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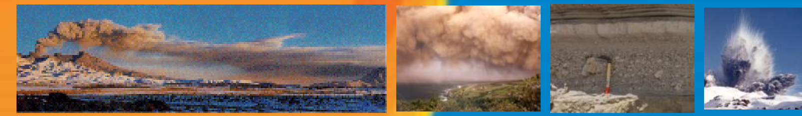
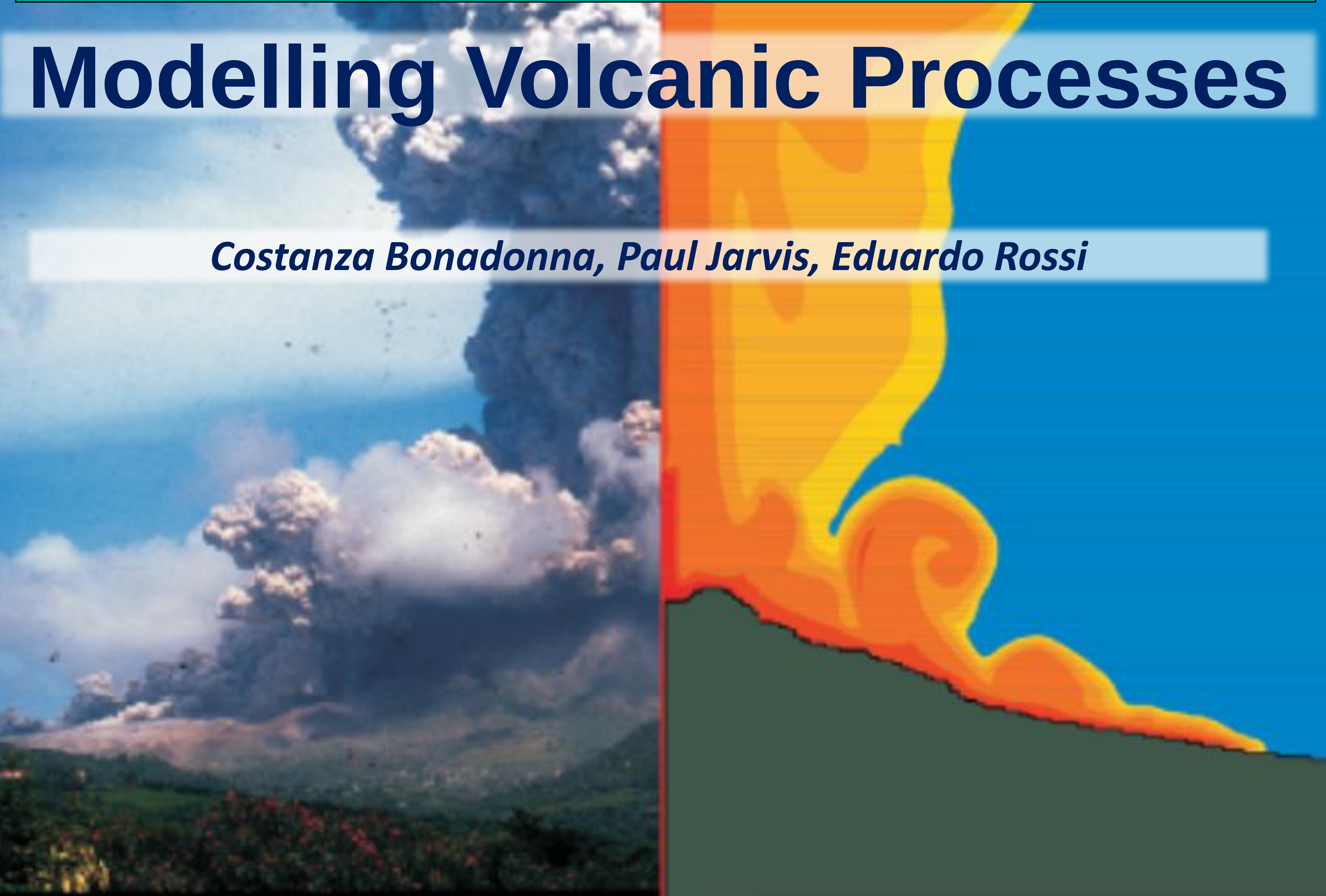


Modelling Volcanic Processes

Costanza Bonadonna, Paul Jarvis, Eduardo Rossi



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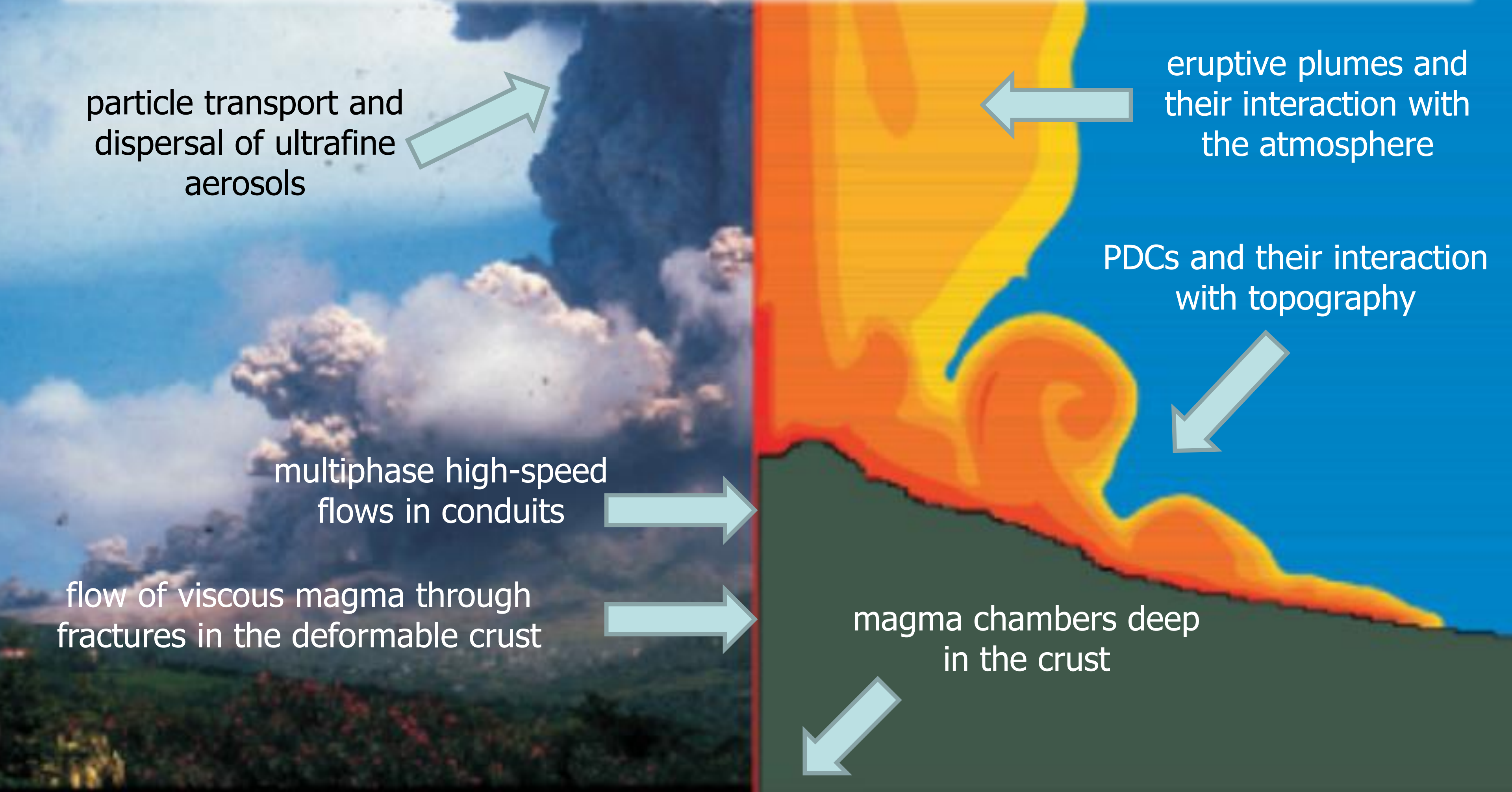
MOTIVATION

Volcanic system → range of scales, material property variations, and complex interacting physical and chemical processes



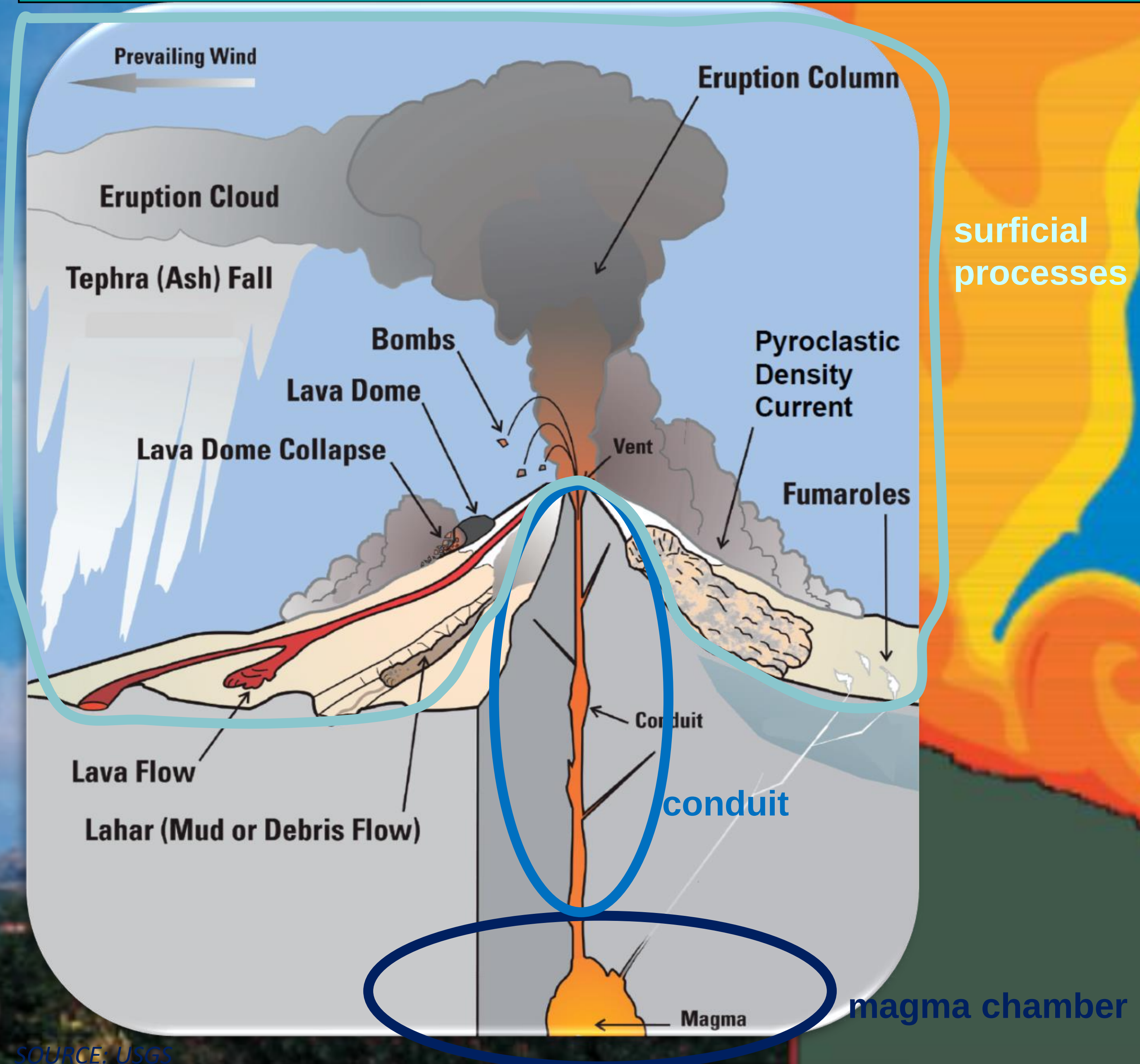
MOTIVATION

Volcanologists have the drive and the responsibility to progress their science to improve **understanding** and **mitigation** of the effects of volcanic eruptions

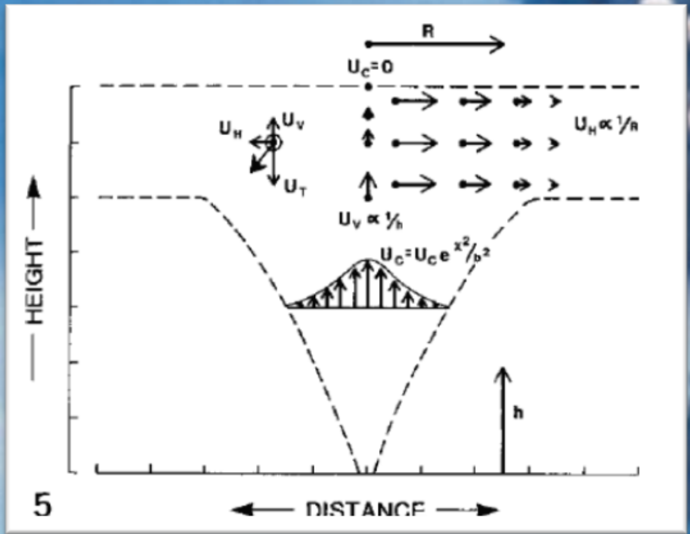
Many key volcanic processes cannot be observed and analysed directly

Hazardous processes are required to be analytically and numerically described for both real-time forecasting and long-term risk reduction strategies

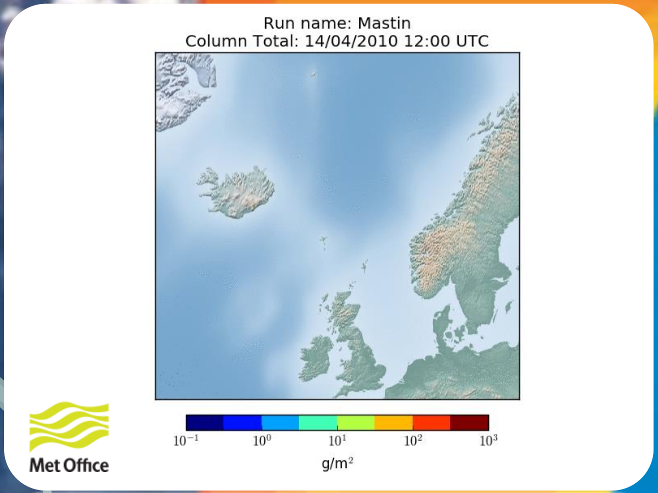
A variety of dedicated models of different complexity needed to be developed at multiple scales that could address different purposes


Prevailing Wind
Eruption Column
Eruption Cloud
Tephra (Ash) Fall
Bombs
Lava Dome
Lava Dome Collapse
Vent
Pyroclastic Density Current
Fumaroles
Conduit
Lava Flow
Lahar (Mud or Debris Flow)
Magma
surficial processes
conduit
magma chamber

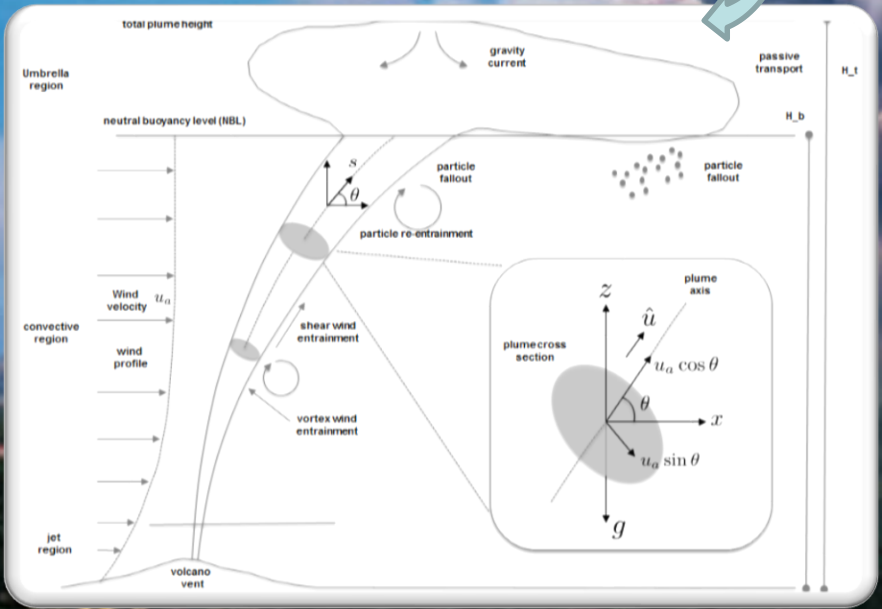
SOURCE: USGS

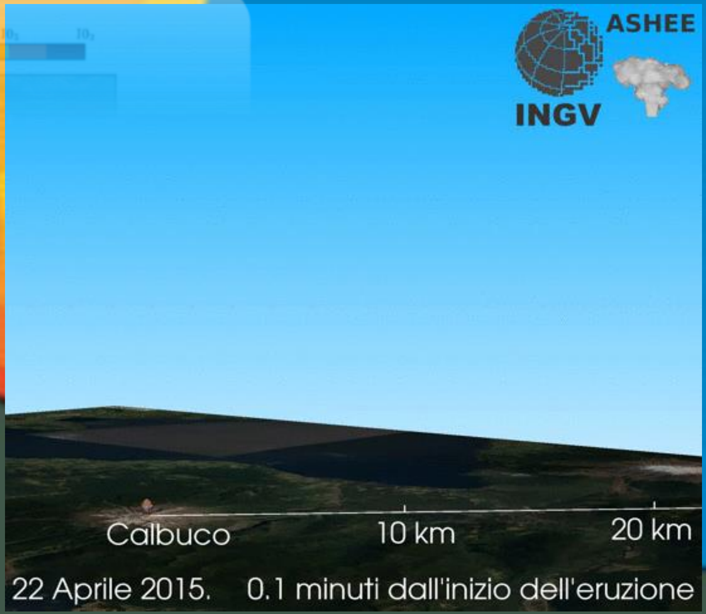


SOURCE: Carey and Spark 1986



SOURCE: Folch et al. 2011





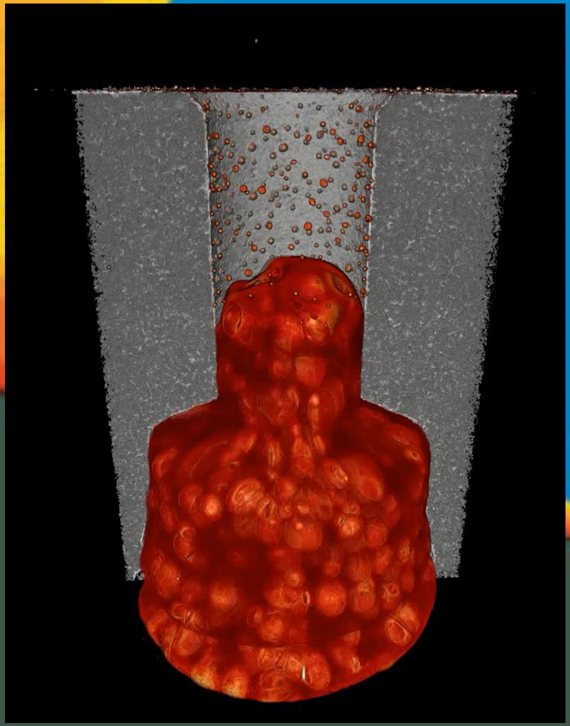
http://193.206.223.51:8081/software-list




Ash cloud spreading and sedimentation

Pyroclastic density currents

Vesiculation and expansion



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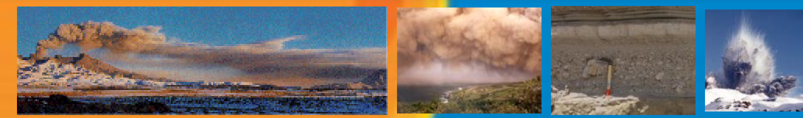
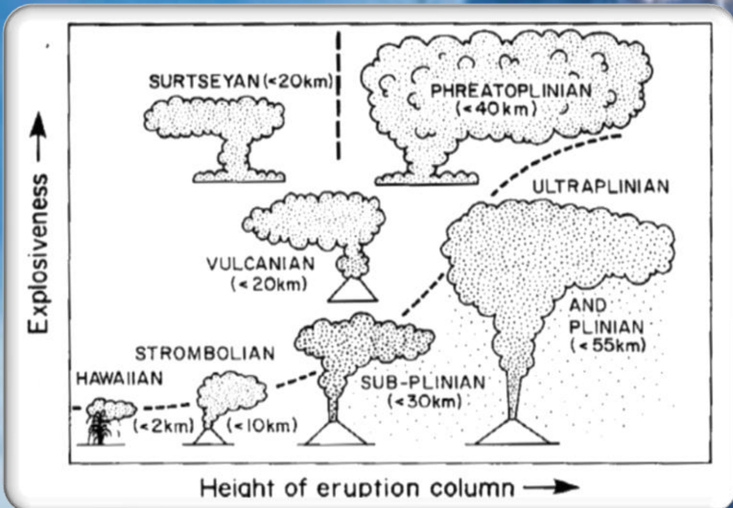
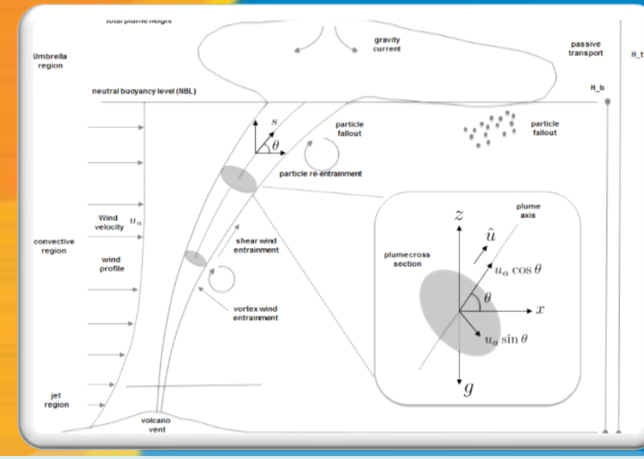




PART 1 – ERUPTION CHARACTERIZATION (Costanza BONADONNA) 12 November 2020

PART 2 – SUBSURFACE DYNAMICS (Paul JARVIS) 13 November 2020



PART 3 – MODELLING: Part I (Eduardo ROSSI) 26 November 2020

PART 4 – MODELLING: Part II (Paul JARVIS) 27 November 2020

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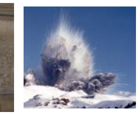
PHYSICAL CHARACTERIZATION OF EXPLOSIVE ERUPTIONS

Costanza Bonadonna
Costanza.Bonadonna@unige.ch



MAIN QUESTION:

How well can we constrain explosive eruptions?

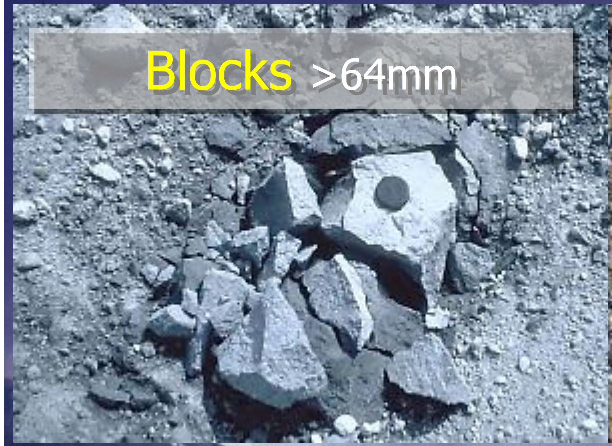
How well can we constrain Eruption Source Parameters (ESP)?

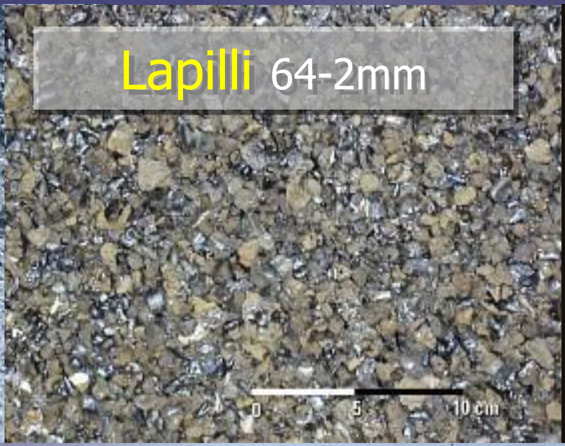
→ **Erupted Volume/Mass, Plume Height, Mass Eruption Rate (MER), Eruption Duration, Total Grain-Size Distribution (TGSD)**

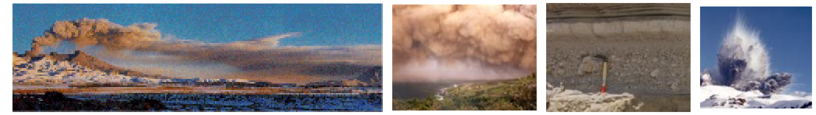



TEPHRA

collective term for airborne volcanic ejecta irrespective of size, composition or shape – *Thorarinsson 1944*







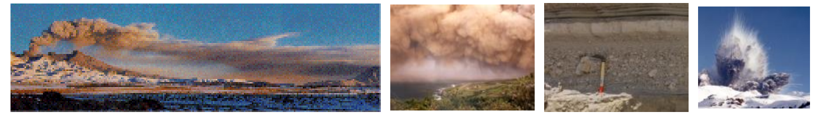
TEPHRA

reticulite

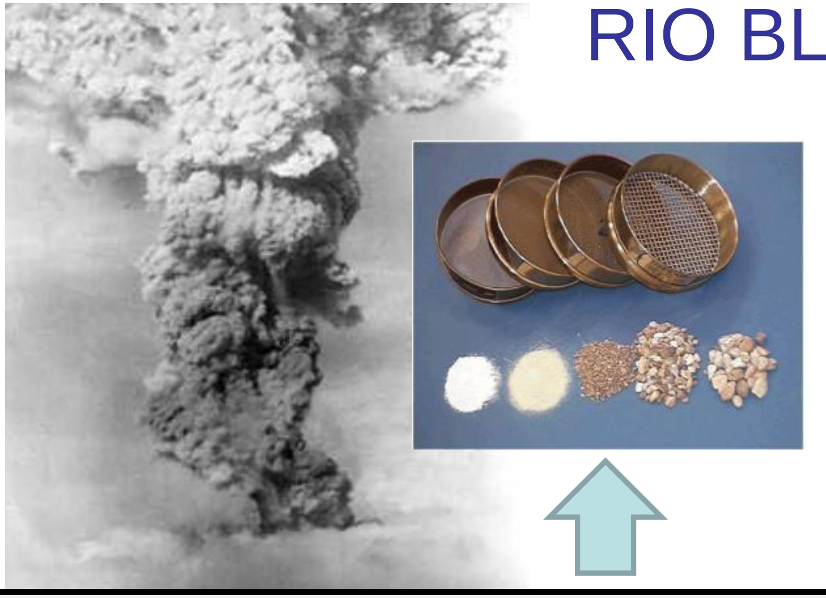
Pele's hair

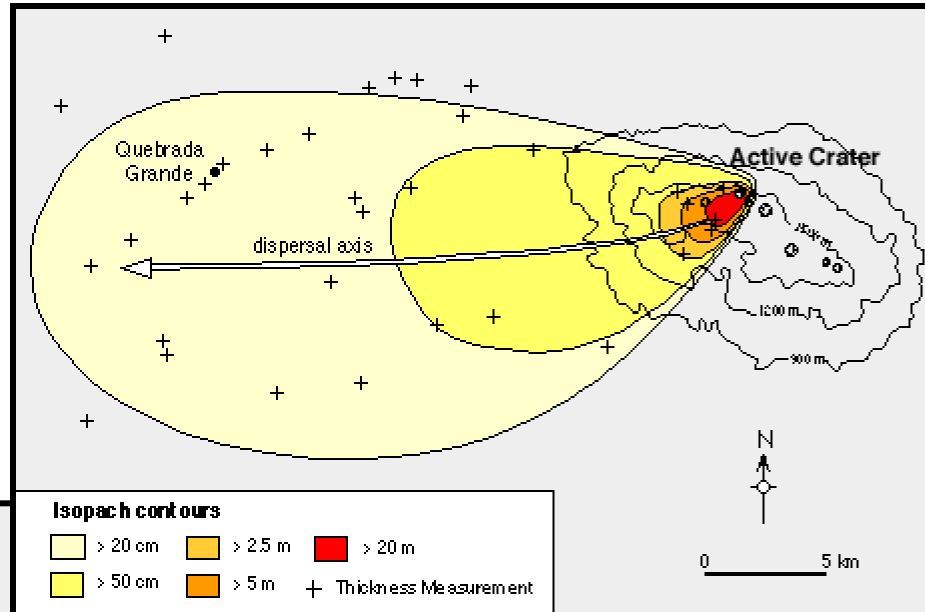
scoria

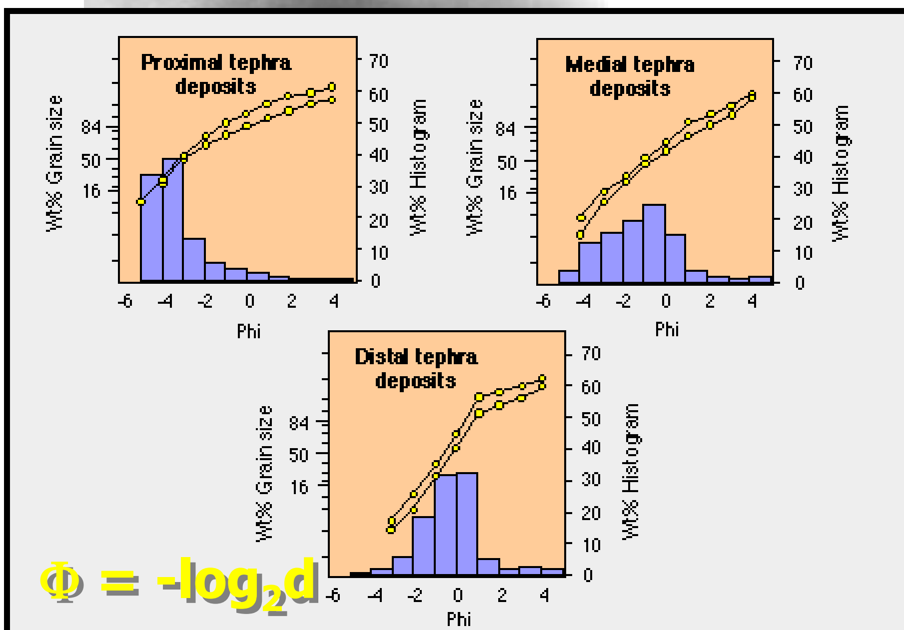
pumice



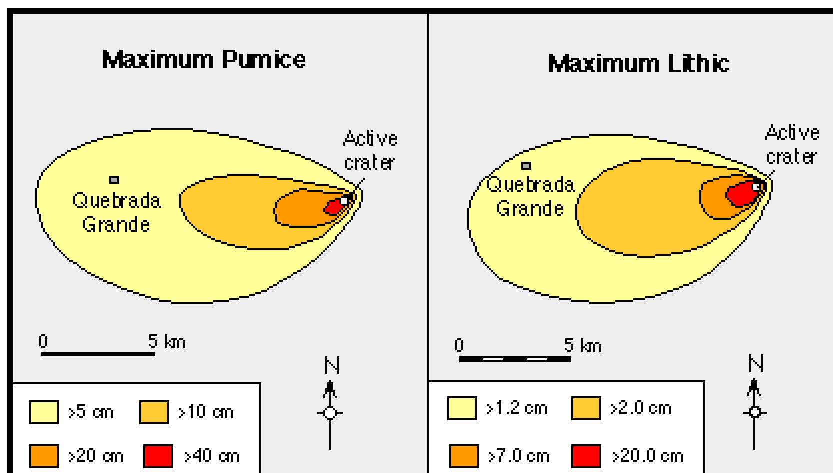
RIO BLANCO TEPHRA (3500 BP)



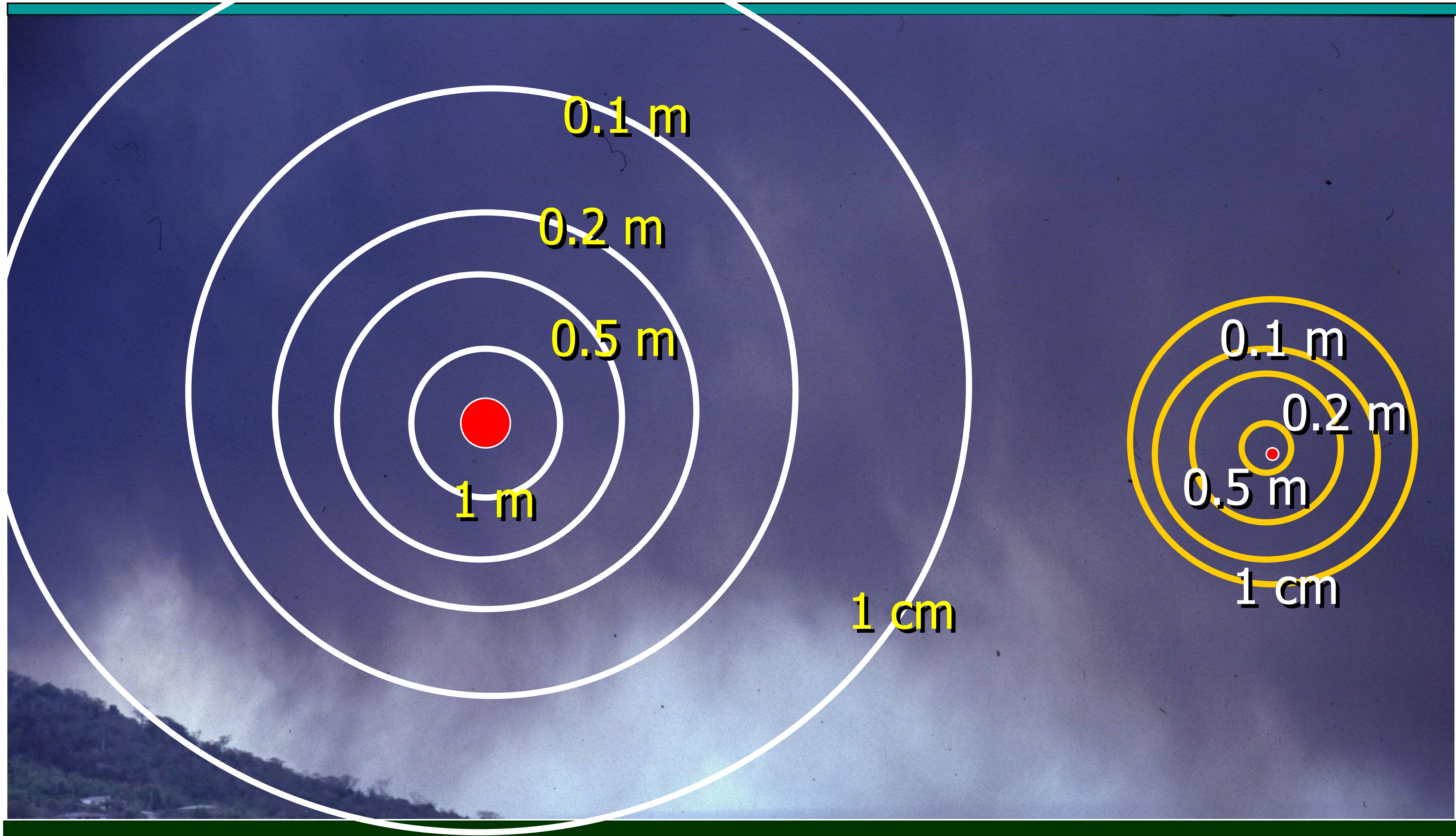




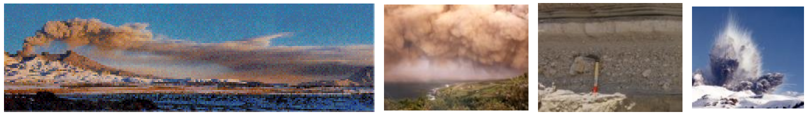
$\Phi = -\log_2 d$

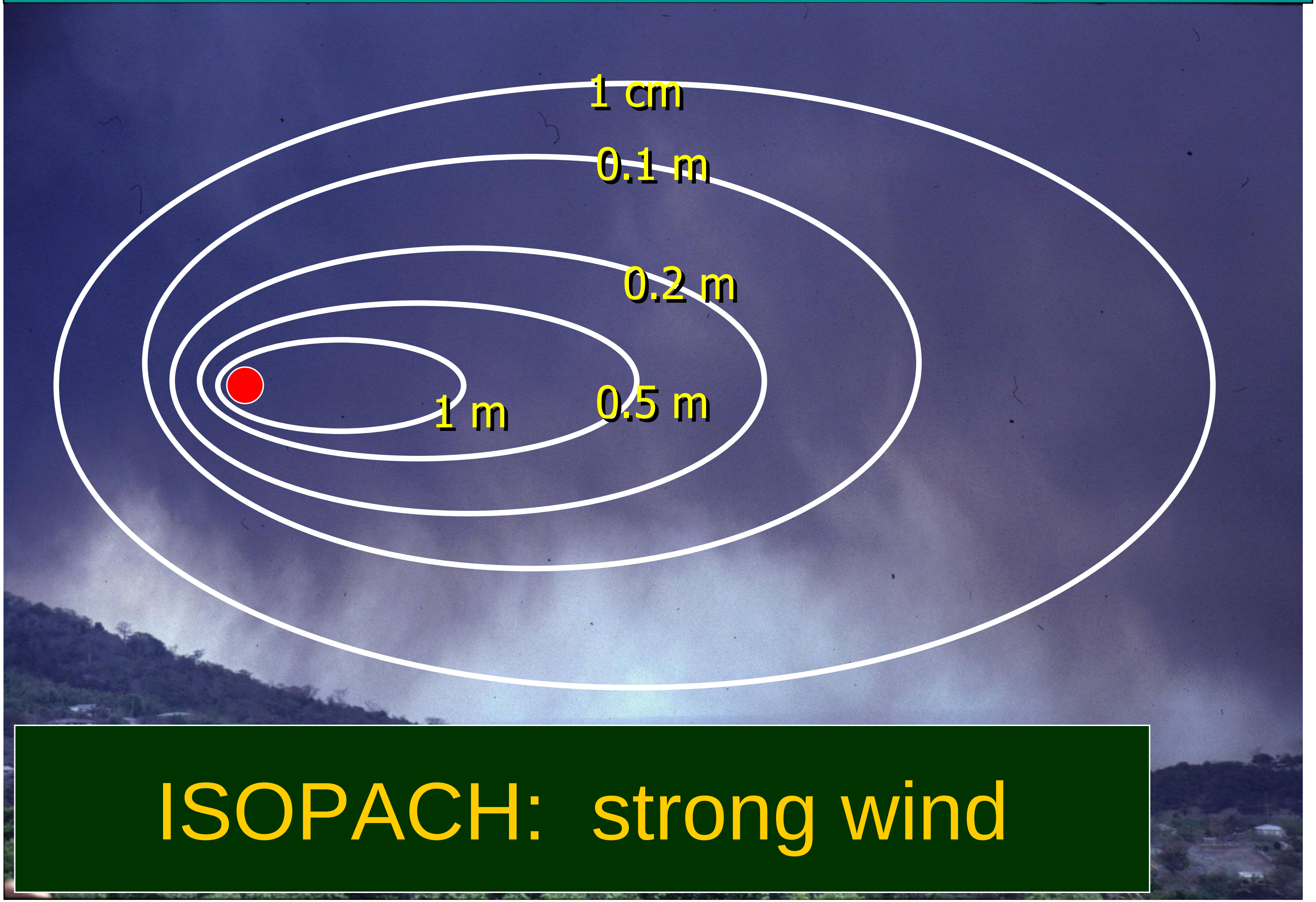


http://www.geo.utexas.edu/faculty/barker/kempter/rbtephra.html



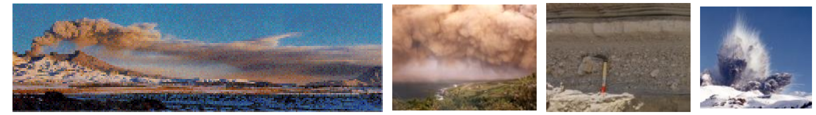
ISOPACH: contrasting intensity

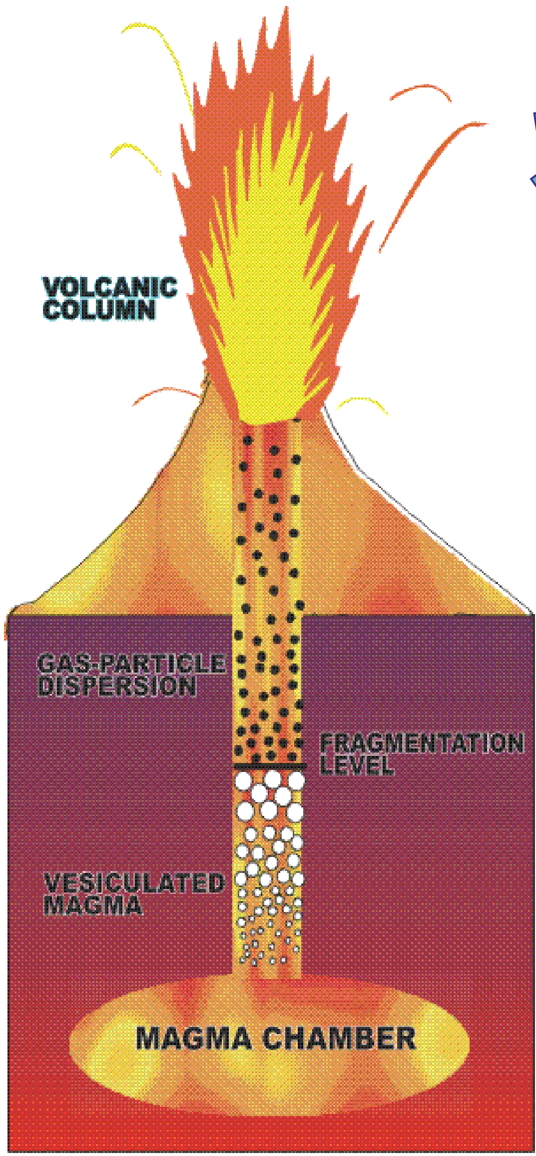


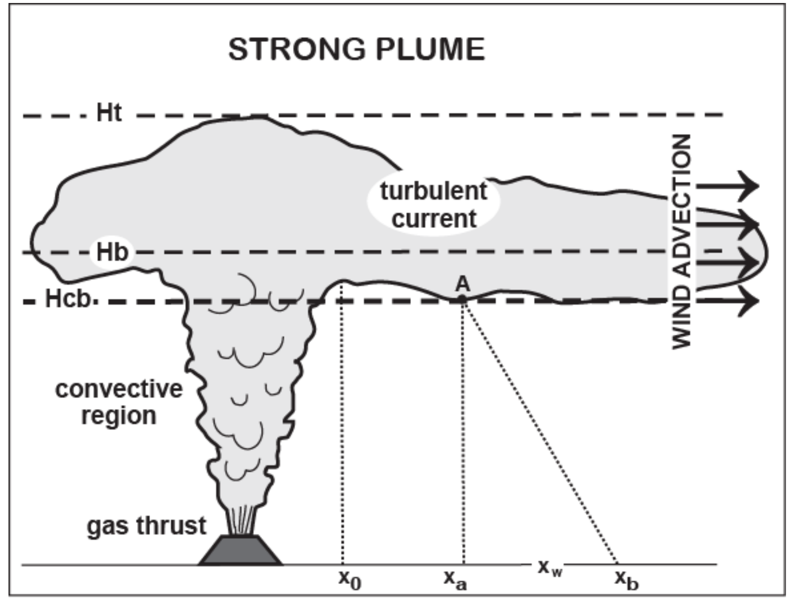


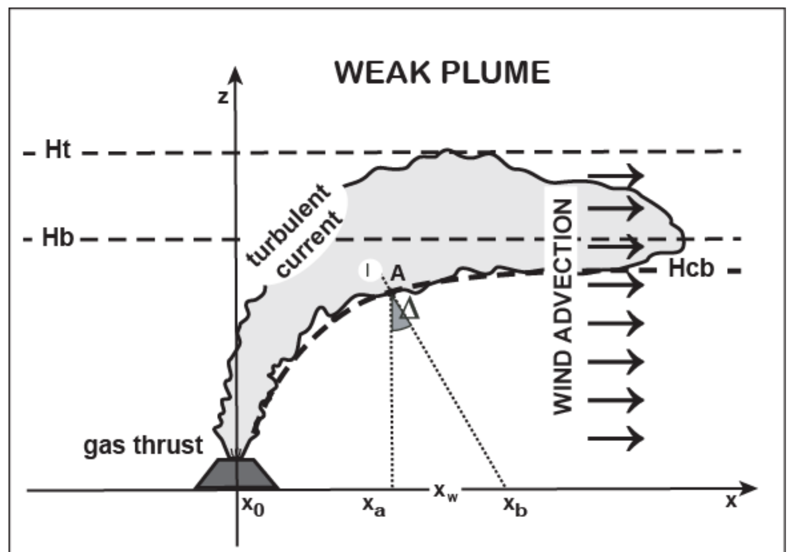
ISOPACH: strong wind

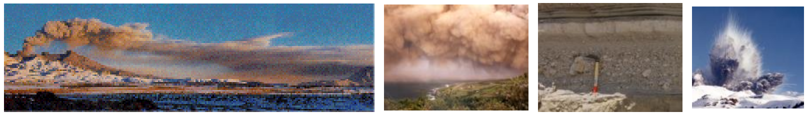
Characterization of tephra deposits: particle sedimentation

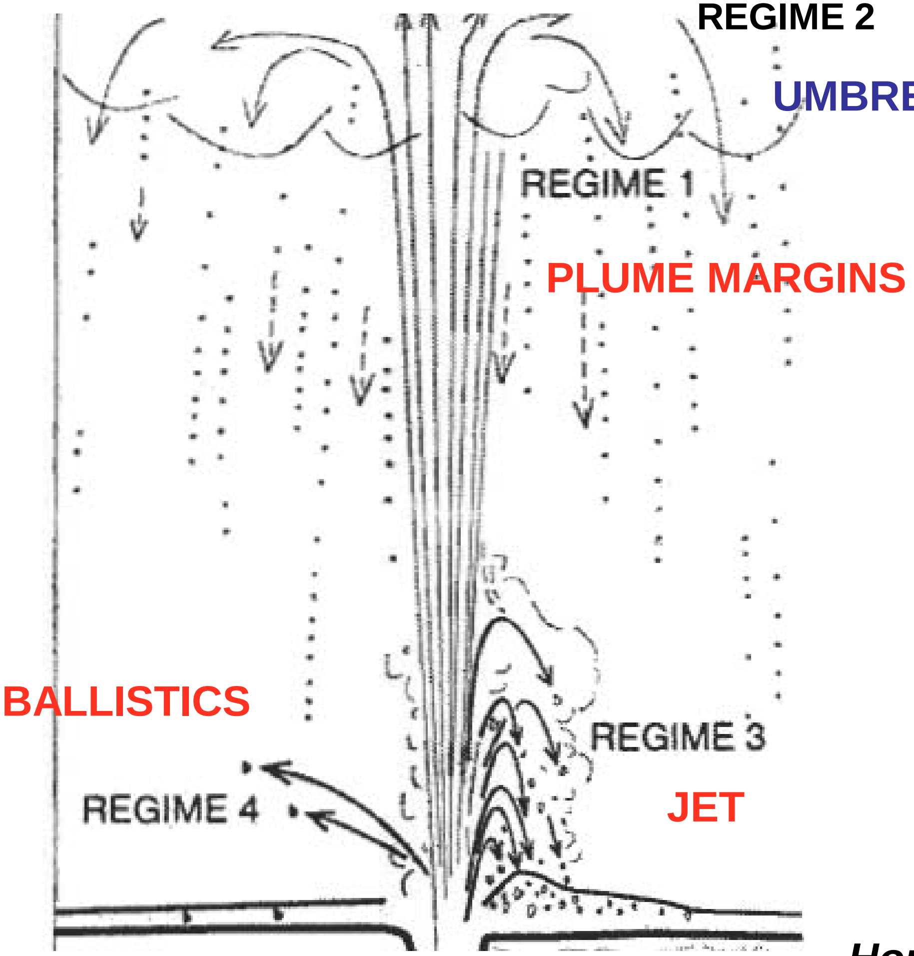




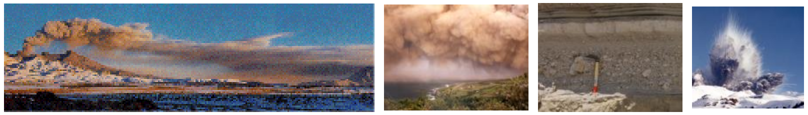


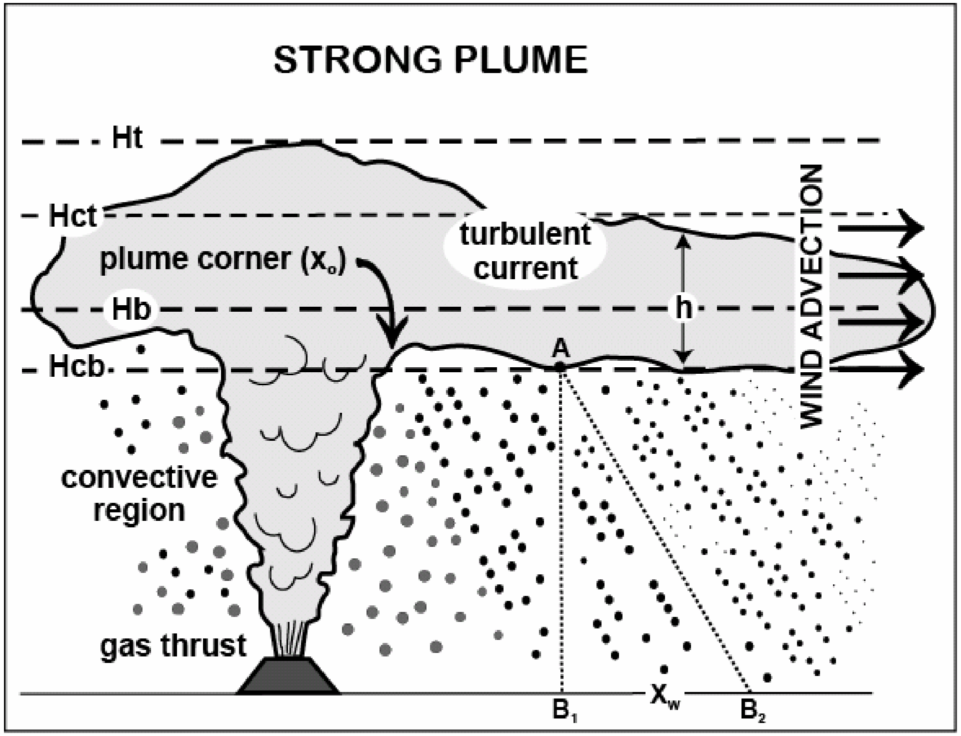






Houghton et al. [2004]





Sedimentation from strong plumes

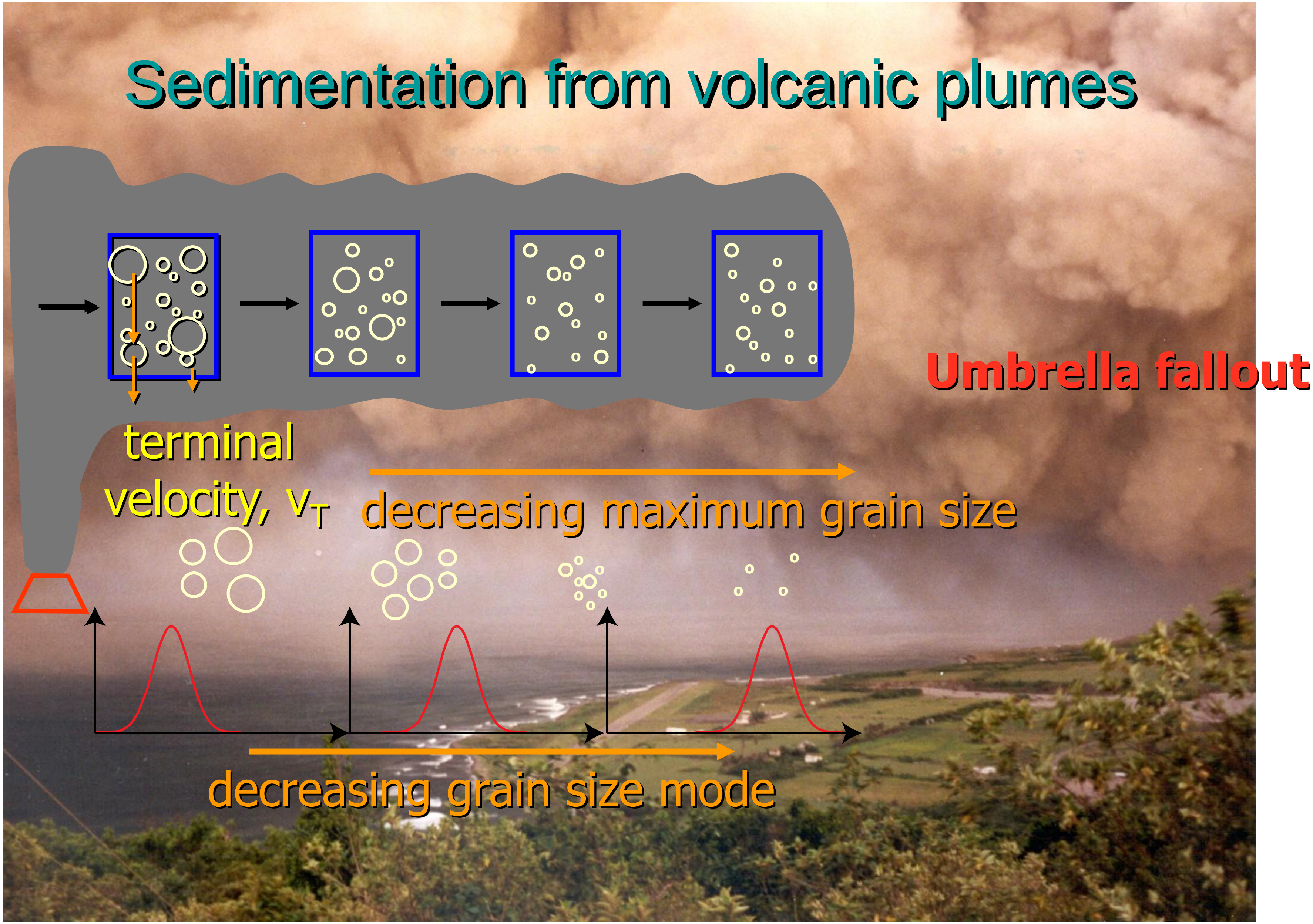
$$X_o = 0.2\ H_t$$

CLASS 1 → coarse fragments ejected from the jet (ballistics)
Typically <4km from vent

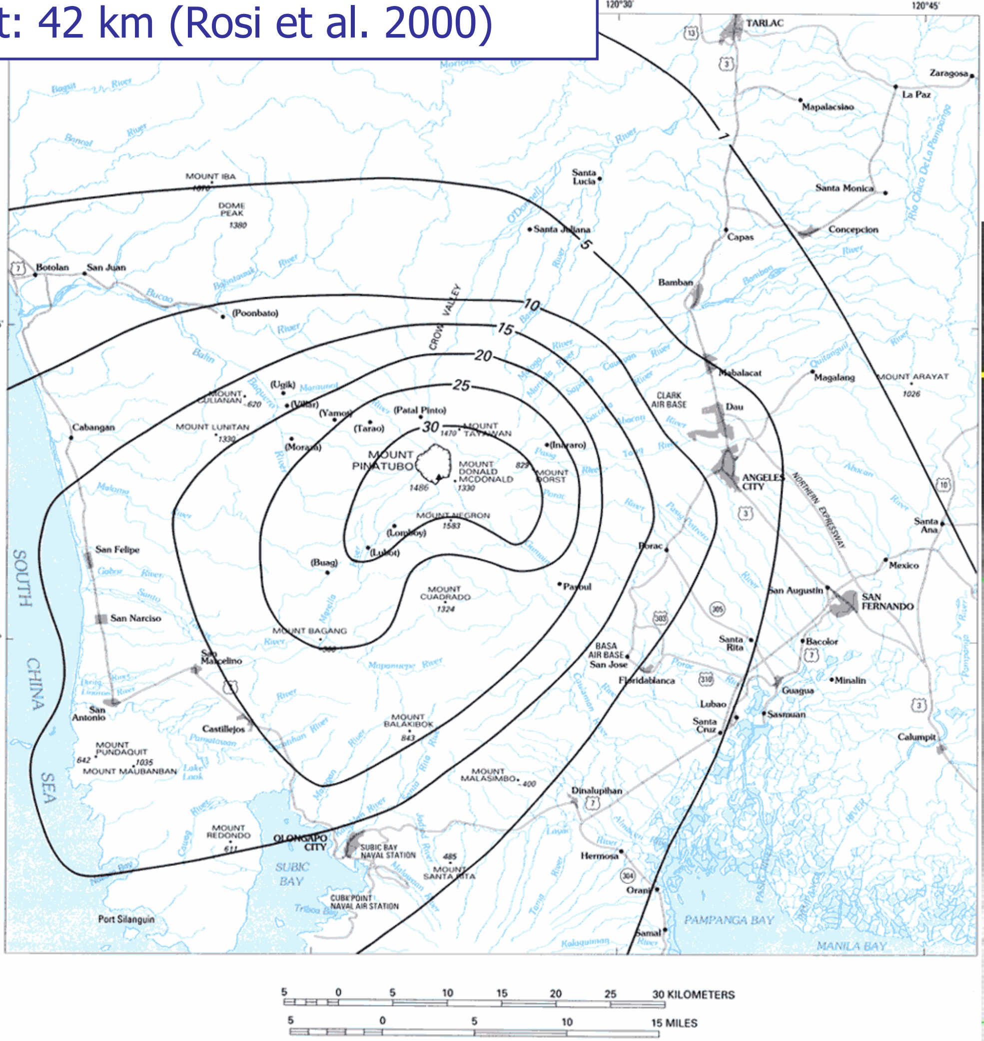
CLASS 2 → convective region. Typically <15km

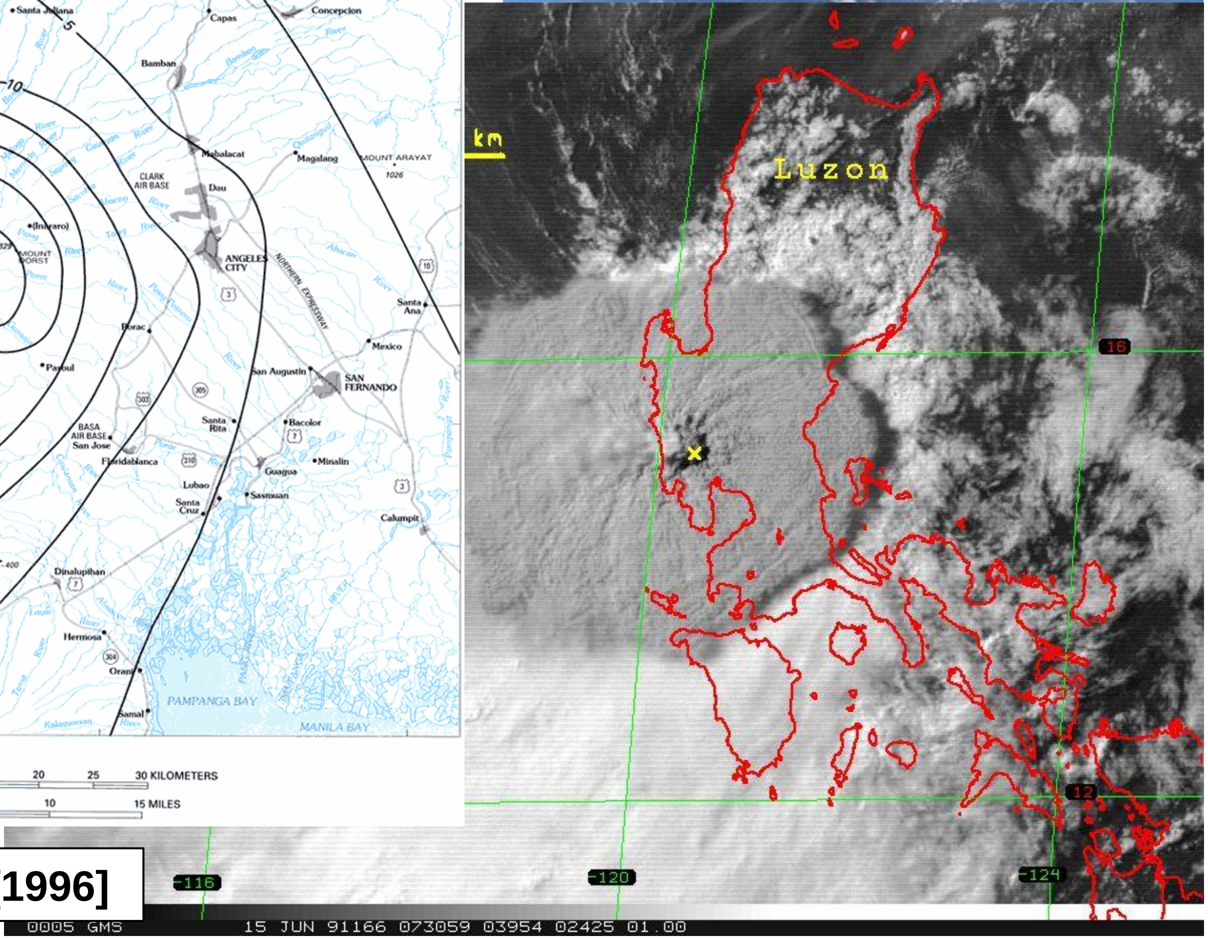
CLASS 3 → umbrella cloud

CLASS 4 → fine particles dispersed in atmosphere


Sedimentation from volcanic plumes
Umbrella fallout
terminal velocity, v_T
decreasing maximum grain size
decreasing grain size mode

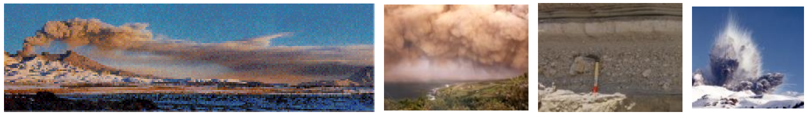
Pinatubo (Philippines) 1991
Ht: 42 km (Rosi et al. 2000)



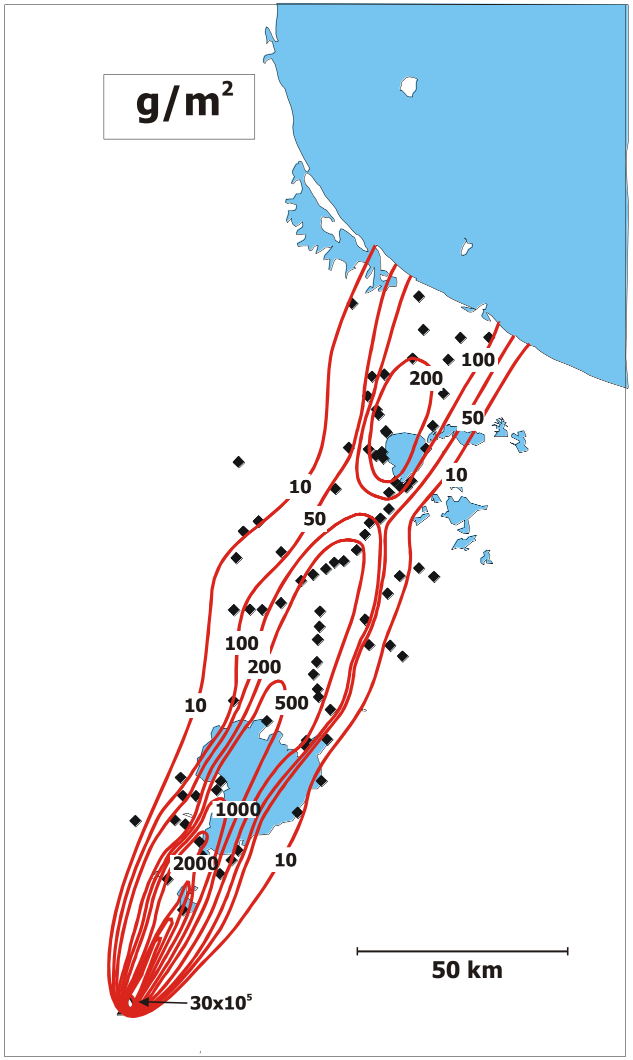


Paladio-Melosantos et al. [1996]

Characterization of tephra deposits: particle sedimentation

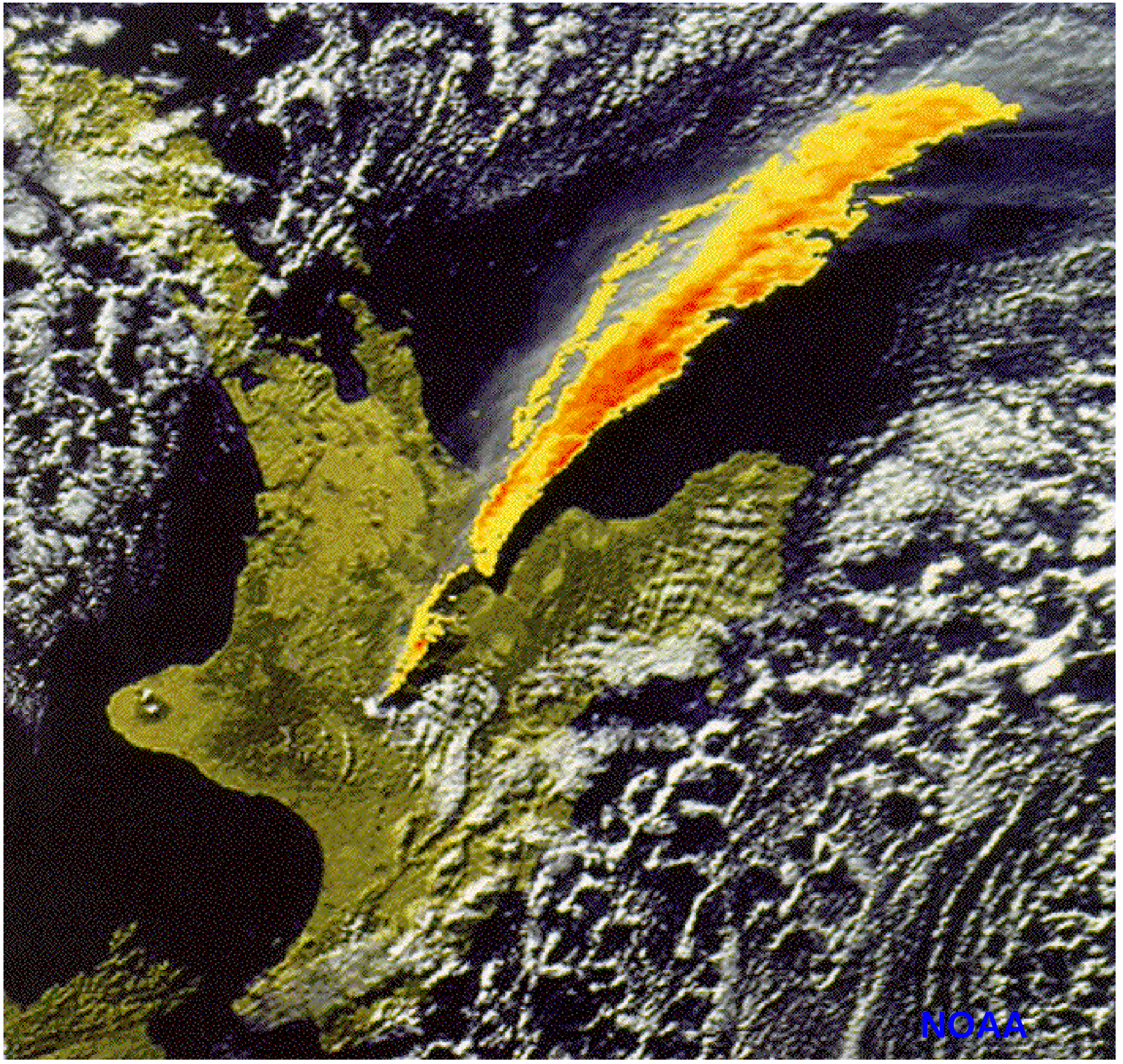


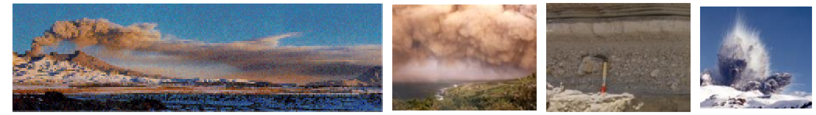
Ruapehu (NZ) 1996
Ht: 8.5 km
Wind speed: 24 m s^{-1}



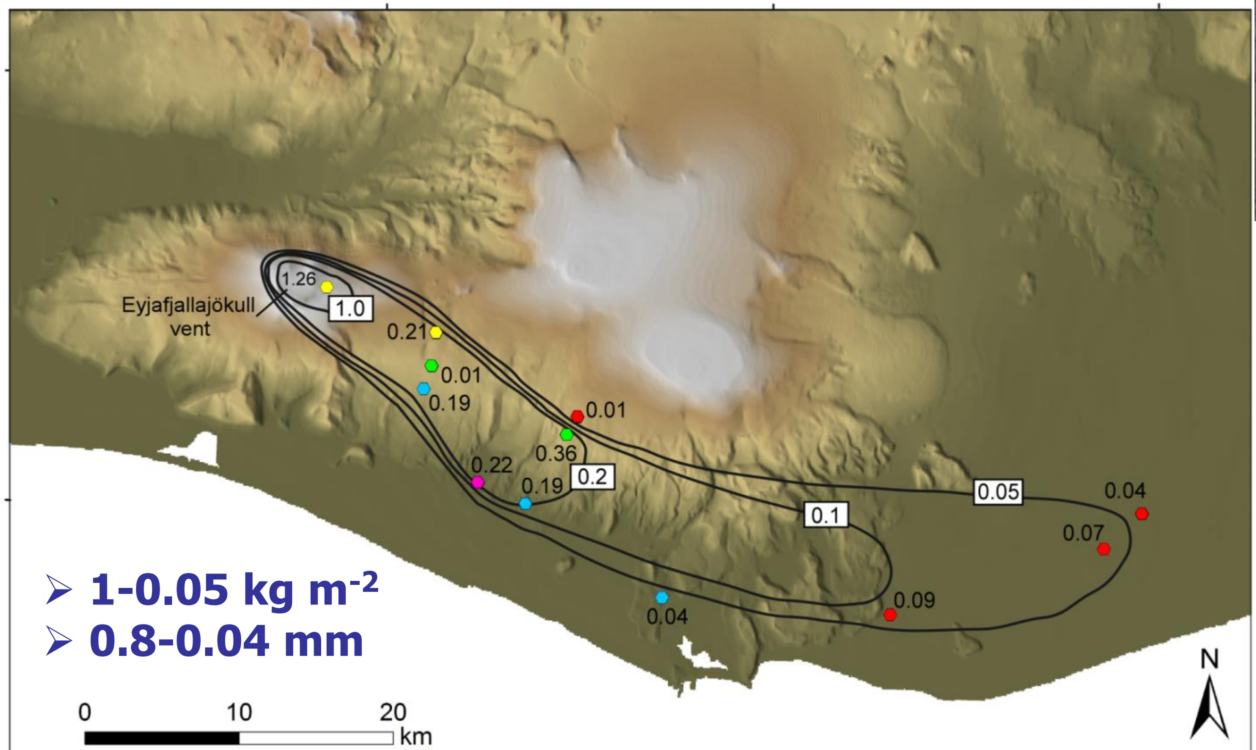
- **no upwind sedimentation**
- **narrow deposit**
- **lapilli sedimentation from inclined sector**

15:30 NZST 17 June 1996





Eyjafjallajökull 4-8 Mayo 2010 (Iceland)
Ht: 5-10 km asl
Wind speed: 10-16 m/s



- $1\text{-}0.05\ kg\ m^{-2}$
- 0.8-0.04 mm

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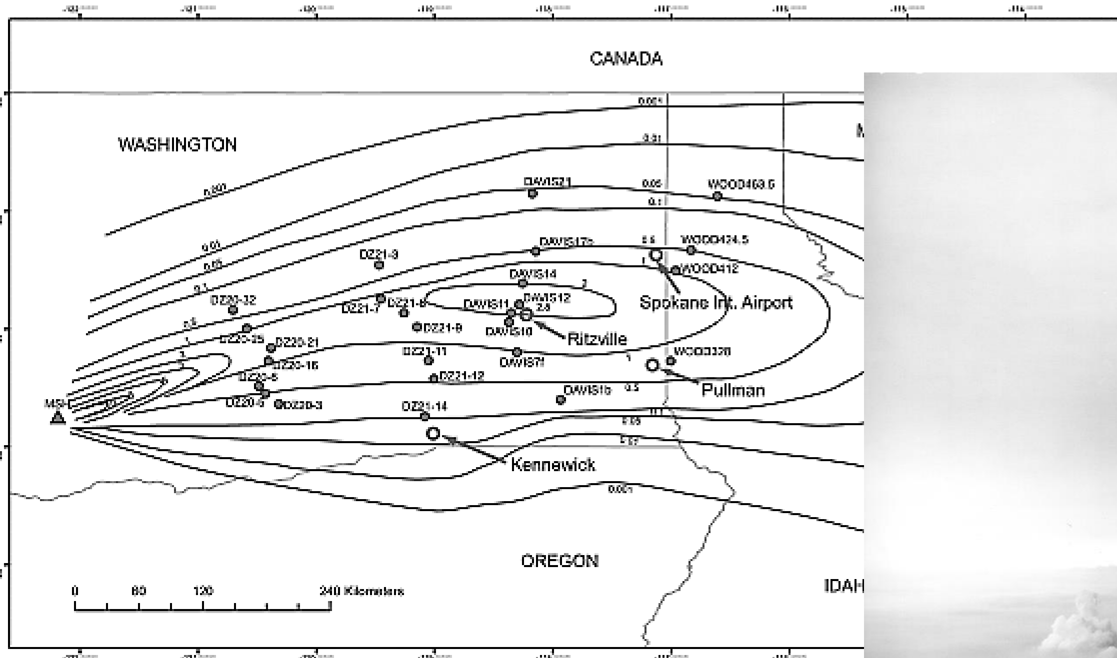
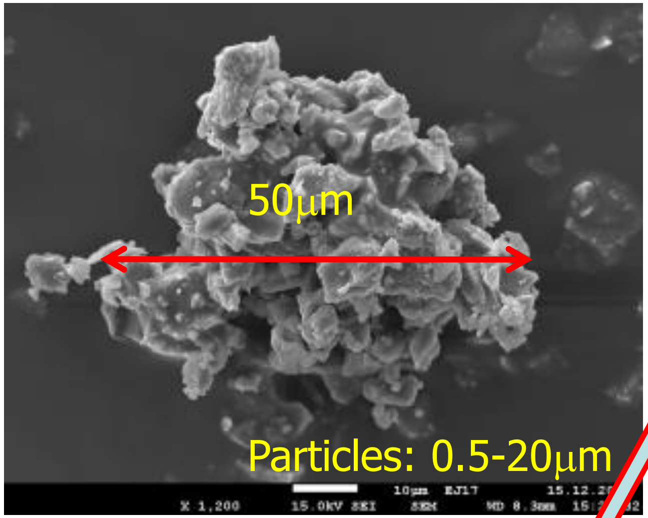
Mt St Helens (18 May 1980)
Ht: 19 km

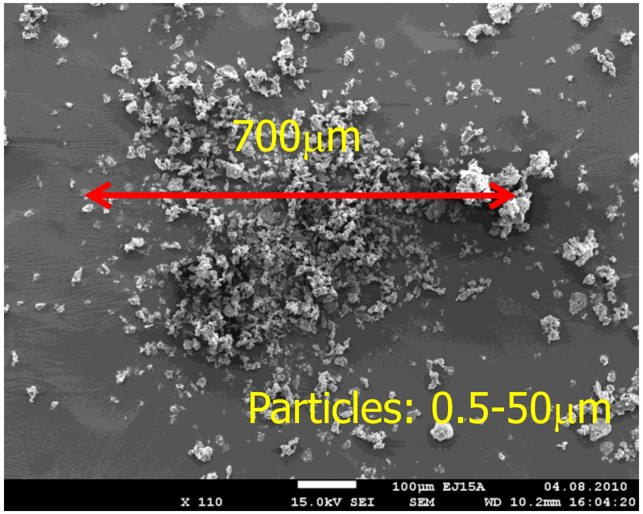


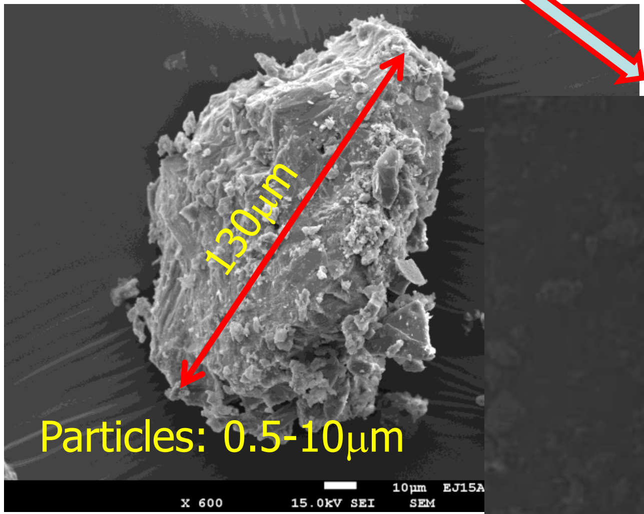
Characterization of tephra deposits: particle sedimentation

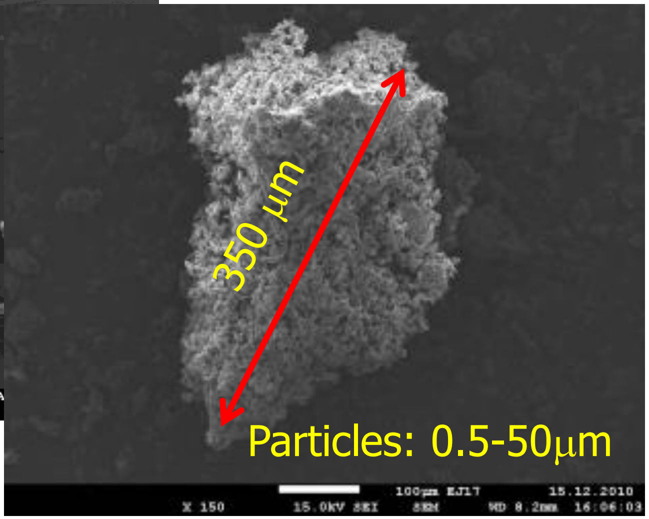



Aggregate type	Typical aggregate size range (diameter)	Particle size range (diameter)
Particle clusters (PC) PC1 (ash clusters)	Few 10 s µm up to ~10 cm	<1–40 µm: rare particles >200 µm
PC2 (coated particles)	0.05–2 mm	<1–40 µm





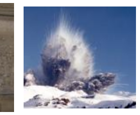


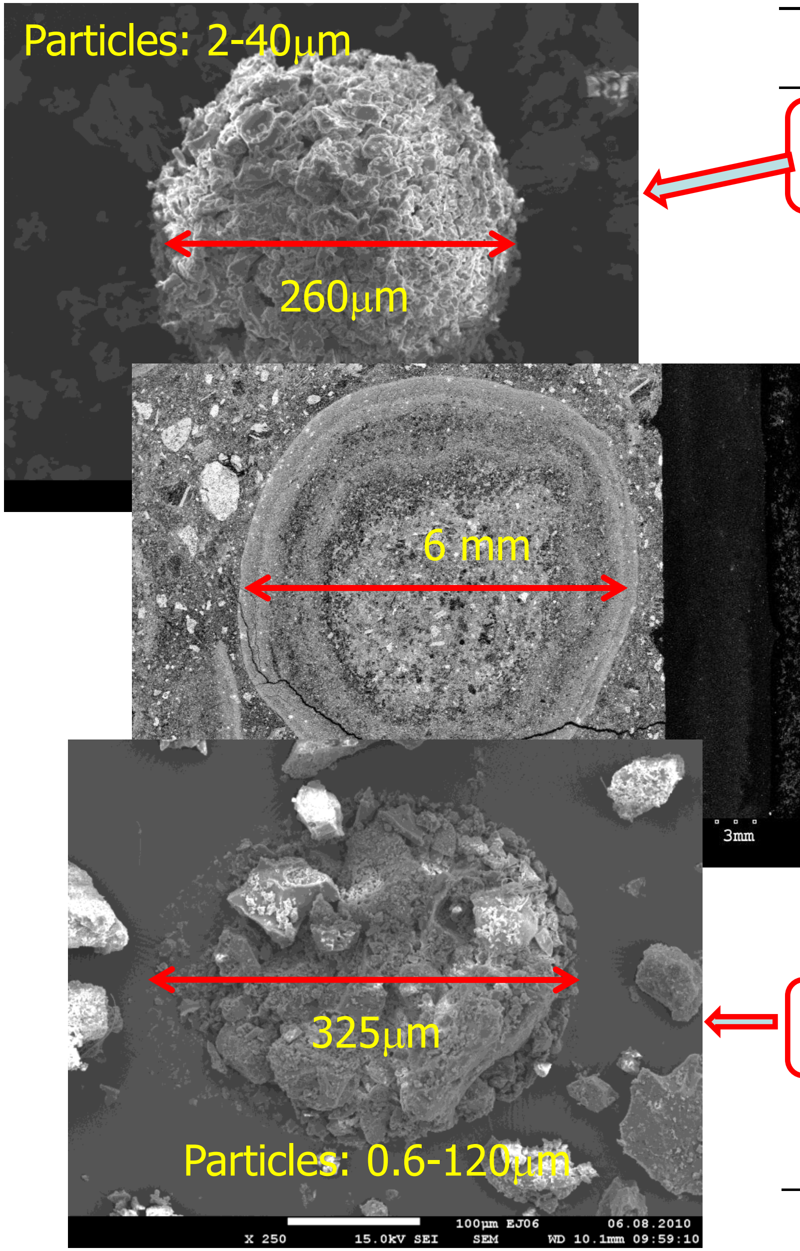


Brown et al. 2012

Characterization of tephra deposits: particle sedimentation



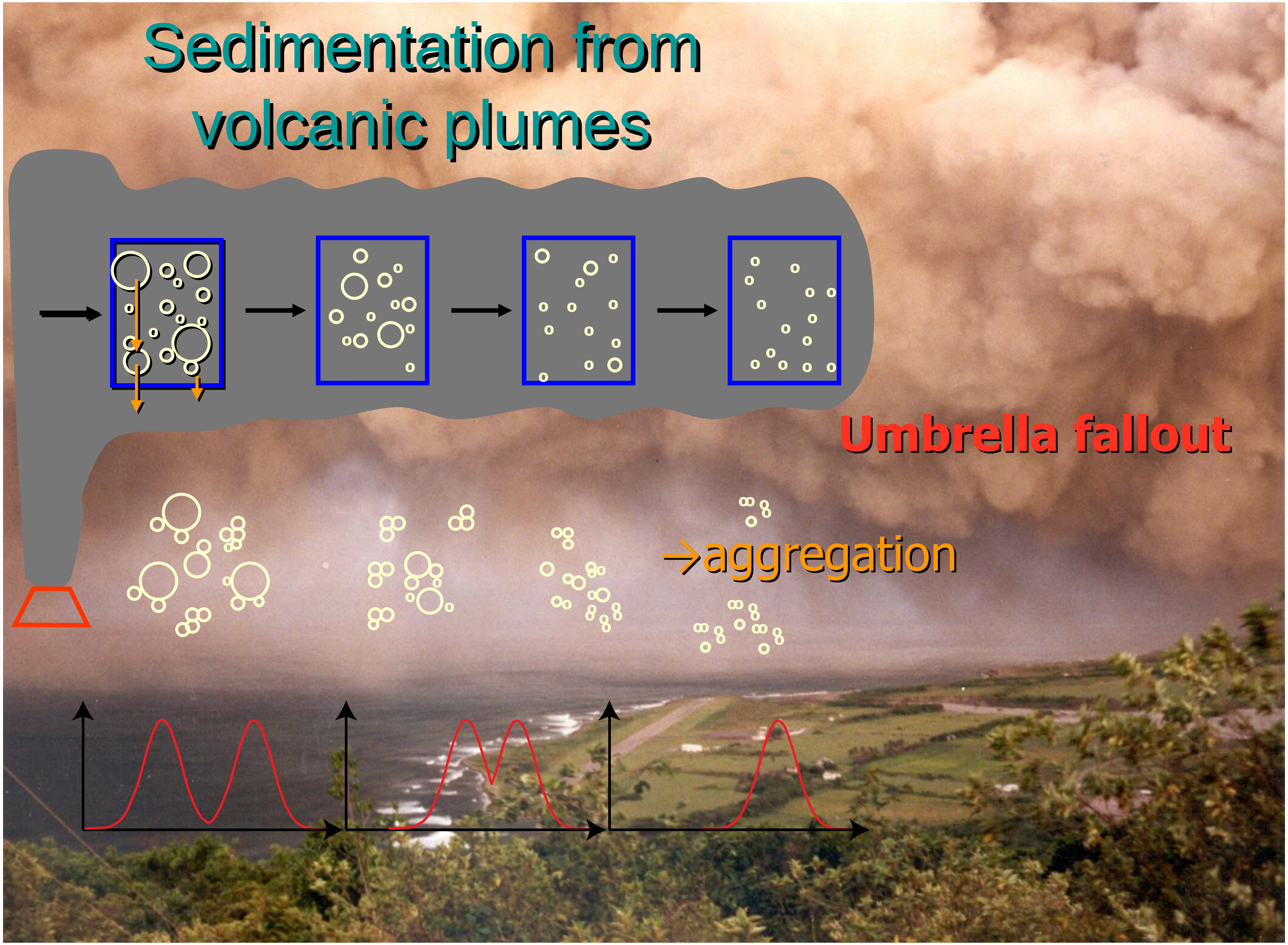



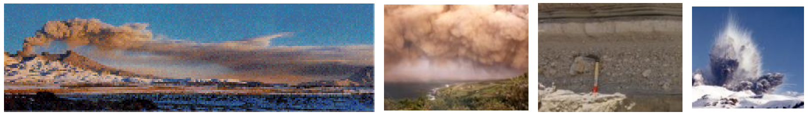


Aggregate type	Typical aggregate size range (diameter)	Particle size range (diameter)
Accretionary pellets (AP)		
AP1 (poorly-structured pellets)	100 μm up to a few mm	<1–400 μm; median values of 30–90 μm; rare particles >1 mm
AP2 (pellets with concentric structure)	2–15 mm; Rarely >30 mm	Cores: typical median values of <25–50 μm Rims: median values of 10–15 μm
AP3 (liquid pellets)	0.1–6 mm	Wide range of particle sizes (0.001–1 mm)

Brown et al. 2012

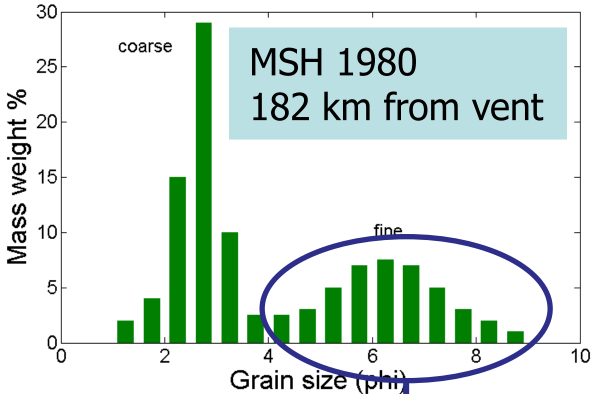
Sedimentation from volcanic plumes





Most particles <125 μm fall within aggregates with greater terminal velocities
$\rightarrow$Premature settling of fine ash from plumes

BIMODAL GRAINSIZE DISTRIBUTIONS

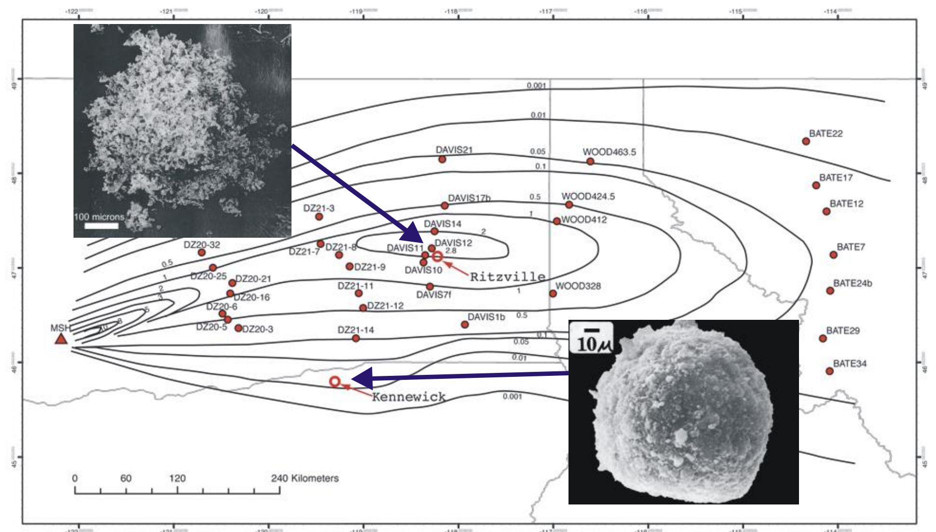


from Carey and Sigurdsson 1982

Particles that fell within aggregates

WIDESPREAD PHENOMENON:
Fine mode: <90 μm
e.g., Mt St Helens, Sakurajima, Soufriere St Vincent, Montserrat, Vesuvius, Laacher See Tephra, Tambora, Eyjafjallajökull

DOUBLE SECONDARY MAXIMA OF ACCUMULATION



Isomass map of the MSH 1980 eruption; from Durant et al. 2009

e.g., Mt St Helens 1980, Quizapu 1932, Hudson 1991




MAIN QUESTION:

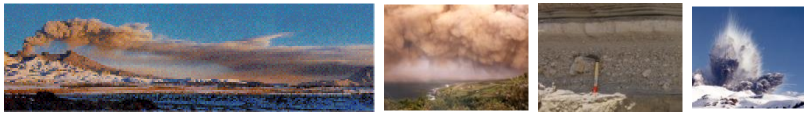
How well can we constrain explosive eruptions?

How well can we constrain Eruption Source Parameters (ESP)?

→ Erupted Volume/Mass, Plume Height, Mass Eruption Rate (MER), Eruption Duration, Total Grain-Size Distribution (TGSD)

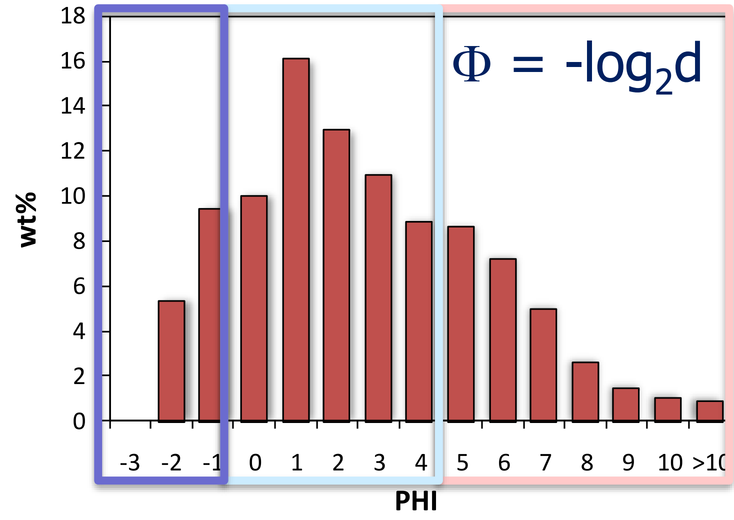
GRAIN SIZE PROPERTIES

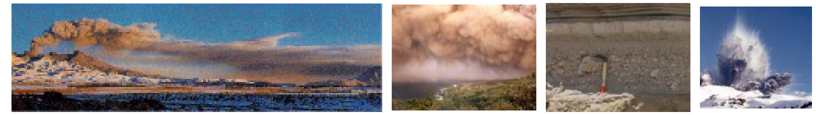
median, sorting, whole deposit



GRAIN SIZE PROPERTIES

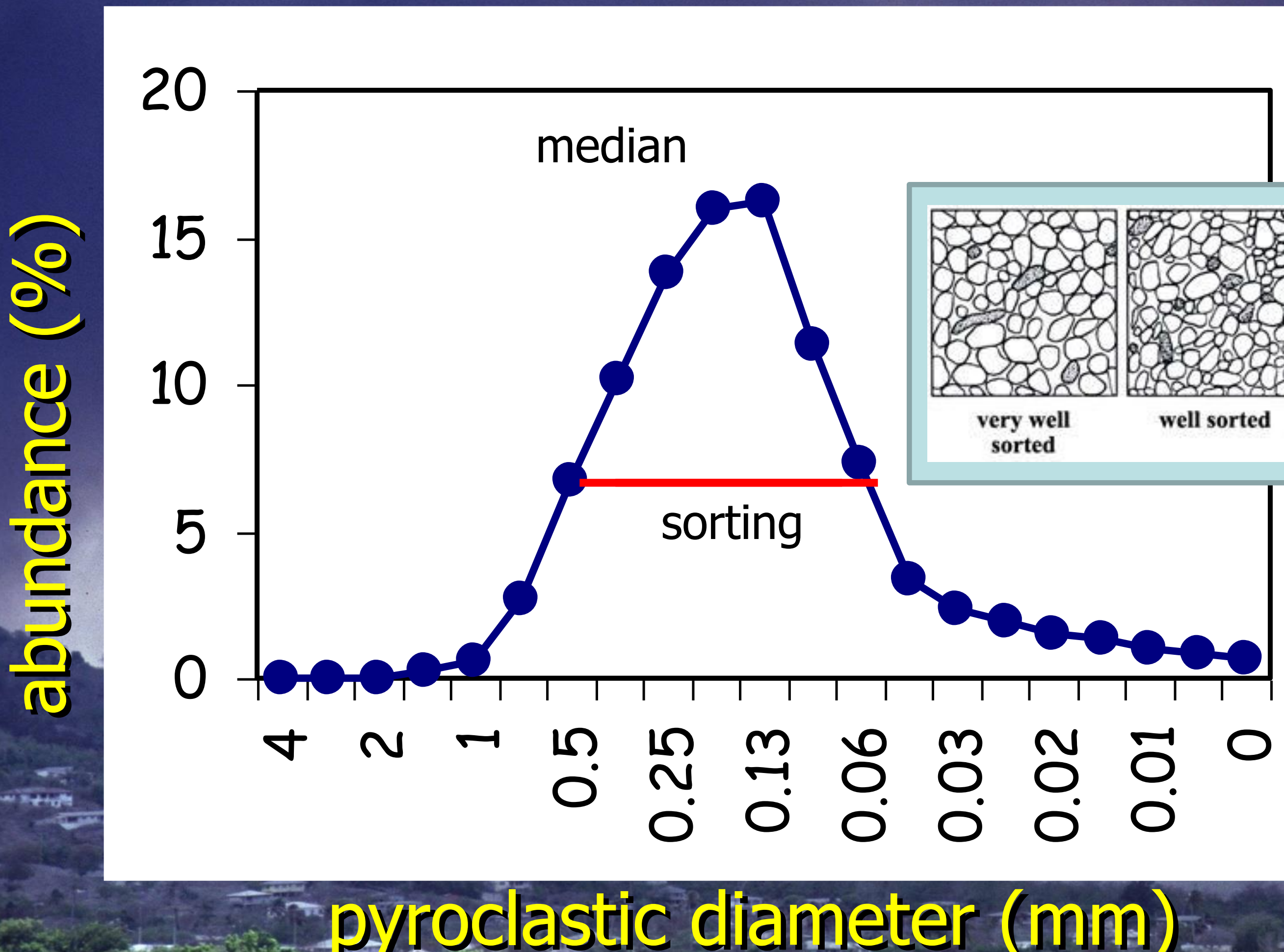
	bombs / blocks	lapilli	coarse ash	**fine ash**
Diameter	> 64 mm	(64 mm-2 mm)	(2 mm-63µm)	**(<63µm)**
Residence time	≈ sec	≈ min	≈ hours to few days	**several days**
Travel distance	proximal <10km	medial <50 km	distal <100 km	**very distal <1000 km**

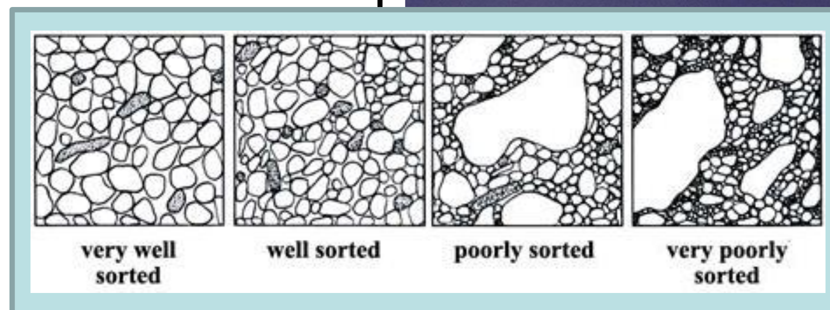


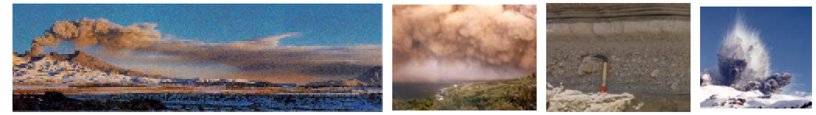


GRAIN SIZE PROPERTIES

absolute size = median   range of size = sorting







GRAIN SIZE PARAMETERS

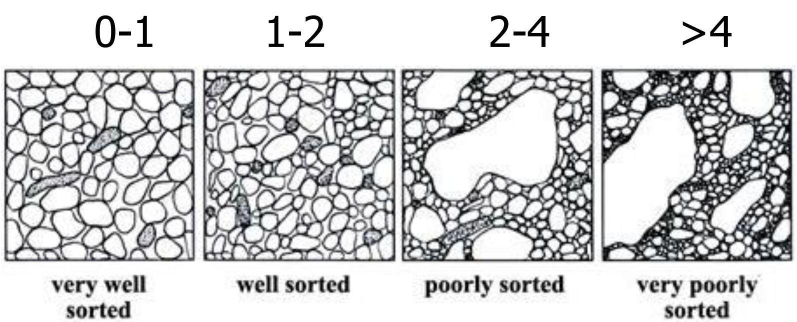
percentiles ϕ 16, 50, 84

Log scale: bin data geometrically phi scale = $-\log_2 D$

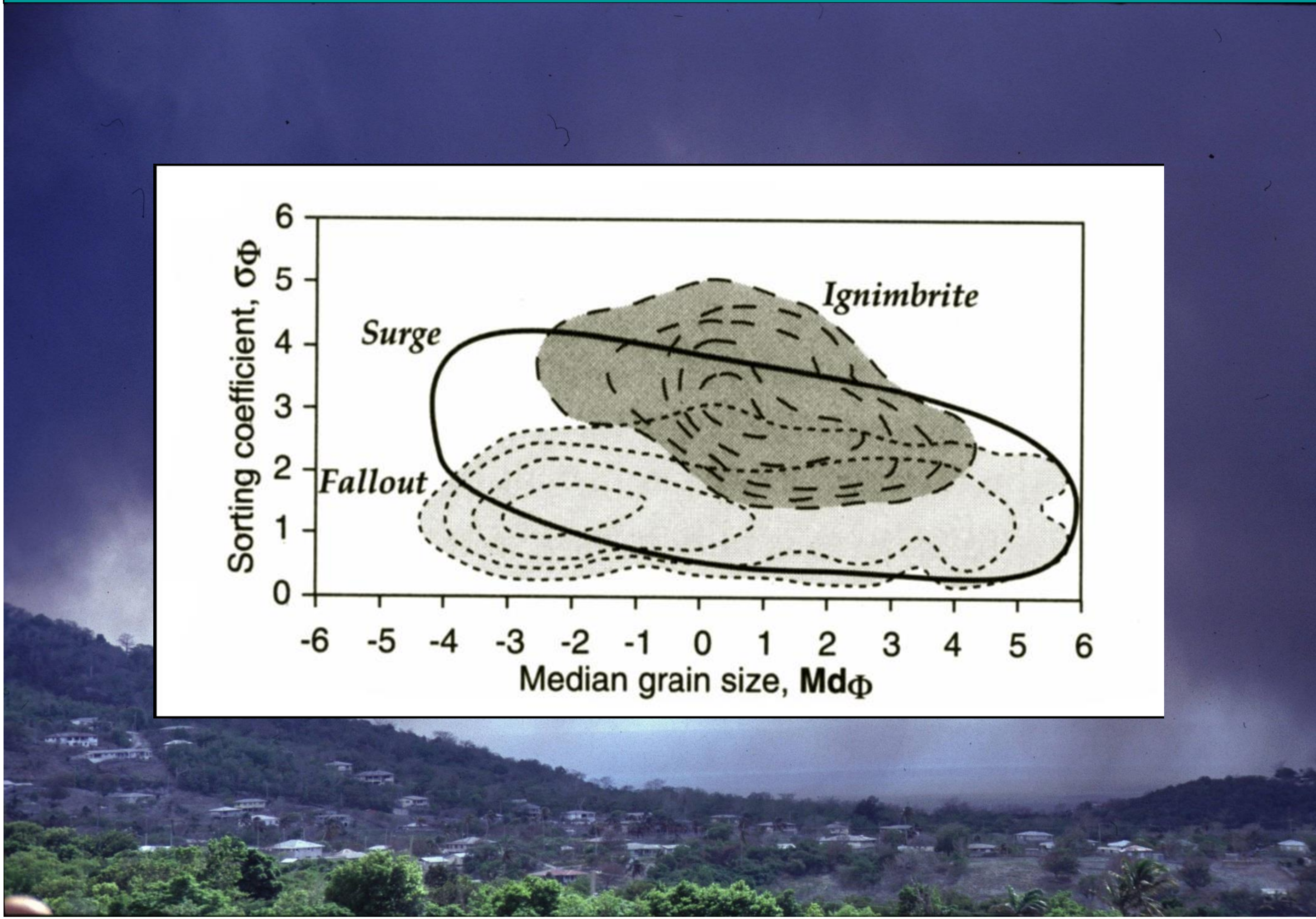
Inman (1952)

ABSOLUTE SIZE: median (M_d) = $\phi 50$

SORTING: graphic SD (s_G) = $(\phi 84 - \phi 16)/2$



Characterization of ESPs: grainsize distribution






Fall vs. Surge vs. Flow

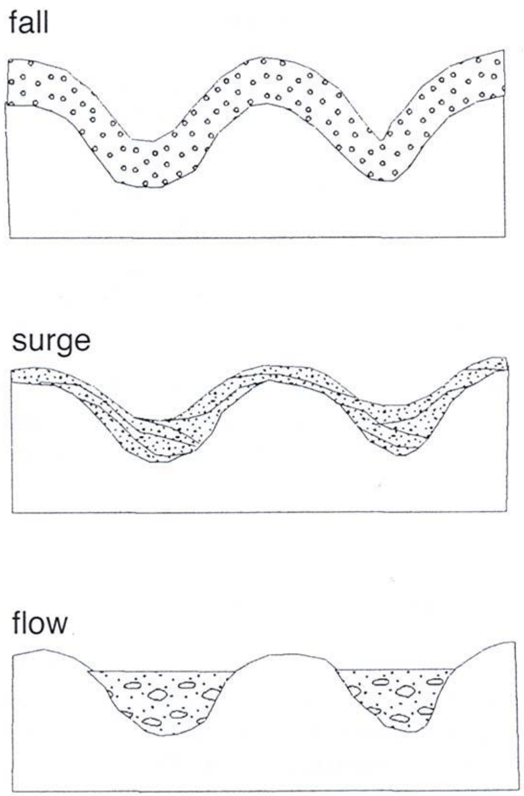
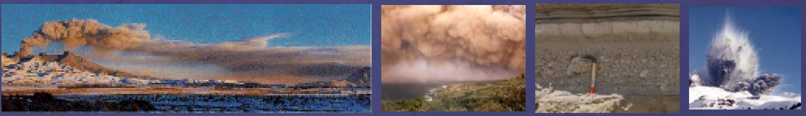


FIGURE 5 Schematic diagram of the archetypal characteristics of the three main pyroclastic deposit types. Fall—mantle bedding, with plane parallel beds and no internal erosion, good sorting (except where water is present; see Section III,E), juvenile clasts with angular to ragged shapes. Surge—nonmantling beds, thickening into low-lying areas, with cross-stratification, pinch-and-swell bedding and scoured contacts, moderate sorting, juvenile clasts with some degree of rounding. Flow—landscape-filling units, generally poorly bedded to nonbedded, poor sorting, rounded juvenile clasts.

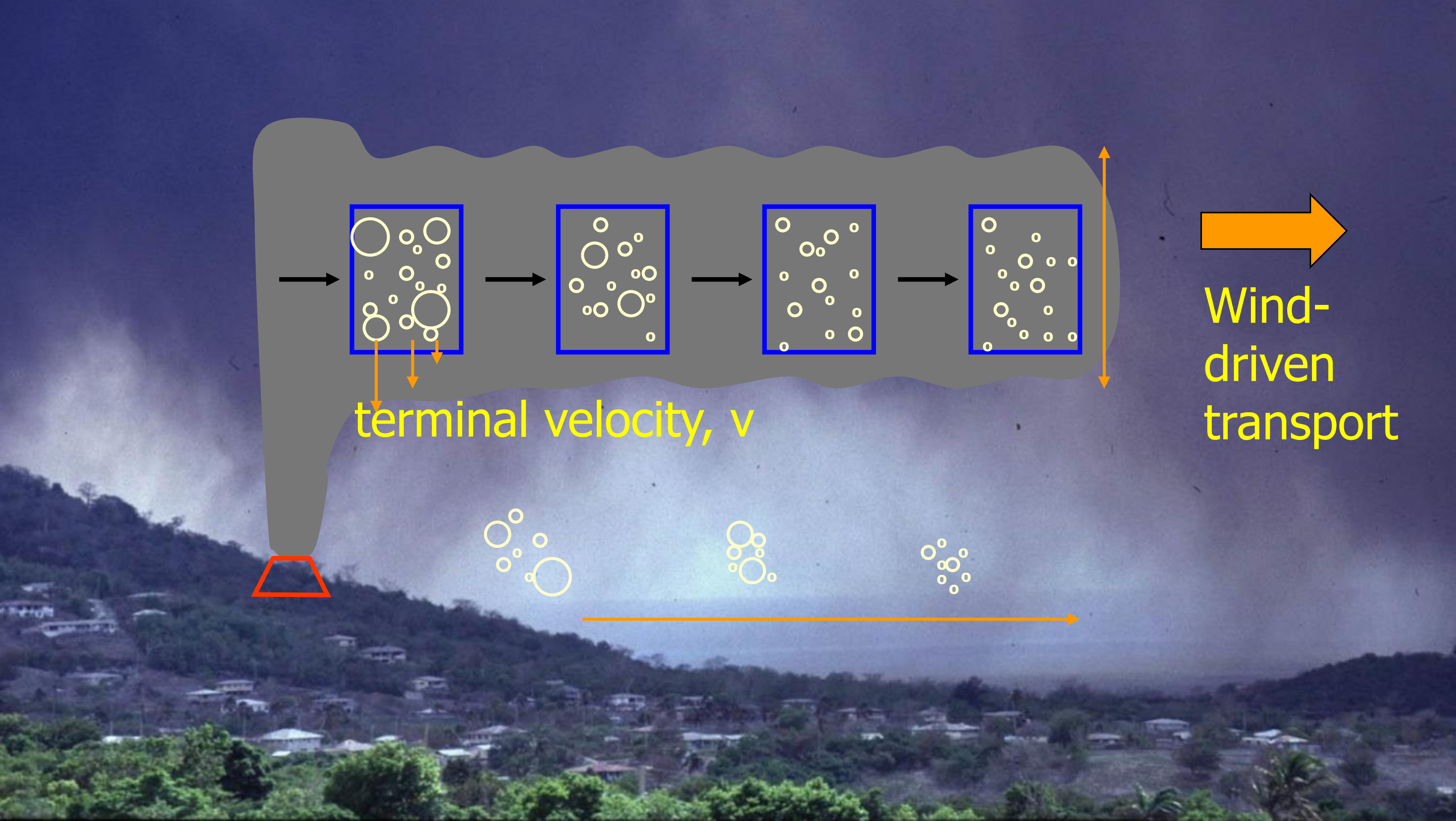
Fall: drape landscape, no cross beds or wave bedforms, well sorted, bedded

Surge: pinch and swell, basal scouring, cross bedding, (i.e., features that express lateral transport), good to poor sorting, sustained high temperatures rare.

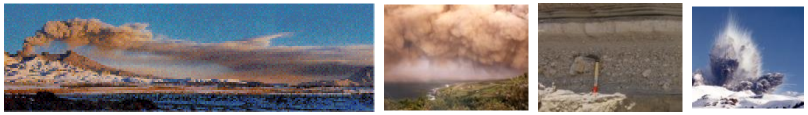
Flow: thicken into or are confined in valleys because flow is gravity driven, show basal scouring but lack internal bedforms, poor sorting. Sustained high temperatures (welding) typical. High T indicative of efficient transport (little mixing with ambient air).

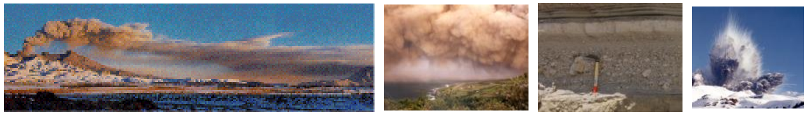


FALLOUT

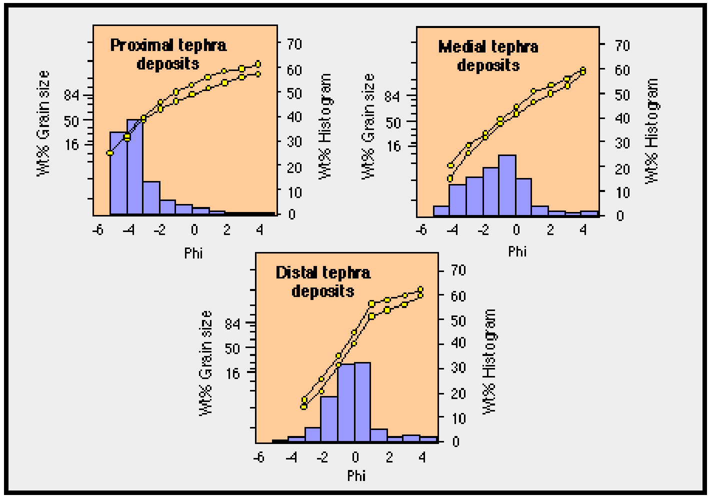


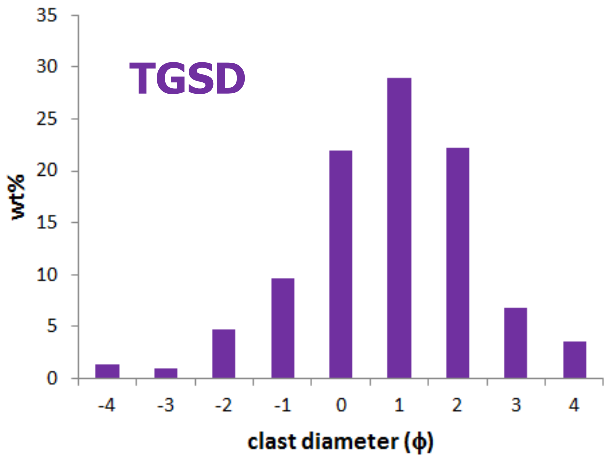
Characterization of ESPs: grainsize distribution

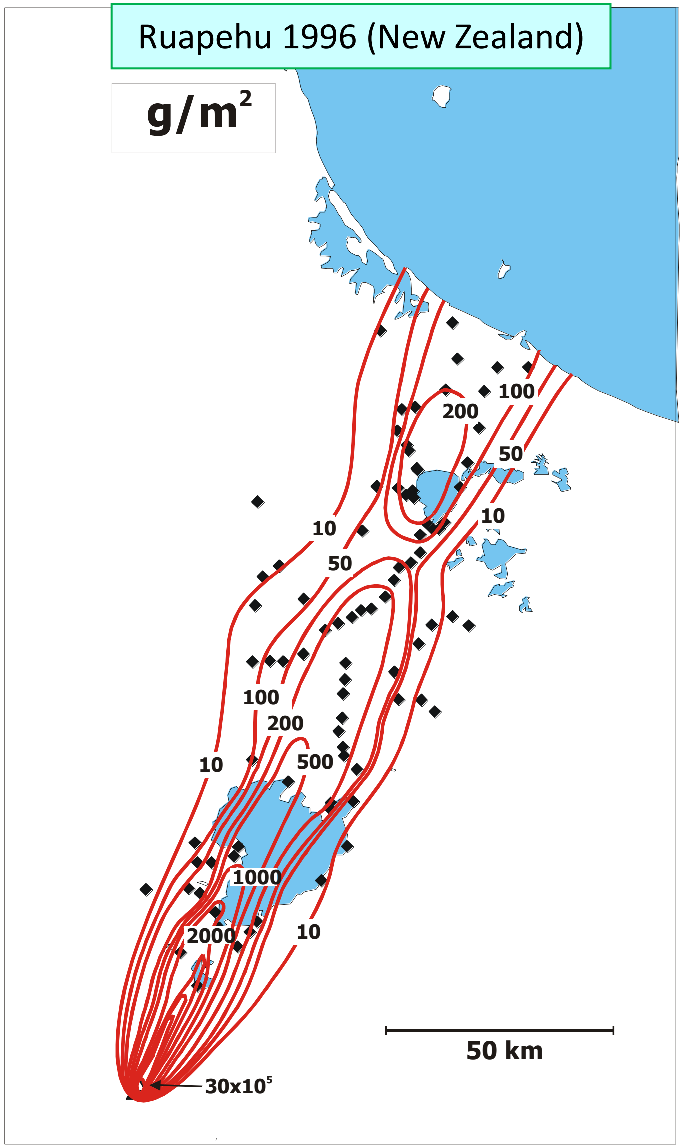


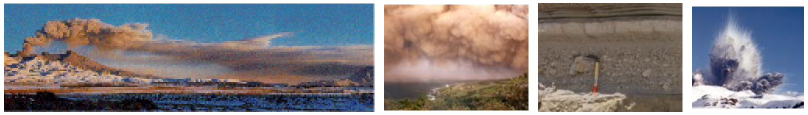


TOTAL GRAINSIZE DISTRIBUTION









TOTAL GRAINSIZE DISTRIBUTION

The determination of the total grainsize distribution relies on the averaging technique used and on the exposure of the deposit

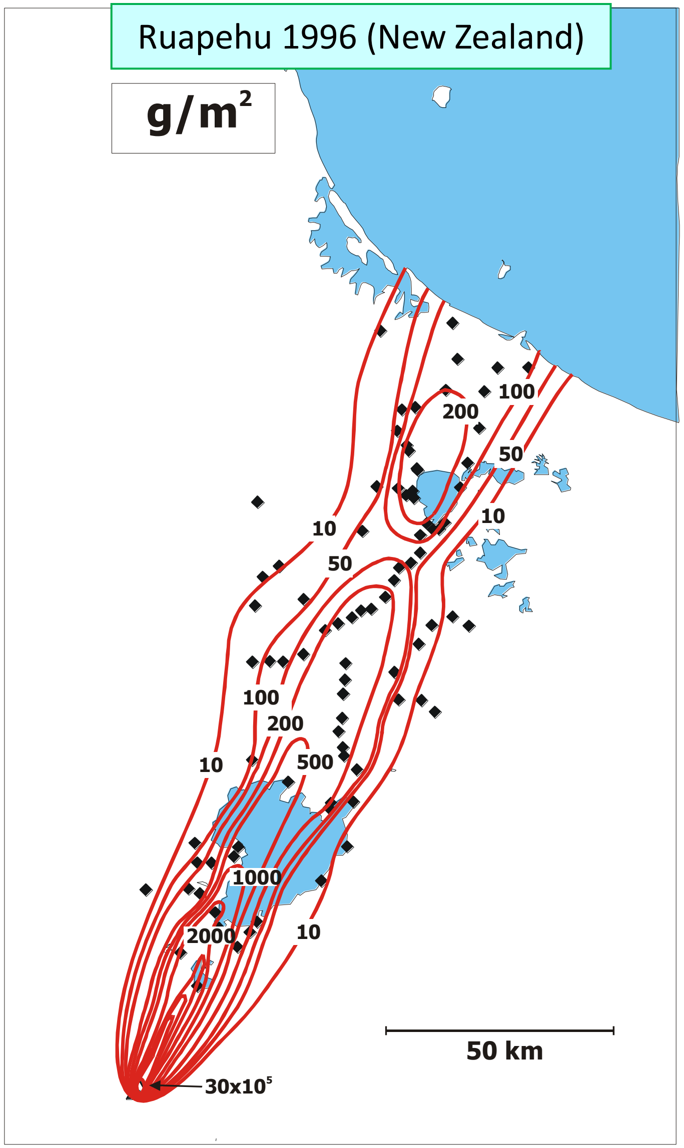
MOST COMMON TECHNIQUES:

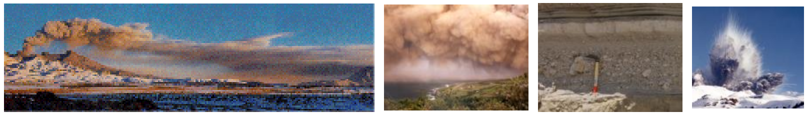
- **weighted average**

  does not deal well with deposits with non-uniform distributions of data

- **various types of sectorisation**

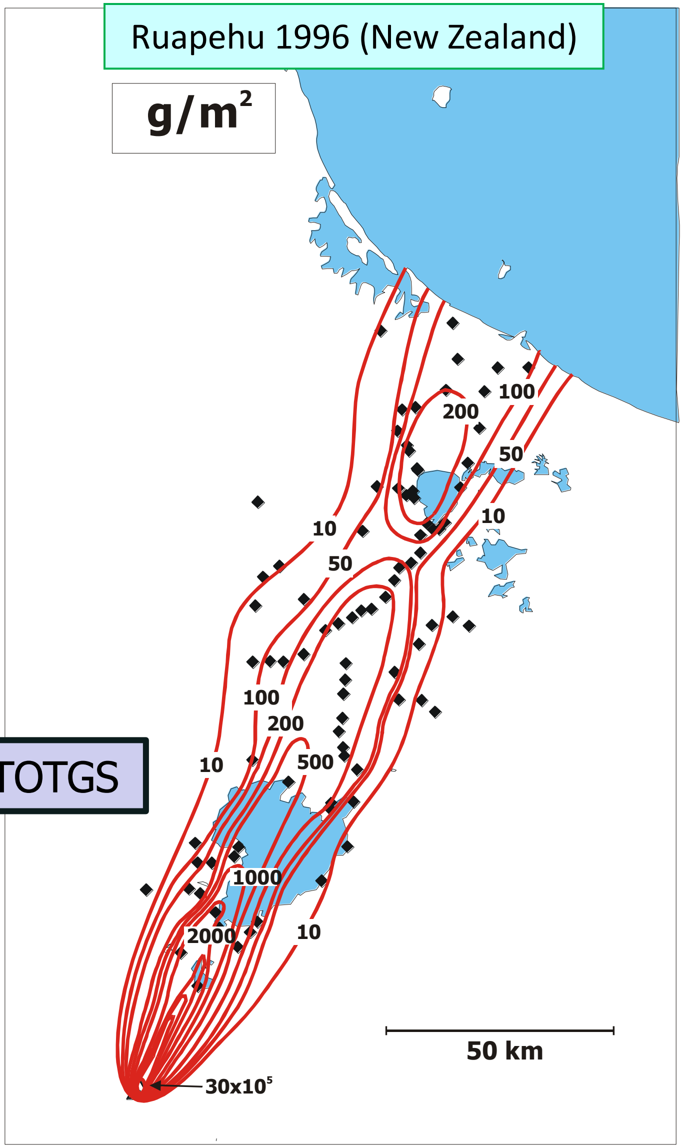
  biased due to the arbitrary choice of sectors



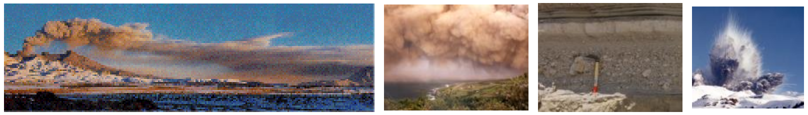


TOTAL GRAINSIZE DISTRIBUTION

https://github.com/e5k/TOTGS



Bonadonna and Houghton 2005



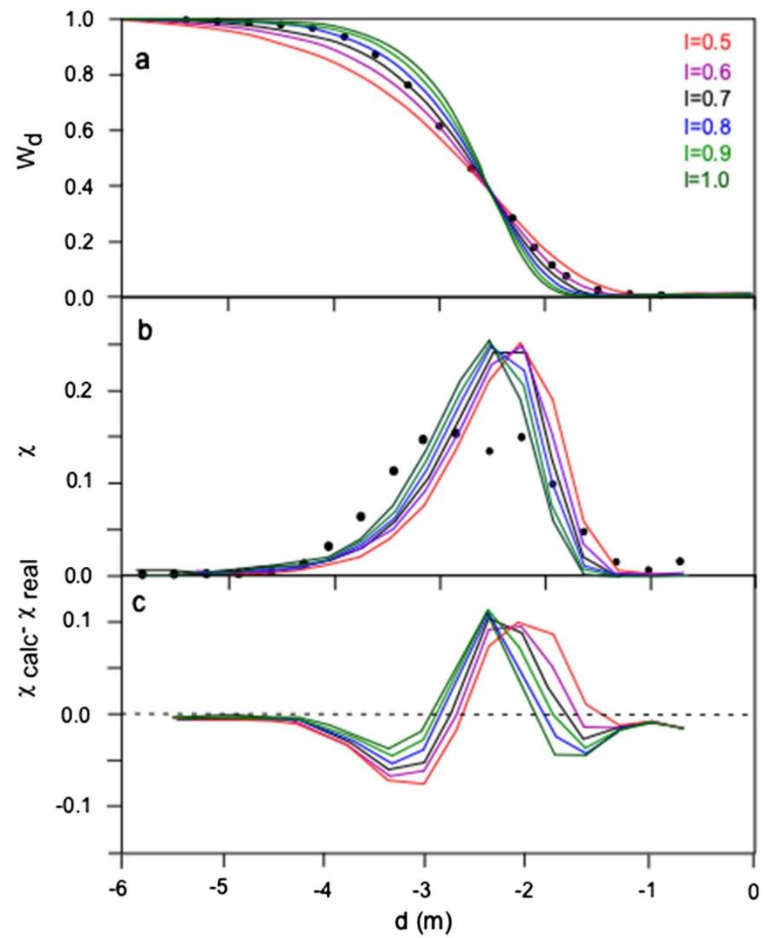
TOTAL GRAINSIZE DISTRIBUTION

Ruapehu 1996 (New Zealand)

How can we deal with missing data?

The **Rosin-Rammler** distribution:

- shows the best compromise between fitting capacity and stability with respect to sampling bias (50 field tephra deposits + 20 synthetic tephra deposits)
- can reconstruct TGSD only based on x_o (which is related to the median grain-size) within a given range of the shape parameter l provided by literature data
- can best reproduce the TGSD tails even when the parameter l is not well constrained
- but cannot well describe bimodality!

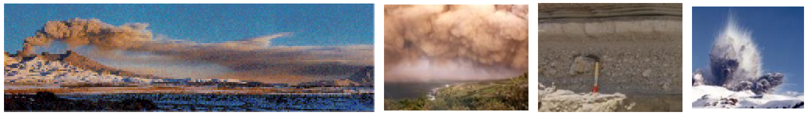


Pioli, et al. 2019




EXERCISE 1

DETERMINATION OF
GRAIN-SIZE PARAMETERS AND TGSD



MAIN QUESTION:

How well can we constrain explosive eruptions?

How well can we constrain Eruption Source Parameters (ESP)?

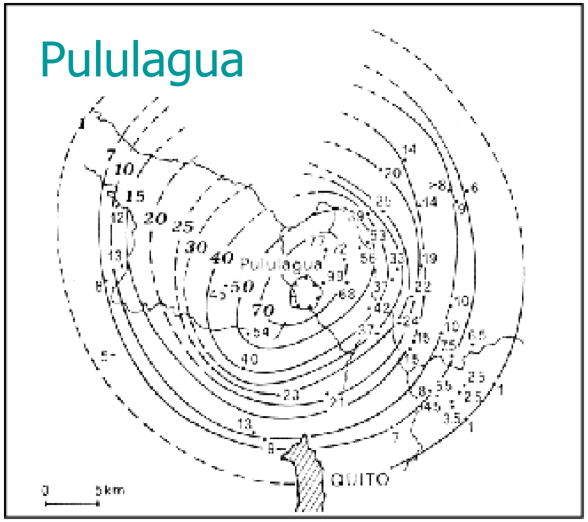
→ **Erupted Volume/Mass**, **Plume Height,**
Mass Eruption Rate (MER), Eruption Duration,
Total Grain-Size Distribution (TGSD)

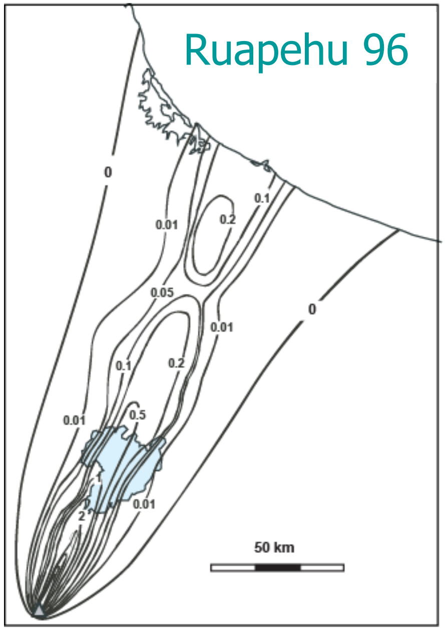


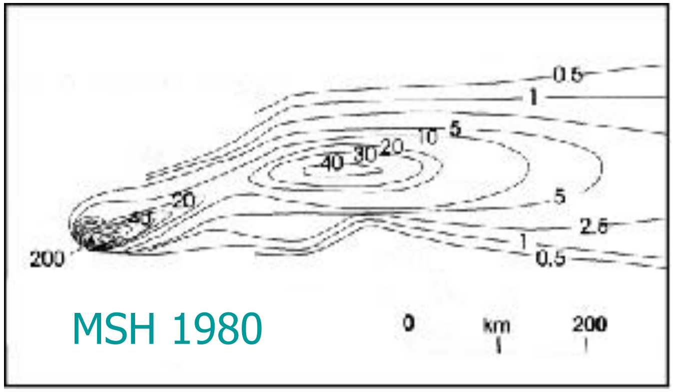

ERUPTED MASS/VOLUME

PROBLEMS:

- Non-linearity of the functions linking thickness and area
- General scarcity of data
- Lack of distal and very proximal data

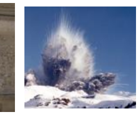






Isopach/isomass maps



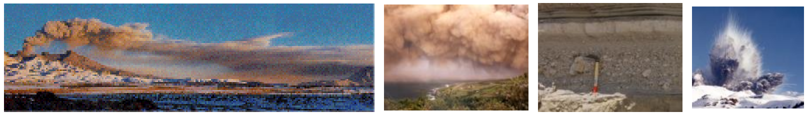
ERUPTED MASS/VOLUME

PROBLEMS:

- Non-linearity of the functions linking thickness and area
- General scarcity of data
- Lack of distal and very proximal data

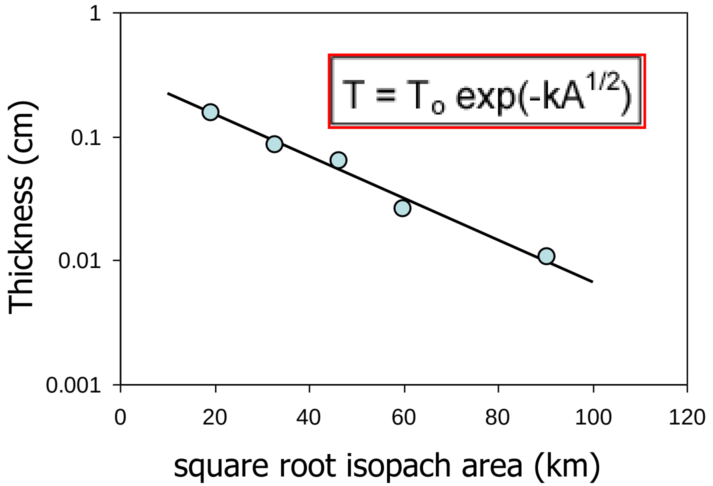
MOST COMMONLY USED MODEL:

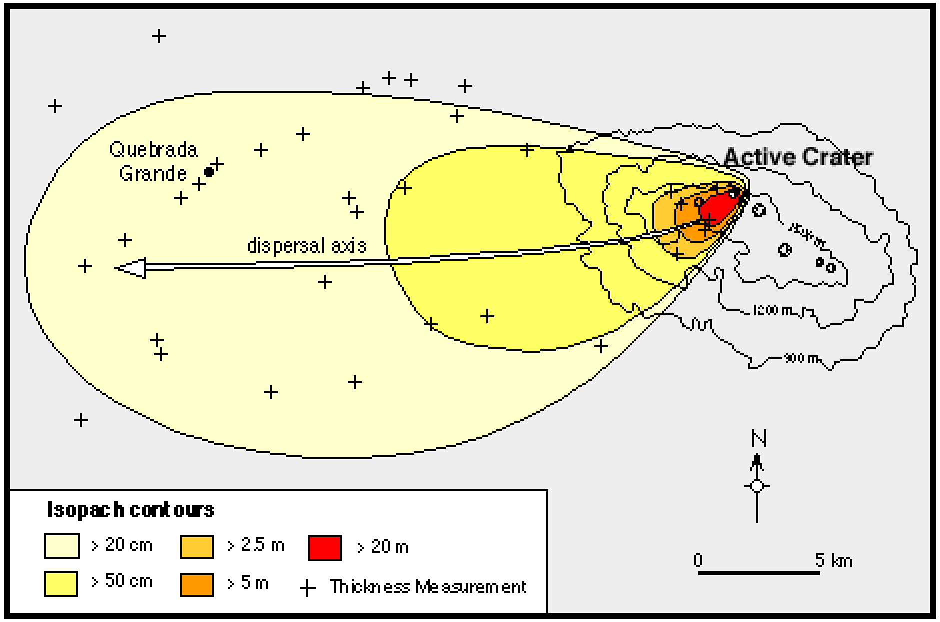
→ **EXPONENTIAL CURVE FITTING**



ERUPTED MASS/VOLUME

EXPONENTIAL FITTING

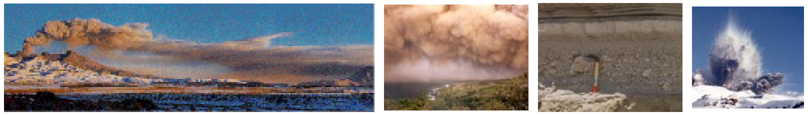




$$V = To\ 13.08\ bt^2$$

$$bt = \ln 2/(k\sqrt{\pi})$$

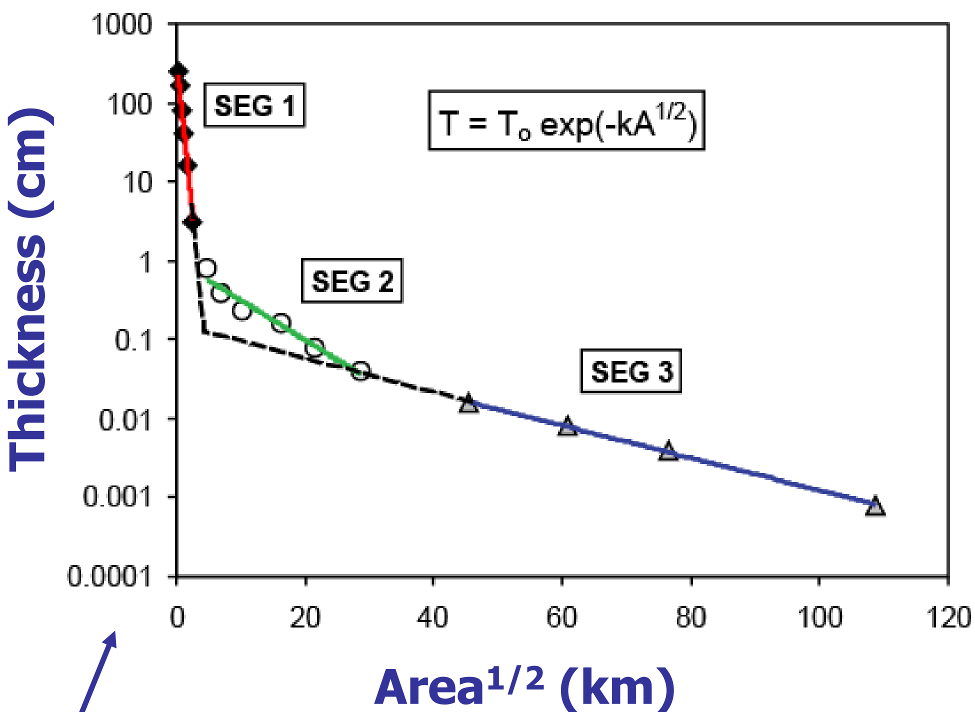
Pyle, 1989



ERUPTED MASS/VOLUME

EXPONENTIAL CURVE FITTING

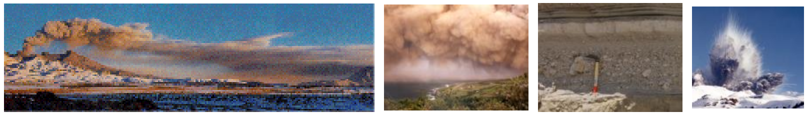
(multiple exp. segments: Fierstein and Nathenson 1992; Pyle 1995; Bonadonna and Houghton 2005)



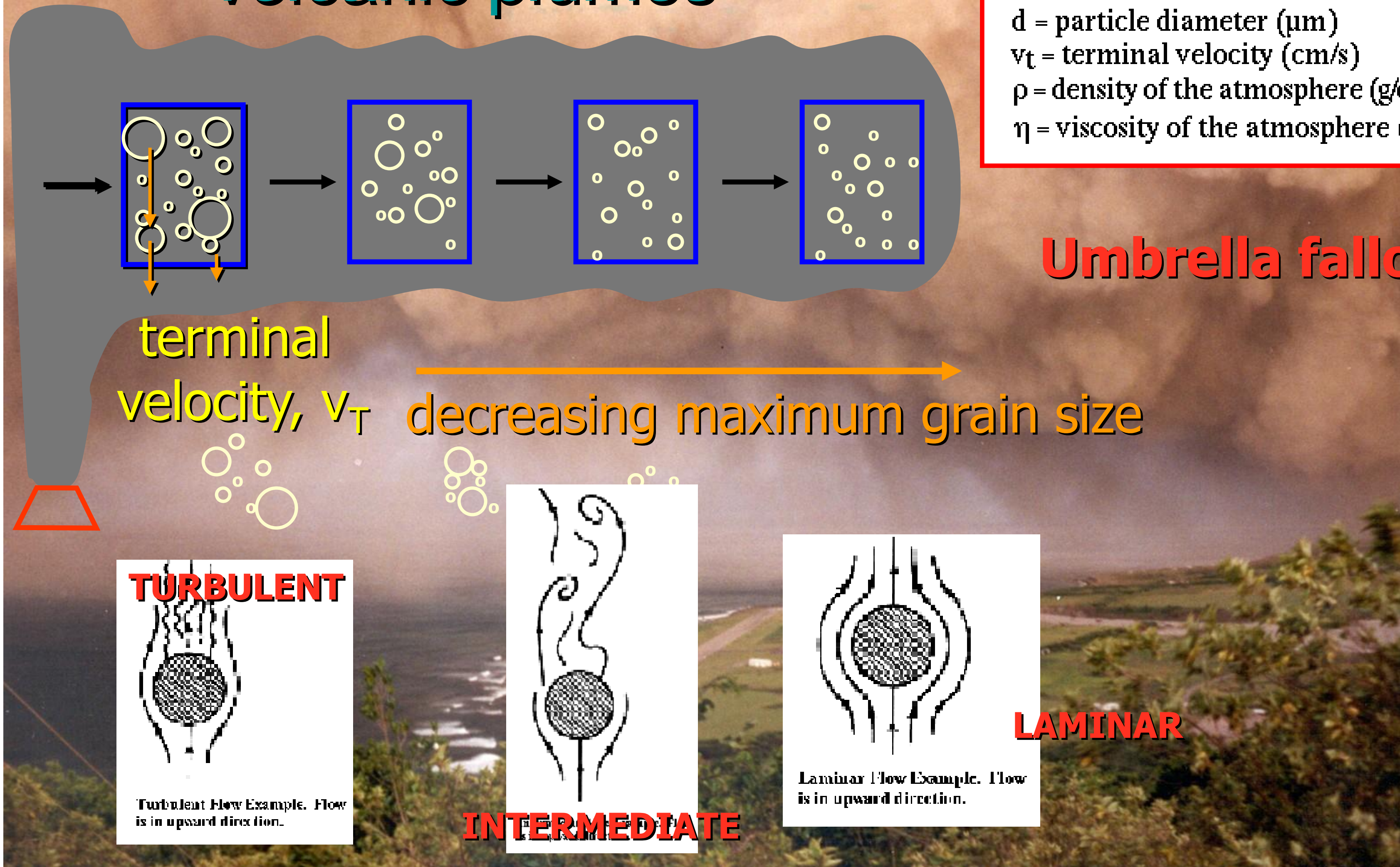
e.g. Ruapehu 1996 eruption

Empirical and theoretical support:

- **Thinning of well exposed deposits**
  e.g. Hildreth and Drake 1992 (Novarupta 1912); Fierstein and Hildreth 1992 (Quizapu 1932); Scasso et al. 1994 (Hudson 1991); Bonadonna and Houghton 2005 (Ruapehu 1996)
- **Rose 1993; Bonadonna et al. 1998**



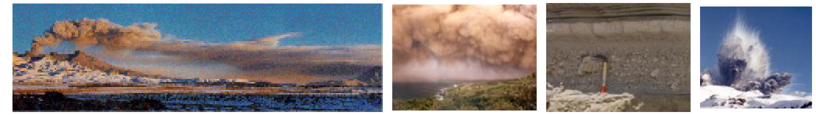
Sedimentation from volcanic plumes



Reynolds number (Re):

$$Re = (d * v_t * \rho)/\eta$$

d = particle diameter (µm)
v_t = terminal velocity (cm/s)
ρ = density of the atmosphere (g/cm^3)
η = viscosity of the atmosphere (g/cm-s)



Sedimentation from volcanic plumes

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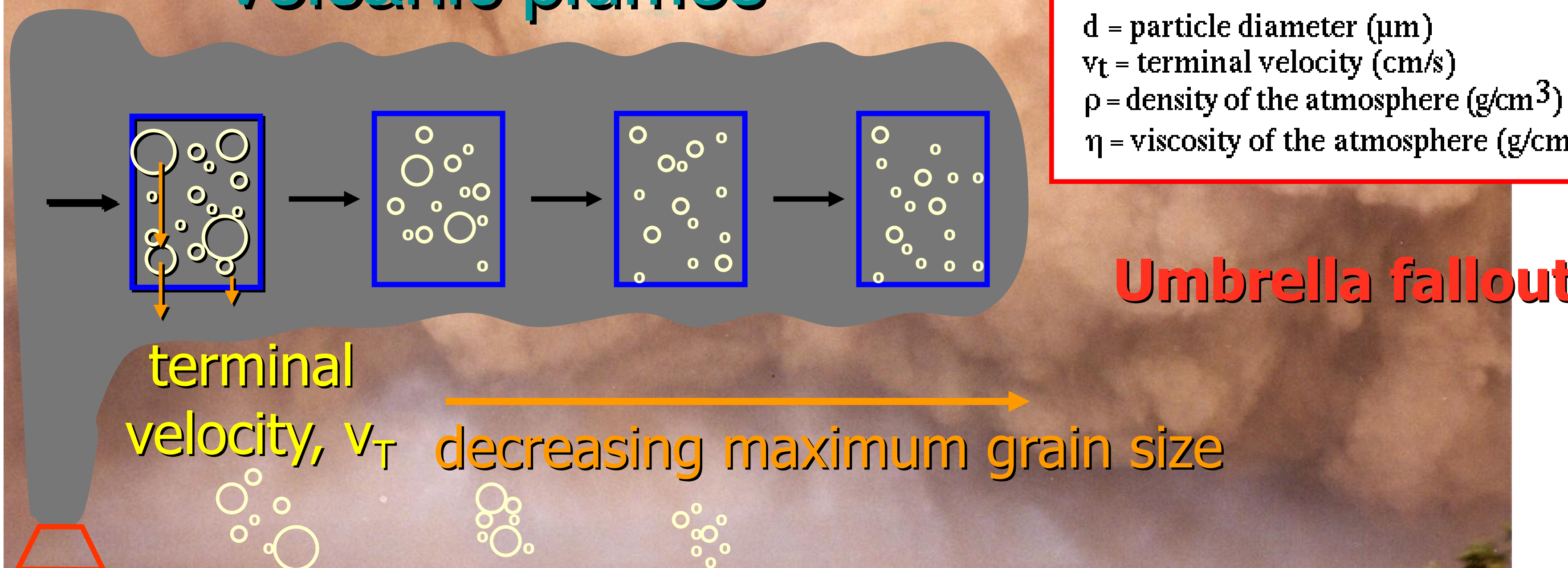
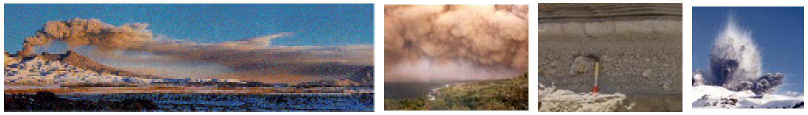


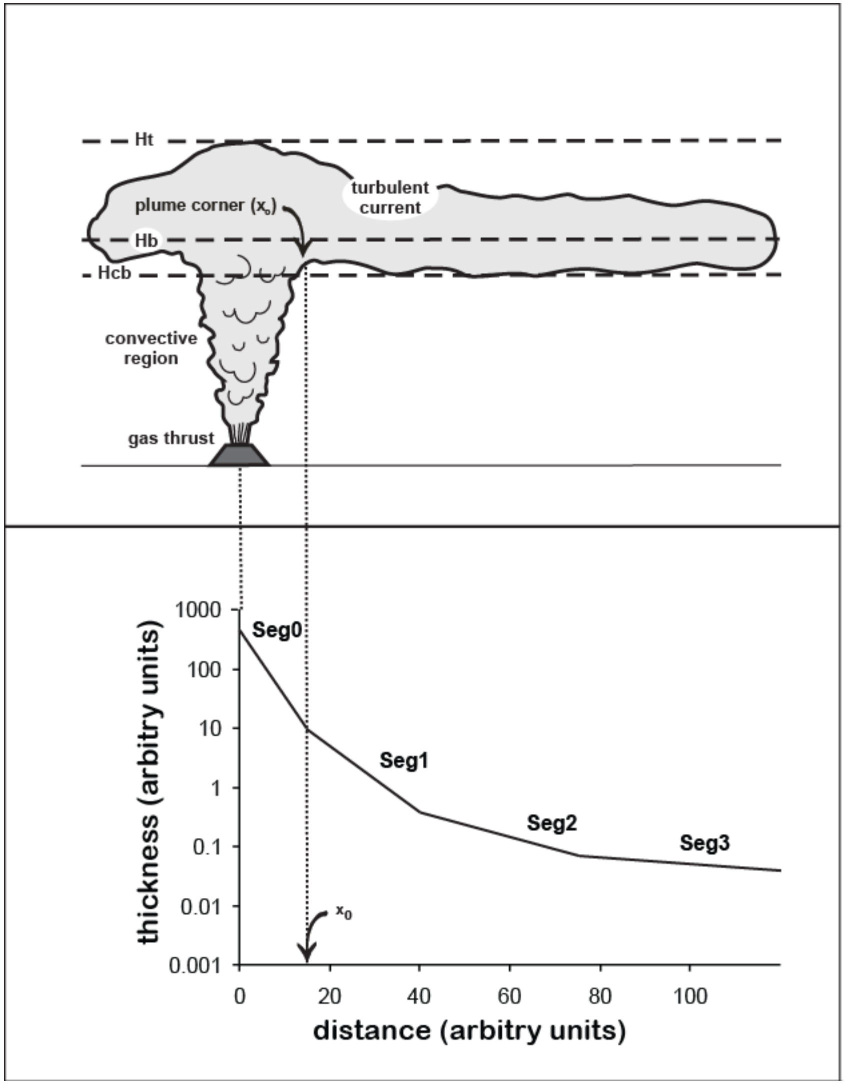
$V_t \approx (3.1\, g\rho d/\sigma)^{1/2}$ (for Reynolds numbers 500-200,000)

$V_t \approx d(4\rho^2 d^2/225\mu\sigma)^{1/3}$ (for Reynolds numbers 6-500)

$V_t \approx (g\rho d^2/18\mu)$ (for Reynolds numbers <6)



Sedimentation from strong plumes



Segment 0: sedimentation from plume margins

Segment 1: high-Re particle settling

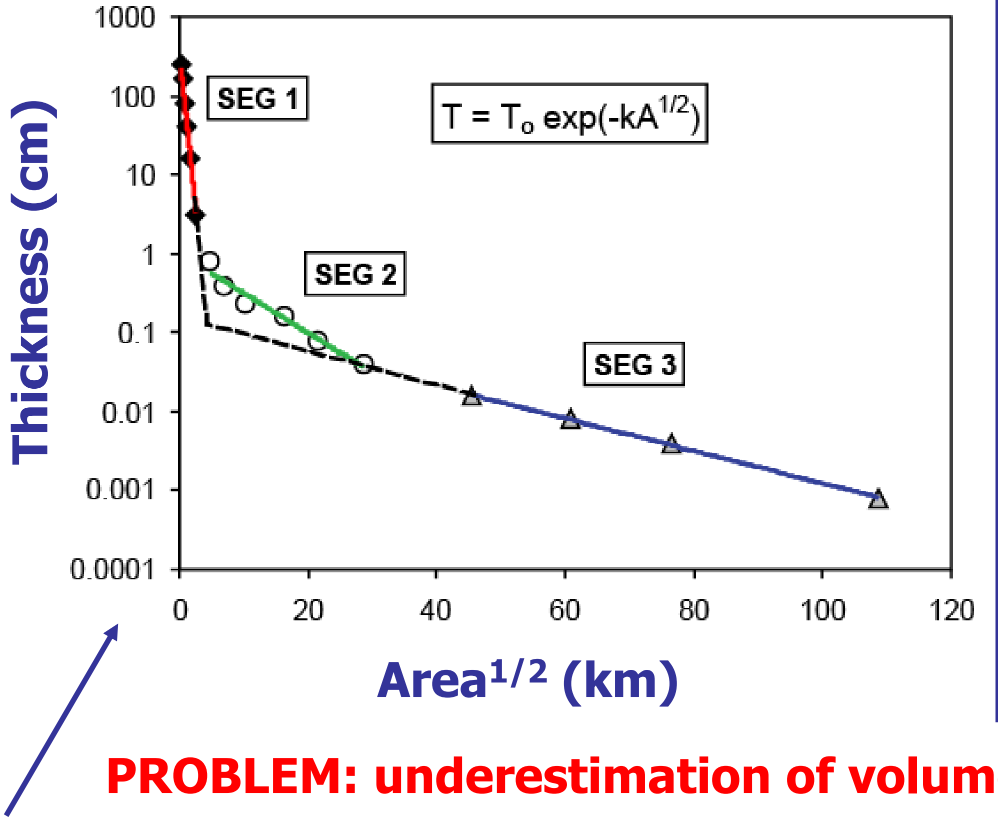
Segment 2: intermediate-Re particle

Segment 3: low-Re particle settling




ERUPTED MASS/VOLUME

EXPONENTIAL CURVE FITTING

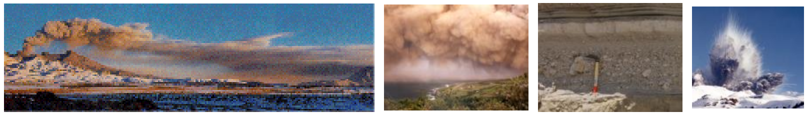


Variations:

- **1 exp. segment (Pyle 1989)**
- **2 exp. segments** (Fierstein and Nathenson 1992, Pyle 1995)
- **>2 exp. segments** (Bonadonna and Houghton 2005)
- **One proximal isopach line** (Legros 2000)
- **Missing distal data** (Sulpizio 2005)
- **Thickness measurements** (Burden et al. 2013)

PROBLEM: underestimation of volume in case of missing distal data

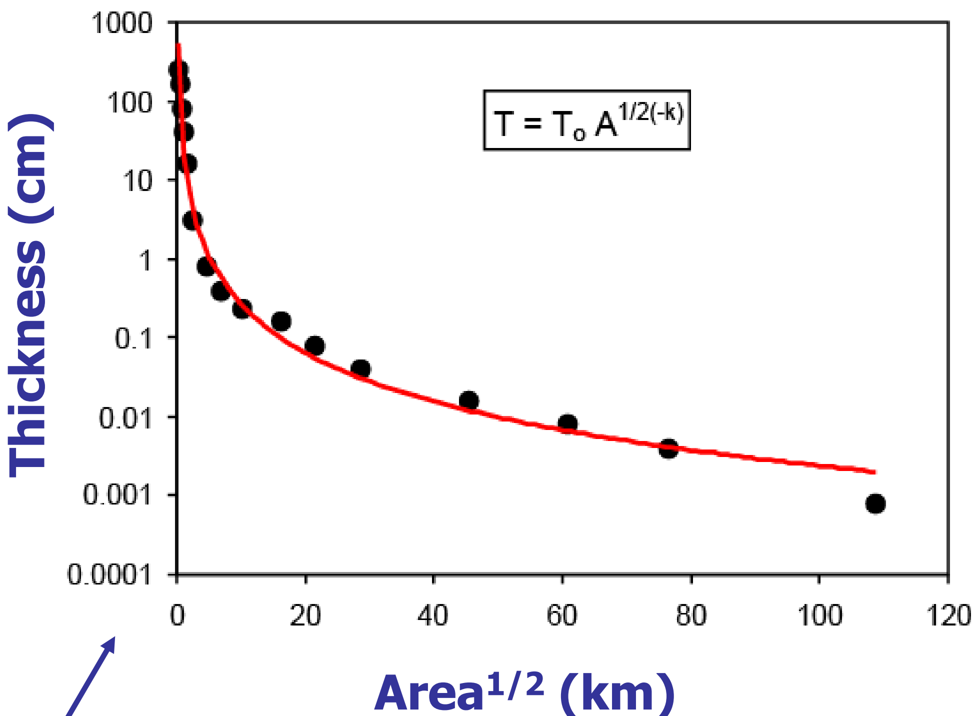
e.g. Ruapehu 1996 eruption



ERUPTED MASS/VOLUME

POWER-LAW CURVE FITTING

(Bonadonna and Houghton 2005)

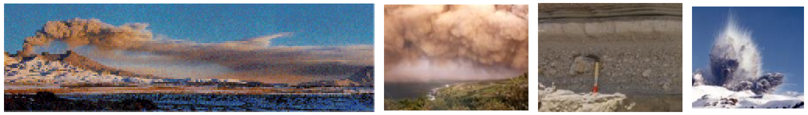


e.g. Ruapehu 1996 eruption

Empirical and theoretical support:

- **Thinning of well exposed deposits**
  e.g. Hildreth and Drake 1992 (Novarupta 1912); Fierstein and Hildreth 1992 (Quizapu 1932); Scasso et al. 1994 (Hudson 1991); Bonadonna and Houghton 2005 (Ruapehu 1996)
- **Rose 1993; Bonadonna et al. 1998**

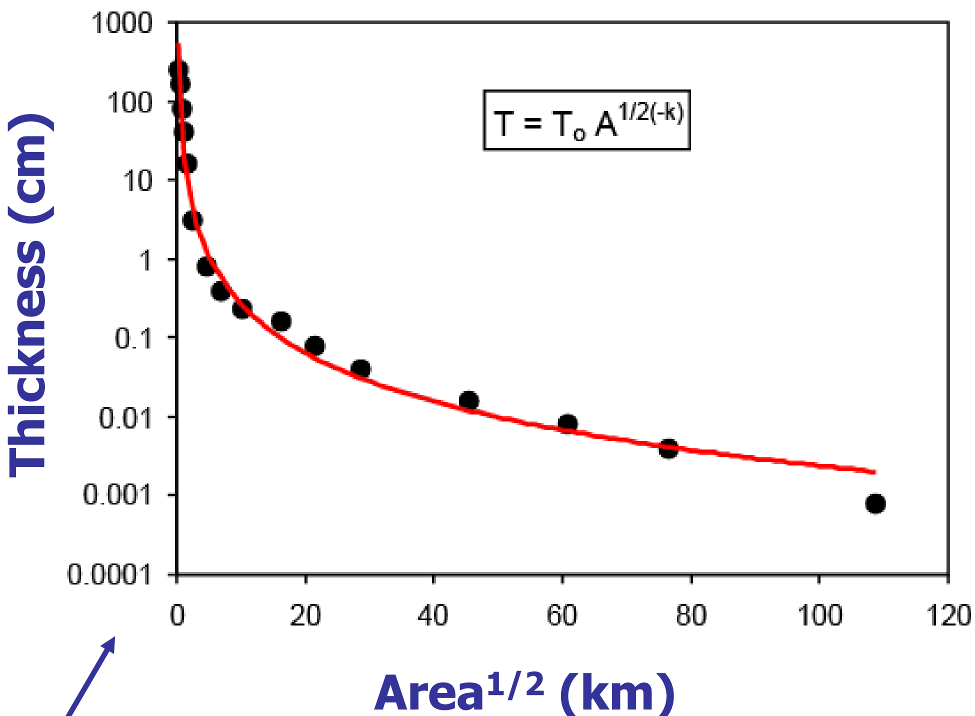
PROBLEM: choice of integration limits



ERUPTED MASS/VOLUME

POWER-LAW CURVE FITTING

(Bonadonna and Houghton 2005)



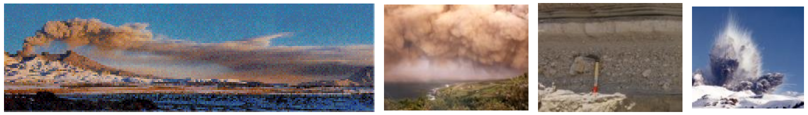
$$Volume = \frac{2T_{pl}}{2-k}\left(C^{(2-k)} - B^{(2-k)}\right)$$

k>2 → rapid thinning → small deposit → Volume sensitive to prox integration limit

k<2 → gradual thinning → large deposit → Volume sensitive to distal integration limit

e.g. Ruapehu 1996 eruption

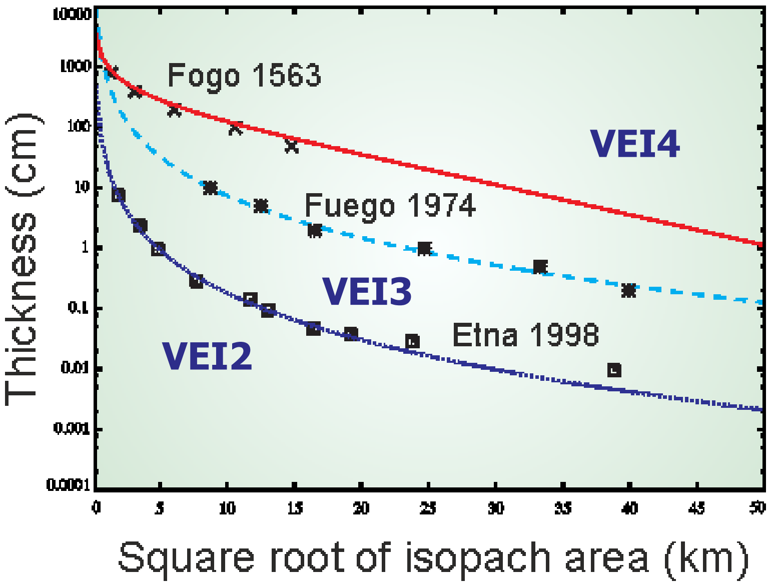
PROBLEM: choice of integration limits



ERUPTED MASS/VOLUME

WEIBULL CURVE FITTING

(Bonadonna and Costa 2012)



$$V = \int_0^\infty T\,\mathrm{d}A = 2\int_0^\infty T(x)\,x\,\mathrm{d}x = \frac{2\theta\lambda^2}{n}\left[1 - \mathrm{e}^{-\left(\frac{x}{\lambda}\right)^n}\right]_0^\infty = \boxed{\frac{2\theta\lambda^2}{n}}$$

λ: characteristic decay length scale of deposit thinning (km);
θ: thickness scale (cm)




ERUPTED MASS/VOLUME

PROS and CONS

Exponential fit:

☺ Easy to apply

☺ Can be integrated between 0 and ∞

☹ Underestimates the volume in case of missing distal and/or proximal data

☹ The choice of exponential segments can be arbitrary and biased by available data

Power-law fit:

☺ Easy to apply

☺ Can better describe the natural gradual deposit thinning with distance

☹ Cannot be integrated between 0 and ∞ (→ difficult choice of extremes of integration)

Weibull fit:

☺ Can better describe the natural gradual deposit thinning with distance

☺ Can be integrated between 0 and ∞

☹ Is of more complex application and needs > 3 data points (because it has 3 free par.)



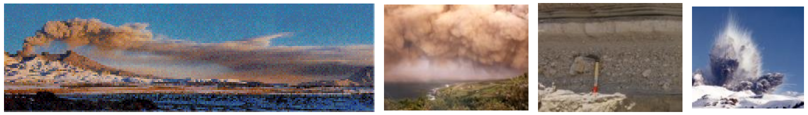

ERUPTED MASS/VOLUME

Free softwares for the determination of erupted mass/volume:

Biass S, Bonadonna C, Houghton BF (2019) A step-by-step evaluation of empirical methods to quantify eruption source parameters from tephra-fall deposits, Journal of Applied Volcanology, "TephraFit" **https://github.com/e5k/TephraFits**

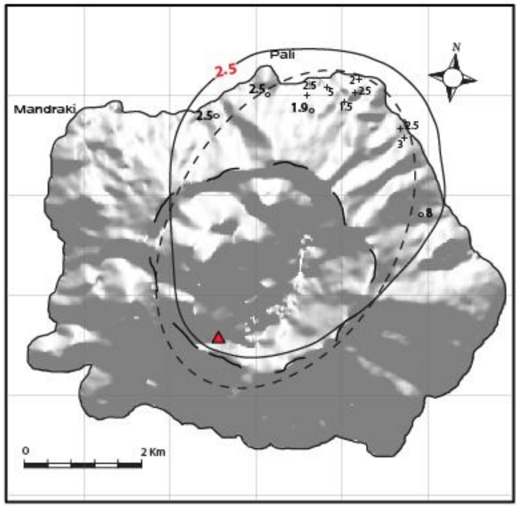
Biass S, Bagheri GH, Aeberhard W and Bonadonna C (2014). TError: towards a better quantification of the uncertainty propagated during the characterization of tephra deposits, Statistics in Volcanology "TError" **https://github.com/e5k/TError**

Daggitt M, Pyle D, Mather TA (2014) AshCalc – a new tool for the comparison of the exponential, power-law and Weibull models of tephra deposition, Journal of Applied Volcanology "AshCalc", **https://vhub.org/resources/ashcalc**

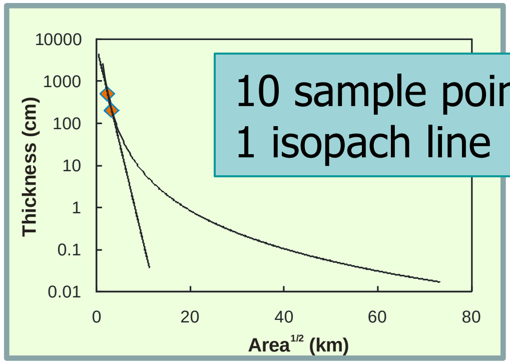


ERUPTED MASS/VOLUME

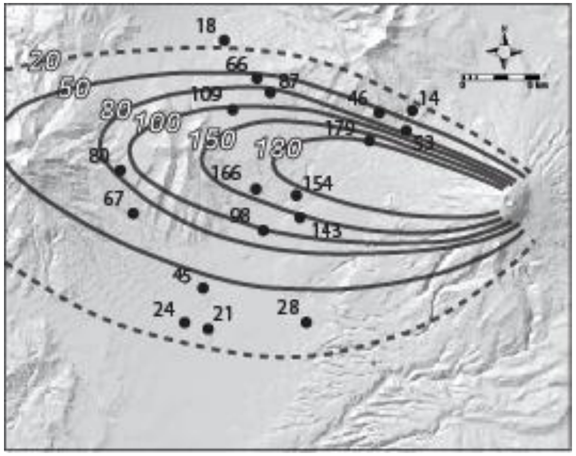
Lower Pumice, Nisyros (Greece)



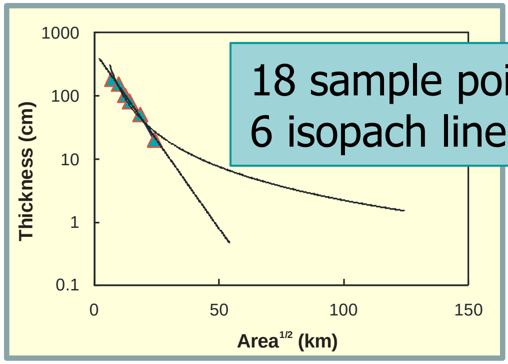
Longchamp et al. , 2011



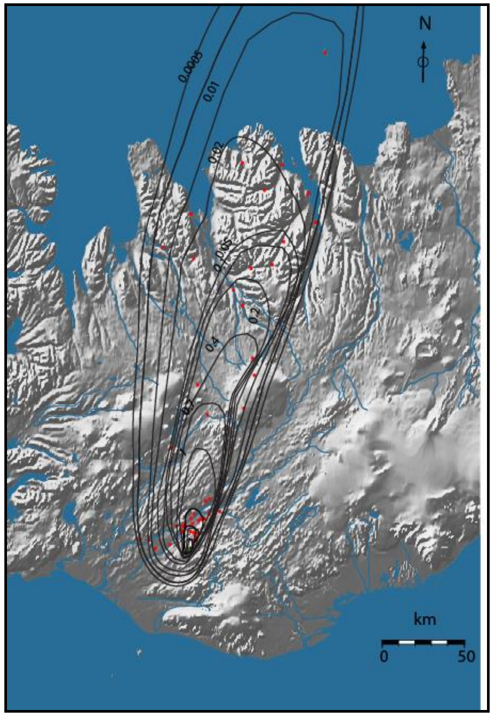
Layer 3 (820BP), Cotopaxi (Ecuador)



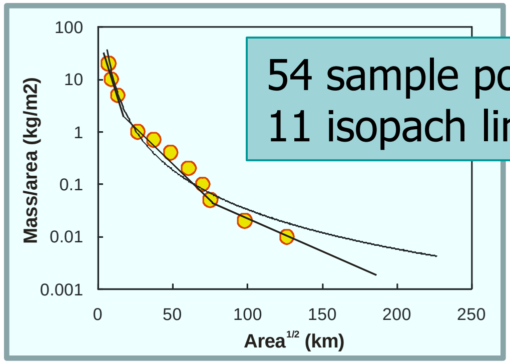
Biass and Bonadonna , 2011

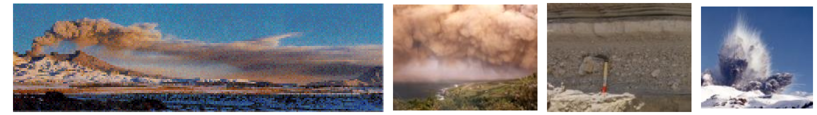


Hekla 2000 (Iceland)



Data from Árnasson (2002), Haraldsson (2001), Haraldsson et al. (2002) and Larsen (unp. data)





ERUPTED MASS/VOLUME

Lower Pumice, Nisyros (Greece)

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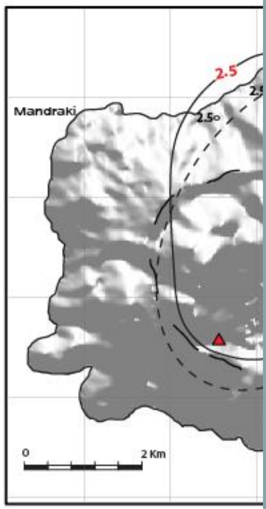

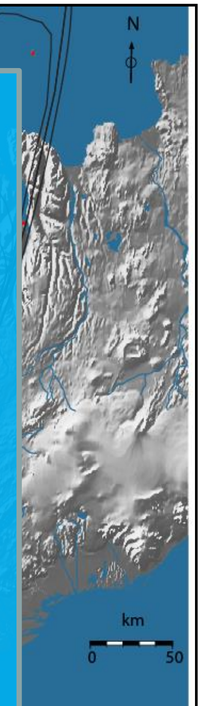

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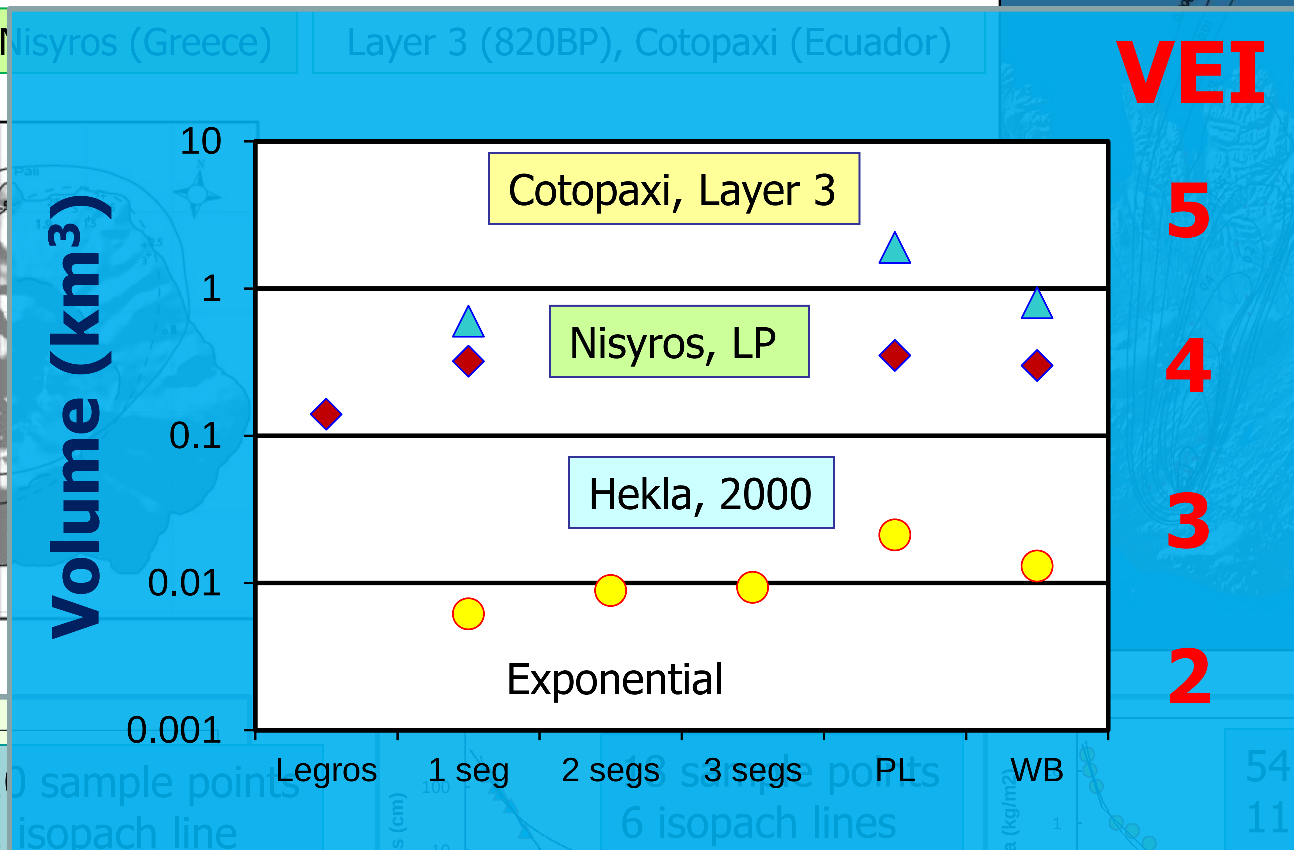
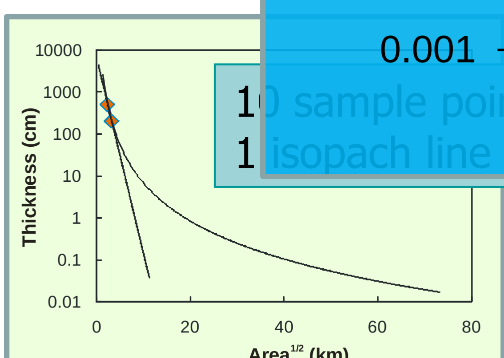
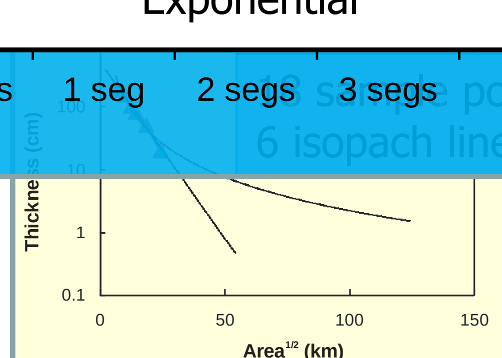
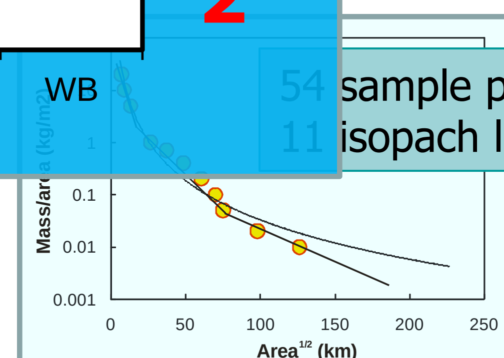




10 sample points
1 isopach line

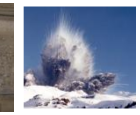


10 sample points
6 isopach lines



54 sample points
11 isopach lines



Volcanic Explosivity Index (VEI)

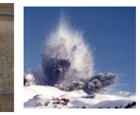
Volume of ejecta: both tephra and igmbrite if present

Criteria VEI:	0	1	2	3	4	5	6	7	8
DESCRIPTION	non-explosive	small	moderate	mod-large	large	very large-------------------------------------			
VOLUME OF EJECTA (M^3)	$<10^4$	10^4-10^6	10^6-10^7	10^7-10^8	10^8-10^9	10^9-10^{10}	10^{10}-10^{11}	10^{11}-10^{12}	$>10^{12}$
(TSUYA CLASSIFICATION)+	(I)	(II-III)	(IV)	(V)	(VI)	(VII)	(VIII)	(IX) ---------	
COLUMN HEIGHT (KM)*	<0.1	0.1-1	1-5	3-15	10-25	>25--			
QUALITATIVE DESCRIPTION	---"gentle, effusive"--		-----"explosive"-------		-----"cataclysmic, paroxysmal, colossal"----------------------				
				"severe, violent, terrific"--					
CLASSIFICATION		----------"Strombolian"-			-------"Plinian"--				
	-------"Hawaiian"------		-------------"Vulcanian"-------------			--------"Ultraplinian"--------------------------			
DURATION (hours)	------------------------<1--------------				------------>12--				
of continuous blast			-----------------1-6----------------						
				--------------------6-12---------------					
CAVW MAX EXPLOSIVITY**	--lava flows-	---------------explosion or nuée ardente--							
		------phreatic---------..							
	-----dome or mudflow---								
TROPOSPHERIC INJECTION	negligible	minor	moderate	substantial--					
STRATOSPHERIC INJECTION	none	none	none	possible	definite	significant-----------------------------------			

+If all eruptive products were pyroclastic ejecta
*For VEI's 0-2, uses km above crater; for VEI's 3-8, uses km above sea level.
**The most explosive activity indicated for the eruption in the Catalogue of Active Volcanoes

Newhall and Self 1982



Volcanic Explosivity Index (VEI)

Volume of ejecta: both tephra and igmbrite if present

TABLE I. The Volcanic Explosivity Index[a]

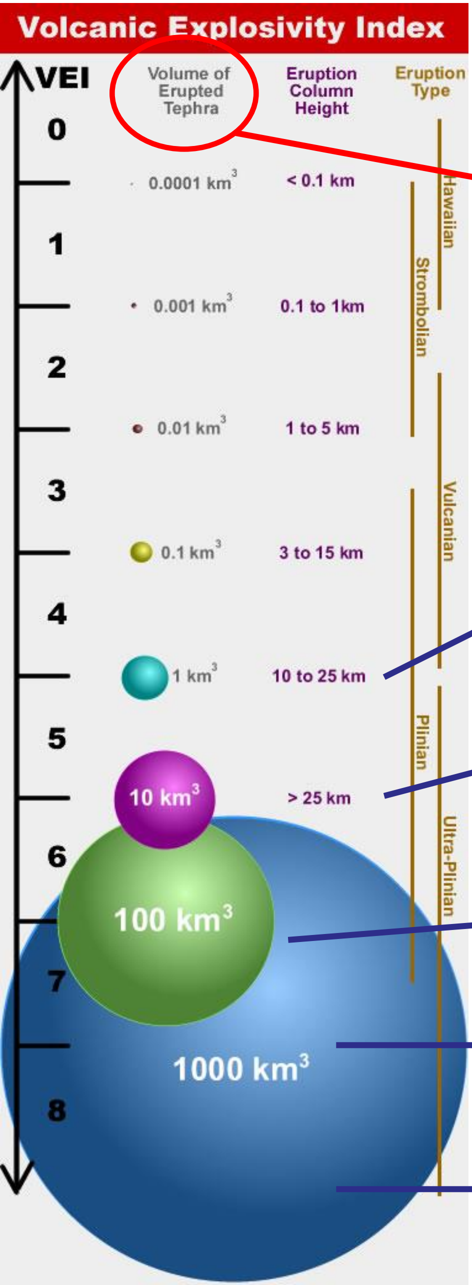
VEI index	*0*	*1*	*2*	*3*	*4*	*5*	*6*	*7*	*8*
General description	Nonexplosive	Small	Moderate	Moderate–large	Large	Very large			
Qualitative description	Gentle	Effusive	← Explosive →		← Cataclysmic, paroxysmal →				
Maximum erupted volume of tephra (m^3)	10^4	10^6	10^7	10^8	10^9	10^{10}	10^{11}	10^{12}	10^{13}
Eruption cloud column height (km)	<0.1	0.1–1	1–5	3–15	10–25	>25			

[a] Adapted from Newhall and Self, 1982.

- ❖ In practice, modern eruptions defined by plume height
- ❖ Ancient eruptions defined by maximum erupted volume
- ❖ Not useful for effusive (lava) eruptions, which by default get VEI of 0, 1




Volcanic Explosivity Index (VEI)



Volume of ejecta: both tephra and igmbrite if present

Mono-Inyo Craters (past 5,000 years)

Mount St Helens (May 18, 1980) (bulk deposit =1.3 km^3, of which tephra is 1.2 km^3) **(VEI 5)**

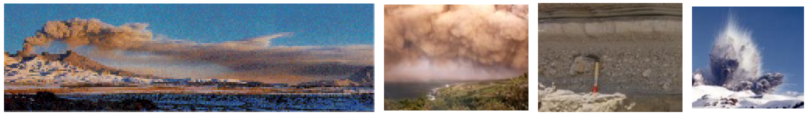
Pinatubo (June 1991) (bulk deposit ~ 11 km^3, of which tephra is 6.5 km^3) **(VEI 6)**

Tambora (1815) (bulk deposit ~ 110 km^3, of which tephra is 101 km^3) **(VEI 7)**

Long Valley Caldera (760ka) (bulk deposit ~1380 km^3, of which tephra is 300 km^3) **(VEI 8)**

Yellowstone Caldera (640ka) (bulk deposit ~1000 km^3, of which tephra is 500 km^3) **(VEI 8)**

VOGRIPA: http://www.bgs.ac.uk/vogripa/

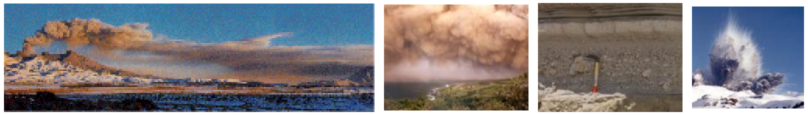


Mass vs. Volume of Volcanic Eruptions

Deposit volume depends on deposit vesicularity → deposits with same mass but different vesicularity have different volume

To alleviate this problem → deposit volume is converted to Dense Rock Equivalent volume

This is an estimate of the volume of dense, unvesiculated magma that produced deposit



Dense Rock Equivalent (km^3)

Mass deposit = Mass magma

$$\rho_{deposit} \times V_{deposit} = \rho_{magma} \times V_{magma}$$

$$V_{magma}\,(DRE) = \rho_{dep} \times V_{dep} / \rho_{magma}$$




Dense Rock Equivalent (km^3)

$V_{deposit} = 1.8\ km^3$

$\rho_{deposit} = 0.9\ g\ cm^{-3}$

$\rho_{magma} = 2300\ kg\ m^{-3}$

V_{magma} (DRE) ?




EXERCISE 2

DETERMINATION OF ERUPTED VOLUME:

https://vhub.org/resources/531/download/Eruption_Magnitude_and_Deposit_Volume.pdf



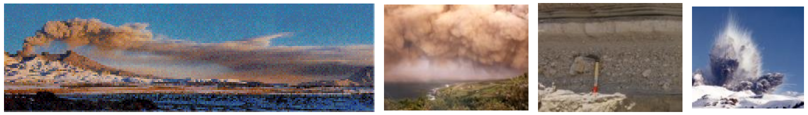

MAIN QUESTION:

How well can we constrain explosive eruptions?

→ Erupted Volume/Mass, **Plume Height**,
Mass Eruption Rate (MER), Eruption Duration,
Total Grain-Size Distribution (TGSD)

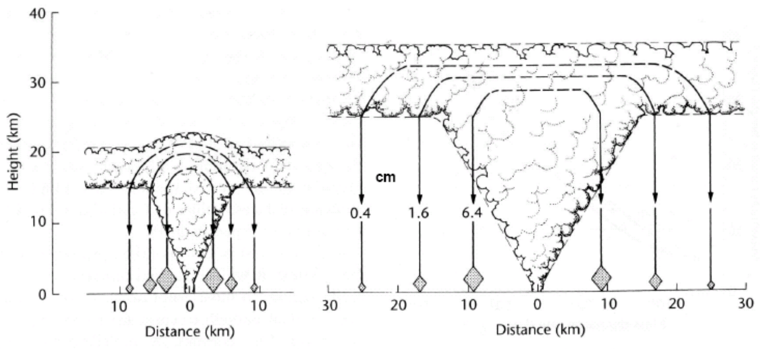
PLUME HEIGHT

- Direct observations
- Satellite images (thermal infrared data; cloud shadow clinometry; cloud stereoscopy)
- Radar, Lidar, Thermal cameras, …
- Empirical methods (Carey and Sparks 1986; Rossi et al. 2018)
- Analytical methods (inversion on grainsize)



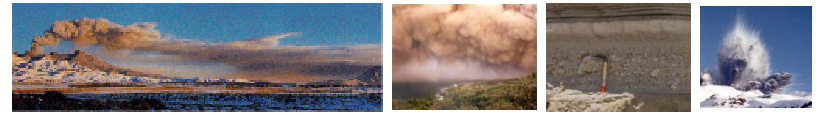
COLUMN HEIGHT

Method of Carey and Sparks (1986)



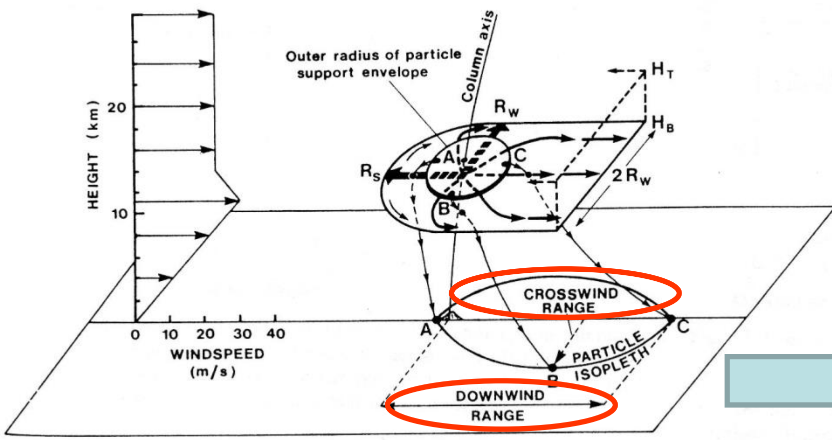
a more intense plume carries a clast higher:

- **it takes longer to fall**
- **the wind carries it further away**
- **it falls further from the vent**



COLUMN HEIGHT

Method of Carey and Sparks (1986)

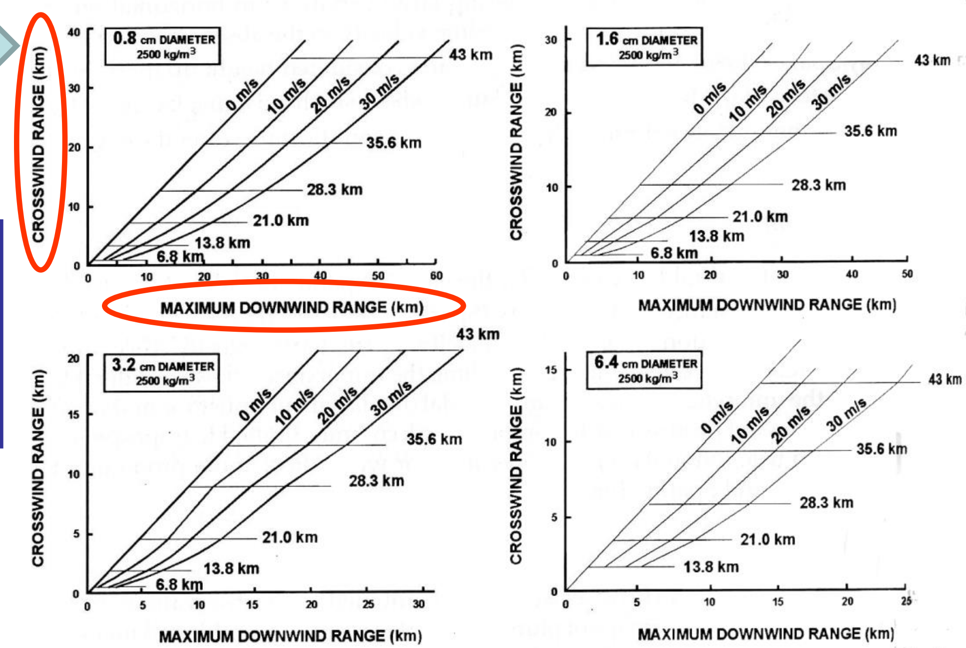


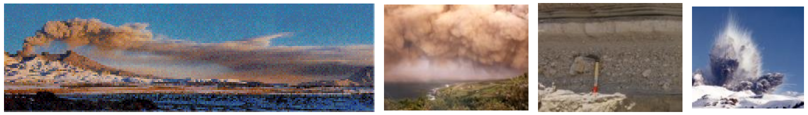
a more intense plume carries a clast higher:

- it takes longer to fall
- the wind carries it further away
- it falls further from the vent

Main assumptions:

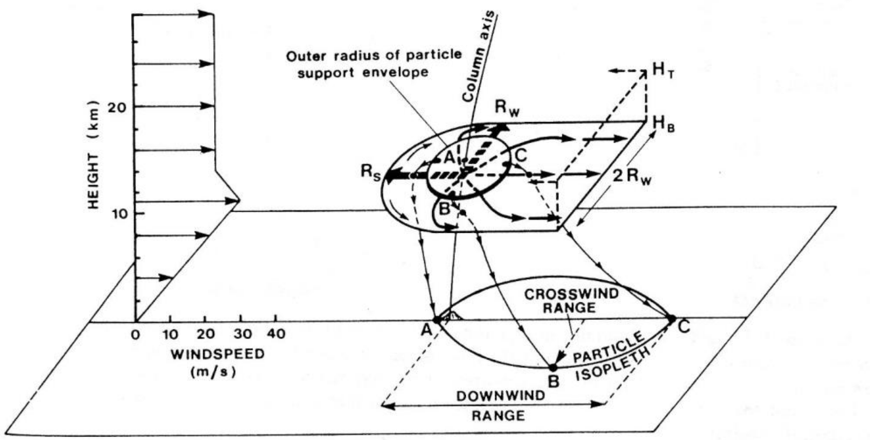
- Sustained strong plumes
- Monotone plume vertical velocity
- Fixed wind profile → trop.=11km

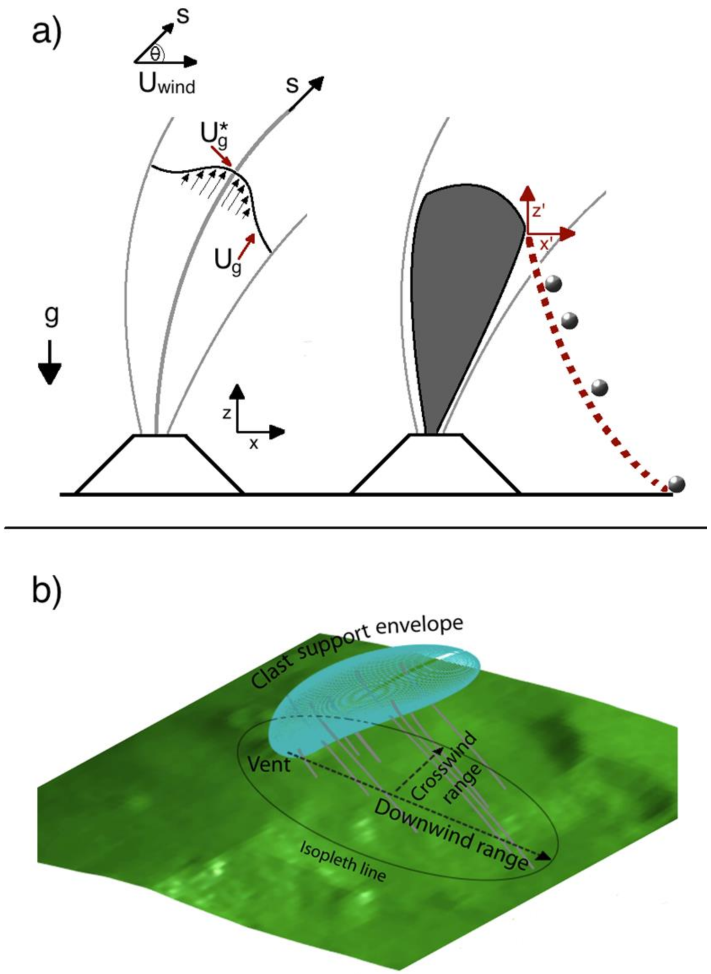




COLUMN HEIGHT

Method of Rossi et al. (2018)

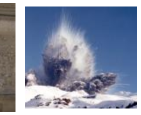




Implementations:

- Better description of plume dynamics (e.g. better description of gravitational spreading) and of the effect of wind on plume rise
- Wind profiles used to construct the nomograms are the same as for C&S86 but real 3D wind profiles can also be considered
- Effect of particle shape on settling velocity



COLUMN HEIGHT

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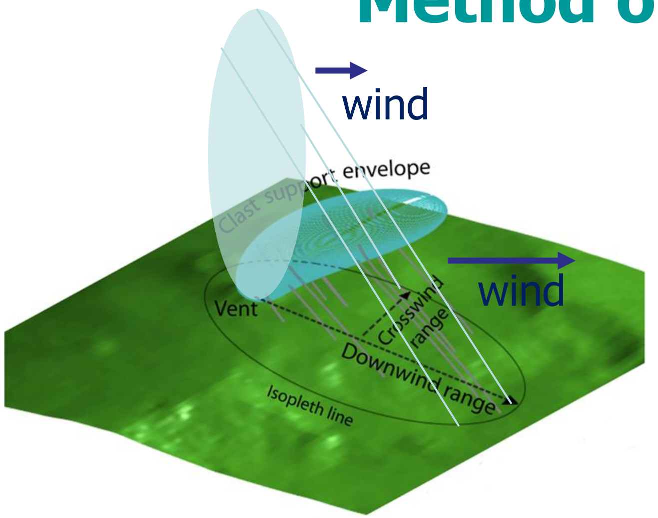

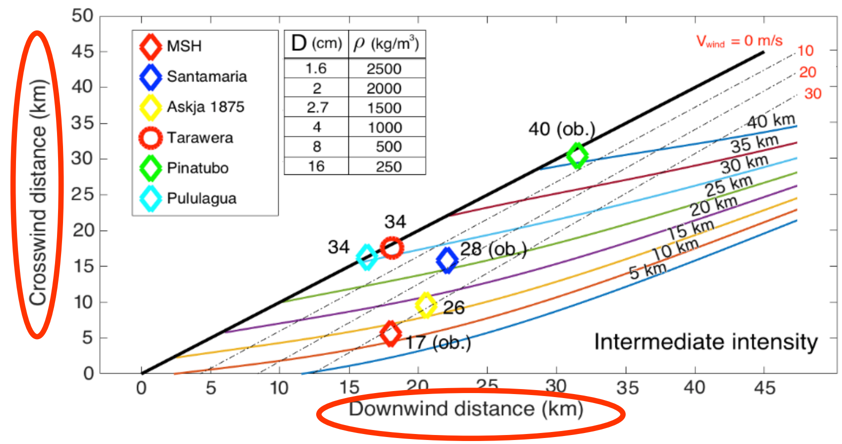

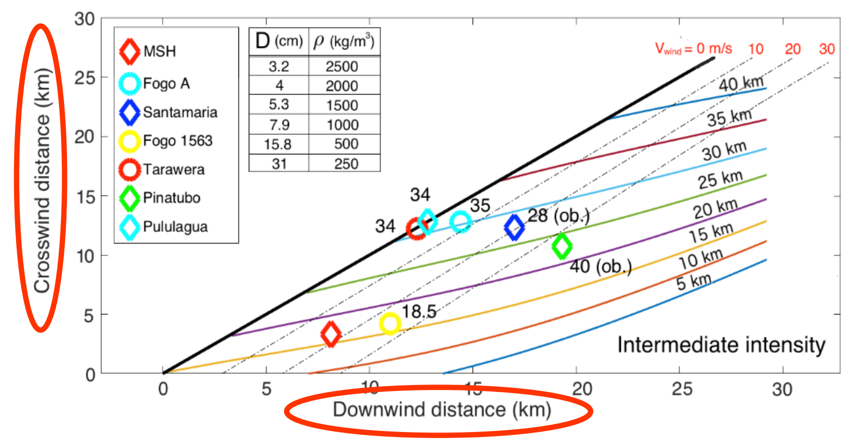

Implementations:

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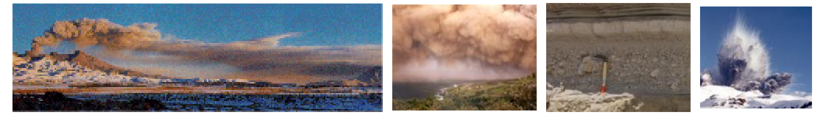
- Wind profiles used to construct the nomograms are the same as for C&S86 but real 3D wind profiles can also be considered
- Effect of particle shape on settling velocity

Relation between height and CW range is not longer univocal for a given DW range (effect of wind)



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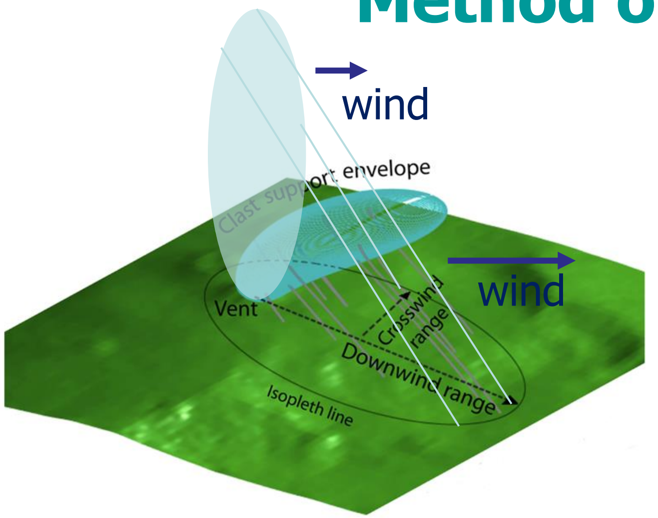
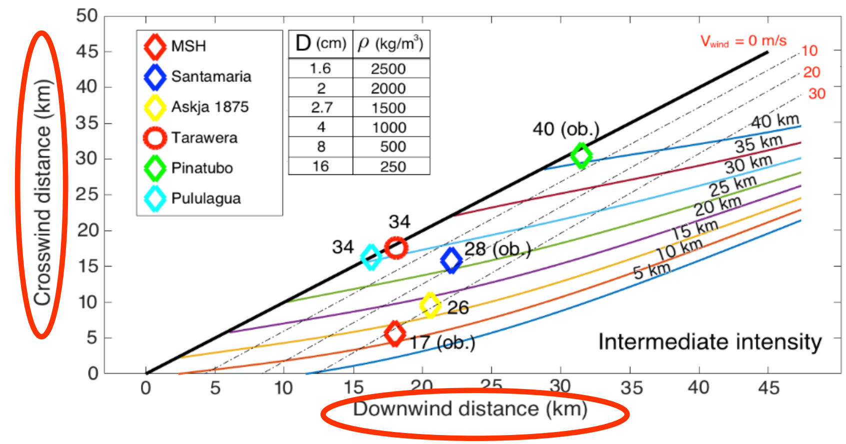
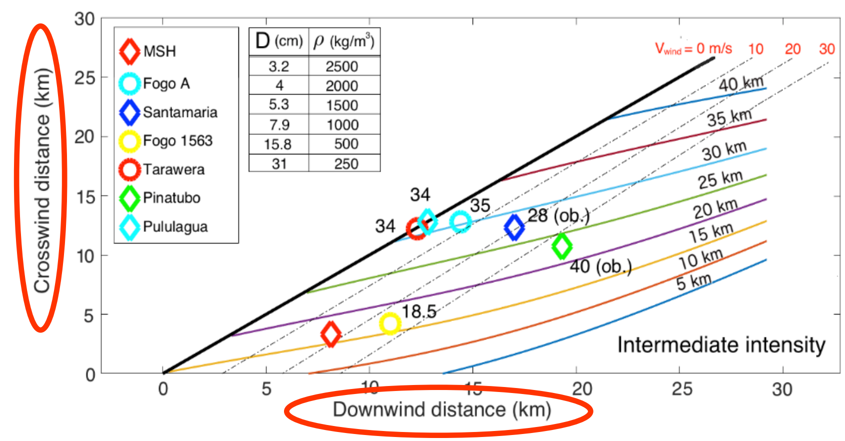






Implementations:

- Better description of plume dynamics (e.g. better description of gravitational spreading) and of the effect of wind on plume rise
- Wind profiles used to construct the nomograms are the same as for C&S86 but real 3D wind profiles can also be considered
- Effect of particle shape on settling velocity

Resulting heights are generally lower than those predicted by C&S86 because also the wind is accounted for




MAIN QUESTION:

How well can we constrain explosive eruptions?

How well can we constrain Eruption Source Parameters (ESP)?

→ **Erupted Volume/Mass, Plume Height,**

Mass Eruption Rate (MER), Eruption Duration,

Total Grain-Size Distribution (TGSD)




MASS ERUPTION RATE

$$\mathbf{Ht}\ (\mathrm{km}) = \mathbf{K}\ [\mathbf{MER}(\mathrm{Kg/s})]^{1/4}$$

Wilson and Walker 1987

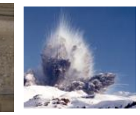
K= 0.236

- Central vent
- Silicic eruptions
- No wind
- Dry atmosphere

K= 0.295
K= 0.247
K= 0.244

Fontana Lapilli (Masaya, Nicaragua); Etna 2001, 2002 (Italy)
Wehrmann et al. 2006, Scollo et al. 2007, Andronico et al. 2008



MASS ERUPTION RATE

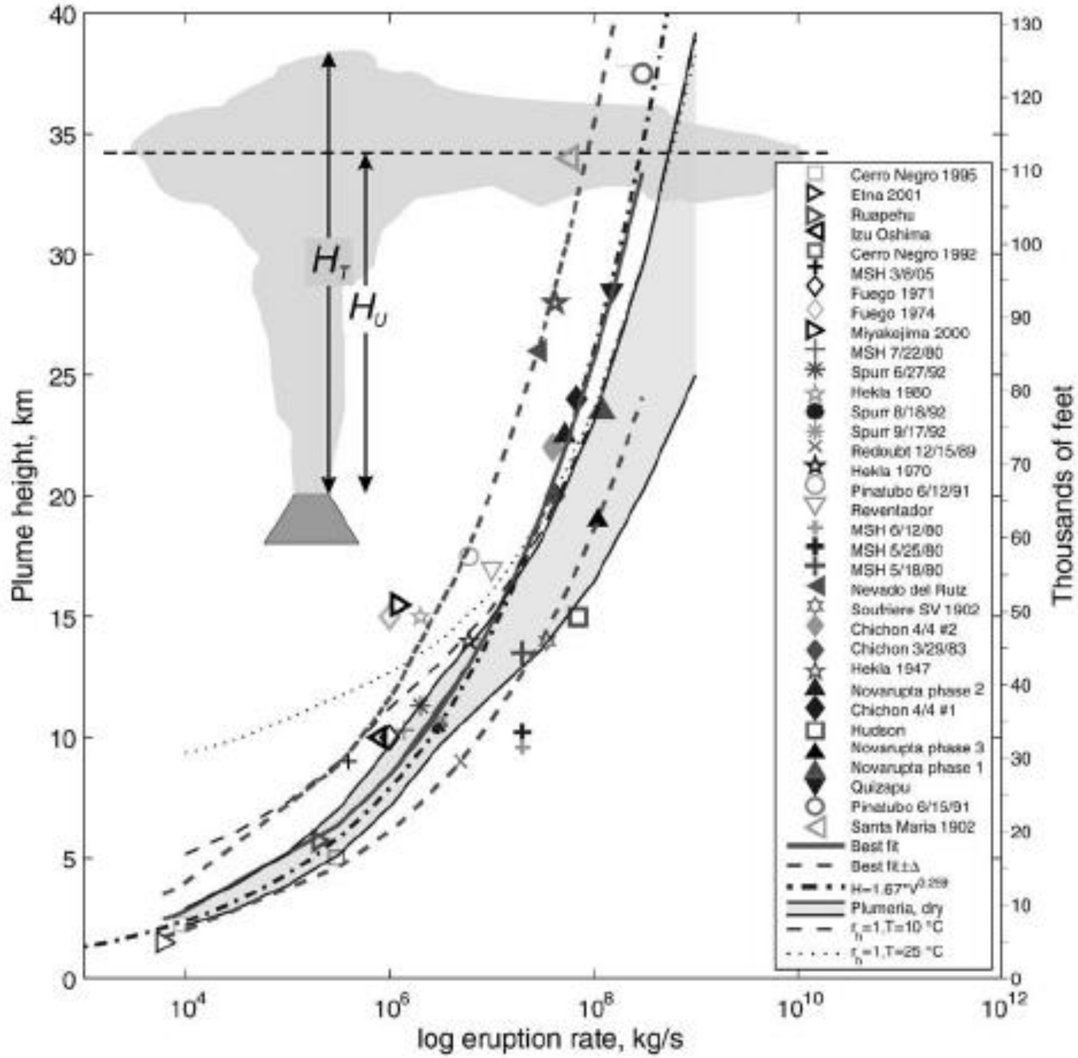
Ht (km)= **K [MER**(Kg/s)**]$^{1/4}$**

Wilson and Walker 1987

➢ Mass eruption rate from empirical equations can only be constrained within a factor of 10

$Ht = 2.00\, Q^{0.241}$

Mastin et al. 2009

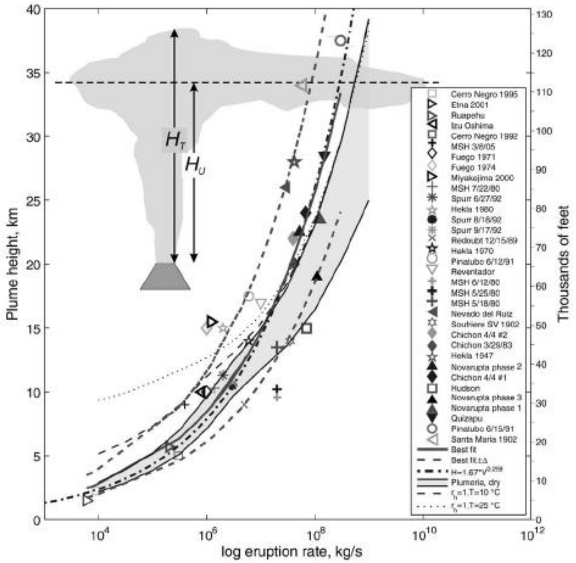


Q=volumetric flow rate (m^3/s)
MER=Q x magma density
= Q x 2500 kg m^{-3}




MASS ERUPTION RATE

$$MER = \pi \frac{\rho_{a0}}{g'} \left(\frac{\alpha^2 \overline{N}^3}{10.9} H^4 + \frac{\beta^2 \overline{N}^2 \overline{v}}{6} H^3 \right)$$



H Plume height

$\overline{N}$ Average buoyancy frequency across the plume height

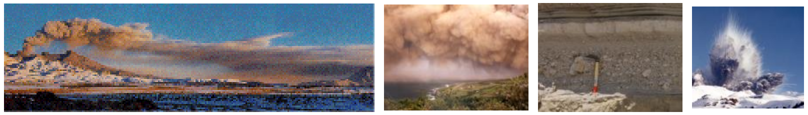
$\overline{v}$ Average wind velocity across the plume height

α Radial entrainment coefficient (best fit: 0.1)

β Wind entrainment coefficient (best fit: 0.5)

ρ_{a0} Density of the atmosphere

Degruyter and Bonadonna (2012)



MASS ERUPTION RATE

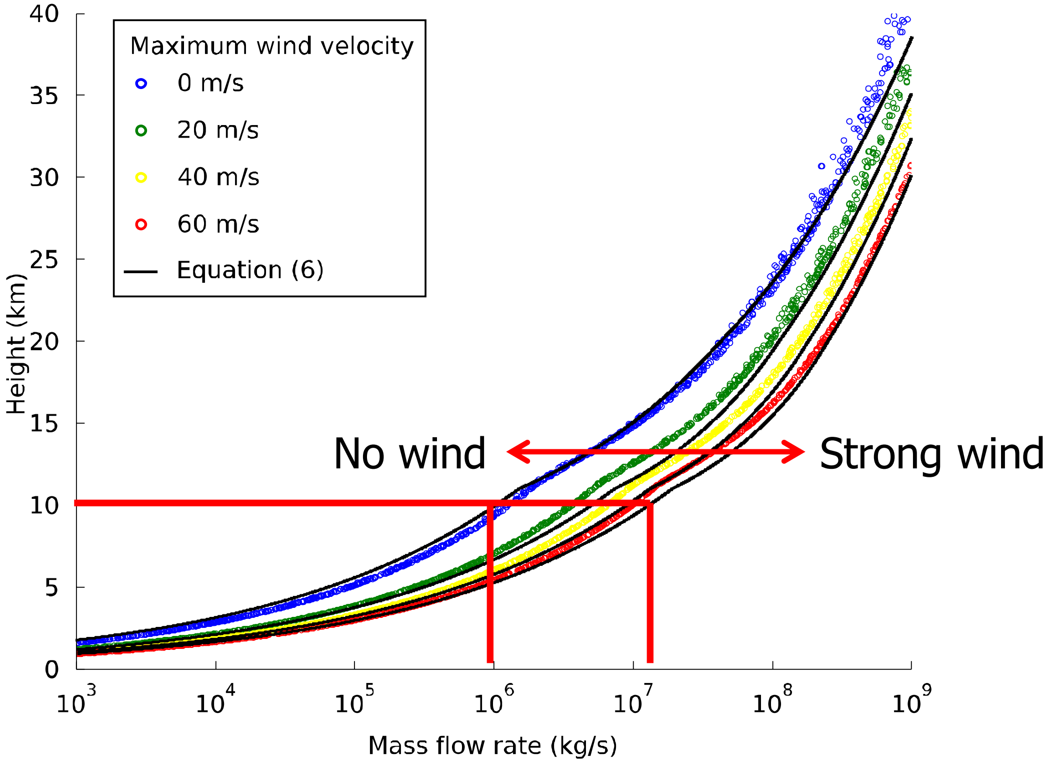
$$MER = \pi \frac{\rho_{a0}}{g'}\left(\frac{\alpha^2 \overline{N}^3}{10.9} H^4 + \frac{\beta^2 \overline{N}^2 \overline{v}}{6} H^3\right)$$

- can account for variability of both source and atmospheric conditions
- remains accurate within a factor of two compared with a 1D plume model

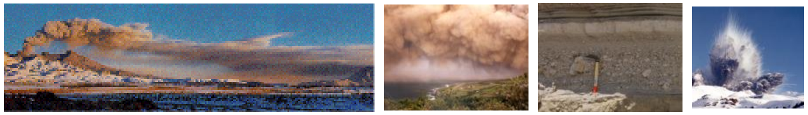
$$\Pi = \frac{\overline{N} H}{1.8\overline{v}}\left(\frac{\alpha}{\beta}\right)^2$$

wind becomes dominant if the height, buoyancy frequency, and radial entrainment are small and the wind speed and wind entrainment are large

$$\Pi << 1$$



Degruyter and Bonadonna (2012)



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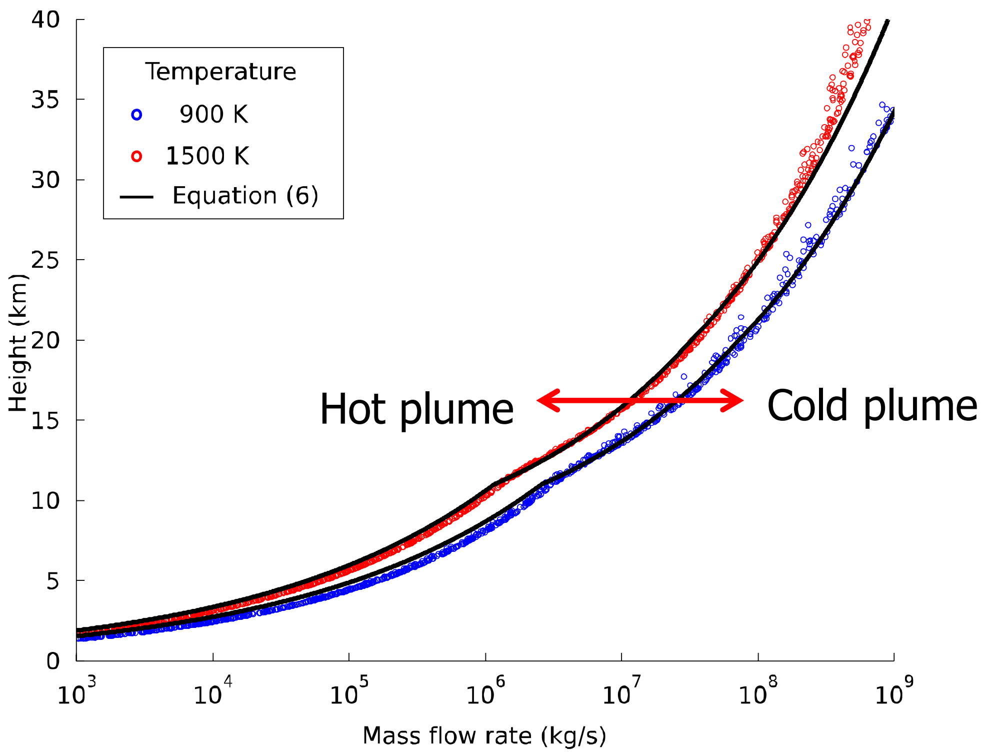
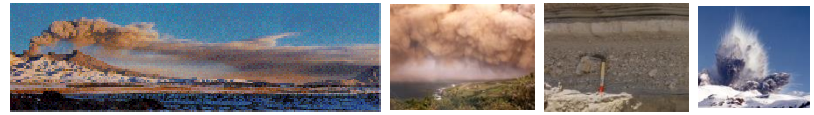


Degruyter and Bonadonna (2012)



MASS ERUPTION RATE AND DURATION

$$Ht = K\,MER^{1/4}$$

Wilson and Walker 1987

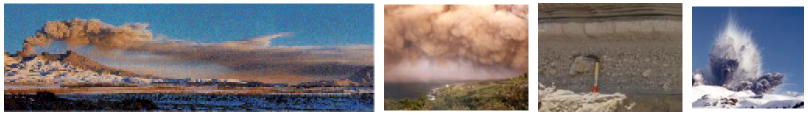
$$Ht = 2.00\,Q^{0.241} \rightarrow MER = Q \times 2500\ kg\ m^{-3}$$

Mastin et al. 2009

$$MER = \pi \frac{\rho_{a0}}{g'} \left(\frac{\alpha^2 \overline{N}^3}{10.9} H^4 + \frac{\beta^2 \overline{N}^2 \overline{v}}{6} H^3 \right)$$

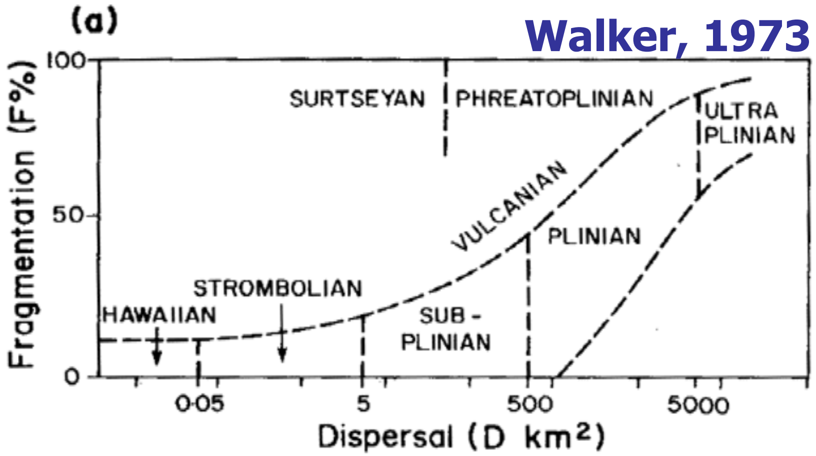
Degruyter and Bonadonna (2012)

Eruption duration (s) = Erupted Mass (kg) / MER (kg/s)

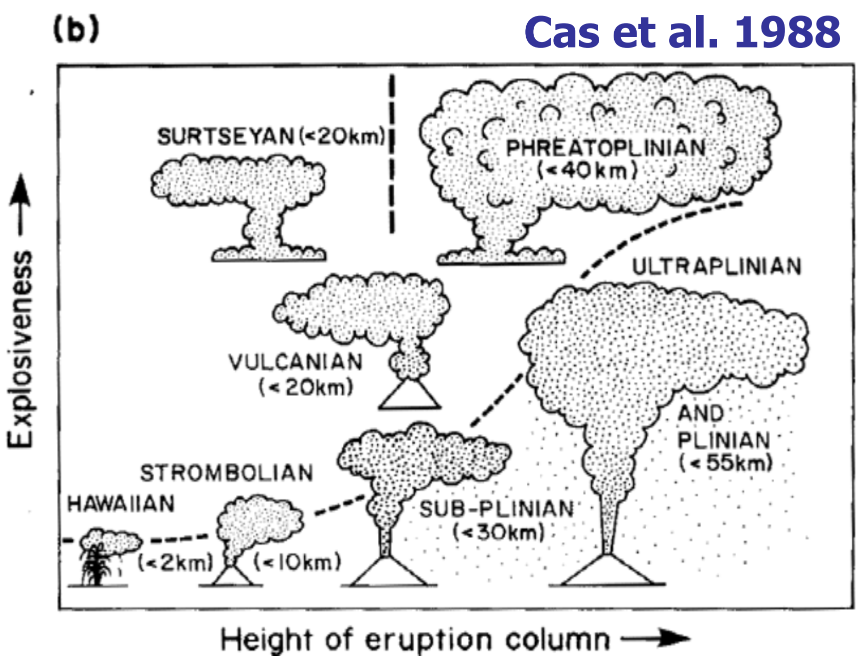


CLASSIFICATION

Walker, 1973



Cas et al. 1988

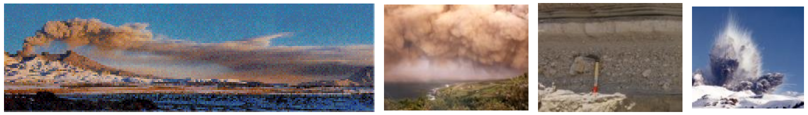


D (area of pyroclastic dispersal) → indicator of the column height.

Area enclosed by an isopach contour representing 1% of the maximum thickness (0.01 Tmax).

F (fragmentation) → indicator of the explosiveness of the eruption.

Percent of tephra <1mm, measured along an axis of dispersal where the isopach is 10% of Tmax (0.1 Tmax).

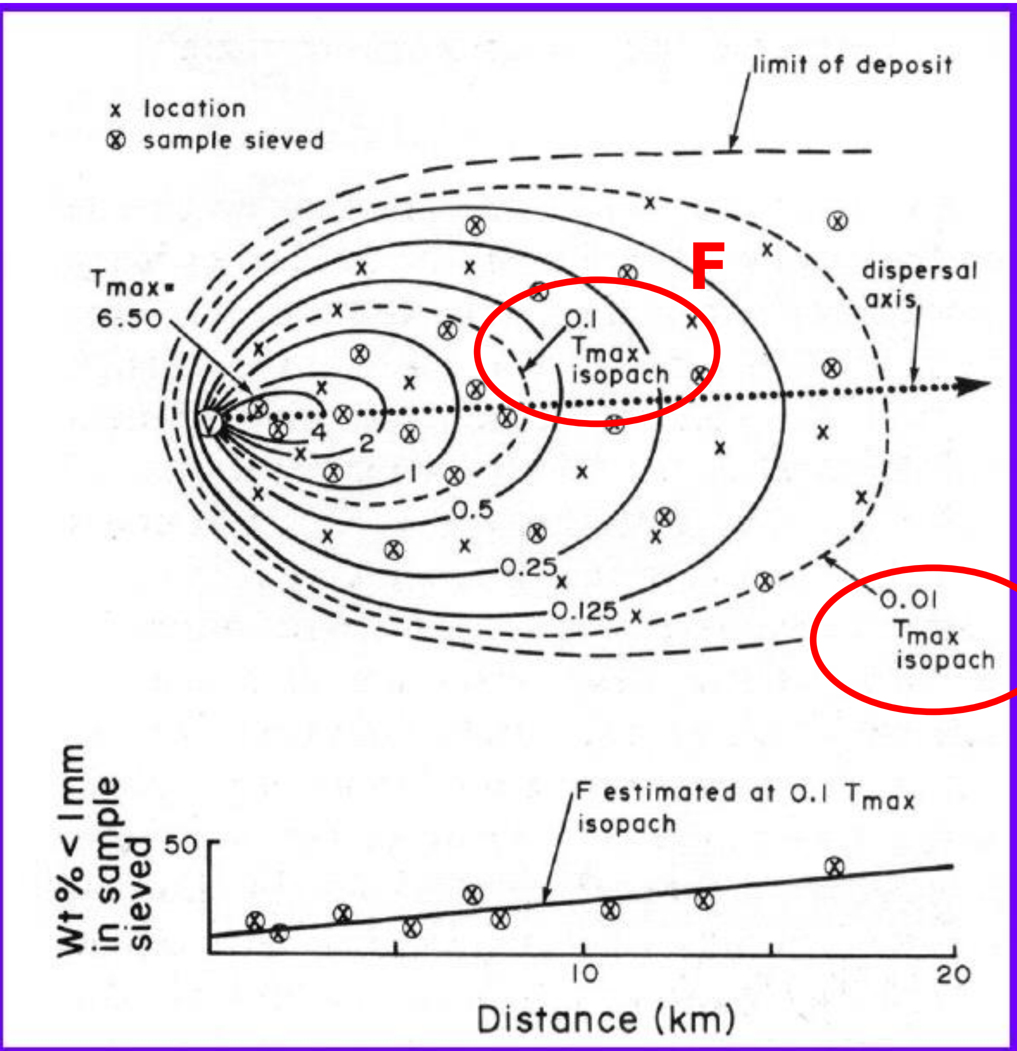


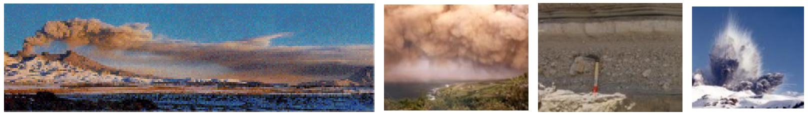
CLASSIFICATION

Fragmentation (F), a measure of how thoroughly the magma is blown apart.

It's one measure of the intensity of an explosion.

Walker, 1973

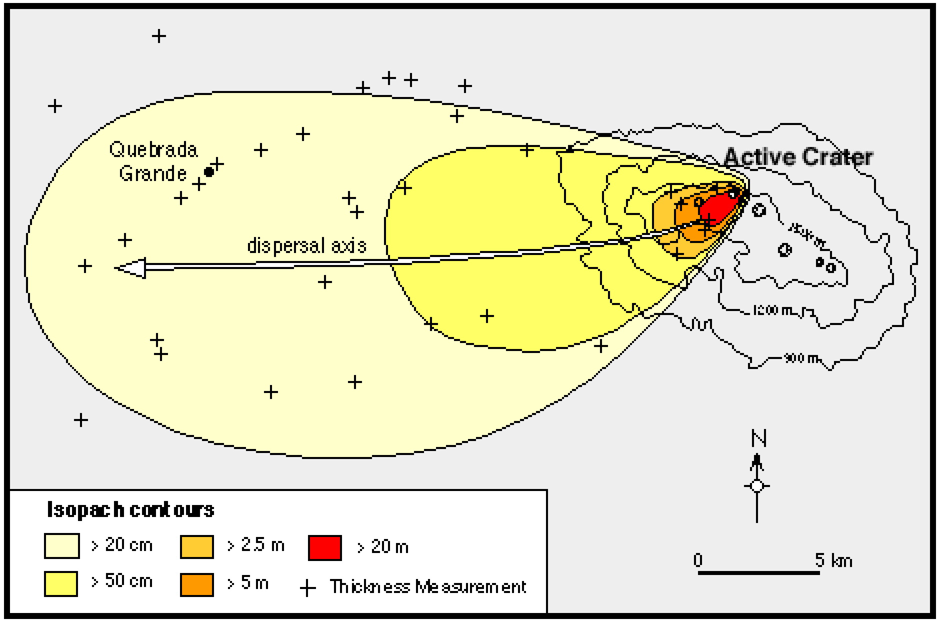


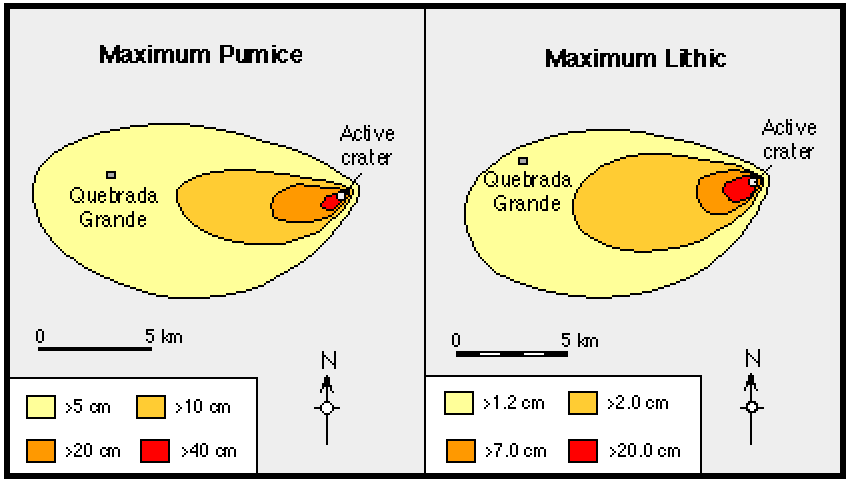


CLASSIFICATION

Pyle 1989

Classification based on the distribution of
largest clasts (bc) and on deposit thinning (bt)

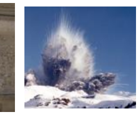




Isopach maps

Isopleth maps



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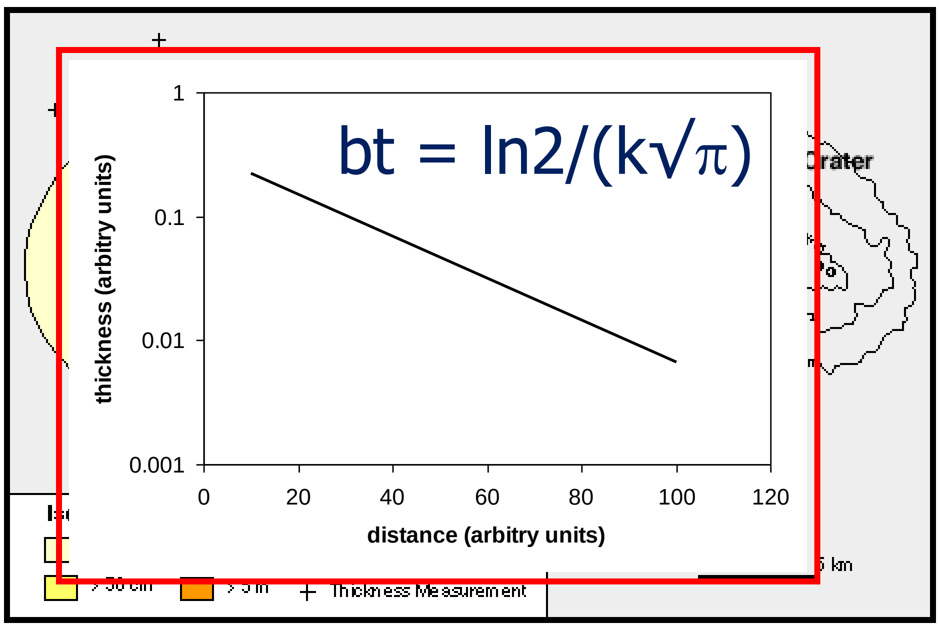
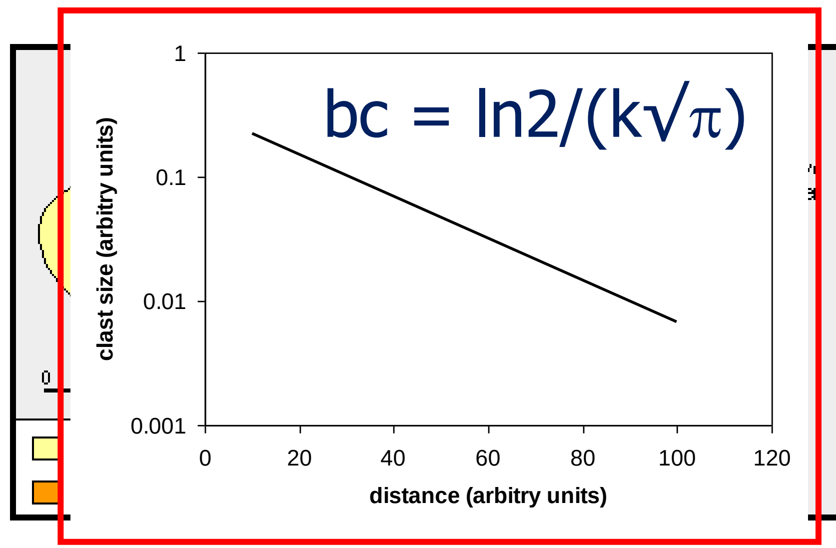
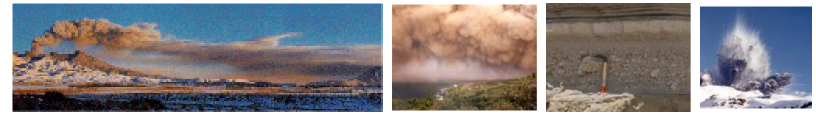




Isopach maps

Isopleth maps



CLASSIFICATION

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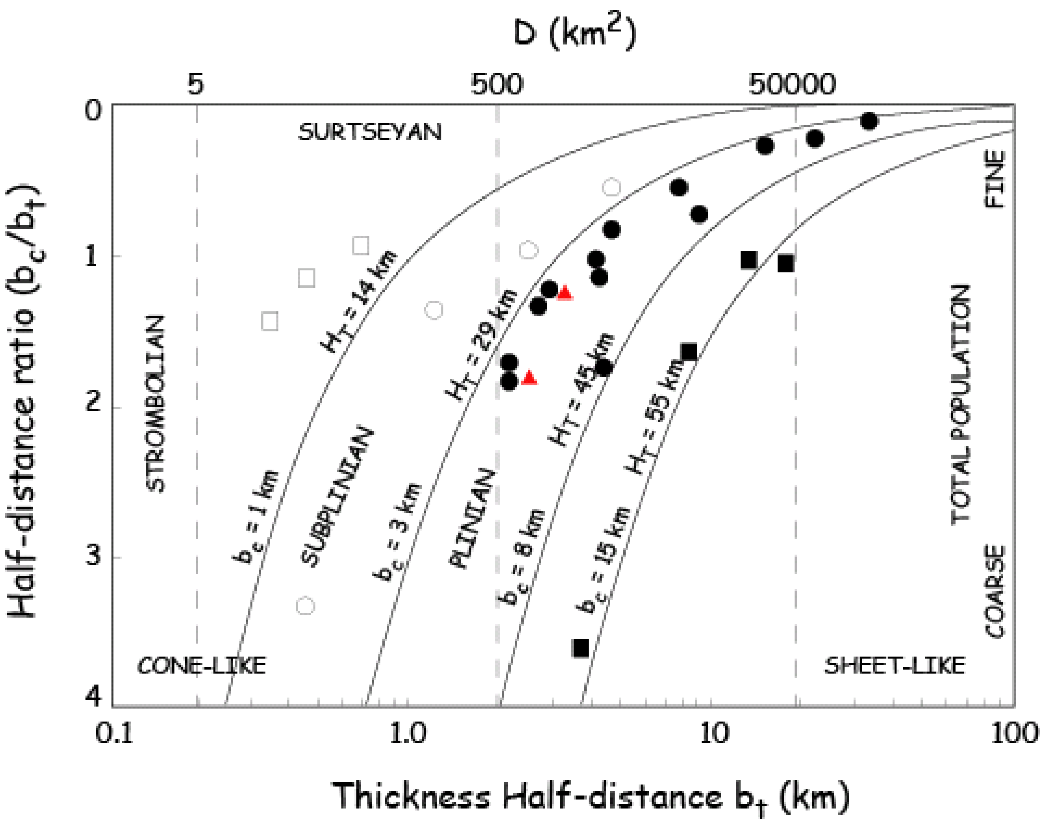
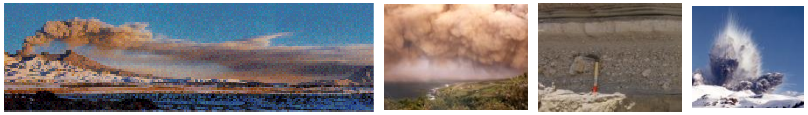
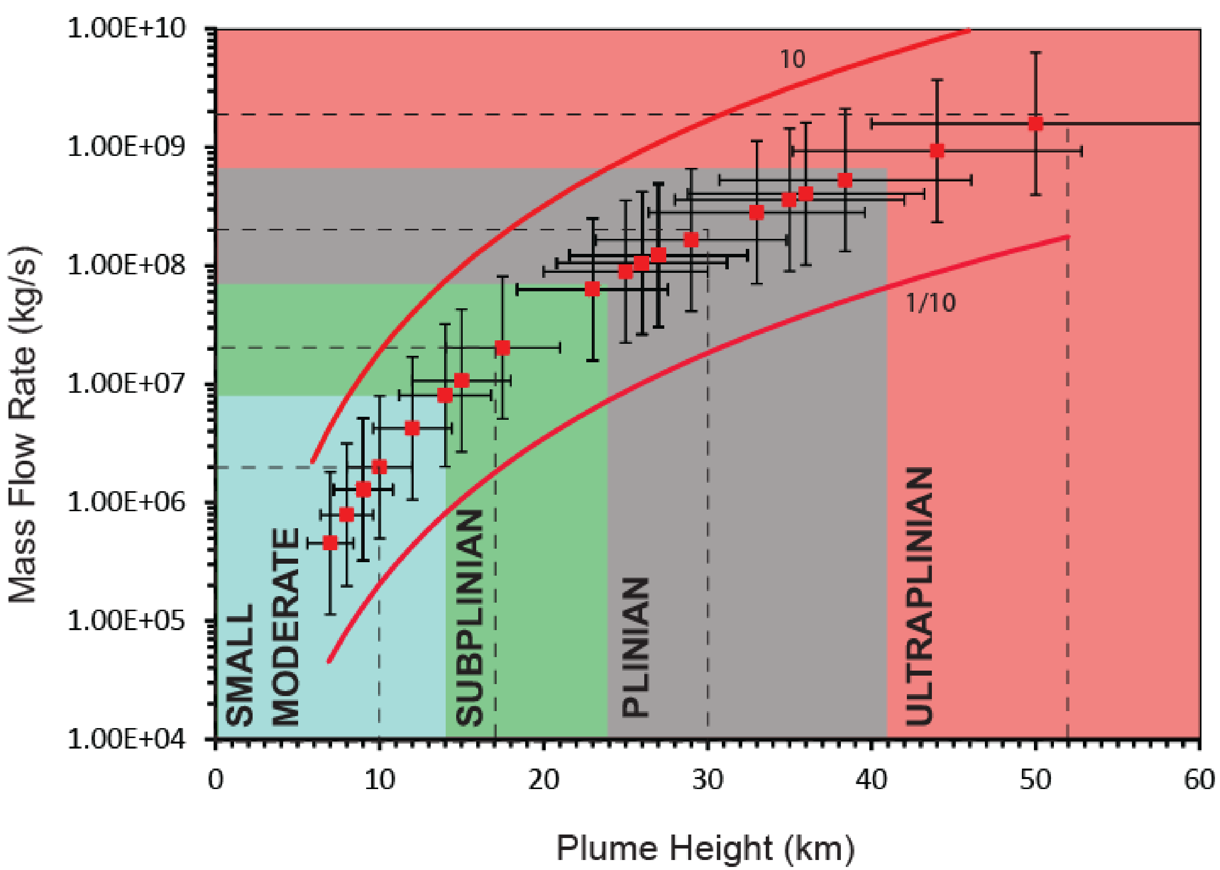


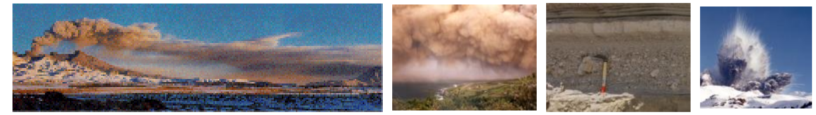


CLASSIFICATION

Bonadonna and Costa 2013

Eruptive style based on Plume Height and Mass Flow Rate





Physical characterization of explosive eruptions

- **Isopach maps/Isomass maps** → erupted volume/mass (+VEI and DRE) + wind direction
- **Isopleth maps** → Plume height → Mass Eruption Rate → Eruption duration (when combined with erupted mass)
- **Isopleth maps** → Wind speed
- **Individual grainsize distributions** → Total grainsize distribution → Infos on eruptive style and fragmentation efficiency
- **Thickness vs grainsize data** → eruption classification




EXERCISE 3

DETERMINATION OF:

Column height

Mass Eruption Rate

Volume

ERUPTION CLASSIFICATION

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