

Water in Subsurface

Geohydraulics| CE60113

Lecture:01-02

Learning Objectives

- To identify the subsurface zones
- To use the hydrologic balance equation
- To characterize a groundwater hydrograph

Water Cycle

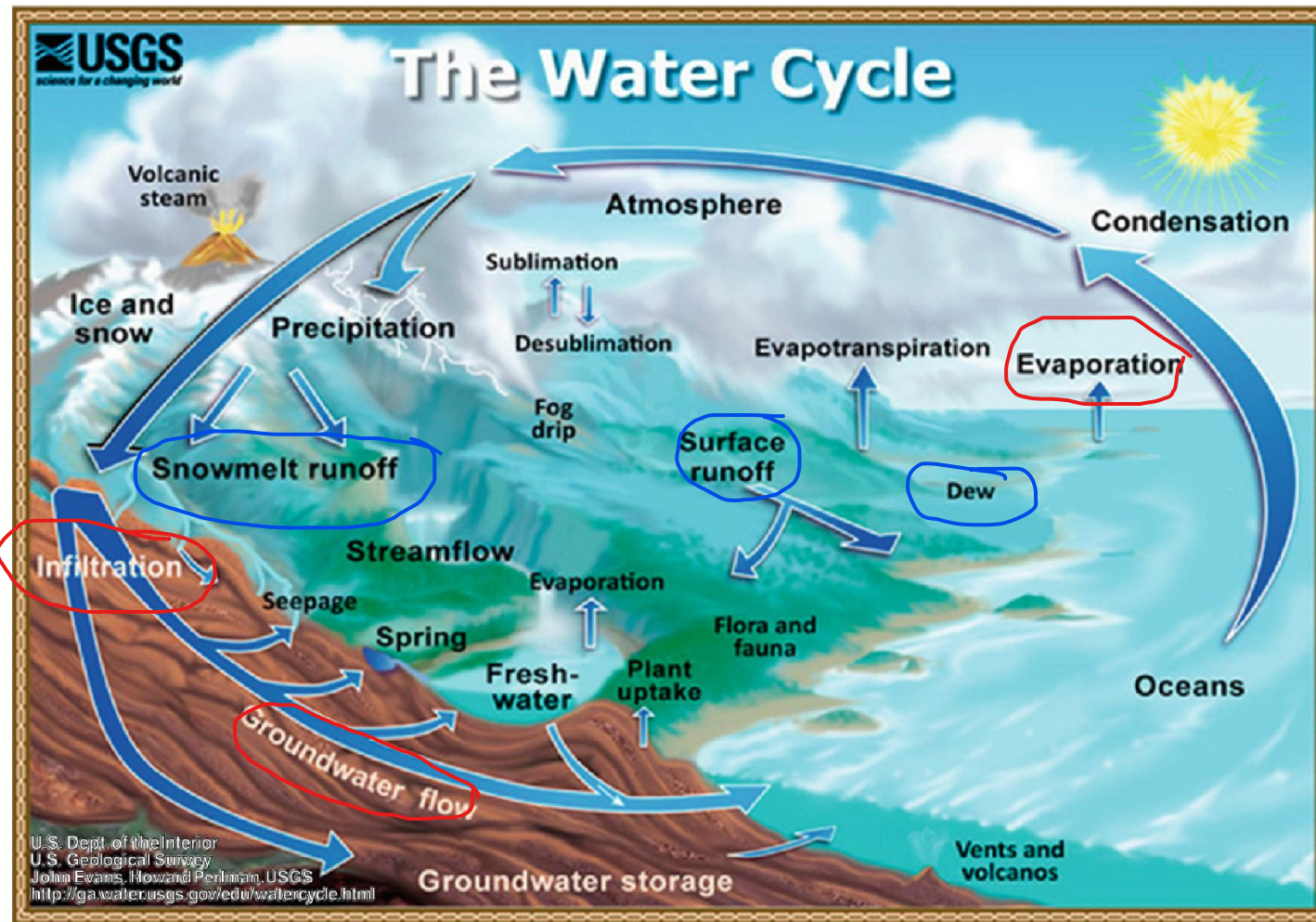
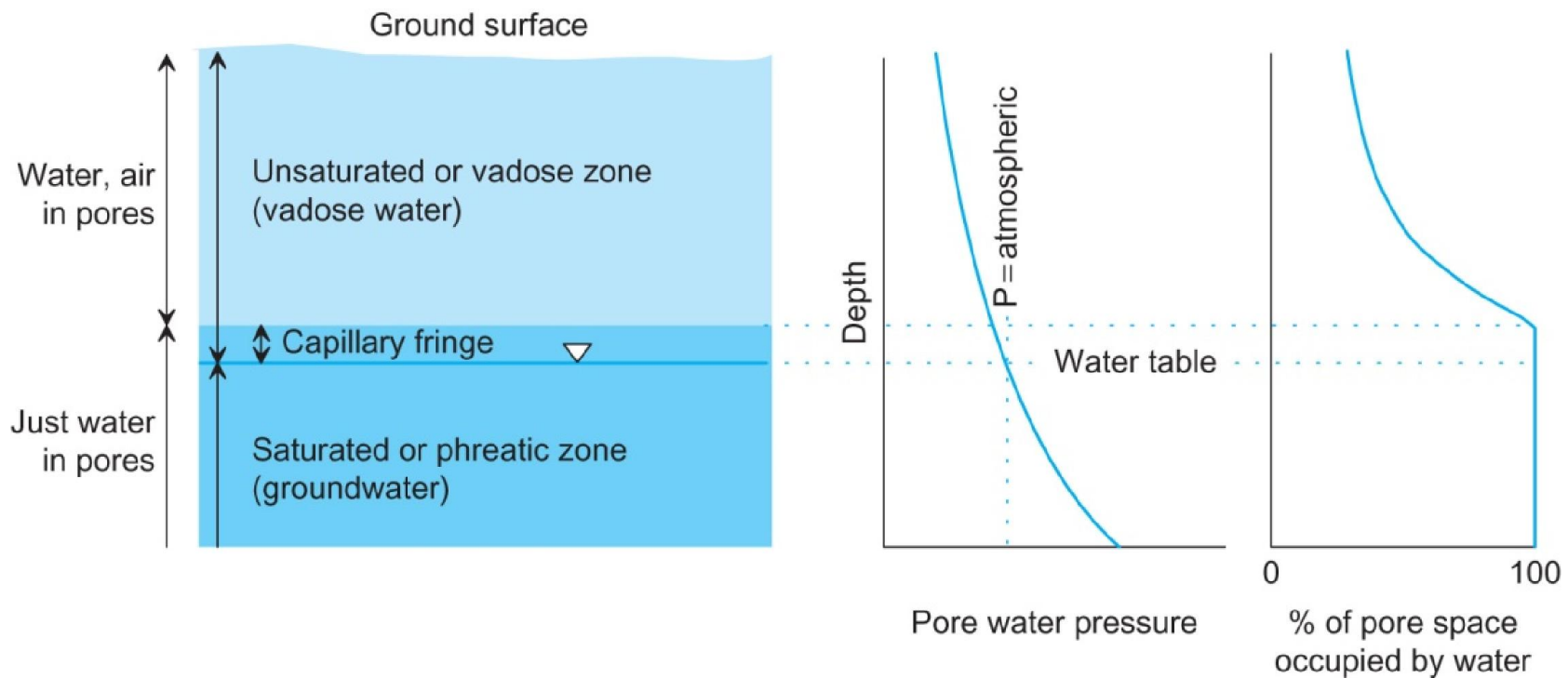


Fig. 1.3 The water cycle (Evans and Perlman 2014)

Subsurface Zones



Subsurface Zones (Contd.)

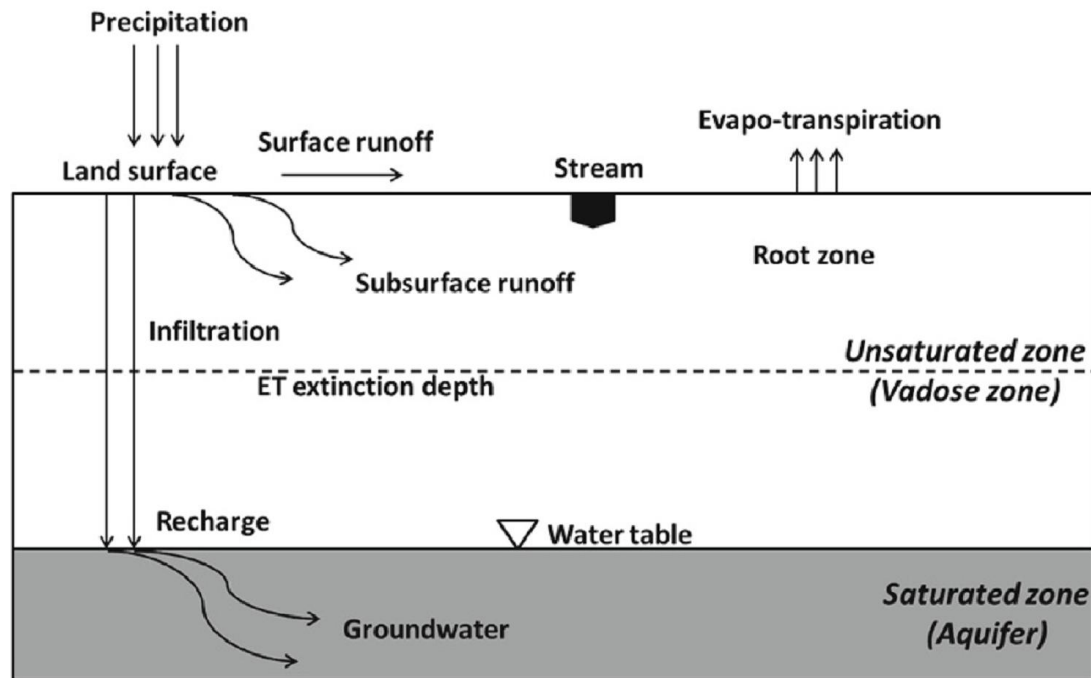
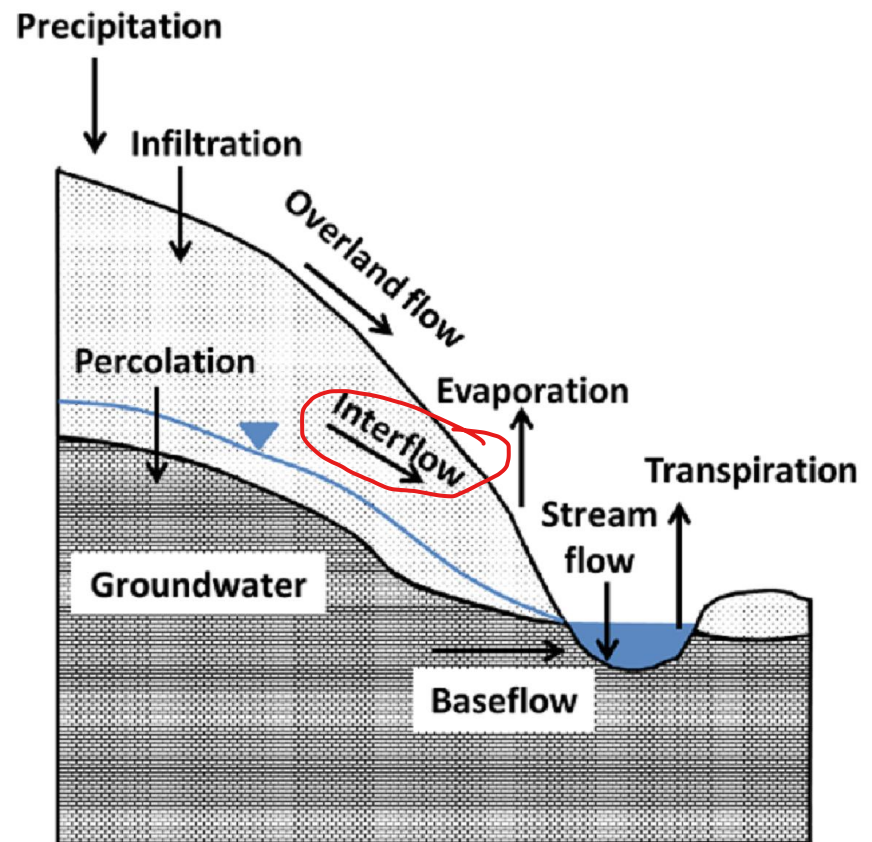


Fig. 1.4 Terrestrial hydrosystems divided into vadose and phreatic zones

Connection between surface and subsurface water

Fig. 1.6 Basic runoff processes



Consolidation Problem

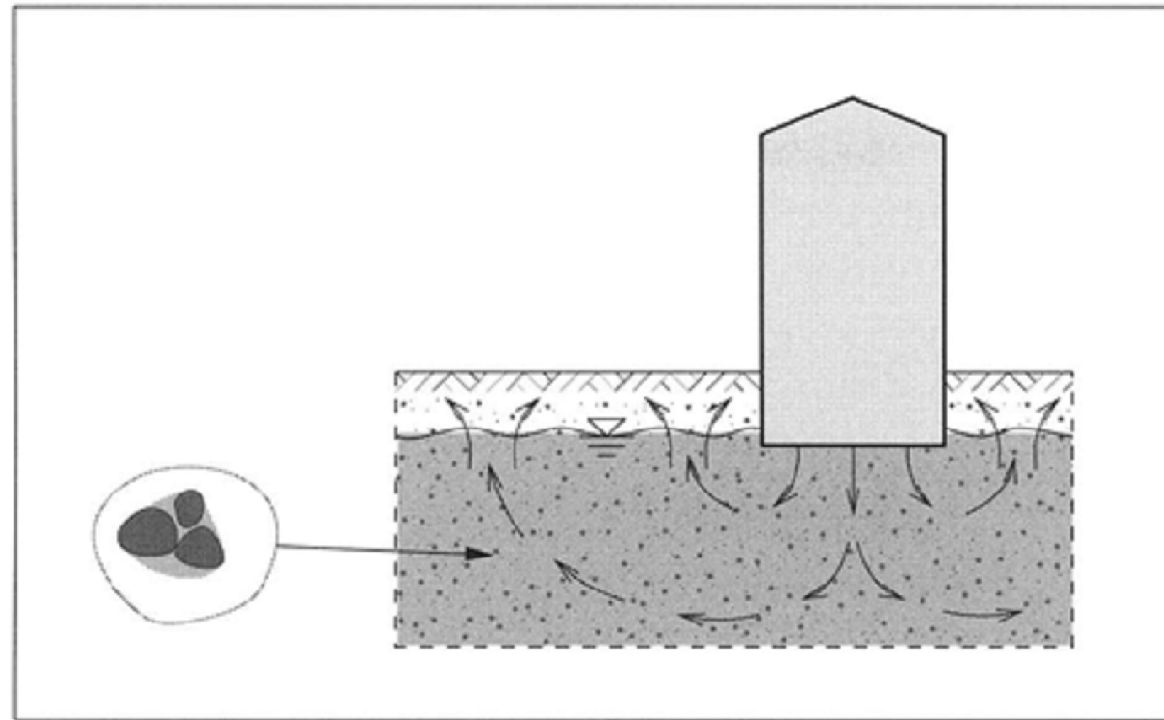
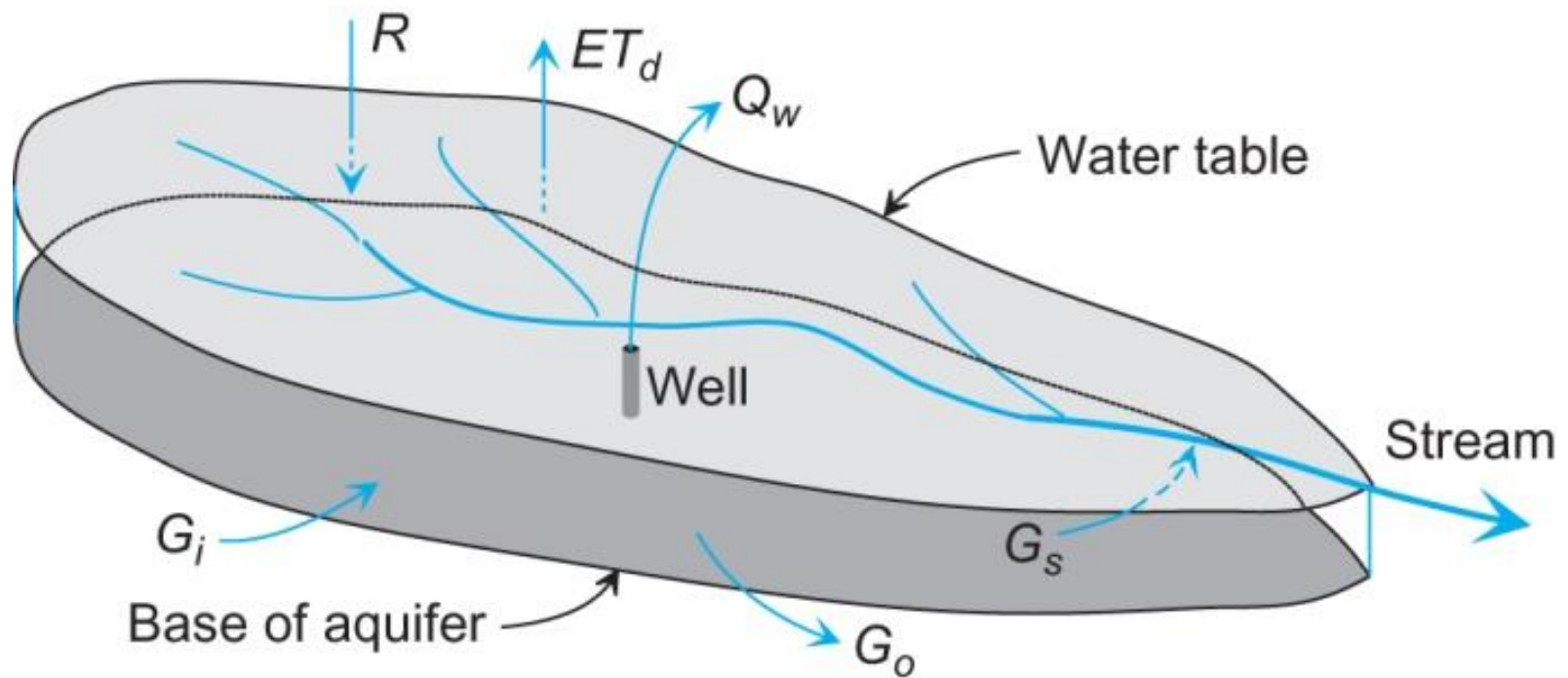


Fig. 4. Consolidation problem.

Hydrologic Balance



Hydrologic Balance (Contd.)

flux in – flux out = rate of change in water stored within

Fluxes into this region include the inlet stream flow (I) and the net groundwater discharge (G). Fluxes out of this region include the outlet stream flow (O) and evaporation (E) from the surface. Hydrologic balance in this case requires

$$I + G - O - E = \frac{dV}{dt}$$

This is for a particular problem only.
example 1.1

where dV/dt is the change in reservoir volume per time.

Hydrologic Balance (Contd.)

- Hydrologic Balance

$$R + G_i - G_o - G_s - ET_d - Q_w = \frac{dV}{dt}$$

- If, over a long time span, there is an approximate steady-state balance where flow in equals flow out, then the transient term disappears and the balance equation becomes

$$R + G_i - G_o - G_s - ET_d - Q_w = 0$$

Hydrologic Balance (Contd.)

$Q_w = 0$ describes the balance:

$$Q_w > 0 \quad Q_w = -dV/dt$$

R and G_i may increase, while G_o , G_s , and ET_d may decrease.

$$R + G_i - G_o - G_s - ET_d - Q_w = 0$$

Safe yield of aquifers

- The yield of an aquifer is defined as the rate at which water can be withdrawn without depleting the supply to such an extent that withdrawing at a rate higher than that is no longer economically feasible.

Groundwater Balance

- Input to the system – outflow from the system = Change in storage of the system

$$\Delta S = R_r + R_s + R_i + R_t + S_i + I_g - E_t - T_p - B_f - O_g$$

where:

ΔS is the change in groundwater storage

R_r is the recharge from rainfall

R_s is the recharge from canal seepage

R_i is the recharge from field irrigation

R_t is the recharge from tanks

S_i is the influent seepage from rivers

I_g is the inflow from other basins

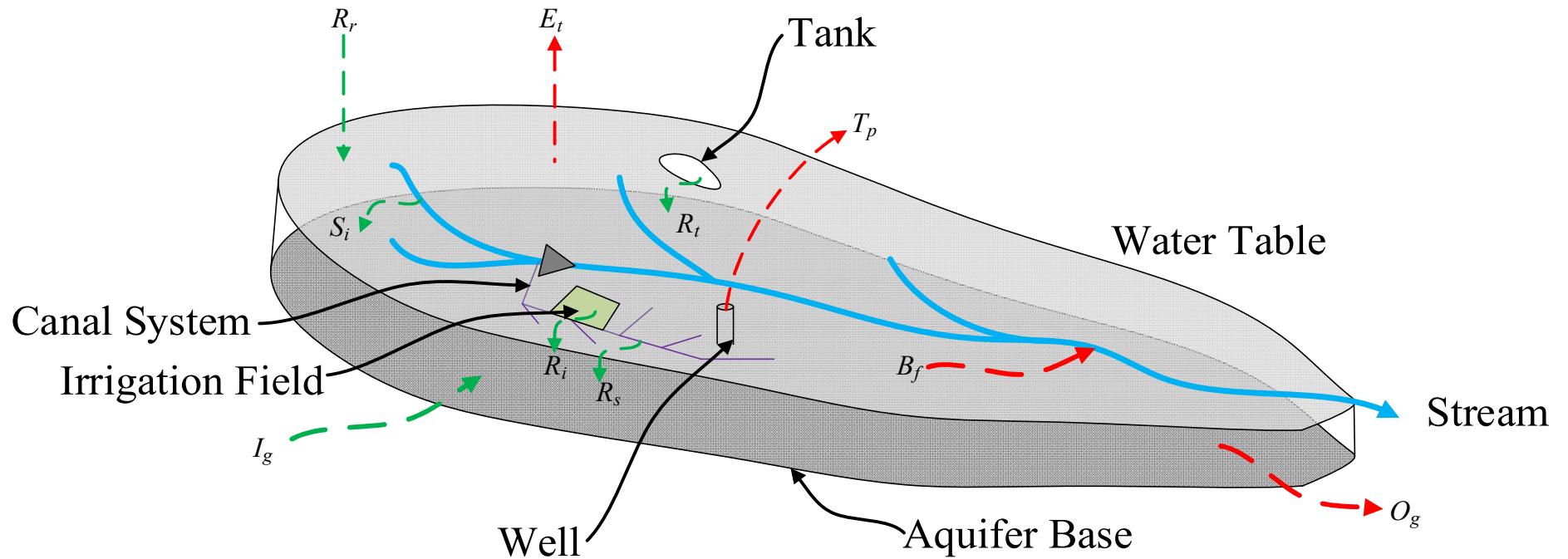
E_t is the evapotranspiration from groundwater

T_p is the draft from groundwater

B_f is the baseflow, the part of the groundwater inflows to rivers

O_g is the outflow to other basins.

Groundwater Balance (Contd.)



$$\Delta S = R_r + R_s + R_i + R_t + S_i + I_g - E_t - T_p - B_f - O_g$$

Data requirement

- Rainfall data
- Land use data and cropping patterns
- River stage and discharge data
- River cross sections
- Canals discharge and distributaries data, and the possible seepage from canals
- Tanks depth, capacity, area, and seepage data
- Water table data
- Groundwater draft (the number of each type of well operating in the area, their corresponding running hours each month, and their discharge)
- Aquifer parameters

GEC

Reprint

DETAILED GUIDELINES FOR IMPLEMENTING
THE GROUND WATER ESTIMATION
METHODOLOGY

GEC-97



CENTRAL GROUND WATER BOARD
MINISTRY OF WATER RESOURCES
GOVERNMENT OF INDIA
2009

REPORT OF THE
GROUND WATER RESOURCE ESTIMATION COMMITTEE
(GEC-2015)

METHODOLOGY

Ministry of Water Resources, River Development & Ganga Rejuvenation
Government of India

NEW DELHI
October, 2017

Water balance in confined aquifers

$$W_{sc,i} = W_{sc,i-1} + W_{per} - W_{pc}$$

where:

$W_{sc,i}$ is the amount of water stored in the confined aquifer on day i (mm)

$W_{sc,i-1}$ is the amount of water stored in the confined aquifer on day $i-1$ (mm)

W_{per} is the amount of water percolating from the unconfined aquifer into the confined aquifer on day i (mm)

W_{pc} is the amount of water removed from the confined aquifer by pumping on day i (mm).

Water balance in unconfined aquifers

$$W_{su,i} = W_{su,i-1} + R_r - B_f - W_{sd} - W_{per} - W_{pu}$$

where:

$W_{su,i}$ is the amount of water stored in the unconfined aquifer on day i (mm)

$W_{su,i-1}$ is the amount of water stored in the unconfined aquifer on day $i-1$ (mm)

R_r is the amount of recharge entering the aquifer on day i (mm)

B_f is the base flow to the main channel on day i (mm)

W_{sd} is the amount of water moving into the soil zone in response to water deficiencies on day i (mm)

W_{per} is the amount of water percolating from the unconfined aquifer into the confined aquifer on day i (mm)

W_{pu} is the amount of water removed from the unconfined aquifer by pumping on day i (mm).

Water balance in unsaturated zone

$$\Delta S_{\text{unsat}} = W_{\text{inf}} + W_{\text{cap}} + I_g - W_{\text{per}} - E_t - B_f$$

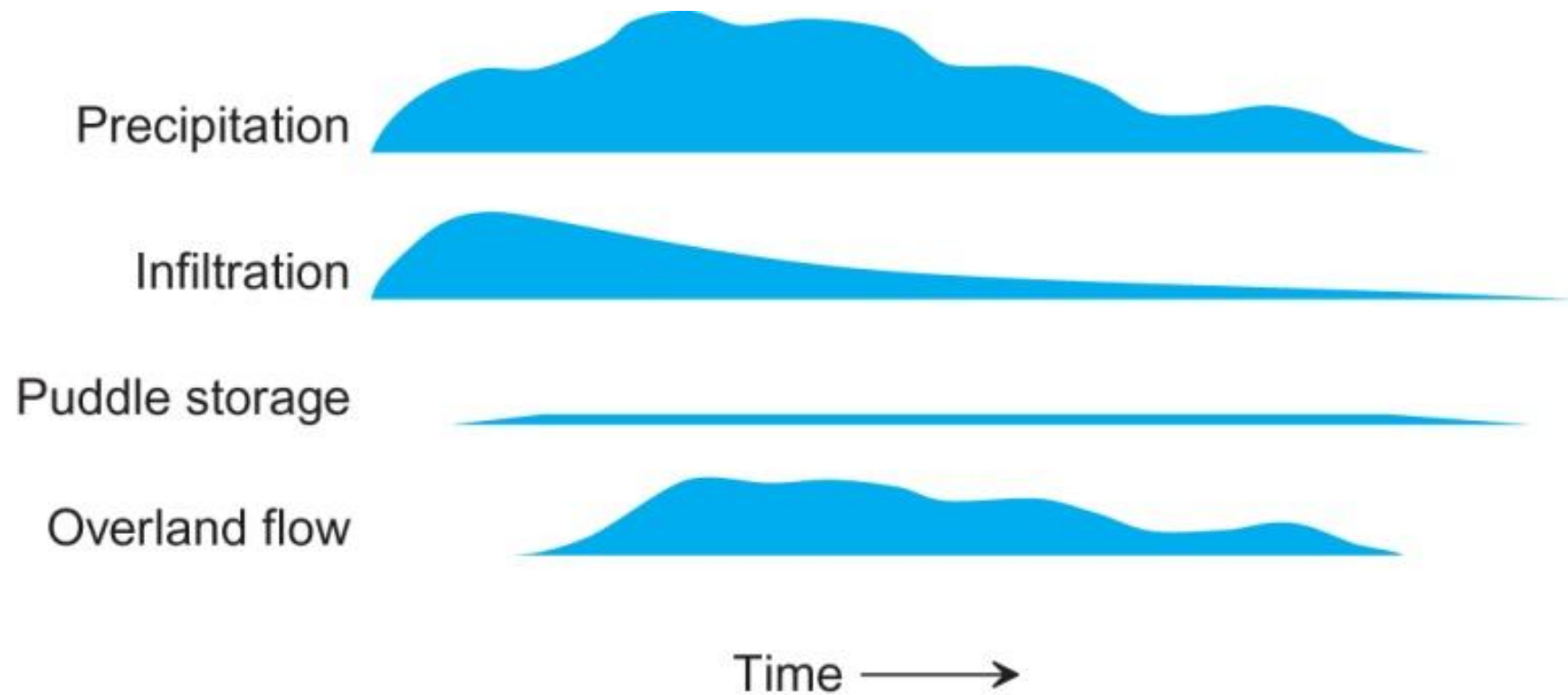
where:

ΔS_{unsat} is the storage change in unsaturated zone for the reflected time interval (mm)

W_{inf} is the amount of recharge entering the zone by infiltration (mm)

W_{cap} is the amount of recharge entering the zone by capillary rise (mm)

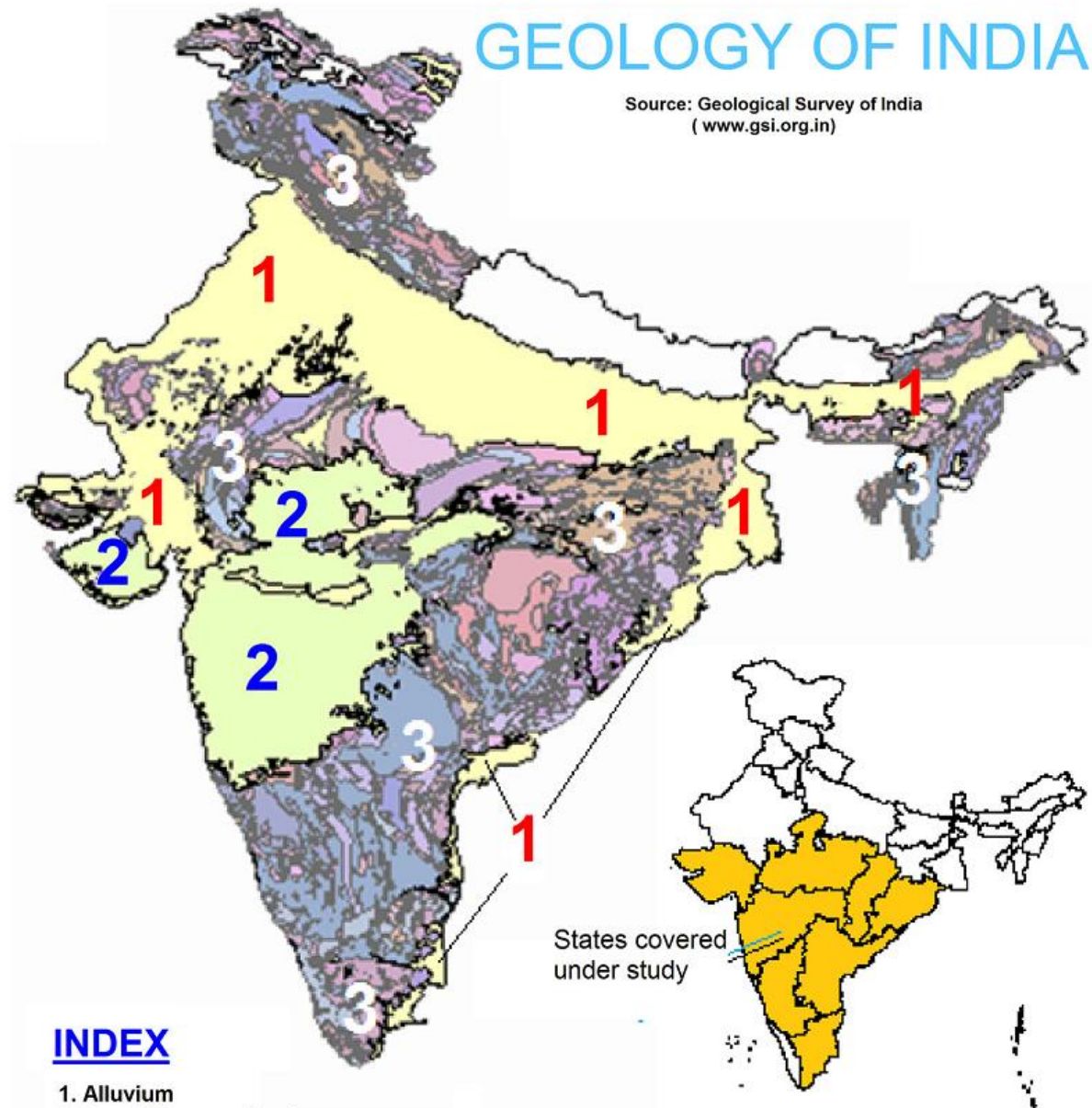
Infiltration and Recharge



INDIA

GEOLOGY OF INDIA

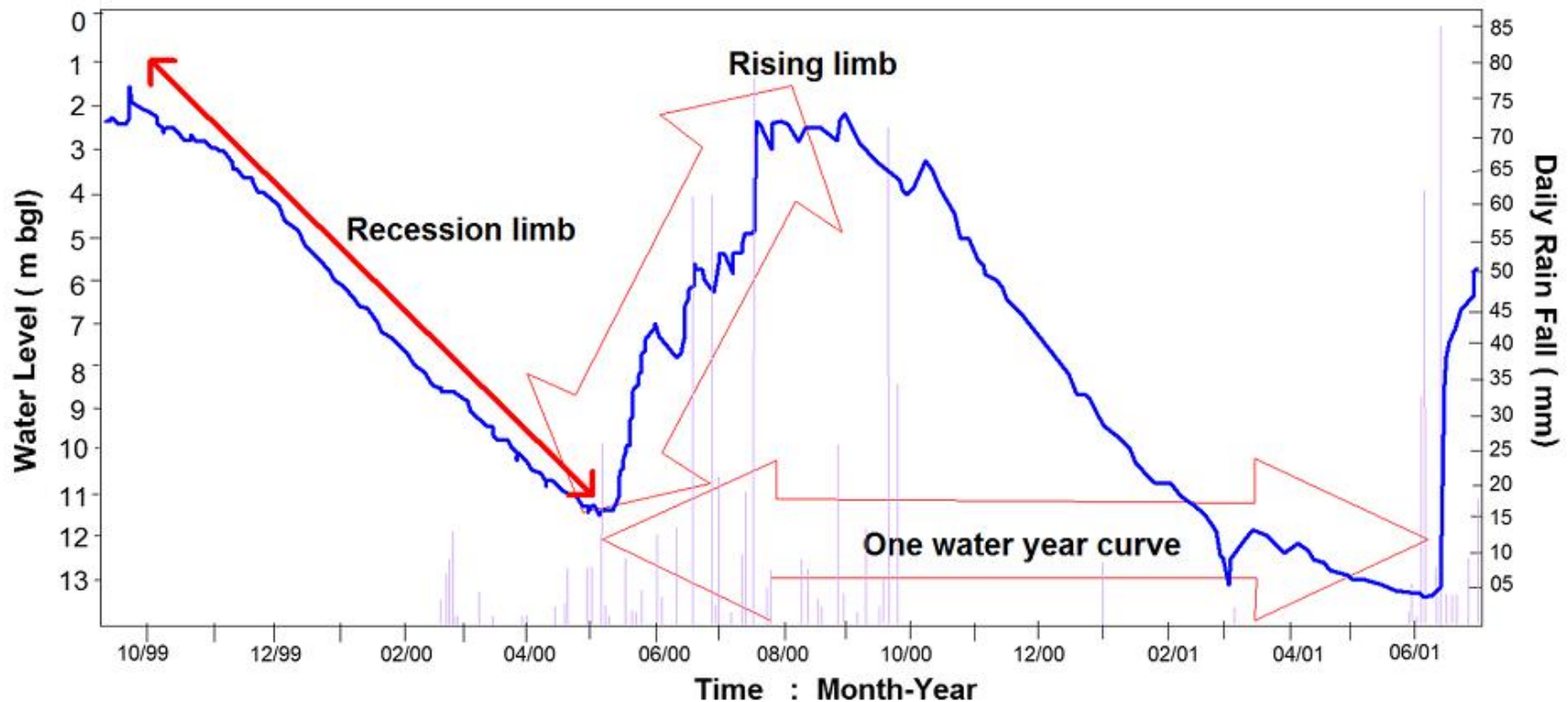
Source: Geological Survey of India
(www.gsi.org.in)



INDEX

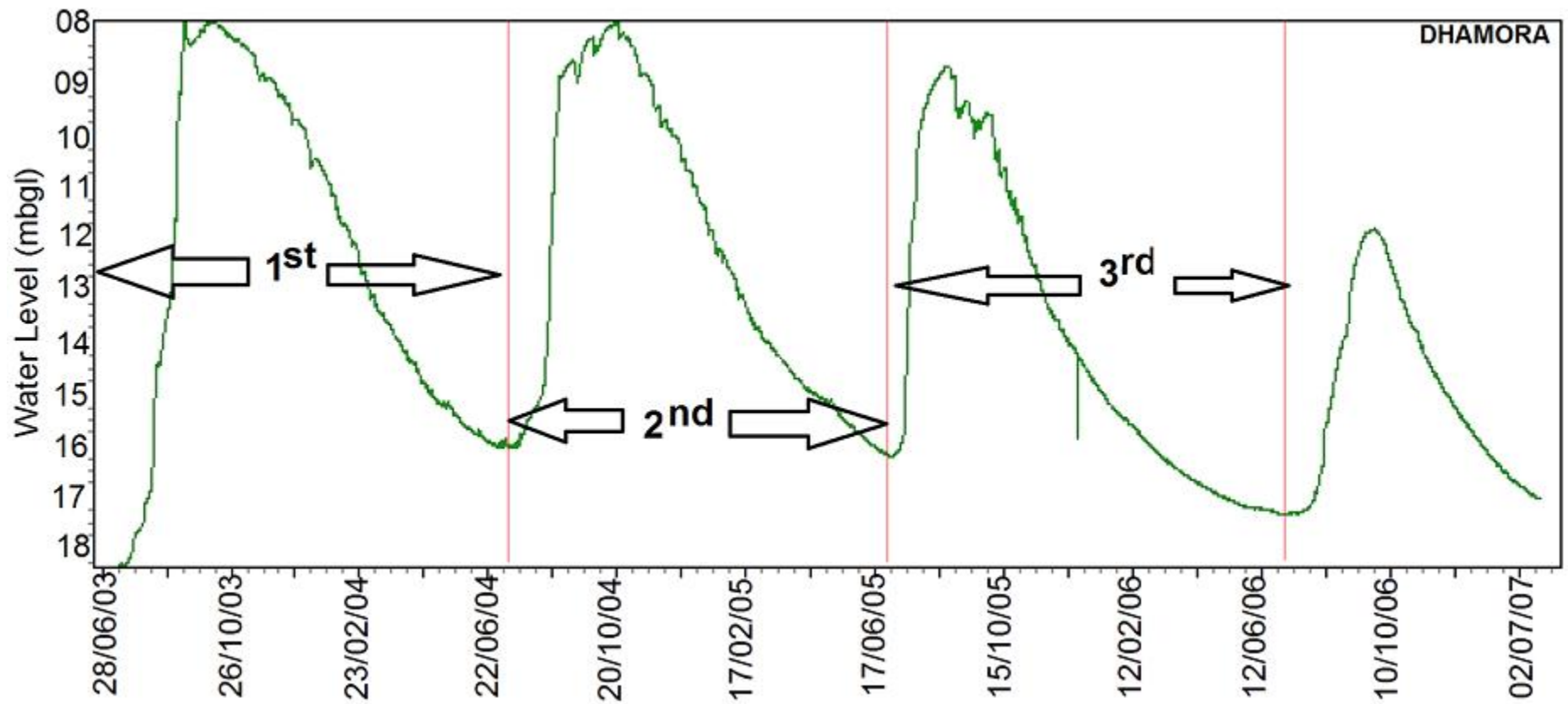
1. Alluvium
2. Basalt (Deccan trap)
3. Crystallines and other consolidated sedimentary rocks

Hydrograph

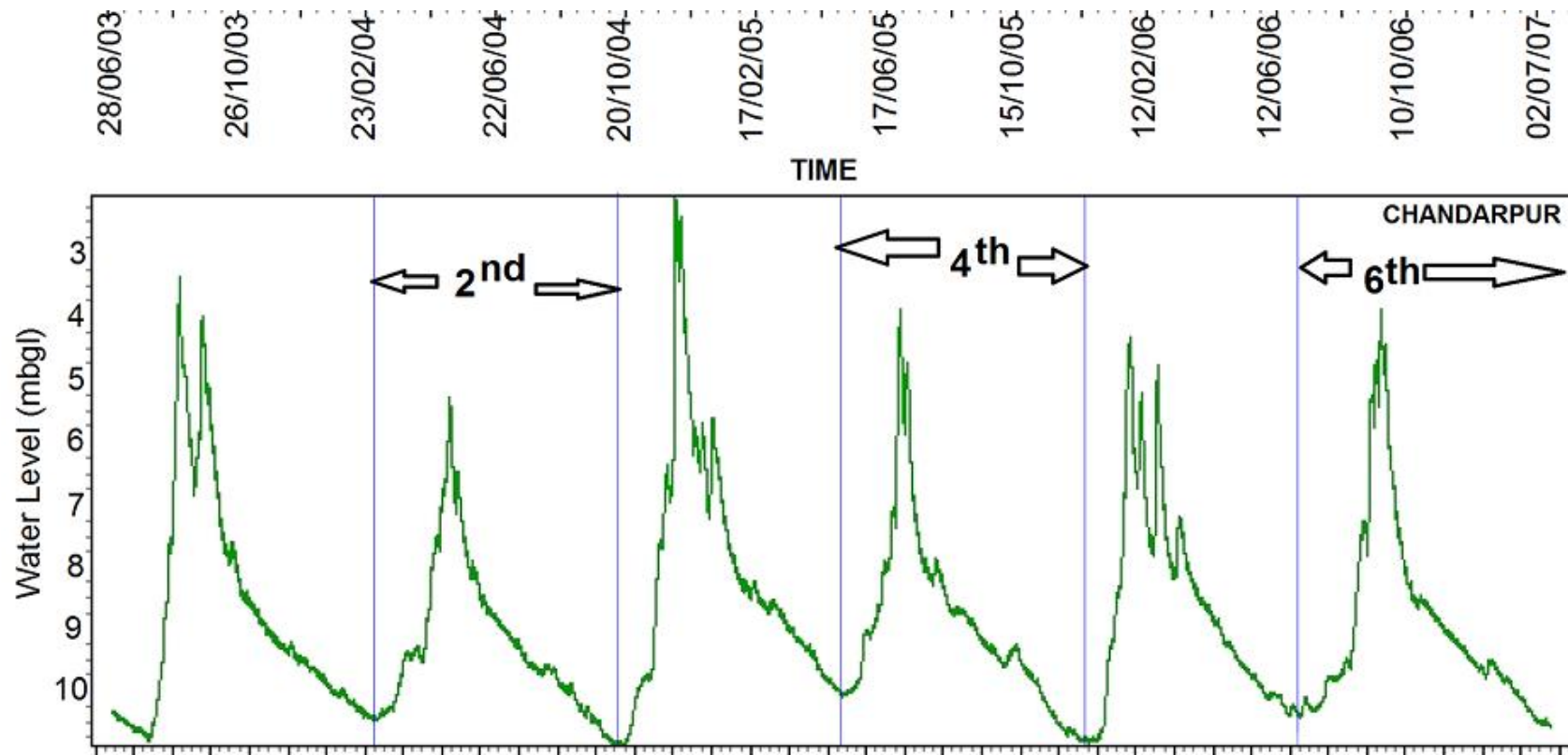


A typical true hydrograph generated from data collected by an automatic water level recorder fitted in a well (piezometer). Note the daily rainfall hyetograph and corresponding rise in water level during monsoon period. Location: Raipur

Hydrograph (Contd.)



Hydrograph (Contd.)

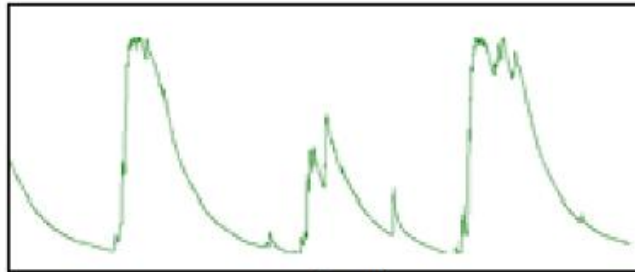


Recharge Phenomenon

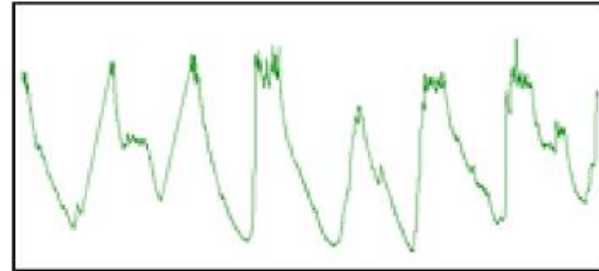
- In response to monsoon recharge groundwater level rises to its shallowest level, generating the rising limb.
- Based on the rainfall pattern and amount during monsoon, the peak takes its shape as per the aquifer character.
- Hydrograph crests can be described as pointed, flat or rounded in shape.
- The order of abundance of crest shape in the Peninsular hard rocks is pointed > flat > rounded
- The crest shapes are characteristic and do not change for a particular well/piezometer, however the magnitude or width may vary annually.
- Pointed crests are either single peak or multi peaked.
- The number of peak/peaks and its width varies annually based on rainfall pattern and quantity.

Crest of True Hydrograph

Pointed Crest

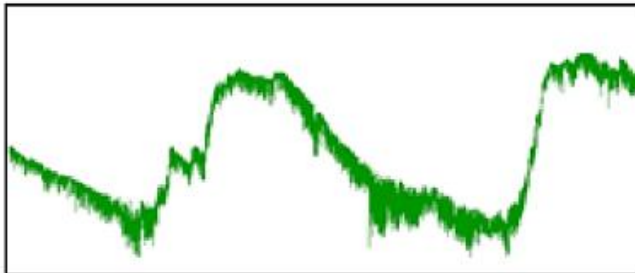


Boath

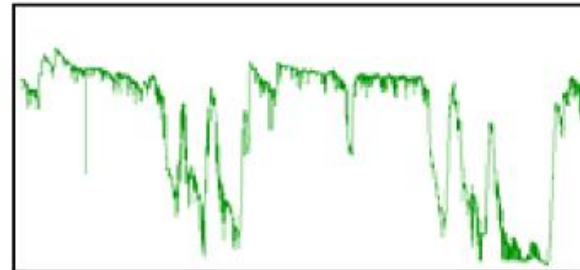


Dandepali

Flat top Crest

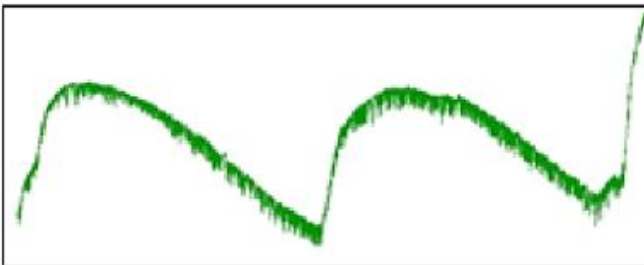


Tongpal



Belvandi

Rounded Crest

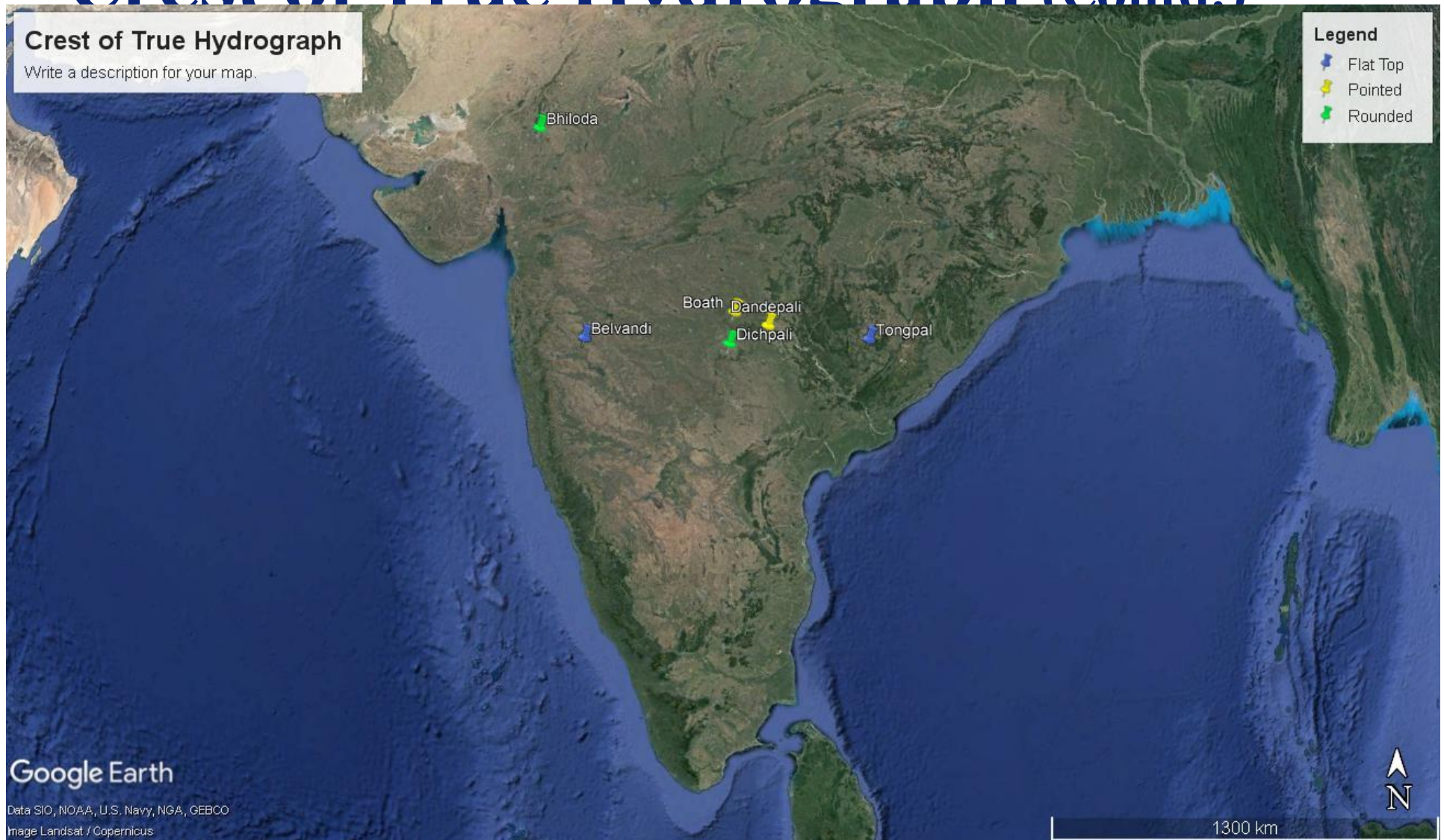


Bhiloda



Dichpali

Crest of True Hydrograph (Contd.)

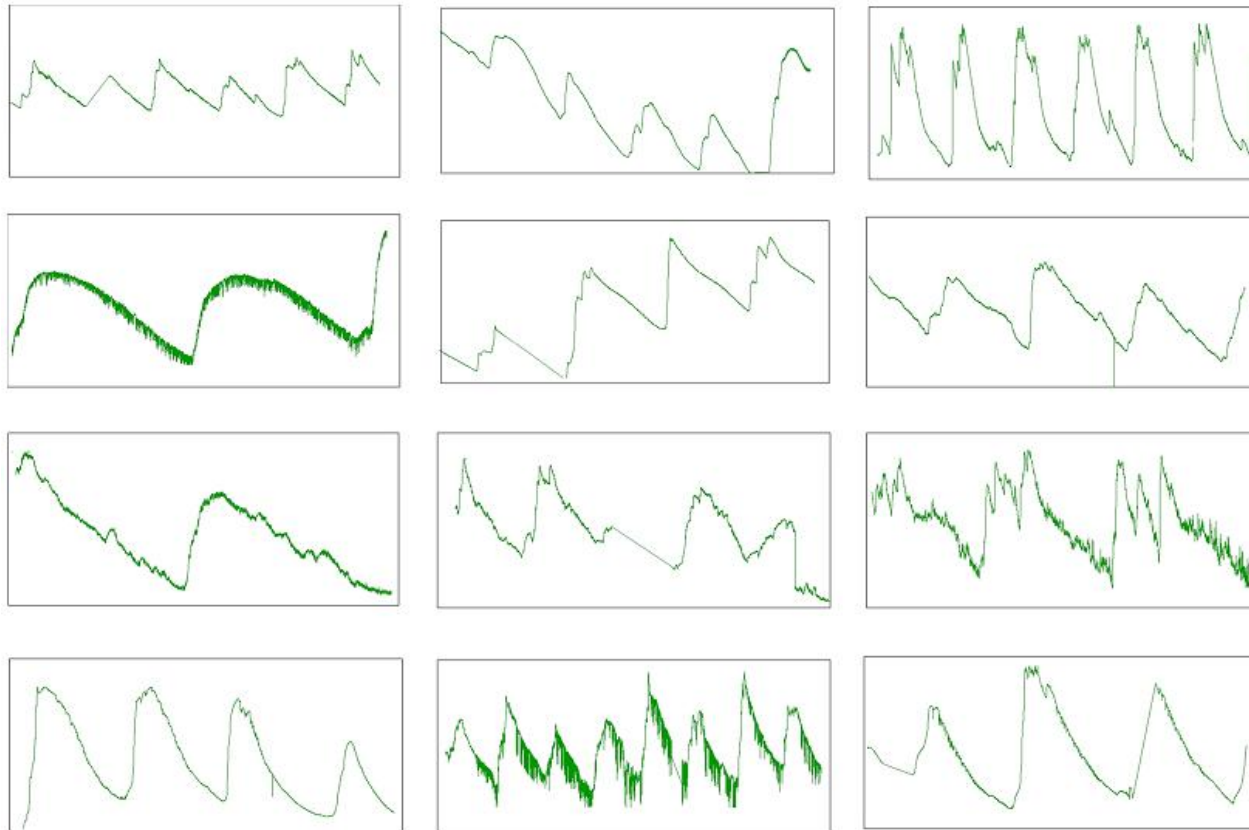


Discharge Phenomenon

- In response to the base flow and or the draft, the water level starts declining after the monsoon thus generating the recession limb.
- The combination of rising and recession limb pattern can be classified into three types, viz.
 - (a) V type curve
 - (b) U type curve
 - (c) S type curve.

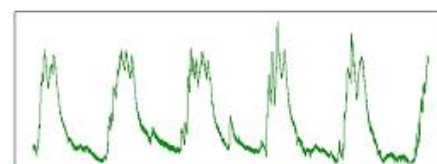
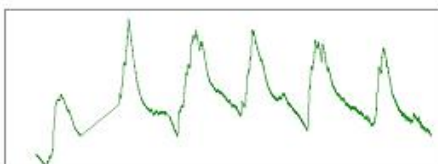
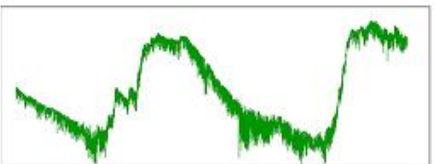
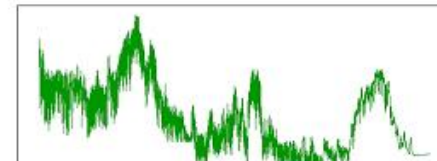
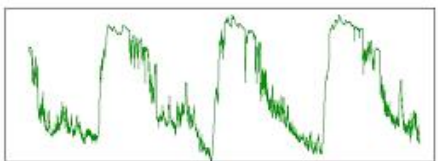
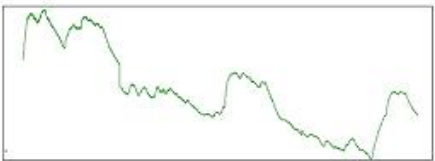
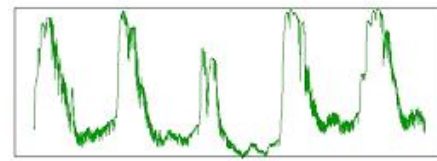
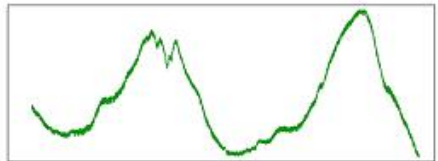
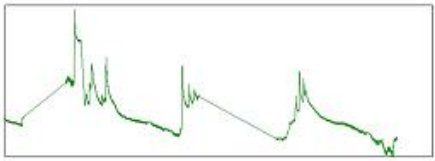
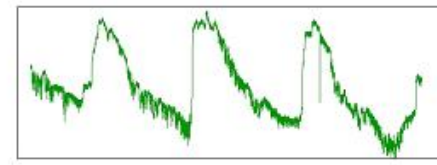
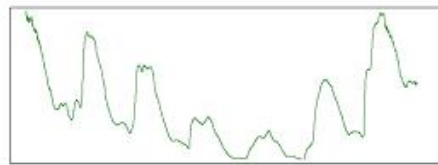
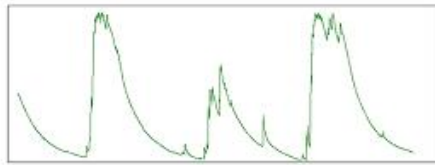
V type curves

- V type curves are those curves where both rising and recession limbs have nearly same slope angle though the rising limbs may sometimes be steeper than recession limbs. Also, they generally don't show any break in slope trend throughout the limb



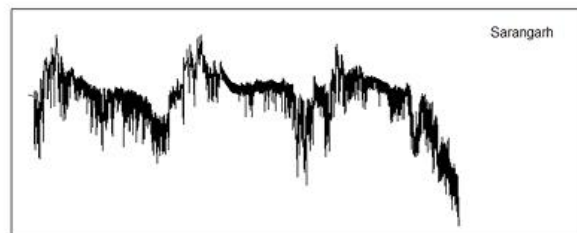
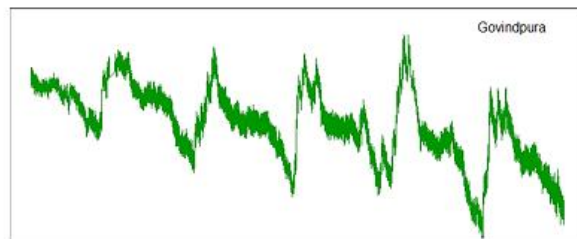
U type curves

- U type curves are those curves where rising limb has a nearly uniform slope, but the recession limb shows a definite break in the slope. The initial discharge is gentle followed by a fast phase of discharge and then again a slowdown giving rise to U shaped curve.
- Sometimes the slope is initially steep and subsequently gentle after some period of discharge and a combination of U and V type is developed. The lowest part of the trough is not very sharp in angle, rather it is nearly flat or curved.

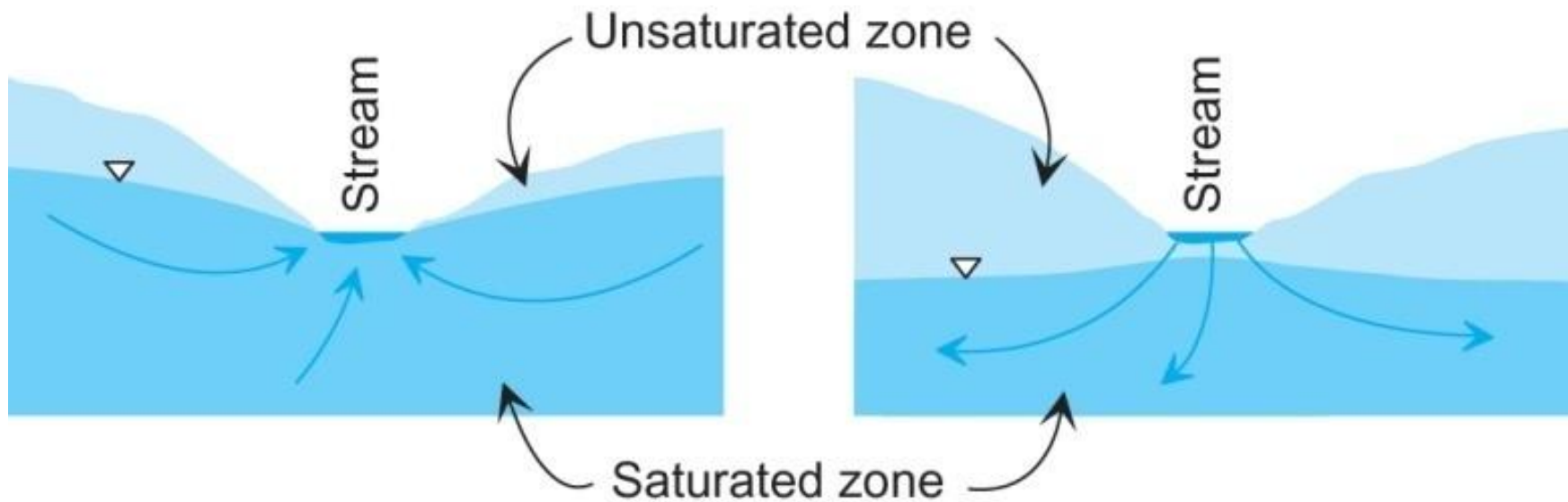


S type curves

- S type curves are those, which have sharp initial and ultimate-slope of recession limb and in between flat or very gentle slope giving rise to two breaks in the actual slope of the recession limb.
- The S type curves indicate low seasonal fluctuation and are not very common

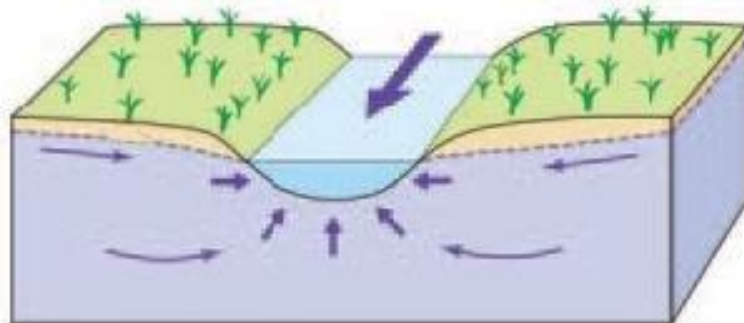


Groundwater Discharge to Surface Water Bodies

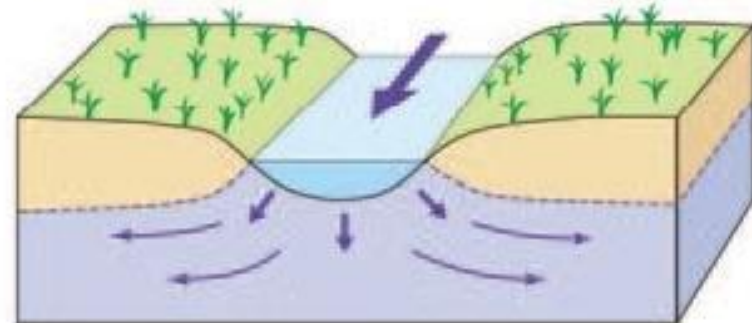


Stream-Aquifer Interaction

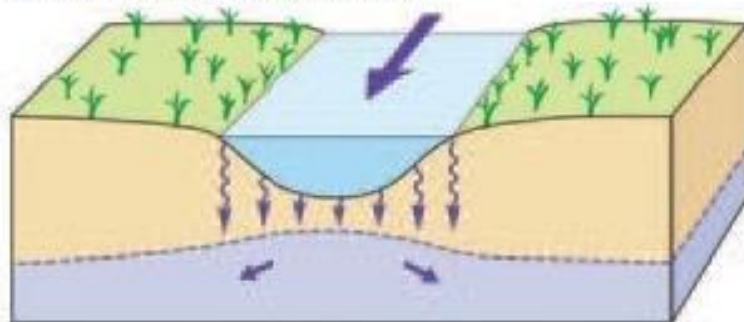
(a) Gaining stream



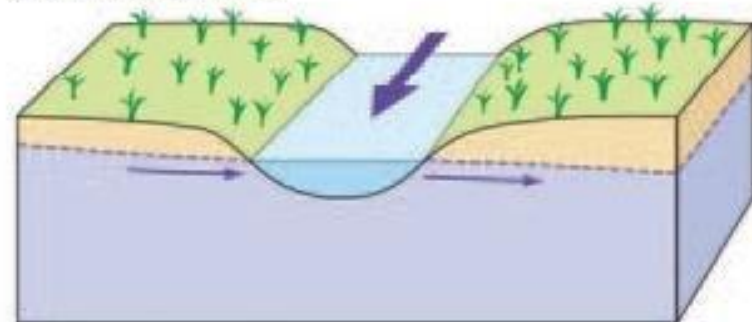
(b) Losing stream



(c) Disconnected stream



(d) Throughflow



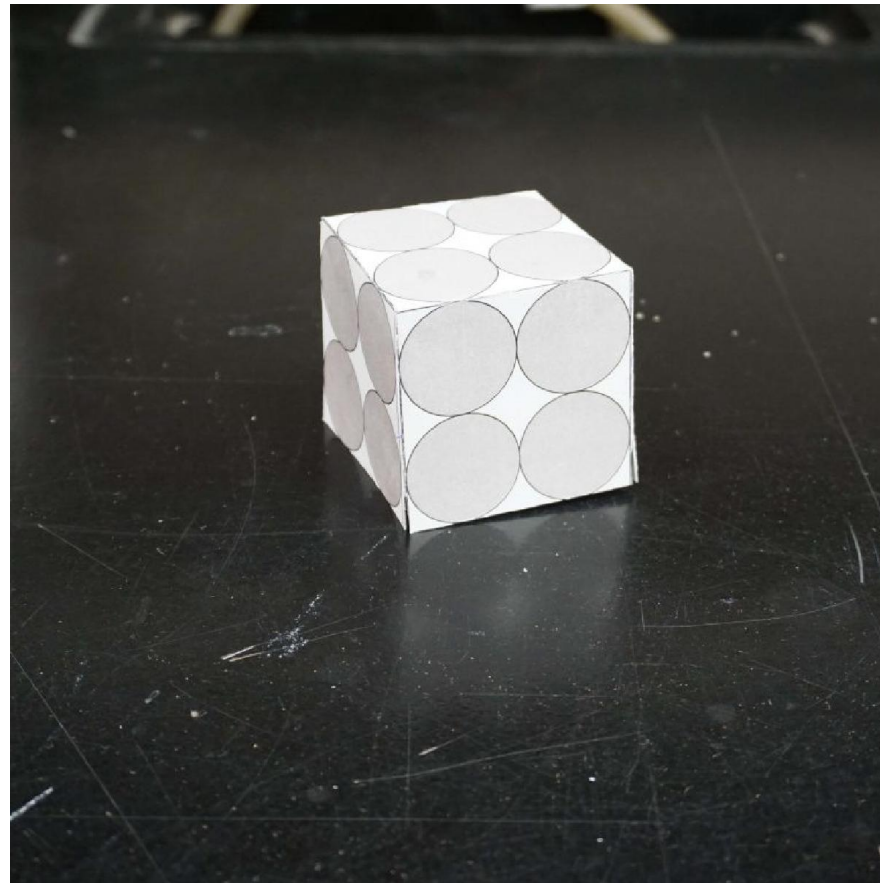
Land surface
Unsaturated zone

Shallow aquifer
Stream

Flow direction
Water table

Home Lab

- Foldable Aquifer Project -<http://aquifer.geology.buffalo.edu/>
- Paper aquifer model
 - Porosity and Grain Packing



Learning Strategy

Chapter 1: Groundwater: The Big Picture




Appl Water Sci (2017) 7:801–812
DOI 10.1007/s13201-015-0293-z



ORIGINAL ARTICLE

Aquifer response to recharge–discharge phenomenon: inference from well hydrographs for genetic classification

Arunangshu Mukherjee^{1,3}  · Anita Gupta¹ · Ranjan Kumar Ray¹ · Dinesh Tewari²

Thank you