Numerical trajectory patterns of floating macroalgae

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The ocean currents act as an important dispersal mechanism of material on the ocean surface, both natural and anthropogenic. Macroalage are one of the most important natural dispersal mechanisms in the ocean, and are regarded as the tumble-weed of the ocean. Despite the many studies that exist, the relative role of wind and surface currents in macrolage trajectory is still unknown. Past studies also tend to focus on kelp-rafts which can vary greatly in size and making it difficult to form a concensus as to the role of wind and surface currents. In addition past studies have not conisidered surface-area in relation to drag characterisics of floating macroalgae, which have been shown to play a role in the trajectory and accumulation of other flotsam. The advances in Lagrangian trajectory modeling has been exponential in recent years and allows the use of numerical experiments to investigate trajectory of flotsam in the ocean. This study aimed to shed light on the relative role of wind versus surface currents, as well as the role of drag, in macrolagae trajectory and accumulation. This study did not investigate kelp-raft behaviour but instead focused on solitary kelp individuals. This was achieved by comparing simulations of virtual kelp particles which incorporate drag with that of purley lagrangian particles. The results showed that the virtual kelp particles do follow the ocean currents in general, but that the inclusion of drag into the scnearios causes accumulation in eddies and vortices within the study domain. The results also point to the importance of sub-mesoscale features, such as eddies and vortices, as important surface accumlation zones for solitary floating macroalage.

# Introduction

There are a range of objects, both natural and anthropogenic, floating in the ocean of which macroalgae are regarded as one of the most important passive dispersal mechanisms of marine taxa. Floating kelp acts as a passive dispersal mechanism for a range of marine taxa and is sometimes referred to as the ‘tumble-weed’ of the ocean (Edgar, 1987; Norton, 1992; Bushing, 1994; Helmuth et al., 1994; Holmquist, 1994; Smith, 2002; McCormick et al., 2008). Some macroalgae species are negatively buoyant and sink to the seafloor when detached from the substrate, while other species of macroalgae have air-filled pneumatocysts or stipes which allow the plants to reach the surface where light is more abundant. In turn, the positively buoyant pneumatocysts cause plants to float to the surface when dislodged from the substratum. The structure and number of pneumatocysts varies between species. For example species from the genera *Macrocystis*, *Sargassum*, *Ascophyllum*, and *Fucus* have thalli with many small pneumatocysts, while other kelp species such as *Nereocystis luetkeana*, *Pelagophycus porra*, *Ecklonia radiata*, *Ecklonia maxima* have a single, large pneumatocysts (Dayton, 1985; Smith, 2002; Thiel and Gutow, 2005; Graiff et al., 2016; Batista et al., 2018). In some cases the stipe itself is air-filled, such as with *E. maxima* and *E. radiata*. Although there have been reports of floating Chlorophyta species and some Rhodophyta species, Phaeophyceae species are the most commonly reported forms of floating algae. This is most likely because the green and red species reported floating are not actually positively buoyant, but instead are kept at the surface by gas trapped inbetween or in the thalli (Dromgoole, 1982; Bäck et al., 2000). The giant kelp *Macrocystis pyrifera* (Helmuth et al., 1994; Kingsford, 1995; Hobday, 2000; Macaya et al., 2005; Graiff et al., 2016; Batista et al., 2018), the bull kelp *Durvillaea antarctica* (Smith, 2002; Collins et al., 2010; Wichmann et al., 2012; Tala et al., 2013 , 2017; Saunders, 2014; Batista et al., 2018), and pelagic *Sargassum* species (Brooks et al., 2018, 2019; Putman et al., 2018, 2020) have been the focus of much of the research regarding spatial and temporal dispersal patterns, the ecological role of rafting, marine connectivity and raft-time.

Past research points to macroalgae trajectory being largely determined by prevailing wind conditions and surface currents (Hobday, 2000; Thiel and Gutow, 2005; Fraser et al., 2011; Rothäusler et al., 2011, 2011). Although ocean currents are regarded as the primary influence, the relative importance of wind versus surface current is still unknown although the role of wind has been recognised as important in several studies. For example, a study by Harrold and Lisin (1989) investigated the seasonal trajectories of radio-transmitter tagged *M. pyrifera* in nearshore Monterrey Bay. The results showed that kelp rafts, with little surface, area exposed to the wind were largely driven by a combination of wind and wind waves, however the relative importance of wind and wind waves was not clear. In addition, the tagged kelp trajectories were more consistent with the formation of eddies during winter. Previous studies have identified wind as an important mechanism of dispersal in wind dominated ocean systems. For example, the subAntartic latitudes the West Wind Drift causes continuous unidirectional surface flow and is regarded as an important potential mechanism for dispersal of floating kelp. Other studies have used genetic approaches to determine macroalgae raft trajectory characteristics by inferring source location from genetically similar populations (Nikula et al., 2013). For example, a study by Fraser et al. (2011) on the rafting capabilities of *Durvillaea antarctica* used a combination of population genetics and relative age estimate of ‘goosebarnacles’ attached to the raft. The presence of goosebarnacles suggests a long raft time as these species have a slow growth rate; while the genetic analyses showed these species are able to raft up to 390km from their local origin. The authors suggested that wind and water-movement were the primary influences of trajectory, however, this was only inferred from the genetic results and local climatology data (Fraser et al., 2011). Past research has shown that buoyancy and drag also play a role in determining the trajectory and rate of transport for surface floating material (onink2019; Shen et al., 2019; Furnans et al., 2008; Beron-Vera et al., 2016; Trinanes et al., 2016; D’Asaro et al., 2018; Gates et al., 2018). The “sail” area of an object floating at sea is the surface area exposed to the wind which results in air form drag, while the area of the object below the surface of the water is exposed to surface currents which result in hydrodynamic form drag (Breivik et al., 2011; Griffin et al., 2017). In addition, past research conducted within the maritime industry has shown that the size and shape of a vessel determine the relative importance of waves or wind as drivers of trajectory, as well as orientation of the object. If the length of the object is longer than the significant wave height then waves will be the primary driver of trajectory while the opposite is true for smaller objects where the effect of waves is regarded as negligible (Breivik et al., 2011; Griffin et al., 2017). Drag is ultimately determined by the size and shape of the object which are properties that vary considerably with macroalgae species.

To accurately determine the trajectory of marine macroalgae these aspects need to be taken into account. Past research by Allen and Plourde (1999) and Breivik et al. (2011) have provided estimates of drag for various objects based on experimental work, which consisted of both direct and indirect methods of the Leeway model. However, this work does not consider biological material such as macroalgae. Although drag estimates do exist, these have been applied for macroalgae not detached from the substratum and are regarded as fixed-point estimates. Although various approaches have been used in the past to investigate floating macroalagae trajectory, very few studies have employed the use of Langrangian trajectory modeling. Furthermore, none of the existing studies using this modeling approach have considered macroalgal morphology cross-sectional area (i.e. shape) as an aspect of drag and ultimately trajectory. Studies by Brooks et al. (2019), Putman et al. (2018), Putman et al. (2020) used lagrangian approaches to investigate the effects of inertia, raft-size, and windage on macroalgae trajectory. Brooks et al. (2019) used a custom growth model to estimate changes in biomass and ultimately radial size, while a customised Hybrid Coordinate Ocean Model (HYCOM) for the trajectory simulations was used. The results showed that trajectory of pelagic *Sargussum* was significantly influenced by inertia and the radial size of the rafts. Other work by Putman et al. (2020) investigated the effect of including windage in macroalgae trajectory simulations. The authors showed that including ad-hoc windage factors greatly improved accuracy regarding entrainment or expulsion from eddies when compared to tracked sargassum mats in the ocean. Langrangian trajectory modeling is a useful tool, however, the various models available often assume the object to be a spherical particle, which is obviously not the case for both biological and anthropogenic forms of marine debris in nature. The radial size included by Brooks et al. (2019) is an improvement, however this approach does not take into account the complex morphology on an individual level into account.

In addition, past studies focus on kelp-rafts, not solitary floating kelp, and therefore are not able to investigate the role of eddies and vortices in the accumulation of kelp at the surface. Gyres, eddies and converging surface currents have been identified as important in the accumulation and dispersal of plastics and other marine debris (Brach et al., 2018; D’Asaro et al., 2018; Onink et al., 2019). Past studies investigating the dispersal of marine debris have also included the possible effects of inertia in accumulation of ocean drifter instruments in mesoscale features. For example, a study by Beron-Vera et al. (2016) used a modeling approach to investigated the role of eckman transport and inertia to explain the observed behaviour of undrogued drifters in the ocean gyres. Undrogued drifter trajectory is affected by wind unlike the drogue counter parts which makes the effects of wind neglible. The results showed that the combined effect inertia, not eckman transport, played a significant role in the observed drifter behaviour (Beron-Vera et al., 2016). The authors also noted that the presence of other flotsam along with the undrogued drifters in ocean gyres also suggests that aspects inertia also play a role in accumulation of other marine debris (Beron-Vera et al., 2016). The approach by Beron-Vera et al. (2016) has also been improved upon by Beron-Vera et al. (2019) but also shows the significant role of inertia in flotsam trajectory and accumulation.

In this study, a numerical approach is used to shed light on the role of hydrodynamic and direct wind drag on macroalgal drift trajectory using lagrangian based trajectory simulations. It should be noted that this study does not attempt to characterise the dispersal of drifting macroalage within the study domain. The authors argue that too little is know about the drift characteristics of solitary floating macroalgae to characterise dispersal effectivley. Instead, this study will aim to shed light on the ‘behaviour’ of surface drifting macroalgae from a coastal environment into the open ocean, which will inform future characterisation studies. This will be achieved by comparing the trajectory outputs of virtual ‘passive’ particles (no drag characteristics added) with kelp particles (drag characteristics added) of varying degrees of hydrodynamic and wind drag exposure. The different degrees of wave and wind exposure are meant to reflect the different buoyancy characteristics, as well as identify the role of magnitude of the drag components on overall trajectory. In addition, the modeling of single kelp offers the opportunity to study the role of eddies and vortices in accumulation points on the surface. Therefore, separate simulations will be run to compare the entrained characteristics of passive particle with a kelp particle.

# Methods

## Study site

The region of interest for this study was the Cape Peninsula in the Western Cape (Figure 1A). A release site was chosen for this study, Kommetjie, which is located on the Atlantic side of the Cape Peninsula (Figure 2B). The hydrodynamics in the study region are driven by a complex interaction of wind and waves, which are also influenced by larger oceanic processes. In South Africa, the larger processes can largely be attributed to the Agulhas Current (AC) and the Benguela Current (BC)(**???**; Garzoli et al., 1996; Lutjeharms et al., 2000; Lutjeharms, 2007; Rubio et al., 2009). The AC is a component of the South-west Indian Ocean sub-gyre, flowing predominantly south-westward following the continental shelf edge and eventually retroflecting eastward back into the South Indian Ocean (@ Lutjeharms et al., 2000; Lutjeharms, 2007). The region where the retroflection occurs is known as the Agulhas retroflection and consists of high levels of mesoscale variability (Garzoli et al., 1996). The variability results in shedding of rings and filaments known as the Agulhas Rings and have an average diameter of ~324km (Lutjeharms and Van Ballegooyen, 1988; Lutjeharms, 2006, 2007). The Agulhas Rings are the main contributors to the Agulhas leakage which are the transport of warm water from the Indian Ocean to the South Atlantic Ocean (Garzoli et al., 1996; Lutjeharms, 2006; Beal et al., 2015). The Agulhas leakage is also a source of meso-scale variability within the Benguela region as anticyclonic rings are shed and drift northwestward along the southern Benguela slope region (Rubio et al., 2009). The Benguela region is characterised by strong coastal upwelling cells which are primarily driven by equatorward winds (Shannon and Nelson, 1996; Hardman-Mountford et al., 2003), and is known as the Benguela upwelling region. The Benguela upwelling region starts at 27S and extends southward to 35S (Shannon and Nelson, 1996). The region is also characterised by high amounts of mesoscale variability in the form of eddies and filaments (Blanke et al., 2002, 2005; Rubio et al., 2009). The Benguela inshore region is dominated by a northwestward flow as a result of topographical steering, wind stress and interactions with passing Agulhas rings and eddies (Veitch et al., 2010).

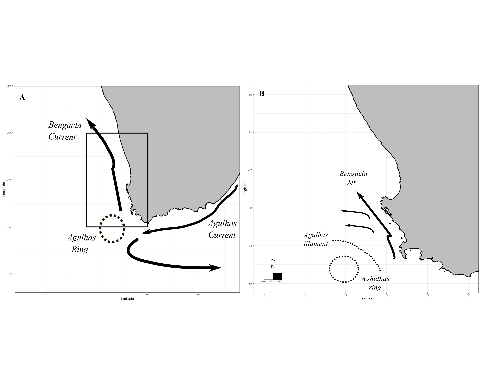


Figure 1:Map of the study domain and the main features pertaining to transport in the region. In panel A, the map of the entire South African coastline and the main currents presented. The resulting eddies and filaments from the Agulhas retroflection and study area are presented. In panel B, the splitting of the Benguela Jet, eddie features and general flow of the study domain are presented.

The study domain falls within the southern Benguela upwelling ecosystem, which is a region of high mesoscale variability and activity. These occur mostly in the form of eddies, filaments and vortices which are generated from the Agulhas retroflection (Veitch and Penven, 2017; Veitch et al., 2018), generally flowing northward along the coastline and begin to move offshore near Cape Columbine. The Benguela Jet is a coastal flowing current and is driven by a complex interaction of eddies, upwelling, eckman transport and geostrophic currents (Stramma and Peterson, 1989; Hutchings et al., 2009; Veitch and Penven, 2017; Veitch et al., 2018). The stability of the jet varies seasonally with more stable flows in winter and splitting of the jet during summer and spring. In January, the jet begins to split between Cape Point and Cape Columbine (**???**; Veitch et al., 2018) which results into a weak narrow inshore branch which flows along the shelf edge and a strong flowing north-westwards offshore branch. The area is also an area of high upwelling activity driven by strong westerly winds (Hutchings et al., 2009). These winds are most intense during summer (December-March) which favours upwelling conditions in the region, and also act to reinforce the flow of the Benguela Jet; which occurs due to cold upwelling waters contributing to the across-shore density gradient (Ragoasha et al., 2019). In winter the wind direction is reversed which decreases upwelling favorable conditions, but the core of the Benguela jet is maintained by the intrusion of warm agulhas waters into the Benguela region over the shelf edge (Hutchings et al., 2009; Veitch et al., 2018; Ragoasha et al., 2019).  
## Particle tracking model

The model is based on parameterisations that have been used for previous studies investigating iceberg, capsized marine vessels and microplastics. It should be noted that most of the previous studies regarding macroalgae trajectory have investigated rafts and not solitary floating individuals. A kelp raft will vary in size and can range from a few meters across to large dense mats the size of a sports field. Furthermore, some authors suggest that most macroalgae become entangled through ocean currents and not through the hydrodynamic forces that dislodge them. In addition, large kelp rafts are not a characteristics around the coast of the Cape Peninsula but rather solitary kelp are (Pers. Obs). Therefore, the trajectory of only solitary *E. maxima* individuals will be considered. Numerical calculations of particle trajectory with a Lagrangian method (Delandmeter and Van Sebille, 2019), and be described by the equation

where is the three-dimensional position of a particle, is the three-dimensional velocity field at the location in the ocean model at particle point in time, is time, is a change in position due to ‘behavior’ of the particle at a particular point in time. This can range from swimming in fish to sinking or beaching. In order to investigate the effects of drag components of trajectory of floating macroalgae, this study incorporated drag as the custom behavior.

This study simulated two scenarios; one where the plant is fully submerged and only exposed to hydrodynamic drag; and the other where parts of the plants are partially exposed resulting in a combination of hydrodynamic and air form drag. Drag based on shape-based coefficients were used in the calculations for determining the overall kelp velocity vectors. A momentum energy equation was used to calculate the drag forces for the relevant simulation. This approach has been employed in modeling iceberg drift trajectory (Lichey and Hellmer, 2001; Eik, 2009; Andersson et al., 2017) and was adapted to suit this particular study. The energy momentum equation used to calculate the drag force exerted on the virtual kelp particle was,

where is kelp mass, the virtual kelp particle velocity, the hydrodynamic drag, is the wind drag, the surface-current flow, and is the time step. The assumption made was that the mass of the kelp did not change for each virtual kelp particle over the course of the simulations. Hydrodynamic drag, wind drag and surface current-flow are two-dimensional vector quantities and were calculated for each time step for each simulation.

## Simulations

Virtual particles were released with no drag or wind effects from the study site in order to act as a reference for comparison to simulations with varying degrees of hydrodynamic and wind drag, and are referred to hereafter as ‘passive particles’. All particles that have any form of drag exposure are hereafter referred to as ‘kelp’ particles. All simulations have the same release location (Kommetjie) and were run for 30 days with hourly outputs using a fourth order Runge-Kutte advection scheme (**???**; North et al., 2009). Brownian motion was also incorporated in order to account for any stochasticity not resolved by the model. These are standard practices for Lagrangian trajectory studies (**???**; North et al., 2009; Delandmeter and Van Sebille, 2019). To test the different magnitudes that hydrodynamic and wind drag might play in the overall trajectory, simulations were run with varying degrees of drag exposure scenarios. These scenarios were meant to act as proxy for buoyancy, which determines the surface area exposed to hydrodynamic and wind drag. These estimates were expressed as percentages which were used to calculate hydrodynamic and wind drag for applicable simulation. For example, if 85% of the plant is submerged and 15% of above water, then the overall hydrodynamic and wind drag forces would be 85% and 15% of the total drag forces respectively. This approach has been used in past research investigating pelagic Sargassum rafts (Putman et al., 2018, 2020). The exposure scenarios used in this study were 100%, 95%, 90%, and 85% for hydrodynamic drag; 0 %, 5%, 10%, and 15% for wind (see table 1). Each scenario consisted of a particle set with a size of 1000 where released from Kommetjie at the start of simulation. The particles sets were advected with a Runge-Kutte fourth order scheme for 30 days and incorporated brownian motion. Since the aim of this investigation was to investigate the relative surface behaviour of virtual kelp particles, only the surface layer was used for simulations. It should also be noted that the Copernicus model used places velocity vectors over some of the land covering the coastal domain. This was not considered a concern, as once again, the aim was to investigate surface drift behavior. The time period chosen to run the simulations was 01-01-2018 un 30-01-2018. This particular time period was chosen as the splitting of the Benguela Jet occurs in Janurary, which also conincides with a period of higher wind intensity, upwelling along the coast and mesoscale activity (Rubio et al., 2009; Veitch et al., 2018; Ragoasha et al., 2019).

Table 1:Details for each simulation executed for the study.

| **Scenario** | **ID** | **Type** | **Parameters** | **Cross-sectional area** | **Drag Exposure level** |
| --- | --- | --- | --- | --- | --- |
| 1 | RK4\_Pfloat | Passive | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | N/A | N/A |
| 2 | Min\_Kfloat\_H100W00 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Minimum | Hydrodynamic drag only |
| 3 | Min\_Kfloat\_H095W005 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Minimum | Hydrodynamic drag 95%, wind drag 5% |
| 4 | Min\_Kfloat\_H090W10 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Minimum | Hydrodynamic drag 90%, wind drag 10% |
| 5 | Min\_Kfloat\_H085W15 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Minimum | Hydrodynamic drag 85%, wind drag 15% |
| 6 | Mean\_Kfloat\_H100W00 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Mean | Hydrodynamic drag only |
| 7 | Mean\_Kfloat\_H095W005 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Mean | Hydrodynamic drag 95%, wind drag 5% |
| 8 | Mean\_Kfloat\_H090W10 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Mean | Hydrodynamic drag 90%, wind drag 10% |
| 9 | Mean\_Kfloat\_H085W15 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Mean | Hydrodynamic drag 85%, wind drag 15% |
| 10 | Max\_Kfloat\_H100W00 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Maximum | Hydrodynamic drag only |
| 11 | Max\_Kfloat\_H095W005 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Maximum | Hydrodynamic drag 95%, wind drag 5% |
| 12 | Max\_Kfloat\_H090W10 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Maximum | Hydrodynamic drag 90%, wind drag 10% |
| 13 | Max\_Kfloat\_H085W15 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Maximum | Hydrodynamic drag 85%, wind drag 15% |
| 14 | RK4\_Pfloat\_acc | Passive | Runge-Kutte (4th), Brownian motion, 30 days, repeated release every hour | N/A | N/A |
| 15 | Max\_Kfloat\_H085W15\_acc | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, repeated release every hour | Maximum | Hydrodynamic drag 85%, wind drag 15% |
| 16 | Max\_ShapeFloat\_H085W15 | Kelp | Runge-Kutte (4th), Brownian motion, 30 days, 1000 particles | Maximum area, cylinder shape | Hydrodynamic drag only |

To determine if vortices and eddies play a role in kelp accumulation on the surface, a separate set of simulations were performed. Unlike the previous simulations these simulation released a particle every hour. These were run over the same time period for 30 days with hourly outputs using a fourth order Runge-Kutte advection scheme and incorporated Brownian motion. In order to determine if the shape co-efficient used plays a role in model used, a simulation of a kelp with the maximum surface area and drag exposure was run using the shape coefficient for a cylinder to compare to a simulation using a sphere, and was run for 30 days using a fourth order Runge-Kutte advection scheme with brownian motion and a particle set of size 1000.

## Model inputs

### Ocean and wind model

Both passive and kelp particles were simulated around the South African coastline within hind-cast outputs from the Copernicus Marine Environment Monitoring Service. The Copernicus outputs used in this study contains 3D daily current information from the top layer to the bottom (Global Analysis Forecast PHY\_001\_024). The model has a spatial resolution of 0.08 and is interpolated on an Arakawa C native grid.

In order to incorporate the effects of direct wind drag within the simulations, the Copernicus Marine Environment Monitoring Service global wind product was used (WIND\_L4\_NRT\_OBSERVATIONS\_012\_004). The outputs from the model is composed of 6-hourly averaged fields of surface 10m wind velocity vectors with a spatial resolution of 0.25. Both averaged ocean current and wind data are stored as meridonal and zonal velocity vectors.

### Hydrodynamic and wind drag

The drag coefficient for a spherical particle was used in the calculation of hydrodynamic drag as the trajectory model assumes the particles are spherical. The drag equation used was: , where is the density of seawater or air, is the cross-sectional area exposed to hydrodynamic or wind drag, is the drag coefficient for a spherical particle and is the surface velocity vector of the flow-/wind- field. Hydrodynamic and wind drag components were calculated separately for both the meridonal and zonal velocity vectors, as per Furnans et al. (2008). Since drag force is dependent on the velocity vectors which vary with time, the meridonal and zonal velocities were interpolated and used in the drag force calculation for each time-step. The same approach was used for the wind drag force. The equation that was used to calculate hydrodynamic drag, was also used to calculate wind drag.

### Cross-sectional area

In order to incorporate hydrodynamic and wind drag, the cross-sectional area of the kelp was calculated first. Known geometric shapes reflecting the relevant plant sections were used to estimate the surface area for various parts of the plant, for details please refer to figure 2. The dimensional data needed was estimated in cases were data was not available for that particular morphological characteristic. The bulb/pneumatocyst is a highly variable morphological characteristic and in some cases can appear absent, the same is true for the holdfast. Therefore, a standard cross-sectional area was used for both the bulb and holdfast, for all simulations.

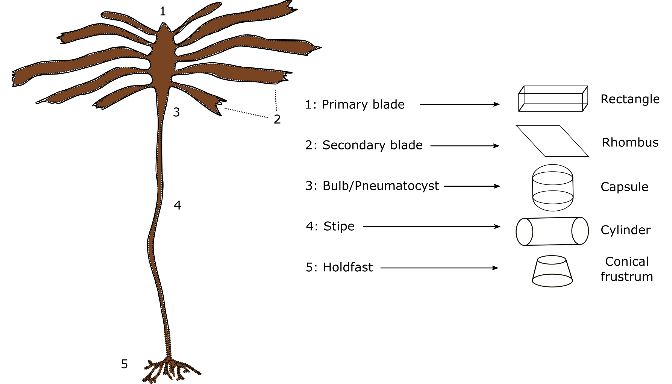


Figure 2:Diagram of the *E.maxima* morphological characteristics considered in this study, along with the associated geometric shapes used to calculate the overall surface area of each morphological characteristic.

The cross-sectional area calculated was site specific and the dimensions needed were garnered from morphology data from a previous study (Coppin et al., 2020). The minimum, mean and maximum were calculated for each morphological characteristic, which were used for calculating the overall minimum, mean and maximum cross-sectional areas needed to run the various simulations (see table 1). The trajectory of these “types” of cross-sectional areas (minimum, mean and maximum overall cross-sectional area) were used to determine the influence of drag, both hydrodynamic and wind, on overall trajectory.

Table 2:Summary of the atrributes and estimates used to calculate the overall surface area of Kommetjie kelp individuals used in the various simulation.

| **Plant characteristic** | **Approximate shape** | **Equation Used** | **Plant dimensions** | **Ac Minimum** | **Ac Mean** | **Ac Maximum** |
| --- | --- | --- | --- | --- | --- | --- |
| Secondary blade | Rectangle | Ac = 2lw + 2lh + 2wh | frond length, frond width\*,frond thickness\* | 1,136 | 1,993 | 2,550 |
| Primary blade | Rhombus | Ac = (lxb)/2 | primary length, primary width | 96 | 262 | 536 |
| Bulb/pneomatocyst | Capsule | Ac = 4?r2 + 2?rh | bulb length\*, bulb base radius\* | 628 | 628 | 628 |
| Stipe | Cylinder | Ac = 2?r(r + h) | stipe radius from stipe circumference, stipe length | 9,797 | 19,338 | 34,273 |
| Holdfast area | Conical frustrum | Ac =?(R2 + r2) + ?(R+r)?(R-r)2 + h2 | top radius\*, bottom radius\*, height\* | 640 | 640 | 640 |
|  |  |  | Total Area (centimeters) | 12,297 | 22,861 | 38,628 |
|  |  |  | Total area (meters) | 123 | 229 | 386 |
| Total weight | N/A | mass (kg) = stipe mass + (frond mass x 6) | Total weight (kg) | 17 | 34 | 49 |

## Analysis

In order to compare trajectories between cross-sectional area types and wind exposure scenarios, mean trajectories were calculated and density distribution maps were produced. Density distribution maps are an effective method for assessing pathways and ocean connectivity (**???**). In addition, mean trajectory was calculated and included in the density distribution maps to further aid in comparing simulation outputs. The distance that particles traveled was also measured and used to produce boxplots. Distance traveled from the release site can help reveal topographic steering and momentum energy transfer from the wind field to kelp particle (**???**). If the wind direction is in the same direction as the current the virtual particles will cover larger distances due to an increase in the velocity as a result of momentum energy transfer (**???**; Putman et al., 2018, 2020).

# Results

## Trajectory comparison

The mean trajectories of the different cross-sectional areas are similar to the mean trajectory of the passive particles (Figure 2A). When comparing mean trajectories between the different cross-sectional areas for the 100% hydrodynamic drag simulations, no differences can be seen with only slight variation. The results from the simulations including different hydrodynamic and wind drag exposures and cross-sectional area types show differences in mean trajectory when compared to the passive simulation (Figure 2B). The end points of the simulations including hydrodynamic and wind drag were different compared to the passive simulations, as well as being further from the release site. When comparing the different hydrodynamic and wind drag simulations to each other, only slight variation in mean trajectories was evident. In all simulations particles flowed in a north-westward direction.

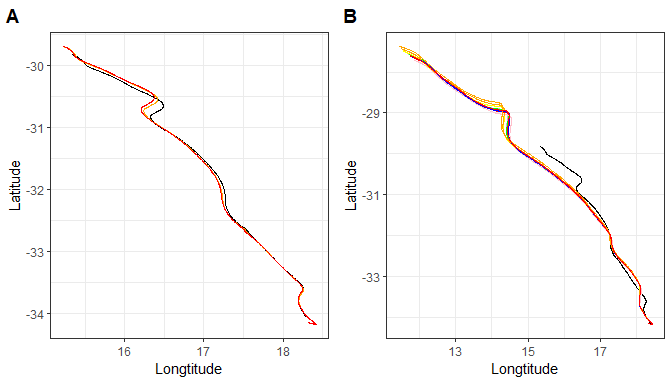


Figure 2: Comparison of mean trajectories. Plot A is the comparison of mean trajectories for only hydrodynamic drag and plot B is the comparison of mean trajectories for simulations with varying degrees of hydrodynamic and wind drag. In plot A, the black line represents simulation 1, the blue line 2, the orange line 6 and the red line 10. In plot B, the black line represents simulation 1, the grey line simulation 3, the green 4, the pink 5, the yellow 7, orange 8, red 9, blue 11, bronze 12, and cherry 13. The reader is referred to tables 1 and 2 for details pertaining to each scenario.

## Comparison of density distributions

Across all simulations particles flowed in a north-westward direction (Figure 3). The density plots for simulations that considered any form of drag showed a higher density of particles along the mean trajectory path compared to the passive simulation which has an almost even density of particles across grids (Figure 3A, 3E, 3I). In addition the density plots show that the particles in simulations that considered any form of drag got entrained in a vorticy while particles in the passive simulation deflected away (Figure 3). Comparison of cross-sectional areas and different drag exposure scenarios show no differences in mean trajectory and spatial patterns in density, with only slight variations across types.

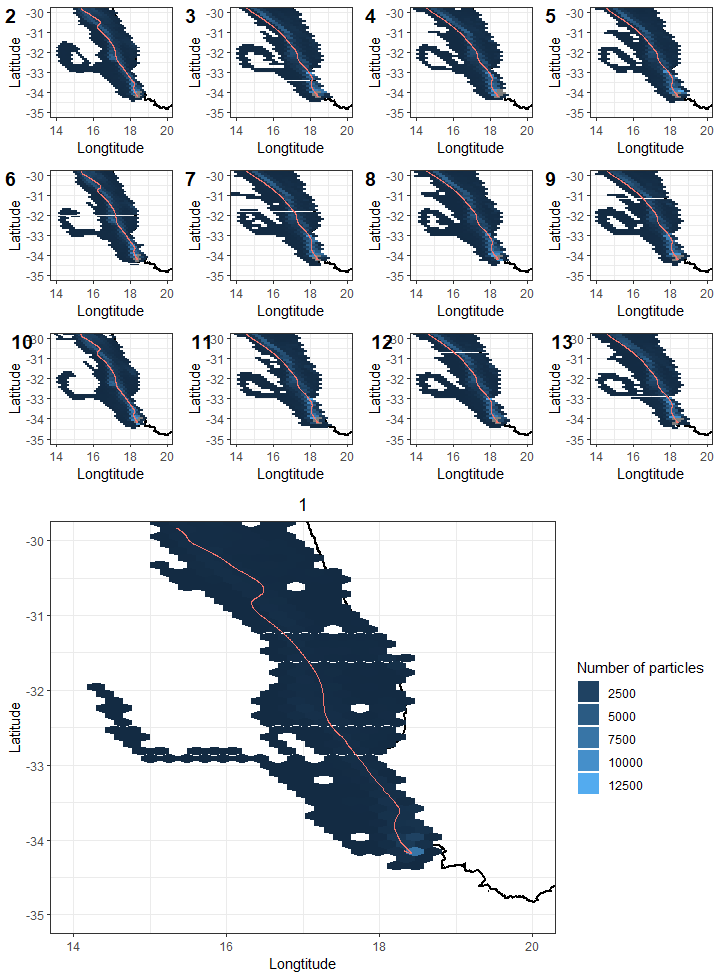
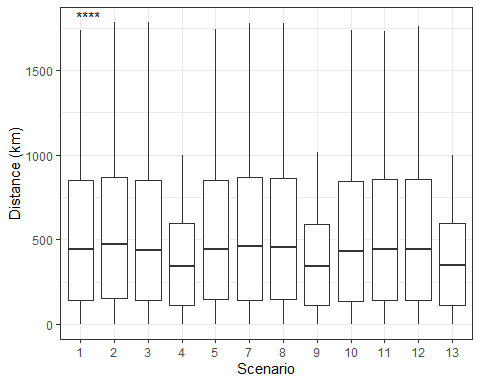


Figure 3: Comparison of density plots at the end run time of each scenario. The plots are regular hexagons within the density of particles is calculated. Plots A-D represents the minimum, plots E-H are the mean and plots I-L the maximum cross-sectional areas. The bottom plot is the density plot for the passive simulation. The mean trajectories for each scenario are also plotted for comparison. The reader is referred to tables 1 and 2 for details pertaining to each scenario.

## Comparison of distances

There were significant differences among the median distances travelled for all particles in each simulation (Figure 4; p < 0.05). The simulations that considered for 100% hydrodynamic drag travelled less distance than particles for simulations which considered a combination of hydrodynamic and wind drag scenarios. Comparison between simulations that considered hydrodynamic and wind drag show no significant differences with only slight variations among simulations.

 ## Accumulation

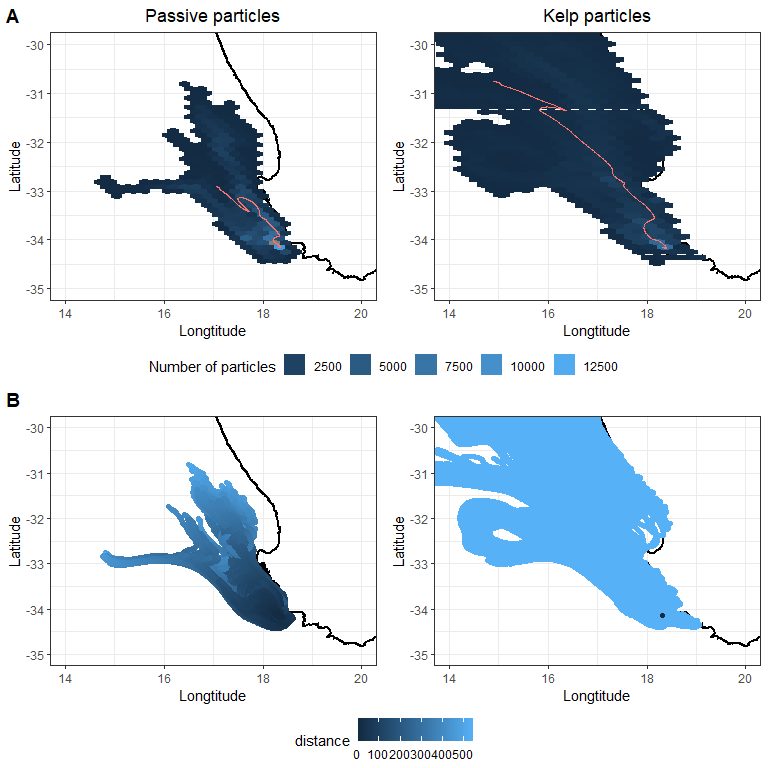


Figure 6: Comparison of trajectory data between passive and kelp accumulation scenarios. Panel A are a comparison of density plots and panel B is a comparison of age.

The role of eddies that exist within the study domain over the time period play a clear role in the accumulation and dispersal of kelp particles. Passive particles tend to remain on the outer edge of the eddy and then are expelled on the other side and continue to follow the ocean current (see Figure 5A). The opposite is true for kelp particles which have the highest exposure level of hydrodynamic and wind drag. The kelp particles are entrained in the eddy (see Figure 5A). This is also reflected when comparing the age of particles at the end of the simulation. In Figure 5B, younger particles are near the release site and older particles further away from shore, while older kelp particles are retained in the eddy.

## Comparison of coefficents

The comparison of mean trajectories between the sphere and cylindrical shape coefficients revealed no dissimilarity (Figure 6A). This was also reflected when comparing distance travelled by particles and particle density plots (Figure 6B). The simulation which used a cylindrical shape coefficient showed higher density in an eddy compared to the simulation which used a sphere for the shape coefficient (Figure 6C-D).

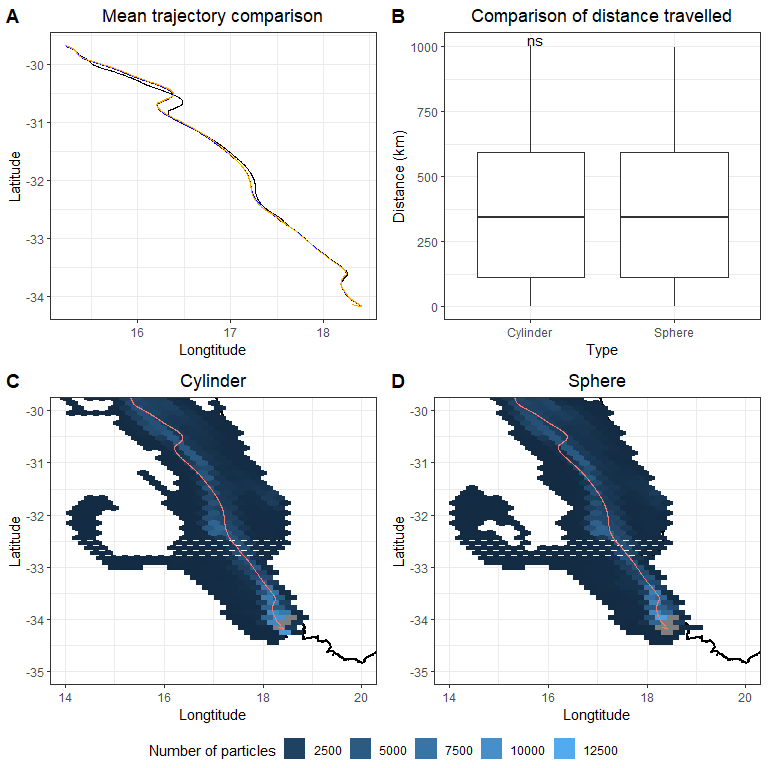


Figure 6: Comparison of trajectory data between a kelp sphere and kelp cylinder. Plot A is the comparison of mean trajectories, plot B is the comparison of total distance travelled by all particles, and plot C and D are the density plots for simulations using a sphere and cylinderical shape coefficient. The reader is referred to tables 1 and 2 for details pertaining to each scenario.

# Discussion

The treatment of drifting macroalgae as purely lagrangian will not be able to accurately determine patterns of passive dispersal, such as entrainment in eddies or vortices (Fossette et al., 2012; Miron et al., 2020). Although past research has shown the inclusion of inertia and windage greatly increase observed macroalgae trajectory and entrainment characteristics (Brooks et al., 2018, 2019; Putman et al., 2018, 2020), past research has focused on large macroalgae-rafts which can vary greatly in size and surface area. None of the past research has taken advantage of the advancement in numerical ocean models and lagrangian trajectory modeling to investigate how aspects of drag (hydrodynamic and wind) affect the trajectory patterns of solitary drifting macroalgae. This study compared particles trajectories with simulations that included no forms of drag with simulations that included various forms of both hydrodynamic and wind drag. The mean virtual kelp trajectory deviated from the mean passive trajectory when including varying levels of hydrodynamic and wind drag scenarios. The results from this study show the inclusion of hydrodynamic and wind drag improves entrainment patterns, and also our understanding of how surface area plays a role in solitary drifting macroalgae trajectory, in this case *E.maxima*.

Past research has inferred from direct and indirect techniques that drifting macroalgae tend to follow the prevailing surface currents (Hobday, 2000; Thiel and Gutow, 2005; Fraser et al., 2011; Rothäusler et al., 2011, 2011). The results from this numerical experiment confirm that drifting macroalgae do follow the surface current in the study region when considering only hydrodynamic drag. In addition, not only are the trajectories similar but so are the end points at the end of simulation. However, when ad-hoc wind drag (windage) is considered in combination with hydrodynamic drag the end locations differ greatly when compared to purely lagrangian particle trajectories. The inclusion of wind drag causes the particles to flow further away from the coast and with a greater distance travelled from the release location. The difference in trajectory and distance covered can be attributed to the inclusion of wind drag. Wind drag adds to the effects of hydrodynamic drag when the wind direction and current direction are opposite to each other. However, when the wind direction is in the same or similar direction to that of the current, momentum energy from the wind is transferred to the object (Hackett et al., 2006; Putman et al., 2020). This causes the particles to travel further as well as exposed to different time varying flows. Past studies have shown this when including windage as a function of inertia (Putman et al., 2018, 2020; Brooks et al., 2019). A study by Putman et al. (2020) used ad-hoc windage factors that was based on the surface current velocity to assess the approach for improving transport predictions of pelagic Sargassum. The results showed that including ad-hoc windage factors improved virtual trajectories compared to tracked Sargassum mats which was partly due to the inclusion of momemtum energy transfer from wind. This is also reflected in the analysis comparing distances travelled by all the particles in each simulation. The particles in the simulations that included wind drag traveled significantly greater distances compared to the passive or hydrodynamic only drag.

The addition of drag forces causes the particles to cluster together along the mean trajectory compared to that of the passive particles which are more evenly dispersed. This is most likely due to the shared drag characteristics among the particles which included the plants weight and surface area into the calculation. Other studies investigating flotsam trajectory characteristics have found similar results (Breivik et al., 2011; Miron et al., 2020; Olascoaga et al., 2020). A study by Miron et al. (2020) conducted field experiments using a range of objects such as spheres, cubes, cuboids to compare to investigate the effect of inertia on particles dynamics. The results from the study showed that objects tended to cluster according to the shape of the object which the authors attributed to shared characteristics of inertia.

The ocean is made up of varying time-varying flows and changes in an objects velocity can cause the object to be exposed to different flow patterns over time. The lack of dissimilarity in trajectory between different cross-sectional area types (minimum, mean and maximum) suggests that reduction in velocity in relation to cross-sectional area is negligible. A similar result was found by Le Gouvello et al. (2020) who investigated the effects of swimming behavior on sea turtle hatchling dispersal in the Agulhas region. The results suggested that the hatchlings trajectories are mostly influenced by the ocean currents in the first year of hatching due to as juveniles do are not swim fast enough to alter thier overall trajectory. Therefore, the differences in cross-sectional area types may not be significant enough in order to have an effect on the relevant velocity vectors (zonal and meridonal). Another possible reason is the resolution of the ocean model used. The resolution of the underlying ocean model is an important aspect of lagrangian ocean modeling and must be able to resolve sub-grid scale ocean processes for finer scale applications. A study by Hart-Davis et al. (2018) assessed the inclusion of stochastic motion, wind and currents into forecasting for search and rescue using the same model ocean model used in this study. The authors found that the inclusion of brownian motion greatly increased the accuracy in representing sub-grid scale processes and that the inclusion of wind, currents and stochastic motion greatly improved forecasting applications. Therefore, the authors argue that the lack of dissimilarity between simulations of different cross-sectional area and drag exposure scenarios is not as a result of the resolution of the model.

The same is also true when comparing different hydrodynamic and wind drag exposure scenarios. These scenarios reflect different buoyancy situations which result in different amounts of surface area exposed to hydrodynamic and/or wind drag. The similarities between the different combined drag scenarios suggests that higher magnitudes of wind exposure do not significantly alter trajectories. Instead, the results from this study suggest that the inclusion of drag forces in simulating macroalgal trajectory may result in improving accuracy of entrainment patterns. All the simulations that included any form of drag resulted in a fairly high density of particles entrained in a vortex. The entrainment and expulsion from eddies and vortices is an important characteristic of floating macroalgae. Accuracy in predicting entrainment and expulsion of macroalgae-rafts from eddies has been investigated previously by Putman et al. (2018) and Putman et al. (2020), who show that including windage greatly improves entrainment patterns when comparing virtual and real Sargassum rafts. However, the aforementioned studies investigated large rafts and not floating individuals. Inertia is based on the size/weight of the raft, while hydrodynamic and wind drag are based on current velocity and surface area. Therefore, the role of inertia may be negligible for a solitary individual and rather other forms of drag, such as in this study, should be considered. The results from this study suggest that the inclusion of hydrodynamic and/or wind drag is an important component when simulating virtual kelp particles will greatly increase the accuracy of entrainment and expulsion from sub-grid mesoscale oceanographic features, such as eddies and vortices.

The role of eddies and vortices in the entrainment of kelp is further supported by the accumulation simulations. These results cleary show that the inclusion of drag greatly increases the number of particles entrained in the eddy. Eddies and vortices may act as accumulation zones for kelp where the formation of rafts may occur, or the kelp is entrained for long enough to eventually sink in area below the eddy. In a study by Woodborne et al. (1989) references a site offshore approximately 120km from Cape Columbine where dredged kelp was found. The area referenced by Woodborne et al. (1989) coincides with the area where virtual kelp are entrained by a vorticy. The use of a cylindrical shape coefficient instead of a sphere in the calculation of drag force also had no effect on the mean trajectory. Instead, the inclusion of a different shape coefficient increased to the density of particles entrained in an eddy field. This suggest that inclusion of different shape coefficients play a role in further improving estimates and hindcasts of macroalgal dispersal patterns.

# Conclusion

Most of the past research has focused on large macroalgae rafts which can vary greatly in size and shape, however no research exists attempting to clarify the drifting characteristics of solitary macroalgae. Understanding how to accurately model the distribution of solitary kelp can lead to be understanding of accumulation zones and sinks and ultimately the ecological links pertaining to those processes. The findings from this study indicate that solitary floating virtual *E. maxima* particles tend to follow the prevailing surface currents and the inclusion of wind causes the trajectories to differ greatly from that of a purely lagrangian particle. Furthermore, the inclusion of both hydrodynamic and wind drag cause clustering of particles along the trajectory as well causing particles to become entrained in eddies. Different cross-sectional areas exposed to hydrodynamic and wind drag had no effect on overall trajectory suggesting that those differences are negligible when investigating dispersal patterns of drifting *E.maxima*. In addition, the use of a cylindrical shape coefficient has no effect on mean trajectory but also causes higher density of particles to become entrained in an eddy. Overall the inclusion of drag forces is an important aspect to consider when investigating the dispersal patterns of solitary macroalgae, while differences in cross-sectional areas are not. This study also provides an approach which can be adapted to model any floating solitary macroalgae, provided the surface area can be estimated accurately. Gaps in the research exist when considering what oceanographic processes play a role in solitary macroalgal dispersal as well as how these vary seasonally. The identification of biological and physical factors that play a role in accumulation zones is also needed, as well as better estimates for sinking rates and raft-times.

# Appendix

## Domain features

library(jpeg)  
library(grid)  
library(gridExtra)  
  
vector\_2 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day2.jpg")))  
vector\_5 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day5.jpg")))  
vector\_10 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day10.jpg")))  
vector\_15 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day15.jpg")))  
vector\_20 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day20.jpg")))  
vector\_25 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day25.jpg")))  
vector\_30 <- rasterGrob(as.raster(readJPEG("OceanParcels/trajectory\_data/vector\_outputs/jpeg/vector\_day30.jpg")))  
  
ggarrange(vector\_2, vector\_5, vector\_10, vector\_15, vector\_20, vector\_25, vector\_30, labels = "AUTO")

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1. University of the Western Cape [↑](#footnote-ref-20)
2. South African Weather Service [↑](#footnote-ref-21)
3. University of the Western Cape [↑](#footnote-ref-22)