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Quantifying the relative importance of greenhouse gas emissions from current and future savanna land use change across northern Australia

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Abstract. The clearing and burning of tropical savanna leads to globally significant emissions of greenhouse gases (GHGs); however there is large uncertainty relating to the magnitude of this flux. Australia's tropical savannas occupy the northern quarter of the continent, a region of increasing interest for further exploitation of land and water resources. Land use decisions across this vast biome have the potential to influence the national greenhouse gas budget. To better quantify emissions from savanna deforestation and investigate the impact of deforestation on national GHG emissions, we undertook a paired site measurement campaign where emissions were quantified from two tropical savanna woodland sites; one that was deforested and prepared for agricultural land use and a second analogue site that remained uncleared for the duration of a 22-month campaign. At both sites, net ecosystem exchange of CO₂ was measured using the eddy covariance method. Observations at the deforested site were continuous before, during and after the clearing event, providing high-resolution data that tracked CO₂ emissions through nine phases of land use change. At the deforested site, post-clearing debris was allowed to cure for 6 months and was subsequently burnt, followed by extensive soil preparation for cropping.

During the debris burning, fluxes of CO₂ as measured by the eddy covariance tower were excluded. For this phase, emissions were estimated by quantifying on-site biomass prior to deforestation and applying savanna-specific emission factors to estimate a fire-derived GHG emission that included both CO_2 and non- CO_2 gases. The total fuel mass that was consumed during the debris burning was $40.9\,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}$ and included above- and below-ground woody biomass, course woody debris, twigs, leaf litter and C_4 grass fuels. Emissions from the burning were added to the net CO_2 fluxes as measured by the eddy covariance tower for other post-deforestation phases to provide a total GHG emission from this land use change.

The total emission from this savanna woodland was $148.3\,\mathrm{Mg\,CO_2}$ -e $\mathrm{ha^{-1}}$ with the debris burning responsible for $121.9\,\mathrm{Mg\,CO_2}$ -e $\mathrm{ha^{-1}}$ or $82\,\%$ of the total emission. The remaining emission was attributed to $\mathrm{CO_2}$ efflux from soil disturbance during site preparation for agriculture (10 % of the total emission) and decay of debris during the curing period prior to burning (8 %). Over the same period, fluxes at the uncleared savanna woodland site were measured using a second flux tower and over the 22-month observation period, cumulative net ecosystem exchange (NEE) was a net carbon sink of $-2.1\,\mathrm{Mg\,C\,ha^{-1}}$, or $-7.7\,\mathrm{Mg\,CO_2}$ -e $\mathrm{ha^{-1}}$.

Estimated emissions for this savanna type were then extrapolated to a regional-scale to (1) provide estimates of the magnitude of GHG emissions from any future deforestation and (2) compare them with GHG emissions from prescribed savanna burning that occurs across the northern Australian savanna every year. Emissions from current rate of annual savanna deforestation across northern Australia was double

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that of reported (non-CO₂ only) savanna burning. However, if the total GHG emission, CO2 plus non-CO2 emissions, is accounted for, burning emissions are an order of magnitude larger than that arising from savanna deforestation. We examined a scenario of expanded land use that required additional deforestation of savanna woodlands over and above current rates. This analysis suggested that significant expansion of deforestation area across the northern savanna woodlands could add an additional 3% to Australia's national GHG account for the duration of the land use change. This bottom-up study provides data that can reduce uncertainty associated with land use change for this extensive tropical ecosystem and provide an assessment of the relative magnitude of GHG emissions from savanna burning and deforestation. Such knowledge can contribute to informing land use decision making processes associated with land and water resource development.

1 Introduction

An increase in greenhouse gas (GHG) emissions through human-related activities is leading to rapid change in the climate system (IPCC, 2013). It is, therefore, crucial to obtain data describing the net GHG balance at regional to global scales to better characterise anthropogenic forcing of the atmosphere (Tubiello et al., 2015). Emissions from land use change (LUC) are the integral of ecosystem transformations that can include emissions from deforestation and conversion to agriculture, logging and harvest activity, shifting cultivation, as well as regrowth sinks following harvest and/or abandonment of previously cleared agriculture lands (Houghton al., 2012). At present, LUC emits $0.9 \pm 0.5 \,\mathrm{Pg}\,\mathrm{C}\,\mathrm{yr}^{-1}$ to the atmosphere, which is approximately 10% of anthropogenic carbon emissions (Le Quéré et al., 2014). Data sources and methods used to estimate LUC emissions are diverse. These include census-based historical land use reconstructions and land use statistics, satellite estimates of biomass change through time (Baccini et al., 2012), satellite-monitored fire activity and burn area estimates associated with deforestation (van der Werf et al., 2010). In addition, there is increasing use of ecosystem models coupled with remote sensing to estimate emissions from LUC (Galford et al., 2011).

Emissions associated with the LUC sector have the highest degree of uncertainty given the complexity of processes involving net emissions and Houghton et al. (2012) assessed this uncertainty at $\sim\!0.5\,\mathrm{Pg}\,\mathrm{C}\,\mathrm{yr}^{-1}$, which is of the same order of magnitude as the emissions themselves. Uncertainties in estimating GHG emissions arising from savanna clearing, associated debris burning and conversion to agriculture are greater than those for tropical forests (Fearnside et al., 2009). It is important to quantify the emissions and their uncertainties in savannas, particularly because tropical savanna woodland and grasslands occupy a large

area globally (27.6 million km²), greater than tropical forest (17.5 million km², Grace et al., 2006). Deforestation and associated fire from these biomes are the largest contributors to global LUC emissions (Le Quéré et al., 2014). Much of these GHG emissions are from the Brazilian Amazonia, an agricultural area that has been expanding since the 1990s. However, over the last decade, the rate of tropical forest deforestation in this region has decreased from $16\,000\,\mathrm{km}^2$ in the early 2000s to $\sim 6500\,\mathrm{km}^2$ by 2010 (Lapola et al., 2014), but at the expense of the Brazilian Cerrado, a vast savanna biome of some 2.04 million km², where clearing rates have been maintained (Ferreira et al., 2013, 2016; Galford et al., 2013). Given the suitability of the Cerrado topography and soils for mechanised agriculture, the Cerrado may become the principal region of LUC in Brazil (Lapola et al., 2014).

Northern Australia is one of the world's major tropical savanna regions, extending some 1.93 million km² across north-western Western Australia, the northern half of the Northern Territory and Queensland (Fisher and Edwards, 2015). This biome occupies approximately one quarter of the Australian continent and since European arrival, 5 % has been cleared for improved pasture, horticulture and cropping (Landsberg et al. 2011), making it one of least disturbed savanna regions in the world (Woinarski et al., 2007). However, this small percentage equates to a substantial area of 9.2 million ha, and LUC and associated economic development in northern Australia is a government imperative and this is likely to involve expansion and intensification of grazing, irrigated cropping, horticulture and forestry (Committee on Northern Australia, 2014). Drivers of this potential expansion in food and fibre production include the exploitation of the growing markets of Asia as well as domestic factors such as the perception that land and water resources of northern Australia can provide a future agricultural resource base to offset the expected declines in agricultural productivity in southern Australia due to adverse impacts of climate change (Steffan and Hughes, 2013).

Historically, intensive agricultural developments in northern Australia have been implemented based on limited scientific knowledge with dysfunctional policy and market settings, and as a result there has been limited success (Cook, 2009). Future expansion needs to be underpinned by sound understanding of the consequences of regional-scale land transformation on carbon and water budgets and GHG emissions. Any significant expansion in northern agricultural production would require clearance of native savanna vegetation, with unknown increases in GHG emissions. Most LUC studies occur at catchment, regional or biome scales (Houghton et al., 2012) and are not underpinned by good understanding of underlying processes. However, there are an increasing number of plot-scale studies using eddy covariance and chamber methods to provide direct measures of net GHG fluxes from contrasting land uses (Lambin et al., 2013). These studies typically compare microclimate and fluxes of GHGs from pastures and/or crops with adjacent forest ecosystems under a range of management conditions (e.g. Anthoni et al., 2004; Zona et al., 2013) or natural grasslands and different cropping types (e.g. Zenone et al., 2011). In tropical regions, there is a focus on transitions from forest to pasture and from forest to crops for food or bioenergy production (Galford et al., 2011; Wolf et al., 2011; Sakai et al., 2004).

There are few studies that directly measure GHG emissions and sinks prior to, during and after LUC at a single site. Land use change can involve rapid changes in net GHG emissions over varying temporal scales (minutes, hours and seasonal cycles) and continuous flux measurements are essential to capture the magnitude of these events (Hutley et al., 2005). However, there are no direct observations of emissions from savanna clearing in northern Australia, contributing to the uncertainty associated with the LUC sector in Australia's national GHG accounts (Commonwealth of Australia, 2015a).

Our objective is to provide a comprehensive assessment of GHG emissions associated with savanna clearing. Our aims are to (1) quantify the typical rates of CO₂ exchange of intact tropical savanna and make comparative measurements from an analogue site that was to be cleared, (2) quantify CO₂ fluxes before, during and after a clearing event, (3) estimate both CO₂ and non-CO₂ (CH₄ and N₂O) GHG emissions arising from burning of cleared debris and (4) quantify ecosystem-scale GHG balance for this land use conversion and compare it with emissions from savanna fire, a significant source of GHG emissions across northern Australia.

2 Methods

In this study we used a paired site approach, where concurrent fluxes of CO2, water vapour and energy were measured using eddy covariance towers from an uncleared savanna woodland site and a similar savanna woodland site on the same soil type that was to be cleared, burnt and prepared for agricultural production. Fluxes of CO₂ were monitored for 161 days prior to clearing at both sites with observations continuing during the clearing event (deforestation) and for another 507 days through phases of woody debris and grass curing, burning and soil preparation through raking and ploughing. The entire observation period was 668 days. Flux observations of net CO₂ exchange were combined with on-site biomass measurements and regionally calibrated pyrogenic emissions factors to estimate emissions of CO₂, CH₄ and N₂O (Meyer et al., 2012; Commonwealth of Australia, 2015b) from burning of the cleared debris that was a key component of the land conversion. Fire-derived emissions were combined with net CO2 fluxes from the land conversion phases to provide a total net emission in units of CO₂-e for this LUC. In this paper, we use the term "deforestation" to describe savanna clearing. Deforestation is defined under Australia's National Greenhouse Accounts as the loss of forest/woodland cover due to direct human-induced actions that fail to regenerate cover via natural regrowth or restoration planting (Commonwealth of Australia, 2015a).

2.1 Study sites

Both savanna woodland sites were located within the Douglas-Daly river catchment approximately 300 km south of Darwin, Northern Territory (Fig. 1). Both sites are OzFlux sites (www.ozflux.org.au), with flux observations ongoing at the uncleared savanna (UC) site since 2007 (Beringer et al., 2011, 2016a; Hutley et al., 2011). OzFlux is the regional Australian and New Zealand flux tower network that aims to provide continental-scale monitoring of CO₂ fluxes and surface energy balance to assess trends and improve predictions of Australia's terrestrial biosphere and climate (Beringer et al., 2016a). The UC site is broadly representative of Australian tropical savanna woodland found on deep, well drained sandy loam soils at sites with $\sim 1000 \, \mathrm{mm}$ MAP (Table 1). The cleared savanna site (CS) was carefully selected to ensure the vegetation and soils were as similar to the UC site as possible and with topography suitable for eddy covariance measurements.

Both sites were classified as savanna woodland (type 4B2, Aldrick and Robinson 1972, 1:50000 mapping) with an overstorey cover of 30 %, equivalent to the Eucalypt woodland Major Vegetation Group (MVG) of the National Vegetation Information System (NVIS, Commonwealth of Australia, 2003). The sites were dominated by an overstorey of Eucalyptus tetrodonta (F. Muell.), Corymbia latifolia (F. Muell.). Soils at both the UC and CS sites were red kandosols of the haplic mesotrophic great group (Isbell, 2002), characterised as deep, sandy loams (Table 1). The long-term mean annual precipitation (MAP) (\pm SD) at the UC site was estimated at 1180 ± 225 mm (1970–2012, Australian Water Availability Project (AWAP), www.csiro.au/awap), similar to the CS site at 1107 ± 342 mm (1985–2013, Bureau of Meteorology station, Tindal, NT). Slopes at both sites were < 2 % with a fetch of ~ 1.5 km at the UC site and ~ 1 km at the CS site. At both sites, 23 m guyed masts were installed to support eddy covariance instruments at 21.5 m above-ground. The tower at the CS site was moved 3 times to ensure adequate fetch was maintained according to seasonal wind direction during clearing and phases of the land use conversion. Instrument height was also adjusted given the height of the surface post-clearing and during the soil tillage phase (Table 2).

Satellite-derived burnt area mapping is available across northern Australia at 250 m resolution (North Australian Fire Information system, NAFI, www.firenorth.org.au) and indicated that fires had occurred within the flux footprint of the UC flux tower in 5 out of the last 13 years (2000–2013), whereas no fires had occurred within the footprint of the CS site. The average fire return time for the entire Australian savanna biome is 3.1 years (Beringer et al., 2015).

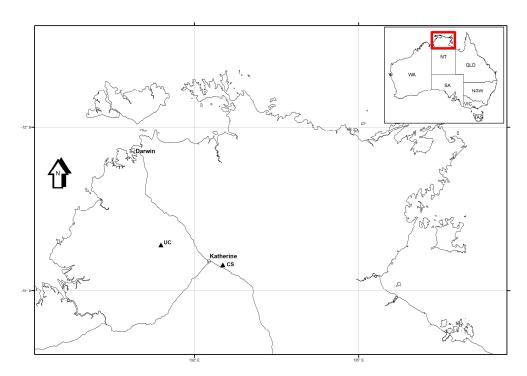


Figure 1. Location of the uncleared site (UC) and the cleared savanna (CS) sites south of Darwin, Northern Territory. The inset figure shows the distribution of the savanna biome across northern Australia as defined by Fox et al. (2001).

Table 1. Site characteristics for the uncleared savanna (UC) and cleared (CS) sites. Site soil orders are given as in Isbell (2002) with savanna vegetation classified using Fox et al. (2001). Fire frequency was estimated from fire mapping taken from the North Australian Fire Information system (NAFI, www.firenorth.org.au) for 2000–2012. The fire frequency estimate for the CS site excluded the debris fires in August 2012. Basal area and stem density is provided for all woody stems > 2 cm DBH at both sites. Mean site LAI for the UC is taken from Hutley et al. (2011) and for the CS site, was estimated from canopy hemispherical photos, see text for details.

Site	UC	CS
Location	14°09′33.12″ S, 131°23′17.16″ E	14°33′48.71″S, 132°28′39.47″ E
Soils	Red Kandosol	Red Kandosol
Vegetation type	Savanna woodland with mixed grasses	Savanna woodland with mixed grasses
	Map unit D4. E. tetrodonta, C. latifolia,	Map unit D4. E. tetrodonta, Ery-
	Terminalia grandiflora, Sorghum spp.,	throphleum chlorostachys, Corymbia.
	Heteropogon triticeus	bleeseri, Sorghum spp., H. triticeus
Map unit area (km ²)	59 986	59 986
Fire frequency (yr ⁻¹)	0.23	0.07
Basal area $(m^2 ha^{-1})$	8.3	6.8
Canopy height (m)	16.4	14.2
Above-ground biomass ($Mg C ha^{-1}$)	30.6 ± 9.2	26.2 ± 7.0
Stem density (ha ⁻¹)	330 ± 58	643 ± 102
Overstorey LAI (wet/dry)	NA/0.8	0.9/0.5
MODIS LAI (wet/dry)	1.5/0.9	1.6/1.0
MAP (mm)	1372 ^a /1180 ^b	1107 ^c
$\operatorname{Max} T_{\operatorname{air}} (^{\circ} C)$	37.5 (Oct)/31.2 (Jun)	37.5 (Oct)/29.7 (Jun)
$\operatorname{Min} T_{\operatorname{air}} (^{\circ} \mathbf{C})$	23.8 (Jan)/12.6 (Jul)	25.0 (Nov)/13.7 (Jul)

^a On-site observations, 2007–2012. ^b Gridded precipitation (AWAP, 1970–2012). ^c Tindal BoM station (14.52°S, 132.38° E, data from 1985 to 2013). NA stands for not available.

2.2 Flux measurements and data processing

Eddy covariance systems at both sites consisted of CSAT3 3-D ultrasonic anemometers (Campbell Scientific Inc., Logan, USA) and a LI-7500 open-path CO_2/H_2O analyser (Licor Inc., Lincoln, USA). Flux variables were sampled at 10 Hz and covariances were stored every 30 min. The LI-7500 gas analysers were calibrated at approximately 6-month interval for the duration of the data collection period and were highly stable. Mean daily rainfall, air temperature, relatively humidity, soil heat flux (F_g , W m⁻²) and volumetric soil moisture (θ_v , m³ m⁻³) from surface to 2.5 m depths were measured at both sites. The radiation balance was measured using a CNR4 net radiometer (F_n , W m⁻²) (Kipp and Zonen, Zurich).

Thirty minute covariances were stored using data loggers (CR3000, Campbell Scientific, Logan), and data postprocessing and quality control was undertaken using the OzFluxQC system as described by Isaac et al. (2016). In this system, data are processed through three levels: Level 1 is the raw data as collected by the data logger, Level 2 are qualitycontrolled data and Level 3 are post-processed and corrected but not gap-filled data. Quality control measures at Level 2 include checks for plausible value ranges, spike detection and removal, manual exclusion of date and time ranges and diagnostic checks for all quantities involved in the calculations to correct the fluxes. Quality checks make use of the diagnostic information provided by the sonic anemometer and the infrared gas analyser. Level 3 post-processing includes 2-dimensional coordinate rotation, low- and high-pass frequency correction, conversion of virtual heat flux to sensible heat flux $(F_h, W m^{-2})$ and application of the WPL correction to the latent heat $(F_e, W m^{-2} \text{ and } CO_2 \text{ fluxes } (F_c)$ (Isaac et al., 2016). Level 3 data also include the correction of the ground heat flux for storage in the layer above the heat flux plates (Mayocchi and Bristow, 1995).

Gap filling of meteorology and fluxes along with flux partitioning of net ecosystem exchange (NEE) into gross primary productivity (GPP) and ecosystem respiration (R_e) was performed on the Level 3 data using the Dynamic INtegrated Gap filling and partitioning for Ozflux (DINGO) system as described by Beringer et al. (2016b). In summary, DINGO gap fills meteorological variables (air temperature, specific humidity, wind speed and barometric pressure) using nearby Bureau of Meteorology (BoM, www.bom.gov.au) automatic weather stations that were correlated with tower observations. All radiation streams were gap-filled using a combination of MODIS albedo products (MOD09A1) and BoM gridded global solar radiation and gridded daily meteorology from the Australian Water Availability Project (AWAP) data set (Jones et al., 2009). Precipitation was gap-filled using either nearby BoM stations or AWAP data. Soil temperature and moisture were filled using the BIOS2 land surface model (Haverd et al., 2013) run for each site and forced with BoM or AWAP data. Energy balance closure was examined using standard plots of $F_{\rm h}+F_{\rm e}$ vs. $F_{\rm n}-F_{\rm g}$ using 30 min flux data from both sites (data not shown). For the CS site, closure was examined using data grouped according to the nine LUC phases as given in Table 2. For the UC site, all 30 min data from 2007 to 2015 were used.

Gap filling of fluxes was undertaken using DINGO, which uses an artificial neural network (ANN) model following Beringer et al. (2007). Model training uses gradient information in a truncated Newton algorithm. NEE and fluxes of sensible, latent and ground heat fluxes were modelled using the ANN with incoming solar radiation, VPD (vapour pressure deficit), soil moisture content, soil temperature, wind speed and MODIS EVI as inputs. The ustar threshold for each site was determined following Reichstein et al. (2005) and nighttime observations below the ustar threshold were replaced with ANN modelled values of Re using soil moisture content, soil temperature, air temperature and MODIS EVI as inputs. The ANN R_e model was then applied to daylight periods to estimate daytime respiration and GPP was calculated as the difference between NEE and R_e . For data collected at the CS site, a unique ANN model was developed for each LUC phase given the differing canopy and microclimatology of each phase. At each site, daily NEE, Re and GPP were calculated for each day of each phase.

2.3 Leaf area index

Canopy leaf area index (LAI) at the CS site in the surrounding intact savanna was measured using a 180° hemispherical lens (Nikon 10.5 mm, f/2.8) after Macfarlane et al. (2007). Three savanna transects were photographed seasonally on nine occasions over 2.1 years from the preclearing phase (October 2011) to December 2013. Along each 100 m transect, 11 hemispherical pictures were taken at 10 m intervals (33 photos for each measure occasion). At both sites the LAI was also estimated using MODIS Collection 5 LAI (MOD15A2) for a 1 km pixel around each tower. The 8-day product was interpolated to daily time series using a spline fit. Only MODIS values with a quality flag of 0 for FparLai_QC were used in the estimate, indicating the main algorithm that was used (http://lpdaac.usgs.gov/sites/default/files/public/modis/ docs/MODIS-LAI-FPAR-User-Guide.pdf).

2.4 Land use conversion

The specific sequence and timing of clearing, burning and land preparation phases is given in Table 2. Conversion of woodland to agricultural land in northern Australia is typically achieved by pulling trees over using large chains held under tension between two bulldozers. Clearing occurs at the end of the wet season when soil moisture is still high and soil strength low as under these conditions trees are easily pulled over, with a large fraction of the tree root mass extracted when pulled. At the CS site, 295 ha of savanna were

Table 2. Characteristics of land conversion phases during the 668-day observation period at the savanna clearing site (CS). Also given are the canopy heights following LUC phases and flux instrument heights that were adjusted following clearing, burning and then soil preparation phases.

Season	Period	LULUC phases	Canopy height (m)	Instrument height (m)
Late dry season	Sep-Oct 2011	Intact savanna	16	21.5
Wet season pre-clearing	Oct 2011-Feb 2012	Intact savanna	16	21.5
Wet season clearing	Mar–May 2012	Savanna deforested using bulldozers, followed by de- bris decomposition, understory grass germination	3	7
Dry season pre-burn	May-Aug 2012	Vegetation debris curing, understorey grass growth	2	7
Debris burning	Aug 2012	Debris and grasses burnt, soil ripped to 60 cm to remove roots, roots and remaining debris stockpiled, reburnt	2	7
Dry season post-burn	Aug-Nov 2012	Grass and shrubs germination and resprouting	1	7
Early wet season	Nov 2012–Jan 2013	Removal remaining below-ground biomass. Wet season rains stimulates grass growth, shrub resprouting and growth	1	7
Wet season	Jan-Mar 2013	All regenerated vegetation removed, soil bed preparation	0	3
Dry season	Apr-Jul 2013	Soil cultivation in stages	0	3

deforested between 2 and 6 March 2012 using this technique. A permit for this land conversion had been issued by the regional land management agency following an impact assessment and erosion control planning. Chains were under tension and intercepted tree boles at 0.1–0.2 m height above the ground, which assisted in pulling the trees and limited damage to the soil surface. As a result, grasses, woody resprouts and shrubs of the understorey remained largely intact following deforestation (Plate 1a). Mechanised ripping of soil to 60 cm depth was also undertaken to remove remaining coarse root material.

A cost-effective method of removing cleared vegetation is curing (drying) and subsequent burning and the land managers at the CS site left debris on site to for 5 months through the dry season (March to August, 2011). Burning of debris occurred over a 22-day period in the late dry season, August 2012 (Plate 1b), a period of consistent southerly trade winds of low relative humidity (10-20%, BoM, Tindal station, NT). Prior to ignition, 100 m fire breaks were installed around the entire 295 ha block and then lit in blocks of ~ 80 ha in size. There was an initial ignition of the fine and coarse fuels (grasses, litter and twigs, defined below) and woody debris (heavy fuels). Heavy fuels that were not completely consumed following the initial burn were then stock-piled in rows \sim 1–2 m in height and reignited until the fuel was consumed (Plate 1c). Inspection of debris post-fire suggested \sim 5 % of fine fuels remained as ash and \sim 10 % of the heavy fuels remained as charcoal, and these were subsequently incorporated into the top soil during soil bed preparation (Plate 1d).

2.5 GHG emissions from debris burning

Emissions of CO₂, CH₄ and N₂O from the debris burning were estimated following the approach as outlined in the IPCC Good Practice Guidelines (IPCC, 2003), which uses country or region-specific emission factors for fire activity (indicated by burnt area) and the mass of fuel pyrolysed to estimate the emission of each trace gas. This approach is well developed for the fire regime of the northern Australian savanna and is described by Russell-Smith et al. (2013) and Murphy et al. (2015a). These authors describe a novel GHG emissions abatement methodology for savanna burning that combines indigenous fire practices with an emissions accounting framework, the Emissions Abatement through Savanna Fire Management (Commonwealth of Australia, 2015b, www.comlaw.gov.au/Series/ F2013L01165). This methodology is a legislative instrument that establishes procedures for abatement projects for prescribed savanna burning and defines emission factors for four fuel classes; fine (grass and litter < 6 mm diameter fragments), coarse (6 mm-5 cm), heavy (>5 cm diameter) and shrubs fuels (Russell-Smith et al., 2013). Emissions of GHGs are estimated based on vegetation type, fuel mass per area for each fuel type, burn area, the burning efficiency (BEF) for each fuel type, defined as the mass of fuel exposed to fire that is pyrolysed, the fuel carbon content (%), elemental C: N ratios and emission factors (EFs) for each GHG (CO₂, CH₄ and N₂O) and global warming potentials for each gas. Across the northern Australian savanna, values for BEFs and EFs have been determined for both high (> 1000 mm MAP) and low precipitation zones (1000-600 mm MAP) and for both early and late dry season fires, which are fires occurring after 1 August which typically have higher intensity and

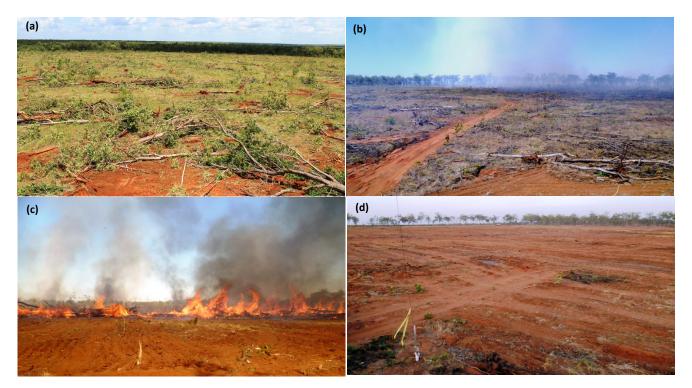


Plate 1. Key LUC phases associated with: (a) the clearing event, Phase 3; (b) debris burning of the cured grass, litter and woody fuels following the 5-month curing period, Phase 5; (c) stockpiling and ignition of remaining unburnt debris and (d) post-fire site preparation with all biomass consumed, Phase 9.

combustion efficiencies than early dry season fires (Russell-Smith et al., 2013).

We used these definitions of vegetation fuel type (wood-land savanna with mixed grass) and associated EFs, carbon contents and N:C ratio values as defined in the methodology to estimate GHG emissions from the debris fire using the following equation:

$$E = \sum_{i} (FL_{j} \times BEF_{j} \times CC_{j} \times N : C_{N_{2}O} \times EF_{i,j} \times GWP_{i}), \quad (1)$$

where E is the sum of emissions in Mg CO₂-e ha⁻¹ for each GHG i (CO₂, CH₄ and N₂O), FL $_j$ is the fuel load for fuel type j (fine, coarse, heavy) in Mg C ha⁻¹, BEF $_j$ is the burning efficiency factor, CC $_j$ is the fractional carbon content, N: CN₂O is the fuel nitrogen to carbon ratio for N₂O emissions, EF $_{i,j}$ is the emission factor for GHG i and fuel type j, and GWP $_i$ is the global warming potential for each GHG i (after Commonwealth of Australia, 2015b). The debris fire differed from a typical savanna fire in that there was a significantly higher heavy fuel load present and it was of high intensity which consumed the vast majority of fuel (Plate 1c, d), reflected in the assumed BEFs we used. The fire-derived emissions were combined with tower-derived NEE data from the post-clearing phases (Table 3) to give a total emission in CO₂-e for this LUC.

2.6 Quantifying fuel loads

To accurately quantify emissions from the debris fire, fine, coarse and heavy fuels were estimated using plots and transects established across the 295 ha deforestation area. For fine fuels, six 100 m transects were randomly located and at 20 m intervals along each transect, all fine (grass, woody litter) and coarse (twigs, sticks) fuels were harvested from 1 m² quadrats, dried and weighed to give a mean fine and coarse fuel mass for the site. We assigned on-site coarse woody debris (CWD), above-ground and below-ground biomass estimates to the heavy fuel class (> 5 cm diameter fragments). To quantify CWD, an additional six 100 m transects were randomly located across the deforestation area and along each transect the length and diameter of all intersected CWD fragments were recorded to estimate fragment volume. In these savannas, large fragments (> 10 cm diameter) are frequently hollowed from the action of termites and fire and the diameter and length of the annulus of such fragments were measured to estimate this missing volume. In addition, large fragments that were tapered were treated as a frustum of a cone and a second diameter was taken at the fragment end to improve volume estimation. Fragment volumes were calculated and converted to mass using rot classes (RCs) and associated wood densities (g cm⁻³). Five rot classes (RCs) were defined and assigned to each CWD fragment to cap-

Table 3. Cumulative precipitation and mean NEE, R_e and GPP (Mg C ha⁻¹ month⁻¹) for each of the LUC phases at the CS site as measured by the flux tower. These fluxes are given for the UC site for these same periods. One-way ANOVA was used to test for differences between mean daily NEE for each LUC phase with significantly different means labelled with an asterisk. On the days of ignition during the debris burning phase, flux data at the CS site were excluded. Integrated fluxes are given for the post-clearing period (507 days) and the entire observation period (668 days) for both sites in Mg C ha⁻¹.

				CS				UC		
	Phase	Period	Rainfall	NEE	$R_{\rm e}$	GPP	Rainfall	NEE	$R_{\rm e}$	GPP
LULUC phases	number	(d)	(mm)	(Mg C l	na^{-1} mo	nth^{-1})	(mm)	(Mg C ha	a^{-1} mor	nth^{-1})
Intact canopy cover	1	161	736.6	-0.23	1.57	-1.79	1076.8	-0.25	1.45	-1.70
Clearing event	2	4	59.4	0.23*	1.95	-1.73	59.8	0.38*	1.80	-1.50
Wet-dry debris curing, decomposition	3	59	143.2	0.98**	1.39	-0.41	412.0	0.32**	1.53	-1.22
Dry season pre-burn	4	94	0	0.34**	0.57	-0.23	2.4	0.15**	0.94	-0.79
Fire emissions late dry	5	22	0	0.90**	0.76	0.0	0.0	-0.01**	0.71	-0.72
Dry season post-burn	6	67	2.2	0.31**	0.37	-0.06	64.4	-0.28	0.64	-0.91
Early wet regrowth	7	80	361.0	0.03**	0.99	-0.96	345.8	-0.32	1.80	-2.12
Wet season site prep.	8	91	701.7	0.62**	0.99	-0.37	914.4	-0.20**	1.67	-1.88
Dry season final bed prep. and cultivation	9	90	0	0.29**	0.32	-0.02	10.8	0.06**	0.91	-0.85
				(M	Ig C ha	-1)		(Mg	g C ha ⁻¹	¹)
Total post-clearing		507	1267.5	7.2**	12.8	-5.6	1809.6	-0.78**	20.7	-21.5
Total all phases		668	2004.1	6.0**	21.2	-15.2	2886.4	-2.1**	28.5	-30.6

^{*} Denotes significantly different mean NEE at the 5 % level, ** significant at 1 %.

ture the decay gradient of fragments. These were defined as recently fallen, solid wood (RC1), solid wood with or without branches present but with signs of aging (RC2), obvious signs of weathering, still solid wood, bark may or may not be present (RC3), signs of decay with the wood sloughed and friable (RC4) and severely decayed fragments with little structural integrity remaining (RC5). A wood density was assigned to each RC and species (where identifiable) after Rose (2006) and Brown (1997) to provide an accurate estimate of CWD mass that included decay and hollowing. For the dominant *Eucalyptus* and *Corymbia* species wood densities ranged from 0.7 g cm⁻³ (RC1) to 0.56 g cm⁻³ (RC5).

Above-ground biomass was quantified by surveying all woody plants > 1.5 m in height or > 2 cm DBH across eight 50×50 m plots. All woody individuals were identified to species and stem diameter at 1.3 m height (DBH) and tree height were measured. Region-specific allometric equations are available for tree species found at the CS site (Williams et al., 2005) and these were used to estimate above-ground biomass for each individual tree and shrub based on DBH and height. Below-ground biomass was calculated using the root / shoot ratio estimate of Eamus et al. (2002) for these savanna stands, which was 0.38. These trees have large lateral roots in the top 30 cm of soil, with no tap root and 90 % of root biomass is found in the top 50 cm of soil (Eamus et al., 2002). As such, we assumed that chaining and bulldozer clearing of all above-ground biomass followed by soil ripping (ploughing) to 60 cm soil depth, plus mechanised removal of root biomass associated with tree boles and subsequent burning, resulted in a near-complete removal of both above- and below-ground woody biomass pools (Plate 1d).

2.7 Deforestation and savanna burning emissions at catchment to regional scales

The potential impact of any expanded deforestation across the northern Australian savanna landscapes was assessed relative to historic deforestation rates and resultant GHG emissions and arising from prescribed savanna burning. This land management activity contributes $\sim 3\%$ to Australia's national GHG emissions (Whitehead et al., 2014) and is 25 % of the Northern Territory's annual emissions (Commonwealth of Australia, 2015a). Annual emissions from these activities (historic and future savanna deforestation and prescribed burning) were estimated at three spatial scales: (1) catchment, (2) state/territory and (3) regional. Emissions estimates from deforestation and savanna burning were compiled for (1) the Douglas–Daly river catchment where the UC and CS sites are located (area 57 571 km²), a catchment with less than 5 % of the native vegetation deforested to date (Lawes et al., 2015) but earmarked for future development; (2) the savanna area of Northern Territory (856 000 km²) and (3) the savanna region of northern Australia as defined by Fox et al. (2001) with MAP>600 mm, an area of 1.93 million km² (Fig. 1, insert).

Emissions of GHGs from historic deforestation from the Douglas–Daly catchment were estimated using our estimates for savanna land conversion combined with satellite-derived annual deforestation area (1990–2013) as reported by Lawes et al. (2015) for this catchment to give a catchment-scale mean annual estimate of emissions from deforestation in $Gg CO_2$ -e yr^{-1} . Annual deforestation emissions data for the Northern Territory and the northern Australian savanna re-

gion were taken from the National Greenhouse Gas Inventory (NGGI) for the same period 1990-2013. The Department of Environment is responsible for reporting sources of greenhouse gas emissions and removals by sinks in accordance with UNFCCC Reporting Guidelines on Annual Inventories and the supplementary reporting requirements under the Kyoto Protocol. State and Territory GHG Inventories are reported for 1990 to 2013 (Commonwealth of Australia, 2015a) and we used data for the Land Use, Land-Use Change and Forestry sector, Activity A.2 Deforestation. These emissions are reported for each state, but are not biome based and for our regional savanna estimate, emissions data for Western Australia, the Northern Territory and Queensland were used but were calculated using the area within each state that was defined as savanna by Fox et al. (2001, Fig. 1). Mean annual deforestation emissions from the savanna area of each state and territory (1990-2013) were summed to calculate a mean (±SD) annual deforestation rate for the northern Australian savanna area (1.92 million km²) in $Gg CO_2$ -e yr⁻¹.

Emissions from savanna burning were calculated using the online Savanna Burning Abatement Tool (SAVBat2, www. savbat2.net.au) using the predefined vegetation fuel types (VFTs) mapping for the northern Australian savanna (Fisher and Edwards, 2015; Thackway, 2014), both components of the Emissions Abatement through Savanna Fire Management methodology. SAVBat2 combines satellite-derived burnt area mapping (www.firenorth.org.au) with fuel load estimates from VFT mapping, GHG emission factors and burn efficiencies to estimate annual emissions from burn areas. In accordance with IPCC accounting rules, only non-CO₂ emissions are reported for savanna burning as it is assumed that CO₂ emissions from dry season burning is offset by regrowth of vegetation (mostly C₄ grasses) in subsequent wet season(s) (IPCC, 1997). However, for comparisons with deforestation emissions, we calculated emissions of CO2 as well as non-CO₂ emissions. SAVBat2 estimates were compiled for the same areas as savanna deforestation estimates; the Douglas-Daly river catchment, savanna of the NT and the northern Australian savanna. Mean annual burning emissions for 1990-2013 were calculated and are reported as non-CO₂ (CH₄, N₂O) and total emissions (CO₂, CH₄ and N₂O) in $Gg CO_2$ -e yr^{-1} .

2.8 Emissions from expanded deforestation across northern Australia

Emissions from expanded deforestation across northern Australia was estimated by upscaling our estimate of deforestation emissions per hectare from catchment areas identified as having future clearing potential. These areas were based on the land use assessment of northern Australian catchments by Petheram et al. (2014) and identified catchments with development potential based upon surface water storage and proximity of land resources suitable for irrigation development for agriculture, horticulture or improved pastures.

Using these criteria, suitable catchments were identified in Western Australia (Fitzroy River, Ord Stage 3; 75 000 ha potential area), the Northern Territory (Victoria, Roper Rivers, Ord Stage 3, Darwin-Wildman River area; 114 500 ha) and Queensland (Archer, Wenlock, Normanby, Mitchel Rivers; 120 000 ha). This gives a potential savanna deforestation area of 311 000 ha, equivalent to an additional 16 % of cleared land over and above the 1886512 ha that has been cleared across the savanna biome since 1990 (Commonwealth of Australia, 2015a). Projected emissions included mean annual emissions from historic deforestation rates plus emissions from this expanded deforestation scenario. Expanded deforestation areas were calculated assuming any such clearing would occur over a 5-year period and are reported as non-CO₂ (CH₄, N₂O) and total emissions (CO₂, CH₄ and N₂O) in $Gg CO_2$ -e yr^{-1} .

3 Results

3.1 Pre-clearing site comparisons

Pre-clearing meteorology, flux observations and energy balance closure for UC and CS sites were compared (Fig. 2). Mean monthly NEE, R_e and GPP for each LUC phase for both sites are given in Table 3. Flux measurements prior to clearing were made for 161 days, a period spanning the late dry to early wet season transition (September–December) through to the middle of the wet season (January-February, Table 2). Flux data at the CS site were validated by assessing energy balance closure, with a regression between energy balance components suggesting closure was high with a slope of 0.91 and an R^2 of 0.95 (n = 4778). Site differences for each phase were tested using one-way ANOVA using daily mean NEE with days as replicates. For Phase 1, mean daily NEE was not significantly different between the two sites during (P < 0.64, df = 321). Seasonal patterns of T_{air} , VPD (Fig. 2b), LAI (Fig. 2c) and C fluxes (NEE, GPP, R_e , Fig. 2d) were similar when both sites were intact, although precipitation was 340 mm higher at the UC site (Table 3).

At both sites, NEE shifted from being a weak sink of less than $-1 \,\mu \text{mol CO}_2 \,\text{m}^{-2} \,\text{s}^{-1}$ during the late dry season to a net source of CO₂ during the early wet season (Fig. 2d). During this period, R_e increased rapidly from +2 to +5 μ mol m⁻² s⁻¹ in early October with the onset of wet season rain, but then remained relatively constant for the remainder of the wet season. As the wet season progressed, temporal patterns of GPP were similar at both sites, then steadily increased to -6 to -7 μ mol m⁻² s⁻¹ and remained at this rate until they cleared (March 2012). R_e was relatively stable during this period and NEE increased to -2 μ mol m⁻² s⁻¹ through the wet season (December to February). Despite the higher precipitation received at the UC site, mean monthly NEE, GPP and R_e differed by <10 % (Table 3, intact canopy phase). Normalising fluxes by MODIS LAI for each site fur-

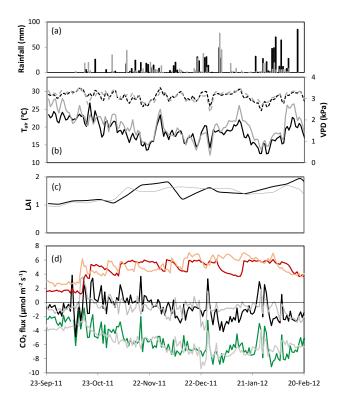


Figure 2. Comparative meteorology and fluxes for the uncleared (UC) and cleared savanna CS sites prior to the clearing event. Data spans the late dry season (September 2011) through to the middle of the wet season prior to the clearing event of 2–6 March 2012. Plots include (a) daily precipitation (black bars UC site, grey bars CS site), mean daily $T_{\rm air}$ (black lines UC, grey CS), (b) mean daily VPD (dashed lines; black UC, grey CS), (c) interpolated 8-day MODIS LAI (black UC, grey CS), (d) NEE (black UC, grey CS) partitioned into $R_{\rm e}$ (red UC, pink CS) and GPP (dark green UC, pale green CS).

ther reduced differences to 2% (data not shown), suggesting site differences were small and the UC site provides a suitable control for the CS site.

3.2 Fluxes following clearing

Clearing of the 295 ha block commenced on 2 March 2012 and the bulldozers reached the footprint of the flux tower at \sim 09:00 h local time on 6 March (Fig. 3). As for Phase 1, energy balance closure of flux tower data for LUC Phases 2 to 4 (post-clearing phases) was high, with a slope > 0.9 and R^2 > 0.92. Over all phases at the CS site, closure was lower, with a slope of 0.81 (R^2 = 0.95, n = 26 395), similar to that of the UC site at 0.87 (R^2 = 0.93, n = 99 998).

The 4-day clearing event occurred during relatively high soil moisture conditions, with surface (5 cm depth) $\theta_{\rm v}$ ranging from 0.08 to 0.10 m³ m⁻³ and subsoil $\theta_{\rm v}$ (50 cm depth) ranging from 0.12 to 0.14 m³ m⁻³. As a result, pre-clearing fluxes were high and NEE reached

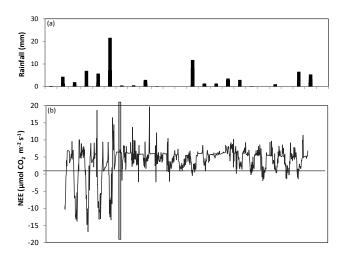


Figure 3. (a) Daily precipitation and **(b)** diurnal patterns of NEE at the CS site for the week prior to the clearing event of 2–6 March 2012 (vertical bar) and 3 weeks post-clearing.

 $-15\,\mu mol\,CO_2\,m^{-2}\,s^{-1}$ during the middle of the day (Fig. 3). Mean daily NEE for the week prior to clearing was a net CO₂ sink of $-0.60\pm0.63\,\mu mol\,m^{-2}\,s^{-1}$ and was not significantly different to mean daily NEE at the UC site of $-0.80\pm0.93\,\mu mol\,m^{-2}\,s^{-1}$ (ANOVA, P<0.03). For the 3 weeks following clearing, the CS site rapidly became a net source of CO₂ with a mean daily NEE of $+4.38\pm0.24\,\mu mol\,m^{-2}\,s^{-1}$, with a much reduced diurnal amplitude and no response to precipitation events (Fig. 3a, b). High closure (slope > 0.9) was observed during Phases 2 to 4, although this was reduced (slope = 0.75) for the post-fire and soil preparation, Phases 6–9.

Table 3 provides values of precipitation and monthly NEE, $R_{\rm e}$ and GPP for the seven LUC phases following clearing, namely debris decomposition and curing (153 days), burning (22 days), wet season regrowth (80 days), followed by soil tillage and preparation of irrigated raised soil beds (181 days). For each phase, the comparable flux estimate from the UC site is estimated for all post-clearing phases and for the entire observation period. Following clearing, GPP at the CS site was reduced by a factor of 3.5 when compared to the UC for the same period (March 2012–January 2013, Table 3). While greatly reduced, GPP still occurred at the CS site during this 13.7-month period $(-0.38 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}\,\mathrm{month}^{-1})$ via resprouting of felled overstorey and subdominant trees and shrubs, as well as grass germination and growth stimulated by early wet season precipitation (November 2012–January 2013, 361 mm, Table 3). Ecosystem respiration during this period was higher at the UC site $(+1.12 \,\mathrm{Mg} \,\mathrm{C} \,\mathrm{ha}^{-1} \,\mathrm{month}^{-1})$ when compared to the CS site $(+0.82 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}\,\mathrm{month}^{-1})$ and, given the large decline in GPP, the CS site was a small net C source at $+0.51 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}\,\mathrm{month}^{-1}$, compared to the UC site which was a weak sink of $-0.03 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}\,\mathrm{month}^{-1}$.

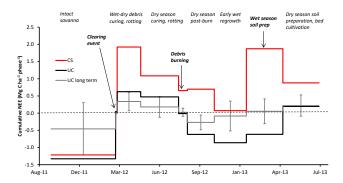


Figure 4. Cumulative NEE from the CS (red line) and UC sites (black line) for each land use phase (see Table 2 for details) over the entire observational period, September 2011 to July 2013. The UC site is a long-term savanna site of the Australian flux network (OzFlux, see Beringer et al., 2016a) and using the sites' 8-year flux record (2007–2013), the long-term cumulative mean NEE is plotted for each land use phase of (grey line; $\pm 95\%$ CI). The dashed line indicates zero net CO₂ flux.

Cumulative NEE over all the post-clearing LUC phases was $+7.2 \,\mathrm{Mg} \,\mathrm{Cha}^{-1}$ at the CS site compared to a net sink of $-0.78 \,\mathrm{Mg} \,\mathrm{C} \,\mathrm{ha}^{-1}$ at the UC site (Table 3). The temporal dynamics of cumulative NEE across all LUC phases (note differences in phase duration) is summarised in Fig. 4, which compares fluxes from both sites for the complete observation period. Three significant periods of C emission are evident in Fig. 4. Firstly, the clearing event and the subsequent switch from a C sink to a net source of 1.9 Mg C ha⁻¹ due to soil disturbance and the decomposition of biomass. Secondly, this was followed by a reduction in source strength over the dry season of 2012, attributable to a reduction in R_e during the dry season (2012 dry season pre-burn phase, Table 3). Thirdly, there were other major emissions attributed to soil tillage and bed preparation in the wet and dry seasons of 2013, a cumulative net emission of $+2.75 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}$ that occurred over the final 6 months (Fig. 4) in preparation for cropping. Over this phase, the UC site was a net sink of $-0.62 \,\mathrm{Mg}\,\mathrm{C}\,\mathrm{ha}^{-1}$.

3.3 Emissions from debris burning

Table 4 gives fuels loads, BEFs, EFs, carbon content and N:C ratios for each fuel type used to estimate the GHG emission from the debris burning. Fuel load was dominated by heavy fuels with a mean (\pm SD) above-ground biomass of $26.9\pm7.0\,\mathrm{Mg\,C\,ha^{-1}}$ and a range of 14.4 to $39.3\,\mathrm{Mg\,C\,ha^{-1}}$ across the eight biomass plots. The mean below-ground biomass was estimated at $9.0\pm2.4\,\mathrm{Mg\,C\,ha^{-1}}$ and CWD was $1.4\pm0.6\,\mathrm{Mg\,C\,ha^{-1}}$. Fine and coarse fuels were 1.4 ± 0.7 and $0.5\pm1.0\,\mathrm{Mg\,C\,ha^{-1}}$. Using these fuel loads with savanna EFs and the BEFs estimated for the site gave emissions of CO₂, CH₄ and N₂O for each fuel type and the

emission from debris burning totalled 121.9 Mg CO_2 -e ha⁻¹, with 9.5 % of this total comprising non- CO_2 emissions (Table 4).

3.4 Total GHG emission

Emissions derived from debris burning need to be combined with the post-clearing NEE as measured by the EC system to provide a total GHG emissions estimate from this LUC in units of CO₂-e. The LUC phases following clearing spanned a 502-day period (Table 3), and NEE was +7.2 Mg C ha⁻¹ or +26.4 Mg CO₂-e ha⁻¹. In comparison, NEE from the UC site over the same period was -0.78 Mg C ha⁻¹ or -2.9 CO₂-e ha⁻¹. Adding NEE from post-clearing phases (Phases 2–9, Table 3) to emissions from debris burning (Table 4) gave a total emission of +148.3 Mg CO₂-e ha⁻¹ for the CS site. The CO₂-only emission from debris burning plus post-clearing NEE was +136.7 Mg CO₂ ha⁻¹, which was a flux 45 times larger than the observed savanna CO₂ sink at the UC site over the post-clearing period.

3.5 Upscaled and projected emissions from deforestation and savanna burning

Table 5 provides mean (±SD) GHG emissions estimates for savanna burning and deforestation for 1990-2013. At all spatial scales, annual mean burnt area dwarfed the mean annual land area deforested. For the Douglas-Daly catchment area, reported non-CO₂ emissions from savanna burning were $577 \pm 124 \,\mathrm{Gg}\,\mathrm{CO}_2$ -e yr⁻¹, almost 4 times larger than emissions from the mean annual savanna deforestation rate of $163 \pm 162 \,\mathrm{Gg}\,\mathrm{CO}_2$ -e yr⁻¹. For the Northern Territory savanna, mean annual burning emissions were an order of magnitude larger than mean annual deforestation emissions (Table 4) and 2 orders of magnitude larger if CO₂ emissions were included. At a regional scale, the mean annual deforestation rate across the northern Australian savanna was $16161 \pm 5601 \,\mathrm{Gg}\,\mathrm{CO}_2$ -e yr⁻¹, with emissions from Queensland savanna area dominating this amount at $15762 \pm 5566 \,\mathrm{Gg}\,\mathrm{CO}_2\,\mathrm{yr}^{-1}$. This is double that of the reported (non-CO₂ only) emission from prescribed burning at $6740 \pm 1740 \,\mathrm{Gg} \,\mathrm{CO}_2$ -e yr⁻¹ (Table 5).

Emissions estimates that include future deforestation rates would be equivalent to savanna burning, at least for the duration of the additional deforestation. For the Douglas–Daly catchment, this future emission is estimated at 756 Gg CO₂-e yr $^{-1}$ and across the Northern Territory savanna area, this would be 3413 Gg CO₂-e yr $^{-1}$, rates of emission that are equivalent to burning emissions catchment (Douglas–Daly, 577 ± 124) and state scales (Northern Territory savanna, 3490 ± 922 Gg CO₂-e yr $^{-1}$). Emissions that include future deforestation rates for the northern Australian savanna region were estimated at 24 393 Gg CO₂-e yr $^{-1}$ and would be 3 times the reported savanna-burning annual emissions (Table 5).

Table 4. Measured fuel loads, assumed burning efficiencies (BEFs), carbon contents, N: C ratio and emissions factors (EFs) used to estimate GHG emissions from the burning of the post-deforestation fine, coarse and heavy fuel debris. Emission factors, carbon content and C: N ratio were assumed for the vegetation fuel type woodland savanna with mixed grass (code hWMi) as given in the Emissions Abatement through the Savanna Fire Management methodology (Commonwealth of Australia, 2015b), available at www.legislation.gov.au/Details/F2015L00344 and Meyers et al. (2012).

Fuel type	Fuel load	BEF	Carbon	N:C	EF CO ₂	EF CH ₄	EF N ₂ O			g CO ₂ -e	
	$(Mg C ha^{-1})$		content	ratio				CO_2	CH ₄	N_2O	Total
Fine	1.1 ± 0.70	0.95	0.46	0.0096	0.97	0.0031	0.0075	3.9	0.1	0.04	4.0
Coarse	0.5 ± 1.0	0.9	0.46	0.0081	0.92	0.0031	0.0075	1.5	0.0	0.01	1.6
Heavy – AGB	26.2 ± 7.0	0.9	0.46	0.0081	0.87	0.01	0.0036	75.2	7.9	0.32	83.4
Heavy - CWD	1.4 ± 0.6	0.9	0.46	0.0081	0.87	0.01	0.0036	4.0	2.7	0.11	28.5
Heavy – BGB	9.0 ± 2.4	0.9	0.46	0.0081	0.87	0.01	0.0036	25.7	0.0	0.02	4.4
Total								110.2	11.1	0.50	121.9

4 Discussion

Australia has lost approximately 40 % of its native forest and woodland since colonisation (Bradshaw, 2012), with most of this clearing for primary production in the eastern and south-eastern coastal region. Attention has now turned to the productivity potential of the largely intact northern savanna landscapes, which will involve trade-offs between management of land and water resources for primary production and biodiversity conservation (Adams and Pressey, 2014; Grundy et al., 2016). Globally and in Australia, savanna fire ecology and fire-derived GHG emissions have been reasonably well researched (Beringer et al., 1995; Cook and Meyer, 2009; Livesley et al., 2011; Meyer et al., 2012; Walsh et al., 2014; van der Werf et al., 2010) and the impacts of fire on the functional ecology of the Australian savanna has been recently reviewed by Beringer et al. (2015). In this study, we focussed on savanna deforestation and land preparation for agricultural use. These phases result in a series of events that may lead to pulsed GHG emissions that would otherwise be missed or greatly underestimated by episodic measurements taken at a weekly or monthly frequency after an initial treefelling event (Neill et al., 2006; Weitz et al., 1998).

We used the eddy covariance methodology as it provides a direct and non-destructive measurement of the net exchange of CO₂ and other GHG gases at high temporal resolution, ranging from 30 min intervals to daily, monthly, seasonal and annual estimates. The method is useful as a full carbon accounting tool as all exchanges of CO₂ from autotrophic and heterotrophic components of the ecosystem undergoing change are quantified (Hutley et al., 2005). This approach provides essential data for bottom-up GHG and carbon accounting studies as micrometeorological conditions and associated fluxes can be tracked through time for the duration of a land use conversion.

At the CS site, burning of post-clearing debris comprised 82 % of the total emission of 148.4 Mg $\rm CO_2$ -e ha⁻¹, with the remainder attributed to NEE as measured by the flux tower. This flux comprised significant $\rm CO_2$ losses via respiration of

debris, enhanced soil CO2 efflux from soil disturbance and tillage, which was partially offset by net uptake of CO₂ from woody resprouting post-clearing and periods of grass growth following wet season rainfall (Fig. 4). Soil disturbance via ripping, tillage and preparation was responsible for 10 % of the CO₂ emission from the conversion. The EC flux tower was in operation during the clearing event, demonstrating the utility of this method as the switch of the ecosystem from being a net CO₂ sink to being a net source. This occurred over a number of hours as the clearing event was completed (Fig. 3). During the LUC phase changes, there was little evidence of major pulses of CO2 flux, instead there was a rapid transition to a new diurnal pattern following the clearing (Fig. 3) or the commencement of soil preparation (data not shown). This is in contrast to non-CO₂ flux emissions, in particular N₂O, with short-term emissions often following disturbances (Grover et al., 2012; Zona et al., 2013) and can account for a significant fraction of annual emissions.

The net CO₂ source measured by the flux tower represents an emission that would be missed if vegetation biomass density alone was used to estimate LUC emissions, which is the approach currently used in remote sensing LUC studies at regional and continental scales. The total GHG emission we report in this study is more accurately described as a land conversion, as it includes the oxidation of biomass plus emissions associated with soil disturbance and tillage required for a conversion to a cropping or grazing system.

The emission estimate from this study does not include non-CO₂ soil-derived fluxes of CH₄ and N₂O, which can be significant for LUC events in certain ecosystems (Tian et al., 2015). Grover et al. (2012) compared soil CO₂ and non-CO₂ fluxes from native savanna with young pasture and old pastures (5–7 and 25–30 years old) in the Douglas–Daly river catchment. Soil emissions of CO₂-e were 30 % greater on the pasture sites compared with native savanna sites, with this change being dominated by increases in CO₂ emission and soil CH₄ exchange shifting from a small net sink to a small net source at the pasture sites. Non-CO₂ soil fluxes were generally small, especially N₂O emissions, although these mea-

area) and regional scales (northern Australian savanna area, Fig. 1). For savanna burning, burnt area and associated mean annual emissions (±SD) are given for both reported non-CO₂ (CH₄, N₂O) and total emissions (CO₂, CH₄ and N₂O). For the identical areas as used for savanna burning, mean annual GHG emissions from deforestation (±SD) are given. For the Douglas–Daly river catchment, deforestation area was taken from Lawes et al. (2015) and combined with deforestation emissions from the CS site. Deforestation emissions (1990–2013) for the NT and the northern Australian savanna area are taken from the State and Territory Greenhouse Gas Inventories (Commonwealth of Australia, 2015a). In bold text the emissions associated with the current deforestation rate plus expanded deforestation areas as identified by Petheram et al. (2014), which are combined with emissions from the able 5. Greenhouse gas emissions for 1990–2013 from prescribed savanna burning and savanna deforestation at catchment (Douglas–Daly rivers), state/territory (Northern Territory CS site to give an upscaled estimate of potential emissions with agricultural development at the three spatial scales.

Savanna region		Savanna burning			Savaı	Savanna deforestation	
	Burnt area a (ha yr $^{-1}$)	Emissions non- CO_2^a I $(GgCO_2-e yr^{-1})$ (Emissions total ^a Deforestati (Gg CO ₂ -e yr ⁻¹) (ha yr ⁻¹)	Deforestation area $(ha ext{ yr}^{-1})$	Emissions total Expanded $GgCO_2$ -e yr ⁻¹) area ^d (ha)	missions non-CO ₂ Emissions total ^a Deforestation area Emissions total Expanded deforestation Expanded emissions $(\operatorname{Gg} \operatorname{CO}_2 \operatorname{-e} \operatorname{yr}^{-1})$ $(\operatorname{Ida} \operatorname{yr}^{-1})$ $(\operatorname{Ida} \operatorname{yr}^{-1})$ $(\operatorname{Ida} \operatorname{yr}^{-1})$ $(\operatorname{Ida} \operatorname{CO}_2 \operatorname{-e} \operatorname{yr}^{-1})$ area ^d (ha) total ^d ($\operatorname{Ida} \operatorname{CO}_2 \operatorname{-e} \operatorname{yr}^{-1}$	Expanded emissions total ^d (Gg CO_2 -e yr ⁻¹)
Douglas-Daly river catchment 2482100 ± 490400 Northern Territory 13419410 ± 487300 Northern Australian 32249254 ± 11176	2 482 100 ± 490 400 13 419 410 ± 487 300 32 249 254 ± 11 176 004	577 ± 124 3490 ± 922 6740 ± 1729	$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	$163 \pm 162^{\text{b}}$ $398 \pm 128^{\text{c}}$ $16 \ 161 \pm 5601^{\text{c}}$	20 000 114 500 311 000	756 3413 24393

Burnt area and emissions data estimated using the on-line Savanna Burning Abatement Tool (SAVBat2), 1990–2013. These emissions are CH₄ and N₂O only. ^b Deforestation area data taken from Lawes et al. (2015), upscaled using the emissions from the CS sie from this study, 1990–2013. [©] Deforestation area and emissions data taken from the State and Territory Greenhouse Gas Inventories (Commonwealth of Australia, 2015a), 1990–2013. ^d Expanded deforestation area and at taken from catchments as identified surements were made many years after the LUC event and there is uncertainty as to their relevance for a recently deforested and converted savanna site. An additional pathway for CH_4 and N_2O emissions in these savannas is via termite activity (Jamali et al., 2011a, b). In our study, termite mounds were abundant across the CS site but were largely destroyed by clearing and soil preparation, potentially reducing the net non- CO_2 emission following conversion. Further work is required to quantify these non- CO_2 fluxes not associated with debris burning to refine our total emission estimate for savanna deforestation.

This land conversion represents the loss of decades of carbon accumulation in these mesic savanna (> 1000 mm MAP), ecosystems which are currently thought to be a weak carbon sink (Beringer et al., 2015). The 8-year ensemble mean NEE for the UC site was $-0.11\pm0.16\,\mathrm{Mg\,C\,ha^{-1}\,yr^{-1}}$ and is representative of a savanna site at a near-equilibrium state in terms of carbon balance given the low fire frequency (3 in 13 years, Table 1) with high severity fires uncommon (1 in 8 years of flux measurements). The annual increase in tree biomass at this UC site is $0.6\,\mathrm{t\,C\,ha^{-1}\,yr^{-1}}$ (Rudge, Hutley, Beringer, unpublished data) and, given an above-ground standing biomass of $28\,\mathrm{t\,C\,ha^{-1}}$, suggests a regeneration period of approximately four to five decades after stand replacement disturbance event such as deforestation for this savanna type.

Even after the large pool of carbon is lost following oxidation of biomass, carbon loss may continue on cleared land via continued soil carbon mineralisation, leading to a slow decline in soil carbon storage that is frequently reported for forest to cropping LUC (Jarecki and Lal, 2003; Lal and Follett, 2009). Conversion of forest or woodland to improved pasture grazing may result in either increases or decreases in soil carbon (Sanderman et al., 2010). Alternatively, it is possible that carbon sequestration may occur post-clearing via woody regrowth if a cleared site is abandoned and not further prepared for cultivation. This has actually been a relatively common transition and a significant sequestration pathway that needs to be included in savanna LUC assessments (Henry et al., 2015). Admittedly, if savanna-cleared land does fully transition to a cropping system, some fraction of the lost carbon could also be replaced or sequestered by new horticultural or forestry land uses.

There are few detailed, plot-scale studies of GHG emissions from savanna clearing in northern Australia. Several studies (Law and Garnett, 2009, 2011) used the Full Carbon Accounting Model (FullCAM Ver 3.0, Commonwealth of Australia, 2015a; Richards and Evans, 2004) to generate spatial maps of above- and below-ground biomass and soil organic carbon pools across the NT. The FullCAM model uses spatial and temporal soil, climate, precipitation data with NVIS major vegetation classes to simulate carbon losses (as GHG emissions) and uptake between the terrestrial biological system and the atmosphere. Land use change scenarios can be run within the model and Law and Gar-

nett (2009) examined deforestation emissions from the Eucalypt woodland NVIS vegetation class, as per UC and CS site classification. Modelled emissions were $136 \pm 42 \,\mathrm{Mg}\,\mathrm{CO}_2$ -e, comparable to our deforestation estimate of 121.4 Mg CO₂e. Henry et al. (2015) used a life cycle assessment approach to quantify GHG emissions from LUC associated with beef production in eastern Australia. Australia's major beef-producing areas across central and southern Queensland and northern central New South Wales were classified into 11 bioregions, with the northernmost bioregion, the northern Brigalow Belt, falling within the savanna biome. Vegetation biomass from this bioregion was estimated at $84.7 \pm 7.1 \,\mathrm{Mg} \,\mathrm{ha}^{-1}$ or $\sim 41.4 \,\mathrm{Mg} \,\mathrm{Cha}^{-1}$, with an emission estimated at 129 Mg CO₂-e (Henry et al., 2015), similar to the woodland biomass density and resultant emission with deforestation from the CS site of this study.

Our emissions estimate is robust for this vegetation class and can be upscaled and compared with other land sector activities such as prescribed savanna burning. At a regional scale, current levels of savanna burning dominate emissions compared to land clearing rates (Table 5). The cumulative deforestation area across the savanna region since 1990 (1 886 512 ha) is 17 times smaller than the mean annual savanna burn area (32 Mha, Table 5), as approximately 30 to 70 % of the savanna area is burnt annually (Russell-Smith et al., 2009a). Modelling NEP for savanna biome for 1990-2010 (Beringer et al., 2015; Haverd et al., 2013) suggests the northern Australian savanna is near carbon neutrality or is a weak source of CO₂ to the atmosphere once regionalscale fire emissions are included. As such, the IPCC assumption that CO₂ emissions from the previous year's burning are recovered by the following year's wet season growth may have some validity for regional-scale GHG accounting. This assumption at plot-to-catchment scales may not be valid, as localised interannual variability in rainfall, site history and fire management can result in either net accumulation or loss of carbon (Hutley and Beringer, 2011; Murphy et al., 2014, 2015b). Assuming year-to-year CO₂ emitted from burning is resequestered, assessment of the non-CO₂ only emissions from savanna burning with deforestation is useful. This comparison suggests projected deforestation emissions $(24\,393\,\mathrm{Gg}\,\mathrm{CO}_2\text{-e}\,\mathrm{yr}^{-1},\,\mathrm{Table}\,5)$ could be well in excess of current annual burning emissions (6740 Gg CO₂-e yr⁻¹, Table 5), at least for the period of enhanced clearing, which in this study we assumed to be 5 years.

In 2013, Australia's total reported GHG emission was 548 440 Gg CO₂-e and the impact of expanded savanna deforestation on the national emission can be estimated using data in Table 5, which provide estimates of mean annual emissions from the deforestation area, giving a mean annual deforestation emission per ha averaged for the entire savanna area, which is $221 \pm 50.8 \, \text{Mg CO}_2$ -e ha⁻¹ using 1990 to 2013 data (Commonwealth of Australia, 2015a). This value represents a spatially averaged emission as it is derived from the full range of savanna vegetation types and

above-ground biomass, which across the Northern Territory savanna area ranges from 10 to 70 Mg C ha $^{-1}$ (Law and Garnett, 2011). Assuming this emission per ha, an additional 311 000 ha of savanna deforestation, cleared over a 5-year period, adds 12 099 Gg CO₂-e yr $^{-1}$. For the duration of the expanded deforestation, this is a 2.2 % increase to Australia's nation emission over and above the historic savanna LUC emissions (16 161 Gg CO₂-e yr $^{-1}$), which are 2.9 % of national emissions. Using our finding from flux tower measurements that a land conversion (deforestation followed by site tillage and preparation for cultivation) adds an additional 18 % of GHG emissions to a deforestation event, expansion of northern land development could add an additional 3 % or 33 350 Gg CO₂-e yr $^{-1}$ to the reportable national GHG emissions for the duration of the expanded deforestation period.

This assessment is subject to a number of uncertainties. Firstly, a component of our emissions estimate is based on eddy covariance measurements of CO₂ flux, which typically have an error of 10-20% (Aubinet et al., 2012). In this study, energy balance closure suggested fluxes were underestimated by up to 13% across the entire observation period. Energy balance closure ranged from <10 % flux loss during the intact canopy phase to > 20 % error during the final three LUC phases when the flux instruments were at 3 m height measuring net soil CO₂ emissions from the smoothed, vegetation-free ploughed soil surface during preparation. Secondly, it is difficult to predict the nature of future deforestation (rate, area, specific location) and the emission comparisons presented here are indicative only. Catchments selected by Petheram et al. (2014) regarded as suitable or with potential for future development were based on biophysical properties only, were unconstrained by the regulatory environment and did not account for conservation and cultural values placed on identified land and water resources. In addition, challenges to agricultural expansion in northern Australia include uncertain land and water tenure, high development costs and lack of existing water infrastructure, logistics and technical constraints, lack of human capital and distance to markets, all factors that may restrict land clearing. It is well understood that the availability and cost of water for irrigated, or irrigation-assisted agriculture is critical for viable agriculture in northern Australia (Petheram et al., 2008, 2009). Australian governmental policies currently support small-scale, precinct or project-scale approaches, based on well-understood water and soil resources, where water allocation is capped. The current policy and market instruments are likely to ensure that development remains measured and restricted, unlike development of previous decades in other regions of eastern and southern Australia.

As a result we used a conservative estimate of potential land suitability area (311000 ha over a 5-year clearing period), as estimates of assumed clearable area ranging up to 700000 ha (e.g. Douglas–Daly catchment, Adams and Pressey, 2014) or over 1 million ha across northern Australia (Petheram et al., 2014), areas that may be unlikely

given capital investment requirements as well as conservation and cultural considerations. Our comparison with burning emissions is also influenced by the deforestation period we assume. This was based on patterns of historic rates of clearing as there are periods when deforestation rates have easily exceeded 311 000 ha over 5-year periods, particularly in Queensland (Commonwealth of Australia, 2015a) and a longer duration of deforestation reduces the impact on annual national GHG accounting.

There is also uncertainty arising from our emissions from debris burning. Russell-Smith et al. (2009b) estimated errors associated with emissions estimates from the Western Arnhem Land Fire Abatement (WALFA) project, a savanna burning based GHG abatement scheme operating in the Northern Territory. This is a project area of the 23 893 km² consisting of a wide range of vegetation types including open-forest and woodland savanna and sandstone heaths in escarpment areas. Russell-Smith et al. (2009b) estimated the accountable emissions from savanna burning at $272 \pm 100 \,\mathrm{Gg}\,\mathrm{CO}_2$ $e yr^{-1}$ (95 % CI), an error of 30–35 % of the mean. Uncertainty was ascribed to errors in remotely sensed burn area mapping, fuel load estimation, spatial variation of fire severity, errors in BEF for each fuel class and EFs. At the spatial scale of our study area, there were no uncertainties with the burnt area, vegetation structure or fuel type classification, and we used site-specific fuel load estimations used in our calculations, all of which would reduce the error associated with our fire emissions estimate. Russell-Smith et al. (2009b) also reported low coefficients of variability (CV%) of for BEFs across fine, course and heavy fuel types for high severity fires, ranging from 0.3 to 11 % and 2 % CV for EFs for CH₄ and N₂O. Site-specific sources of error include high spatial variability of on-site fuel loads which had a CV % of $\sim 70\%$ (Table 4) and uncertainty associated with the BEF we assumed for coarse and heavy fuel loads (0.9), which is higher than that derived for late dry season savanna fires (0.36, 0.31 respectively, Russell-Smith et al., 2009b). This value was assumed as repeat burning of coarse and heavy fuels ensured $\sim 10\%$ of biomass remained as ash and charcoal at the CS site. This assumed BEF is also consistent with FullCAM (4.00.3) BEF of 0.98 for forest fire with 100 % of trees killed, although this is setting is based on Surawski et al. (2012) who found little empirical evidence for BEF for stand replacement fires. However, given the detailed on-site measurements of fuel load, error in our fire-derived emissions would be of the order of 20 % or less.

5 Conclusions

While GHG emissions from savanna deforestation are dominated by debris burning, emissions from soil tillage and soil bed preparation are likely to be 20% of the total emission, suggesting that satellite-based emissions based on oxidation of cleared vegetation alone do not capture all phases of LUC

prior to cultivation. Savanna burning, using the area as defined in this study, was 1.5% of Australia's national GHG emissions and is of similar magnitude to emissions associated with historic savanna deforestation. However, the deforestation scenario could increase Australia's GHG emissions by at least 3 % per annum for the duration of the expansion, depending on the area and deforestation rate. These are indicative estimates only, but suggest that the impacts of northern agricultural development will have an impact on the national GHG budget and will need to be considered in northern land use decision making processes. These considerations are also particularly relevant given the emission reduction targets set by Australia following the 21st Conference of Parties to the UN Framework Convention on Climate Change (COP21/CMP11) to reduce GHG emissions by 26 to 28 % from 2005 to 2030.

6 Data availability

Flux data from the UC and CS tower sites is available from Australia's Terrestrial Ecosystem Research Network (TERN, http://tern.org.au/) OzFlux Data Portal. Data can be accessed from http://data.ozflux.org.au/portal/site/datainfo.jspx under a user licence as described at http://data.ozflux.org.au/portal/site/licenceinfo.jspx. Fire-derived GHG emissions data can be estimated via the SavBat2 online calculation tool as described in the text using emissions factors and burning efficacies as given in Table 4. Deforestation data are available via the technical reports as cited, with savanna-specific data available on request to the corresponding author.

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