Outline

Intro

The timing and season of pedogenic carbonate precipitation has been debated in recent years, with significant consequences to paleoclimatic inference (a few citations). Clumped isotope temperatures of modern pedogenic carbonates have found a variety of formation temperatures, relative to MAT, in different environments (Mintz, Peters, Quade, Gallagher). Usually, these temperatures are at or above MAT, indicating the well-cited warm-season bias in most environments. The main drivers of pedogenic carbonate formation are increasing temperatures, decreasing *p*CO2, root-water uptake, and evaporation; however, the relative importance of these drivers is debated (Quade, Gallagher, Meyer). Some studies invoke soil drying, lowering soil *p*CO2, while other studies suggest transpiration could play a larger role in concentration Ca ions the growing season (Gallagher, Meyer). Modelling suggests that pedogenic carbonate precipitates and dissolves in both short and long time periods, but the last precipitation event of the year is usually best preserved (Meyer). Overall, the main season and timing of pedogenic carbonate formation remains unresolved and likely varies with C3 vs. C4 vegetation and different precipitation patterns. Here, we create a model that predicts oxygen and carbonate isotope values of pedogenic carbonate in C3 ecosystems with arid to sub humid climates. We then compare predicted model results with modern studies of pedogenic carbonate that include clumped isotope temperatures to gain insight into the season and mechanisms of carbonate precipitation.

Methods

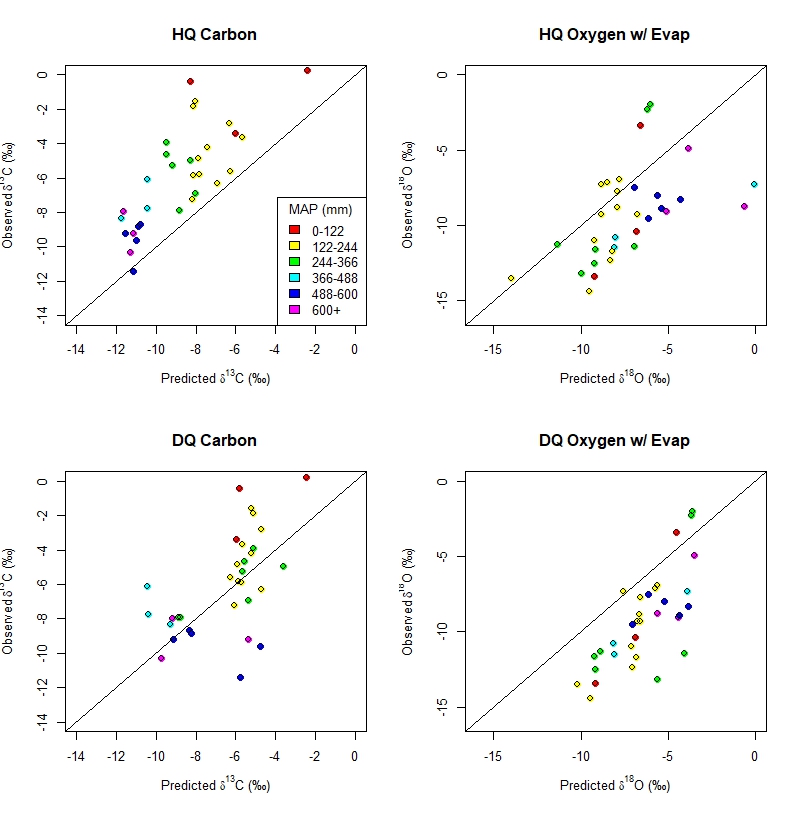
We assume pedogenic carbonate formation takes place over a 3-month season (DJF, MAM, JJA, SON), hereby referred to as the pedogenic carbonate quarter (PCQ). Here we explore a complimentary, model-based approach to evaluating the conditions of soil carbonate formation. We create a mechanistic model connecting carbon and oxygen isotope values of pedogenic carbonate to both seasonal and annual climatic drivers. We explore the effectiveness of the model with different model assumptions and its implications for the timing and mechanisms of pedogenic carbonate formation. The result is a set of hypotheses, which are discussed in depth in the next section.

Results/Discussion

B: Hot and/or dry conditions have been suggested to facilitate pedogenic carbonate precipitation.

Q: Does the modelled d13C and d18O values better match measured d13C and d18O values in hot or dry conditions?

A: Run model using hot quarter and dry quarter conditions.



R: Underpredicts d13C for hot quarter conditions, overpredicts d18O for dry quarter conditions. Does not include evaporative effects on d18O

I: No clear support for one season over the other. More revisions need to be developed in the model to fully match either hot or dry season.

\*Side Note\*

B: Given that the model does not fully resolve season of formation, clumped temperatures might be useful to help inform the season of carbonate formation.

Q: Do clumped temperatures better match warm season or dry season temperatures?

A: Compare gridded climate temperatures of hot and dry seasons to clumped temperatures.

R: Clumped temperatures best match hot season temperatures.

Avg. offsets:

clumped: 17.71 +/- 12.34 C (Kelson 2017 calibration)

dry: -0.22 +/- 5.72 C

hot: 7.02 +/- 1.61 C

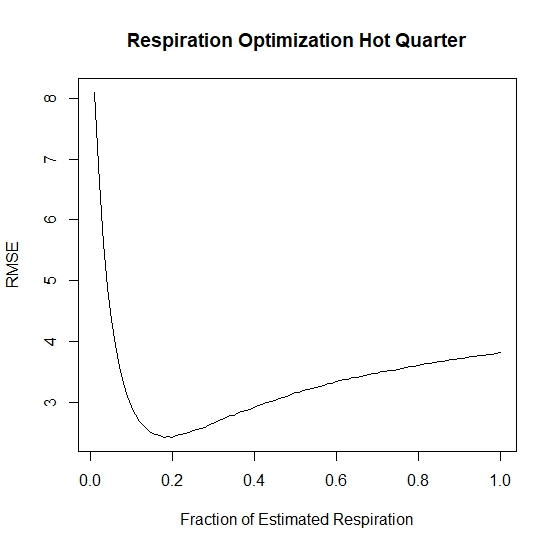
I: Clumped temperatures seem to support hot season precipitation of carbonate

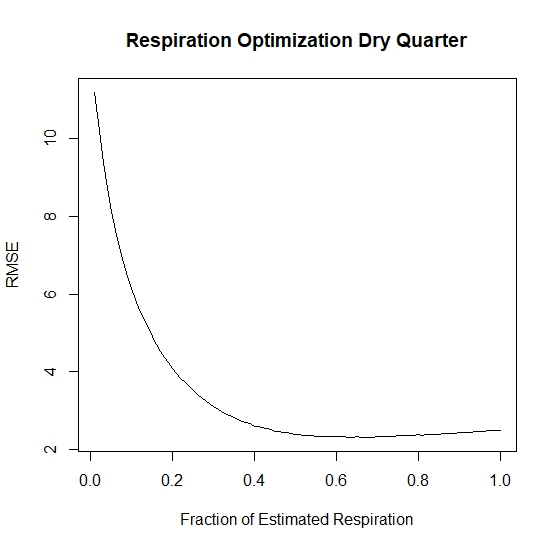
B: Decrease in respiration rate has been suggested as a mechanism of carbonate precipitation

Q: If hot season is assumed to be the season of formation, can we explain the underestimation of d13C values with lower respiration rates?

A: Run the model at successively decreasing fractions of the estimated respiration (respiration ratio).

R: RMSE of measured vs. predicted d13C values is lowest at ~20% of the estimated respiration rate based on mean climate.





I: Respiration rates could decrease at times within the hot season, leading to selective precipitation of pedogenic carbonate at those times, which would not match mean respiration rate over that season. The decrease of soil pCO2 could be the main driver of carbonate precipitation during the hot season.

\* Side Note \*

B: Lower MAP could affect the relationship between soil respiration and climate (temperature and precipitation).

Q: Does filtering respiration sites for lower MAP affect the relationship between respiration rate and climate?

A: Run the JAGS model for the filtered sites (annual respiration, MAP < 760 mm)

R: This filtering affects it only slightly, with a small increase in the effect of precipitation and temperature and small decrease in basal respiration rate. K from 4.25 to 4.87 and Q from 0.054 to 0.055. R0 from 1.25 to 1.24.

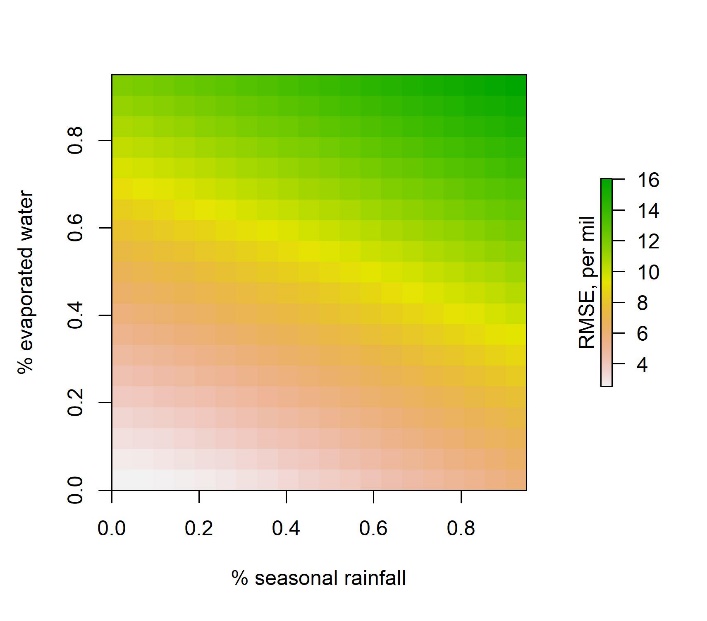
I: The very small change in these parameters does not appreciably affect any results so the relationship holds at low precipitation regimes.

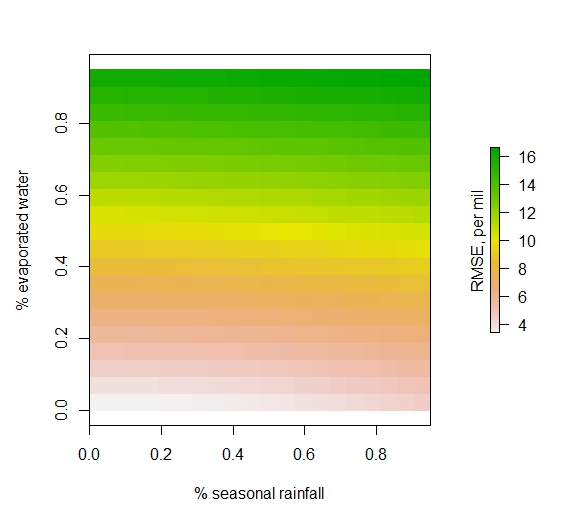
B: The effects of evaporation and seasonal rainfall on d18O values of pedogenic carbonate have been proposed to complicate the relationship between d18O of mean annual precipitation and d18O of pedogenic carbonate.

Q: Do evaporation and/or seasonal rainfall significantly affect d18O values of pedogenic carbonate?

A: Run the model with successively increasing amounts of seasonal rainfall bias and evaporative effects.

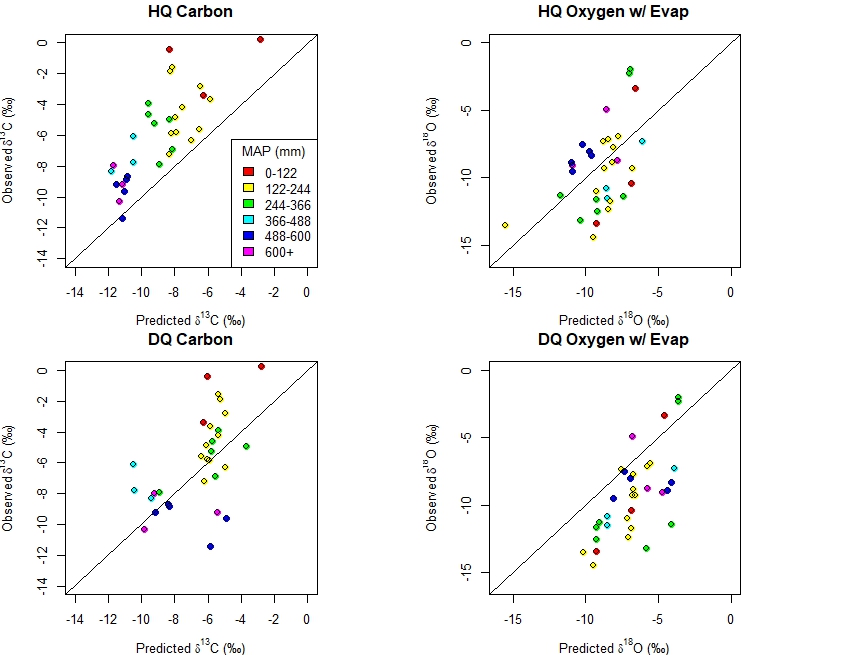
R: RMSE of predicted vs. measured d18O is lowest with no seasonal rainfall bias and no evaporative effects.





I: Therefore, d18O seems to be most influenced by MAT and temperature of formation, with minimal influence of evaporation or seasonal rainfall.

Comparison w/ no evap



Transpiration leading to upward water movement may be more important than evaporation? (e.g. Meyer). Or at 50cm or below, as we have selected samples, evaporation is not prevalent.

Maybe we should add z as an input into the model rather than dependent on MAP.

This method won’t always work in paleoclimate; however, you could put large error bars on current estimates and try to calculate A horizon erosion into the equation.

Working on this

- Both of these processes enrich oxygen isotopes of pedogenic carbonate. If we look at the HQ and DQ plots comparing measured vs. predicted oxygen isotope values, HQ is normally distributed, but the DQ is biased toward higher values, not lower values. Therefore, neither of these processes would correct for this bias.

CO2 degassing as a mechanism of pedogenic carbonate formation

* Soil pCO2 likely too high in the model
  + Optimize the respiration equation in the model to test if consistently lower respiration rates could explain the difference between modelled and observed carbon isotope values of pedogenic carbonate given certain climatic conditions.
* Respiration ratio (fraction of respiration from equation) vs. RMSE
  + This seems to indicate that actual respiration rate at the time of precipitation is about 20 – 25 % (depending on clumped vs. all data) of estimated respiration rate in the warm quarter of the year.
* Could the model equation for respiration be biased for these conditions?
  + Extracted sites from the large annual respiration database where MAP < 760, then re-fit the regression using rjags. The resulting equation constants were K = 4.87 Q = 0.55 Ro = 1.24. When the HQ model is run using this regression, the RMSE for data-model comparison reduced very slightly by ~ 0.1.
    - Working on a plot to compare the parameters with successively lower MAP cutoffs. Then, create a plot with the parameters vs. lower MAPs.
* Discussion points/implications:
  + Soil pCO2 decreasing by ~80% compared to mean rates during that season may be an important driver of carbonate mineralization.
  + Low MAP sites may have slightly different parameters in the respiration rate equation; however, the difference is not enough to matter with respect to the respiration ratio results.

One interesting thing that I have been thinking, which may be out of the scope of this particular paper:

Most of these modern sites are still very low MAP (even when culling for >100mm)

Seasonal precipitation patterns likely drive carbonate formation in most envrs

For paleo-applications, at least in our PETM sites, MAP is predicted to be much higher. So, how are we creating a good model for paleo-application by testing the model with these arid to hyper-arid sites? Seasonal patterns of precipitation and carbonate formation are likely different in these sites vs. higher MAP sites.

What if we attempt to use data only from MAP > 400 mm that still have carbonate in the modern envr? There may not be many of them, but maybe would be a better analog, even if they don’t have clumped isotope temps. There could be a significant difference or “tipping point” of when sites precipitate carbonate in the summer vs. fall/spring controlled by MAP/soil water balance (which is temp, soil texture dependent as well).

Could be a target of future work: find some sites that are higher MAP, C3 dominant, and also have pedogenic carbonate with the target paleoapplication being paleosols that indicate higher MAP and are older than mid Cenozoic (C3 plants only).