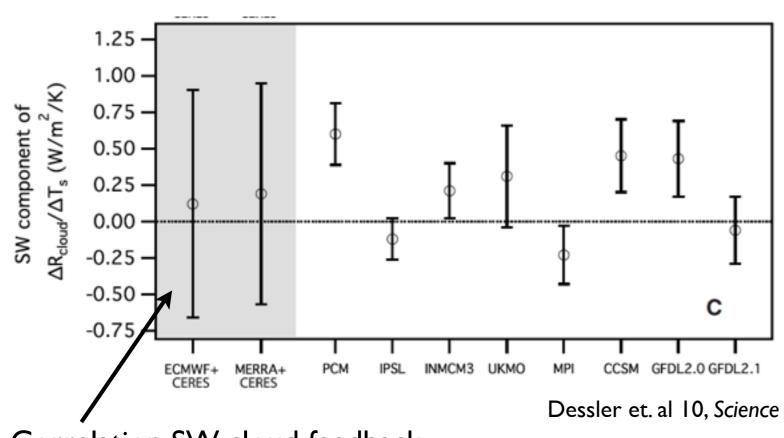
Observational evidence for positive lowcloud feedback and constraints on climate sensitivity

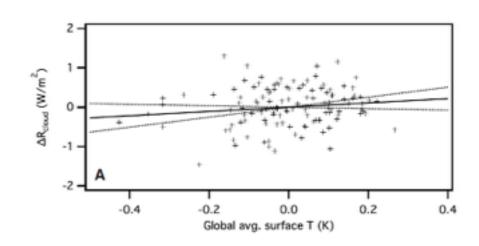
Florent Brient and Tapio Schneider

June 8th, 2015



What did we learn from observations?





Slight positive SW cloud feedback

Strong uncertainty

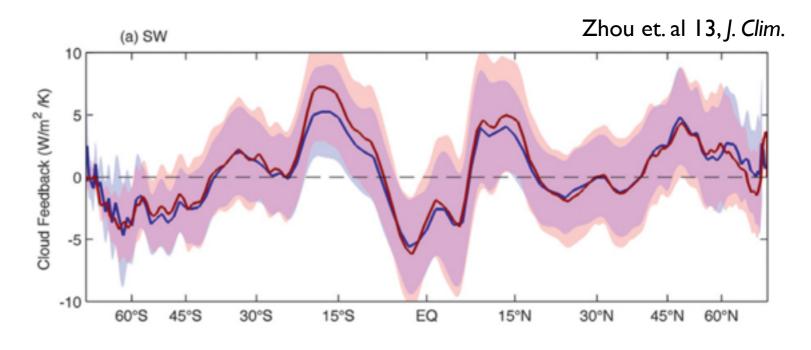
Strong disagreement across models

Correlation SW cloud feedback versus Ts (global)

Strong regional discrepancy

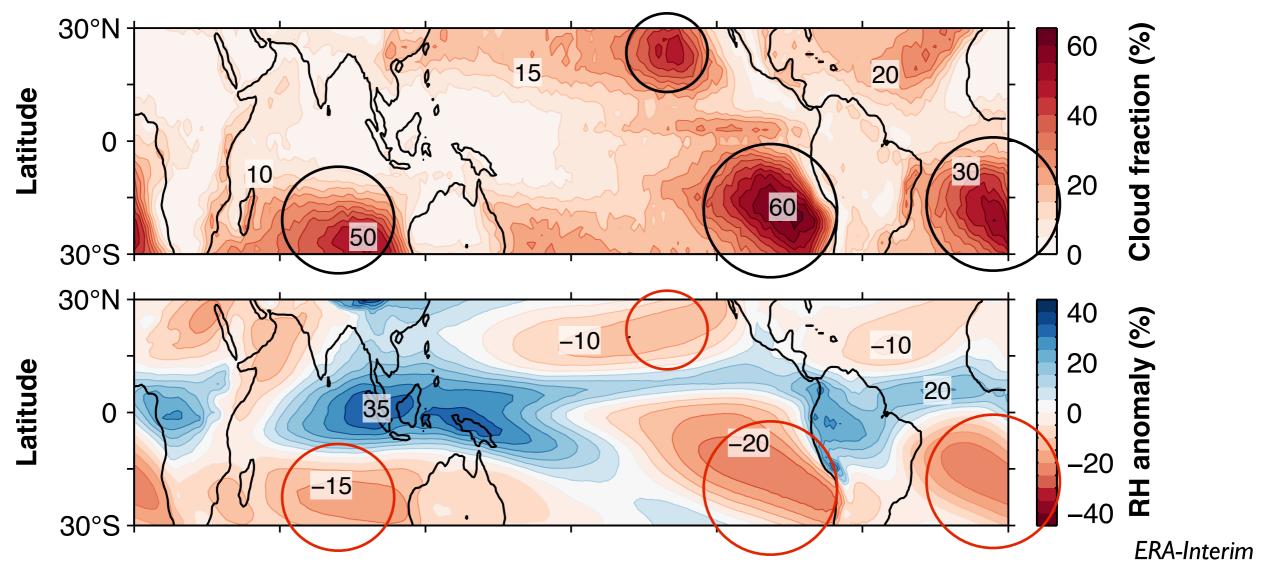
Positive feedback over subtropical regions

Is it possible to constrain cloud feedback based on observed variations?



Focus on low-clouds

Unique Low-Cloud fraction (Calipso/CloudSat)



Low-clouds are present over subtropical regions (eastern parts of oceans)

Drier than the tropical mean

Lower quartile of RH deviation from tropical-ocean averaged RH

Focus on low-clouds

Unique Low-Cloud fraction (Calipso/CloudSat) 30°N Cloud fraction (%) 60 15 Latitude 40 0 30 20 60 50 30°S 30°N 40 RH anomaly (%) -10 -10 20 Latitude 20 35, 0 -20 **-20** -15 **-40** 30°S drier lower quartile, **ERA-Interim** 100 30°N 80 fred occ (%) Latitude 60 0 40 20 30°S 120°W 120°E 0 Frequency of occurrence Longitude

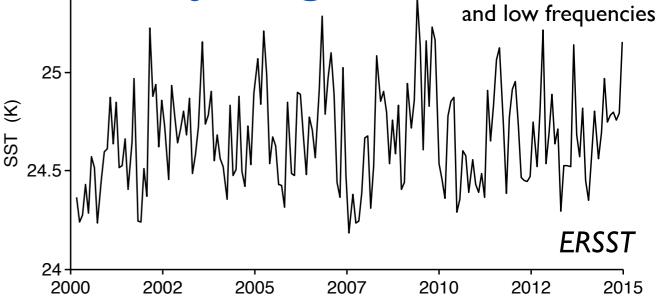
Temporal evolution of dry region

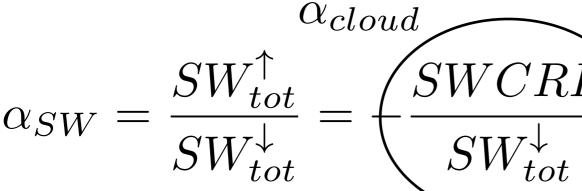
Two observed dataset

Earth SW albedo:

CERES-EBAF (2000-2015)

ISCCP-FD (1983-2009)?





$\frac{\left(SWCRE\right)}{SW_{tot}^{\downarrow}} \frac{\left(SW_{clear}^{\uparrow}\right)}{SW_{tot}^{\downarrow}}$

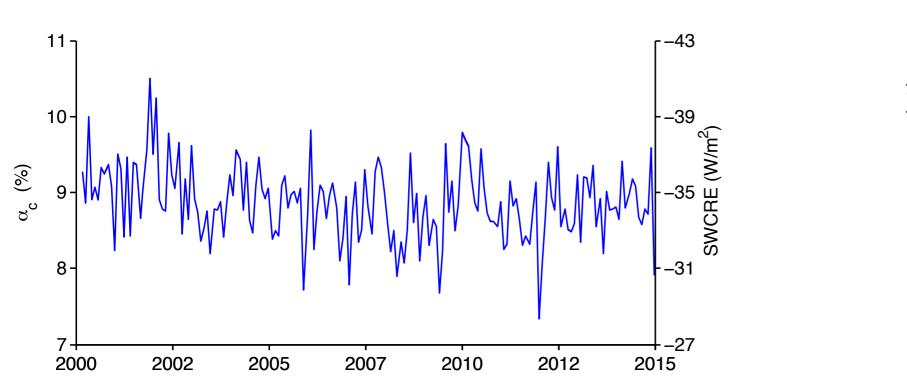
cloud feedback

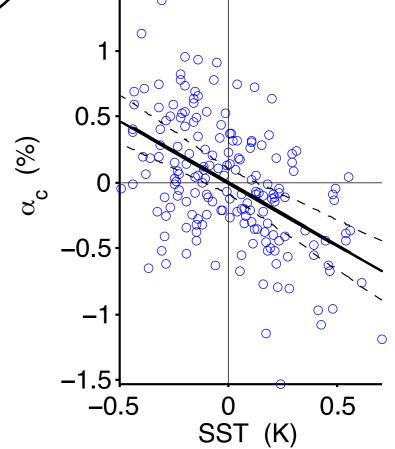
dominated by volcanoes

variabilities at high

 α_{clear}

1.5





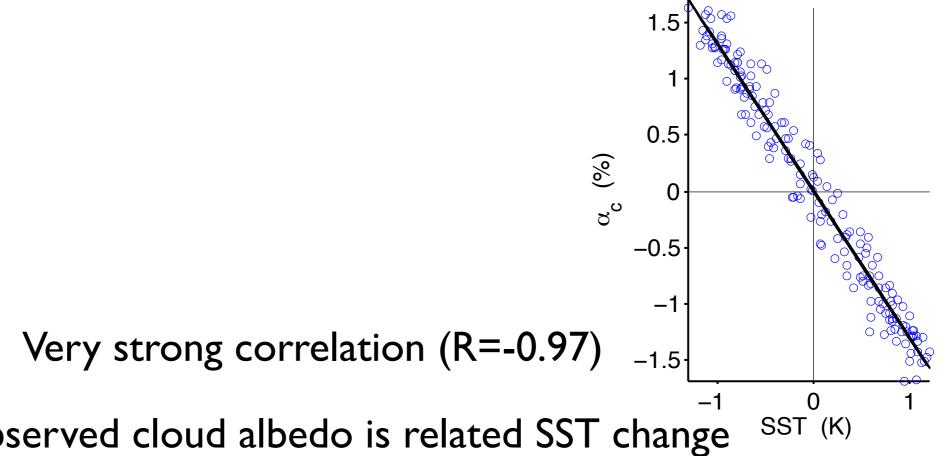
Spectral decomposition seasonal

Seasonal variability

I-yr bandpass filter with I2-order Chebyshev polynomial

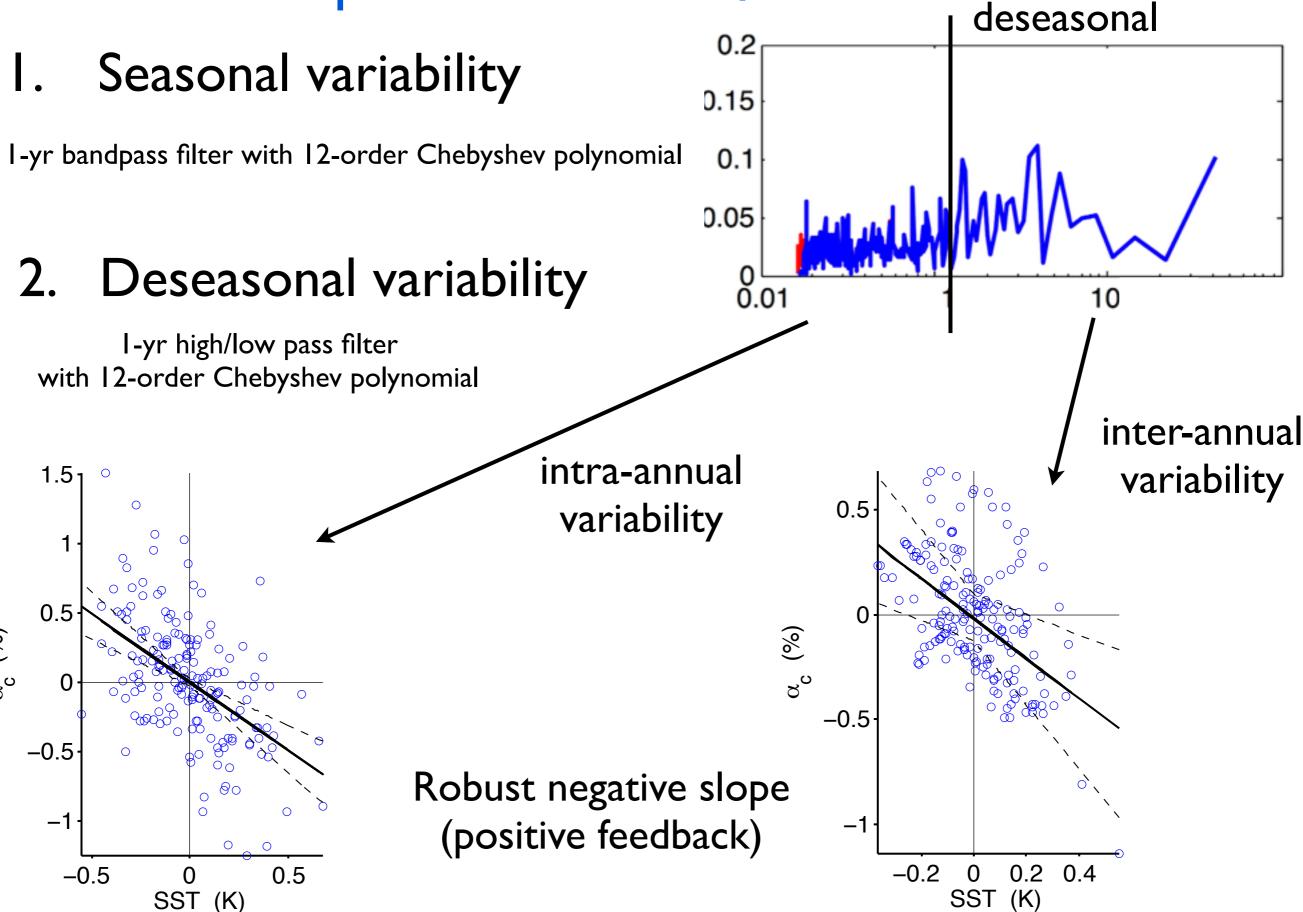
Ŏ.01 10

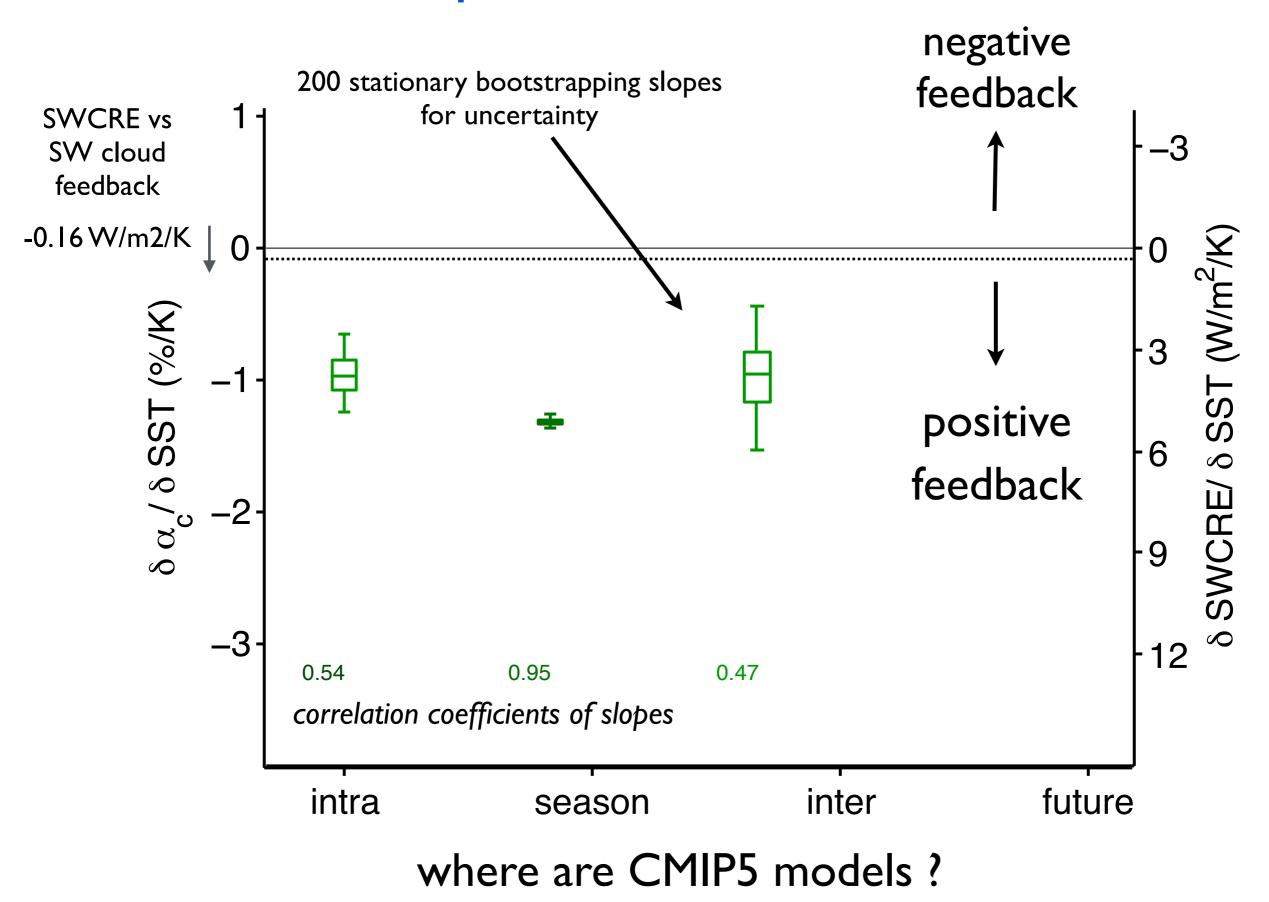
Bandpass filter

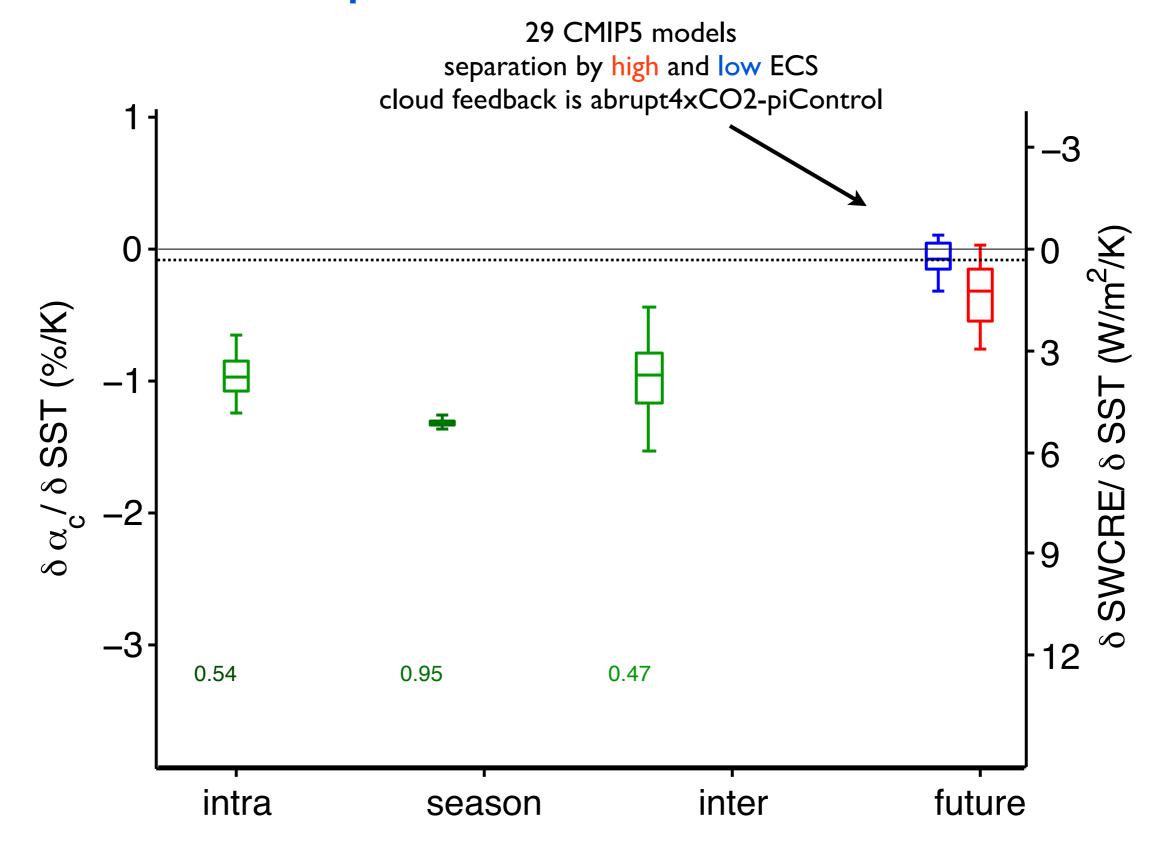


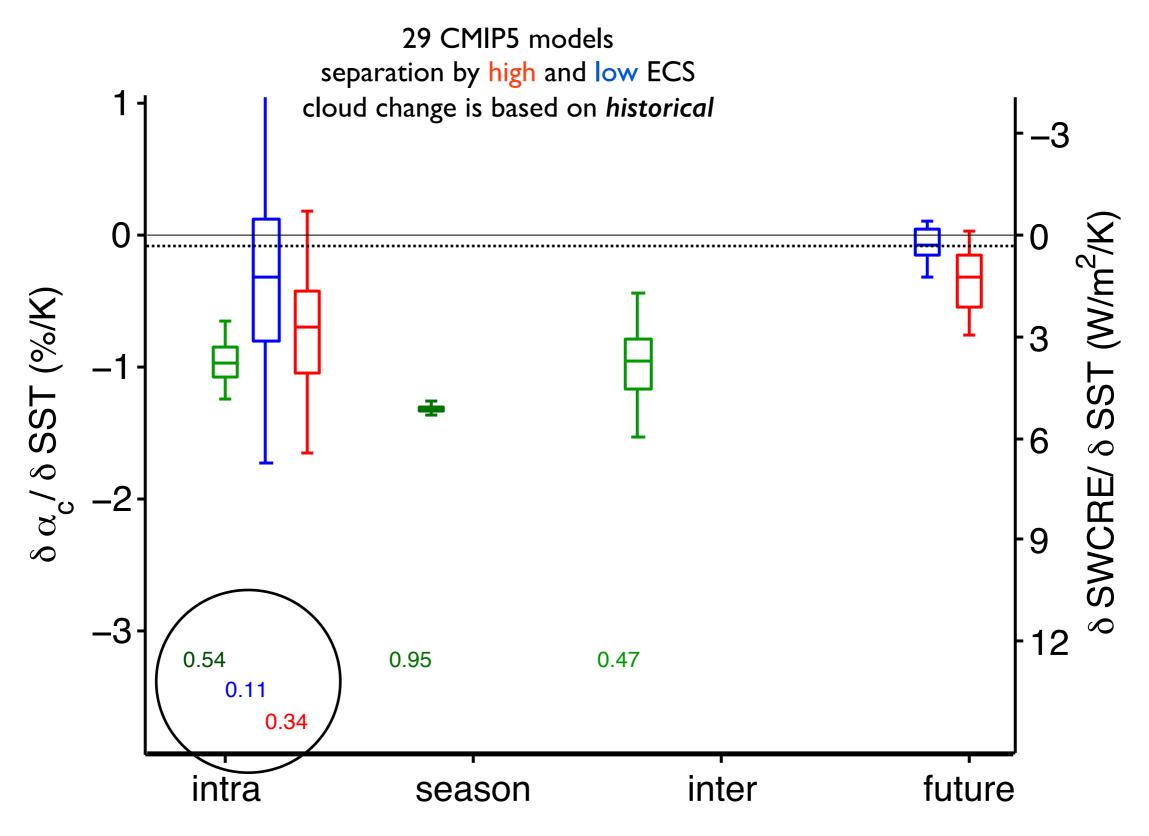
Observed cloud albedo is related SST change

Spectral decomposition

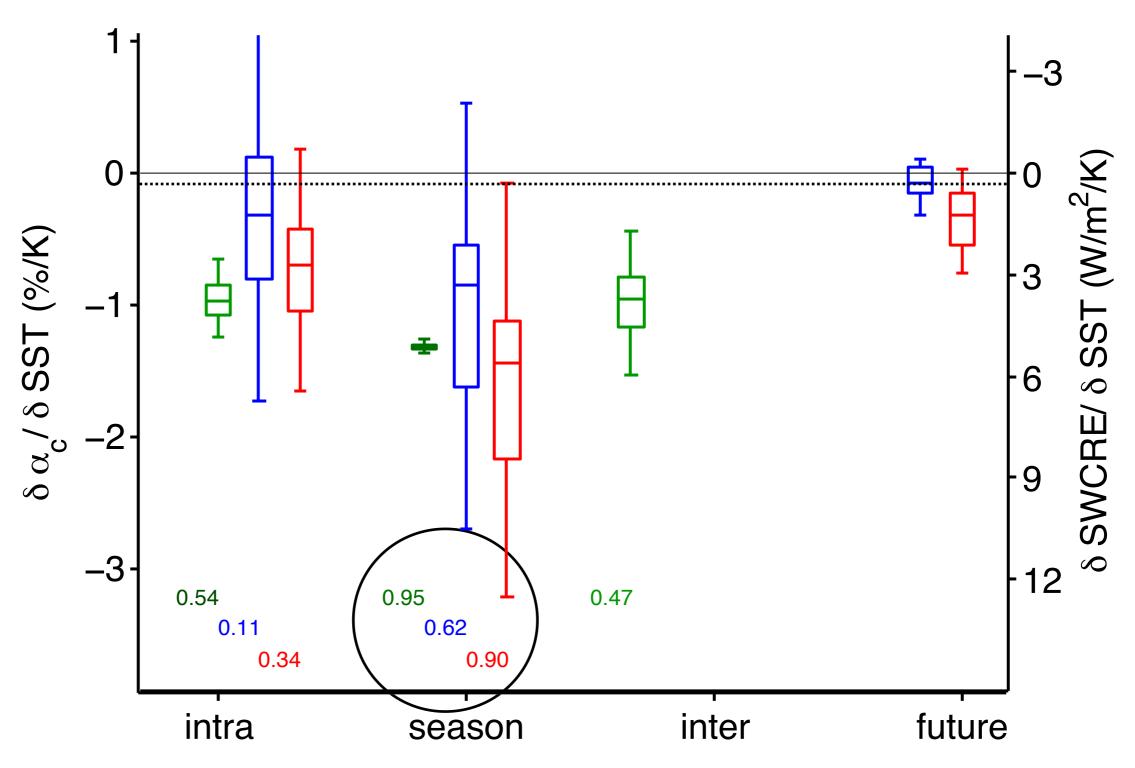




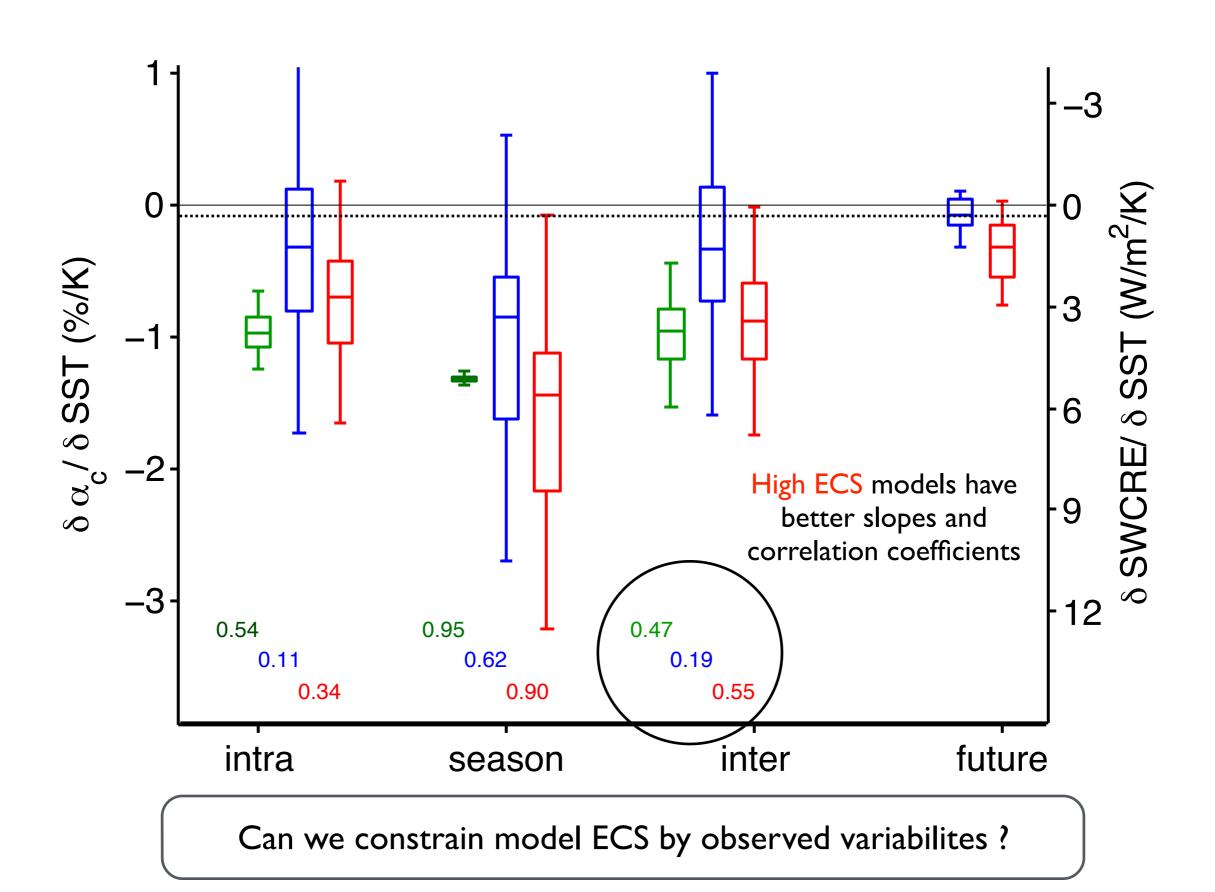


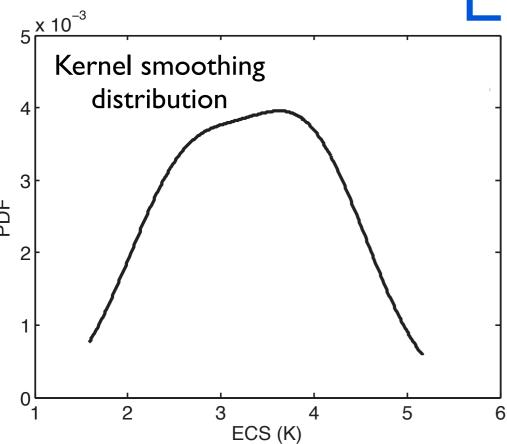


Models have difficulties to represent high-frequency variabilities (MJO?)

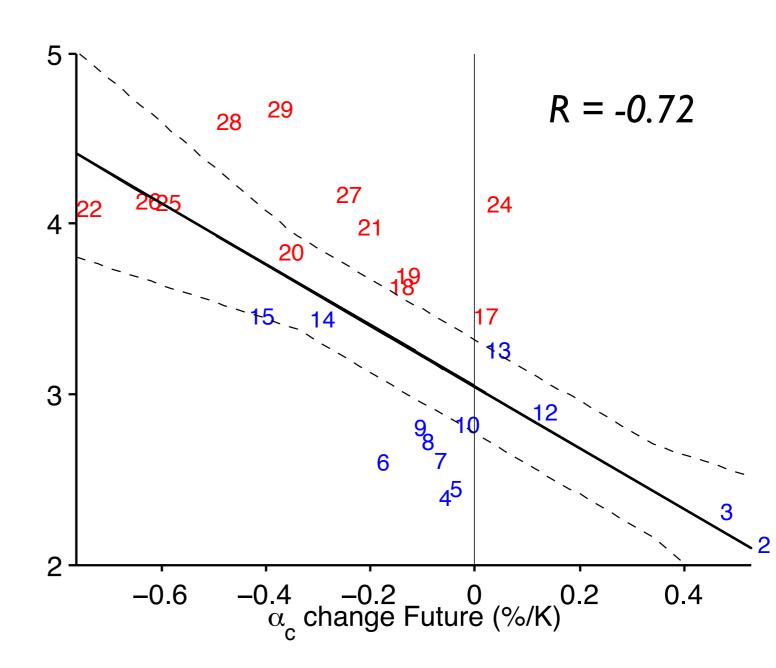


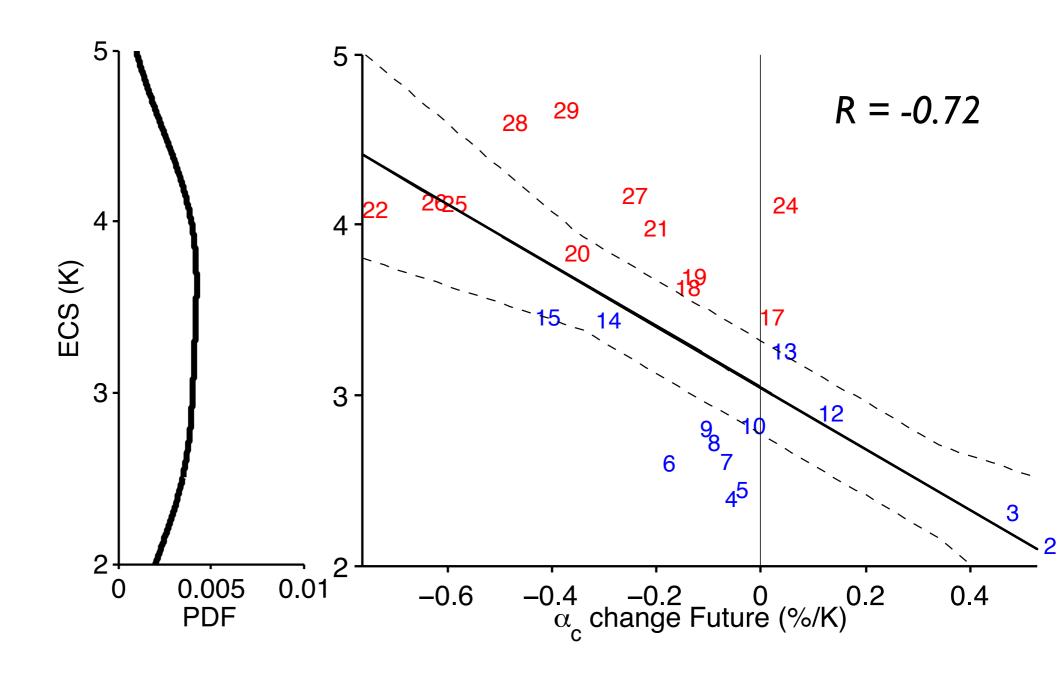
High ECS models have better slope and correlation coefficients

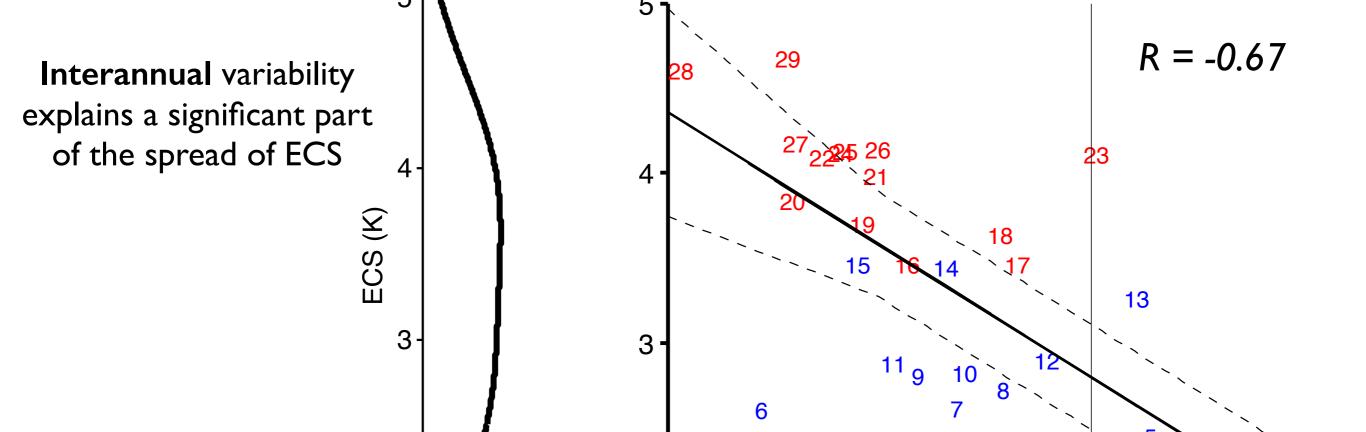




CMIP5 ECS range is [2.1, 4.6] K, with a max probability of 3.6 K







0.01 -1.5

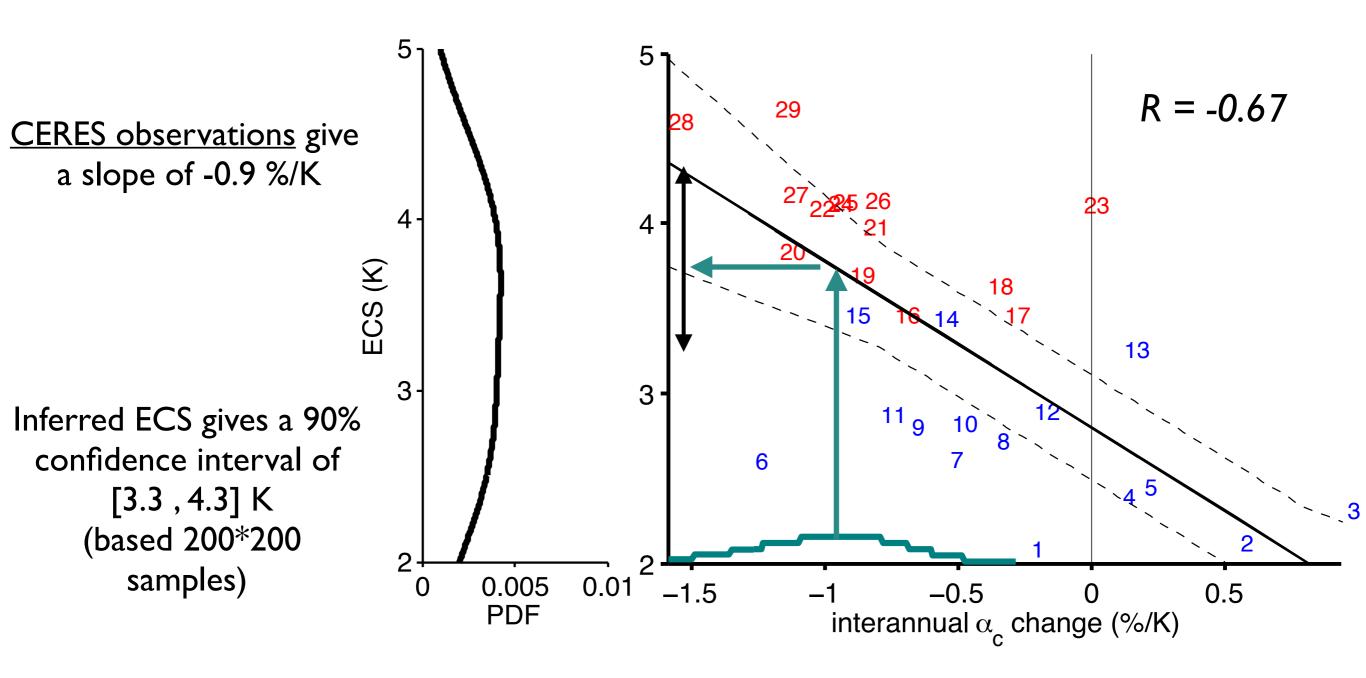
-0.5

interannual $\alpha_{\rm c}$ change (%/K)

0.5

0.005 PDF

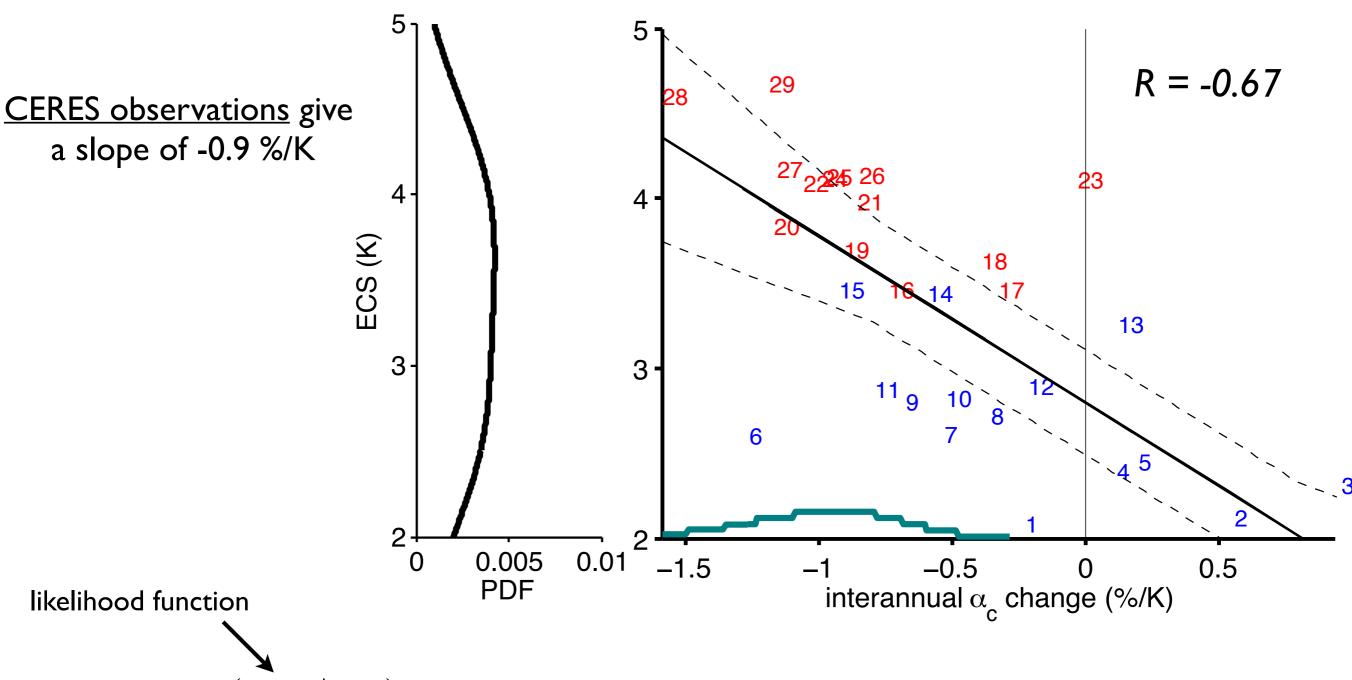
0



Too strong influence of lowest sensitivity models (by their impact on the slope)

Too weak influence of some realistic models

Using Bayesian Model Averaging

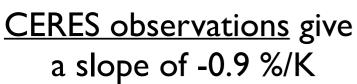


$$w_i = \frac{P(\alpha_{obs}|M_i)}{\Sigma_k P(\alpha_{obs}|M_k)}$$
 Proba of the model given the obs

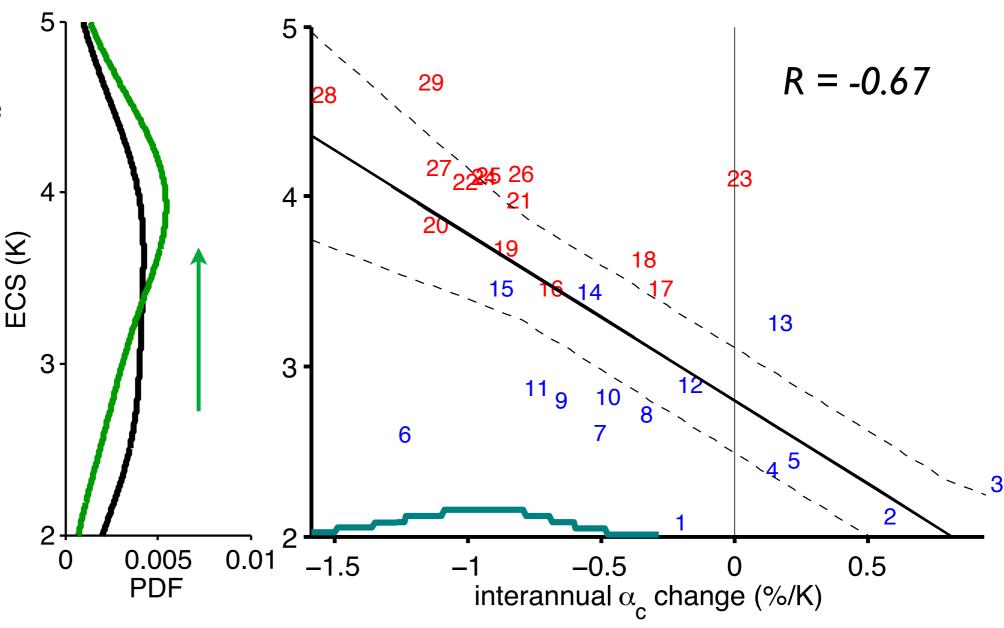
$$P(ECS|\alpha_{obs}) = \Sigma_k \ w_k \ P(ECS|M_k)$$

Proba of ECS given the obs

Strengthen the weight of realistic models



BMA methodology gives a 90% confidence interval of [2.4, 4.9], with a max probability of 3.9K



likelihood function

$$w_i = \frac{P(\alpha_{obs}|M_i)}{\Sigma_k P(\alpha_{obs}|M_k)}$$
 Proba of the model given the obs

$$P(ECS|\alpha_{obs}) = \Sigma_k \ w_k \ P(ECS|M_k)$$

Proba of ECS given the obs

Strengthen the weight of realistic models

Conclusions

Robust positive low-cloud feedback across temporal variability

Seasonal variability of cloud albedo change is related to surface warming misrepresentation by CMIP5 models and possible constrain to improve models

Spread of ECS by CMIP5 models is 2.1-4.5K with a most probable value of 3.3K

ECS lower bounds increases from 2.1K to 2.4 K Most probable ECS value is higher (3.9K)

Need process-oriented analysis of observed low-cloud change (especially by their vertical development)