

## Graphical Abstract

### **Collective sedimentation from ash-clouds: insights from experimental buoyant, particle-laden gravity currents**

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## Highlights

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- Research highlight 1
- Research highlight 2

# Collective sedimentation from ash-clouds: insights from experimental buoyant, particle-laden gravity currents

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## Abstract

*Keywords:*

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## 1. Introduction

The volcanic ash produced by explosive volcanic eruptions is an economic and societal hazard. Fine ash which is breathed in can cause respiratory health problems in humans and animals (Kampa and Castanas, 2007; Anderson et al., 2012; WHO, 2013; Baxter and Horwell, 2015). Agriculturally, ash deposits can contaminate human and livestock food and water supplies, as well as cover crops (Cook et al., 1981; Cronin et al., 1998; Wilson et al., 2011; Craig et al., 2016). More generally, ash can cause buildings to collapse and damage electrical power systems (Wilson et al., 2014; Craig et al., 2016). A well-publicised impact is that on the aviation industry, whereby the risk of

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damage to airplane engines can lead to costly shutdowns of airspace (Budd et al., 2011; Elisondo et al., 2016). The mechanisms of ash dispersal and sedimentation therefore need to be understood in order for accurate hazard and risk assessments to be carried out.

Key tools used for hazard assessment are ash dispersal models which aim to simulate the transport of volcanic ash in the atmosphere (Witham et al., 2007; Bonadonna et al., 2012; Folch, 2012). These have been used to produce operational forecasts of ash clouds (Scollo et al., 2009; Webster et al., 2012) or to produce hazard maps which can be used in decision making (Bonadonna et al., 2005; Macedonio et al., 2005; Folch and Sulpizio, 2010). These models rely on accurate parameterisations of ash sedimentation processes. However, it is often assumed that ash settles at its terminal fall velocity (Hazen, 1904), as determined from its size and the density contrast with the surrounding atmosphere (Clift and Gauvin, 1971; Ganser, 1993). Such a scheme predicts that the model grain size of deposited ash should decrease monotonically with distance from the vent (Bursik et al., 1992; Sparks et al., 1992). Deposits, however, sometimes record more complicated features such as bimodal size distributions whilst remotely sensed GSDs of the ash plume from the 2010 eruption of Eyjafjallajökull also showed that the effective ash radius did not monotonically decrease with distance from the vent (Bonadonna et al., 2011). Indeed, it is increasingly apparent that collective sedimentation processes can strongly control the sedimentation of ash. One commonly cited mechanism is ash aggregation, whereby ash particles stick together, increasing their effective size and thus their fall velocity (Carey and Sigurdsson, 1982; Sorem, 1982; Lane et al., 1993; Bonadonna et al., 2011; Brown et al., 2012). Another

possibility though is that collective sedimentation can occur through the occurrence of convective instabilities which drive larger scale fluid motions that can drive ash sedimentation through downward-propagating finger structures (Bonadonna et al., 2002, 2011; Carazzo and Jellinek, 2012; Manzella et al., 2015; Scollo et al., 2017). These arise due to the formation of a gravitationally unstable interfacial layer at the base of ash clouds, which has previously been called the particle boundary layer or PBL (Carazzo and Jellinek, 2012).

Despite recent progress in understanding how convective instabilities form and evolve, there remains a number of key unanswered questions. In particular, no consideration has been given to the role of background velocity shear, which will be present due to buoyant spreading of the cloud and/or wind drag (Johnson et al., 2015). In this study, we explore how the presence of shear at the base of volcanic clouds, due to buoyant spreading, can influence collective settling behaviour. In particular, we show that there is a regime where Kelvin-Helmholtz shear instabilities (von Helmholtz, 1868; Kelvin, 1871) can interact with settling-driven gravitational instabilities (Hoyal et al., 1999; Blanchette and Bush, 2005) to produce larger fingers that can enhance the sedimentation of fine ash.

### *1.1. Settling-driven gravitational instabilities*

## **2. Methods**

Experiments were performed in a perspex flume of internal length 353 cm, width  $(12.2 \pm 0.5)$  cm and depth 50 cm (Figure 1). The uncertainty on the width is due to bowing of the tank walls. During the experimental setup, two removable gates can be placed at 24 and 53 cm, respectively, from

the left-hand end, creating three sections. The left-most section takes no part in the experiment, whilst the short middle section is where the particle suspension is prepared, and is referred to as the gated section (length 27 cm). The remaining length of the flume is called the environment (length 3m).

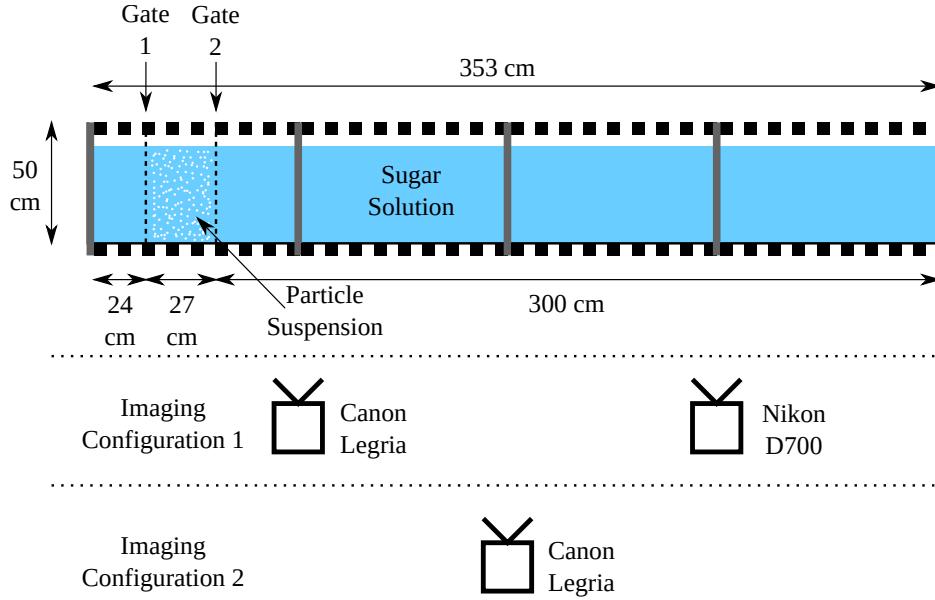


Figure 1: Sketch showing the experimental setup. The flume is separated into three section separated by two gates. The leftmost section is not involved in the experiment. The rightmost section is a sugar solution of constant density whereas the section between gates 1 and 2 is a mixture of fresh water and ballotini. The concentration of particles is varied between experiments. The experiment is initiated by removing gate 2. Experiments were imaged using one of two configurations.

The flume sits within a metal framework. Three vertical supports, each 3cm thick, at distances of 87 cm, 173.5 cm and 260 cm from the left-hand end prevent bowing. This effectively separates the flume into four, nearly equal, section. Behind each section a backing board is placed. For experiments with no particles, red or blue food colouring is added to the current and

the backing board is white. Otherwise the boards are black. The top 5 cm of each board is a row of  $(5 \times 5)$  cm<sup>2</sup> squares, alternating in colour from black-to-white. Meanwhile, at the base of the tank, tape is used to create a similar scale.

The day before an experiment, the flume is filled up to a depth of 47.5 cm. Separately, the desired mass of sugar is completely dissolved in approximately 15 l of water. The flume and sugar solution are then both allowed to equilibrate to room temperature overnight. The next day, the flume is imaged (single frame) in its current configuration. In some experiments, the whole flume was imaged using a Canon Legria HFG40. In others, this camera was position closer to the tank, but only imaged the left-hand half of the flume, whilst the right-hand half was imaged using a Nikon D700 (DSLR). The captured frame(s) are used as reference images using the top(back) and bottom(front) scales. By capturing the image with the back scale partially submerged, it is possible to correct position measurements for distortion due to the refractive index (RI) of the water.

The flume water level is then lowered until it is of a depth of approximately 36.5 cm. Gates 1 and 2 are then put in place. The sugar solution is then added to the environment section. The gated section is then topped up with water until the water depth there is the same as the environment. This procedure results in a water depth of approximately 40 cm. The environment section is rigorously stirred. The RI of fluid from four positions in the environment section (top and bottom, near and far from the gate) is measured using a refractometer to ensure a uniform density. The RI of the gated section is also measured to check that there has been no significant

leakage of sugar solution through the gate. A calibrated digital thermometer is used to measure the temperature of both sections. In all experiments, the maximum difference in temperature between the sections was 0.1 °C.

Recording of the experiment then begins. The required mass of particles is added to the gated section which is thoroughly stirred to ensure a uniform particle concentration. Finally, gate 2 is removed and the particle suspension spreads along the free surface as a buoyant gravity current. Once the current head reaches the end of the tank, gate 2 is returned to its position and recording stops.

The particles used were glass spheres with a density of  $(2.519 \pm ??)$  g cm<sup>-3</sup>, as measured by helium pycnometry using an Ultrapyc 1200e. They had a unimodal size distribution centred on a mode of 36 µm and a standard deviation of 12 µm, as determined from static light scattering using a Bettersizer S3 Plus. Figure 2 shows the measured size distribution of the particles.

### 3. Results

#### 3.1. Currents without particles

Figure 3 shows the evolution of the current in experiment GC3.

Figure 4 shows the evolution of the current in experiment GC9.

#### 3.2. Particle-bearing currents

Figure 5 shows the evolution of the current in experiment GC42.

Figure 6 shows the evolution of the current in experiment GC48.

Figure 7 shows the evolution of the current in experiment GC45.

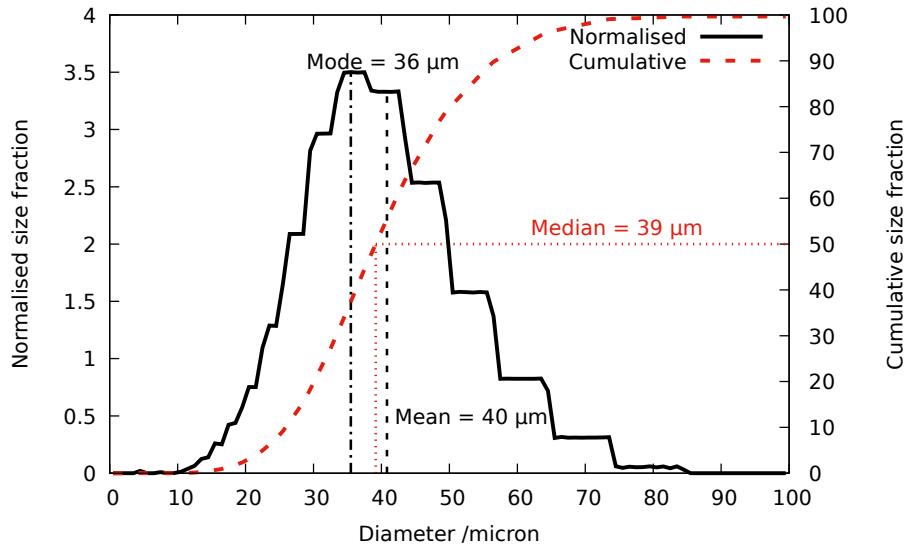


Figure 2: Normalised and cumulative volume-weighted size distributions of the particles used in the experiments.

#### 4. Discussion

#### 5. Conclusions

#### Acknowledgements

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Figure 3: Sequence of images showing experiment GC3 ( $\phi = 0$ ,  $g\tau = 0.162 \text{ m s}^{-2}$ ).

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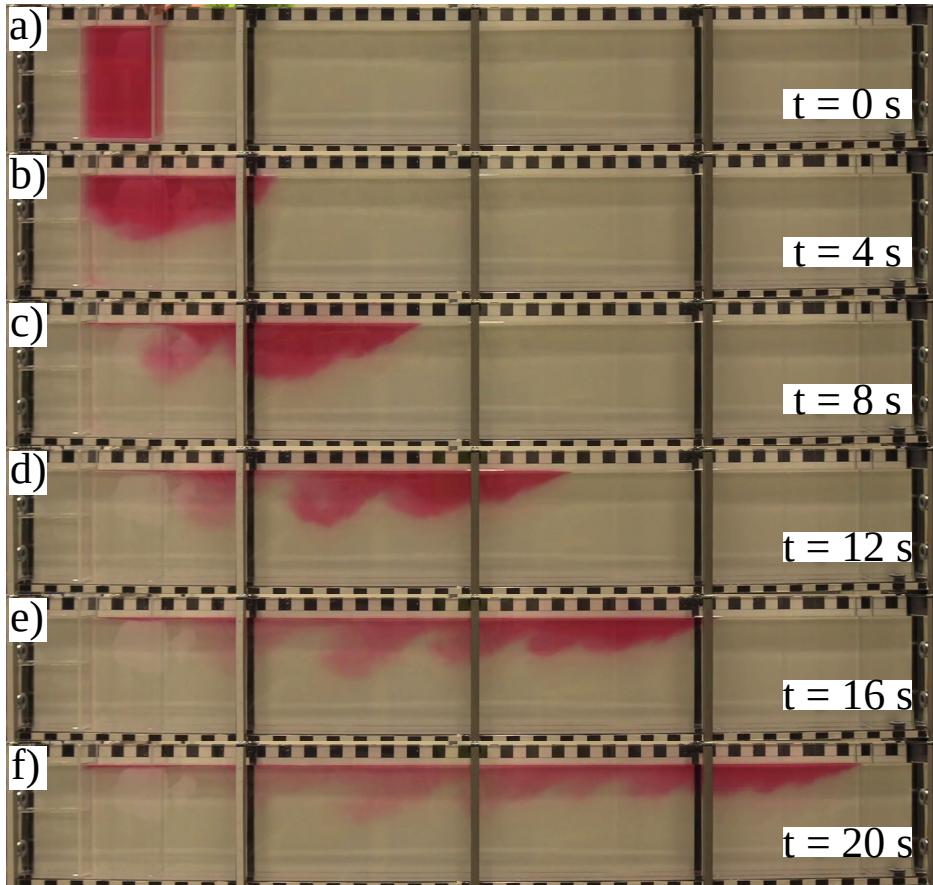


Figure 4: Sequence of images showing experiment GC9 ( $\phi = 0$ ,  $g' = 0.563 \text{ m s}^{-2}$ ).

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Figure 5: Sequence of images showing experiment GC42.

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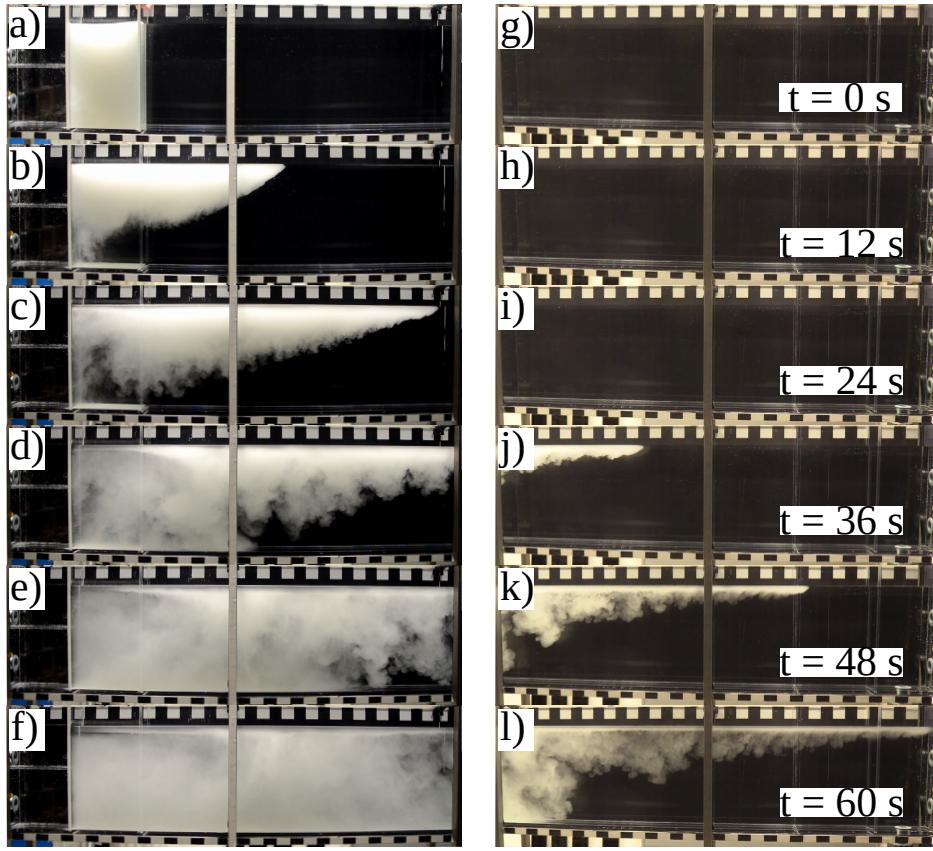


Figure 6: Sequence of images showing experiment GC48. a-f) show the left hand side of the tank whilst g-l) show the right hand side.

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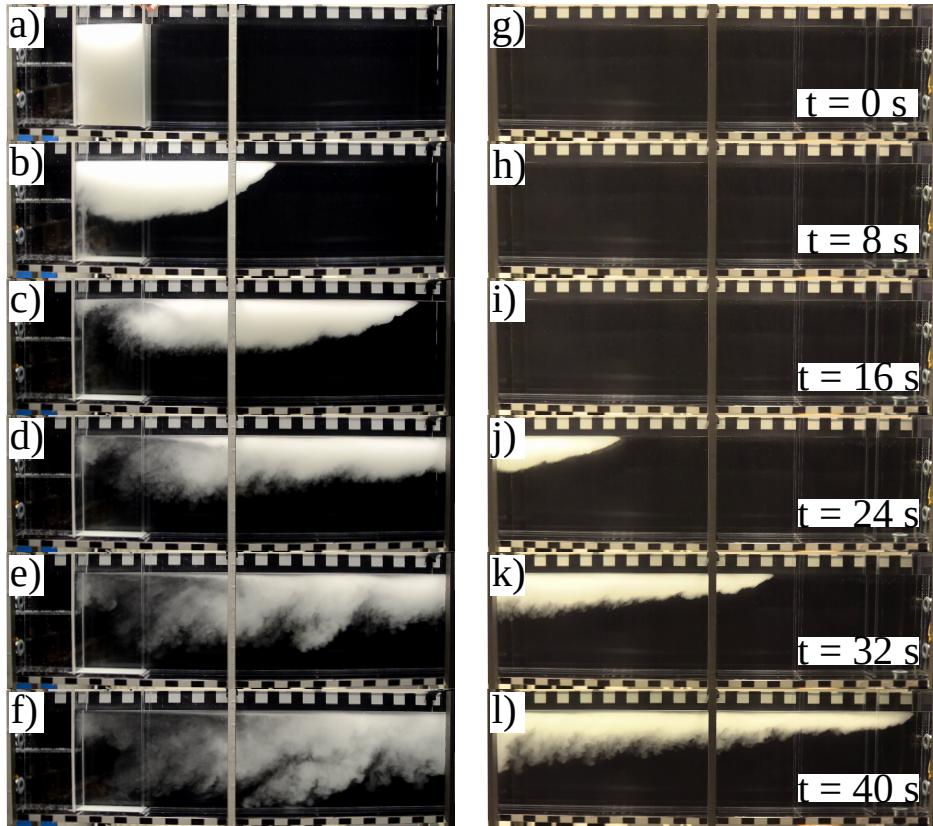


Figure 7: Sequence of images showing experiment GC45. a-f) show the left hand side of the tank whilst g-l) show the right hand side.

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