

# Parameter Estimation

Geohydraulics| CE60113

Lecture:13

# Learning Objective(s)

- To estimate aquifer parameter under steady unconfined flow condition
- To estimate aquifer parameter under steady leaky confined flow condition

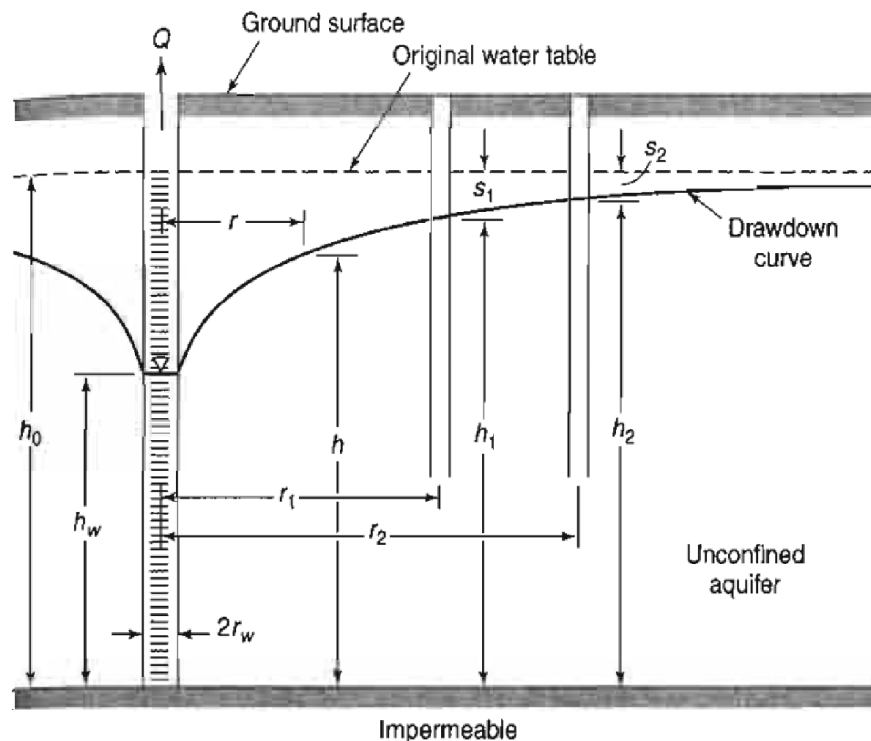
# Basic Aquifer Types

- Confined Aquifer
- Unconfined Aquifer
- Semiconfined/Leaky Aquifer

# Steady Radial Flow in Unconfined Aquifer

- Well discharge can be written as

$$Q = Aq_r = (2\pi rH) \left( -K_r \frac{\partial H}{\partial r} \right) = -2\pi r K_r H \frac{\partial H}{\partial r}$$



- Head difference is

$$H_0^2 - H_w^2 = \frac{Q}{\pi K_r} \ln \left( \frac{r_0}{r_w} \right)$$

- Discharge can be written as

$$Q = \pi K_r \frac{H_0^2 - H_w^2}{\ln \left( \frac{r_0}{r_w} \right)}$$

- Hydraulic conductivity

$$K_r = \frac{Q}{\pi(H_2^2 - H_1^2)} \ln \left( \frac{r_2}{r_1} \right)$$

# Steady Radial Flow in Unconfined Aquifer (Contd.)

$$\frac{2S_y}{K} \frac{\partial H}{\partial t} = \nabla^2 H^2 + \frac{2N}{K}$$
$$\frac{1}{r} \frac{\partial}{\partial r} \left( r \frac{\partial H^2}{\partial r} \right) = 0$$

Solution

$$H^2(r) = A \ln(r) + B$$

$$Q = -2\pi r H q_r(r) = 2\pi r K_r H \frac{\partial H}{\partial r} = \pi K_r A$$

$$H^2(r) = \frac{Q}{\pi K_r} \ln(r) + B$$

Radius of influence:  $H = H_0$  at  $r = r_0$

$$H^2(r) = H_0^2 - \frac{Q}{\pi K_r} \ln\left(\frac{r_0}{r}\right)$$

# Steady Radial Flow in Unconfined Aquifer (Contd.)

- Let us consider  $s_1 = H_0 - H_1$  and  $s_2 = H_0 - H_2$

$$H_1^2(r) = (H_0 - s_1)^2 = \frac{Q}{\pi K_r} \ln(r_1) + B$$

$$H_2^2(r) = (H_0 - s_2)^2 = \frac{Q}{\pi K_r} \ln(r_2) + B$$

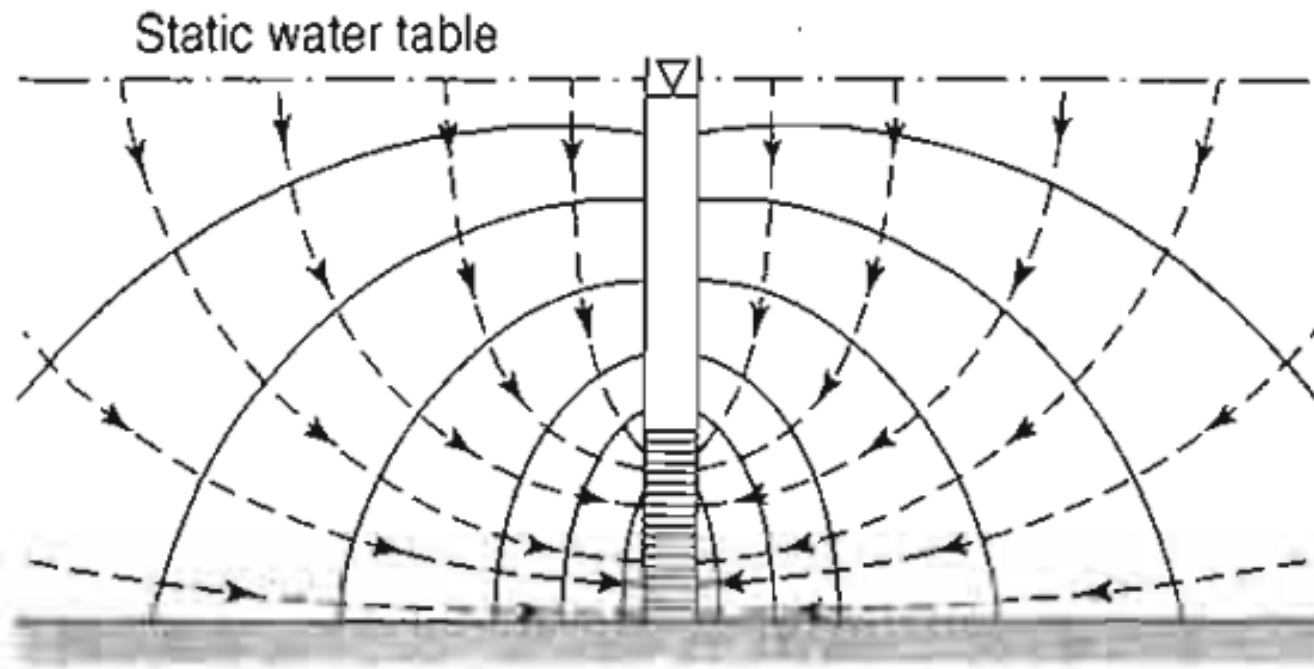
Subtracting the expressions

$$2H_0 \left[ \left( s_1 - \frac{s_1^2}{2H_0} \right) - \left( s_2 - \frac{s_2^2}{2H_0} \right) \right] = \frac{Q}{\pi K_r} \ln \left( \frac{r_2}{r_1} \right)$$

Transmissivity for the full thickness

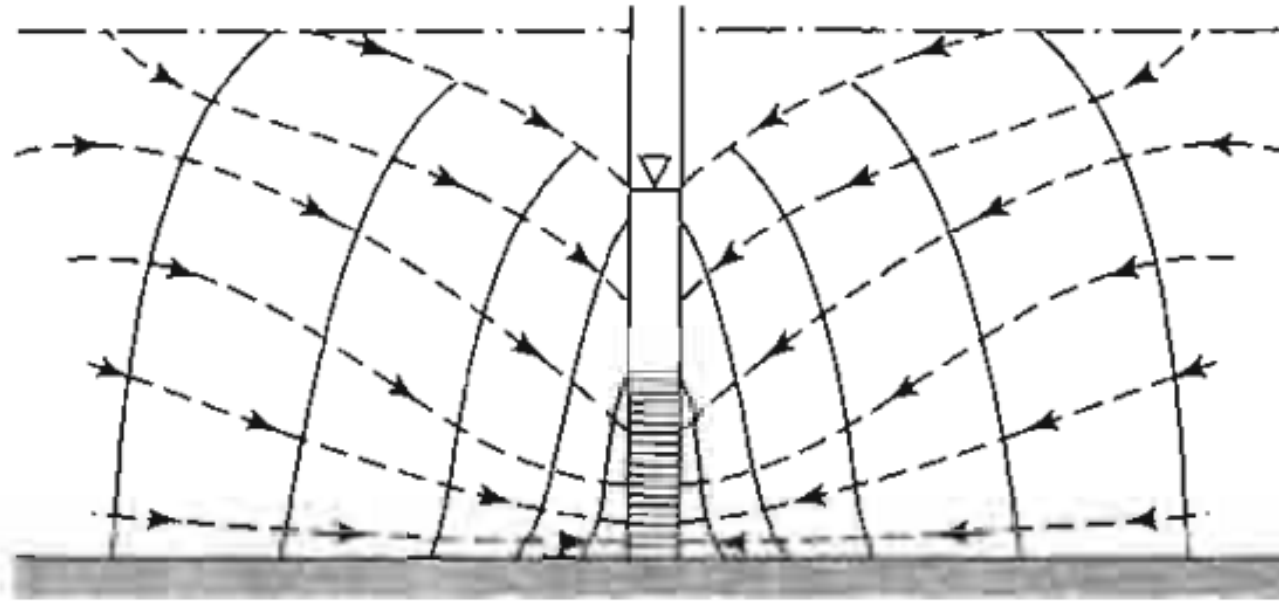
$$T = K_r H_0 = \frac{Q}{2\pi \left[ \left( s_1 - \frac{s_1^2}{2H_0} \right) - \left( s_2 - \frac{s_2^2}{2H_0} \right) \right]} \ln \left( \frac{r_2}{r_1} \right)$$

# Steady Radial Flow in Unconfined Aquifer (Contd.)



(a) Initial stage in pumping a free aquifer. Most water follows a path with a high vertical component from the water table to the screen.

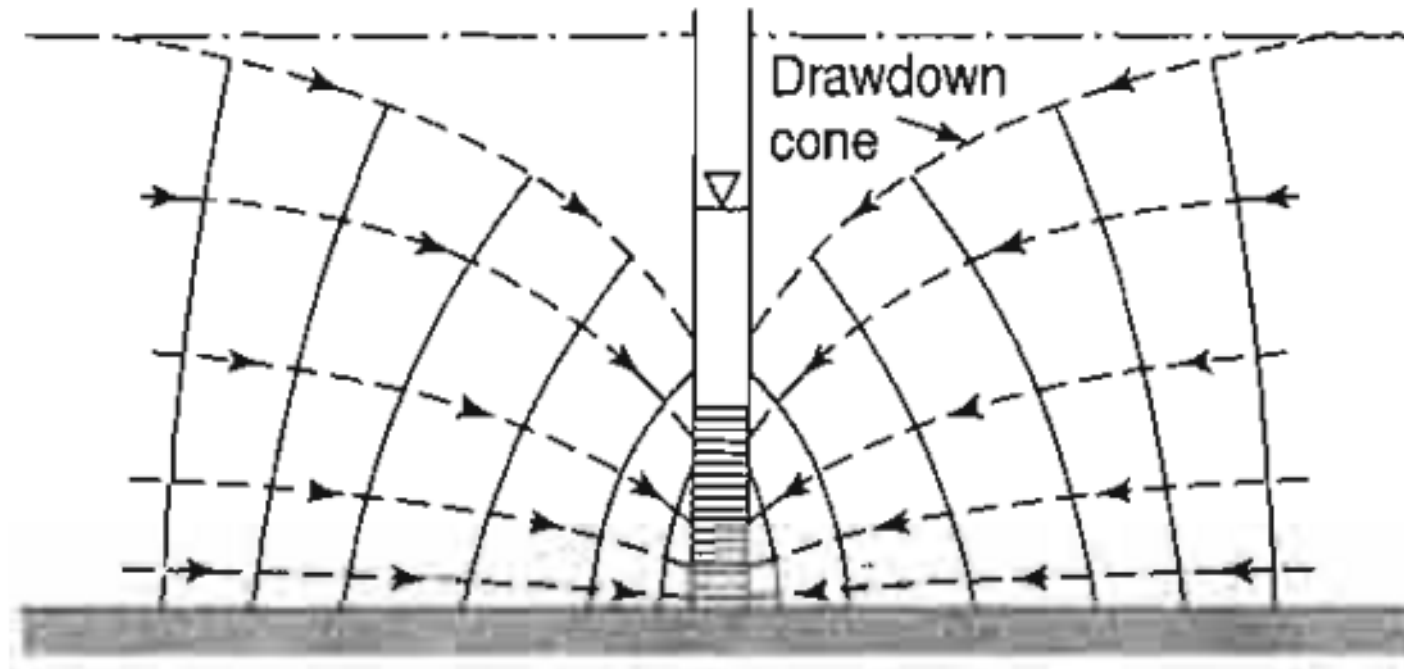
# Steady Radial Flow in Unconfined Aquifer (Contd.)



(b) Intermediate stage in pumping a free aquifer. Radial component of flow becomes more pronounced but contribution from drawdown cone in immediate vicinity of well is still important.



## Steady Radial Flow in Unconfined Aquifer (Contd.)



(c) Approximate steady-state stage in pumping a free aquifer. Profile of cone of depression is established. Nearly all water originating near outer edge of area of influence and stable primarily radial flow pattern established.

# Steady Radial Flow in Unconfined Aquifer (Contd.)

- River or surface water body (constant head)

$$H_0^2 - H^2(r) = \frac{Q}{\pi K_r} \ln \left( \frac{r_0}{r_1} \right) - \frac{Q}{\pi K_r} \ln \left( \frac{r_0}{r_2} \right)$$

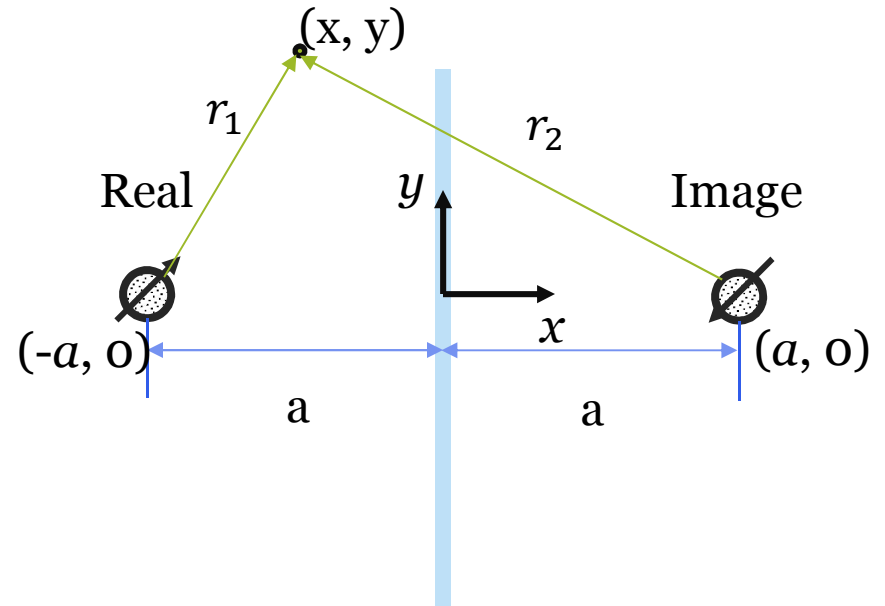
$$= \frac{Q}{\pi K_r} \ln \left( \frac{r_2}{r_1} \right)$$

$$r_1 = \sqrt{(x + a)^2 + y^2}$$

$$r_2 = \sqrt{(x - a)^2 + y^2}$$

Along  $y$ -axis,  $r_1 = r_2$

$$H^2 = H_0^2$$



# Steady Radial Flow in Unconfined Aquifer (Contd.)

- Impermeable faults (no-flow boundaries)

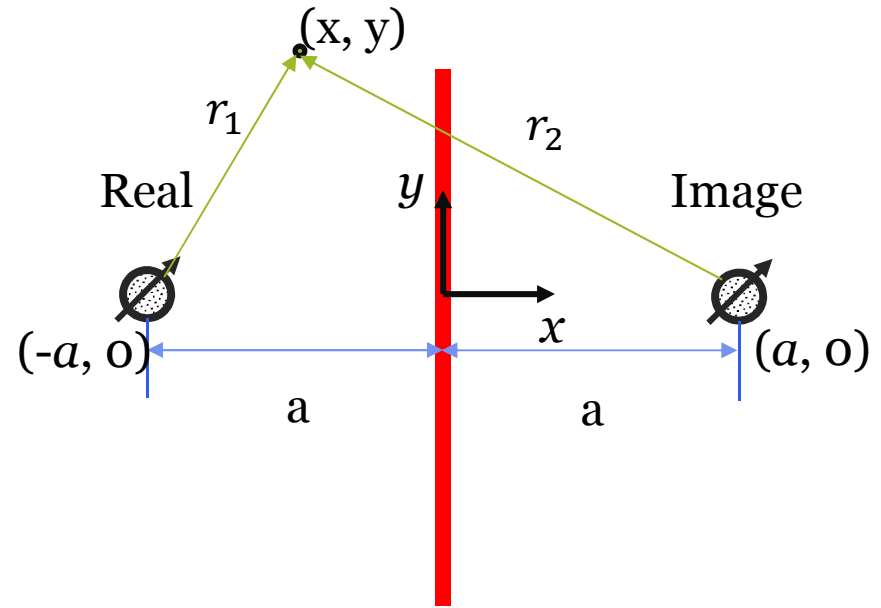
$$\begin{aligned}
 r_1 &= \sqrt{(x+a)^2 + y^2} \\
 r_2 &= \sqrt{(x-a)^2 + y^2} \\
 H^2(r) &= \frac{Q}{\pi K_r} \ln(r_1) + \frac{Q}{\pi K_r} \ln(r_2) + B \\
 &= \frac{Q}{\pi K_r} \ln(r_1 r_2) + B
 \end{aligned}$$

x-component of flow

$$\begin{aligned}
 U_x &= H q_x = -KH \frac{\partial H}{\partial x} = -\frac{K}{2} \frac{\partial H^2}{\partial x} \\
 &= -\frac{Q}{2\pi} \left( \frac{d}{dr_1} (\ln(r_1)) \frac{\partial r_1}{\partial x} + \frac{d}{dr_2} (\ln(r_2)) \frac{\partial r_2}{\partial x} \right) \\
 &= -\frac{Q}{2\pi} \left( \frac{1}{r_1} \frac{x+a}{r_1} + \frac{1}{r_2} \frac{x-a}{r_2} \right) \\
 &= -\frac{Q}{2\pi} \left( \frac{x+a}{(x+a)^2 + y^2} + \frac{x-a}{(x-a)^2 + y^2} \right)
 \end{aligned}$$

Along  $x=0$ , the aquifer flux

$$U_x(0, y) = -\frac{Q}{2\pi} \left( \frac{0+a}{(0+a)^2 + y^2} + \frac{0-a}{(0-a)^2 + y^2} \right)$$



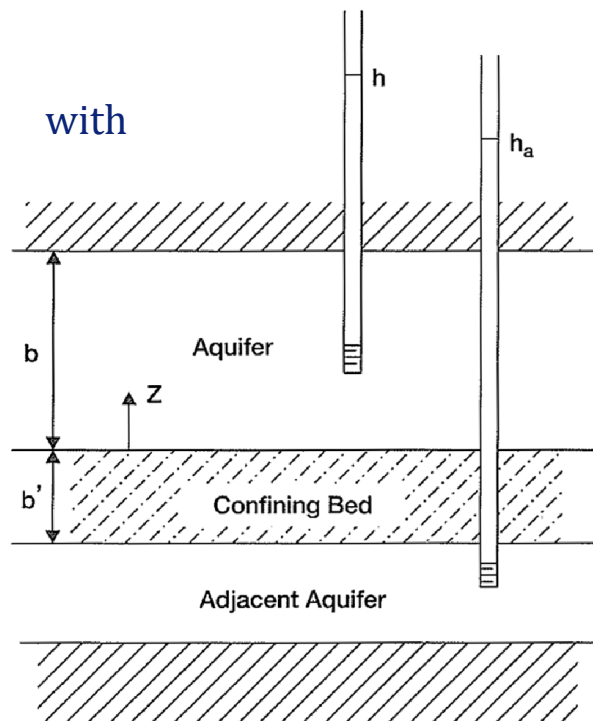
# Steady Radial Flow in Leaky Confined Aquifer

$$S \frac{\partial \bar{h}}{\partial t} = \frac{1}{r} \frac{\partial}{\partial r} \left( T_r \frac{\partial \bar{h}}{\partial r} \right) - \frac{K'}{b'} (\bar{h} - \bar{h}_a)$$

Under steady state radial flow condition in semiconfined aquifer

$$T \frac{1}{r} \frac{\partial}{\partial r} \left( r \frac{\partial h}{\partial r} \right) - \frac{K'}{b'} (h - h_a) = 0$$

$$\frac{d^2 h}{dr^2} + \frac{1}{r} \frac{dh}{dr} - \frac{h}{B^2} = -\frac{h_a}{B^2}$$



$$B = \sqrt{\frac{Tb'}{K'}}$$

- Bessel's modified differential equation of order zero

$$Q = -2\pi r b q_r(r) = 2\pi r K_r b \frac{\partial h}{\partial r}$$

Solution

$$s(r) = h_a - h(r) = \frac{Q}{2\pi T} K_0 \left( \frac{r}{B} \right)$$

$K_0 \left( \frac{r}{B} \right)$  is modified **Bessel function** of the second kind of order zero

# Steady Radial Flow in Leaky Confined Aquifer (Contd.)

- Transmissivity

$$T = \frac{Q}{2\pi s} K_0 \left( \frac{r}{B} \right) = \frac{Q}{2\pi s} K_0 \left( \frac{r}{\sqrt{T \frac{b'}{K'}}} \right)$$

- Nonlinear function can be written as

$$F(T) = T - \frac{Q}{2\pi s} K_0 \left( \frac{r}{\sqrt{T \frac{b'}{K'}}} \right) = 0$$

$$\frac{dF}{dT} = 1 - \frac{Q}{4\pi s} \frac{\left( r \frac{b'}{K'} \right)}{\left( T \frac{b'}{K'} \right)^{3/2}} K_1 \left( \frac{r}{\sqrt{T \frac{b'}{K'}}} \right)$$

- For  $p^{th}$  iteration

$$T^{(p)} = T^{(p-1)} - \frac{F(T^{(p-1)})}{\left( \frac{dF}{dT} \right)^{(p-1)}}$$

where  $p \geq 1$

**Thank you**