

¹ Fishing in a warming ocean: influence of changing temperature
² and upwelling intensity on Indian oil sardine (*Sardinella*
³ *longiceps*) landings

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⁵ Eli E. Holmes¹, Smitha B.R.², Nimit Kumar³, Sourav Maity³, David M.
⁶ Checkley⁴, Mark L. Wells⁵, Vera L. Trainer¹

- ⁷ 1. Northwest Fisheries Science Center, NOAA, Seattle, WA.
⁸ 2. Centre for Marine Living Resources and Ecology, MoES, Kochi, India.
⁹ 3. Indian National Centre for Ocean Information Services, Hyderabad, India.
¹⁰ 4. Scripps Institution of Oceanography, UC San Diego, San Diego, CA.
¹¹ 5. School of Marine Sciences, University of Maine, Orono, ME.

¹² **Running title:** Modeling Indian oil sardine landings

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Abstract

Commercial landings of sardine are known for strong year-to-year fluctuations. A key driver is thought to be environmental variability, to which small forage fish are especially sensitive. We examined the environmental drivers associated with landings fluctuations in the Indian oil sardine using a 32-year time series off Kerala state. Past research suggested a variety of influential variables: precipitation, upwelling intensity, ocean temperature (SST), chlorophyll density and large-scale coupled atmosphere–ocean phenomena (ENSO). Using the life-history of the Indian oil sardine, we developed hypotheses concerning how these environmental variables might affect landings and tested them using generalized additive models which allow non-linear response curves and dynamic linear models which allow time-varying responses. We found significant correlation for only two variables: upwelling intensity and the multi-year average nearshore SST. Both monsoon and post-monsoon landings were correlated with upwelling intensity in June-September. Upwelling intensity has both a positive effect (fueling higher food availability) and a negative effect at extreme intensity (bringing poorly oxygenated water to the surface). However, the most significant correlation (adjusted R^2 of 67.5%) was between the 2.5 year average nearshore SST and post-monsoon landings. The multi-year average SST also been identified as a predictor for Pacific sardine and southern African sardine fluctuations, suggesting that the average SST over the sardine life-span successfully captures a variety of factors which predict future abundance. The temperature in the Western Indian Ocean has been increasing in the last decade and the warming has been most extreme during the summer monsoon. Our work highlights that these changes in sea temperature are likely to affect future oil sardine landings.

Keywords: Indian oil sardine, catch prediction, GAM modeling, climate, sea surface temperature, remote sensing, Southeastern Arabian Sea

51 **Introduction**

52 Environmental variability is known to be a key driver of population variability of small forage
53 fish such as sardines, anchovy and herring (Alheit & Hagen, 1997; Checkley et al., 2017; Cury
54 et al., 2000). In particular, ocean temperature and upwelling dynamics, along with density-
55 dependent feedback, have been identified as important in affecting recruitment success and
56 biomass of European and Pacific sardines (*Sardina pilchardus* and *Sardinops sagax*) (Alheit et
57 al., 2012; Garza-Gil et al., 2015; Jacobson & MacCall, 1995; Lindegren & Checkley, 2012;
58 Lindegren et al., 2013; Rykaczewski & Checkley, 2008). Like other sardines, the Indian oil
59 sardine (*Sardinella longiceps* Valenciennes, 1847) shows strong inter-annual fluctuations and
60 larger decadal booms and busts. The Indian oil sardine offers an instructive case study to
61 investigate the effects of environmental variability, particularly temperature and upwelling dy-
62 namics, as they occupy an ocean system that is warmer than that occupied by other sardines
63 and have a strong seasonal cycle driven by the Indian summer monsoon.

64 The Indian oil sardine is one of the most commercially important fish resources along the
65 southwest coast of India (Figure 1) and historically has comprised approximately 25% of the
66 marine catch in the Indian fisheries (Vivekanandan et al., 2003). Landings of the Indian oil sar-
67 dine are highly seasonal, peaking after the summer monsoon period in October-December and
68 reaching a nadir in spring before the summer monsoon in April-June (Figure 2). At the same
69 time, the landings of this small pelagic finfish are highly variable from year to year. Small
70 pelagics are well known to exhibit high variability in biomass due to the effects of environ-
71 mental conditions on survival and recruitment, but in this fishery, environmental conditions
72 also affect exposure of sardines to the fishery. Until recently, the Indian oil sardine fishery was
73 artisanal and based on small human or low powered boats with no refrigeration. The fishery
74 was confined to nearshore waters, and thus migration of sardines in and out of the coastal zone
75 greatly affected exposure to the fishery and hence landings.

76 Researchers have examined a variety of environmental variables for their correlation with
77 landings of the Indian oil sardine. Precipitation during the monsoon and the day of the mon-
78 soon arrival are thought to act as either a direct or indirect cue for spawning (Antony Raja,
79 1969, 1974; Jayaprakash, 2002; Murty & Edelman, 1966; Pitchaikani & Lipton, 2012; Sri-
80 nath, 1998; Xu & Boyce, 2009). Although many studies have looked for and some have found
81 correlations between precipitation and landings, the reported relationships are positive in some
82 studies and negative in others (Madhupratap et al., 1994). Researchers have also looked for and
83 found correlations with various metrics of upwelling intensity (Jayaprakash, 2002; Longhurst

& Wooster, 1990; Madhupratap et al., 1994; Murty & Edelman, 1966; Srinath, 1998; Thara, 2011), with direct measures of productivity such as nearshore zooplankton and phytoplankton abundance (George et al., 2012; Madhupratap et al., 1994; Menon et al., 2019; Nair, 1952; Nair & Subrahmanyam, 1955; Piontkovski et al., 2015; Pitchaikani & Lipton, 2012), and with near-shore sea surface temperature (SST) (Annigeri, 1969; Pillai, 1991; Prabhu & Dhulkhed, 1970; Supraba et al., 2016). SST can affect both somatic growth rates and juvenile survival but in this system also can cause fish to move off-shore and away from the shore-based fishery. The multi-year average sea temperature is postulated to have effects on recruitment and the survival of larval and juvenile sardines, which affect the later overall abundance (Checkley et al., 2017; Takasuka et al., 2007). The El Niño/Southern Oscillation (ENSO) has a cascading effect on all the aforementioned environmental parameters (SST, precipitation, upwelling) and correlations have been found between ENSO indices and sardine landings (Rohit et al., 2018; Supraba et al., 2016) and coastal anoxia events which affect sardines (Vallivattathillam et al., 2017).

In this paper, we study the utility of environmental covariates from remote sensing to explain year-to-year variability in oil sardine landings using a long-term time series of quarterly Indian oil sardine landings from the southwest coast of India. This time series is derived from a stratified sampling design that surveys the fishery landing sites along the southeast Indian coast and was first implemented in the 1950s (Srinath et al., 2005). The goal of the work presented here is to identify environmental covariates which can explain catch variability and improve the accuracy of short-term catch forecasts. Landings of oil sardines are determined by a combination of biomass, catchability, and effort. An auto-correlated catch model (ARIMA) can capture smooth changes in biomass and effort that cause landing variability, but the environmental variability adds a large component of year-to-year variability to small pelagic landings. In the Indian oil sardine system, catchability is strongly affected by the inshore versus offshore distribution of sardine. The fishery is restricted to the nearshore < 50 km offshore (Rohit et al., 2018) and when the sardines move offshore to spawn or to avoid hypoxic or excessive warm water, they are no longer available to the fishery. Thus the environment has a strong impact on catchability. Biomass is driven by recruitment and survival are linked to environmental factors which determine food resources. The covariates which we study (Table 1) are linked to aspects of oil sardine life-history that are expected to affect catch via catchability or biomass. Covariates from remote sensing are the focus because they are available over a wide spatial extent at a daily and monthly resolution thus are practical for use in operational forecasts. A better understanding of how and whether remote sensing data explains variation in seasonal

¹¹⁸ catch will support future efforts to use satellite data to improve catch forecasts.

¹¹⁹ **Catch modeling versus biomass modeling**

¹²⁰ Modeling and forecasting landings data using statistical models fit to annual or seasonal catch
¹²¹ time series has a long tradition in fisheries and has been applied to many species (Cohen &
¹²² Stone, 1987; Farmer & Froeschke, 2015; Georgakarakos et al., 2006; Hanson et al., 2006;
¹²³ Lawer, 2016; Lloret et al., 2000; Mendelssohn, 1981; Nobel & Sathianandan, 1991; Prista
¹²⁴ et al., 2011; Stergiou & Christou, 1996), including oil sardines (Srinath, 1998; Venugopalan
¹²⁵ & Srinath, 1998). These models can be used to identify the variables correlated with catch
¹²⁶ fluctuations and can be used to provide landings forecasts which are useful for fishery managers
¹²⁷ and the fishing industry. An example of the former is using catch forecasts to set or give
¹²⁸ warnings of seasonal fishery closures if the forecasted catch is higher than the allowed catch
¹²⁹ limits (Farmer & Froeschke, 2015). An example of the latter is the annual Gulf and Atlantic
¹³⁰ menhaden forecast produced by NOAA Fisheries (Hanson et al., 2006; Schaaf et al., 1975).
¹³¹ This multiple regression model has been used for the last 45 years to produce an annual forecast
¹³² of menhaden landings, which is used for planning purposes by the industry, not only the fishers
¹³³ but also fish catch sellers and buyers, businesses which provide fisheries gear, and banks which
¹³⁴ provide financing (Hanson et al., 2006).

¹³⁵ Unfortunately long-term biomass estimates are not possible for the Indian oil sardine.
¹³⁶ Length- or age-structured models (e.g. virtual population analysis) which produce biomass
¹³⁷ estimates are not possible due to the lack of effort and catch-at-age information for the fishery.
¹³⁸ The available long-term effort data are indirect (boat composition of the fishery a multi-year
¹³⁹ intervals) and estimates of number of trips or hours fishing are only available in a few recent
¹⁴⁰ years, and the data available are approximate given the vessel diversity of the fishery and
¹⁴¹ sampling constraints. Nonetheless it is the case that the number and size of boats involved in
¹⁴² the fishery has been increasing. Oil sardines are caught primarily by ring seines, which were
¹⁴³ introduced in the early 1980s. Ring seines of different sizes are used on both traditional small
¹⁴⁴ boats and on large mechanized ships (Das & Edwin, 2018). Since 1985, the ring seine fishery
¹⁴⁵ has expanded steadily in terms of horsepower, size of boats, and length of nets. There are
¹⁴⁶ concerns that over-fishing is a factor in the most recent oil sardine declines after 2015 (Kripa
¹⁴⁷ et al., 2018). Our base auto-correlated catch model is used to capture smooth trends due to
¹⁴⁸ changes due to steadily increasing effort.

¹⁴⁹ For the purpose of our study, the assumption of a tight relationship between landings and

abundance is not necessary. The objective is to understand what drives landings variability, whether it be due to abundance variability or due to exposure to the fishery (by being closer to shore). That said, Indian oil sardine landings are often assumed to reflect the total abundance for reasons specific to the species and the fishery (cf. Kripa et al., 2018). Historically, the fishery was artisanal: small boats with small motors, no refrigeration, and limited to the near shore. The ring seine was introduced in the 1980s, but widespread mechanization of the fleet is a very recent development. Fishers with small boats have limited ability to target the stock, at least not to the degree that landings remain constant as a stock declines. That pattern can be observed in a large, mobile, highly mechanized fleet. The fishery is unregulated, except for a brief closure during the monsoon months, thus the landings are not being affected by area closures and catch limits. Finally, the fishery is dispersed along the entire coastline rather than being focused from a few large ports. Again, for our objectives, it is not necessary that landings be a tight index of biomass, but there are many reasons to assume that this relationship is strong.

Study Area

Our analysis focuses on the Kerala coast (Figure 1) region of India, where the majority of the Indian oil sardines are landed and where oil sardines comprise ca. 40% of the marine fish catch (Srinath, 1998; Vivekanandan et al., 2003). This area is in the Southeast Arabian Sea (SEAS), one of world's major upwelling zones, with seasonal peaks in primary productivity driven by upwelling caused by winds during the Indian summer monsoon (Habeebrehman et al., 2008; Madhupratap et al., 2001) between June and September. Within the SEAS, the coastal zone off Kerala between 9°N to 13°N has especially intense upwelling due to the combined effects of wind stress and remote forcing (BR, 2010; BR et al., 2008). The result is a strong temperature differential between the near-shore and off-shore and high primary productivity and surface chlorophyll in this region during summer and early fall (BR, 2010; Chauhan et al., 2011; Habeebrehman et al., 2008; Jayaram et al., 2010; Madhupratap et al., 2001; Raghavan et al., 2010). The primary productivity peaks subside after September while mesozooplankton abundances increase and remain high in the post-monsoon period (Madhupratap et al., 2001).

Oil sardine life cycle and fishery

The Indian oil sardine fishery is restricted to the narrow strip of the western India continental shelf, within 20 km from the shore (Figure 1). The yearly cycle of the fishery begins at the

181 start of spawning during June to July, corresponding with the onset of the southwest monsoon
182 (Antony Raja, 1969; Chidambaram, 1950) and the initiation of strongly cooler nearshore SST
183 due to upwelling (Figure 3). At this time, the mature fish migrate from offshore to coastal
184 spawning areas, and the spawning begins during the southwest monsoon period when temper-
185 ature, salinity and suitable food availability are conducive for larval survival (Chidambaram,
186 1950; Jayaprakash & Pillai, 2000; Krishnakumar et al., 2008; Murty & Edelman, 1966; Nair et
187 al., 2016). Although peak spawning occurs in June to July, spawning continues into Septem-
188 ber (Antony Raja, 1969; Hornell, 1910; Hornell & Nayudu, 1923; Prabhu & Dhulkhed, 1970)
189 and early- and late-spawning cohorts are evident in the length distributions of the 0-year fish.
190 Spawning occurs in shallow waters outside of the traditional range of the fishery (Antony Raja,
191 1964), and after spawning the adults migrate closer to the coast and the spent fish become ex-
192 posed to the fishery.

193 After eggs are spawned, they develop rapidly into larvae (Nair, 1959). The phytoplankton
194 bloom that provide sardine larvae food is dependent upon nutrient influx from coastal up-
195 welling and runoff from rivers during the summer monsoon and early fall. The blooms start in
196 the south near the southern tip of India in June, increase in intensity and spread northward up
197 the coast (BR, 2010). Variation in the bloom initiation time and intensity leads to changes in
198 the food supply and to corresponding changes in the growth and survival of larvae and in the
199 later recruitment of 0-year sardines into the fishery (George et al., 2012). Oil sardines grow
200 rapidly during their first few months, and 0-year fish (40mm to 100mm) from early spawn-
201 ing appear in the catch in August and September in most years (Antony Raja, 1970; Nair et
202 al., 2016). As the phytoplankton bloom spreads northward, the oil sardines follow, and the
203 oil sardine fishery builds from south to north during the post-monsoon period. Oil sardines
204 remain inshore feeding throughout the winter months, until March to May when the inshore
205 waters warm considerably and sardines move off-shore to deeper waters (Chidambaram, 1950).
206 Catches of sardines are correspondingly low during this time for all size classes. The age at
207 first maturity occurs at approximately 150 mm size (Nair et al., 2016), which is reached within
208 one year. When the summer monsoon returns, the oil sardine cycle begins anew.

209 Catches along the Kerala coast are high throughout the year except during quarter 2, April-
210 June (Figure 2). The age-distribution caught by the fishery varies through the year. The fishery
211 is closed during June to mid-July during the monsoon and peak spawning, and when it resumes
212 in mid-July, it is first dominated by 1-2.5 year old mature fish (Antony Raja, 1969; Bensam,
213 1964; Nair et al., 2016). In August or September a spike of 0-year (40mm) juveniles from
214 the June spawning typically appears in the catch (Antony Raja, 1969; Nair et al., 2016) and

215 another spike of 0-year fish is sometimes seen in February from the late fall spawning (Prabhu
216 & Dhulkhed, 1967, 1970). From October through June, the catch is dominated by fish from
217 120mm-180mm (Antony Raja, 1970; Nair et al., 2016; Prabhu & Dhulkhed, 1970) which is a
218 mix of 0-year, 1-year and 2-year fish (Nair et al., 2016; Rohit et al., 2018).

219 Materials and Methods

220 Sardine landing data

221 Quarterly fish landing data have been collected by the Central Marine Fisheries Research In-
222 stitute (CMFRI) in Kochi, India, since the early 1950s using a stratified multi-stage sample
223 design (Srinath et al., 2005). The survey visits the fish landing centers along the entire south-
224 east coast of India and samples the catch from the variety of boat types and gear types used
225 in the coastal fishery. Landings estimates are available for all the coastal states, however we
226 model the catch for the state of Kerala only, where the longest time series is available and the
227 overwhelming majority of oil sardines are landed (Figure 2). The quarterly landings (metric
228 tons) for oil sardine landed from all gears in Kerala were obtained from CMFRI reports (1956-
229 1984) and online databases (1985-2015) (CMFRI, 1969, 1995, 2016; Jacob et al., 1987; Pillai,
230 1982). The quarterly landing data were log-transformed to stabilize the variance.

231 Remote sensing data

232 We analysed monthly composites of the following environmental data derived from satellite
233 data: sea surface temperature (SST), chlorophyll-a (CHL), upwelling (UPW), the Oceanic
234 Niño Index (ONI), the Dipole Mode Index (DMI) and precipitation. The monthly time series
235 and means of the covariates are shown in Figure 4.

236 For sea surface temperature, we used Advanced Very-High Resolution Radiometer
237 (AVHRR) data, which provides accurate nearshore SST values. Although the ICOADS
238 product provides SST values for earlier years, ICOADS does not provide accurate nearshore
239 temperatures. For 1981 to 2003, we used the Pathfinder Version 5.2 product on a 0.0417
240 degree grid. These data were developed by the Group for High Resolution Sea Surface
241 Temperature (GHRSSST) and served by the US National Oceanographic Data Center. For
242 2004 to 2016, we used the NOAA CoastWatch SST products derived from NOAA's Polar

243 Operational Environmental Satellites (POES).

244 For chlorophyll-a, we used the chlorophyll-a products developed by the Ocean Biology
245 Processing Group in the Ocean Ecology Laboratory at the NASA Goddard Space Flight Cen-
246 ter. For 1997 to 2002, we used the chlorophyll-a 2014.0 Reprocessing (R2014.0) product from
247 the Sea-viewing Wide Field-of-view Sensor (SeaWiFS) on the Orbview-2 satellite. These data
248 are on a 0.1 degree grid. For 2003 to 2017, we used the MODIS-Aqua product on a 0.05 de-
249 gree grid. These CHL data are taken from measurements by the Moderate Resolution Imaging
250 Spectroradiometer (MODIS) on NASA's Aqua Spacecraft. The SST and CHL data were aver-
251 aged across thirteen 1 degree by 1 degree boxes which roughly parallel the bathymetry (Figure
252 1). The SST and CHL satellite data were retrieved from the NOAA ERDDAP server (Simons,
253 2017).

254 For an index of coastal upwelling, we used the sea-surface temperature differential be-
255 tween near shore and 3 degrees offshore based on the index described by Naidu et al. (1999)
256 and BR et al. (2008). For SST, we used the remote sensing sea-surface temperature data sets
257 described above. This SST-based upwelling index has been validated as a more reliable met-
258 ric of upwelling off the coast of Kerala compared to wind-based upwelling indices (BR et
259 al., 2008). The SST-based upwelling index and chlorophyll-a blooms are strongly correlated
260 (Figure 3).

261 Precipitation data were obtained from two different sources. The first was an estimate
262 of the monthly precipitation (in mm) over Kerala from land-based rain gauges (Kothawale &
263 Rajeevan, 2017); these data are available from the Indian Institute of Tropical Meteorology
264 and the data are available from the start of our landing data (1956). The second was a remote
265 sensing precipitation product from the NOAA Global Precipitation Climatology Project (Adler
266 et al., 2016). This provides estimates of precipitation over the ocean using a global 2.5 degree
267 grid. We used the 2.5 by 2.5 degree box defined by latitude 8.75 to 11.25 and longitude 73.25
268 to 75.75 for the precipitation off the coast of Kerala. These data are available from 1979
269 forward (NCEI, 2017). The land and nearshore ocean precipitation data are highly correlated
270 (Appendix E), supporting the use of the land time series as a proxy for the precipitation over
271 the ocean off the Kerala coast.

272 The Oceanic Niño Index (ONI) is a measure of the SST anomaly in the east-central Pacific
273 and is a standard index of the El Niño/Southern Oscillation (ENSO) cycle. The ONI index is
274 3-month running mean of ERSST.v5 SST anomalies in the Niño 3.4 region, based on centered
275 30-year base periods updated every 5 years. The ONI was downloaded from the NOAA Na-

276 tional Weather Service Climate Prediction Center. The Dipole Mode Index (DMI) is defined
277 by the SSTA difference between the western Indian Ocean (10°S – 10°N , 50°E – 70°E) and the
278 southeastern Indian Ocean (10°S – 0° , 90°E – 110°E). The DMI has been shown to be predict hy-
279 poxic events in our study area (Vallivattathillam et al., 2017). The DMI data were downloaded
280 from the NOAA Earth System Research Laboratory.

281 Hypotheses

282 Our statistical tests were structured around specific hypotheses (Table 1) concerning which re-
283 mote sensing covariates in which months should correlate with landings in specific quarters.
284 These hypotheses were based on biological information concerning how environmental con-
285 ditions affect sardine survival and recruitment and affect exposure of Indian oil sardines to
286 the coastal fishery. The Jul-Sep catch overlaps the summer monsoon and the main spawning
287 months. This is also the quarter where small 0-year fish from early spawning often appear in
288 the catch, sometimes in large numbers. Variables that affect or are correlated with movement of
289 sardines inshore should be correlated with quarter 3 landings. In addition, pre-spawning (Mar-
290 May) environmental conditions should be correlated with the spawning strength as adult oil
291 sardines experience an acceleration of growth during this period along with egg development.
292 The post-monsoon catch (Oct-May) is a mix of 0-year fish (less than 12 months old) and ma-
293 ture fish (greater than 12 months old). Variables that are correlated with spawning strength and
294 larval and juvenile survival should correlate with the post-monsoon catch both in the current
295 year and in future years, one to two years after.

296 Our hypotheses (Table 1) focus mainly on two drivers: upwelling and ocean temperature.
297 We also test hypotheses concerning precipitation as this has historically been an environmental
298 covariate considered to influence the timing of oil sardine landings. More recently, researchers
299 have highlighted the influence of large-scale ocean processes, specifically the El Niño/Southern
300 Oscillation, on sardine fluctuations; therefore we test the Ocean Niño Index (ONI) and Dipole
301 Mode Index (DMI) also. Chlorophyll density is directly correlated with sardine food availabil-
302 ity and chlorophyll fronts are known to influence sardine shoaling. However our chlorophyll
303 time series is short (1997-2015) and the statistical power for testing correlation with landings
304 is low. Tests of chlorophyll are shown in the appendices but are not the focus of our analyses.

305 **Statistical models**

306 We modeled the catches during the late-monsoon season (Jul-Sep) separately from the post-
307 monsoon season (Oct-Mar). Thus there is no seasonality in our catch time series, as we ana-
308 lyzed a yearly time series of Jul-Sep catches separately from a yearly time series of Oct-Mar
309 catches. We divided the catch in this way for biological and statistical reasons. Catch in Jul-
310 Sep captures a mix of spawning age fish as it overlaps with the tail end of the spawning season,
311 is affected by a fishery closure during the summer monsoon, and is periodically inflated by the
312 appearance of small 0-year fish from early summer spawning. In addition, the covariates that
313 affect the timing of spawning, movement of post-spawning mature fish inshore, and early egg
314 and larval survival may be different than those that affect later growth, survival and shoaling
315 that exposes fish to the inshore fishery. Analyzing catch and covariate time series without sea-
316 sonality also had an important statistical benefit—we removed the problem of seasonality in
317 the catch and covariates. The oil sardine life-cycle is seasonal and driven by the strong season-
318 ality in this monsoon influenced system. A simple statistical model with quarters will explain
319 much of the quarterly catch data since most of the yearly variability is due to seasonality and
320 any environmental covariate with a similar seasonality will also show high correlation with the
321 landings. Our goal was to explain year-to-year variability thus eliminating the confounding
322 effect of seasonality in the data was important.

323 Preliminarily, we tested ARIMA models on both the Jul-Sep and Oct-Mar catch time
324 series and found little support for auto-regressive errors (ARIMA models with a MA com-
325 ponent) based on diagnostic tests of the residuals and model selection. The best supported
326 ARIMA models were simple AR models ($x_t = bx_{t-1} + \varepsilon_t$). This lack of strong auto-correlation
327 in residuals has been found in other studies that tested ARIMA models for forecasting small
328 pelagic catch (Stergiou & Christou, 1996). We thus used AR-only models, however we tested
329 both linear and non-linear models using generalized additive models (GAMs, Wood, 2017) of
330 the form $x_t = s(x_{t-1}) + \varepsilon_t$ and tested time-varying linear models with dynamic linear models
331 (DLM). GAMs allow one to model the effect of a covariate as a flexible non-linear function
332 while DLMs allow the effect of the covariate to vary over time. Our GAM approach is analo-
333 gous to that taken by Jacobson and MacCall (1995) in a study of the effects of SST on Pacific
334 sardine recruitment.

335 The first step in our analysis was to determine the catch model: the model for current
336 catch as a function of the past catch. We explored four classes of models: null models with a
337 simple function of prior catch, linear regressive models with one to two years of prior catch,

338 dynamic linear models (DLM) which allow the regression parameters to vary (Holmes et al.,
 339 2012, using the MARSS package in R), and GAMs to allow the effect of prior catch to be
 340 non-linear. One feature of GAMs is that they allow the smoothing parameter of the response
 341 curve to be estimated. We fixed the smoothing parameter at an intermediate value so that
 342 smooth responses were achieved. Multi-modal or overly flexible response curves would not
 343 be realistic for our application. We fit GAMs with smooth terms represented by penalized
 344 regression splines (Wood, 2011, using the mgcv package in R) and fixed the smoothing term
 345 at an intermediate value (sp=0.6).

346 We compared the following catch models:

- 347 • null: $\ln(C_{i,t}) = \ln(C_{j,t-1})$
- 348 • random walk: $\ln(C_{i,t}) = \alpha + \ln(C_{j,t-1}) + \varepsilon_t$
- 349 • linear AR-1: $\ln(C_{i,t}) = \alpha + \beta \ln(C_{j,t-1}) + \varepsilon_t$
- 350 • linear AR-2: $\ln(C_{i,t}) = \alpha + \beta_1 \ln(C_{j,t-1}) + \beta_2 \ln(C_{k,t-2}) + \varepsilon_t$
- 351 • DLM AR-1: $\ln(C_{i,t}) = \alpha_t + \beta_t \ln(C_{j,t-1}) + \varepsilon_t$
- 352 • GAM AR-1: $\ln(C_{i,t}) = \alpha + s(\ln(C_{j,t-1})) + \varepsilon_t$
- 353 • GAM AR-2: $\ln(C_{i,t}) = \alpha + s_1(\ln(C_{j,t-1})) + s_2(\ln(C_{k,t-2})) + \varepsilon_t$

354 $\ln(C_{i,t})$ is the log catch in the current year t in season i . We modeled two different catches: S_t
 355 (Jul-Sep) and N_t (Oct-Jun). The catches were logged to stabilize and normalize the variance.
 356 $s()$ is a non-linear function estimated by the GAM algorithm. The model is primarily statistical,
 357 meaning it should not be thought of as a population growth model. We tested models with prior
 358 year and two years prior Oct-Mar catch (N_{t-1} and N_{t-2}) and Jul-Sep catch (S_{t-1} and S_{t-2}) as
 359 the explanatory catch variable. S_t was not used as a predictor for N_t because S_t is the quarter
 360 immediately prior to N_t and would not be available for a forecast model since time is required
 361 to process landings data. The catch models were fit to 1984 to 2015 catch data, corresponding
 362 to the years where the SST, upwelling and precipitation data were available. F-tests, AIC and
 363 leave-one-out cross-validation (LOOCV) on nested sets of models (Wood et al., 2016) were
 364 used to evaluate the support for the catch models and later for the covariate models. LOOCV
 365 involves leaving out a data point, fitting the model, and then predicting the left-out data point.
 366 After selection of the best model with the 1984-2015 data, the fitting was repeated with the
 367 1956-1983 catch data to confirm the form of the catch models. An influential years test was
 368 done by removing each year and repeating the model selection analysis.

369 Once the catch models were determined, the covariates were studied individually and then
 370 jointly. As with the catch models, F-tests, AIC and LOOCV (leave-one-out cross-validation)

371 on nested sets of models were used to evaluate the support for models with covariates. The
372 smoothing term was fixed at an intermediate value ($sp=0.6$) instead of being treated as an
373 estimated variable. Our models for catch with covariates took the form $\ln(C_{i,t}) = M + \alpha +$
374 $s_1(V_{1,t}) + s_2(V_{2,t-1}) + \varepsilon_t$, $\ln(C_{i,t}) = M + \alpha + \beta_1 V_{1,t} + \beta_2 V_{2,t-1} + \varepsilon_t$, and $\ln(C_{i,t}) = M + \alpha +$
375 $\beta_t V_{1,t} + \varepsilon_t$ where M was the best catch model from step 1 and V is a covariate. Thus models
376 with covariates modeled as a linear, non-linear and time-varying effect were compared. The
377 covariates tested are those hypothesized to drive variability in oil sardine landings (Table 1).
378 We tested both models with one and two covariates and did not use correlated covariates in the
379 same model.

380 Results

381 Catches in prior seasons as explanatory variables

382 Using the 1984-2015 catch data, which is the time-period that overlaps our available environ-
383 mental data, the Jul-Sep catch models were compared against a “naive” model in which the
384 forecasted Jul-Sep catch was simply the Jul-Sep catch in the prior year. The “naive” model
385 has no estimated parameters and is a standard null model for time series modeling. Models
386 with $\ln(N_{t-1})$ (Oct-Mar catch in prior year), whether linear or non-linear, as the explanatory
387 covariate were strongly supported over the naive model and over models with $\ln(S_{t-1})$ (Jul-
388 Sep catch in prior year) as the explanatory variable (Tables A1 and A2). Addition of the catch
389 two years prior, $\ln(N_{t-2})$ or $\ln(S_{t-2})$, did not reduce AIC or LOOCV for either the linear or
390 non-linear models. We tested the support for non-linearity in the effect of the prior year catch
391 by comparing models with $\ln(N_{t-1})$ or $\ln(S_{t-1})$ included as a linear term or as a non-linear
392 function $s()$ using GAMs (Table A2). The residual error decreased using a non-linear response
393 at the cost increased degrees of freedom. The result was only weak (non-significant) support
394 for allowing a non-linear response based on AIC and LOOCV.

395 The results on model structure were similar for models of the Oct-Mar landings (N_t),
396 but the models explained much more of the variance (with a maximum adjusted $R^2 = 56.6$).
397 The most supported model for N_t (Tables A3 and A4) based on AIC and F-tests used a non-
398 linear response to Oct-Mar catch of the previous season $\ln(N_{t-1})$ plus a non-linear response
399 to Jul-Sep catch two years prior $\ln(S_{t-2})$. However the simpler model with only $\ln(N_{t-1})$ had
400 the lowest LOOCV (out of sample prediction accuracy). Thus this simpler model was also
401 included as one of the base models for the Oct-Mar catch. Models with Jul-Sep catch in the

402 current fishing season were not used as these data would not be available by Oct of the current
403 season (for forecast purposes).

404 As diagnostic checks, we did the same model comparison for the landings data set from
405 1956 to 1983. The results were the same for the Jul-Sep catch (Table A5) with the model with
406 $\ln(N_{t-1})$ included as a non-linear covariate with the lowest AIC and LOOCV. For the Oct-Mar
407 catch (Table A6), the results were very similar but not identical. The model with $\ln(N_{t-1})$
408 included as a non-linear covariate had the lowest LOOCV while the models with $\ln(N_{t-1})$ and
409 $\ln(S_{t-2})$ or $\ln(S_{t-1})$ had the lowest AIC (though less than 1 from the AIC of the $\ln(N_{t-1})$
410 model). We also did an influential years test using Leave-One-Out cross-validation (Appendix
411 G). This test involved leaving out one year and repeating the model selection tests. These tests
412 also supported the selected base models. The dynamic linear models (allowing a time-varying
413 effect of prior catch) performed poorly for the Jul-Sep catch with high AIC and LOOCV. For
414 the Oct-Mar catch, the performance was mixed with higher AIC but lower LOOCV.

415 Based on the model selection tests, the following non-linear model was chosen as the base
416 model for Jul-Sep (S_t) catch:

$$M0 : \ln(S_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t.$$

417 Two base non-linear models were chosen for Oct-Mar (N_t) catch:

$$M1 : \ln(N_t) = \alpha + s_1(\ln(N_{t-1})) + s_2(\ln(S_{t-2})) + \varepsilon_t$$

418

$$M2 : \ln(N_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$$

419 Note that although M1 was the best model for Jul-Sep catch, it was only weakly explanatory.
420 The maximum adjusted R^2 was less than 30% (Table A2). For the Oct-Mar catch, M2 and M3
421 were more explanatory with an adjusted R^2 of 45.3% for M2 and 56.6% for M3 (Table A4).

422 Environmental covariates as explanatory variables

423 There was no support for using precipitation during the summer monsoon (June-July) or pre-
424 monsoon period (April-May) as an explanatory variable for the quarter 3 (July-September) or
425 post-monsoon (October-March) catch (hypotheses S1 and S2; Tables B1 and B2). This was
426 the case whether precipitation in the current or previous season was used, if precipitation was
427 included as non-linear or non-linear effect, and if either precipitation during monsoon (June-

428 July) or pre-monsoon (April-May) were used as the covariate. Quarter 3 overlaps with the
429 spawning period and precipitation is often thought to trigger spawning, however we were un-
430 able to find any consistent association of catch during these spawning and early-post spawning
431 months with precipitation. Raja (1974) posited that the appropriate time period for the affect
432 of rainfall is the weeks before and after the new moon when spawning is postulated to occur
433 and not the total rainfall during the monsoon season. Thus the lack of correlation may be due
434 to using too coarse of a time average for the precipitation.

435 The sea-surface temperature before spawning (March-May) has been speculated to be cor-
436 related with successful egg development and spawning behavior (hypothesis S4 and S5) and
437 extreme heat events pre-spawning have been associated with low recruitment. This suggests
438 that March-May in the current and prior years should be associated with low catch. The sea-
439 surface temperature during larval and early juvenile development (October-December) may
440 affect survival and growth in multiple ways and thus could correlate with biomass in future
441 years (hypothesis L1). However we found no support for either of these SST variates as ex-
442 planatory variables for the July-September catch and only weak support (based on AIC) for
443 March-May SST in the current season for explaining variability in post-monsoon catch. The
444 fall average SST in the prior season did not explain variability in either July-September or
445 October-March catch. See Tables B3 and B4.

446 We also found no correlation between the ONI index (hypothesis A2) for either July-
447 September or post-monsoon catch (Tables B1 and B2).

448 Instead we found with the covariates indirectly and directly associated with productiv-
449 ity and food availability: upwelling intensity and surface chlorophyll. The correlation between
450 landings and upwelling was only found for upwelling in the current season. No correlation was
451 found when we used the upwelling index from the prior season. The correlation between land-
452 ings and upwelling was found for both July-September and October-March landings and with
453 either upwelling index: average nearshore SST along the Kerala coast during June-September
454 or the average SST nearshore versus offshore differential (UPW) off Kochi in June-September
455 (Table 2, Table B3 and Table B4). These two upwelling indices are correlated but not identical.
456 The model with average June-September nearshore SST was more supported than the model
457 using the SST differential off Kochi. For July-September catch, this model with a non-linear
458 response had an adjusted R^2 of 41.0 versus an adjusted R^2 of 24.4 for the model with no co-
459 variates (Table B3), and for October-March catch, the adjusted R^2 was 61.8 versus 56.6 (Table
460 B4). Note, that this covariate is June-September in the current season and overlaps with the
461 July-September catch. Thus this model cannot be used to forecast July-September catch but

462 does help us understand what factors may be influencing catch during the monsoon.

463 Chlorophyll-a density is speculated to be an important predictor of larval sardine sur-
464 vival and growth. In addition, sardines shoal in response to coastal chlorophyll blooms, which
465 brings them in contact with the coastal fisheries. Thus chlorophyll-a density is assumed to be
466 an important driver of future or current sardine catches. We had chlorophyll-a remote sensing
467 data only from 1998 onward. Our simplest covariate model required 5 degrees of freedom,
468 thus we were limited in the analyses we could conduct. In addition, the years, 1998-2014,
469 have relatively low variability in catch sizes; the logged catch sizes during this period range
470 from 10-11 during quarter 3 and 11-12 during the other three quarters. Second degree polyno-
471 mial models were fit (Appendix C) to the average log chlorophyll-a density in the current and
472 prior season from quarter 3 (July-September), 4 (October-December), and 1 (January-March).
473 Chlorophyll-a density was not a significant predictor for the July-September catch for any of
474 the tested combinations of current or prior season and quarter. The only significant effect was
475 seen for post-summer monsoon catches using chlorophyll-a density in October-December of
476 the prior season (Table C1). This is in contrast to the results with monsoon upwelling indices,
477 which found a correlation with the current season but not prior seasons.

478 The strongest correlation however was found with the multi-year average sea surface tem-
479 perature for the nearshore waters off Kerala (latitude 8 to 11). The average sea surface tem-
480 perature over multiple prior years has been found to be correlated with sardine recruitment in
481 Pacific sardines (Checkley et al., 2017; Jacobson & MacCall, 1995; Lindegren et al., 2013)
482 and southern African sardines (Boyer et al., 2001). We tested as a model covariate the average
483 SST for 2.5 years prior to the July-September catch, so January-June in the current calendar
484 year and the two prior calendar years for a 30-month average. This covariate can be used
485 for forecasting since it does not overlap with either July-September or October-March catch.
486 This variate with a non-linear response was best covariate for both the July-September and the
487 post-monsoon catch. For post-monsoon catch, the model with SST had an adjusted R^2 of 67.5
488 versus 56.6 without. For the July-September catch, the adjusted R^2 was 41.0 with SST and 24.4
489 without. The response curve was step-like with a negative effect at low temperatures and then
490 an positive flat effect at higher temperatures (Figure 5). This is similar to the step-response
491 found in studies of the correlation between average SST and recruitment in Pacific sardines
492 (Jacobson & MacCall, 1995).

493 **Discussion**

494 Sardines in all the world's ecosystems exhibit large fluctuations in abundance (Baumgartner et
495 al., 1992). These small forage fish are strongly influenced by natural variability in the ocean
496 environment. Upwelling, influenced by both large-scale forces such as regimes shifts and El
497 Niño/Southern Oscillation patterns (Alheit & Hagen, 1997; Schwartzlose et al., 2010) and by
498 seasonal wind and current patterns, brings nutrient rich and oxygen rich waters to the surface.
499 This drives the seasonal variability in phytoplankton resources and in turn sardine prey (Bakun
500 et al., 2008). Local variability in temperature, salinity, and oxygen levels have both direct and
501 indirect on sardine reproduction, recruitment and survival (Checkley et al., 2017). Sardines
502 are also influenced by competition and predation by other species and well-known for their
503 sensitivity to over-fishing which has been linked to many fishery collapses (Kripa et al., 2018).

504 Many studies on Pacific sardines have looked at the correlation between ocean temperature
505 (SST) and recruitment. Temperature can have direct effect on larval survival and growth and
506 an indirect effect on food availability. Studies in the California Current System, have found
507 that SST explains year-to-year variability in Pacific sardine recruitment (Checkley et al., 2009,
508 2017; Jacobson & MacCall, 1995; Lindegren & Checkley, 2012) and that the average nearshore
509 temperature over multiple seasons is the explanatory variable. Similar to these studies, we
510 found that the average nearshore SST over multiple seasons was the covariate that explained
511 the most variability in catch both in the monsoon and post-monsoon months. McClatchie et
512 al. (2010) found no SST relationship with SST and Pacific sardine recruitment, however their
513 analysis used a linear relationship while the other studies, and ours, that found a relationship
514 (Checkley et al., 2017; Jacobson & MacCall, 1995) allowed a non-linear relationship. Both
515 Jacobson and MacCall (1995) and Checkley et al. (2017) found a step-like response function
516 for temperature: below a threshold value the effect of temperature was linear and above the
517 threshold, the effect was flat and at lower temperatures the effect was negative and became
518 positive as temperature increased. Our analysis found a similar pattern with a negative effect
519 when the 2.5-year average temperature was below 28.35°C and positive above and with the
520 positive effect leveling off above 28.5°C (Figure 5).

521 There were four outlier years when catch were much lower than expected based on prior
522 catches: 1986, 1991, 1994 and 2013. The 2.5-year average SST predicted the collapses in
523 1986 and 1991 (Figure 6); the size of the residual with the covariate was much smaller than
524 without the covariate. The largest collapse was in 1994 and the most recent, in our dataset,
525 was 2014. The 2.5-year average SST did not predict the 1994 nor 2013 collapse. There was

526 no change in the size of the residual with and without the covariate. In fact, none of the
527 covariates we tested changed the size of the model residuals for 1994 nor 2013. The causes of
528 these unusual declines appear either unrelated to the environmental factors we studied. This
529 suggests either that other factors, biological or anthropogenic, drove these declines or that a
530 particular combination of environmental factors led to the declines. It should also be noted
531 that our upwelling indices captured on one aspect of upwelling: the nearshore intensity. Other
532 aspects of upwelling also affect oil sardines, such as the spatial extent both along the coast and
533 off the coast and the timing of the start of upwelling.

534 Seasonal productivity in the SE Arabian Sea upwelling is driven by the summer monsoon,
535 which causes strong coastal upwelling that moves from the south to the north over the summer.
536 This drives a strong seasonal pattern of zooplankton abundance (Figure 3). Despite the strong
537 connection between sardine recruitment, growth and survival with upwelling, we found no
538 correlation with upwelling in the prior season with landings. We did find a correlation between
539 upwelling in the current season with landings in the current season. The biological reasons
540 behind a positive relationship with upwelling are clear. Upwelling drives productivity and
541 higher food resources in the current season leads to higher recruitment and higher numbers
542 of 0-year fish in the landings or may bring sardines into the nearshore to feed where they are
543 exposed to the fishery. However, the explanatory power of the upwelling indices was mainly
544 due to the negative effect of extremely high upwelling (Figure 5). Extremely high upwelling
545 transports larval sardines offshore and can create regions of low oxygen which sardines avoid.

546 Conclusions

547 Remote sensing satellites can be used to detect changes in ocean physical, biological and chem-
548 ical properties, such as surface temperature, winds, surface height, surface waves, rainfall and
549 surface salinity, as well as the ecosystem and water quality changes. Unlike in-situ measure-
550 ments, environmental measures from remote-sensing can be acquired rapidly and over large
551 regions. However, which environmental covariates will improve forecasts is not obvious from
552 oil-sardine life-history alone. We tested using many of the covariates known or suspected to
553 have an effect on sardine spawning, growth and survival: precipitation, upwelling indices, ocean
554 temperature and chlorophyll-a in various critical months of the sardine life-cycle. We found
555 that the multi-year average nearshore ocean temperature explained the most variability in the
556 landings. This covariate is not as directly tied to stages of the oil-sardine life-cycle as the other
557 covariates we tested, though it does integrate over multiple influences (upwelling strength and

558 temperature) over multiple years.

559 The temperature of the Western Indian Ocean, of which the Southeast Arabian Sea is a
560 part, has been increasing over the last century at a rate higher than any other tropical ocean
561 (Roxy et al., 2014) and the warming has been most extreme during the summer monsoon
562 months. This ocean climate change is affecting oil sardine distributions with significant land-
563 ings now occurring north of Goa (Vivekanandan et al., 2009). Continued warming is expected
564 to affect the productivity of the region via multiple pathways, including both the direct effects
565 of temperature change on the physiology and behavior of organisms and a multiple of indirect
566 effects (Moustahfid et al., 2018). These indirect effects includes changes to salinity, oxygen
567 concentrations, currents, wind patterns, ocean stratification and upwelling spatial patterns, phe-
568 nology, and intensity. Incorporating environmental covariates into landings forecasts has the
569 potential to improve fisheries management for small pelagics such as oil sardines in the face of
570 a changing ocean environment (Haltuch et al., 2019; Tommasi et al., 2016). However, moni-
571 toring forecast performance and covariate performance in models will be crucial as a changing
572 ocean environment may also change the association between landings and average sea surface
573 temperature.

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846 **Figure Legends**

847 Figure 1. Close up of Kerala State with the latitude/longitude boxes used for the satellite data.
848 Kerala State is marked in grey and the oil sardine catch from this region is being modeled.

849 Figure 2. Quarterly catch data 1956-2014 from Kerala. The catches have a strong seasonal
850 pattern with the highest catches in quarter 4 Note that quarter 3 is July-September and that the
851 fishery is closed July 1 to August 15, thus the fishery is only open 1.5 months in quarter 3.
852 The mean catch (metric tonnes) in quarters 1 to 4 are 38, 19.2, 30.9, and 59.9 metric tonnes
853 respectively.

854 Figure 3. Key oil sardine life-history events overlaid on the monthly sea surface tempera-
855 ture in the nearshore and offshore and the nearshore chlorophyll density.

856 Figure 4. Remote sensing covariates used in the analysis. All data are monthly averages
857 over Box 4 in Figure 1 on the Kerala coast off of Kochi. Panel A) Upwelling Index. The
858 upwelling index is the difference between the near-shore sea surface temperature (SST) and
859 the off-shore SST defined as 3 degrees longitude offshore. Panel B) Surface chlorophyll-a
860 (CHL). The CHL data are only available from 1997 onward. Panel C) Sea surface temperature
861 constructed from Advanced Very High Resolution Radiometer (AVHRR). Panel D) Average
862 daily rainfall (mm/day) off the Kerala coast.

863 Figure 5. Effects of covariates estimated from the GAM models. Panel A) Effect of the
864 2.5 year average nearshore SST on catch during the catch during July-September (late spaw-
865 ning and early post-spawning) months. Panel B) Effect of upwelling (inshore/off-shore SST
866 differential) during June-September in the current season on July-September catch. The index
867 is the difference between offshore and inshore SST, thus a negative value indicates warmer
868 coastal surface water than off-shore. Panel C) Effect of the 2.5 year average nearshore SST on
869 catch during the catch during October-March (post-monsoon, age-0, -1, -2 year fish). Panel
870 D) Effect of upwelling (inshore/off-shore SST differential) during June-September in the cur-
871 rent season on October-March catch. Strong upwelling (positive upwelling index) in the larval
872 and juvenile high growth period (Oct-Dec) is associated with higher early survival and larger
873 cohorts of age-0 fish in the catch.

874 Figure 6. Fitted versus observed catch with models with and without environmental co-
875 variates. Panel A) Fitted versus observed log catch in the spawning months with only non-
876 spawning catch in the previous season as the covariate: $S_t = s(N_{t-1}) + \varepsilon_t$. Panel B) Fitted
877 versus observed log catch in July-September with the 2.5-year average nearshore SST added

878 as a covariate to the model in panel A. This model was: $S_t = s(N_{t-1}) + s(V_t) + \varepsilon_t$. Panel
879 C) Fitted versus observed log catch in the post-monsoon months with only post-monsoon
880 catch in the previous season and July-September catch two seasons prior as the covariates:
881 $N_t = s(N_{t-1}) + s(S_{t-2}) + \varepsilon_t$. Panel D) Fitted versus observed log catch in the post-monsoon
882 months with 2.5-year average nearshore SST (V) added as covariates. This model was $N_t =$
883 $s(N_{t-1}) + s(S_{t-2}) + s(V_t) + \varepsilon_t$.

Table 1. Hypotheses for covariates affecting landings. S_t is Jul-Sep catch in the current season, S_{t-1} is Jul-Sep catch in the previous season. N_t is the Oct-Mar catch in the current season and N_{t-1} is the Oct-Mar catch in the prior season. DD = hypotheses related to effects of past abundance (landings) on current abundance. S = hypotheses related to catch during the spawning months. L = hypotheses related to larval and juvenile growth and survival. A = hypotheses affecting all ages. See the introduction for discussion of the literature on the effects of environmental covariates on sardine landings. ns-SST = nearshore SST. r-SST = regional (nearshore and offshore) SST t , $t - 1$, and $t - 2$ indicate current, prior, and two seasons prior.

Hypothesis	Description
DD1 $S_t \sim N_{t-1}$	S_t is age 1+ age fish and reflects the 0-2yr fish in N_{t-1} which have aged 3-6 months (Nair et al. 2016).
DD2 $S_t \sim S_{t-1} + S_{t-2} N_t S_{t-1} + S_{t-2}$	S_t is age 1+ fish and is dominated by spent post-spawning fish (Nair et al. 2016). S_t should be correlated with the abundance of the spawning stock and cohort strength. If cohort strength persists over time, landings in the current season should correlate with Jul-Sep landings in the previous two seasons.
DD3 $N_t \sim N_{t-1}$	N_{t-1} includes age 0, 1, 2 and 2+ fish (Nair et al. 2016). Thus the 0, 1, and 2 yr fish will appear 1 year later in the N_t landings.
S1 $S_t \sim$ Jun-Jul precipitation in t	The magnitude of precipitation in June-July directly or indirectly signals mature fish to spawn offshore, after which the spent adults migrate inshore to the fishery (Anthony Raja 1969, 1974; Srinath 1998).
S2 $S_t \sim$ Apr-Mar precipitation in t	If there is pre-monsoon rain during Apr-May then spawning begins earlier and spent adults return to the fishery sooner and are exposed to the fishery longer.
S3 $N_t \sim$ Apr-Mar precipitation in t	Precipitation is an indicator of climatic conditions during the period of egg development. This affect the success of spawning, thus affecting the 0-year class and the Oct-Mar catch (Antony Raja 1969, 1974; Srinath 1998).
S4 $S_t \sim$ Jun-Sep UPW in t	High rates of upwelling drive mature fish further offshore and thus leads to lower exposure to the fishery. This movement may be driven by advection of high phytoplankton biomass further offshore or increased inputs of low oxygenated water to the surface (Murty and Edelman 1970, Antony Raja 1973, Pillai et al. 1991).
S5 $S_t \sim$ Mar-May r-SST in t $N_t \sim$ Mar-May r-SST in t	Extreme heating events prior to the spring monsoon drives mature fish from the spawning areas, resulting in poor recruitment and fewer age 0 year fish in the Oct-Mar catch (Antony Raja 1973, Pillai et al 1991).

Table 1. Continued.

Hypothesis	Description
L1 $S_t \sim \text{Oct-Dec ns-SST } t - 1$ $N_t \sim \text{Oct-Dec ns-SST } t - 1$	Larval and juvenile growth and survival is affected by temperature. The temperature in the critical post-monsoon window, when somatic growth of the 0-year fish is highest, should be correlated with future abundance.
L2 $S_t \sim \text{Jun-Sep UPW in } t - 1 \text{ & } t$ $N_t \sim \text{Jun-Sep UPW in } t - 1 \text{ & } t$	Higher upwelling rates leads to greater phytoplankton productivity, better larval and juvenile growth, and higher landings in the post monsoon Oct-March catch. However, extreme upwelling may decrease this catch, and future catches, due to poor oxygen conditions or advection of phytoplankton biomass.
L3 $S_t \sim \text{CHL in } t - 1 \text{ & } t$ $N_t \sim \text{CHL in } t - 1 \text{ & } t$	Surface Chl-a is a proxy for phytoplankton abundance and thus food availability, so higher bloom intensities may support greater fish abundance in the current and future years.
A1 $S_t \sim \text{2.5-yr ave. ns-SST}$ $N_t \sim \text{2.5-yr ave. ns-SST}$	Spawning, early survival, and recruitment of sardines depends on a multitude of cascading factors summarized by the multi-year average nearshore SST. The multi-year average SST has been correlated with recruitment and abundance in other sardines (Checkley et al. 2017) and is related to the optimal temperature windows for sardines (Takasuka et al. 2007).
A2 $S_t \sim \text{ONI in } t - 1$ $N_t \sim \text{ONI in } t - 1$	The El Nino/Southern Oscillation (ENSO) impacts a range of environmental parameters (e.g., precipitation, SST, frontal zones, winds, etc.) that may impact spawning and early survival of sardines (Supraba et al. 2016, Rohit et al. 2018).
A3 $S_t \sim \text{DMI in } t - 1$ $N_t \sim \text{DMI in } t - 1 \text{ & } t$	Negative Dipole Mode Index (DMI) values in Sep-Nov are associated with anoxic events on the Kerala coast while positive values are associated with a absence of such events (Vallivattathillam et al. 2017).

Table 2. Top covariates for the monsoon (Jul-Sep) and post-monsoon (Oct-Mar) catch (S_t and N_t) models. M0, M1 and M2 are the base models with only prior catch as covariates. To the base models, the listed environmental covariates are added. The full set of covariate models and the results of the tests on the nested models sets are given in Appendix B. The fitted versus observed catches from the covariate models are shown in Figure 6.

Model	Residual df	Adj. R2	RMSE	AIC	LOOCV RMSE
Jul-Sep catch models with covariates					
V_t = Jun-Sep SST current season					
W_t = Jun-Sep Bakun-UPW current season					
Z_t = 2.5-year average SST					
M0: $\ln(S_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$	28.6	24	1.184	109.52	1.299
$\ln(S_t) = M0 + s(V_t)$	25.9	41	1.007	103.43	1.192
$\ln(S_t) = M0 + \beta W_t$	27.6	28	1.133	108.66	1.404
$\Rightarrow \ln(S_t) = M1 + s(Z_t)$	26.2	41	1.011	103.26	1.338
Oct-Mar catch models with covariates					
V_t = Mar-May SST current season					
W_t = Jun-Sep SST current season					
Z_t = 2.5-year average SST					
X_t = fall DMI prior season					
M1: $\ln(N_t) = \alpha + s(\ln(N_{t-1})) + s(\ln(S_{t-2})) + \varepsilon_t$	24.8	57	0.713	79.53	1.062
$\ln(N_t) = M1 + s(V_t)$	22	63	0.628	76.01	1.002
$\ln(N_t) = M1 + \beta W_t$	23.8	63	0.648	75.57	1.042
$\Rightarrow \ln(N_t) = M1 + s(Z_t)$	22.7	67	0.597	71.88	0.827
$\ln(N_t) = M1 + s(X_t)$	21.1	68	0.58	72.69	0.89
M2: $\ln(N_t) = \alpha + s(\ln(N_{t-1})) + \varepsilon_t$					
$\ln(N_t) = M2 + s(V_t)$	27.6	45	0.836	84.75	0.966
$\ln(N_t) = M2 + \beta W_t$	24.8	47	0.791	85.9	0.981
$\Rightarrow \ln(N_t) = M2 + s(Z_t)$	26.6	52	0.772	81.79	0.927
$\ln(N_t) = M2 + s(X_t)$	25.3	60	0.688	76.34	0.796
	23.7	43	0.8	88.43	0.969

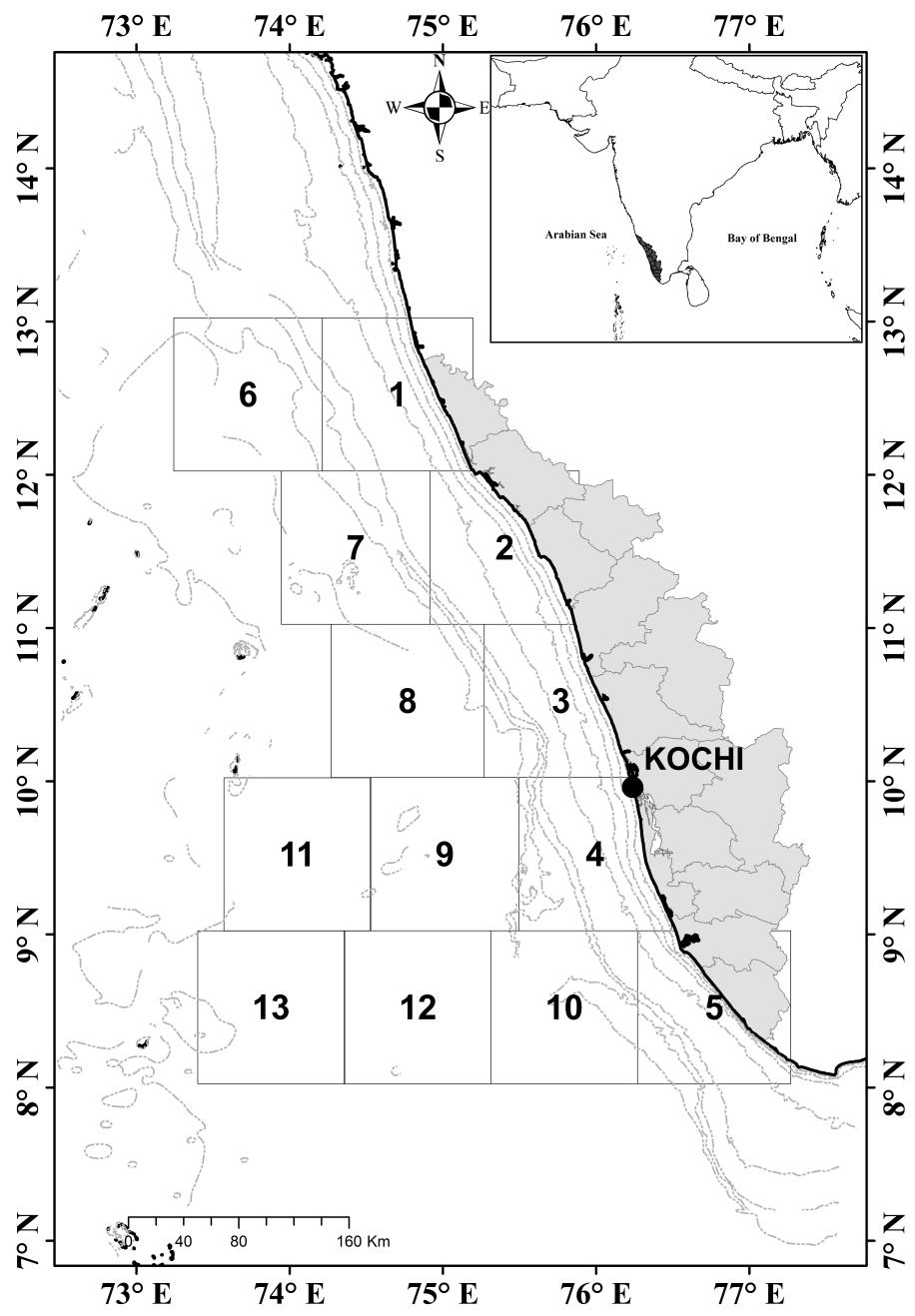


Figure 1

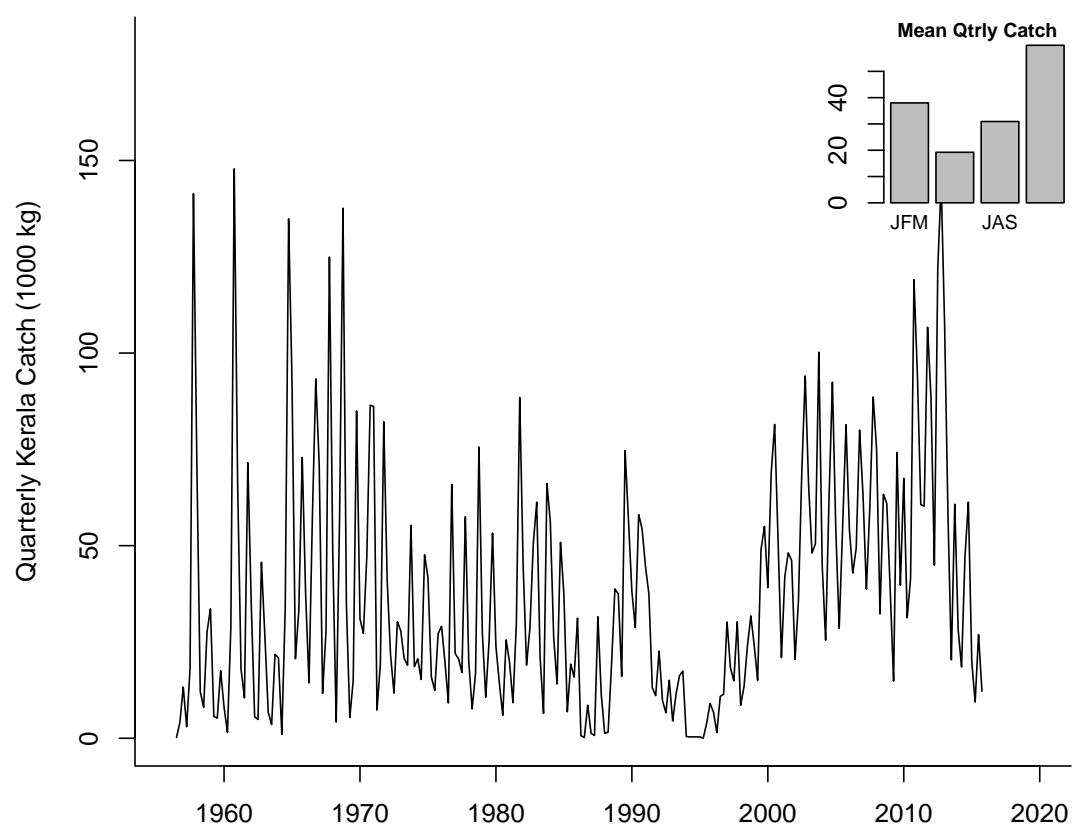


Figure 2

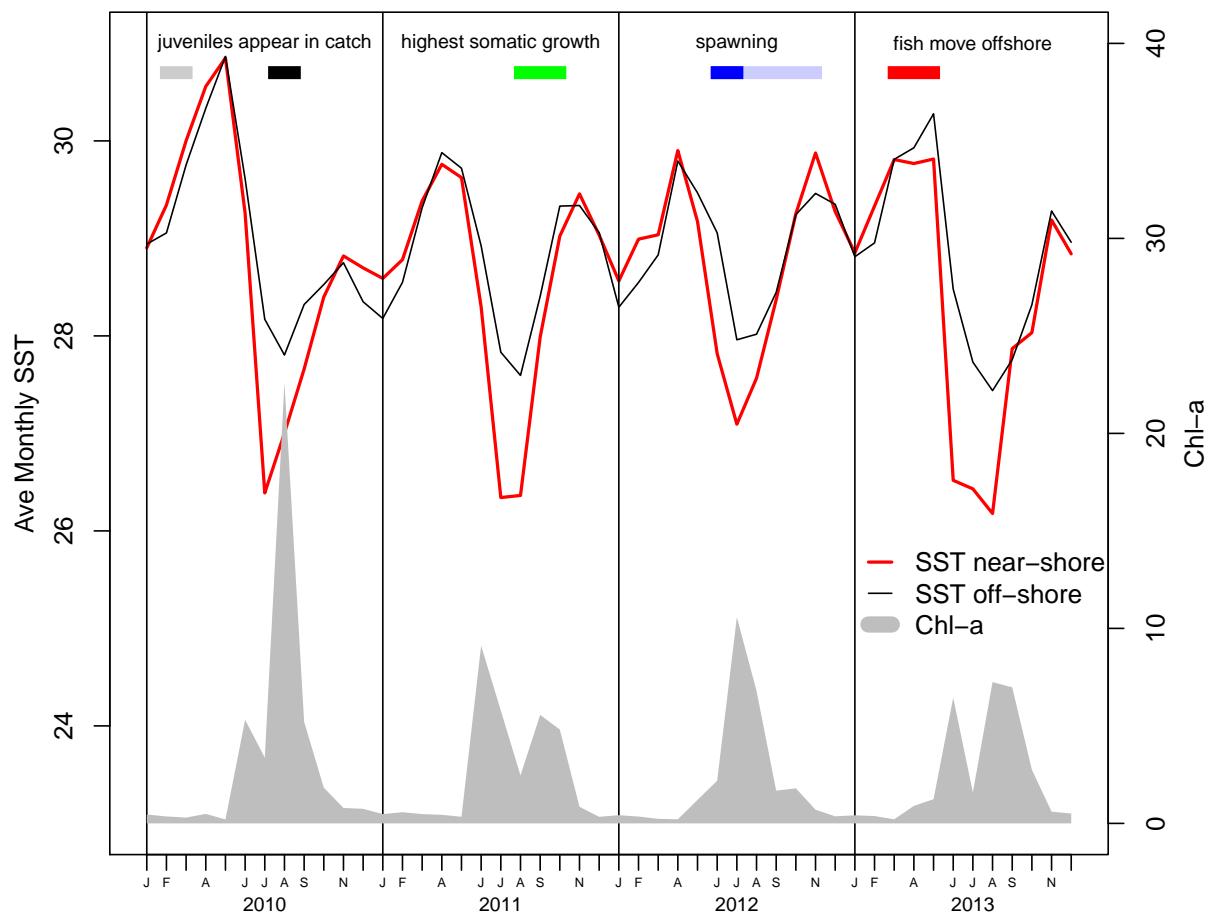


Figure 3

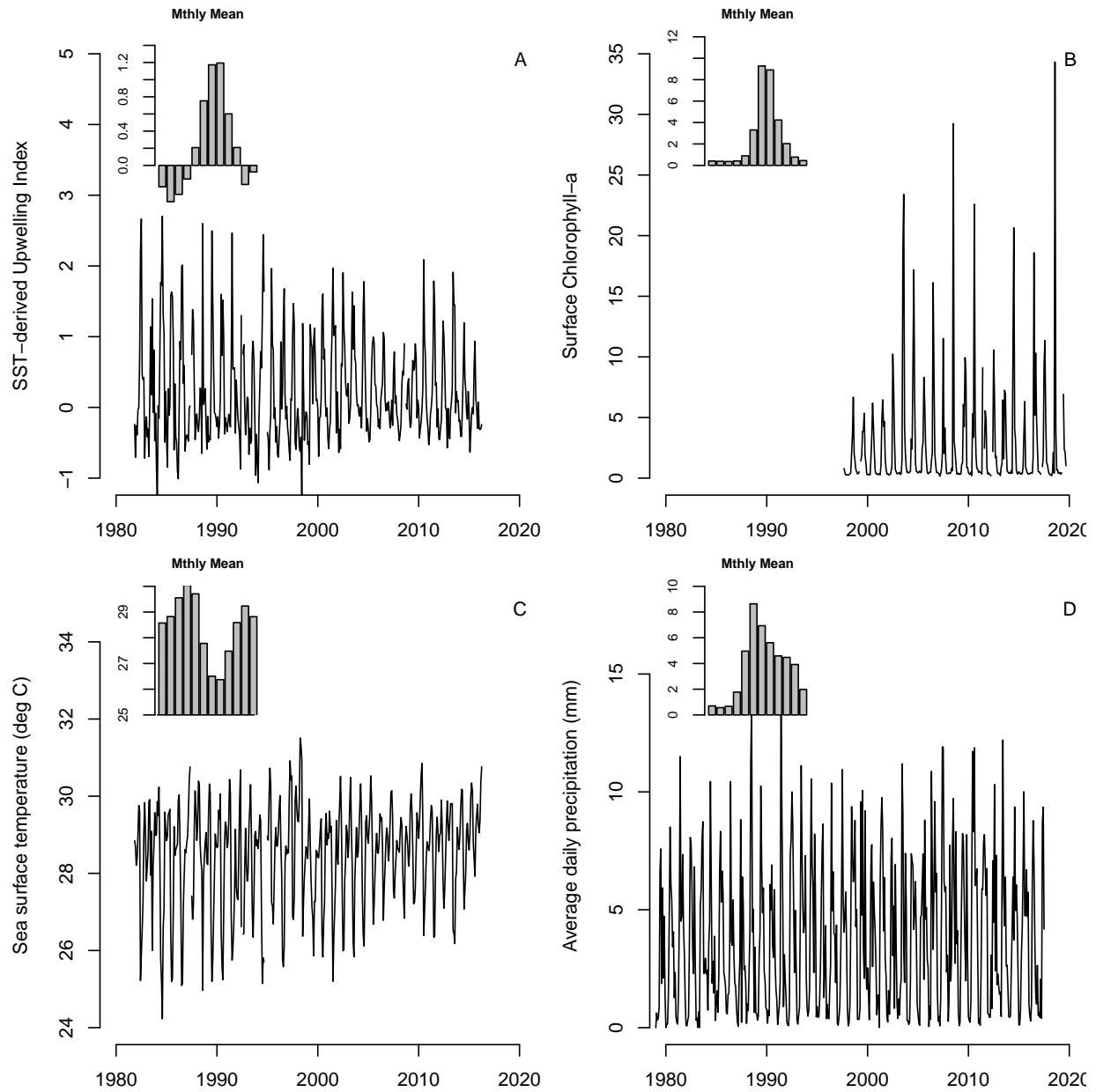


Figure 4

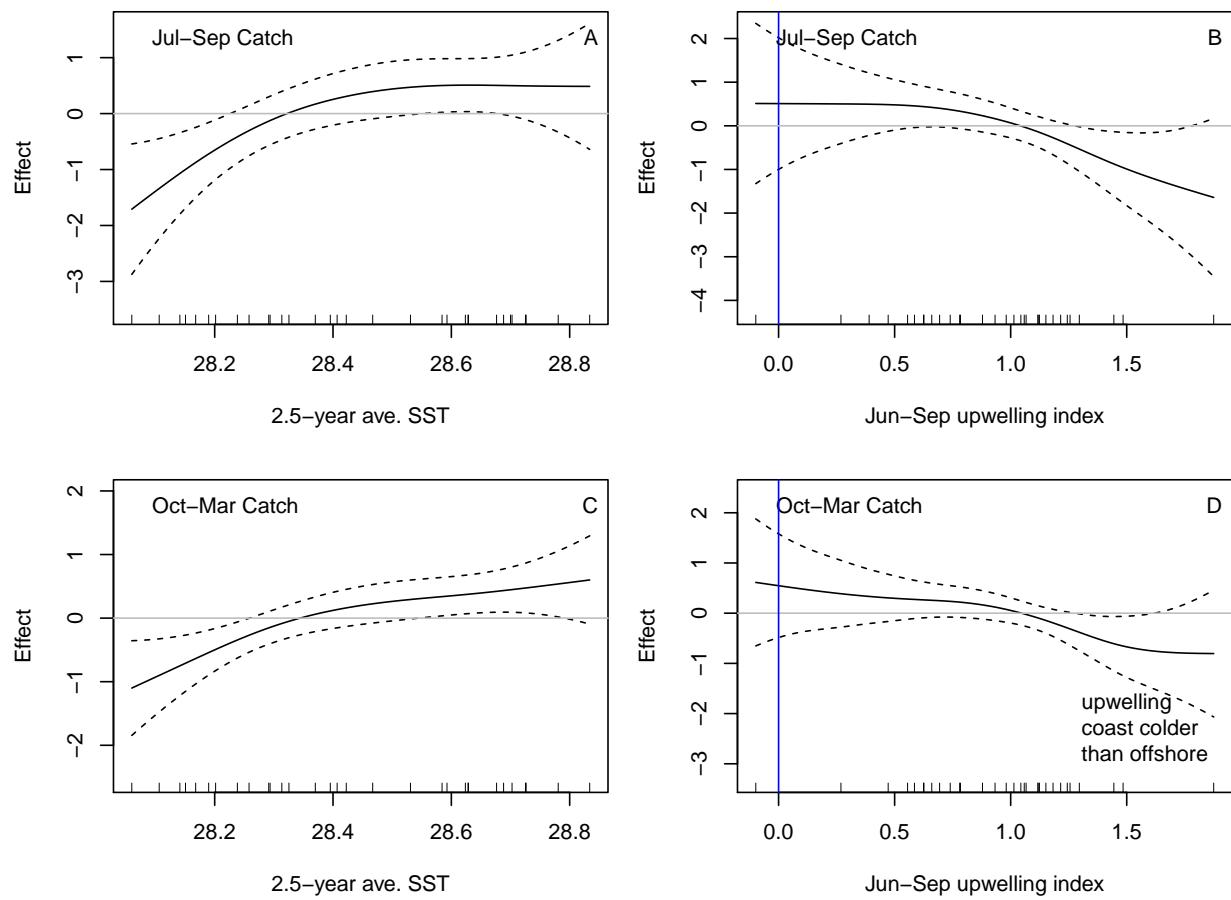


Figure 5

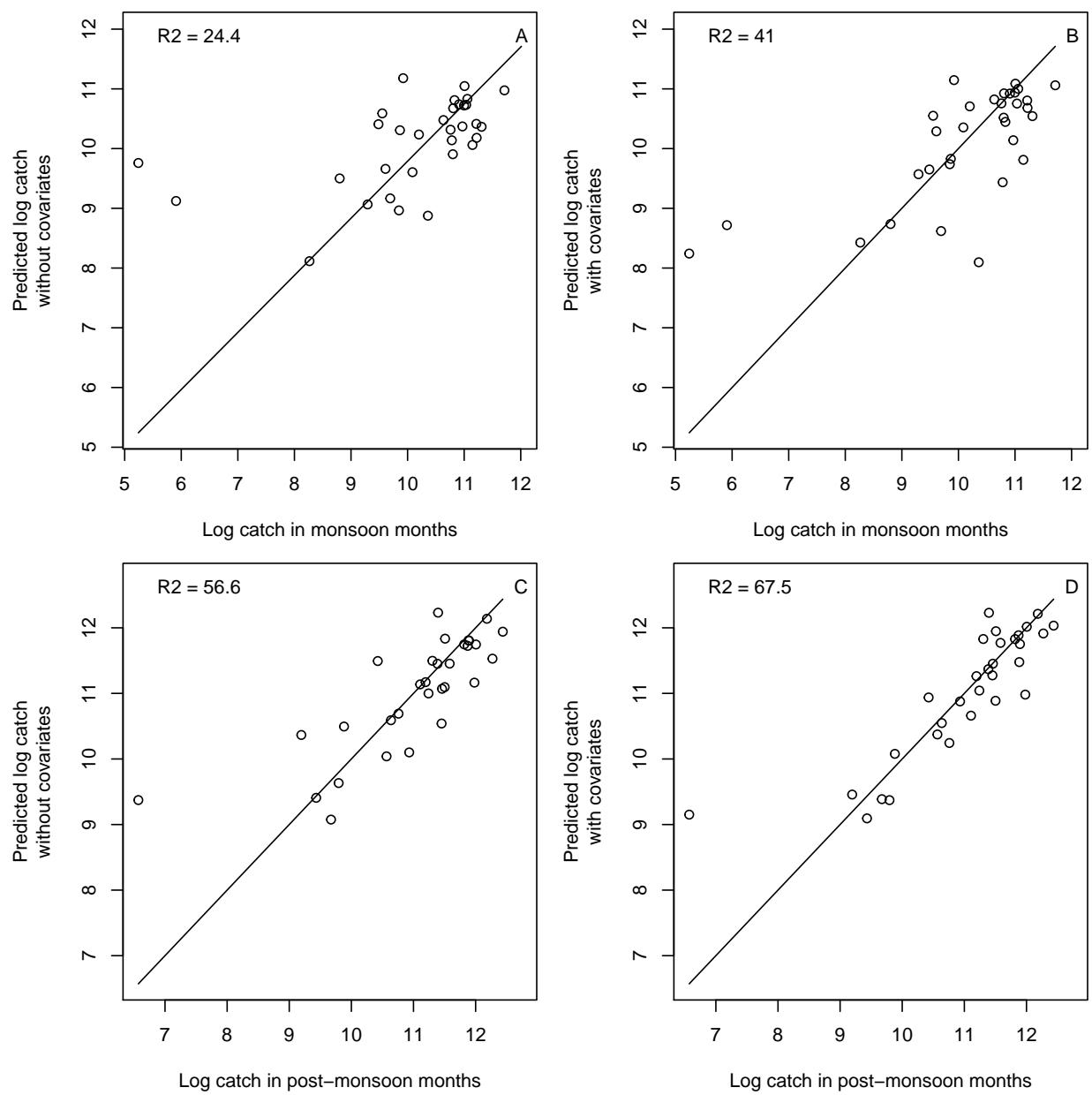


Figure 6