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Virtual Conference on the Digital Subsurface, 16–23 April

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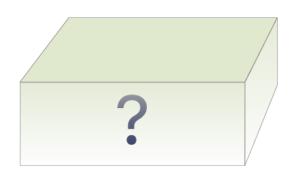




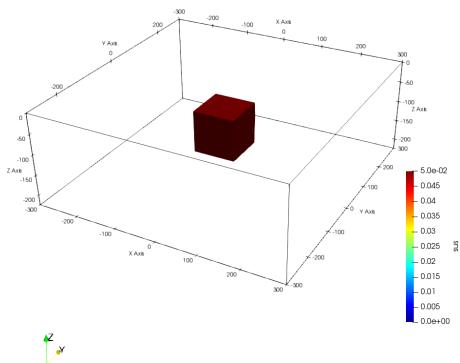
### Outline

- Backgrounds: Magnetics
- Inversion Framework
- 1D Linear Inverse problem
- 3D Magnetic Inversion
- Including Geologic Information
- Summary



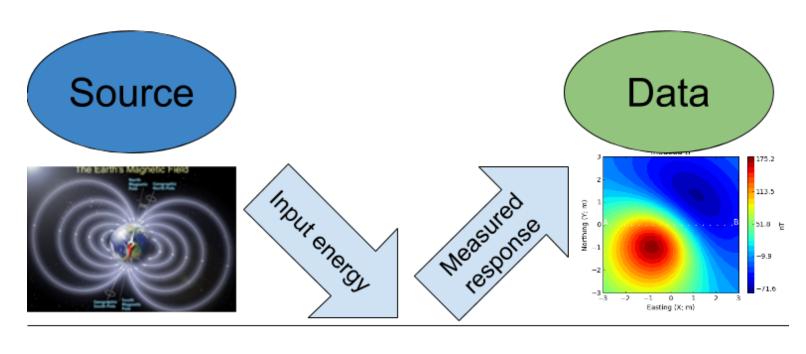


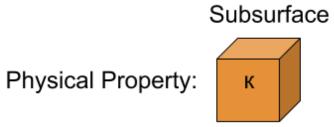
# 3D magnetic inversion





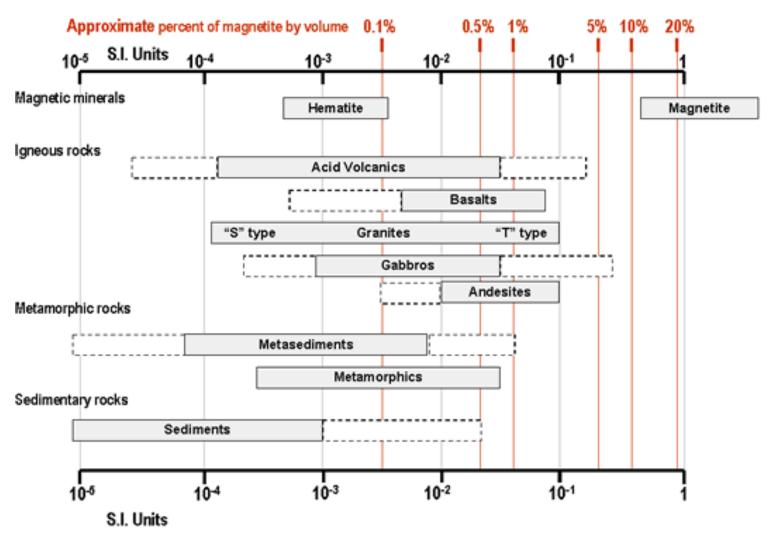
# Survey: Magnetics





к: Magnetic susceptibility

# Magnetic susceptibility

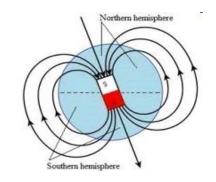


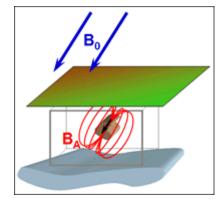
# Magnetic surveying

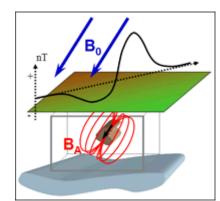
- Earth's magnetic field  $\vec{B}_0$  is the source:
- Materials become magnetized

$$\vec{M} = \kappa \vec{H}_0$$
 (magnetization)  
 $\vec{H}_0 = \vec{B}_0/\mu_0$ 

- Create anomalous magnetic field
- Measure total magnetic field:  $|\vec{B}| = |\vec{B}_0 + \vec{B}_A|$
- Total field anomaly:  $\triangle \vec{B} = |\vec{B}_0 + \vec{B}_A| |\vec{B}_0|$   $\triangle \vec{B} \simeq \vec{B}_A \cdot \hat{B}_0$  where  $\hat{B}_0 = \frac{\vec{B}_0}{|\vec{B}_0|}$





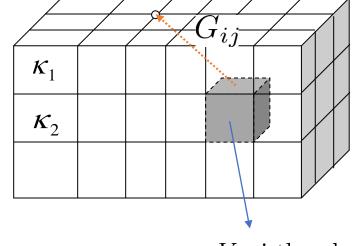


### Forward modelling

Discretize earth

$$\kappa_j \ (j=1,\ldots,M)$$
 susceptibility

Magnetic anomaly data are

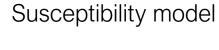


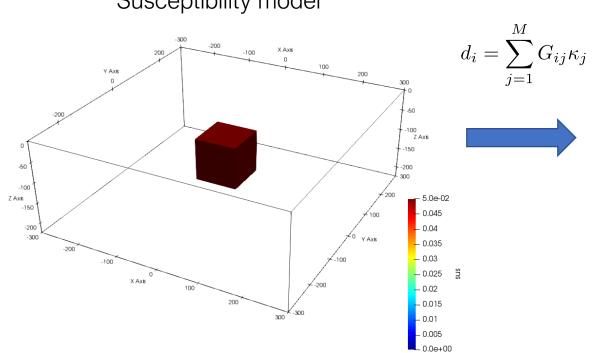
 $V_j$ : j-th volume

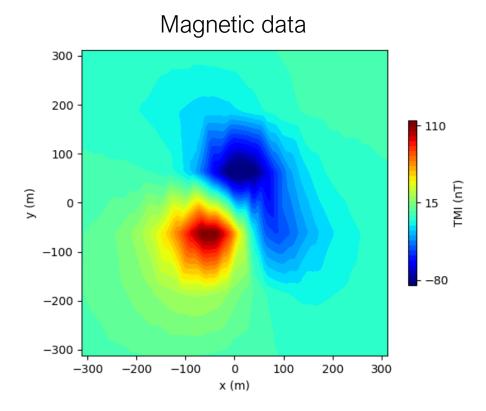
$$d_i = \sum_{j=1}^{M} G_{ij} \kappa_j$$

$$\begin{cases} G_{ij} = \hat{B}_0 \cdot \left\{ \frac{\mu_0}{4\pi} \int_v \kappa \nabla \nabla \left( \frac{1}{r_i - r_j} \right) dV_j \right\} \cdot \hat{B}_0 \\ \hat{B}_0 = \frac{\vec{B}_0}{|\vec{B}_0|} \end{cases}$$

# Forward modelling

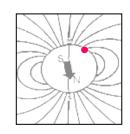


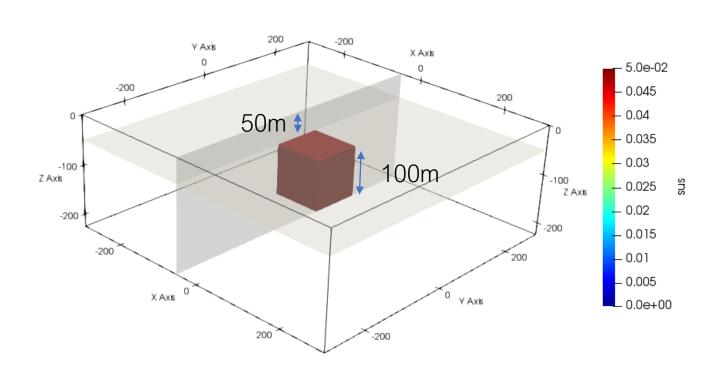






# Synthetic susceptibility model



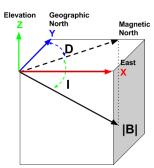


#### Earth field

- Inclination: 30°

- Declination: 45°

 $- |B_0| = 50,000 \text{ nT}$ 



#### Susceptible block

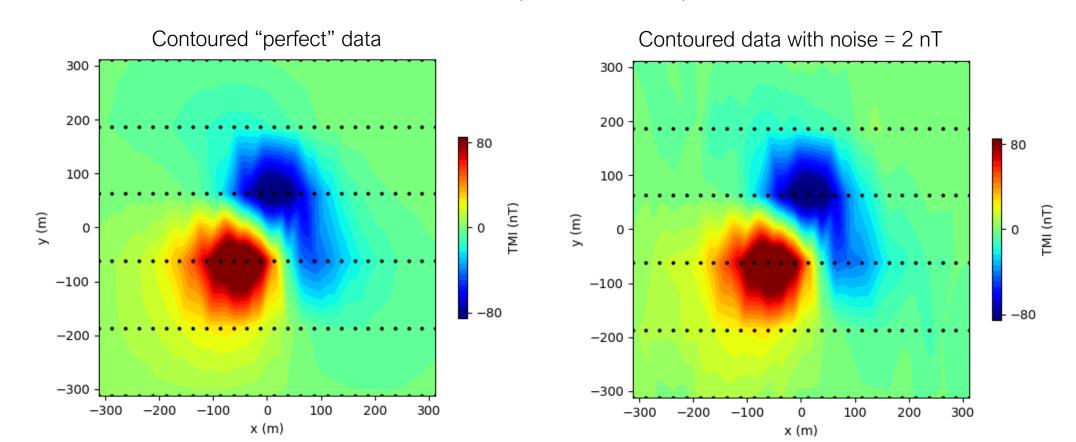
- 100m x 100m x 100m block
- Block susceptibility = 0.5
- Block top = 50m



### Synthetic survey

Survey parameters: - 100 m line spacing.

- 25 station spacing.
- N=156 (elevation= 2m)



### Solving inverse problem

#### Model objective function

$$\phi_m = \alpha_s \int_v w_s \left(\kappa - \kappa_{\text{ref}}\right)^2 dv + \alpha_x \int_v \left(\frac{d\kappa}{dx}\right)^2 dx + \alpha_y \int_v \left(\frac{d\kappa}{dy}\right)^2 dy + \alpha_z \int_v \left(\frac{d\kappa}{dz}\right)^2 dz$$

Data misfit 
$$\phi_d = \sum_{j=1}^N \left( \frac{G_{ij}\kappa_j - d_j^{obs}}{\epsilon_j} \right)$$

Choose 
$$\kappa_{ref}=0, \alpha_s=0.0001, \alpha_x=\alpha_y=\alpha_z=1$$
 
$$L_x=\sqrt{\frac{\alpha_x}{\alpha_s}}=100$$

#### The Inverse problem is:

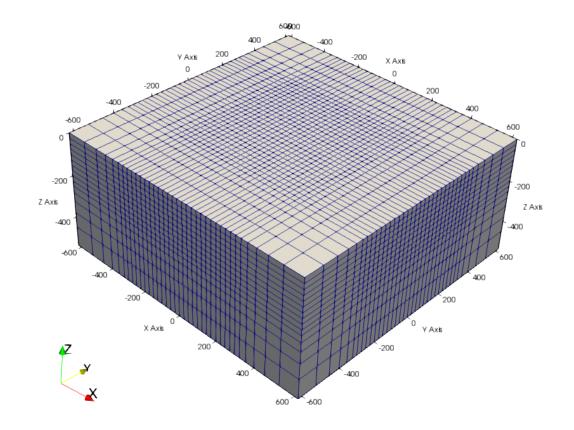
minimize 
$$\phi = \phi_d + \beta \phi_m$$

find  $\beta$  such that  $\phi_d = \phi_d^*$  where  $\phi_d = N$ 

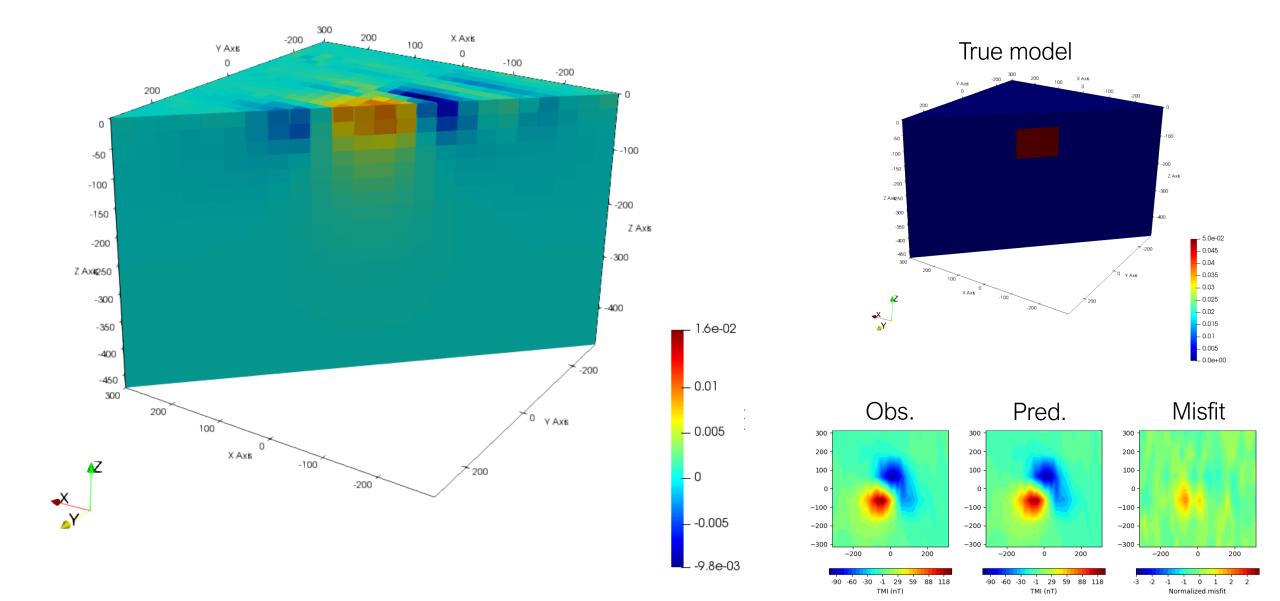
### Discretization

- Earth model for inversion:
  - dx=dy=dz=25m
  - N/S and E/W padding = 300m
  - Number of cells (M) = 19440

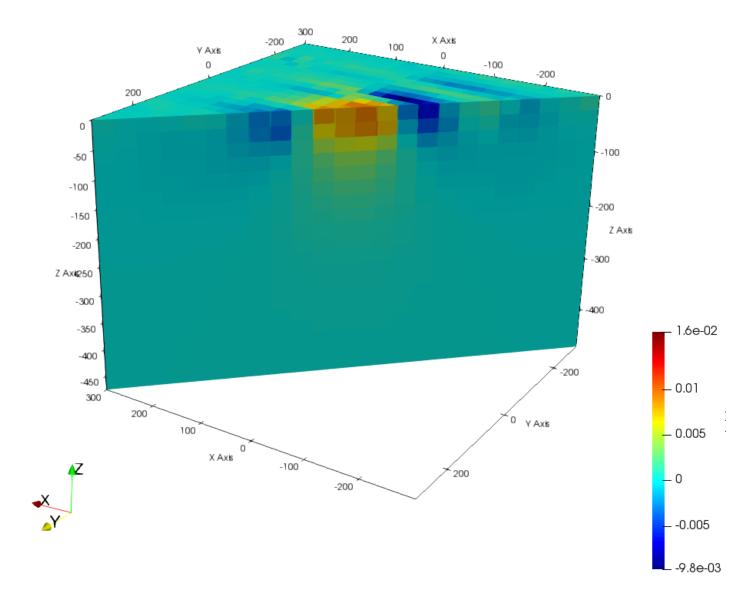
- Therefore:
  - No. data is N = 176
  - No. unknowns is M = 11,492



### Inversion results



### Inversion results



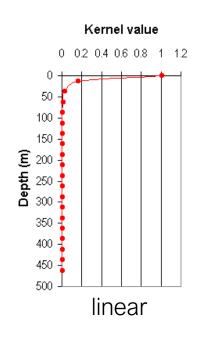
### Two primary difficulties:

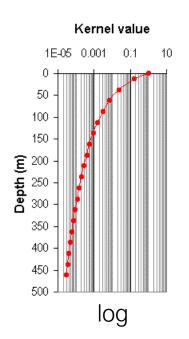
- Concentration of susceptibility is near the surface
- Regions of negative susceptibility

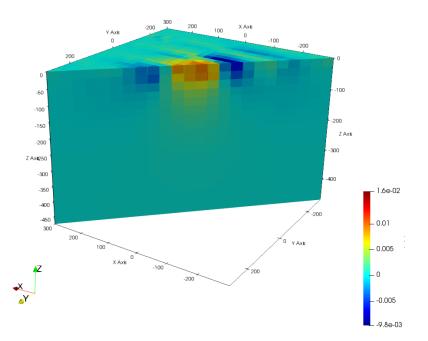
### What went wrong?

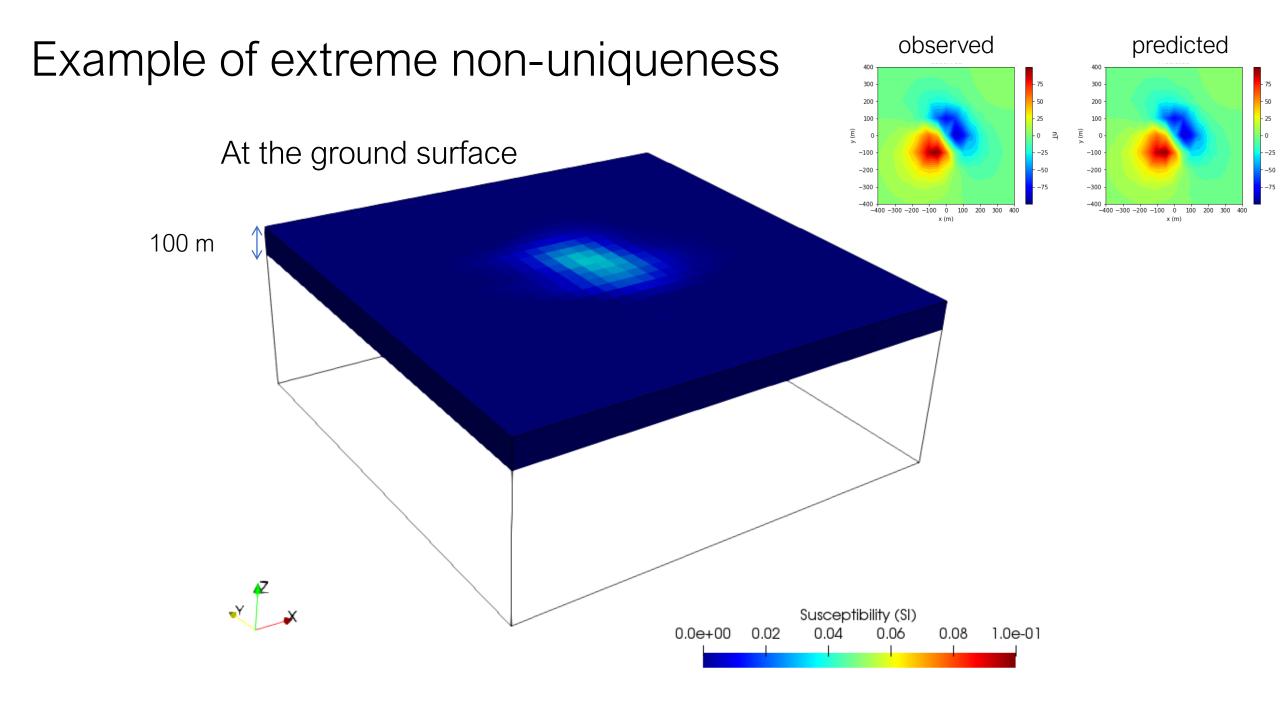
Fundamental non-uniqueness of all potential fields:

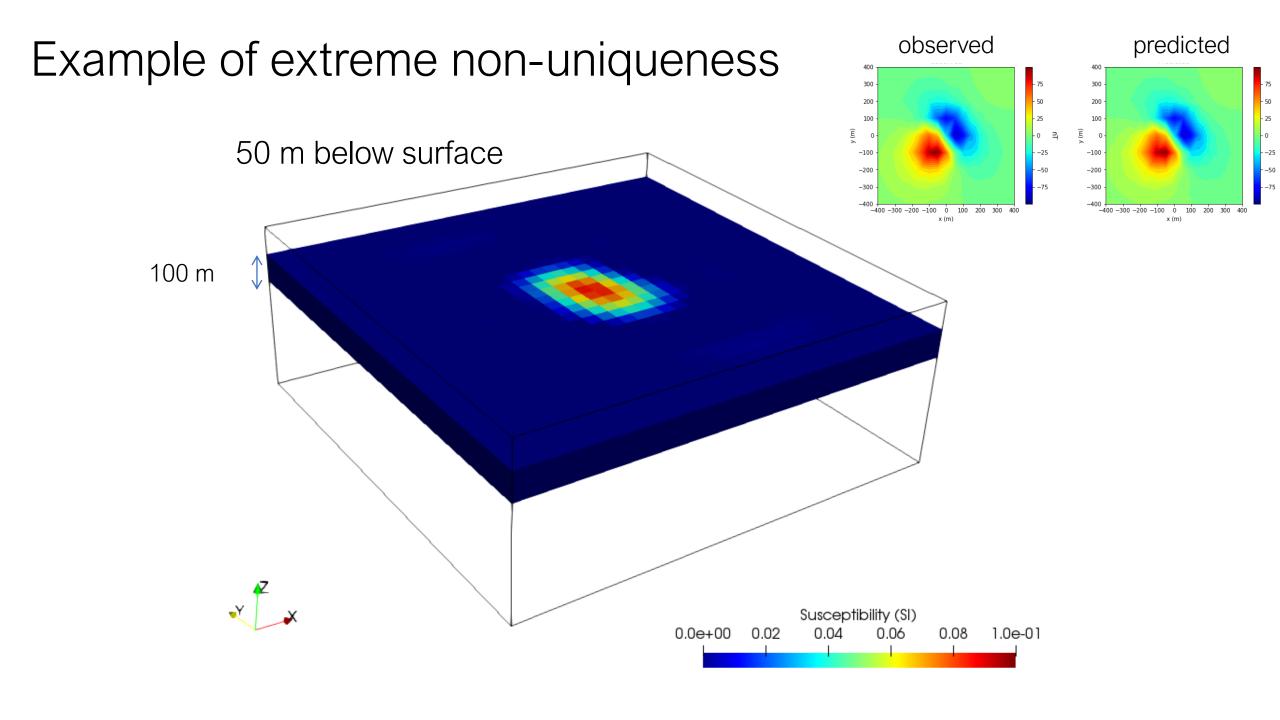
- As a consequence of Green's third identity ...
  - an observed magnetic field can be caused by a thin layer of susceptible material at any arbitrary depth
- The rapid decay of our kernels causes a concentration of  $\kappa$  near the surface to be a preferred solution.

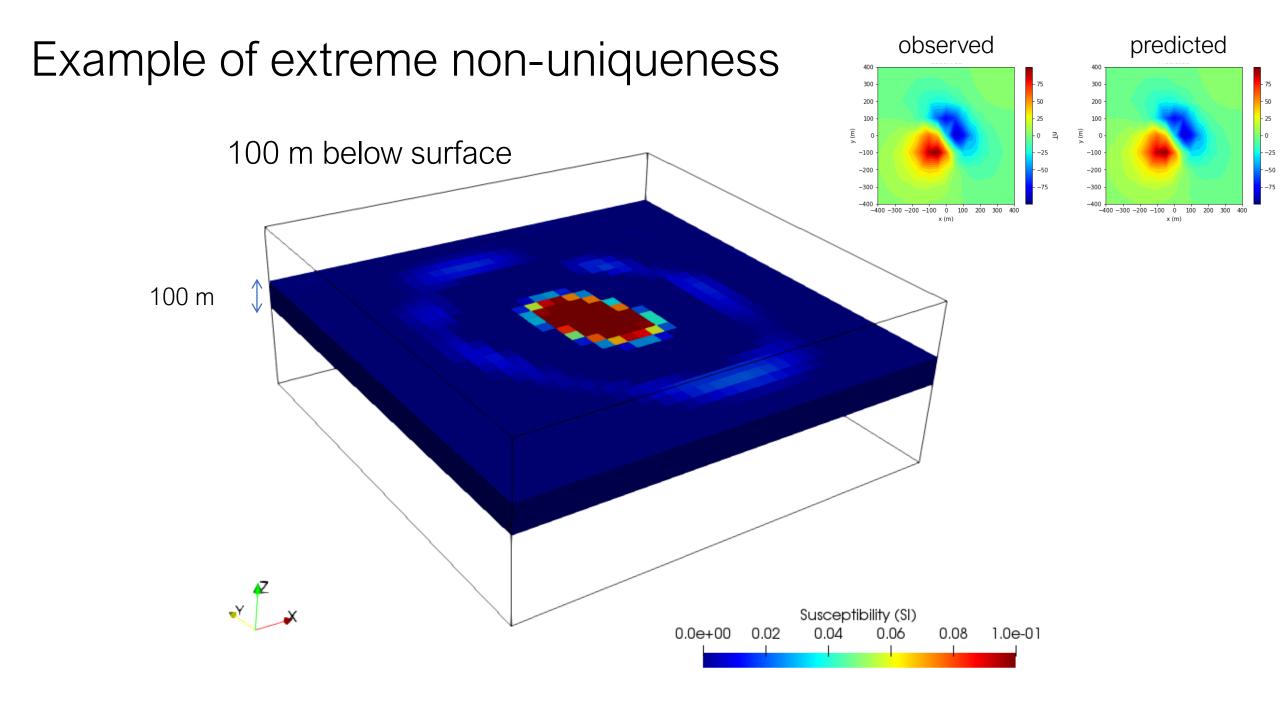


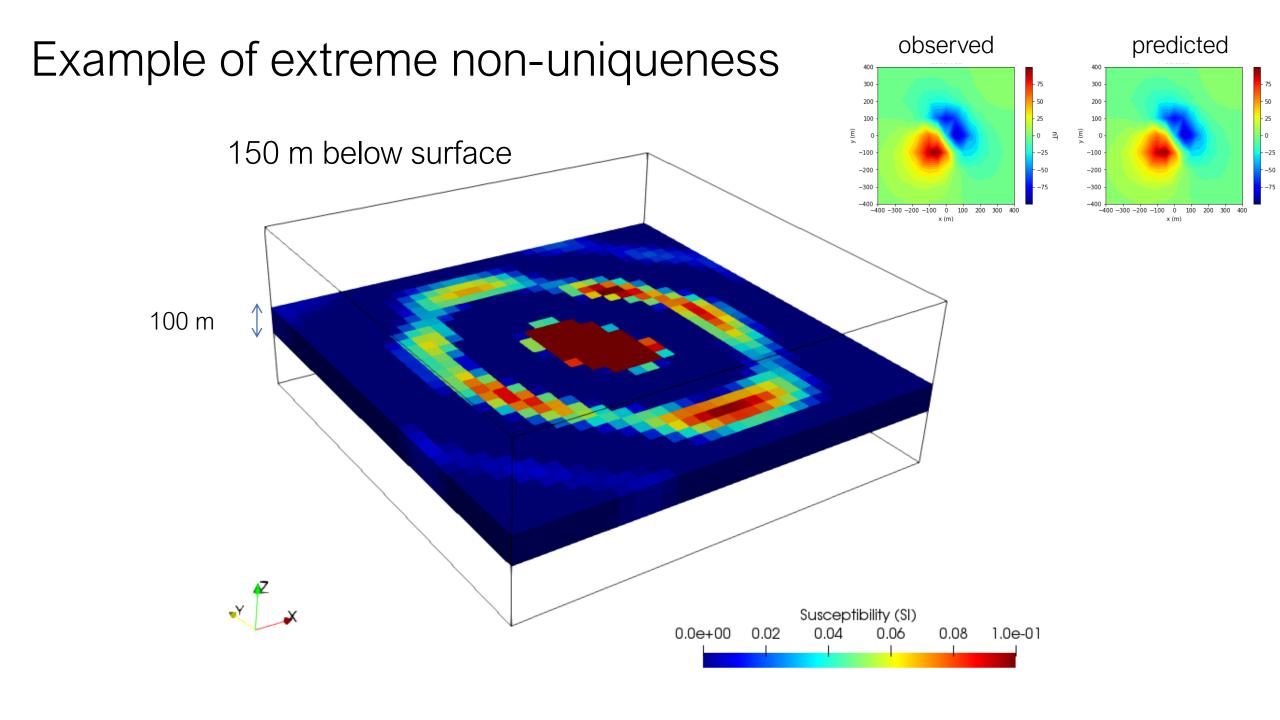












### Inversion with sensitivity weighting

#### Model objective function

$$\phi_m = \alpha_s \int_v w_s \left(\kappa - \kappa_{\text{ref}}\right)^2 dv + \alpha_x \int_v w_x \left(\frac{d\kappa}{dx}\right)^2 dx + \alpha_y \int_v w_y \left(\frac{d\kappa}{dy}\right)^2 dy + \alpha_z \int_v w_z \left(\frac{d\kappa}{dz}\right)^2 dz$$

Data misfit 
$$\phi_d = \sum_{j=1}^{N} \left( \frac{G_{ij} \kappa_j - d_j^{obs}}{\epsilon_j} \right)$$

 $\{w_s, w_x, w_y, w_z\}$ : additional weightings

Choose  $\{w_s, w_x, w_y, w_z\} \propto \frac{1}{(z+z_0)^3}$ 

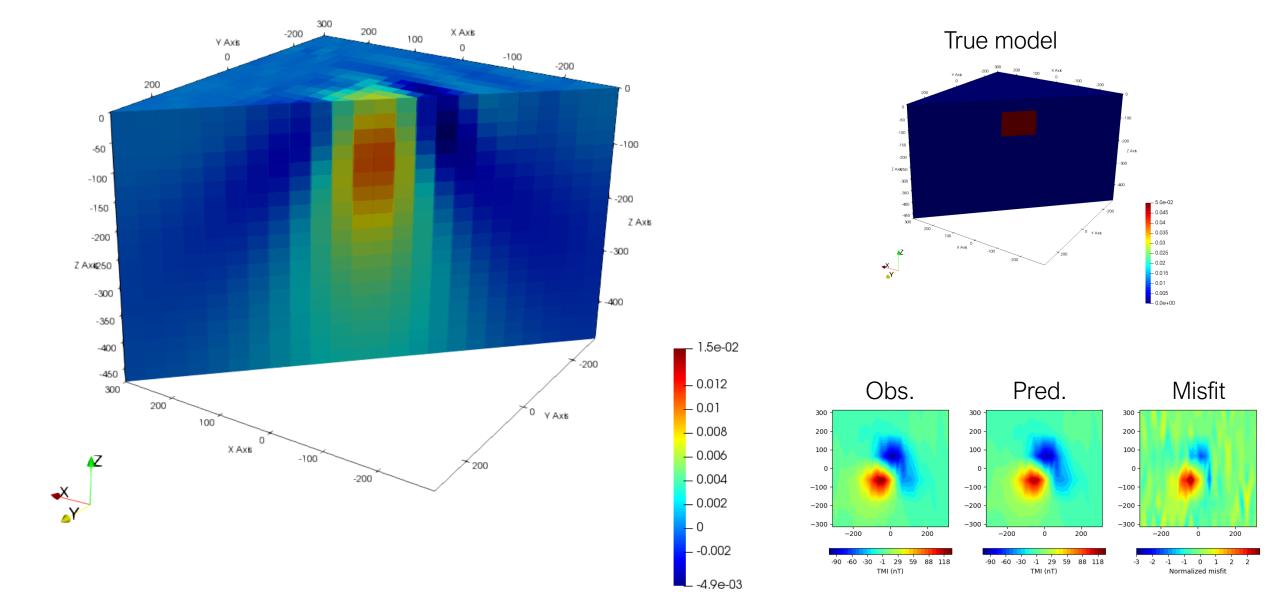
Allows cells at depth to contribute

#### The Inverse problem is:

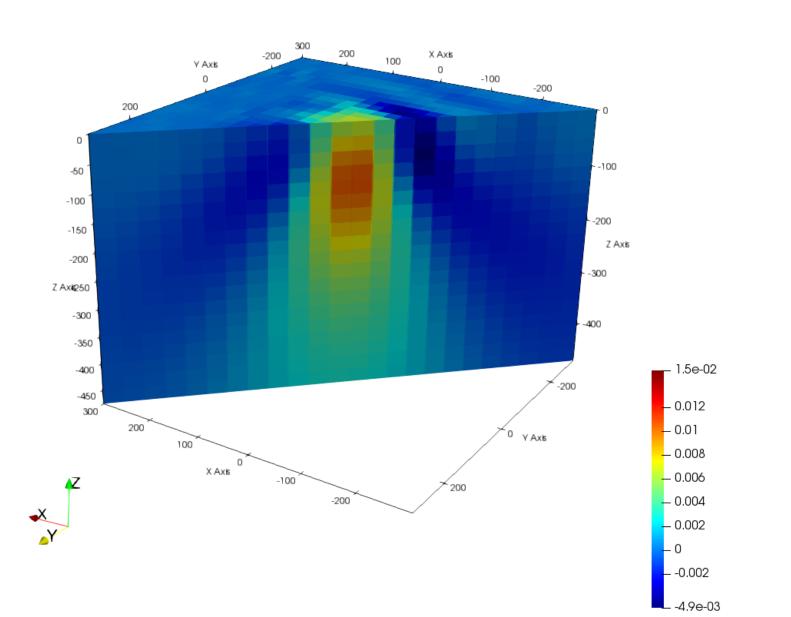
minimize 
$$\phi = \phi_d + \beta \phi_m$$

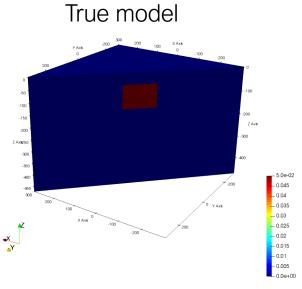
find  $\beta$  such that  $\phi_d = \phi_d^*$  where  $\phi_d = N$ 

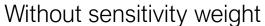
# Inversion with sensitivity weighting

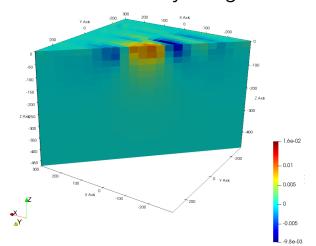


# Inversion with sensitivity weighting



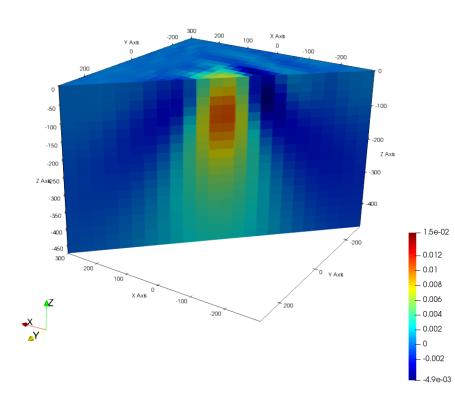






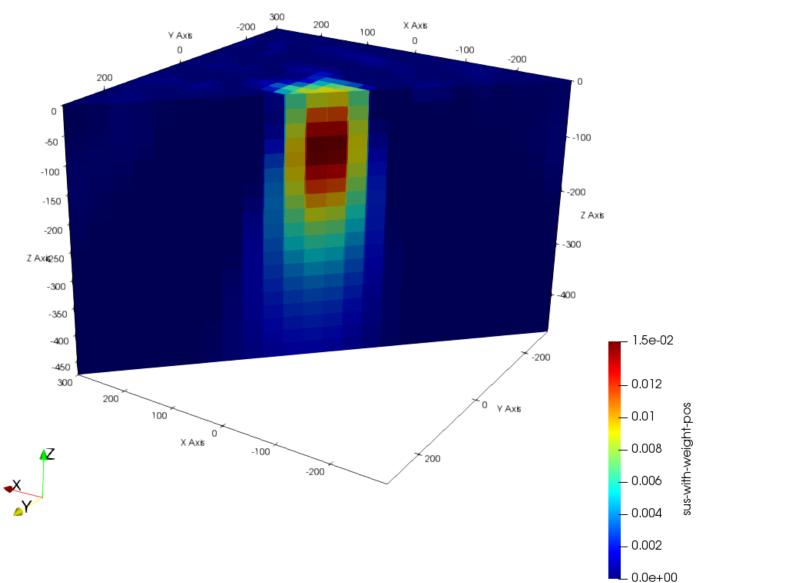
# Summary for sensitivity weighting

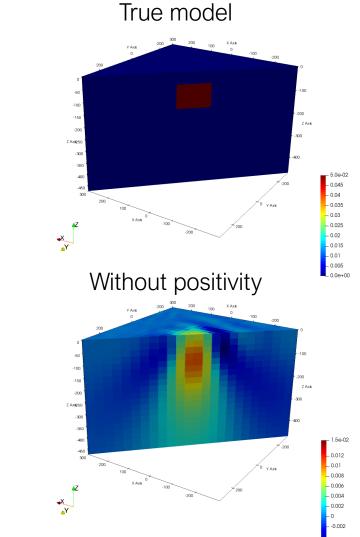
- Structure is no longer concentrated at surface.
- Main anomaly is at a reasonable depth.
- BUT:
  - Negative κ persists
  - There is a long tail extending down and out.
- Require positivity



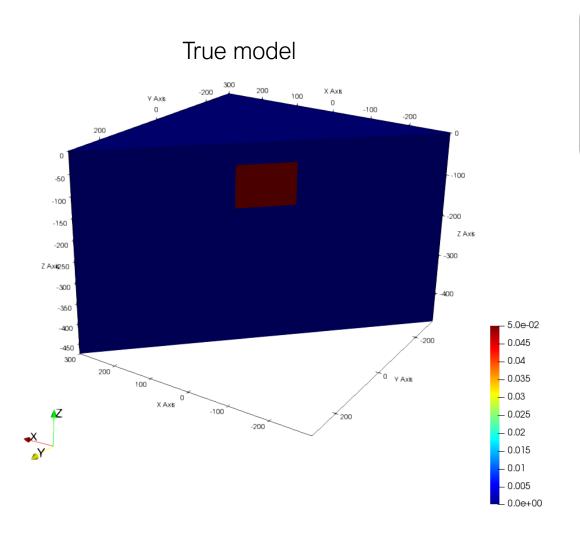
3D magnetic inversion:

- depth (or sensitivity) weighting
- positivity (bounds):  $\mathbf{m} \geq 0$





### Think about the spatial character of the true model



Most of model parameters are zero

Susceptible block and background has sharp boundaries

#### Model norm:

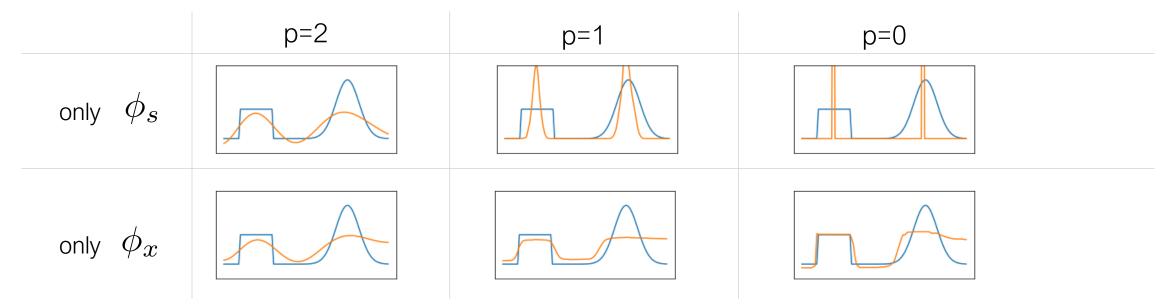
$$\phi_m = \alpha_s \int_v w_s \left(\kappa - \kappa_{\text{ref}}\right)^2 dv + \alpha_s \int_v w_s \left(\frac{d\kappa}{dx}\right)^2 dx + \alpha_y \int_v w_s \left(\frac{d\kappa}{dy}\right)^2 dy + \alpha_z \int_v w_s \left(\frac{d\kappa}{dz}\right)^2 dz$$

### General character

$$\phi_m = \sum_{i=1}^M |m_i|^p v_i$$

- Geometric character
  - p=2: all elements close to zero
  - p=1: sparse solution, # of non-zero elements are ≤ # of data
  - p=0: minimum support, model with the fewest number of elements

### • 1D problem



# Magnetic inversion with Lp norms

Model objective function

$$\phi_m = \alpha_s \int_v w_s \left| \kappa - \kappa_{\text{ref}} \right|^{p_s} dv + \alpha_x \int_v w_x \left| \frac{d\kappa}{dx} \right|^{p_x} dx + \alpha_y \int_v w_y \left| \frac{d\kappa}{dy} \right|^{p_y} dy + \alpha_z \int_v w_z \left| \frac{d\kappa}{dz} \right|^{p_z} dz$$

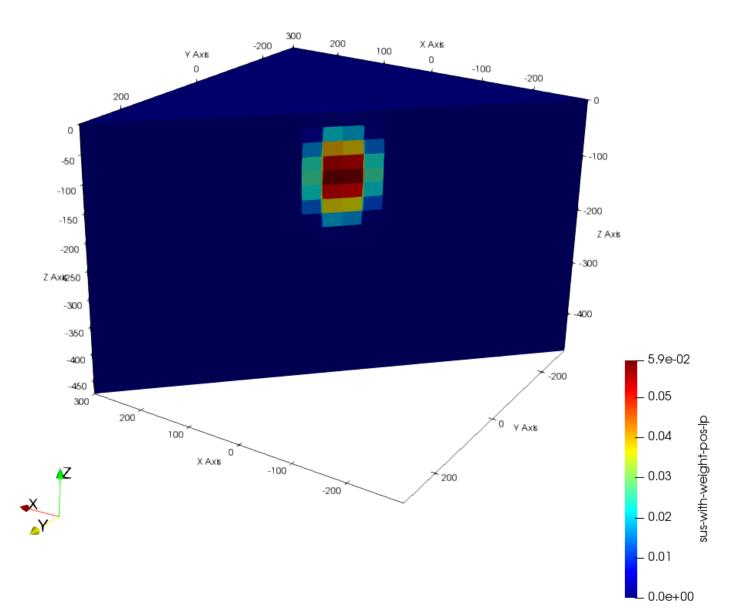
Data misfit 
$$\phi_d = \sum_{j=1}^N \left( \frac{G_{ij} \kappa_j - d_j^{obs}}{\epsilon_j} \right)$$

Fournier and Oldenburg (2019)

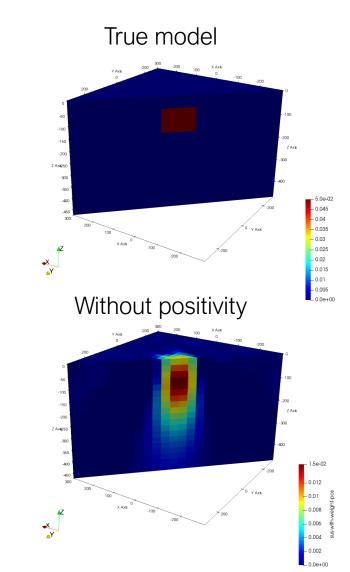
The Inverse problem is:

minimize 
$$\phi = \phi_d + \beta \phi_m$$
  
find  $\beta$  such that  $\phi_d = \phi_d^*$  where  $\phi_d = N$   
subject to  $\kappa \geq 0$ 

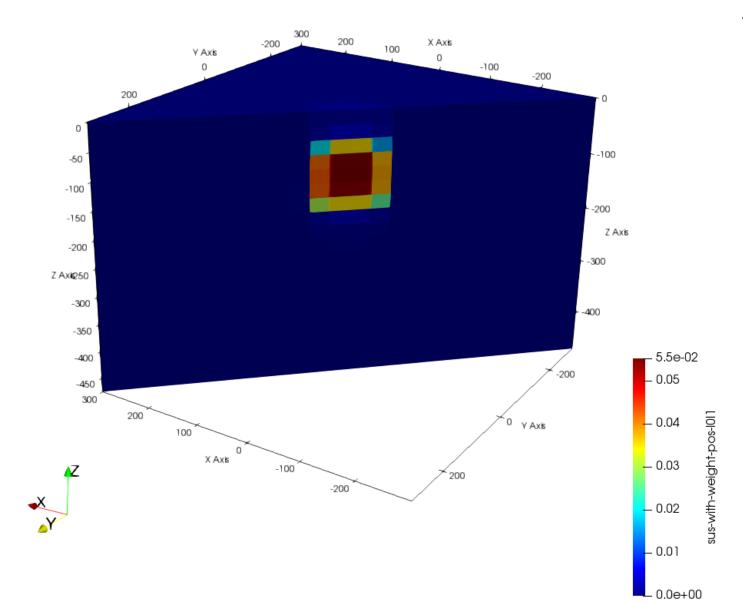
### 3D magnetic inversion:



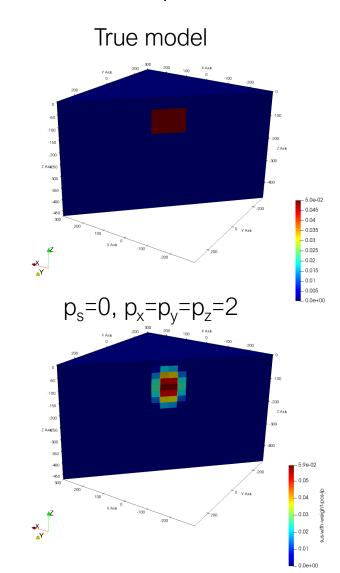
- depth (or sensitivity) weighting
- positivity (bounds):  $\mathbf{m} \geq 0$
- $L_p$  norm ( $p_s$ =0,  $p_x$ = $p_y$ = $p_z$ =2)



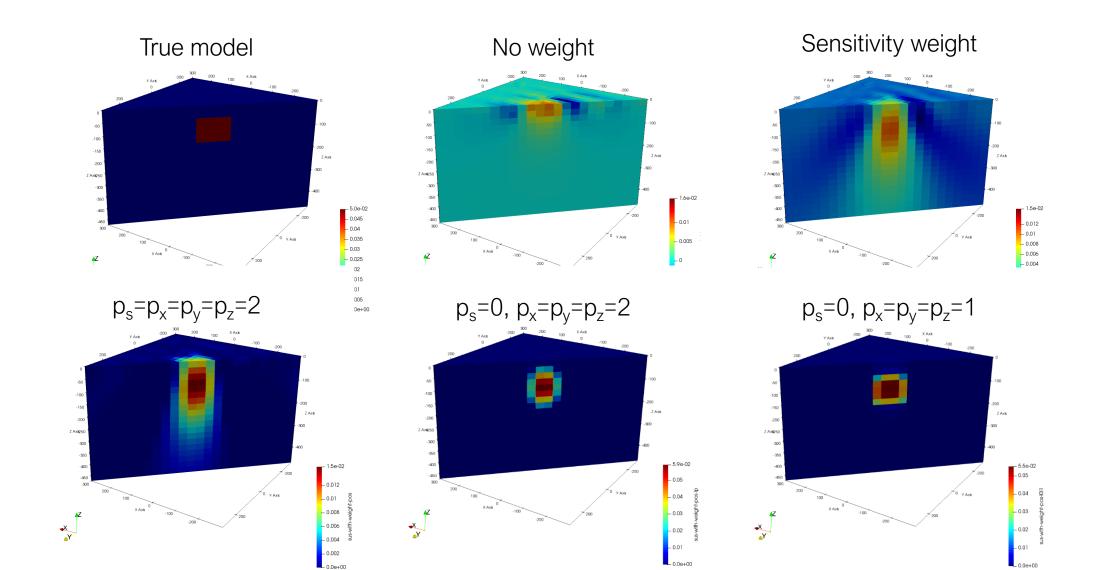
### 3D magnetic inversion:



- depth (or sensitivity) weighting
- positivity (bounds):  $\mathbf{m} \geq 0$
- $L_p$  norm ( $p_s=0$ ,  $p_x=p_v=p_z=1$ )

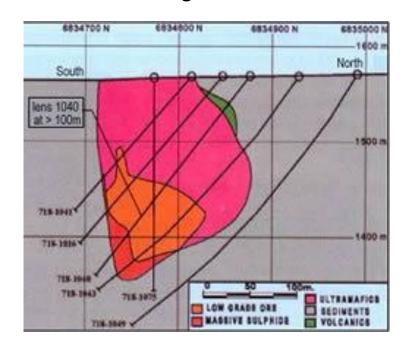


# Summary: magnetic inversion

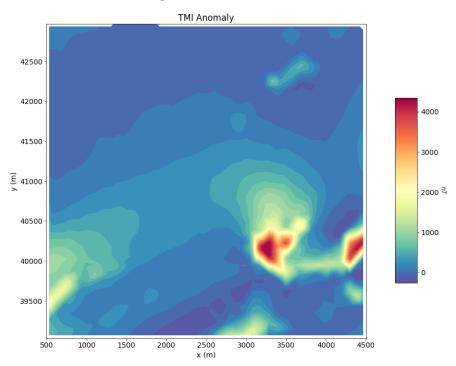


### Revisit the case history

Geologic section



#### Magnetic data

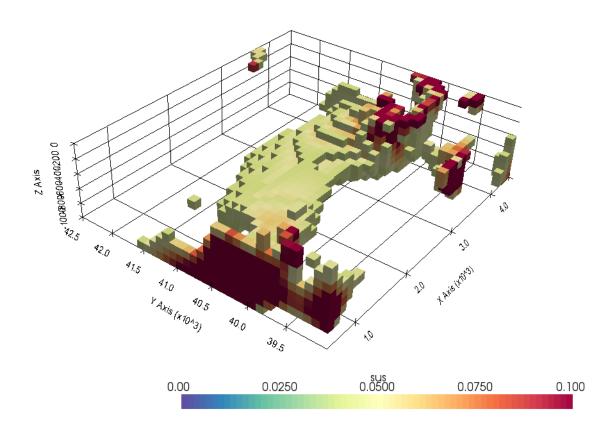


Initial conceptual model: two ultramafic pipes

Can make impact on drilling location and mineral reserve

# Magnetic inversion

The recovered model (SimPEG)



Model norm:

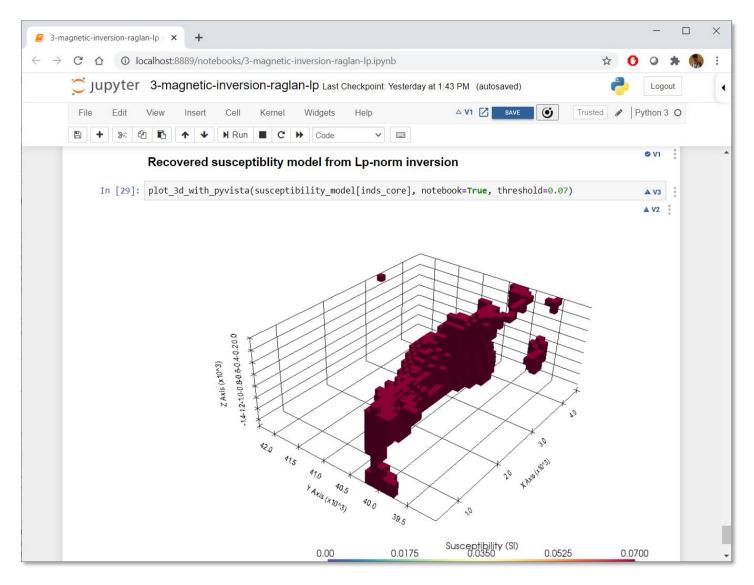
$$\phi_m = \alpha_s \int_v w_s \left(\kappa - \kappa_{\text{ref}}\right)^2 dv +$$

$$\alpha_x \int_v w_x \left(\frac{d\kappa}{dx}\right)^2 dx + \alpha_y \int_v w_y \left(\frac{d\kappa}{dy}\right)^2 dy + \alpha_z \int_v w_z \left(\frac{d\kappa}{dz}\right)^2 dz$$

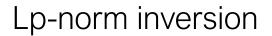
What if the target body is expected to be compact?

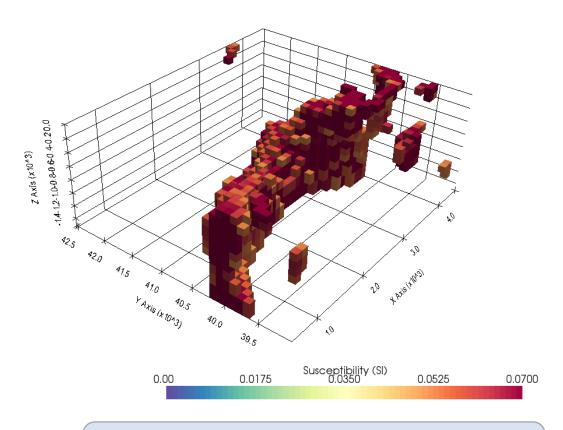
Use a sparse norm (e.g.,  $p_s$ =0)

### Lp-norm inversion of the Raglan magnetic data



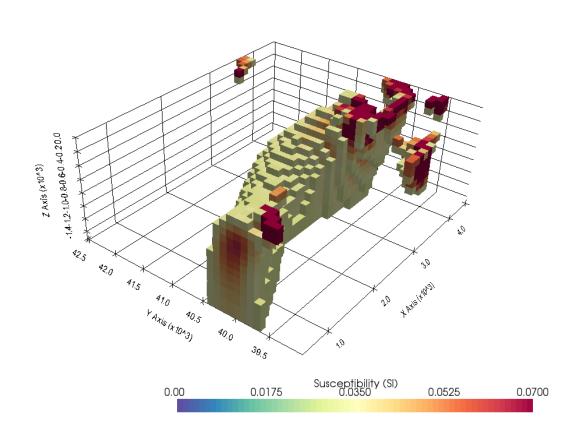
### Comparison: Lp-norm vs L2-norm





#### Beneficial for reserve estimate

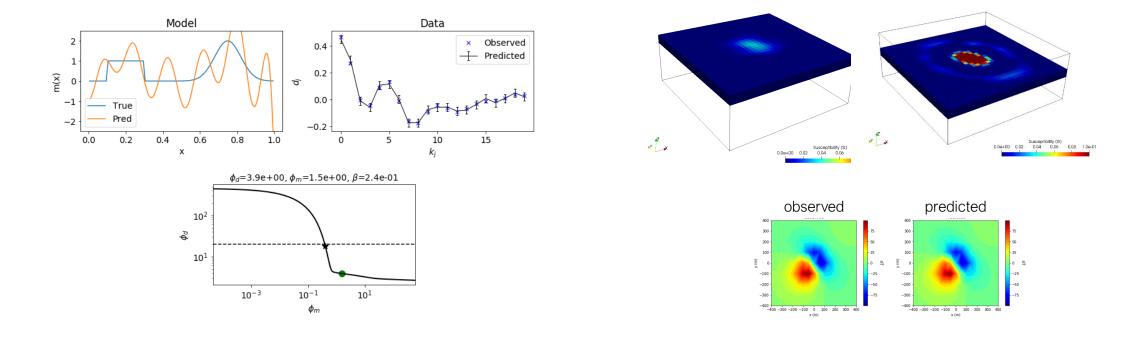
#### L2-norm inversion



### Summary

- Increasing complexity & volume of data
  - Increasing needs of data-driven approaches

Fitting the observed data is not an enough condition!



### Summary

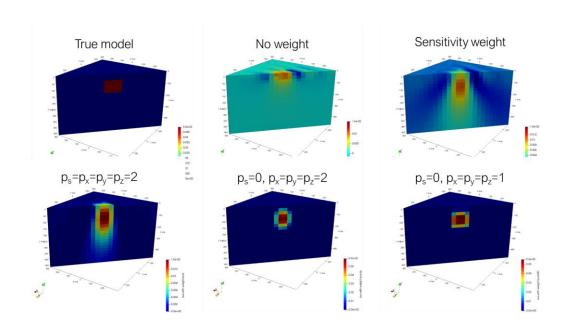
- An inversion framework can play an important role
  - Data-driven + Physics-driven + Prior knowledge

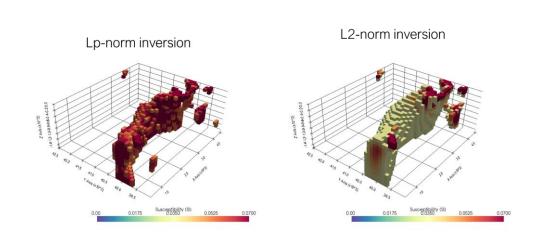
minimize 
$$\phi(m) = \phi_d(m) + \beta \phi_m(m)$$

 $\phi_d$ : data misfit

 $\phi_m$ : model norm

 $\beta$ : trade-off parameter



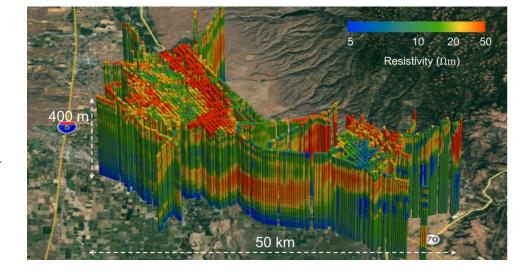


# Multiple airborne geophysics

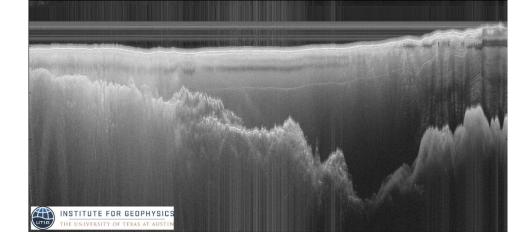
Potential fields
 <u>Magnetics</u>
 Gravity



Electromagnetics

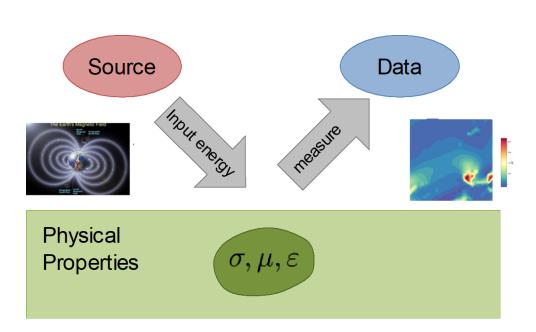


• (Ice) Radar



Increasing Resolution

### But in a generic level, they are very similar ...

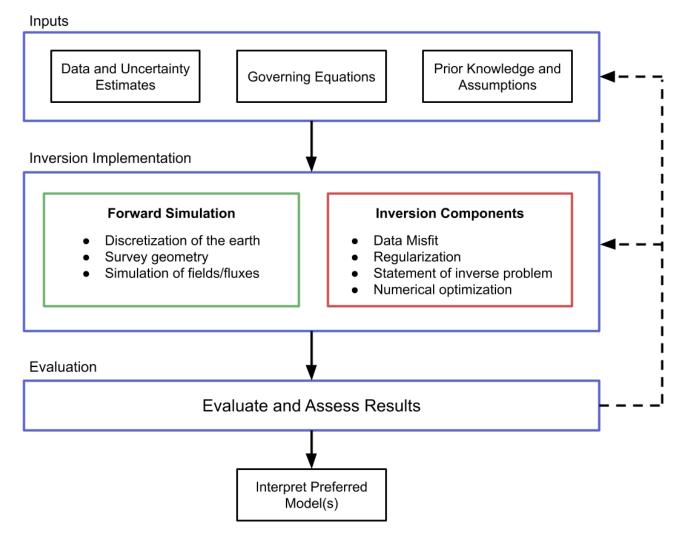


minimize  $\phi(m) = \phi_d(m) + \beta \phi_m(m)$ 

 $\phi_d$ : data misfit  $\phi_m$ : model norm  $\beta$ : trade-off parameter

# SimPEG provides ..

#### A common & modular inversion framework





- Gravity
- Magnetics
- Direct current resistivity
- Induced polarization
- Electromagnetics
- Fluid flow

### There are many other software packages

And they are growing!











### Open data











?













### Challenging geoscience problems



### Open data











?

Inversion framework can provide a "data-driven" approach

### Challenging geoscience problems



### Resources

Slides: <a href="http://bit.ly/transform-2021-slides">http://bit.ly/transform-2021-slides</a>

Notebooks: <a href="https://curvenote.com/@swung/inversion-for-geologists-transform-2021">https://curvenote.com/@swung/inversion-for-geologists-transform-2021</a>

Magnetics: <a href="https://gpg.geosci.xyz/content/magnetics/index.html">https://gpg.geosci.xyz/content/magnetics/index.html</a>

Github: <a href="https://github.com/simpeg/transform-2021-simpeg">https://github.com/simpeg/transform-2021-simpeg</a>

SimPEG: <a href="https://www.simpeg.xyz">https://www.simpeg.xyz</a>

# Thank you!





#### The livestream has ended.

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Virtual Conference on the Digital Subsurface, 16–23 April

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