

¹ Influence of changing temperature and upwelling intensity on
² Indian oil sardine (*Sardinella longiceps*) landings

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¹¹ **Running title:** Modeling Indian oil sardine landings

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Abstract

Commercial landings of sardine are known for strong year-to-year fluctuations. A key driver is thought to be environmental variability, to which small forage fish are especially sensitive. We examined the environmental drivers associated with landings fluctuations in the Indian oil sardine using a 32-year time series of quarterly catch. Past research suggested a variety of influential variables: precipitation, upwelling intensity, ocean temperature (SST), chlorophyll concentration and large-scale coupled atmosphere–ocean phenomena (ENSO). Using the life history of the Indian oil sardine, we developed hypotheses concerning how these variables might affect landings and tested them using generalized additive models which allow non-linear response curves and dynamic linear models which allow time-varying responses. We found significant correlation for three variables: upwelling intensity, an ENSO index, and the multi-year average nearshore SST. Upwelling intensity can have both a positive effect (fueling higher food availability) and a negative effect at extreme intensity (surface anoxia). The negative effect was apparent for both monsoon and post-monsoon catch. However, the most significant correlation (adjusted R^2 of 67.5%) was between the 2.5 year average nearshore SST and post-monsoon landings. The multi-year average SST also been identified as a predictor for Pacific and southern African sardine fluctuations, suggesting that the average SST over the sardine lifespan successfully captures a variety of factors which predict future abundance. The Western Indian Ocean has been steadily warming and changes been most extreme during the summer monsoon. Our work highlights that these changes are likely to affect oil sardine landings.

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Keywords: Indian oil sardine, catch prediction, GAM modeling, climate, sea surface

temperature, remote sensing, Southeast Arabian Sea

49 **Introduction**

50 Environmental variability is known to be a key driver of population variability of small forage
51 fish such as sardines, anchovy and herring (Alheit & Hagen, 1997; Checkley et al., 2017; Cury
52 et al., 2000). In particular, ocean temperature and upwelling dynamics, along with density-
53 dependent feedback, have been identified as important in affecting recruitment success and
54 biomass of European and Pacific sardines (*Sardina pilchardus* and *Sardinops sagax*) (Alheit et
55 al., 2012; Garza-Gil et al., 2015; Jacobson & MacCall, 1995; Lindegren & Checkley, 2012;
56 Lindegren et al., 2013; Rykaczewski & Checkley, 2008). Like other sardines, the Indian oil
57 sardine (*Sardinella longiceps* Valenciennes, 1847) shows strong inter-annual fluctuations and
58 larger decadal booms and busts. The Indian oil sardine offers an instructive case study to
59 investigate the effects of environmental variability, particularly temperature and upwelling dy-
60 namics, as they occupy an ocean system that is warmer than that occupied by other sardines
61 and have a strong seasonal cycle driven by the Indian summer monsoon.

62 The Indian oil sardine is one of the most commercially important fish resources along the
63 southwest coast of India (Figure 1) and historically has comprised approximately 25% of the
64 marine catch in the Indian fisheries (Vivekanandan et al., 2003). Landings of the Indian oil sar-
65 dine are highly seasonal, peaking after the summer monsoon period in October-December and
66 reaching a nadir in spring before the summer monsoon in April-June (Figure 2). At the same
67 time, the landings of this small pelagic finfish are highly variable from year to year. Small
68 pelagics are well known to exhibit high variability in biomass due to the effects of environ-
69 mental conditions on survival and recruitment, but in this fishery, environmental conditions
70 also affect exposure of sardines to the fishery. Until recently, the Indian oil sardine fishery was
71 artisanal and based on small human or low powered boats with no refrigeration. The fishery
72 was confined to nearshore waters, and thus migration of sardines in and out of the coastal zone
73 greatly affected exposure to the fishery and hence landings.

74 Researchers have examined a variety of environmental variables for their correlation with
75 landings of the Indian oil sardine. Precipitation during the monsoon and the day of the mon-
76 soon arrival are thought to act as either a direct or indirect cue for spawning (Antony Raja,
77 1969, 1974; Jayaprakash, 2002; Murty & Edelman, 1966; Pitchaikani & Lipton, 2012; Sri-
78 nath, 1998; Xu & Boyce, 2009). Although many studies have looked for and some have found
79 correlations between precipitation and landings, the reported relationships are positive in some
80 studies and negative in others (Madhupratap et al., 1994). Researchers have also looked for and
81 found correlations with various metrics of upwelling intensity (Jayaprakash, 2002; Longhurst

& Wooster, 1990; Madhupratap et al., 1994; Murty & Edelman, 1966; Srinath, 1998; Thara, 2011), with direct measures of productivity such as nearshore zooplankton and phytoplankton abundance (George et al., 2012; Madhupratap et al., 1994; Menon et al., 2019; Nair, 1952; Nair & Subrahmanyam, 1955; Piontkovski et al., 2014; Pitchaikani & Lipton, 2012), and with nearshore sea surface temperature (SST) (Annigeri, 1969; Pillai, 1991; Prabhu & Dhulkhed, 1970; Supraba et al., 2016). SST can affect both somatic growth rates and juvenile survival but in this system also can cause fish to move offshore and away from the shore-based fishery. The multi-year average sea temperature is postulated to have effects on recruitment and the survival of larval and juvenile sardines, which affect the later overall abundance (Checkley et al., 2017; Takasuka et al., 2007). The El Niño/Southern Oscillation (ENSO) has a cascading effect on all the aforementioned environmental parameters (SST, precipitation, upwelling) and correlations have been found between ENSO indices and sardine landings (Rohit et al., 2018; Supraba et al., 2016) and coastal anoxic events which affect sardines (Vallivattathillam et al., 2017).

In this paper, we study the utility of environmental covariates from remote sensing to explain year-to-year variability using a long-term time series of quarterly Indian oil sardine landings. This time series is derived from a stratified sampling design that surveys the fishery landing sites along the southwest Indian coast and was first implemented in the 1950s (Srinath et al., 2005). The goal is to identify environmental covariates which can explain catch variability and improve the accuracy of short-term catch forecasts. Landings are a product of the biomass, the catchability, and the effort. A traditional auto-correlated catch model (ARIMA) can model smooth changes in landings, such as due to changes in fleet size or multi-year biomass changes, but the environment adds a large component of year-to-year variability that such a model does not capture. The environment affects food resources which affects recruitment through spawning and survival, and thus the biomass available to the fishery. In addition, in the Indian oil sardine system, catchability is strongly affected by the environment by affecting the inshore versus offshore distribution of sardine. The fishery is restricted to the nearshore < 50 km offshore (Rohit et al., 2018). When the sardines move offshore to spawn or to avoid hypoxic or excessively warm water, they are no longer available to the fishery. Thus, through its effects on recruitment and catchability, the environment has the potential to drive year-to-year changes in landings. The covariates which we study (Table 1) are linked to aspects of oil sardine life history that are expected to affect catch via catchability or biomass. Covariates from remote sensing are the focus because they are available over a wide spatial extent at a daily and monthly resolution thus are practical for use in operational forecasts. A better understanding of how and whether remote sensing data explains variation in seasonal catch will

¹¹⁶ support future efforts to use satellite data to improve catch forecasts.

¹¹⁷ **Catch modeling versus biomass modeling**

¹¹⁸ Modeling and forecasting landings data using statistical models fit to annual or seasonal catch
¹¹⁹ time series has a long tradition in fisheries and has been applied to many species (Cohen &
¹²⁰ Stone, 1987; Farmer & Froeschke, 2015; Georgakarakos et al., 2006; Hanson et al., 2006;
¹²¹ Lawer, 2016; Lloret et al., 2000; Mendelssohn, 1981; Nobel & Sathianandan, 1991; Prista
¹²² et al., 2011; Stergiou & Christou, 1996), including oil sardines (Srinath, 1998; Venugopalan
¹²³ & Srinath, 1998). These models can be used to identify the variables correlated with catch
¹²⁴ fluctuations and can be used to provide landings forecasts which are useful for fishery managers
¹²⁵ and the fishing industry. An example of the former is using catch forecasts to set or give
¹²⁶ warnings of seasonal fishery closures if the forecasted catch is higher than the allowed catch
¹²⁷ limits (Farmer & Froeschke, 2015). An example of the latter is the annual Gulf and Atlantic
¹²⁸ menhaden forecast produced by NOAA Fisheries (Hanson et al., 2006; Schaaf et al., 1975).
¹²⁹ This multiple regression model has been used for the last 45 years to produce an annual forecast
¹³⁰ of menhaden landings, which is used for planning purposes by the industry, not only the fishers
¹³¹ but also fish catch sellers and buyers, businesses which provide fisheries gear, and banks which
¹³² provide financing (Hanson et al., 2006).

¹³³ For the purpose of our study, the assumption of a tight relationship between landings and
¹³⁴ abundance is not necessary. The objective is to understand what drives landings variability,
¹³⁵ whether it is due to biomass or catchability variation. That said, Indian oil sardine landings are
¹³⁶ often assumed to reflect the total abundance for reasons specific to the species and the fishery
¹³⁷ (cf. Kripa et al., 2018). Historically, the fishery was artisanal: small boats with small motors,
¹³⁸ no refrigeration, and limited to the near shore. The ring seine was introduced in the 1980s, but
¹³⁹ widespread mechanization of the fleet is a very recent development. Fishers with small boats
¹⁴⁰ have limited ability to target the stock, at least not to the degree that landings remain constant as
¹⁴¹ a stock declines. That pattern can be observed in a large, mobile, highly mechanized fleet. The
¹⁴² fishery is unregulated, except for a brief closure during the monsoon months, thus the landings
¹⁴³ are not being affected by area closures and catch limits. Finally, the fishery is dispersed along
¹⁴⁴ the entire coastline rather than being focused from a few large ports. Again, for our objectives,
¹⁴⁵ it is not necessary that landings be a tight index of biomass, but there are many reasons to
¹⁴⁶ assume that this relationship is strong.

¹⁴⁷ Unfortunately historical biomass estimates are not possible for the Indian oil sardine.

¹⁴⁸ Length- or age-structured models (e.g. virtual population analysis) which produce biomass
¹⁴⁹ estimates are not possible due to the lack of effort and catch-at-age information for the fishery.
¹⁵⁰ The available long-term effort data are indirect (boat composition of the fishery at multi-year
¹⁵¹ intervals) and estimates of number of trips or hours fishing are only available in a few recent
¹⁵² years, and the data available are approximate given the vessel diversity of the fishery and sam-
¹⁵³ pling constraints. Nonetheless the number and size of boats involved in the fishery has been
¹⁵⁴ increasing. Oil sardines are caught primarily by ring seines, which were introduced in the
¹⁵⁵ early 1980s. Ring seines of different sizes are used on both traditional small boats and on
¹⁵⁶ large mechanized ships (Das & Edwin, 2018). Since 1985, the ring seine fishery has expanded
¹⁵⁷ steadily in terms of horsepower, size of boats, and length of nets. There are concerns that
¹⁵⁸ over-fishing is a factor in the most recent oil sardine declines after 2015 (Kripa et al., 2018).
¹⁵⁹ Steadily increasing effort is assumed to have increased the landings, at least prior to 2015.
¹⁶⁰ Our base catch model, an auto-regressive model, will capture smooth landings trends due to
¹⁶¹ increased effort (or multi-year changes in biomass).

¹⁶² Study Area

¹⁶³ Our analysis focuses on the Kerala coast (Figure 1) region of India, where the majority of the
¹⁶⁴ Indian oil sardines are landed and where oil sardines comprise ca. 40% of the marine fish
¹⁶⁵ catch (Srinath, 1998; Vivekanandan et al., 2003). This area is in the Southeast Arabian Sea,
¹⁶⁶ one of world's major upwelling zones, with seasonal peaks in primary productivity driven by
¹⁶⁷ upwelling caused by winds during the Indian summer monsoon (Habeebrehman et al., 2008;
¹⁶⁸ Madhupratap et al., 2001) between June and September. Within the study area, the coastal zone
¹⁶⁹ off Kerala between 9°N to 13°N has especially intense upwelling due to the combined effects
¹⁷⁰ of wind stress and remote forcing (BR, 2010; BR et al., 2008). The result is a strong tempera-
¹⁷¹ ture differential between the nearshore and offshore and high primary productivity and surface
¹⁷² chlorophyll concentration in this region during summer and early fall (BR, 2010; Chauhan et
¹⁷³ al., 2011; Habeebrehman et al., 2008; Jayaram et al., 2010; Madhupratap et al., 2001; Ragh-
¹⁷⁴ van et al., 2010). The primary productivity peaks subside after September while mesozoo-
¹⁷⁵ plankton abundances increase and remain high in the post-monsoon period (Madhupratap et
¹⁷⁶ al., 2001).

177 **Oil sardine life cycle and fishery**

178 The Indian oil sardine fishery is restricted to the narrow strip of the western India continental
179 shelf, within 20 km from the shore (Figure 1). The yearly cycle of the fishery begins at the
180 start of spawning during June to July, corresponding with the onset of the summer monsoon
181 (Antony Raja, 1969; Chidambaram, 1950) and the initiation of strongly cooler nearshore SST
182 due to upwelling (Figure 3). At this time, the mature fish migrate from offshore to coastal
183 spawning areas, and the spawning begins during the summer monsoon period when temper-
184 ature, salinity and suitable food availability are conducive for larval survival (Chidambaram,
185 1950; Jayaprakash & Pillai, 2000; Krishnakumar et al., 2008; Murty & Edelman, 1966; Nair et
186 al., 2016). Although peak spawning occurs in June to July, spawning continues into Septem-
187 ber (Antony Raja, 1969; Hornell, 1910; Hornell & Nayudu, 1924; Prabhu & Dhulkhed, 1970)
188 and early- and late-spawning cohorts are evident in the length distributions of the 0-year fish.
189 Spawning occurs in waters outside of the traditional range of the fishery (Antony Raja, 1964),
190 and after spawning the adults migrate closer to the coast where the spent fish become exposed
191 to the fishery.

192 After eggs are spawned, they develop rapidly into larvae (Nair, 1959). The phytoplankton
193 bloom that provide sardine larvae food is dependent upon nutrient influx from coastal up-
194 welling and runoff from rivers during the summer monsoon and early fall. The blooms start in
195 the south near the southern tip of India in June, increase in intensity and spread northward up
196 the coast (BR, 2010). Variation in the bloom initiation time and intensity leads to changes in
197 the food supply and to corresponding changes in the growth and survival of larvae and in the
198 later recruitment of 0-year sardines into the fishery (George et al., 2012). Oil sardines grow
199 rapidly during their first few months, and 0-year fish (40mm to 100mm) from early spawn-
200 ing appear in the catch in August and September in most years (Antony Raja, 1970; Nair et
201 al., 2016). As the phytoplankton bloom spreads northward, the oil sardines follow, and the
202 oil sardine fishery builds from south to north during the post-monsoon period. Oil sardines
203 remain inshore feeding throughout the winter months, until March to May when the inshore
204 waters warm considerably and sardines move offshore to deeper waters (Chidambaram, 1950).
205 Catches of sardines are correspondingly low during this time for all size classes. The age at
206 first maturity occurs at approximately 150 mm size (Nair et al., 2016), which is reached within
207 one year. When the summer monsoon returns, the oil sardine cycle begins anew.

208 Catches along the Kerala coast are high throughout the year except during quarter 2, Apr-
209 Jun (Figure 2). The age-distribution caught by the fishery varies through the year. The fishery

210 is closed during June to mid-July during the monsoon and peak spawning, and when it resumes
211 in mid-July, it is first dominated by 1-2.5 year old mature fish (Antony Raja, 1969; Bensam,
212 1964; Nair et al., 2016). In August or September a spike of 0-year (40mm) juveniles from
213 the June spawning typically appears in the catch (Antony Raja, 1969; Nair et al., 2016) and
214 another spike of 0-year fish is sometimes seen in February from the late fall spawning (Prabhu
215 & Dhulkhed, 1967, 1970). From October through June, the catch is dominated by fish from
216 120mm-180mm (Antony Raja, 1970; Nair et al., 2016; Prabhu & Dhulkhed, 1970) which is a
217 mix of 0-year, 1-year and 2-year fish (Nair et al., 2016; Rohit et al., 2018).

218 Materials and Methods

219 Sardine landing data

220 Quarterly fish landing data have been collected by the Central Marine Fisheries Research Insti-
221 tute (CMFRI) in Kochi, India, since the early 1950s using a stratified multi-stage sample design
222 (Srinath et al., 2005). The survey visits the fish landing centers along the entire southwest coast
223 of India and samples the catch from the variety of boat types and gear types used in the coastal
224 fishery. Landings estimates are available for all the coastal states, however we model the catch
225 for the state of Kerala only, where the longest time series is available and the overwhelming
226 majority of oil sardines are landed (Figure 2). The quarterly landings (metric tons) for oil
227 sardine landed from all gears in Kerala were obtained from CMFRI reports (1956-1984) and
228 online databases (1985-2015); see Appendix G for data sources. The quarterly landing data
229 were log-transformed to stabilize the variance.

230 Remote sensing data

231 We analyzed monthly composites of the following environmental data derived from satellite
232 data: sea surface temperature (SST), chlorophyll-a (CHL), upwelling (UPW), the Oceanic
233 Niño Index (ONI), the Dipole Mode Index (DMI) and precipitation. The monthly time series
234 and means of the covariates are shown in Figure 4.

235 For sea surface temperature, we used Advanced Very-High Resolution Radiometer
236 (AVHRR) data, which provides accurate nearshore SST values. Although the ICOADS
237 product provides SST values for earlier years, ICOADS does not provide accurate nearshore

238 temperatures. For 1981 to 2003, we used the Pathfinder Version 5.2 product on a 0.0417
239 degree grid. These data were developed by the Group for High Resolution Sea Surface
240 Temperature (GHRSSST) and served by the US National Oceanographic Data Center. For
241 2004 to 2016, we used the NOAA CoastWatch SST products derived from NOAA's Polar
242 Operational Environmental Satellites (POES).

243 We used the chlorophyll-a products developed by the Ocean Biology Processing Group in
244 the Ocean Ecology Laboratory at the NASA Goddard Space Flight Center. For 1997 to 2002,
245 we used the chlorophyll-a 2014.0 Reprocessing (R2014.0) product from the Sea-viewing Wide
246 Field-of-view Sensor (SeaWiFS) on the Orbview-2 satellite. These data are on a 0.1 degree
247 grid. For 2003 to 2017, we used the MODIS-Aqua product on a 0.05 degree grid. These CHL
248 data are taken from measurements by the Moderate Resolution Imaging Spectroradiometer
249 (MODIS) on NASA's Aqua Spacecraft. The SST and CHL data were averaged across thirteen
250 1 degree by 1 degree boxes which roughly parallel the bathymetry (Figure 1). The SST and
251 CHL satellite data were retrieved from NOAA remote-sensing data servers; see Appendix G
252 for data sources and references.

253 For an index of coastal upwelling, we used three indices. The first was the sea surface tem-
254 perature differential between near shore and 3 degrees offshore based on the index described
255 by Naidu et al. (1999) and BR et al. (2008). For SST, we used the remote sensing sea surface
256 temperature data sets described above. This SST-based upwelling index has been validated as
257 a more reliable metric of upwelling off the coast of Kerala compared to wind-based upwelling
258 indices (BR et al., 2008). The SST-based upwelling index and chlorophyll-a concentration
259 have a strong temporal association (Figure 3). The second index was simply average nearshore
260 SST along the Kerala coasts (average of boxes 2-5 in Figure 1). The third index was the Bakun
261 index based on wind stress. The index is computed from the the x- and y- components of
262 Ekman Transport. See Appendix G for data sources and references.

263 Precipitation data were obtained from two different sources. The first was an estimate
264 of the monthly precipitation (in mm) over Kerala from land-based rain gauges; these data are
265 available from the Indian Institute of Tropical Meteorology and the data are available from the
266 start of our landing data (1956). The second was a remote sensing precipitation product from
267 the NOAA Global Precipitation Climatology Project. This provides estimates of precipitation
268 over the ocean using a global 2.5 degree grid. We used the 2.5 by 2.5 degree box defined by
269 latitude 8.75 to 11.25 and longitude 73.25 to 75.75 for the precipitation off the coast of Kerala.
270 These data are available from 1979 forward. The land and nearshore ocean precipitation data
271 are highly correlated (Appendix D). See Appendix G for the precipitation data repositories and

272 references.

273 The Oceanic Niño Index (ONI) is a measure of the SST anomaly in the east-central Pacific
274 and is a standard index of the El Niño/Southern Oscillation (ENSO) cycle. The ONI index is
275 3-month running mean of ERSST.v5 SST anomalies in the Niño 3.4 region, based on centered
276 30-year base periods updated every 5 years. The ONI was downloaded from the NOAA Na-
277 tional Weather Service Climate Prediction Center. The Dipole Mode Index (DMI) is defined
278 by the SSTA difference between the western Indian Ocean (10°S – 10°N , 50°E – 70°E) and the
279 southeastern Indian Ocean (10°S – 0° , 90°E – 110°E). The DMI has been shown to predict anoxic
280 events in our study area (Vallivattathillam et al., 2017). The DMI data were downloaded from
281 the NOAA Earth System Research Laboratory. See Appendix G for the data servers where the
282 ENSO data were downloaded and computation notes and references.

283 **Hypotheses**

284 Our statistical tests were structured around specific hypotheses (Table 1) concerning which re-
285 mote sensing covariates in which months should correlate with landings in specific quarters.
286 These hypotheses were based on biological information concerning how environmental con-
287 ditions affect sardine survival and recruitment and affect exposure of Indian oil sardines to
288 the coastal fishery. The Jul-Sep catch overlaps the summer monsoon and the main spawning
289 months. This is also the quarter where small 0-year fish from early spawning often appear in
290 the catch, sometimes in large numbers. Variables that affect or are correlated with movement of
291 sardines inshore should be correlated with Jul-Sep landings. In addition, pre-spawning (Mar-
292 May) environmental conditions should be correlated with the spawning strength as adult oil
293 sardines experience an acceleration of growth during this period along with egg development.
294 The post-monsoon catch (Oct-May) is a mix of 0-year fish (less than 12 months old) and ma-
295 ture fish (greater than 12 months old). Variables that are correlated with spawning strength and
296 larval and juvenile survival should correlate with the post-monsoon catch both in the current
297 year and in future years, one to two years after.

298 Our hypotheses (Table 1) focus mainly on two drivers: upwelling and ocean tempera-
299 ture. We also test hypotheses concerning precipitation as this has historically been an envi-
300 ronmental covariate considered to influence the timing of oil sardine landings. More recently,
301 researchers have highlighted the influence of large-scale ocean processes, specifically the El
302 Niño/Southern Oscillation, on sardine fluctuations; therefore we test the Ocean Niño Index
303 (ONI) and Dipole Mode Index (DMI) also. Chlorophyll-a concentration is directly correlated

304 with sardine food availability and the location of chlorophyll fronts are known to be associated
305 with sardine shoaling. However remote sensing chlorophyll data are only available for a few
306 years (1997-2014) and the statistical power for testing correlation with landings is low. Tests
307 with chlorophyll are shown in the appendices but are not the focus of our analyses.

308 Statistical models

309 We modeled the catches during the late-monsoon season (Jul-Sep) separately from the post-
310 monsoon season (Oct-Mar). Thus, there is no seasonality in our catch time series as we ana-
311 lyzed a yearly time series of Jul-Sep catches separately from a yearly time series of Oct-Mar
312 catches. We divided the catch in this way for biological and statistical reasons. Catch in Jul-
313 Sep captures a mix of spawning-age fish as it overlaps with the tail end of the spawning season,
314 is affected by a fishery closure during the summer monsoon, and is periodically inflated by the
315 appearance of small 0-year fish from early summer spawning. In addition, the covariates that
316 affect the timing of spawning, movement of post-spawning mature fish inshore, and early egg
317 and larval survival may be different than those that affect later growth, survival and shoaling
318 that exposes fish to the inshore fishery. Analyzing catch and covariate time series without sea-
319 sonality also had an important statistical benefit—we removed the problem of seasonality in
320 the catch and covariates. The oil sardine life-cycle is seasonal and driven by the strong season-
321 ality in this monsoon influenced system. A simple statistical model with quarters will explain
322 much of the quarterly catch data since most of the yearly variability is due to seasonality and
323 any environmental covariate with a similar seasonality will also show high correlation with the
324 landings. Our goal was to explain year-to-year variability and thus eliminating the confounding
325 effect of seasonality in the data was important.

326 Preliminarily, we tested ARIMA models on both the Jul-Sep and Oct-Mar catch time
327 series and found little support for auto-regressive errors (ARIMA models with a MA com-
328 ponent) based on diagnostic tests of the residuals and model selection. The best supported
329 ARIMA models were simple AR models ($x_t = bx_{t-1} + \varepsilon_t$). This lack of strong auto-correlation
330 in residuals has been found in other studies that tested ARIMA models for forecasting small
331 pelagic catch (Stergiou & Christou, 1996). We thus used AR-only models, however we tested
332 both linear and non-linear models using generalized additive models (GAMs, Wood, 2017) of
333 the form $x_t = s(x_{t-1}) + \varepsilon_t$ and tested time-varying linear models with dynamic linear models
334 (DLM). GAMs allow one to model the effect of a covariate as a flexible non-linear function
335 while DLMs allow the effect of the covariate to vary over time. Our GAM approach is analo-

336 gous to that taken by Jacobson and MacCall (1995) in a study of the effects of SST on Pacific
337 sardine recruitment.

338 The first step in our analysis was to determine the catch model: the model for current
339 catch as a function of the past catch. We explored four classes of models: null models with a
340 simple function of prior catch, linear regressive models with one to two years of prior catch,
341 dynamic linear models (DLM) which allow the regression parameters to vary (Holmes et al.,
342 2012, using the MARSS package in R), and GAMs to allow the effect of prior catch to be
343 non-linear. One feature of GAMs is that they allow the smoothing parameter of the response
344 curve to be estimated. We fixed the smoothing parameter at an intermediate value so that
345 smooth responses were achieved. Multi-modal or overly flexible response curves would not
346 be realistic for our application. We fit GAMs with smooth terms represented by penalized
347 regression splines (Wood, 2011, using the mgcv package in R) and fixed the smoothing term
348 at an intermediate value (sp=0.6).

349 We compared the following catch models:

- 350 • null: $\ln(C_{i,t}) = \ln(C_{j,t-1})$
- 351 • random walk: $\ln(C_{i,t}) = \alpha + \ln(C_{j,t-1}) + \varepsilon_t$
- 352 • linear AR-1: $\ln(C_{i,t}) = \alpha + \beta \ln(C_{j,t-1}) + \varepsilon_t$
- 353 • linear AR-2: $\ln(C_{i,t}) = \alpha + \beta_1 \ln(C_{j,t-1}) + \beta_2 \ln(C_{k,t-2}) + \varepsilon_t$
- 354 • DLM AR-1: $\ln(C_{i,t}) = \alpha_t + \beta_t \ln(C_{j,t-1}) + \varepsilon_t$
- 355 • GAM AR-1: $\ln(C_{i,t}) = \alpha + s(\ln(C_{j,t-1})) + \varepsilon_t$
- 356 • GAM AR-2: $\ln(C_{i,t}) = \alpha + s_1(\ln(C_{j,t-1})) + s_2(\ln(C_{k,t-2})) + \varepsilon_t$

357 $\ln(C_{i,t})$ is the log catch in the current year t in season i . We modeled two different catches: S_t
358 (Jul-Sep) and N_t (Oct-Jun). The catches were logged to stabilize and normalize the variance.
359 $s()$ is a non-linear function estimated by the GAM algorithm. The model is primarily statistical,
360 meaning it should not be thought of as a population growth model. We tested models with prior
361 year and two years prior Oct-Mar catch (N_{t-1} and N_{t-2}) and Jul-Sep catch (S_{t-1} and S_{t-2}) as
362 the explanatory catch variable. S_t was not used as a predictor for N_t because S_t is the quarter
363 immediately prior to N_t and would not be available for a forecast model since time is required
364 to process landings data. The catch models were fit to 1984 to 2015 catch data, corresponding
365 to the years when the SST, upwelling and precipitation data were available. F-tests, AIC and
366 leave-one-out cross-validation (LOOCV) on nested sets of models (Wood et al., 2016) were
367 used to evaluate the support for the catch models and later for the covariate models. LOOCV

368 involves leaving out a data point, fitting the model, and then predicting the left-out data point.
369 The root mean squared error (RMSE) is reported for the set of prediction errors (this is also
370 known as the predicted residual error sum of squares or PRESS statistic). After selection of
371 the best model with the 1984-2015 data, the fitting was repeated with the 1956-1983 catch data
372 to confirm the form of the catch models. An influential years test was done by removing each
373 year and repeating the model selection analysis.

374 Once the catch models were determined, the covariates were studied individually and then
375 jointly. As with the catch models, F-tests, AIC and LOOCV (leave-one-out cross-validation)
376 on nested sets of models were used to evaluate the support for models with covariates. The
377 smoothing term was fixed at an intermediate value ($sp=0.6$) instead of being treated as an
378 estimated variable. Our models for catch with covariates took the form $\ln(C_{i,t}) = M + \alpha +$
379 $s_1(V_{1,t}) + s_2(V_{2,t-1}) + \varepsilon_t$, $\ln(C_{i,t}) = M + \alpha + \beta_1 V_{1,t} + \beta_2 V_{2,t-1} + \varepsilon_t$, and $\ln(C_{i,t}) = M + \alpha +$
380 $\beta_t V_{1,t} + \varepsilon_t$ where M was the best catch model from step 1 and V is a covariate. Thus models
381 with covariates modeled as a linear, non-linear and time-varying effect were compared. The
382 covariates tested are those hypothesized to drive variability in oil sardine landings (Table 1).
383 We tested both models with one and two covariates and did not use correlated covariates in the
384 same model.

385 Results

386 Catches in prior seasons as explanatory variables

387 Using the 1984-2015 catch data, the time-period that overlaps our available environmental data,
388 the Jul-Sep catch models were compared against a “naive” model in which the forecasted Jul-
389 Sep catch is the Jul-Sep catch in the prior year. The “naive” model has no estimated parameters
390 and is a standard null model for time series modeling. Models with $\ln(N_{t-1})$ (Oct-Mar catch
391 in prior year) as the explanatory covariate were strongly supported over the naive model and
392 over models with $\ln(S_{t-1})$ (Jul-Sep catch in prior year) as the explanatory variable (Tables
393 A1 and A2). Addition of the catch two years prior, $\ln(N_{t-2})$ or $\ln(S_{t-2})$, was not supported
394 (by AIC or F-tests) for either the linear or non-linear models. We tested the support for non-
395 linearity in the effect of the prior year catch by comparing models with $\ln(N_{t-1})$ or $\ln(S_{t-1})$
396 included as a linear effect or as a non-linear effect using GAMs (Table A2). The residual error
397 decreased using a non-linear response and LOOCV decreased but at the cost increased degrees
398 of freedom. Overall there were three models with almost identical AIC and LOOCV: linear and

399 non-linear with $\ln(N_{t-1})$, and non-linear with $\ln(N_{t-1})$ and $\ln(N_{t-2})$. We choose the non-linear
 400 with $\ln(N_{t-1})$ as the base catch model based on further diagnostic tests (described below) and
 401 to minimize loss of degrees of freedom. The adjusted R^2 of this model was 24.4%.

402 The model selection results were similar for models of the Oct-Mar landings (N_t), but
 403 the models explained much more of the variance (with a maximum adjusted $R^2 = 56.6$). The
 404 most supported model for N_t (Tables A3 and A4) based on AIC and F-tests used a non-linear
 405 response to Oct-Mar catch of the previous season $\ln(N_{t-1})$ plus a non-linear response to Jul-
 406 Sep catch two years prior $\ln(S_{t-2})$, however the LOOCV (out of sample prediction accuracy)
 407 was higher than the naive null model. The simpler model with only $\ln(N_{t-1})$ had the second
 408 lowest AIC and the lowest LOOCV (and lower than the naive null model). This simpler model
 409 was also included as one of the base models for the Oct-Mar catch.

410 As diagnostic checks, we repeated the model comparisons with the landings data set from
 411 1956 to 1983. The results were the same for the Jul-Sep catch (Table A5) with the model
 412 with $\ln(N_{t-1})$ included as a non-linear covariate giving the lowest AIC and LOOCV. For the
 413 Oct-Mar catch (Table A6), the results were very similar but not identical. The model with
 414 $\ln(N_{t-1})$ included as a non-linear covariate had the lowest LOOCV while the models with
 415 $\ln(N_{t-1})$ and $\ln(S_{t-2})$ or $\ln(S_{t-1})$ had the lowest AIC (though less than 1 from the AIC of
 416 the $\ln(N_{t-1})$ model). We also did an influential years test using leave-one-out cross-validation
 417 (Appendix F). This test involved leaving out one year and repeating the model selection tests.
 418 This analysis supported the selected base models using the 1984-2015 data. The dynamic
 419 linear models (allowing a time-varying effect of prior catch) performed poorly for the Jul-Set
 420 catch with high AIC and LOOCV. For the Oct-Mar catch, the performance was mixed with
 421 higher AIC but lower LOOCV for one of the DLMs.

422 Based on the model selection tests, the following non-linear model was chosen as the base
 423 model for Jul-Sep (S_t) catch:

$$M0 : \ln(S_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t.$$

424 Two base non-linear models were chosen for Oct-Mar (N_t) catch:

$$M1 : \ln(N_t) = \alpha + s_1(\ln(N_{t-1})) + s_2(\ln(S_{t-2})) + \varepsilon_t$$

425

$$M2 : \ln(N_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$$

426 Note that although M0 was the best model for Jul-Sep catch, it was only weakly explanatory.

427 The maximum adjusted R^2 was less than 30% (Table A2). For the Oct-Mar catch, M1 and M2
428 were more explanatory with an adjusted R^2 of 45.3% for M1 and 56.6% for M2 (Table A4).

429 **Environmental covariates as explanatory variables**

430 There was no support for using precipitation during the summer monsoon (June-July) or pre-
431 monsoon period (April-May) as an explanatory variable for the Jul-Sep or Oct-Mar catch (hy-
432 potheses S1 and S2; Tables B1, B2 and B3). This was the case whether precipitation in the
433 current or previous season was used, if precipitation was included as a non-linear or linear ef-
434 fect, and if either precipitation during early monsoon (June-July) or pre-monsoon (April-May)
435 was used as the covariate. Jul-Sep catch overlaps with the late spawning period and precip-
436 itation is often thought to trigger spawning, however we were unable to find any consistent
437 association of catch with precipitation. Raja (1974) posited that the appropriate time period for
438 the effect of rainfall is the weeks before and after the new moon when spawning is postulated
439 to occur and not the total rainfall during the monsoon season. Thus the lack of correlation may
440 be due to using too coarse of a time average for the precipitation.

441 The sea surface temperature before spawning (March-May) has been speculated to be
442 correlated with successful egg development and spawning behavior (hypothesis S4 and S5)
443 and extreme heat events in the pre-spawning period have been associated with low recruitment.
444 This suggests that March-May in the current and prior years should be associated with low
445 catch. The sea surface temperature during larval and early juvenile development (October-
446 December) may affect survival and growth in multiple ways and thus would correlate with
447 biomass in future years (hypothesis L1). However we found no support for either of these SST
448 covariates as explanatory variables for the Jul-Sep catch and only weak support (based on AIC)
449 for March-May SST in the current season for explaining variability in Oct-Mar catch. The fall
450 average SST in the prior season did not explain variability in either Jul-Sep or Oct-Mar catch.
451 See Tables B4, B5 and B6. We also found no correlation between the ONI index (hypothesis
452 A2) for either the Jul-Sep or Oct-Mar catch (Tables B7, B8 and B9).

453 Instead we found support for the covariates indirectly and directly associated with pro-
454 ductivity and food availability: upwelling intensity and surface chlorophyll. The correlation
455 between landings and upwelling was only found for upwelling in the current season. No cor-
456 relation was found when we used the upwelling index from the prior season. The correla-
457 tion between landings and upwelling was found for both Jul-Sep and Oct-Mar landings and
458 with either SST-based upwelling index: average nearshore SST along the Kerala coast during

459 June-September or the average SST nearshore versus offshore differential (UPW) off Kochi
460 in June-September (Table 2, B4, B5 and B6). These two upwelling indices are correlated but
461 not identical. The model with average June-September nearshore SST was more supported
462 than the model using the SST differential off Kochi. For Jul-Sep catch, this model with a
463 non-linear response had an adjusted R^2 of 41.0 versus an adjusted R^2 of 24.4 for the model
464 with no covariates (Table B4), and for Oct-Mar catch, the adjusted R^2 was 61.8 versus 56.6
465 (Table B5). Note, that this covariate is June-September in the current season and overlaps with
466 the July-September catch. Thus this model cannot be used to forecast Jul-Sep catch and gives
467 only a month-prior forecast for Oct-Mar, but it does help us understand what factors may be
468 influencing catch.

469 Chlorophyll-a concentration is speculated to be an important predictor of larval sardine
470 survival and growth. In addition, sardines shoal in response to coastal chlorophyll blooms,
471 which brings them in contact with the coastal fisheries. Thus chlorophyll-a is assumed to be an
472 important driver of future or current sardine catches. We only have chlorophyll-a remote sens-
473 ing data from 1998 onward. Our simplest covariate model required 5 degrees of freedom, thus
474 we were limited in the analyses we could conduct. In addition, the years, 1998-2014, have rela-
475 tively low variability in catch sizes; the logged catch sizes during this period range from 10-11
476 during Jul-Sep and 11-12 during the other three quarters. Second degree polynomial mod-
477 els were fit (Appendix C) to the average log chlorophyll-a concentration in July-September,
478 October-December, and January-March in the current and prior year. Chlorophyll-a concen-
479 tration was not a significant predictor for the Jul-Sep catch for any of the tested combinations
480 of current or prior season and quarter. The only significant effect was seen for Oct-Mar mon-
481soon catches using chlorophyll-a concentration in Oct-Dec of the prior season (Table C1). This
482 is in contrast to the results with monsoon upwelling indices, which found a correlation with
483 the current season but not prior seasons.

484 The strongest correlation however was found with the multi-year average sea surface tem-
485 perature for the nearshore waters off Kerala, latitude 8 to 11 (Table 2, B7, B8 and B9). The
486 average sea surface temperature over multiple prior years has been found to be correlated with
487 sardine recruitment in Pacific sardines (Checkley et al., 2017; Jacobson & MacCall, 1995; Lin-
488 degren et al., 2013) and southern African sardines (Boyer et al., 2001). We tested as a model
489 covariate the average nearshore SST for 2.5 years prior to the Jul-Sep catch, so January-June
490 in the current calendar year and the two prior calendar years for a 30-month average. This
491 covariate can be used for forecasting since it does not overlap with either Jul-Sep or Oct-Mar
492 catch. This covariate with a non-linear response was the best covariate for both the Jul-Sep and

493 Oct-Mar catch. For Oct-Mar catch, the model with multi-year average SST had an adjusted
494 R^2 of 67.5 versus 56.6 without. For the Jul-Sep catch, the adjusted R^2 was 41.0 versus 24.4
495 without the multi-year average SST covariate. The response curve was step-like with a nega-
496 tive effect at low temperatures and then an positive flat effect at higher temperatures (Figure
497 5). This is similar to the step-response found in studies of the correlation between average SST
498 and recruitment in Pacific sardines (Jacobson & MacCall, 1995).

499 The only other strong correlation was found for Oct-Mar catch with the DMI in the
500 prior season. The Dipole Mode Index in the prior year has been shown to be correlated with
501 nearshore anoxia off the Kerala coast (Vallivattathillam et al., 2017). However this correlation
502 was only found with the Oct-Mar catch using the base model with both N_{t-1} and S_{t-2} . The
503 correlations with the multi-year SST average was more robust and found with both Oct-Mar
504 base models and also for Jul-Sep catch.

505 Discussion

506 Sardines in all the world's ecosystems exhibit large fluctuations in abundance (Checkley et
507 al., 2017). These small forage fish are strongly influenced by natural variability in the ocean
508 environment. Upwelling, influenced by both large-scale forces such as regimes shifts and El
509 Niño/Southern Oscillation patterns (Alheit & Hagen, 1997; Schwartzlose et al., 2010) and by
510 seasonal wind and current patterns, brings nutrient rich and oxygen rich waters to the surface.
511 This drives the seasonal variability in phytoplankton resources and in turn sardine prey (Bakun
512 et al., 2008). Local variability in temperature, salinity, and oxygen levels have both direct
513 and indirect effects on sardine reproduction, recruitment and survival (Checkley et al., 2017).
514 Sardines are also influenced by competition and predation by other species and well-known for
515 their sensitivity to over-fishing which has been linked to many fishery collapses (Kripa et al.,
516 2018).

517 Many studies on Pacific sardines have looked at the correlation between ocean surface
518 temperature (SST) and recruitment. Temperature can have direct effect on larval survival and
519 growth and an indirect effect on food availability. Studies in the California Current System,
520 have found that SST explains (a portion of) year-to-year variability in Pacific sardine recruit-
521 ment (Checkley et al., 2009, 2017; Jacobson & MacCall, 1995; Lindegren & Checkley, 2012)
522 and that the average nearshore temperature over multiple seasons is the relevant explanatory
523 variable. Similar to these studies, we found that the average nearshore SST over multiple sea-

sons was the covariate that explained the most variability in catch both in the monsoon and post-monsoon months. McClatchie et al. (2010) found no relationship with SST and Pacific sardine recruitment, however their analysis used a linear relationship while the other studies, and ours, that found a relationship allowed non-linearity. Both Jacobson and MacCall (1995) and Checkley et al. (2017) found a step-like response function for temperature: below a threshold value the effect of temperature was linear and above the threshold, the effect was flat and at lower temperatures the effect was negative and became positive as temperature increased. Our analysis found a similar pattern with a negative effect when the 2.5-year average temperature was below 28.35°C and positive above and with the positive effect leveling off above 28.5°C (Figure 5).

There were four outlier years when Oct-Mar oil sardine landings were much lower than expected based on prior catches: 1986, 1991, 1994 and 2013 (Figure 6, Panel C). The 2.5-year average SST explained the collapses in 1986 and 1991; the size of the prediction was much closer to the observed catch (Figure 6, Panel D). The largest collapse was in 1994 and the most recent, in our dataset, was 2013. The 2.5-year average SST did not explain the 1994 nor 2013 collapses. There was no change in the size of the residual with and without the covariate. The same pattern was seen for the Jul-Sep catch, with the exception that 1991 did not have unusually low Jul-Sep catch. The 2.5-year average SST reduced the prediction error for 1986, but did not (appreciably) for 1994 nor 2013. In fact, none of the covariates we tested explained the lower than expected 1994 and 2013 catches. The causes of these unusual declines appear unrelated to the environmental factors we studied, suggesting either that other factors, biological or anthropogenic, drove these declines or that a particular combination of environmental factors led to the declines. It should also be noted that our upwelling indices captured only one aspect of upwelling: the nearshore intensity. Other aspects of upwelling also affect oil sardines, such as the spatial extent both along the coast and off the coast and the timing of the start of upwelling.

Seasonal productivity in the Southeast Arabian Sea upwelling system is driven by the summer monsoon, which causes strong coastal upwelling that moves from the south to the north over the summer. This drives a strong seasonal pattern of zooplankton abundance (Figure 3). There are biological reasons to expect a positive relationship between upwelling intensity and landings: upwelling drives productivity and higher food resources in the current season which leads to higher larval and juvenile survival and higher numbers of 0-year fish in the landings and brings sardines into the nearshore to feed where they are exposed to the fishery. These positive effects on 0-year fish would cascade to future seasons also. Despite the strong connection

558 between sardine recruitment, growth and survival with upwelling, we found that none of our
559 upwelling indices in the prior season explained the year-to-year variation in landings. We did
560 find that the upwelling intensity in the current season explained variability in landings in the
561 current season, however, the effect was negative not positive and the negative effect emerged
562 at extremely high upwelling (Figure 5). This negative effect is not surprising. Extremely high
563 upwelling transports larval sardines offshore and creates regions of low oxygen which sar-
564 dines avoid (Gupta et al., 2016). What was surprising is that the effect was not uni-modal,
565 with a positive effect at low to moderate upwelling and becoming negative for extremely high
566 upwelling.

567 One of the purposes of our research was to investigate environmental covariates that would
568 improve prediction of landings, not simply explain variability. To test this, we used leave-
569 one-out cross-validation to generate out-of-sample prediction errors. The predictions were
570 compared to a standard null prediction: the catch observed in the prior year. With this null
571 model, whatever the catch was in the prior year (same season) is the the prediction. In Table
572 2, the out-of-sample prediction errors are shown in the LOOCV RMSE (leave-one-out cross-
573 validation root-mean-square error) column. All the GAM catch models (M0, M1 and M2) have
574 better predictive performance than the null model. The next question is whether the covariates
575 improve the predictions compared to the GAM catch models (without covariates). For the Oct-
576 Mar catch, the 2.5-year average SST improved the prediction the most; 22.1% for the more
577 complex GAM model (M1) and 17.5% for the simpler GAM model (M2). For Jul-Sep catch,
578 only Jun-Sep SST in the current season reduced the prediction error and only by 8.2%. The
579 Jul-Sep catch is difficult to forecast. It has high variability that is poorly explained by past
580 catch or the environment. In contrast, the Oct-Mar catch is much better explained by prior
581 catch (higher R^2) and the forecast errors (LOOCV RMSE) are smaller.

582 Conclusions

583 Remote sensing via satellites can be used to detect changes in ocean physical, biological and
584 chemical properties, such as surface temperature, winds, surface height, surface waves, rain-
585 fall and surface salinity, as well as the ecosystem and water quality changes. Unlike in-situ
586 measurements, environmental measures from remote sensing can be acquired rapidly and over
587 large regions. However, which environmental covariates will improve forecasts is not obvious
588 from oil sardine life history alone. We tested many of the covariates that are known or have
589 been postulated to have an effect on sardine spawning, growth and survival (Table 1): pre-

590 cipitation, upwelling indices, ocean temperature and chlorophyll-a in various critical months
591 of the sardine life-cycle. We found that the multi-year average nearshore ocean temperature
592 explained the most variability in the landings and addition of this covariate to the catch mod-
593 els improved out-of-sample prediction. This covariate is not as directly tied to stages of the
594 oil sardine life cycle as the other covariates we tested, though it does integrate over multiple
595 influences (upwelling strength and temperature) over multiple years.

596 The temperature of the Western Indian Ocean, of which the Southeast Arabian Sea is a
597 part, has been increasing over the last century at a rate higher than any other tropical ocean
598 (Roxy et al., 2014) and the warming has been most extreme during the summer monsoon
599 months. This ocean climate change is affecting oil sardine distributions, with significant land-
600 ings now occurring north of Goa (Vivekanandan et al., 2009). Continued warming is expected
601 to affect the productivity of the region via multiple pathways, including both the direct effects
602 of temperature change on the physiology and behavior of organisms and multiple of indirect
603 effects (Moustahfid et al., 2018). These indirect effects include changes to salinity, oxygen
604 concentrations, currents, wind patterns, ocean stratification and upwelling spatial patterns, phe-
605 nology, and intensity. Incorporating environmental covariates into landings forecasts has the
606 potential to improve fisheries management for small pelagics such as oil sardines in the face of
607 a changing ocean environment (Haltuch et al., 2019; Tommasi et al., 2016). However, moni-
608 toring forecast performance and covariate performance in models will be crucial as a changing
609 ocean environment may also change the association between landings and average sea surface
610 temperature.

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863 **Figure Legends**

864 Figure 1. Southwest coast of India with the latitude/longitude boxes used for the satellite data.
865 Kerala State is marked in grey and the oil sardine catch from this region is being modeled.
866 For the SST covariate, ‘nearshore’ SST (ns-SST) was the average of boxes 2 to 5 (0 to 80km
867 offshore), and ‘regional’ SST (r-SST) was the average of boxes 2 to 5 and 7 to 10 (0 to 160km
868 offshore).

869 Figure 2. Quarterly catch data 1956-2014 from Kerala. The catches have a strong seasonal
870 pattern with the highest catches in quarter 4 (Oct-Dec). Note that the fishery is closed July 1 to
871 mid-August, thus the fishery is only open 1.5 months in quarter 3 (Jul-Sep). The mean catch
872 (metric tonnes) in quarters 1 to 4 are 38, 19.2, 30.9, and 59.9 metric tonnes respectively.

873 Figure 3. Key oil sardine life history events overlaid on the monthly sea surface tempera-
874 ture in the nearshore and offshore and the nearshore chlorophyll-a concentration.

875 Figure 4. Remote sensing covariates used in the analysis. All data are monthly averages
876 over Box 4 in Figure 1 on the Kerala coast off of Kochi. Panel A) Upwelling Index. The
877 upwelling index is the difference between the nearshore sea surface temperature (SST) and the
878 offshore SST defined as 3 degrees longitude offshore. Panel B) Nearshore surface chlorophyll-
879 a (CHL). The CHL data are only available from 1997 onward. Panel C) Sea surface temper-
880 ature from Advanced Very High Resolution Radiometer (AVHRR). Panel D) Average daily
881 rainfall (mm/day) off the Kerala coast.

882 Figure 5. Effects of the two most influential covariates estimated from the GAM models:
883 2.5 year average nearshore (boxes 2-5) SST and upwelling intensity in June-September (spawn-
884 ing months). Panel A) Effect of the 2.5 year average nearshore SST on Jul-Sep catch (late
885 spawning and early post-spawning months). Panel B) Effect of upwelling (nearshore/offshore
886 SST differential) during June-September in the current season on Jul-Sep catch. The index
887 is the difference between offshore and inshore SST, thus a negative value indicates warmer
888 coastal surface water than offshore. Panel C) Effect of the 2.5 year average nearshore SST
889 on Oct-Mar catch (post-monsoon, age-0, -1, -2 year fish). Panel D) Effect of upwelling
890 (nearshore/offshore SST differential) during June-September in the current season on Oct-Mar
891 catch.

892 Figure 6. Fitted versus observed catch with models with and without the 2.5 year aver-
893 age nearshore SST included as a covariate. The line is the one-to-one line (prediction equals
894 observed). Panel A) Fitted versus observed log catch in Jul-Sep (late monsoon) with only Oct-

895 Mar catch in the previous season as the covariate: $S_t = s(N_{t-1}) + \varepsilon_t$. Panel B) Fitted versus
896 observed log catch in Jul-Sep with the 2.5-year average SST added as a covariate to the model
897 in panel A. This model was: $S_t = s(N_{t-1}) + s(V_t) + \varepsilon_t$. Panel C) Fitted versus observed log Oct-
898 Mar catch with only Oct-Mar catch in the previous season and Jul-Sep catch two seasons prior
899 as the covariates: $N_t = s(N_{t-1}) + s(S_{t-2}) + \varepsilon_t$. Panel D) Fitted versus observed log Oct-Mar
900 catch with 2.5-year average SST (V) added. This model was $N_t = s(N_{t-1}) + s(S_{t-2}) + s(V_t) + \varepsilon_t$.

Table 1. Hypotheses for covariates affecting landings. The left column shows the model tested. The value to the left of \sim is the response: S_t Jul-Sep catch or N_t Oct-Mar catch in the current season. To the right of \sim are the explanatory variables. S_{t-1} is Jul-Sep catch and N_{t-1} is the Oct-Mar catch in the prior season. DD = hypotheses related to density-dependence. S = hypotheses related to catch during the spawning months. L = hypotheses related to larval and juvenile growth and survival. A = hypotheses affecting all ages. See the introduction for discussion of the literature on the effects of environmental covariates on sardine landings. ns-SST = nearshore SST. r-SST = regional (nearshore and offshore) SST. t , $t - 1$, and $t - 2$ subscripts indicate current, prior, and two seasons prior. ONI and DMI are ENSO indices described in the methods.

Hypothesis	Description
DD1 $S_t \sim N_{t-1}$	S_t is age 1+ age fish and reflects the age 0-2 fish in N_{t-1} which have aged 3-6 months (Nair et al. 2016).
DD2 $S_t \sim S_{t-1} + S_{t-2}$ $N_t \sim S_{t-1} + S_{t-2}$	S_t is age 1+ fish and is dominated by spent post-spawning fish (Nair et al. 2016). S_t should be correlated with the abundance of the spawning stock and cohort strength. If cohort strength persists over time, landings in the current season should correlate with Jul-Sep landings in the previous two seasons.
DD3 $N_t \sim N_{t-1}$	N_{t-1} includes age 0, 1, 2 and 2+ fish (Nair et al. 2016). Thus the age 0, 1, and 2 fish will appear 1 year later in the N_t landings.
S1 $S_t \sim$ Jun-Jul precipitation in t	The magnitude of precipitation in June-July directly or indirectly signals mature fish to spawn offshore, after which the spent adults migrate inshore to the fishery (Anthony Raja 1969, 1974; Srinath 1998).
S2 $S_t \sim$ Apr-Mar precipitation in t	If there is pre-monsoon rain during Apr-May then spawning begins earlier and spent adults return to the fishery sooner and are exposed to the fishery longer.
S3 $N_t \sim$ Apr-Mar precipitation in t	Precipitation is an indicator of climatic conditions during the period of egg development. This affects the success of spawning, thus affecting the 0-year class and the Oct-Mar catch (Antony Raja 1969, 1974; Srinath 1998).
S4 $S_t \sim$ Jun-Sep UPW in t	High rates of upwelling (UPW) drive mature fish further offshore and leading to lower exposure to the fishery. This movement may be driven by advection of high phytoplankton biomass offshore or upwelling of hypoxic water to the surface (Murty and Edelman 1970, Antony Raja 1973, Pillai et al. 1991).
S5 $S_t \sim$ Mar-May r-SST in t $N_t \sim$ Mar-May r-SST in t	Extreme heating events prior to the monsoon drive mature fish from the spawning areas, resulting in poor recruitment and fewer age 0 fish in the Oct-Mar catch (Antony Raja 1973, Pillai et al 1991).

Table 1. Continued.

Hypothesis	Description
L1 $S_t \sim \text{Oct-Dec ns-SST } t - 1$ $N_t \sim \text{Oct-Dec ns-SST } t - 1$	Larval and juvenile growth and survival is affected by temperature. SST in the critical post-monsoon window, when somatic growth of the age 0 fish is highest, should be correlated with future abundance.
L2 $S_t \sim \text{Jun-Sep UPW in } t - 1 \text{ & } t$ $N_t \sim \text{Jun-Sep UPW in } t - 1 \text{ & } t$	Higher upwelling rates lead to greater phytoplankton productivity, better larval and juvenile growth, and higher future landings. However, extreme upwelling may decrease catch due to hypoxic conditions or advection of phytoplankton biomass.
L3 $S_t \sim \text{CHL in } t - 1 \text{ & } t$ $N_t \sim \text{CHL in } t - 1 \text{ & } t$	Surface chlorophyll-a is a proxy for phytoplankton abundance and thus food availability, supporting greater fish abundance, and catch, in the current and future years.
A1 $S_t \sim 2.5\text{-yr ave. ns-SST}$ $N_t \sim 2.5\text{-yr ave. ns-SST}$	Spawning, early survival, and recruitment of sardines depends on a multitude of cascading factors summarized by the multi-year average nearshore SST. The multi-year average SST has been correlated with recruitment and abundance in other sardines (Checkley et al. 2017) and is related to the optimal temperature windows for sardines (Takasuka et al. 2007).
A2 $S_t \sim \text{ONI in } t - 1$ $N_t \sim \text{ONI in } t - 1$	The El Niño/Southern Oscillation (ENSO) impacts a range of environmental parameters (e.g., precipitation, SST, frontal zones, winds, etc.) that may impact spawning and early survival of sardines (Supraba et al. 2016, Rohit et al. 2018).
A3 $S_t \sim \text{DMI in } t - 1$ $N_t \sim \text{DMI in } t - 1 \text{ & } t$	Negative Dipole Mode Index (DMI) values in Sep-Nov are associated with anoxic events on the Kerala coast while positive values are associated with an absence of such events (Vallivattathillam et al. 2017).

Table 2. Top covariates for the monsoon (Jul-Sep) and post-monsoon (Oct-Mar) catch (S_t and N_t) models. M0, M1 and M2 are the base models with only prior catch as covariates. To the base models, the environmental covariates are added. ns-SST is nearshore (0-80km) and r-SST is regional (0-160km). The full set of nested covariate models and tests are given in Appendix B. The LOOCV RMSE (root mean square error) is the out-of-sample prediction error. The LOOCV RSME for the null model for Jul-Sep catch was 1.599 and the LOOCV RMSE for the null model for Oct-Mar catch was 1.015. The fitted versus observed catches from the models with and without 2.5-year average ns-SST are shown in Figure 6.

Model	Residual df	Adj. R2	RMSE	AIC	LOOCV RMSE
Jul-Sep catch models with covariates					
V_t = Jun-Sep ns-SST current season					
W_t = Jun-Sep Bakun-UPW current season					
Z_t = 2.5-year ave ns-SST					
$M0: \ln(S_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$	28.6	24	1.184	109.52	1.299
$\Rightarrow \ln(S_t) = M0 + s(V_t)$	25.9	41	1.007	103.43	1.192
$\ln(S_t) = M0 + \beta W_t$	27.6	28	1.133	108.66	1.404
$\ln(S_t) = M1 + s(Z_t)$	26.2	41	1.011	103.26	1.338
Oct-Mar catch models with covariates					
U_t = Mar-May r-SST current season					
V_t = Jun-Sep ns-SST current season					
Z_t = 2.5-year ave ns-SST					
X_{t-1} = fall DMI prior season					
$M1: \ln(N_t) = \alpha + s(\ln(N_{t-1})) + s(\ln(S_{t-2})) + \varepsilon_t$	24.8	57	0.713	79.53	1.062
$\ln(N_t) = M1 + s(U_t)$	22	64	0.623	75.68	1.008
$\ln(N_t) = M1 + \beta W_t$	23.8	63	0.648	75.57	1.042
$\Rightarrow \ln(N_t) = M1 + s(Z_t)$	22.7	67	0.597	71.88	0.827
$\ln(N_t) = M1 + s(X_{t-1})$	21.1	68	0.58	72.69	0.89
$M2: \ln(N_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$	27.6	45	0.836	84.75	0.966
$\ln(N_t) = M2 + s(V_t)$	24.7	47	0.791	85.97	0.976
$\ln(N_t) = M2 + \beta W_t$	26.6	52	0.772	81.79	0.927
$\Rightarrow \ln(N_t) = M2 + s(Z_t)$	25.3	60	0.688	76.34	0.796
$\ln(N_t) = M2 + s(X_t)$	23.7	43	0.8	88.43	0.969

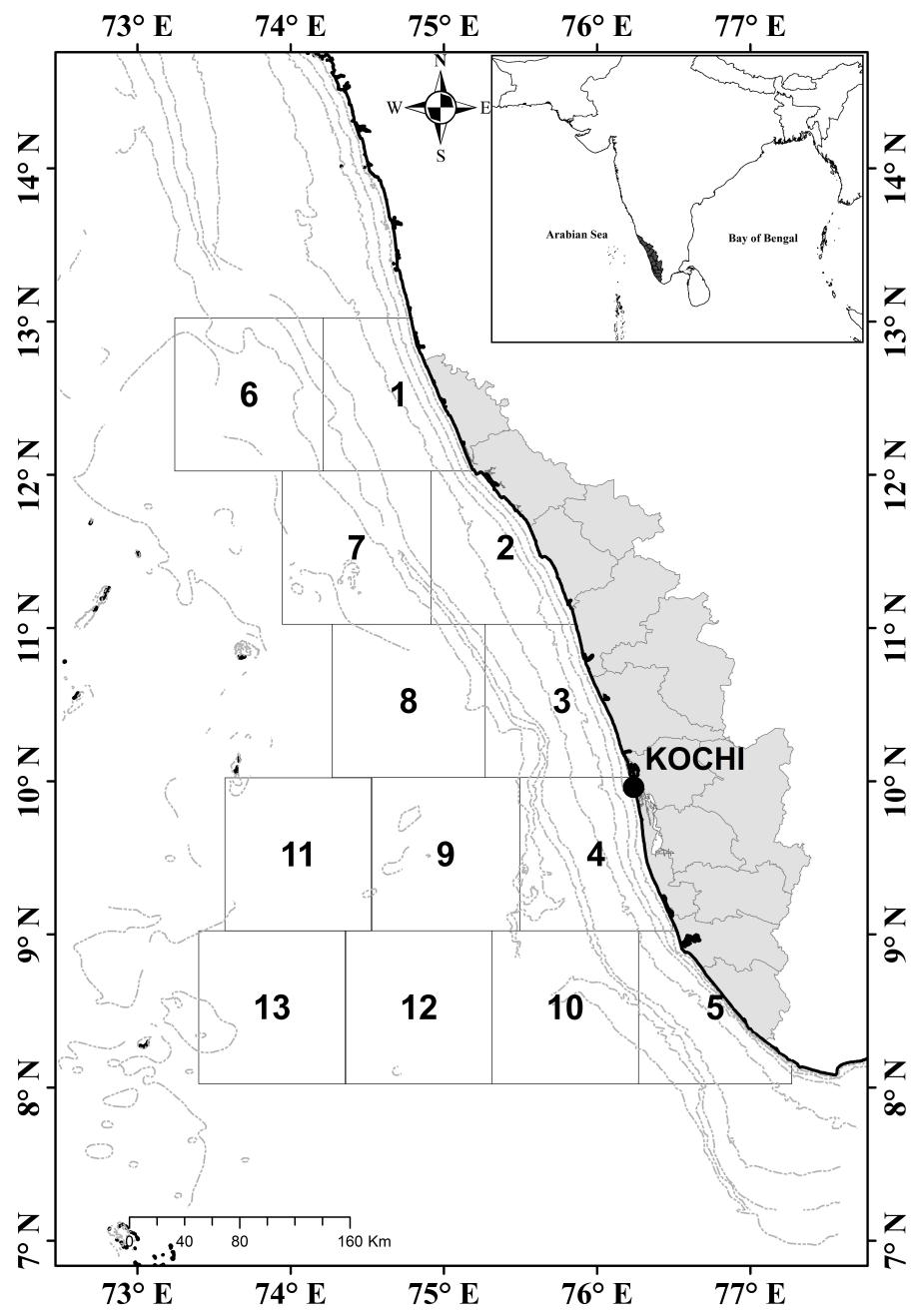


Figure 1

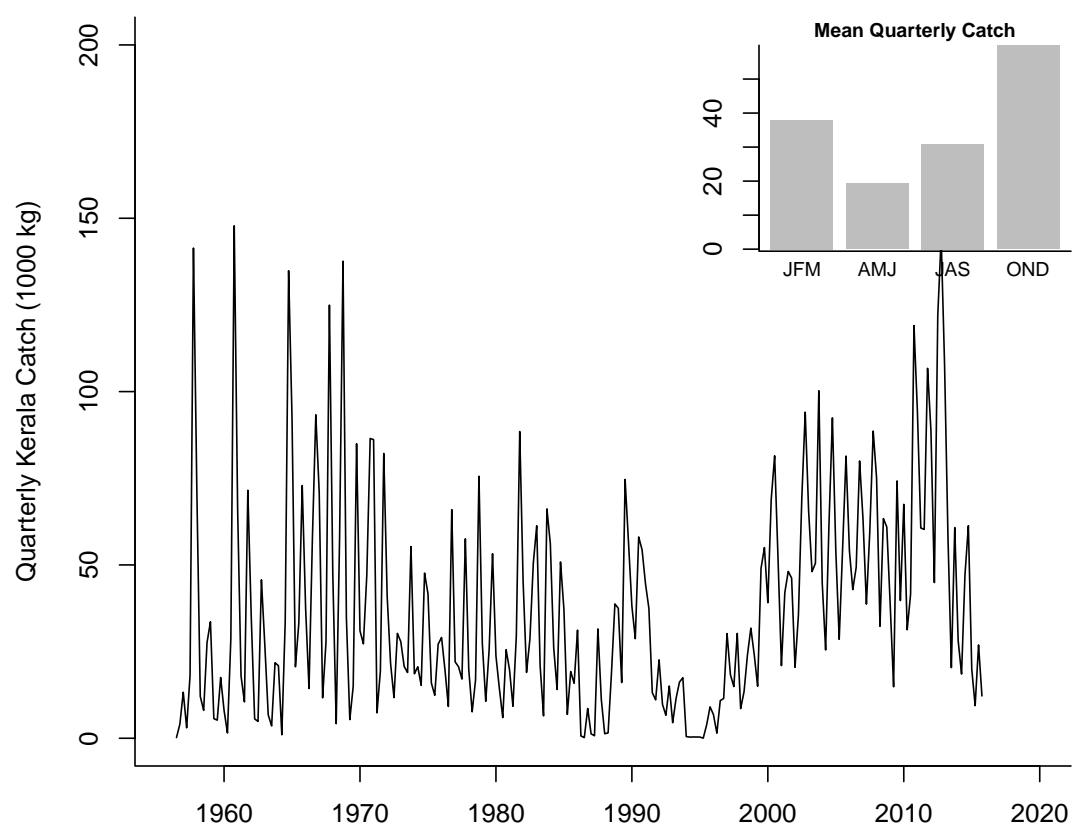


Figure 2

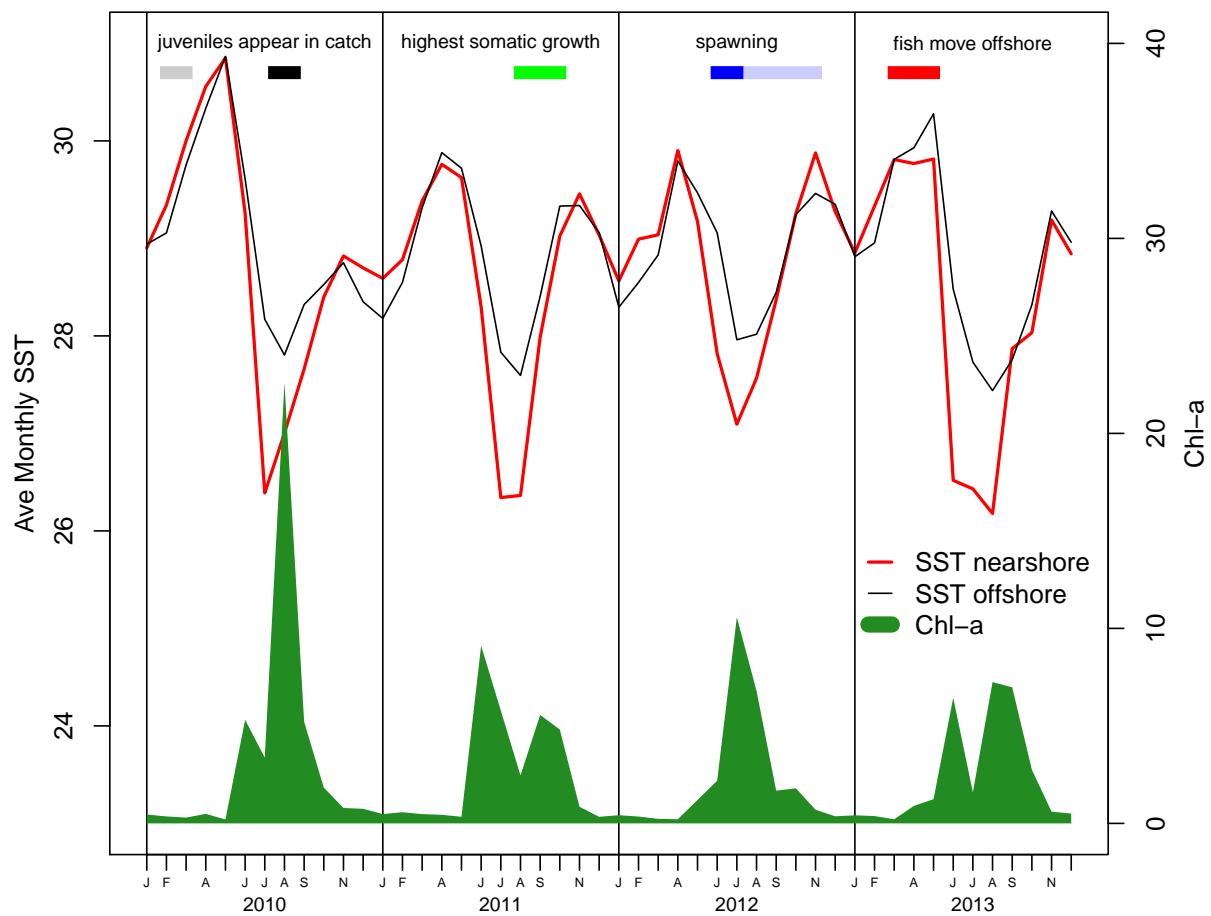


Figure 3

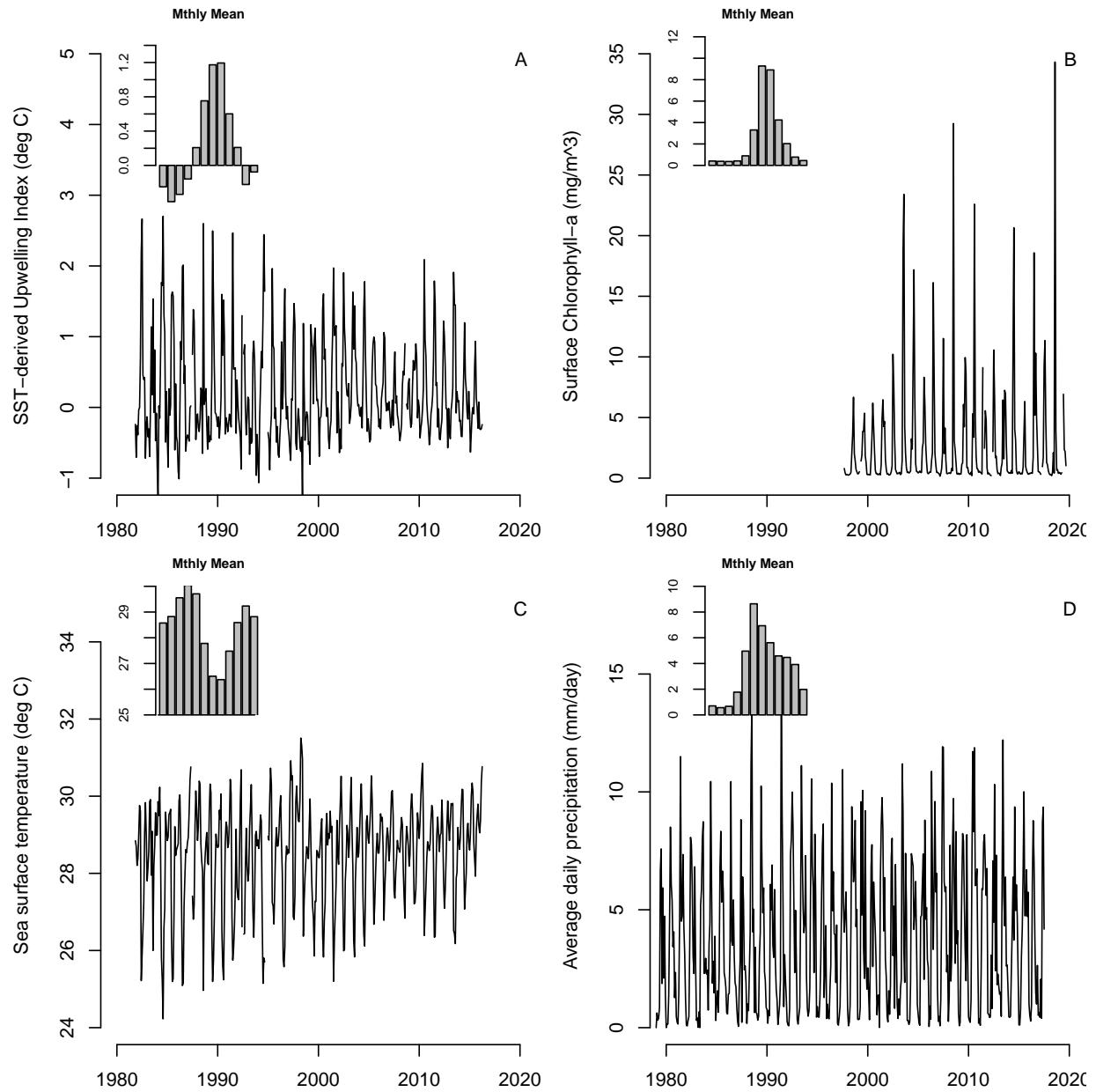


Figure 4

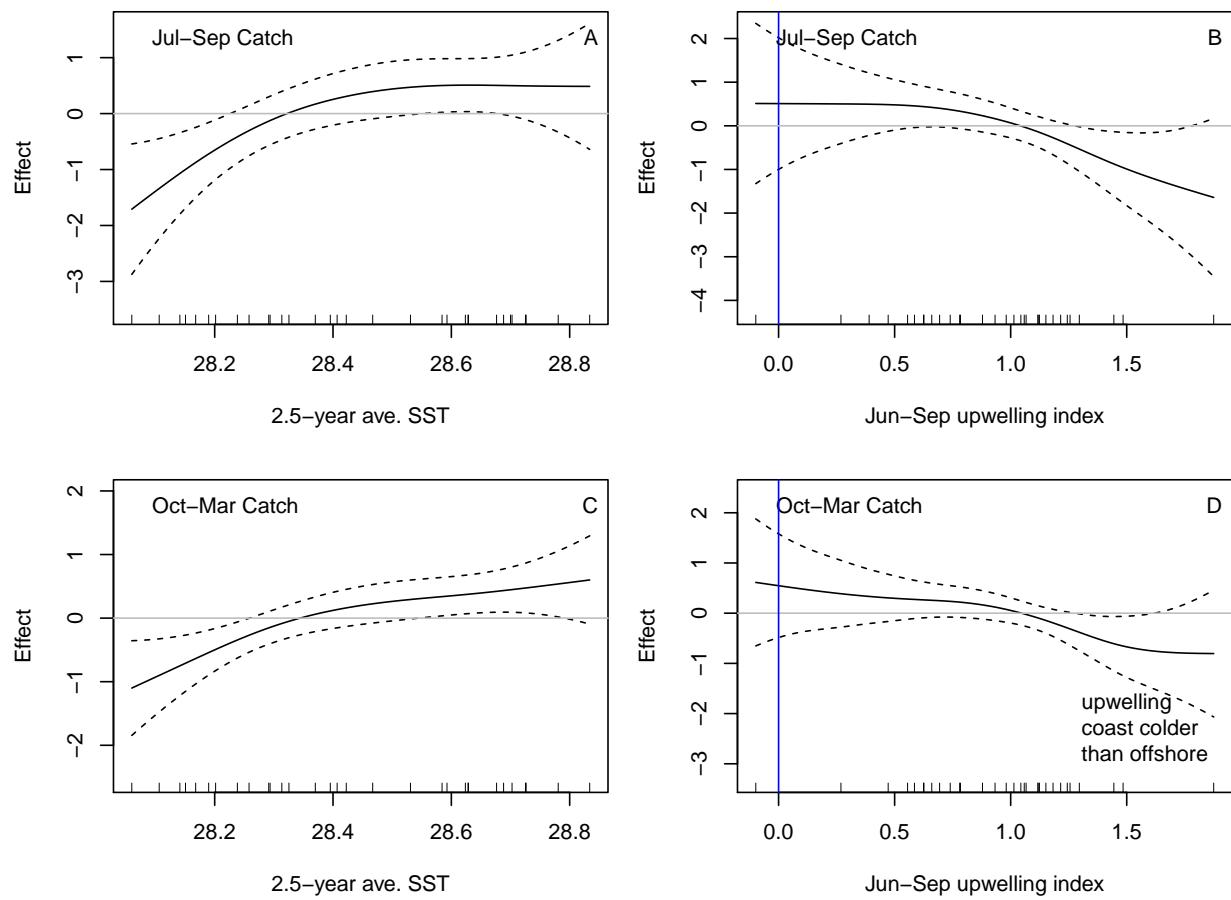


Figure 5

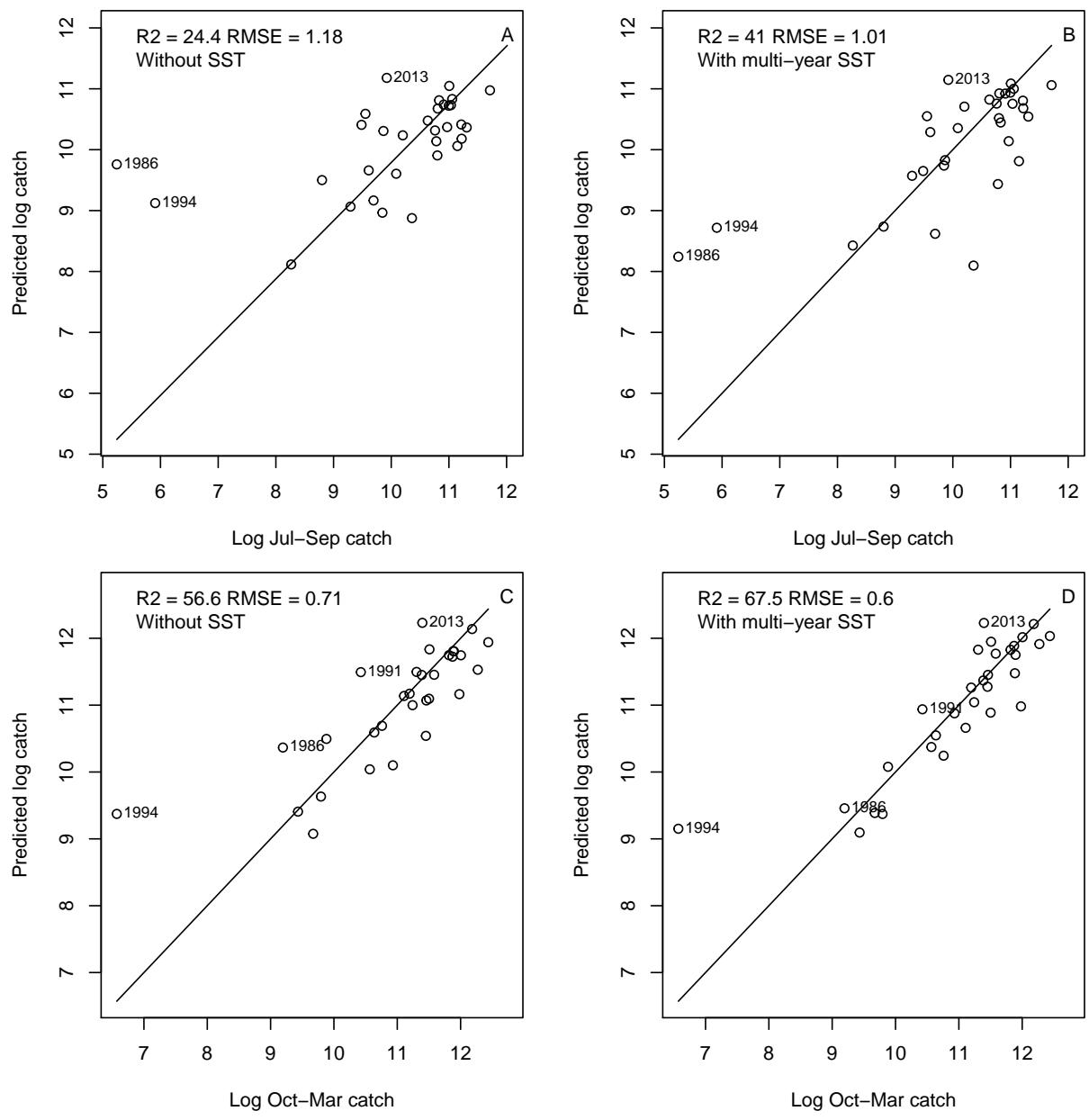


Figure 6