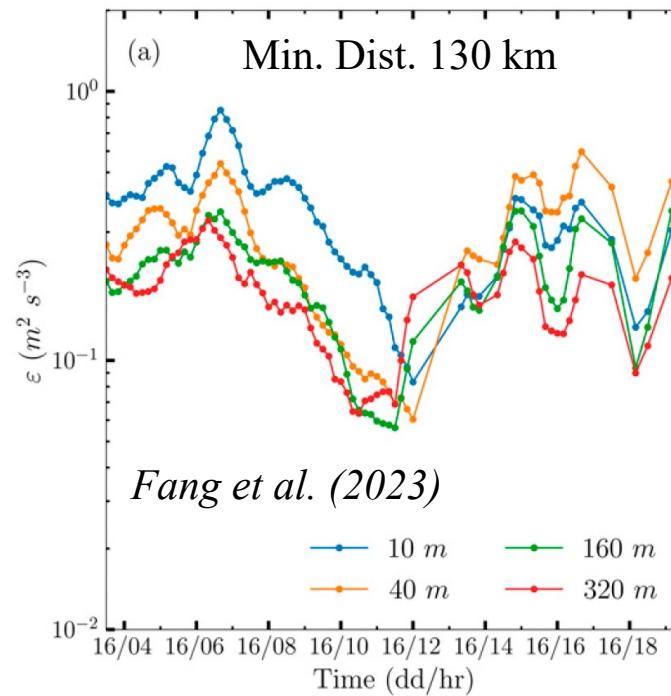


## Turbulent dissipation rate during hurricane landfall

For **pressure** spectra:

Assume  $U_c \approx 1.8 U_{10m}$  from our LES result

Directly use  $U_{10m} \approx 4.2\sqrt{p_{rms}}$  in their units

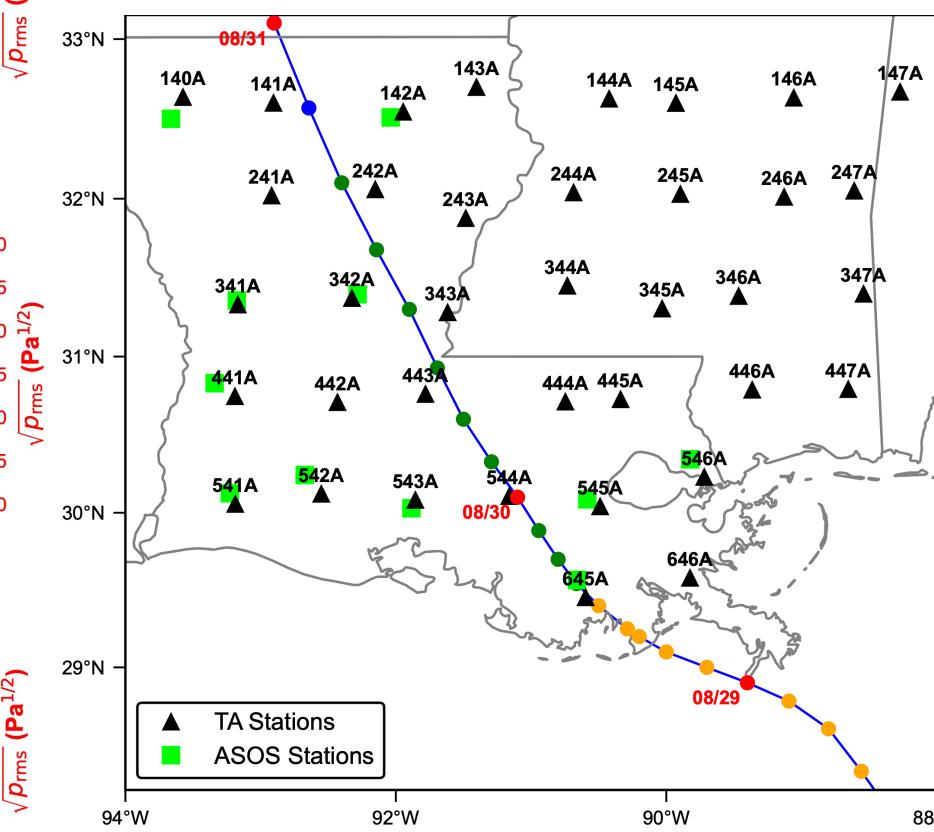
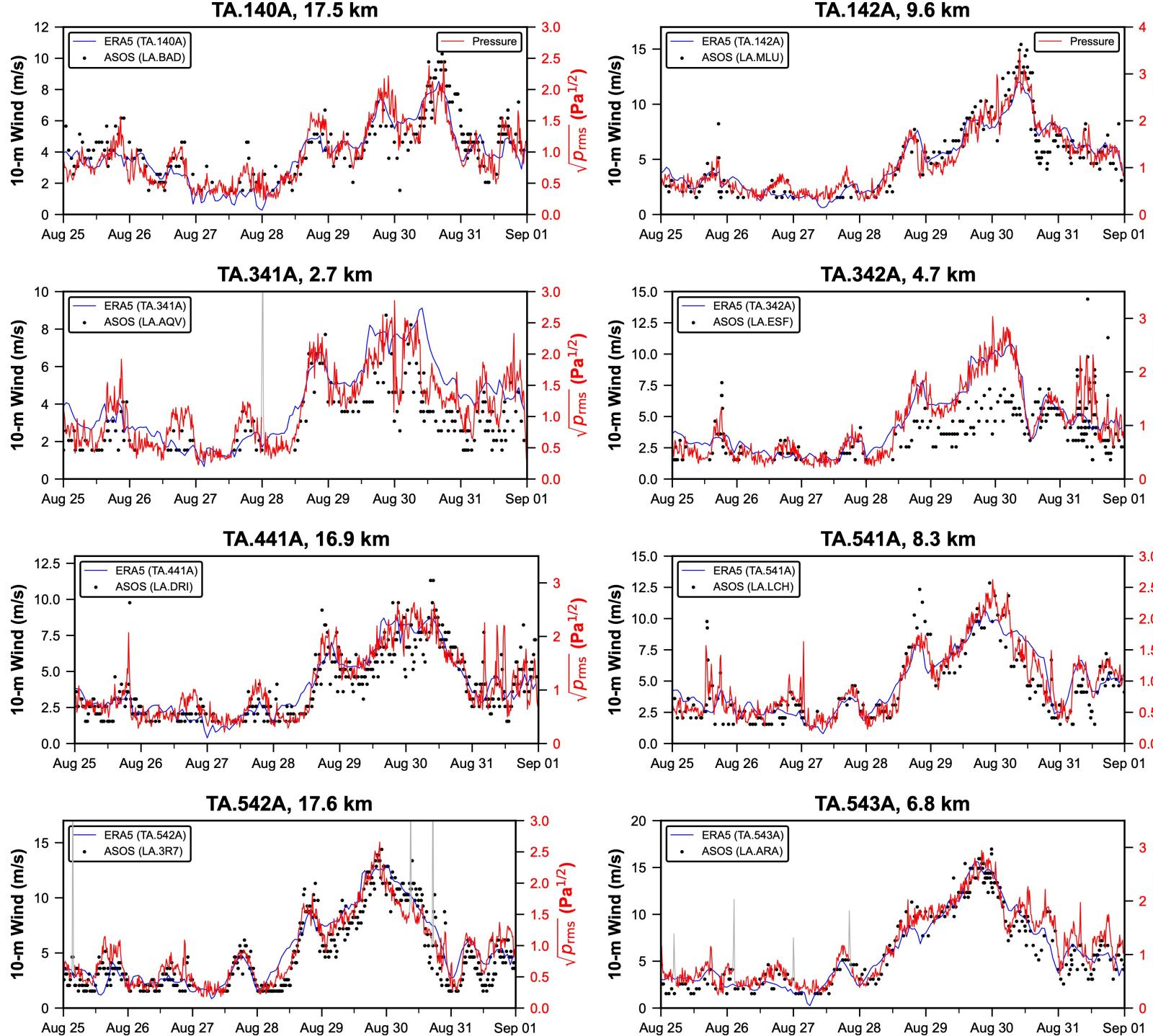


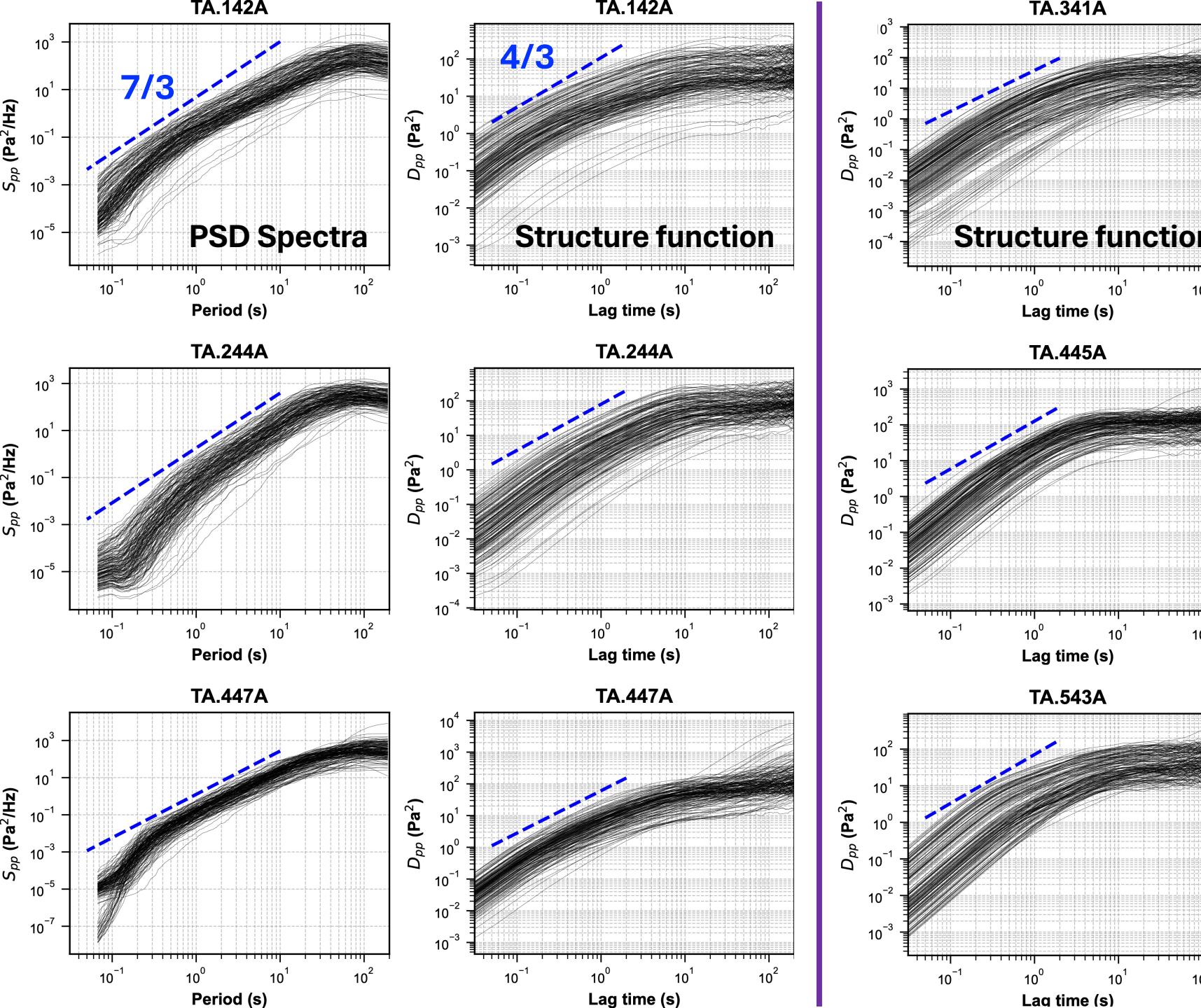
**Wind tower** for 15 m high, while **pressure** imprints from  $\sim 100$  m high (where  $U = U_c$ )

Magnitude of  $\varepsilon$  agrees with other studies on landfalling storms.

Scaling  $\varepsilon \propto U^3$  considered for dominance of shear production, but less certain for storms.

# Infrasound data as proxy of surface wind





## Turbulent pressure

Structure function

$$D_{pp}(\tau) = \overline{[p(t + \tau) - p(t)]^2}$$

$$\frac{1}{\rho^2} D_{pp}(\tau) = C_p (\varepsilon U_c)^{4/3} \tau^{4/3}$$

Deviation from the slope:

- Violation of locally isotropic turbulence assumption
- Interaction of different scales

George et al. (1984), Katul et al. (2025)

**Surprise #1:** Seismic stations record in-situ data of Hurricane Issac after landfall

**Distinct seismic ground motion contributed by ocean and atmosphere**

**Surprise #2:** Turbulence explains the seismoacoustic signatures (in the atm. band)

**Interdisciplinary modeling framework to explain observations**

**Surprise #3:** Potential of seismoacoustic stations for environmental monitoring

**Station networks with years of continuous data, especially infrasound**

Thanks to Stanford Virgil Kauffman

Fellowship for interdisciplinary research