

Turbulent seismoacoustic signals from a hurricane landfall

Qing Ji

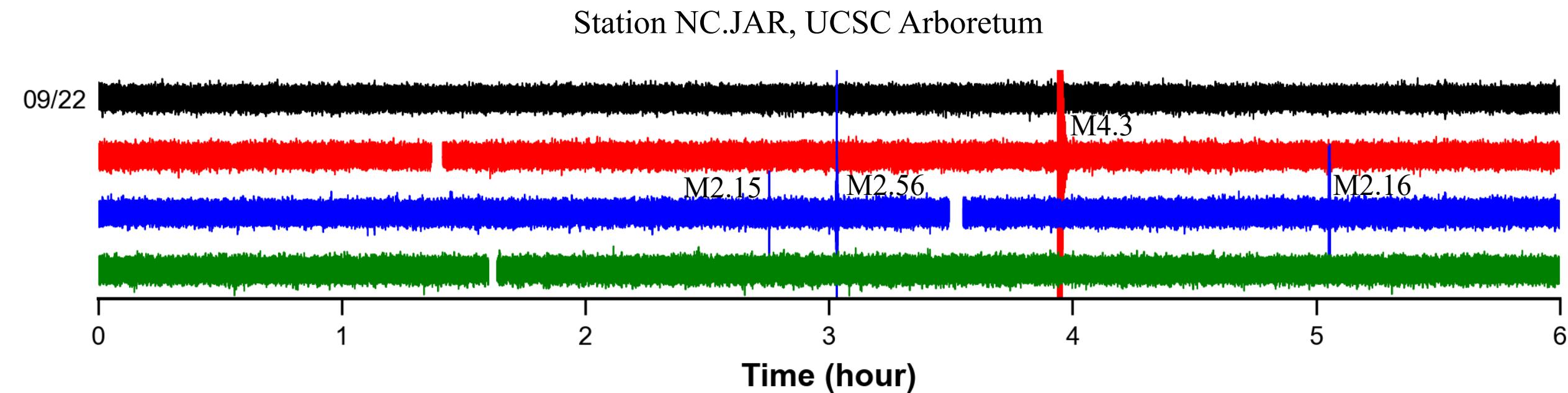
Stanford University

Whole Earth Seminar at UC Santa Cruz, October 1

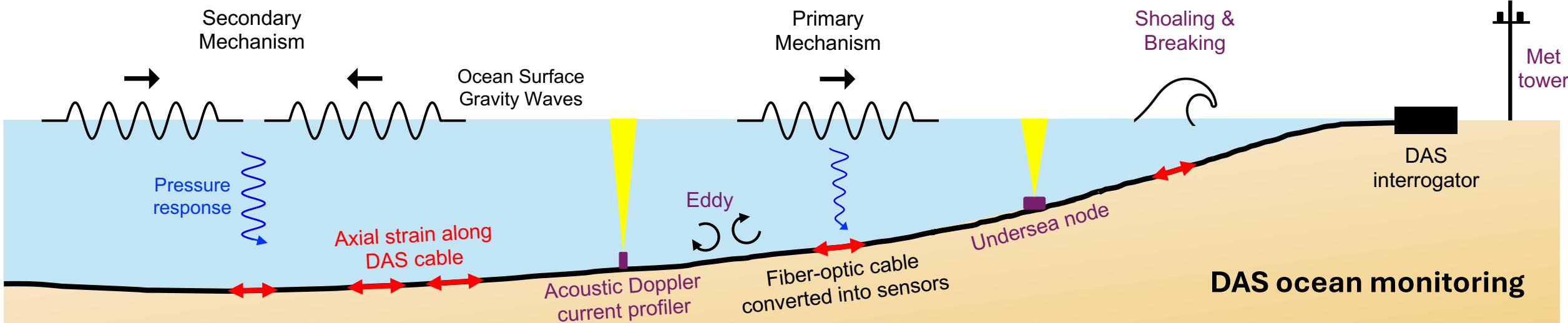
Environmental Seismology

Seismic (and acoustic) sensors continuously record signal from various natural processes, beyond impulsive events such as earthquakes.

New datasets for other fields (e.g., ocean and atmospheric sciences), given advantages of seismic stations.



Seismic instruments for ocean monitoring

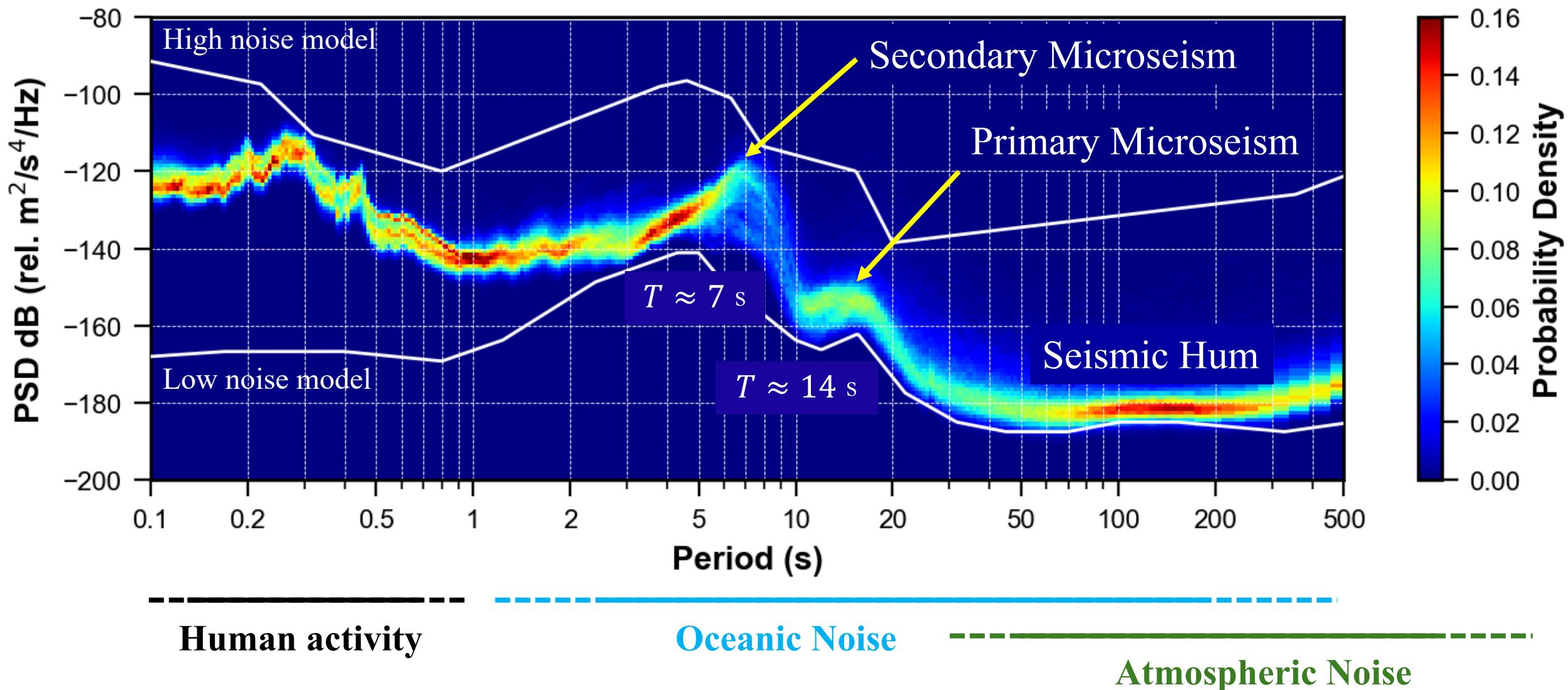


- Continuous monitoring with high resolution
 - e.g., Distributed Acoustic Sensing (DAS) with temporal and spatial resolution
- How ocean processes generate signals recorded by seismic sensors?
 - e.g., strain directly caused by bottom pressure of ocean waves
 - e.g., primary microseism (period ~14 s) and secondary microseism (~7 s)
- How can seismic observations be related / converted into ocean variables?
 - e.g., compare with co-located oceanographic data or ocean models

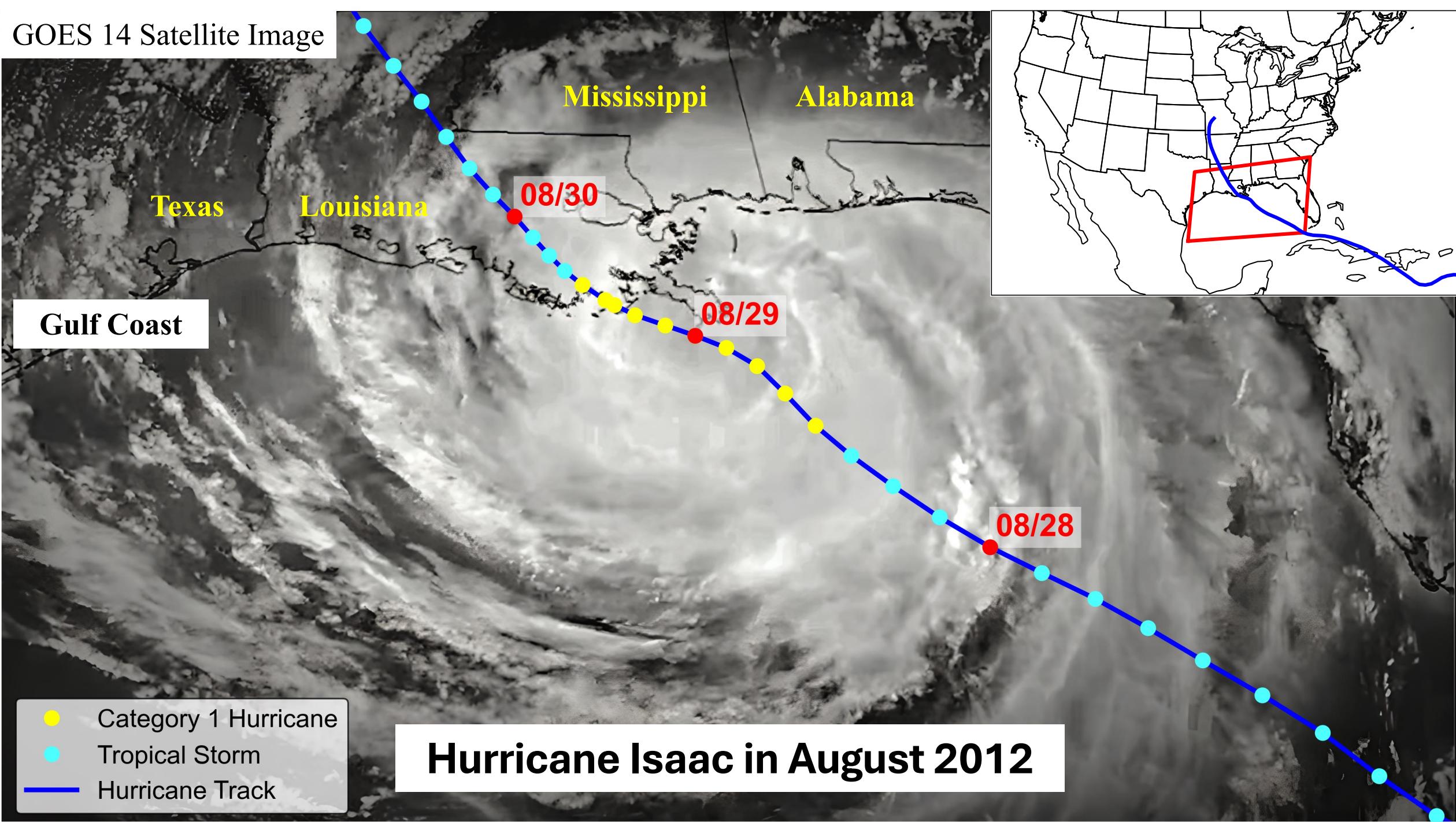
** Example: Microseism from ocean waves

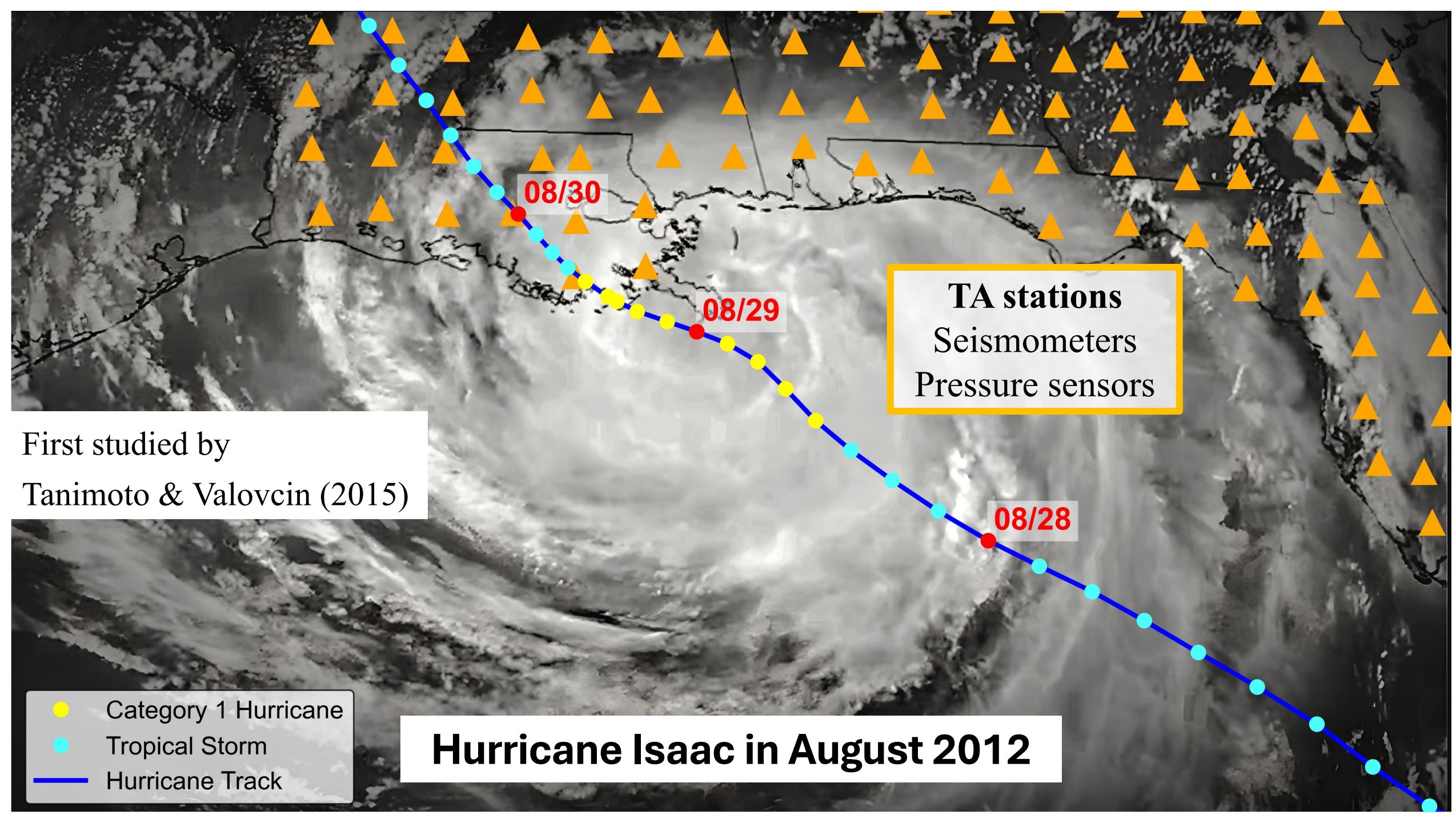
Seismic station CI.JPLS

One-year data, 3-hr window for each PSD curve



GOES 14 Satellite Image





ASOS station

Brown, NC

Automated Surface
Observing Systems

Anemometer/
Wind Tower

Temperature/
Dew Point
Sensor

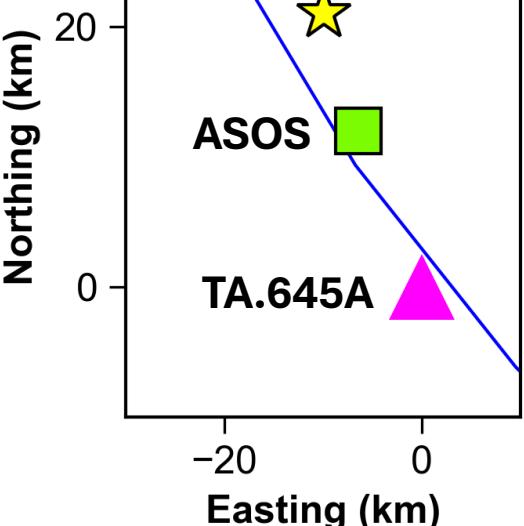
Precipitation
Gauge Precipitation
Identifier

Freezing
Rain
Sensor

Ceilometer Visibility
Sensor

Portable wind tower

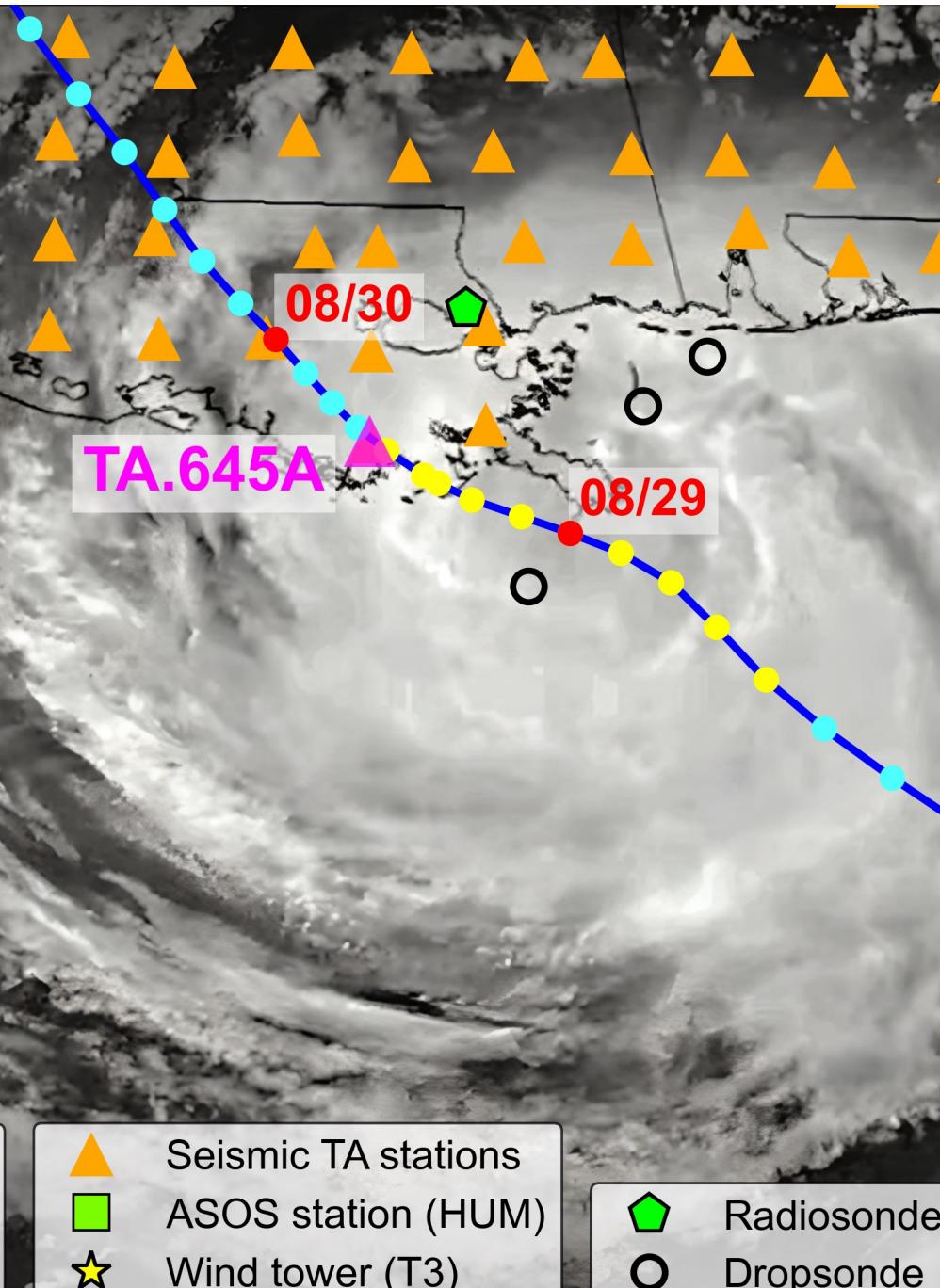
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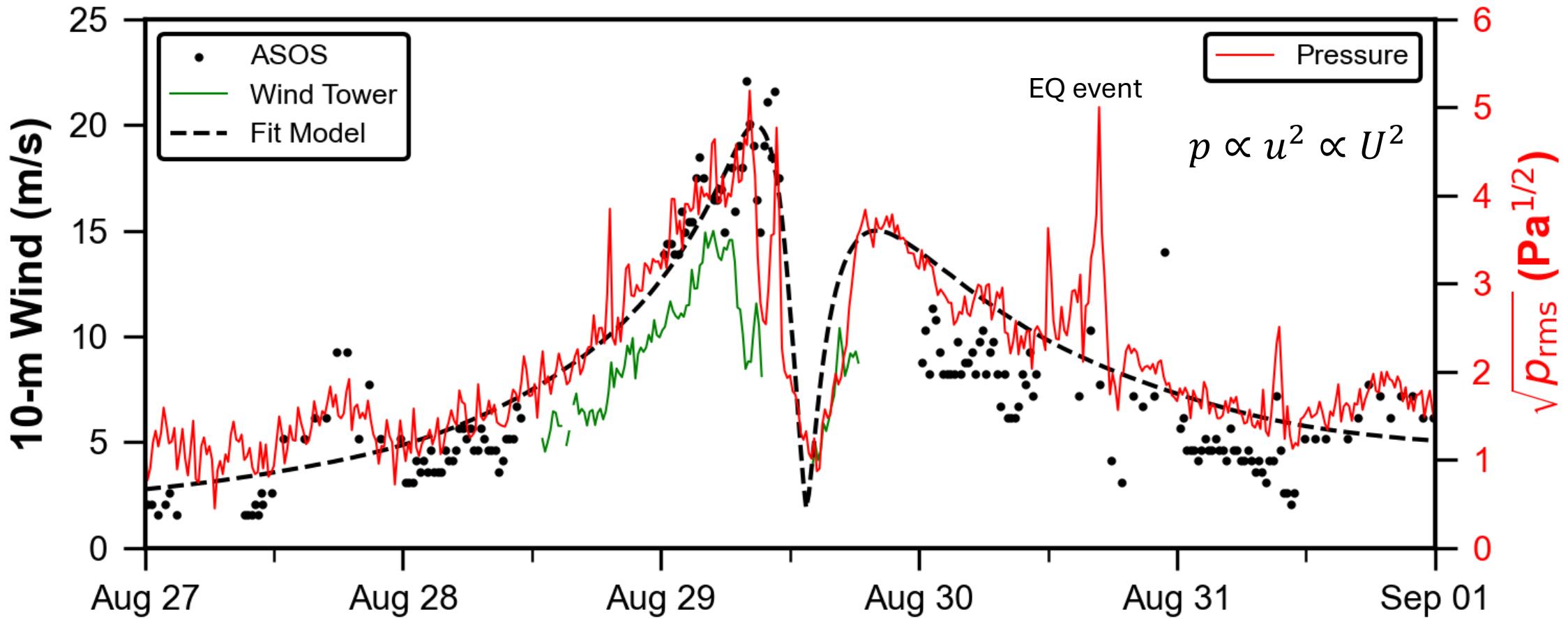
- Category 1 Hurricane
- Tropical Storm
- Hurricane Track

- Seismic TA stations
- ASOS station (HUM)
- Wind tower (T3)

- Radiosonde
- Dropsonde



Infrasound data for hurricane study



Continuous monitoring of the evolution of atmospheric quantities
(e.g., wind speed, wind turbulence)

1. Observation

Seismoacoustic imprints of Hurricane Isaac in 2012 during landfall

2. Interdisciplinary modeling

**How atmospheric processes
generate these observed signals?**

Large-eddy simulation (LES) of turbulent surface pressure

Quasi-static seismic modeling of elastic response under turbulent pressure

3. Turbulent pressure spectrum from infrasound data

Estimate turbulent dissipation rate ε from pressure spectrum

**How to relate these observations
with atmospheric variables?**

Seismic station with environmental sensors

Channel

LH[ZNE]

Observation

Three-component seismic
ground motion

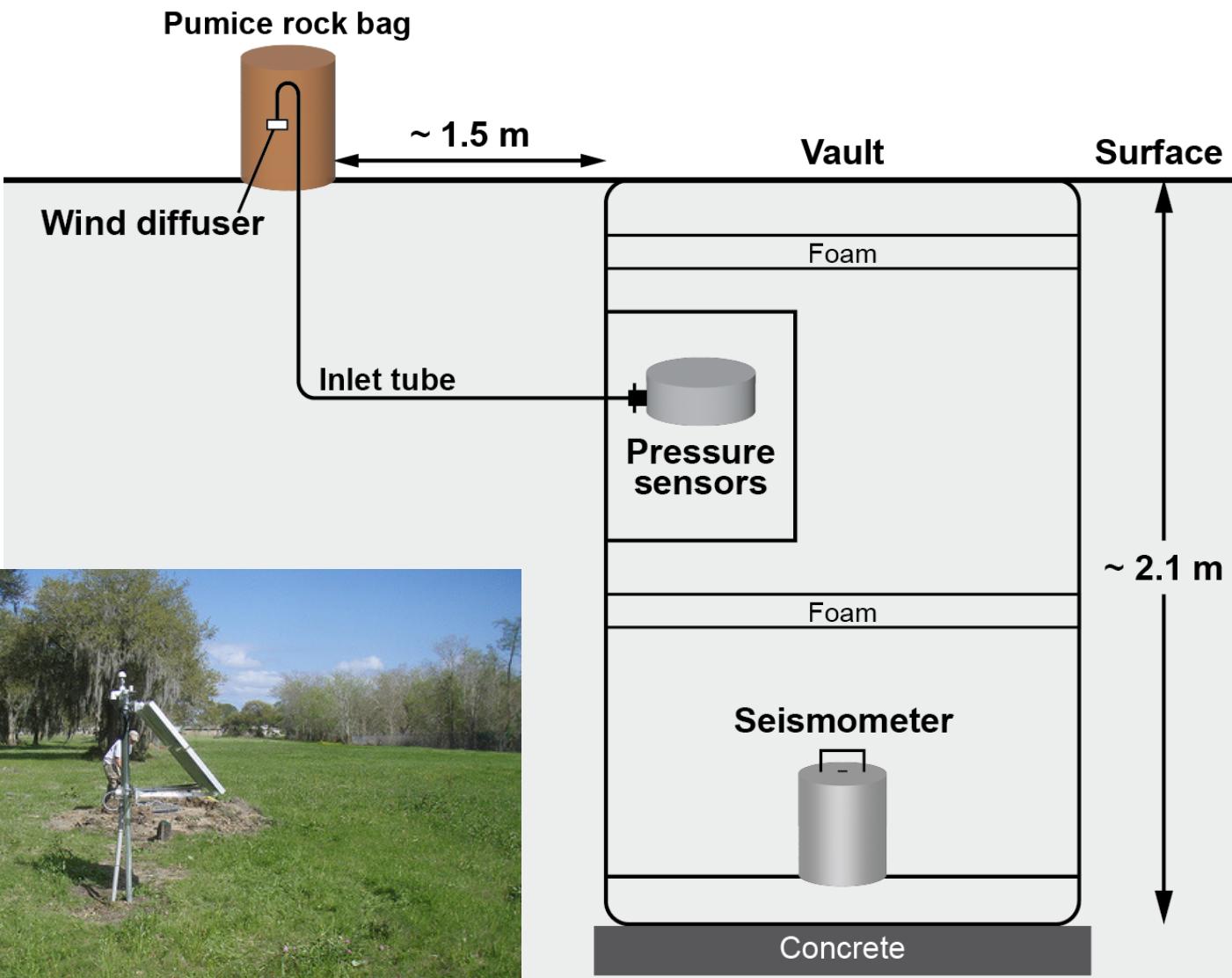
LDO

Barometric pressure

LDF

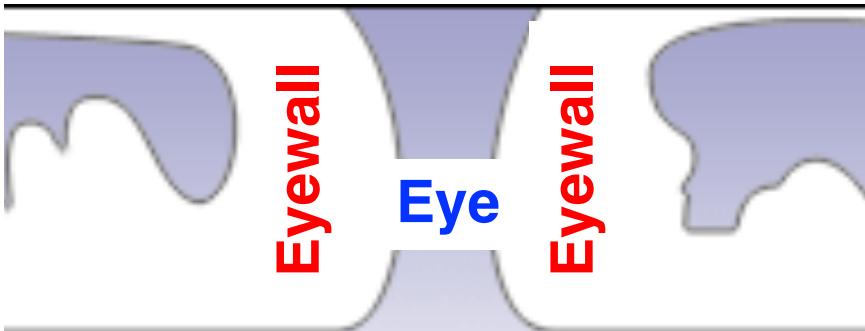
Infrasound pressure

L: Long-period (1 Hz) B: Broadband (40 Hz)



Modified from Tytell et al. (2016)

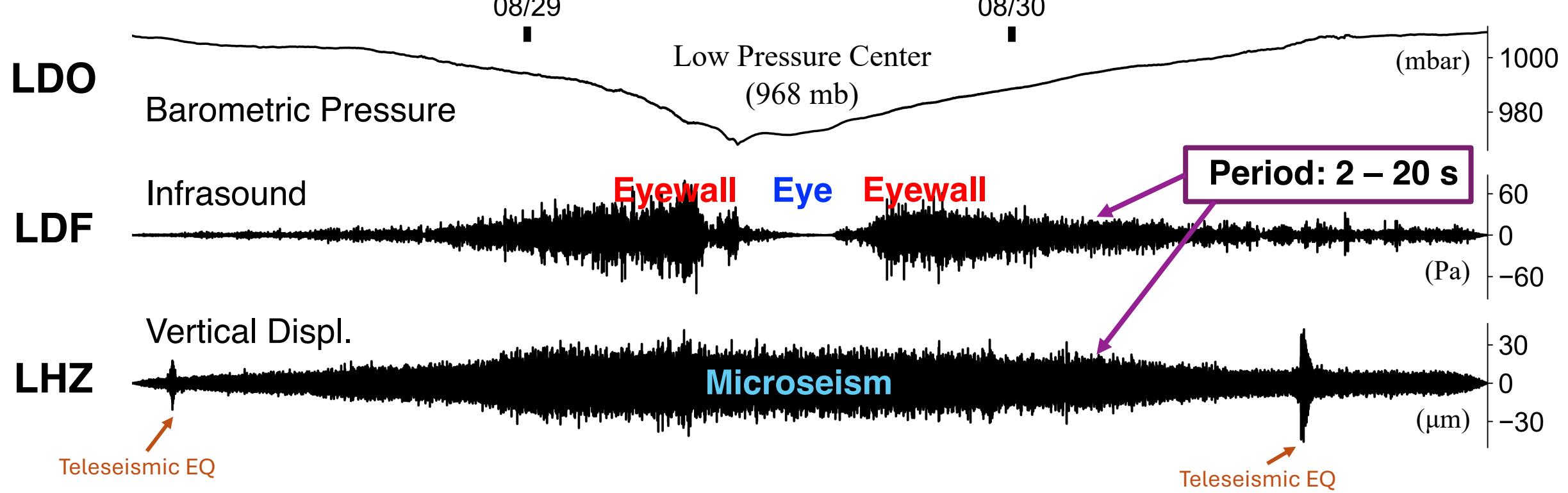
As hurricane passes the station



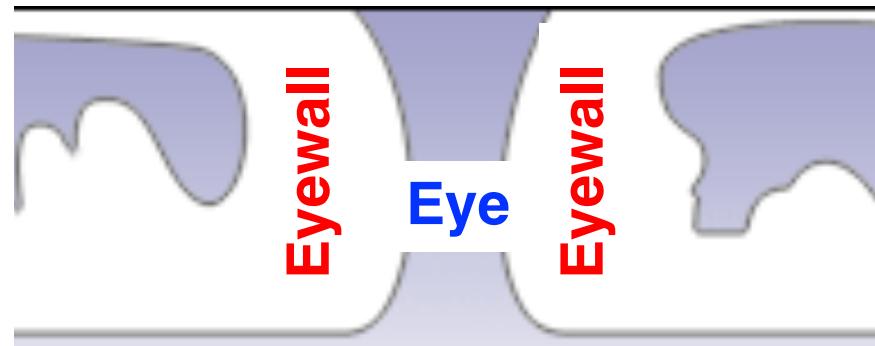
TA Station

L: Long period channel

D: Pressure



As hurricane passes the station



TA Station

L: Long period (1 Hz sampling)

D: Pressure

LDO

Barometric Pressure

08/29

Low Pressure Center
(968 mb)

08/30

(mbar)

1000
980

LDF

Infrasound

Period: 20 – 100 s

40
0
-40

LHZ

Vertical Displ.

Eyewall Eye Eyewall

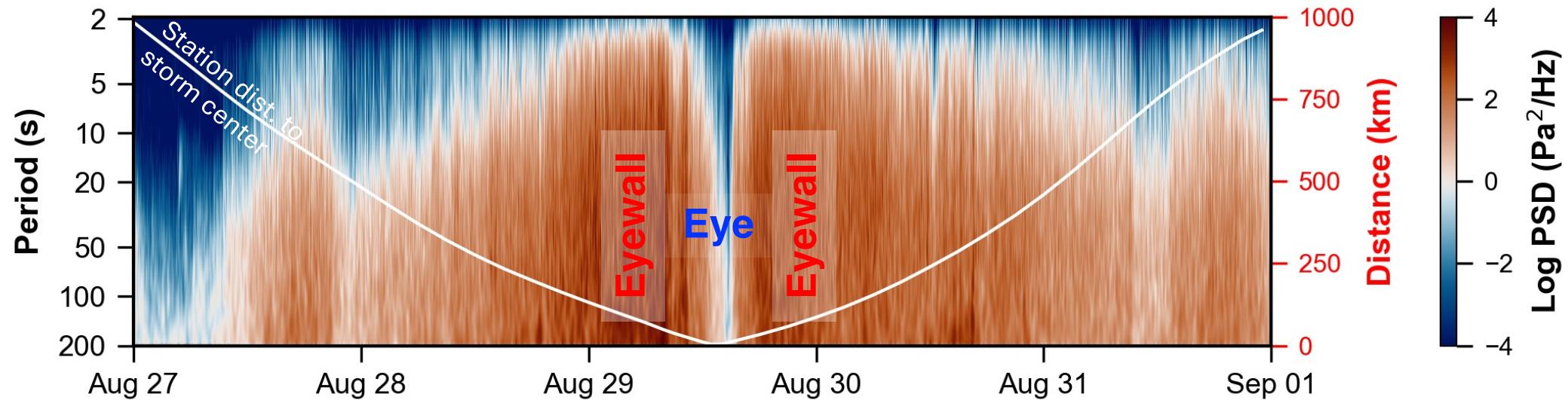
15
0
-15

Teleseismic EQ

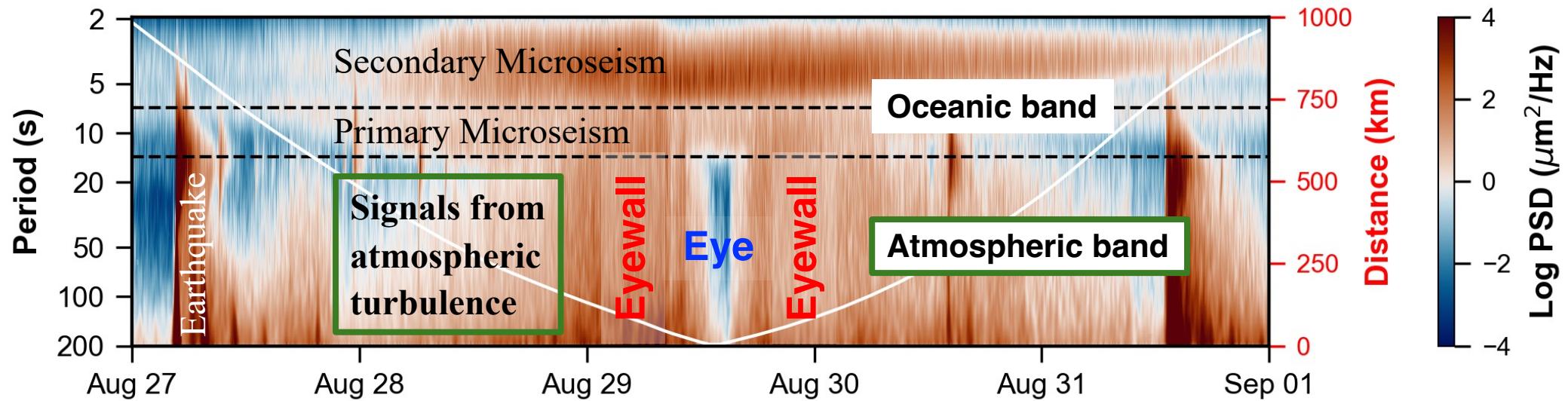
Teleseismic EQ

Wavelet spectrograms of infrasound & seismic data

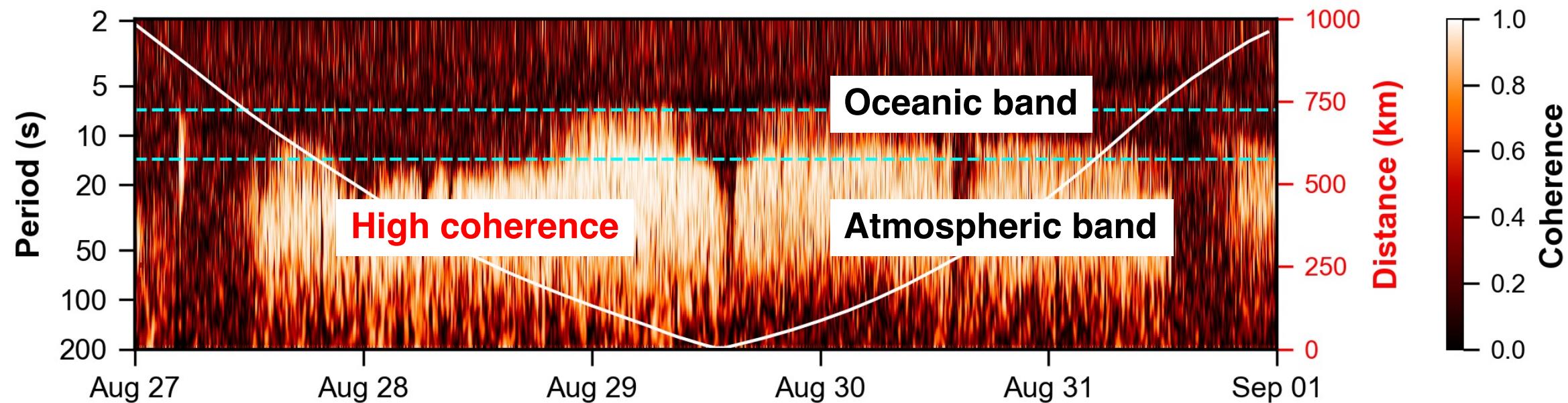
Surface Pressure



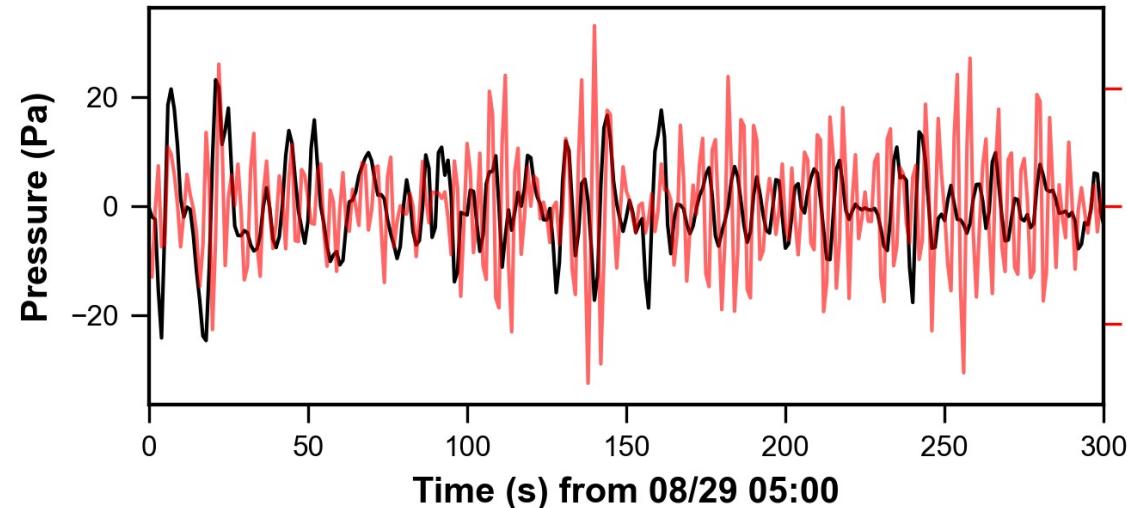
Seismic Vertical Displacement



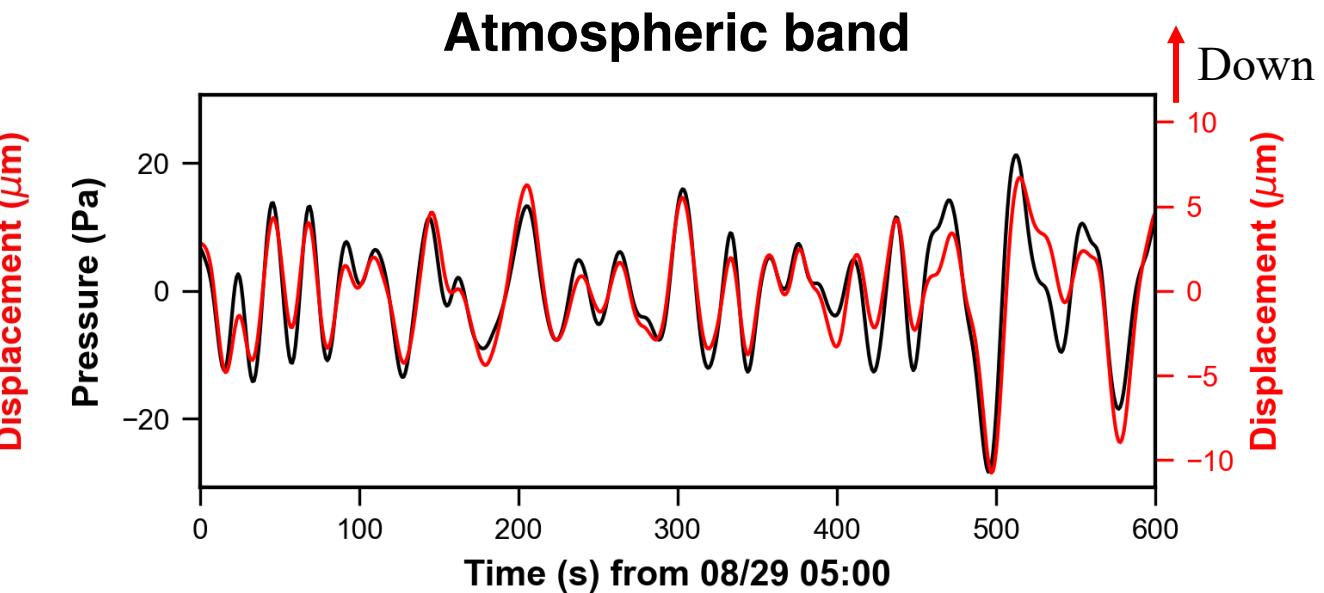
High coherence indicates local quasi-static response



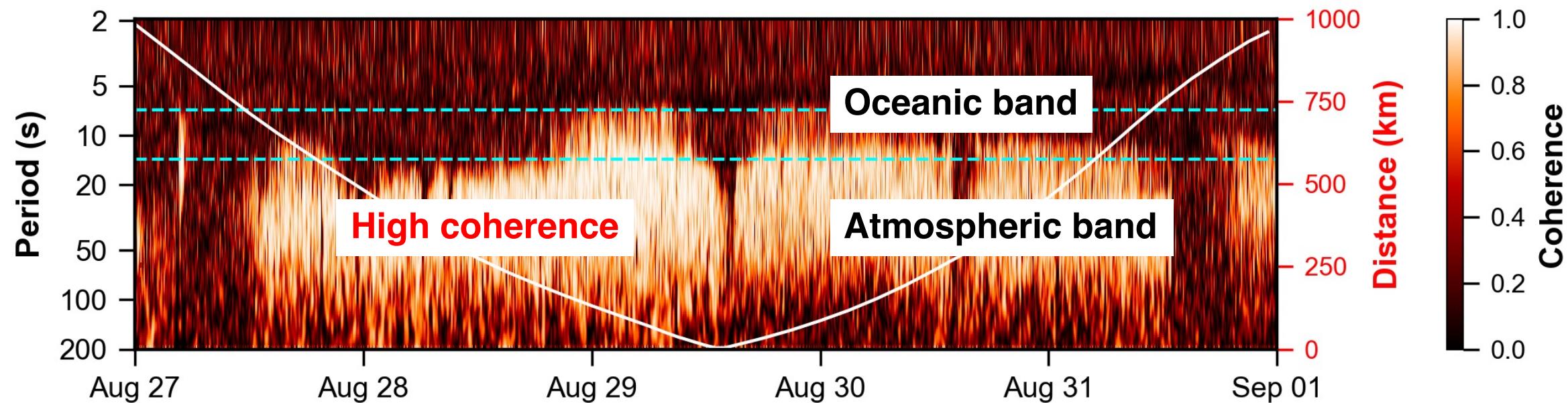
Oceanic band



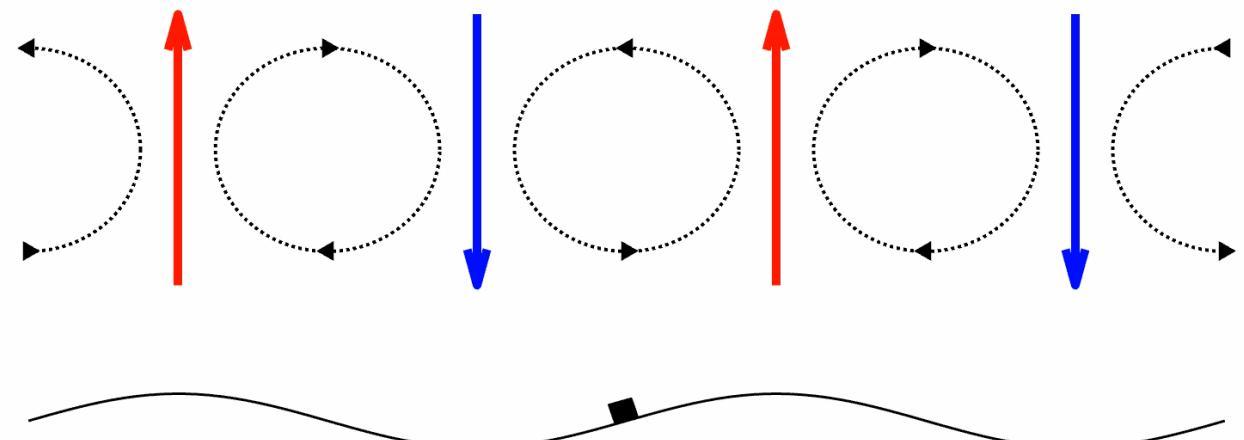
Atmospheric band



High coherence indicates local quasi-static response



Simplified illustration
Sorrells (1971) theory
Pressure wave model



1. Observation

Seismoacoustic imprints of Hurricane Isaac in 2012 during landfall

2. Interdisciplinary modeling

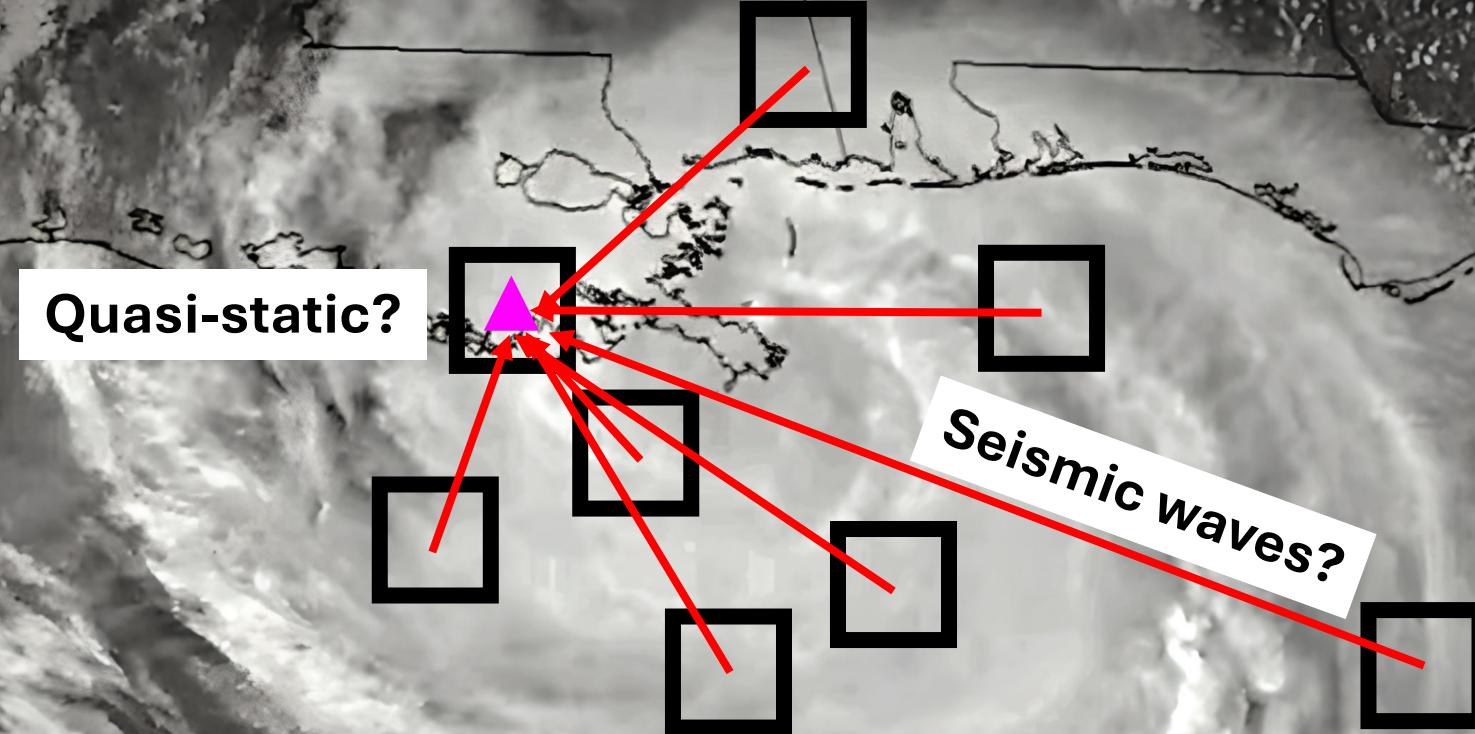
Large-eddy simulation (LES) of turbulent surface pressure

Quasi-static seismic modeling of elastic response under turbulent pressure

3. Turbulent pressure spectrum from infrasound data

Estimate turbulent dissipation rate ε from pressure inertial subrange

Dominant contribution to seismic power?



Follow the framework in Tanimoto & Valovcin (2015): Decompose hurricane into independent sources

Seismic response is “local”



Consistent with

- Observed high coherence
- Shallow compliant sediments
- Quasi-static limit $\frac{\omega r}{c_{seis}} \ll 1$

Follow the framework in Tanimoto & Valovcin (2015): Decompose hurricane into independent sources

Dominant source is ~ km around the station (Ji & Dunham, 2024)

Propagating waves from far regions are negligible, not as previously hypothesized

Quasi-static seismic response

Modeling
(Surface field)

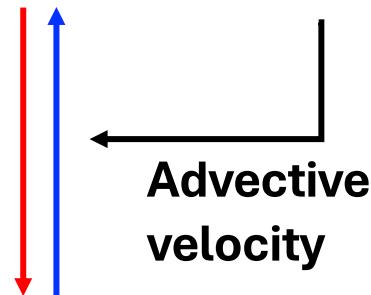
Vertical
displacement

Static Green's function
(laterally homogeneous)

Surface
pressure

$$u_z(\mathbf{k}, \omega) = G(|\mathbf{k}|) p(\mathbf{k}, \omega)$$

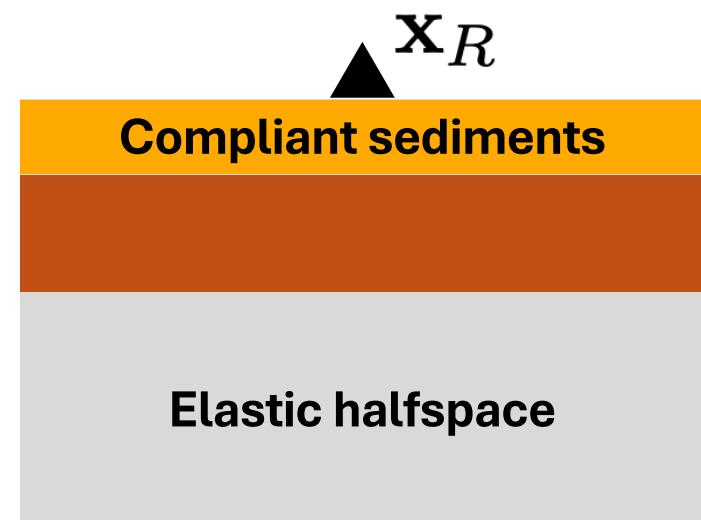
Space-time
conversion



Observation
(Single point)

$$u_z(\mathbf{x}_R, \omega) = L(\omega) p(\mathbf{x}_R, \omega)$$

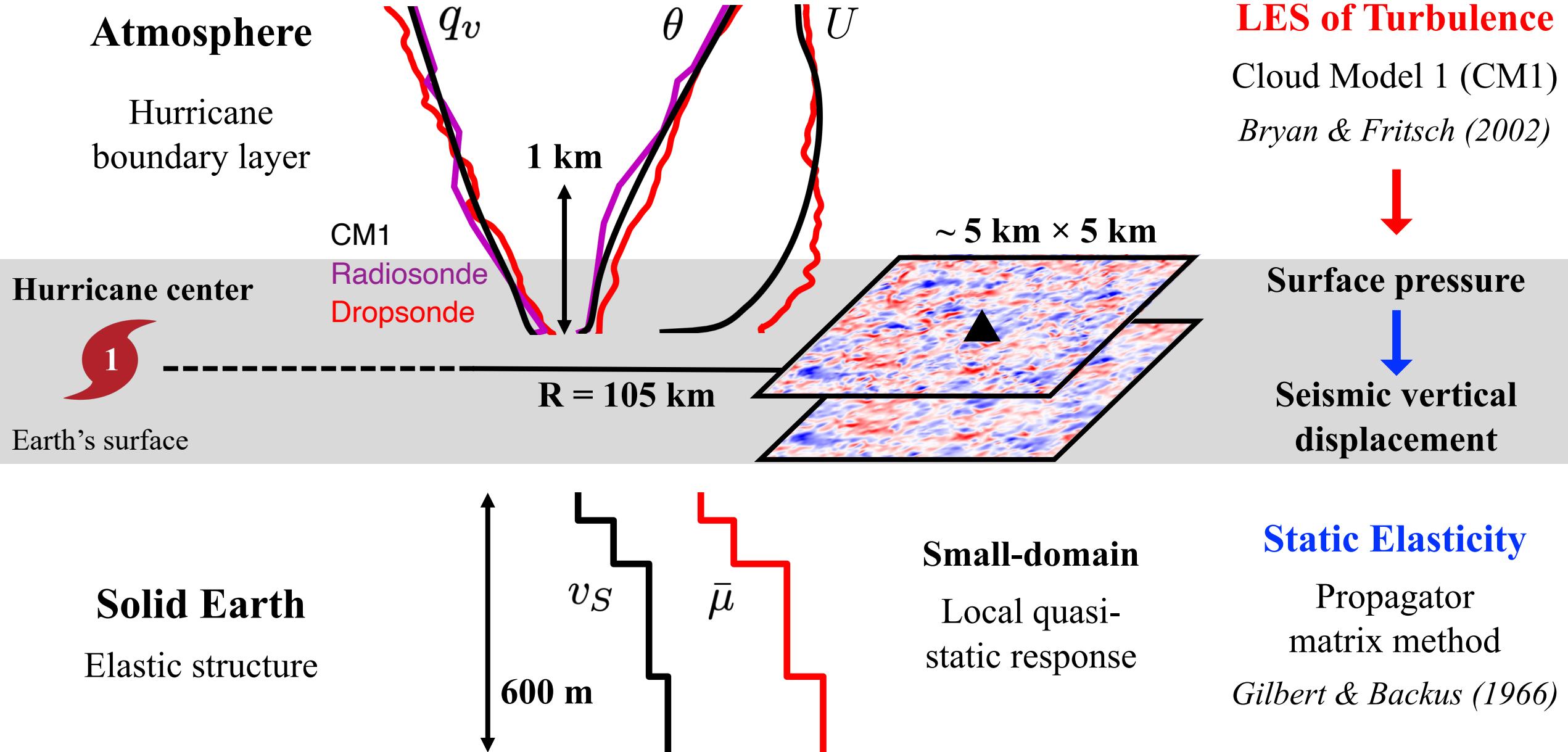
Transfer function
(i.e., linear estimator)



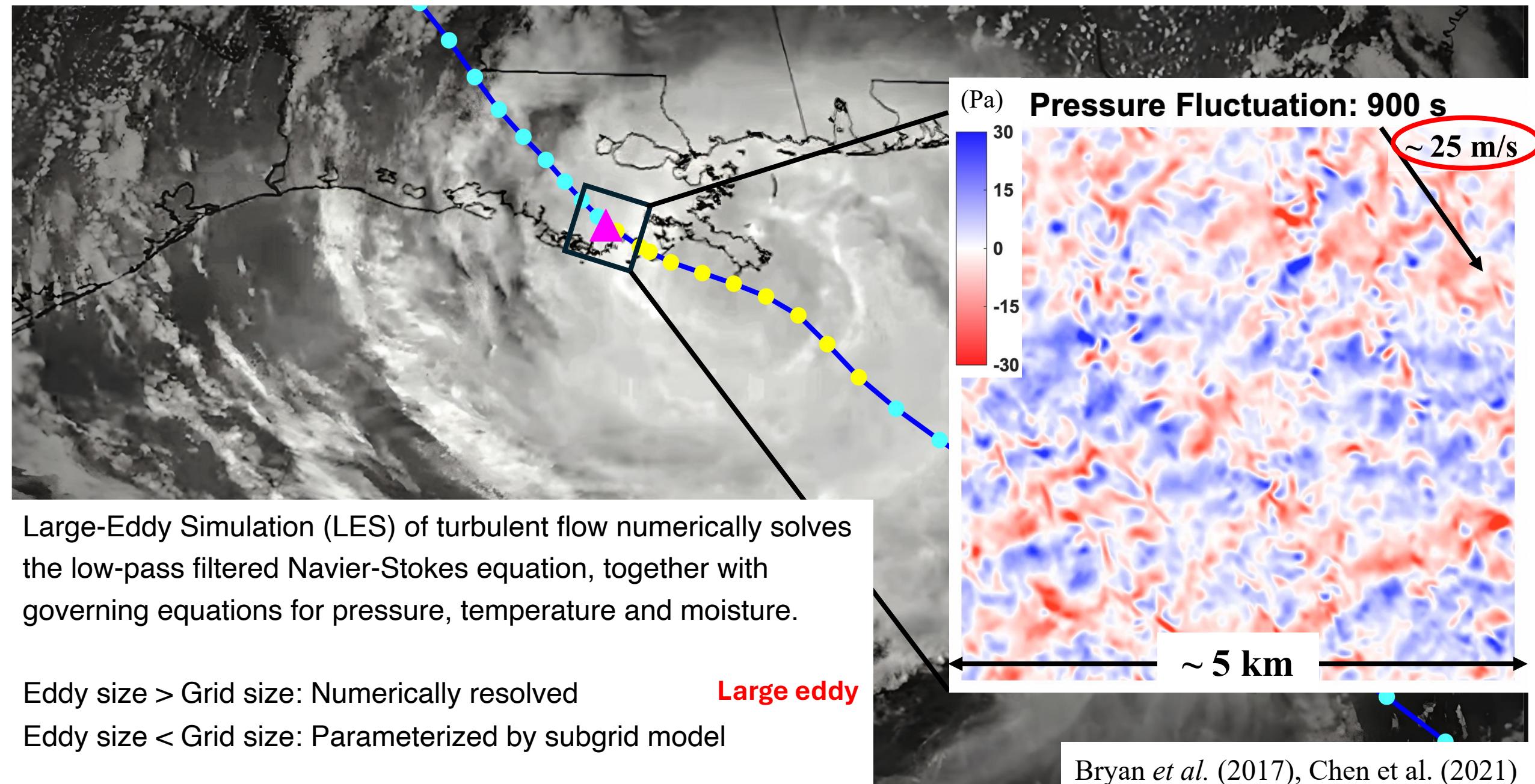
k Horizontal wavenumber

ω Angular frequency

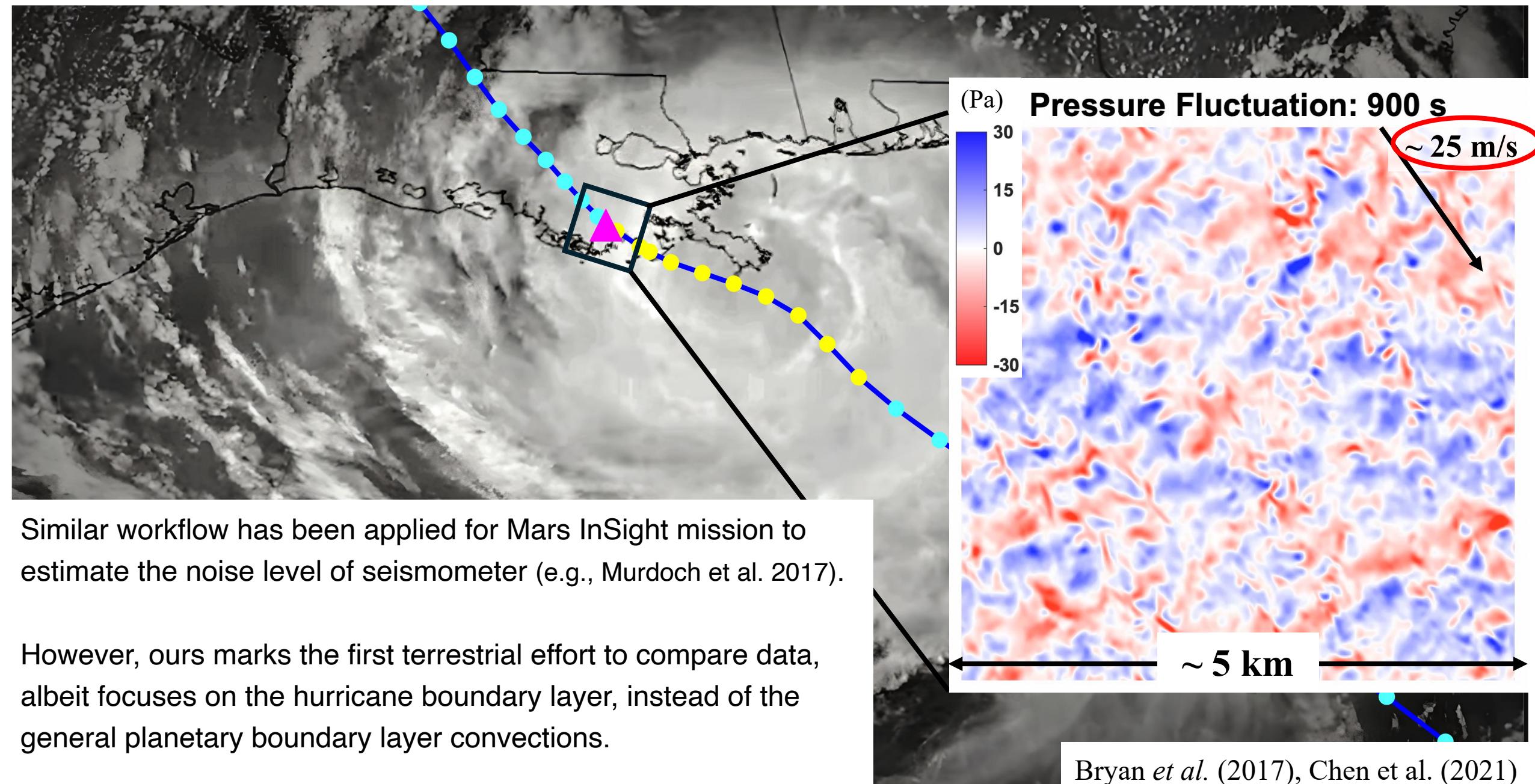
Interdisciplinary modeling



LES of Hurricane Boundary Layer (HBL) on land



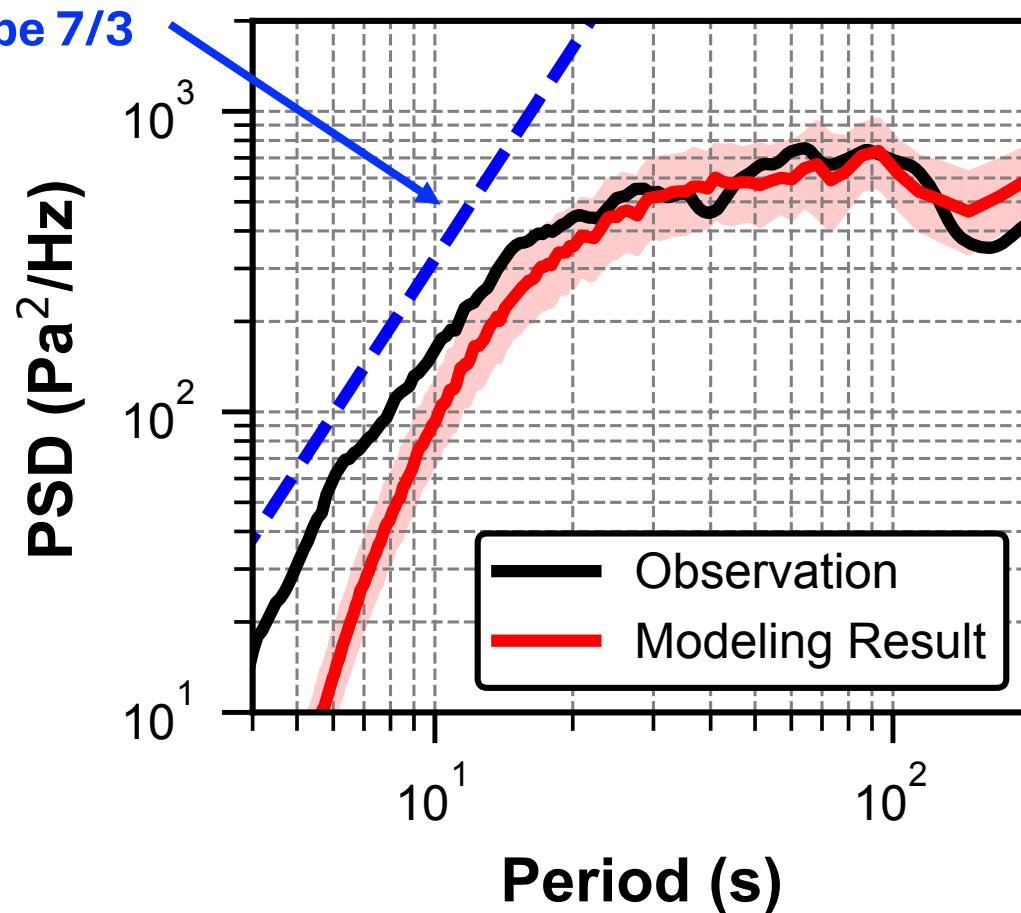
LES of Hurricane Boundary Layer (HBL) on land



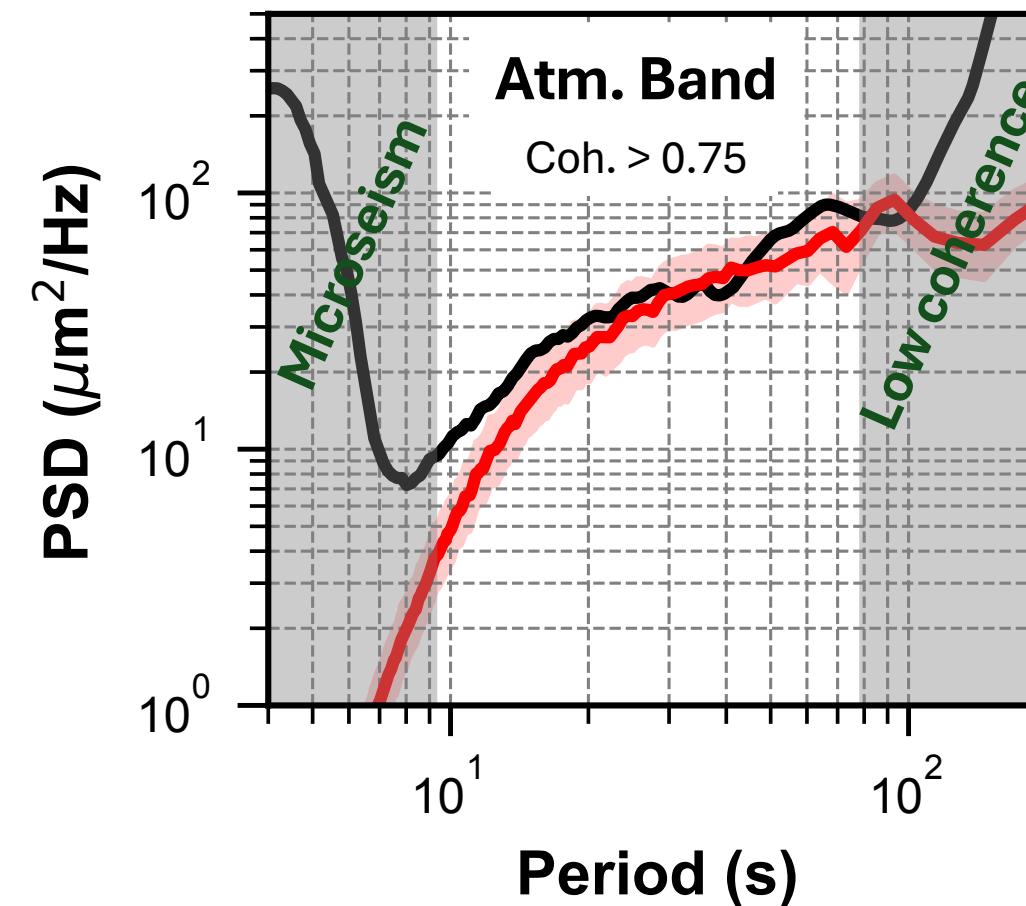
Infrasound & seismic spectra

Inertial subrange

Surface Pressure



Vertical Seismic Displ.



Infrasound data can be used for turbulent analysis.

Seismic signals originate from turbulence in the atmospheric band.

Summary: Generation mechanisms of seismic ambient noise

Natural processes from ocean, atmosphere, ...



Fourier mode
 $e^{i(\mathbf{k} \cdot \mathbf{x} - \omega t)}$

Seismic noise sources
(at Earth's surface)

$$p(x, y, t)$$

$$c_{\text{source}} \ll c_{\text{seis}}$$

$$c_{\text{source}} \approx c_{\text{seis}}$$

Local quasi-static response

Turbulent imprints

Ocean wave imprints on OBS ...

Dynamic seismic waves

Microseism, seismic hum

Background free oscillations ...

1. Observation

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Large-eddy simulation (LES) of turbulent surface pressure

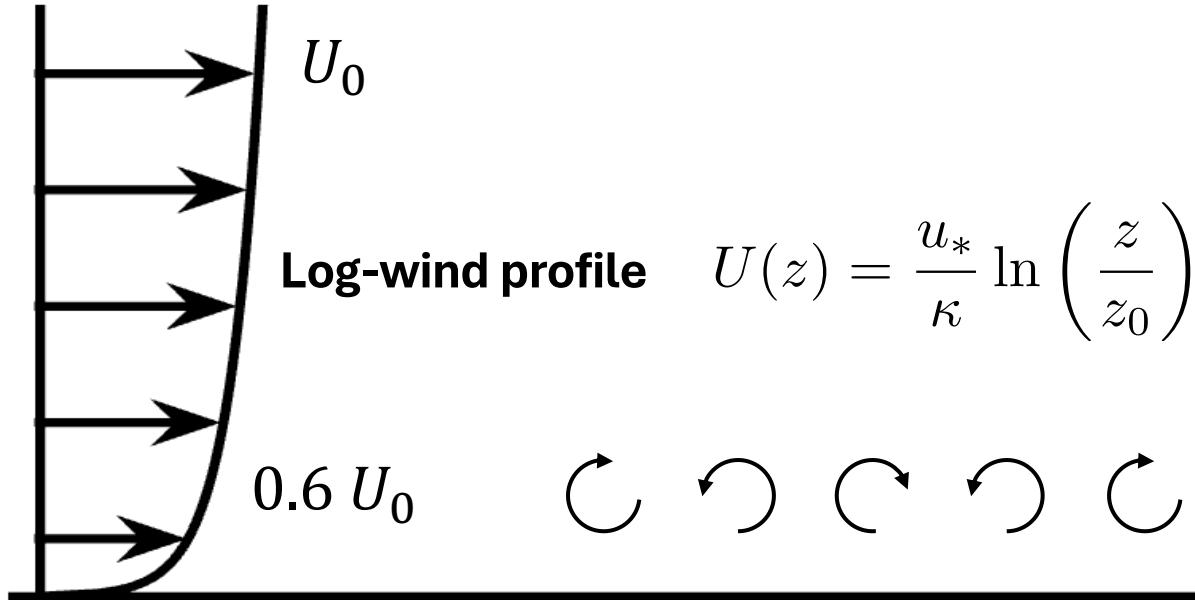
Quasi-static seismic modeling of elastic response under turbulent pressure

3. Turbulent pressure spectrum from infrasound data

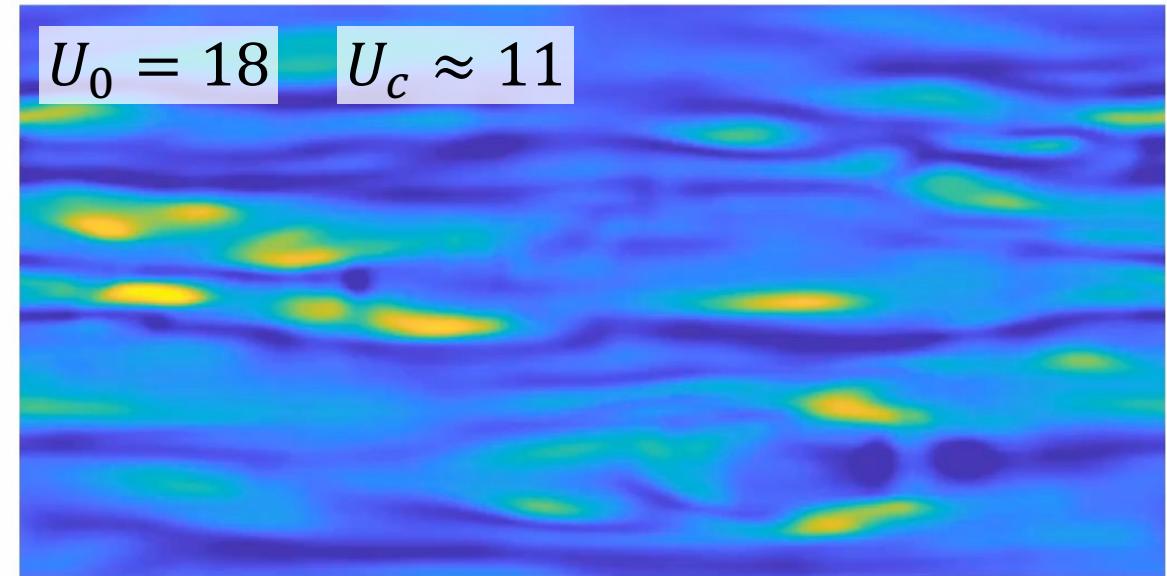
Estimate turbulent dissipation rate ε from pressure spectrum

** Wall stress / pressure of turbulent boundary layers

Turbulent boundary layer



Wall shear stress of a channel flow



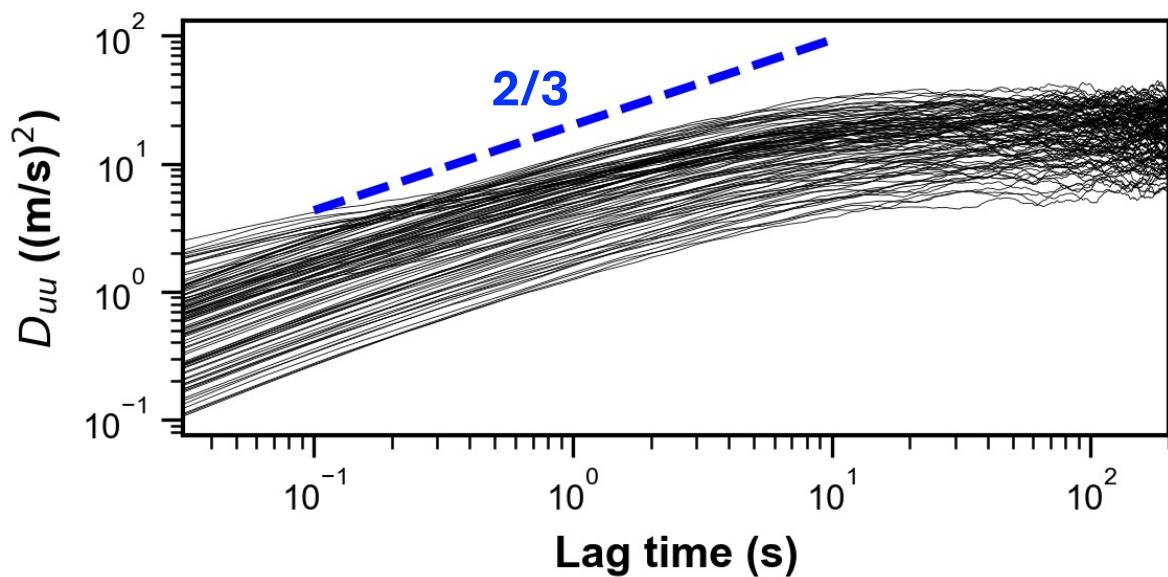
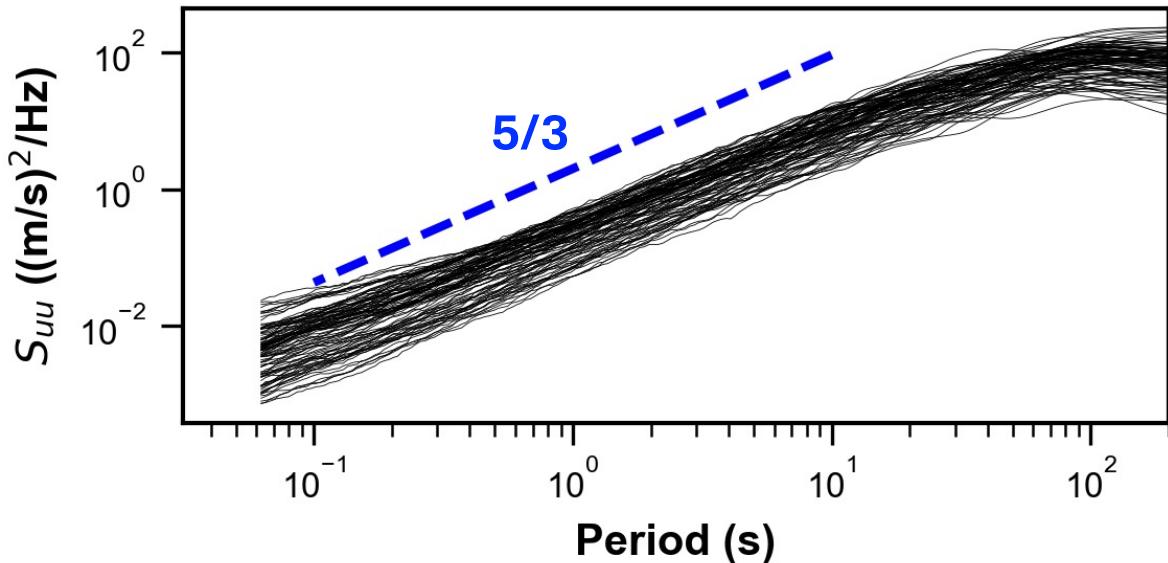
Velocity is zero at the wall, but wall stress / pressure advects with U_c (advective velocity)

Theoretical and experimental studies show $U_c \approx 0.6 U_0$ with free-stream velocity U_0

Assuming a log-wind profile, U_c can further be related to $U_{10\text{ m}}$ at a specific height

Turbulent wind spectrum from tower data

Inertial subrange



Streamwise velocity PSD

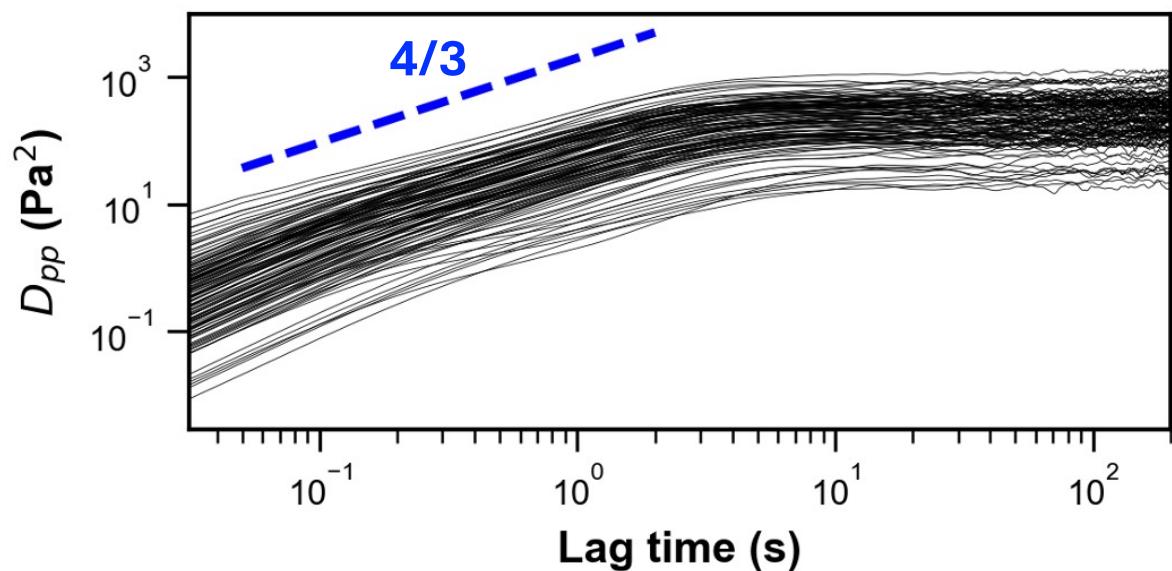
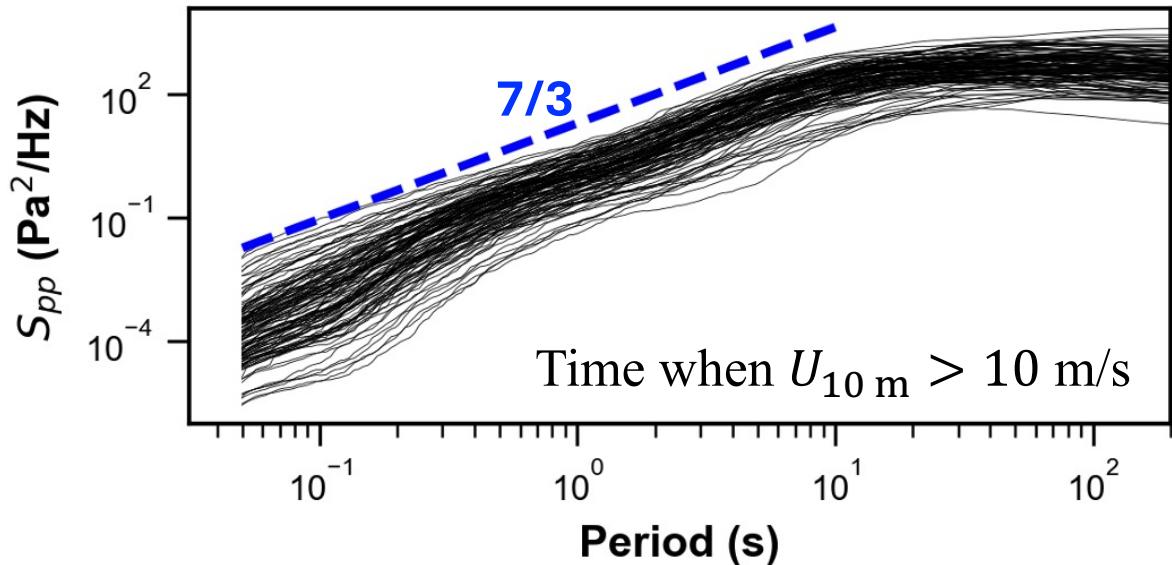
$$S_{uu}(f) = \alpha_u \left(\frac{\varepsilon U}{2\pi} \right)^{2/3} f^{-5/3}$$

Dissipation rate ε is a key parameter describing the turbulence statistic and contributes to an important energy source for hurricanes.

Streamwise structure function

$$D_{uu}(\tau) \equiv \overline{[u(t + \tau) - u(t)]^2} = C_u (\varepsilon U)^{2/3} \tau^{2/3}$$

Turbulent pressure spectrum from infrasound data



Pressure PSD

$$\frac{1}{\rho^2} S_{pp,t}(f) = \tilde{\alpha}_p \left(\frac{\varepsilon U_c}{2\pi} \right)^{4/3} f^{-7/3}$$

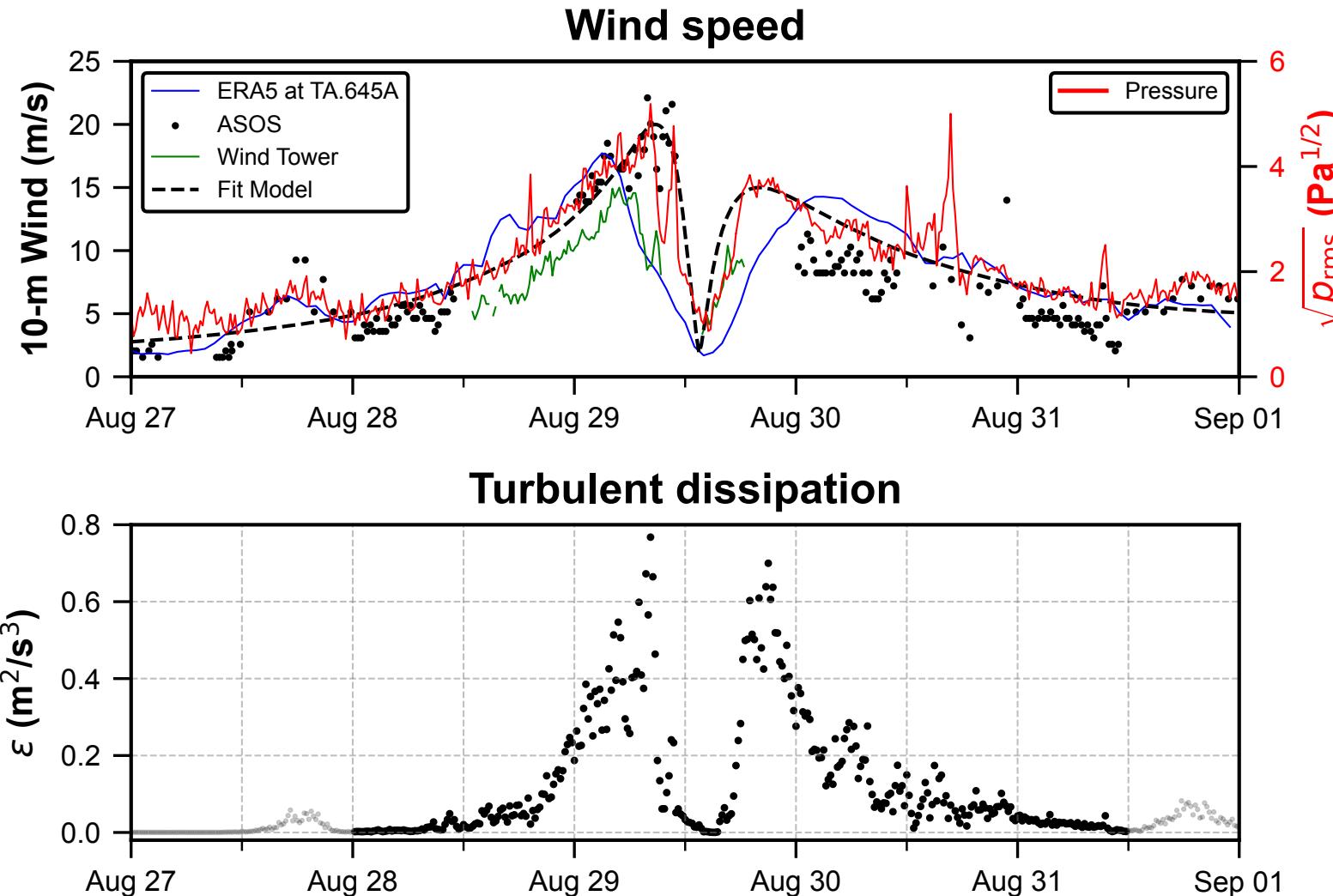
The 7/3 slope is turbulence-turbulence interaction or “slow-term” contribution (George et al. 1984)

Assume $U_c \approx 1.6 U_{10\text{ m}}$

Pressure structure function

$$\frac{1}{\rho^2} D_{pp,t}(\tau) = C_p (\varepsilon U_c)^{4/3} \tau^{4/3}$$

Continuous monitoring of hurricane landfall



Fluctuation: $p_{rms} \propto u_{rms}^2$

Turbulent intensity: $u_{rms} \propto U$

Fit wind model: $U_{10 \text{ m}} = U(r)$

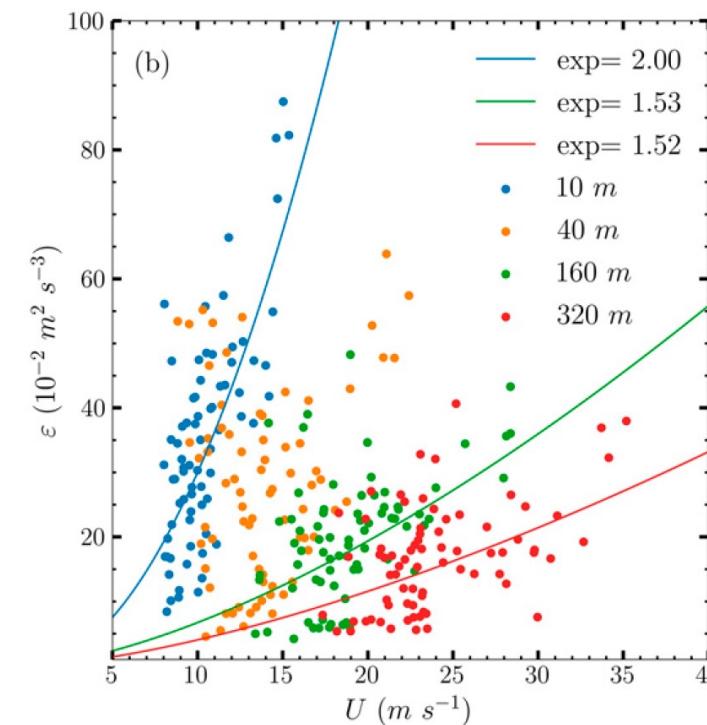
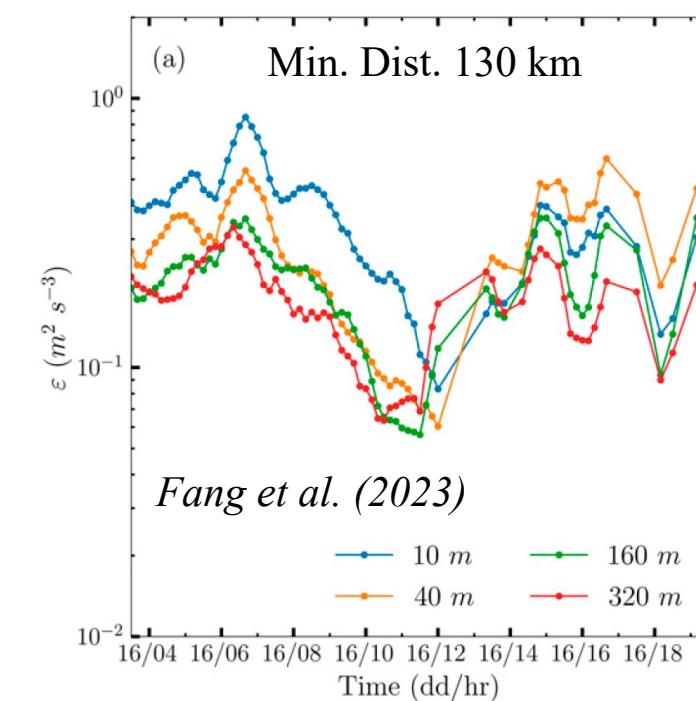
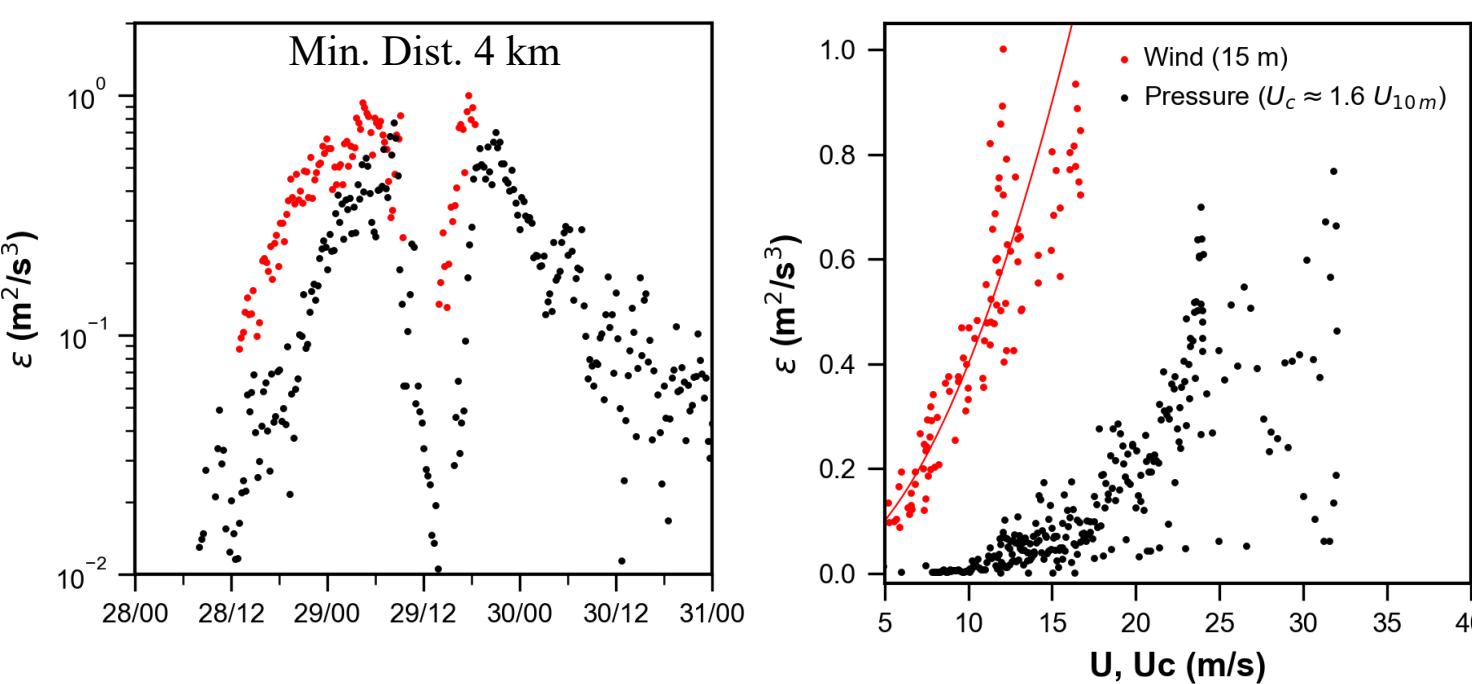
ERA5 reanalysis has a wider radius of the hurricane eye.

$O(0.1 \text{ m}^2/\text{s}^3)$ dissipation rate

Interpreted as ϵ for height $\sim 100 \text{ m}$

Wind model not accounted for diurnal cycles (daytime convection)

Turbulent dissipation rate during hurricane landfall

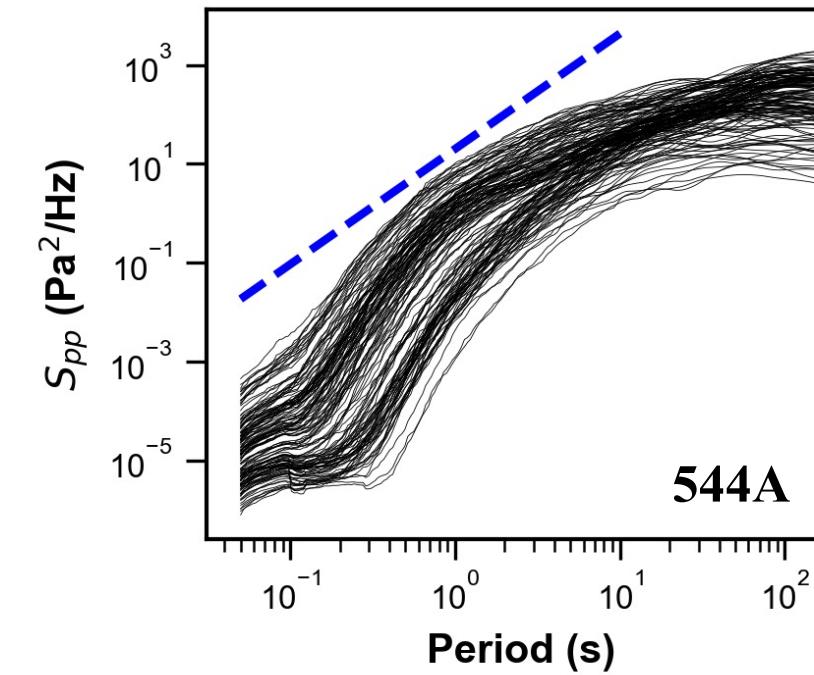
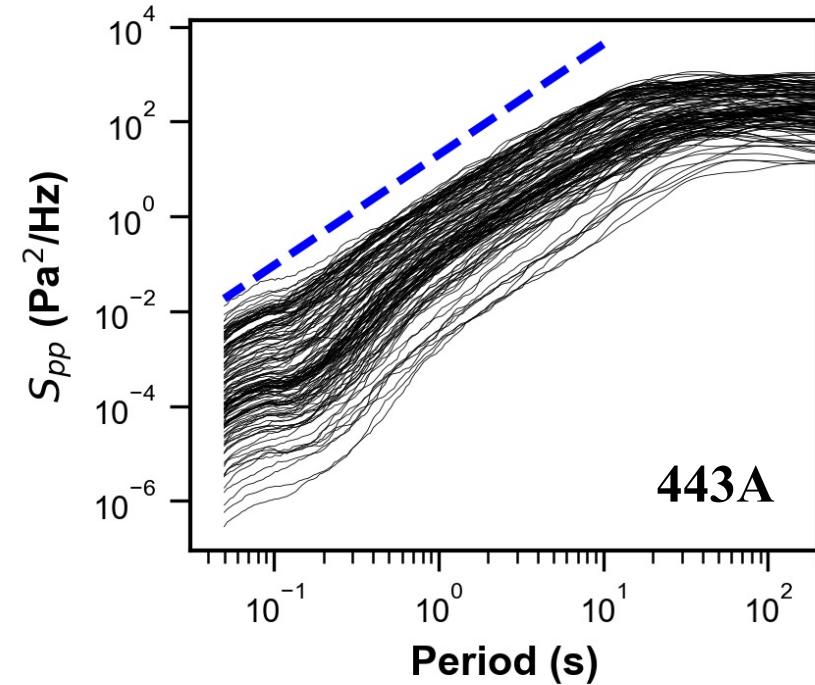
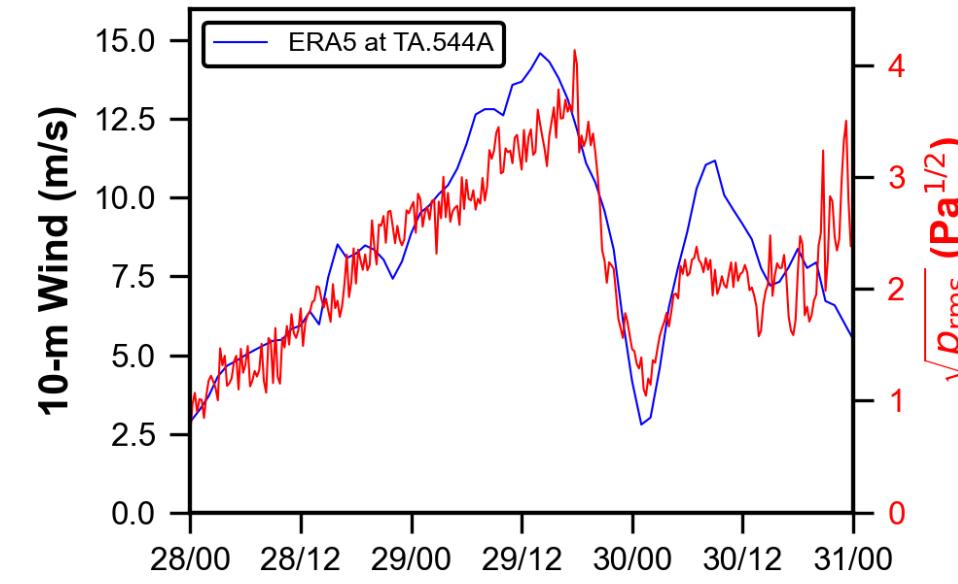
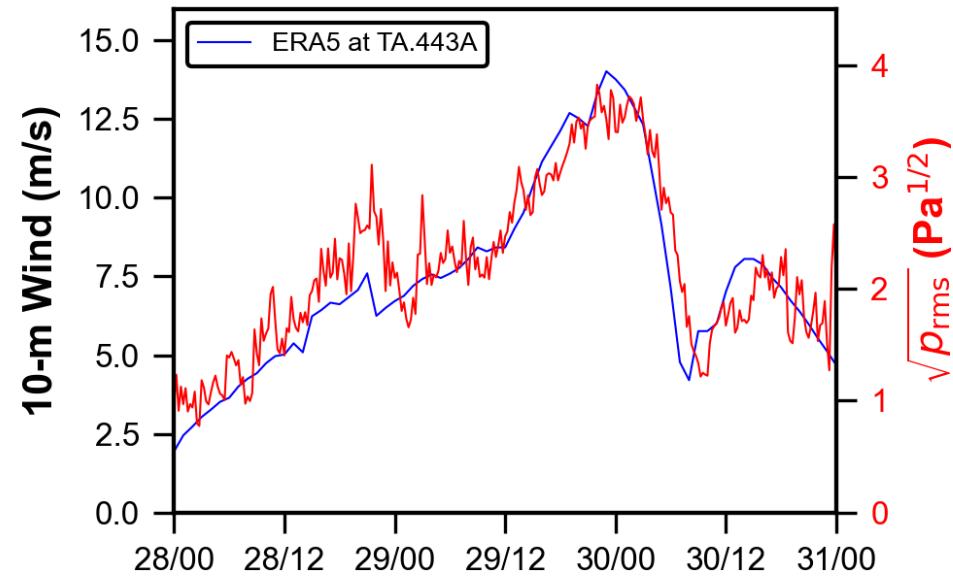


Estimation of ε from pressure uses advective velocity U_c . The result is interpreted as local ε at height where wind speed is U_c .

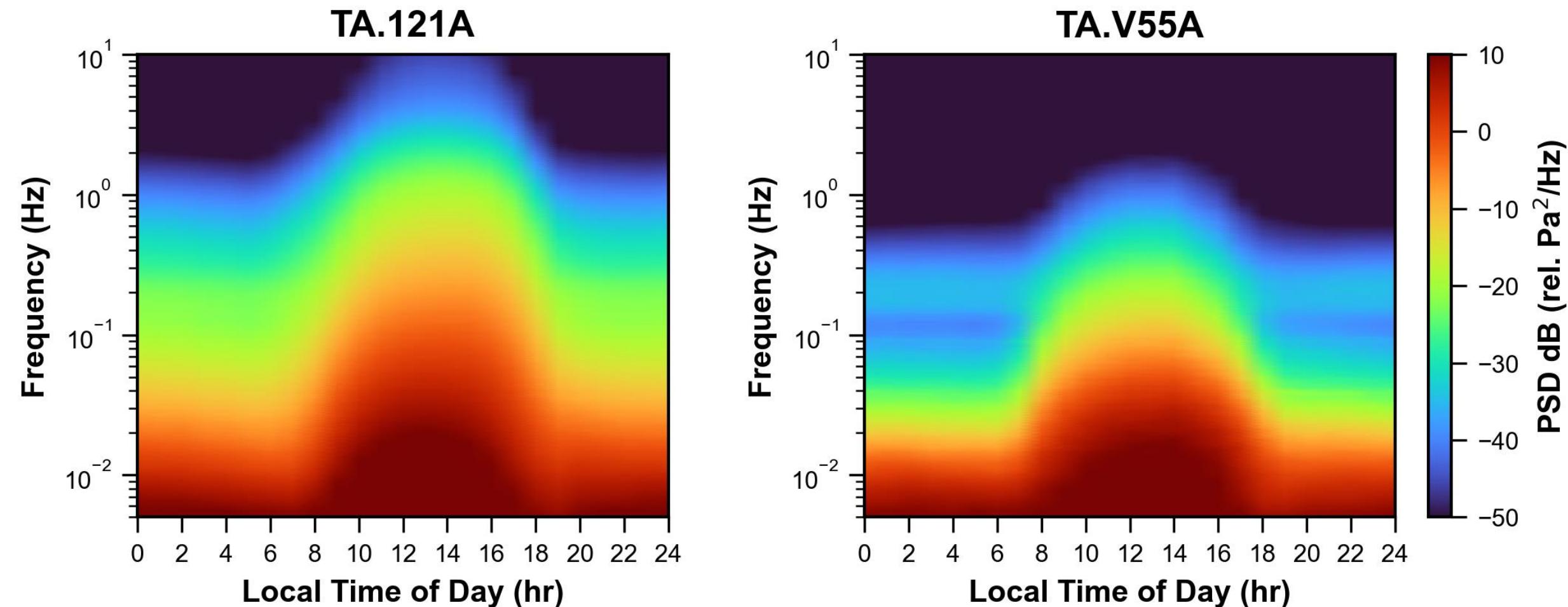
Outliers of low ε at high wind speed is due to using the fit wind model not including a transient drop in wind.

Infrasound data can provide a lower bound of ε of the surface layer.

Potential for monitoring general atmospheric conditions



Diurnal cycles shown in pressure spectral amplitude



Diurnal variation in turbulence in response to solar heating (Stull, 1988)

Surprise #1: Seismic stations record in-situ data of Hurricane Issac after landfall

Distinct seismic ground motion contributed by ocean and atmosphere

Surprise #2: Turbulence explains the seismoacoustic signatures (in the atm. band)

Interdisciplinary modeling framework to explain observations

Surprise #3: Potential of seismic stations for environmental monitoring

Station networks with years of continuous data