Bias Adjustment of ESPO-R5 v1.0

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The temperature and precipitation data from the simulations were first extracted over an area covering North America and, if necessary, converted into daily values. Then using the ESMF software, accessed through its python xESMF interface, all the extracted simulation data is interpolated bilinearly to the ERA5-Land grid.

The ESPO-R5 v1.0 bias adjustment procedure then uses xclim algorithms to adjust simulation bias following a quantile mapping procedure. In particular, the algorithm used is inspired by the "Detrended Quantile Mapping" method described by [2]. The procedure is univariate, bipartite, acts differently on the trends and the anomalies and is applied iteratively on each day of the year (grouping) and on each grid point.

1 Variables

Adjustments are applied separately for each of the 3 variables. Adjusting `tasmax` and `tasmin` independently can lead to physical inconsistencies in the final data (i.e. cases with `tasmin > tasmax`) ([4], [1]). To ensure a physically consistent dataset (for this aspect at least), we compute the daily temperature range (or amplitude; `dtr = tasmax – tasmin`) and adjust this variable, in addition to `tasmax` and `pr`. `tasmin` is reconstructed after the bias-adjustment and this is the variable we store.

While `tasmax` has no physical bounds in practice, this is not the case for `pr` and `dtr` where a lower bound, at 0, exists. Because of this, the adjustment process explained below exists in two flavours : additive and multiplicative. In the latter, it is mathematically impossible for adjusted data to go under 0.

2 Bias-adjustment

The bias adjustment acts independently on each day of the year and each grid point. To make the procedure more robust a window of 31 days around the current day of year is included in the inputs of the calibration (training step). For example, the adjustment for February 1 (day 32) is calibrated using data from January 15 to February 15, over the 30 years of the reference period. For leap years, this would mean that there are 4 times fewer datapoints for the 366th day of the year. To circumvent this issue, we convert all inputs to a "noleap" calendar by dropping data on the 29th of February, except for simulations using the "360 day" calendar. In the latter case, the simulations are untouched but the reference data is converted to that calendar by dropping extra days taken at regular intervals[1]

2.1 Detrending

For each day of the year and each grid points, we first compute the averages and "anomalies" of the reference data and the simulations over the reference period, 1981-2010. Depending on the variable, anomalies are either taken additively or multiplicatively:

$$Y_r = \begin{cases} \overline{Y_r} + Y'_r & \texttt{tasmax} \\ \overline{Y_r} \cdot Y'_r & \texttt{dtr, pr} \end{cases} \tag{1}$$

and similarly for X_c, $\overline{X_c}$ and X'_c.

Instead of a simple moving mean, X_s is detrended with a locally weighted regression (LOESS; [3]). We chose this method for its slightly heavier weights given at the center of the moving window, reducing impacts of abrupt interannual changes on the trend and anomalies. It also has a more robust handling of the extremeties of the

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[1]On a normal year, February 6th, April 20th, July 2nd, September 13th and November 25th are dropped. For a leap year, it is January 31st, April 1st, June 1st, August 1st, September 31st and December 1st.

timeseries. The LOESS window had a 30-year width and a tricube shape, the local regression was of degree 0 and only one iteration was performed. The detrending was applied on each day of the year but after averaging over the 31-day window and it yielded the trend $\overline{X_s}$ and the residuals X'_s. Here again, the process can be additive or multiplicative.

2.2 Adjustment of the residuals

With $F_{Y'_r}$ and $F_{X'_c}$ the empirical cumulative distribution functions (CDF) of Y'_r and X'_c respectively, an adjustment factor function is first computed:

$$A_+(q) := F_{Y'_r}^{-1}(q) - F_{X'_c}^{-1}(q) \qquad\qquad A_\times(q) := \frac{F_{Y'_r}^{-1}(q)}{F_{X'_c}^{-1}(q)} \qquad (2)$$

Where q is a quantile (in range $[0, 1]$), $A_+(q)$ is the additive function used with `tasmax` and $A_\times(q)$ the multiplicative one, used with `pr` and `dtr`. The CDFs are estimated from the 30 31-day windows. In the implementation, maps of A are saved to disk by sampling q with 50 values, going from 0.01 to 0.99 by steps of 0.02. The adjustment is then as follows:

$$X'_{ba} = X'_s + A_+\left(F_{X'_c}(X'_s)\right) \qquad\qquad X'_{ba} = X'_s \cdot A_\times\left(F_{X'_c}(X'_s)\right) \qquad (3)$$

Nearest neighbor interpolation is used to map $F_{X'_c}(X'_s)$ to the 50 values of q. Constant extrapolation is used for values of X'_s outside the range of X'_c.

2.3 Adjustment of the trend

In the training step, as simple scaling or offset factor is computed from the averages:

$$C_+ = \overline{Y_r} - \overline{X_c} \qquad\qquad C_\times = \frac{\overline{Y_r}}{\overline{X_c}} \qquad (4)$$

This factor is applied to the trend in the adjustment step:

$$\overline{X_{ba}} = \overline{X_s} + C_+ \qquad\qquad \overline{X_{ba}} = \overline{X_s} \cdot C_\times \qquad (5)$$

2.4 Final scenario

Finally, the bias-adjusted timeseries for this day of year, grid point and variable is:

$$X_{ba} = \overline{X_{ba}} + X'_{ba} \qquad (6)$$

3 Pre-processing of precipitation

However, the multiplicative mode is prone to division by zero, especially with precipitation where values of 0 are quite common. This problem is avoided by modifying the inputs of the calibration step where the zeros of precipitation are replaced by random values between 0 (excluded) and 0.01 mm/d. The `dtr` timeseries are not modified since it is almost impossible to have zeros for that variable and the few that appear are dissolved by the aggregations of the calibration step.

As observed by [5], when the model has an higher dry-day frequency than the reference, the calibration step of the quantile mapping adjustment will incorrectly map all dry days to precipitation days, resulting in a wet bias. The frequency adaptation method finds the fraction of "extra" dry days:

$$\Delta P_{dry} = \frac{F_{X_c}(D) - F_{Y_r}(D)}{F_{X_c}} \qquad (7)$$

Where D is the dry-day threshold, taken here to be 1 mm/d. This fraction of dry days is transformed into wet days by injecting random values taken in the interval $[D, F_{Y_r}^{-1}(F_{X_c}(D))]$.

Both pre-processing function are applied only on the calibration step inputs (Y_r and X_c) before the division between average and anomalies. As such, only the adjustment factors are impacted by them and there is no explicitly injected precipitation in the final scenarios.

References

[1] Agbazo, M. N., & Grenier, P. (2020). Characterizing and avoiding physical inconsistency generated by the application of univariate quantile mapping on daily minimum and maximum temperatures over Hudson Bay. International Journal of Climatology, 40(8), 3868–3884. https://doi.org/10.1002/joc.6432

[2] Cannon, A. J., Sobie, S. R., & Murdock, T. Q. (2015). Bias Correction of GCM Precipitation by Quantile Mapping: How Well Do Methods Preserve Changes in Quantiles and Extremes? Journal of Climate, 28(17), 6938–6959. https://doi.org/10.1175/JCLI-D-14-00754.1

[3] Cleveland, W. S. (1979). Robust Locally Weighted Regression and Smoothing Scatterplots. Journal of the American Statistical Association, 74(368), 829–836. https://doi.org/10.1080/01621459.1979.10481038

[4] Thrasher, B., Maurer, E. P., McKellar, C., & Duffy, P. B. (2012). Technical Note: Bias correcting climate model simulated daily temperature extremes with quantile mapping. Hydrology and Earth System Sciences, 16(9), 3309–3314. https://doi.org/10.5194/hess-16-3309-2012

[5] Themeßl, M. J., Gobiet, A., & Heinrich, G. (2012). Empirical-statistical downscaling and error correction of regional climate models and its impact on the climate change signal. Climatic Change, 112(2), 449–468. https://doi.org/10.1007/s10584-011-0224-4