Vertical Accretion in Microtidal Regularly and Irregularly Flooded Estuarine Marshes

C. B. Crafta, E. D. Seneca and S. W. Broome

^aDuke Wetland Center, School of the Environment, Duke University, Durham, North Carolina 27706, U.S.A., ^bDepartment of Botany, North Carolina State University, Raleigh, North Carolina 27695-7612, U.S.A. and ^cDepartment of Soil Science, North Carolina State University, Raleigh, North Carolina 27695-7619, U.S.A.

Received 26 May 1992 and in revised form 14 December 1992

Keywords: sea-level changes; accretion; sedimentation; cesium; North Carolina coast

Vertical accretion rates were measured in microtidal (tidal amplitude less than 0·3 m) regularly (flooded twice daily by the astronomical tides), and irregularly flooded (inundated only during spring and storm tides) estuarine marshes in North Carolina to determine whether these marshes are keeping pace with rising sea-level and to quantify the relative contribution of organic matter and mineral sediment to vertical growth. Accretion rates in streamside and backmarsh locations of each marsh were determined by measuring the Cesium-137 (137Cs) activity in 2 cm soil depth increments. Soil bulk density, organic carbon (C), total nitrogen (N) and particle density also were measured to estimate rates of accumulation of organic matter (OM), mineral sediment and nutrients.

With the exception of the backmarsh location of the regularly flooded marsh, vertical accretion rates in the marshes studied matched or exceeded the recent (1940–80) rate of sea-level rise (1·9 mm year⁻¹) along the North Carolina coast. Accretion rates in the irregularly flooded marsh averaged 3.6 ± 0.5 mm year⁻¹ along the streamside and 2.4 ± 0.2 mm year⁻¹ in the backmarsh. The regularly flooded marsh had lower accretion rates, averaging 2.7 ± 0.3 mm year⁻¹ along the streamside and 0.9 ± 0.2 mm year⁻¹ in the backmarsh.

Vertical accretion in the irregularly flooded marsh occurred via in situ production and accumulation of organic matter. Rates of soil OM (196–280 g m $^{-2}$ year $^{-1}$), organic C (106–146 g m $^{-2}$ year $^{-1}$) and total N (6·9–10·3 g m $^{-2}$ year $^{-1}$) accumulation were much higher in the irregularly flooded marsh as compared to the regularly flooded marsh (OM=51–137 g m $^{-2}$ year $^{-1}$, C=21–59 g m $^{-2}$ year $^{-1}$, N=1·3–4·1 g m $^{-2}$ year $^{-1}$). In contrast, vertical accretion in the regularly flooded marsh was sustained by allochthonous inputs of mineral sediment. Inorganic sediment deposition contributed 677–1139 g m $^{-2}$ year $^{-1}$ mineral matter to the regularly flooded marsh as compared to only 147–311 g m $^{-2}$ year $^{-1}$ to the irregularly flooded marsh.

The irregularly flooded marsh exhibited an accretionary balance (vertical accretion rate minus local apparent sea-level rise) of 0.5-1.7 mm year⁻¹ suggesting that these microtidal marshes would be capable of maintaining their elevation if the rate of sea-level rise increased. In the regularly flooded marsh, the stream-side zone had an accretionary balance of 0.8 mm year⁻¹ while the backmarsh had a deficit of 1.0 mm year⁻¹. Our results suggest that microtidal regularly flooded

marshes would be susceptible to submergence by an increase in the rate of sealevel rise unless accompanied by an increase in mineral sediment deposition.

Introduction

Coastal marshes are an important component of estuarine ecosystems. Regularly (inundated twice daily by astronomical tides) and irregularly (inundated only during spring and storm tides) flooded marshes provide habitat and nursery grounds for finfish, shellfish, wildlife and waterfowl (Bellrose & Trudeau, 1988; Odum et al., 1988), support estuarine productivity by contributing detritus to the estuarine foodweb (Marinucci, 1982), serve as sites of nutrient transformations (Nixon, 1980; Jordan et al., 1983) and reservoirs of sediments (Frey & Bason, 1978; Bastian & Benforado, 1988), organic carbon (Armentano & Menges, 1986) and other nutrients (Craft et al., 1988).

Coastal marshes also slow shoreline erosion by buffering storm tides and floodwaters (Rosen, 1980; Knutson, 1988). The root-rhizome mat produced by marsh vegetation increases sediment shear strength, reducing shoreline erosion compared to unvegetated shorelines (Knutson, 1988).

There is increasing concern over the effects of anthropogenic carbon dioxide (CO₂) inputs to the atmosphere on global warming and sea-level rise. It has been estimated that the level of the earth's oceans may rise 30–250 cm by the year 2075 due to melting of polar ice and thermal expansion of ocean surface waters caused by global warming (Revelle, 1983; Titus & Barth, 1984). Since coastal marshes occupy the transition zone between terrestrial and marine ecosystems, these wetlands are extremely susceptible to erosion and submergence by rising sea-level. Maintenance of estuarine wetlands threatened by rising sea-level will depend on landward migration and/or vertical growth.

Many coastal marshes grow upward in response to rising sea-level. For the past century, sea-level has risen 10–20 cm (Gornitz et al., 1982; Barnett, 1983; Revelle, 1983). Estimates of accretion in estuarine marshes of the north-eastern (Redfield, 1972; Armentano & Woodwell, 1975; Bricker-Urso et al., 1989), mid-Atlantic (Kearney & Ward, 1986; Griffin & Rabenhorst, 1989) and south-eastern (Benninger & Chanton, 1985; Sharma et al., 1987) U.S.A. coast suggests these wetlands are keeping pace with sea-level rise. In contrast, many marshes in the Mississippi delta have not kept pace with sea-level rise (DeLaune et al., 1978, 1983; Hatton et al., 1983).

Vertical accretion occurs via sedimentation and accumulation of organic matter derived from net primary production of marsh vegetation (Redfield, 1972; DeLaune et al., 1978; Hatton et al., 1983; Bricker-Urso et al., 1989). In the Mississippi delta, sediment deposition is important, contributing approximately 80% of the mass and 50% of the volume of soil solids (DeLaune & Patrick, 1980; DeLaune et al., 1983). However, many estuarine marshes of the Atlantic coast do not receive large inputs of sediment and must rely on accumulation of organic matter to support their growth (Stevenson et al., 1986; Bricker-Urso et al., 1989).

Accretion in coastal marshes is controlled, to a large extent, by hydrology. Frequency and depth of flooding are important determinants of organic matter accumulation and sediment deposition in marshes (Gosselink & Turner, 1978; Stevenson et al., 1986). Several researchers have observed that accumulation of organic matter in marsh soils is greater in marshes that are irregularly flooded by spring and storm tides as compared to regularly flooded marshes (Gosselink & Turner, 1978; Craft et al., 1988). In contrast, sediment deposition generally is lower in marshes that are irregularly flooded (Gosselink

& Turner, 1978). Depth of flooding also may regulate the rate of marsh accretion. Stevenson *et al.* (1986) observed that accretionary balance (vertical accretion rate minus apparent sea-level rise) was positively correlated with mean tidal range (r=0.83). Stevenson *et al.* (1986) observed that coastal marshes in microtidal (tidal amplitude less than 0.5 m) environments exhibited an accretionary deficit and, thus, were susceptible to submergence under the current rate of sea-level rise.

We compared rates of vertical accretion in microtidal regularly and irregularly flooded estuarine marshes in North Carolina to: (1) determine whether these marshes are keeping pace with sea-level rise, and (2) quantify the relative contribution of organic matter and mineral sediment to the accretionary process. The regularly flooded marsh was characterized by twice daily inundation by astronomical tides with a tidal amplitude of 0·3 m. The irregularly flooded marsh was inundated only during spring and storm tides and had a tidal range of less than 10 cm. Recent (25 years) rates of accretion were estimated in streamside and backmarsh locations of each marsh by measuring ¹³⁷Cs in soil depth increments. Soil bulk density, particle density, organic carbon (C) and total nitrogen (N) also were measured to determine the role of organic and inorganic material in vertical growth and to estimate rates of accumulation of organic C and total N in marsh soils.

Methods

Site description

Soil cores were collected from a regularly and an irregularly flooded estuarine marsh along the North Carolina coast in the autumn of 1988 (Figure 1). The two marshes varied in tidal inundation, salinity and plant species composition. The regularly flooded marsh, Oregon Inlet, was inundated twice daily by the astronomical tides with a tidal amplitude of 0.3 m (Craft *et al.*, 1988). The irregularly flooded marsh, Jacob's Creek, was inundated only during spring and storm tides with a tidal range of less than 10 cm (Craft *et al.*, 1988). Regularly flooded marshes such as Oregon Inlet also are flooded more often (40-50% of the time) and to a greater depth (≥ 30 cm above the soil surface at high tide) than irregularly flooded marshes which are inundated less than 10% of the time with a flooding depth of less than 10 cm (Jernigan, 1990).

Surface water salinity at the regularly flooded marsh was 20–35 ppt as compared to salinities of 0–15 ppt at the irregularly flooded marsh (Craft et al., 1988). The regularly flooded marsh was vegetated by a monotypic stand of Spartina alterniflora Loisel, while the irregularly flooded marsh contained a mosaic of emergent macrophytes dominated by Juncus roemerianus Scheele, Distichlis spicata (L.) Greene, S. patens (Ait.) Muhl. and S. cynosuroides (L.) Roth.

Sample collection and analyses

Ten soil cores were collected from each of the two marshes using an 8.5 cm diameter by 30 cm deep stainless steel corer. The samples were stratified by location with five cores collected along the marsh edge (streamside) and five cores taken from the backmarsh, near the upland border. The soil cores were sectioned into 2 cm depth increments and the increments were air-dried, ground, sieved through a 2 mm-mesh diameter screen and weighed.

The depth increments were analysed for ¹³⁷Cs to estimate rates of vertical accretion over the past 25 years. The ¹³⁷Cs maximum in the soil profile corresponds to the soil surface in approximately 1964 (the period of maximum deposition of ¹³⁷Cs from above-ground thermonuclear testing). Cesium-137 has been used successfully to estimate recent rates of

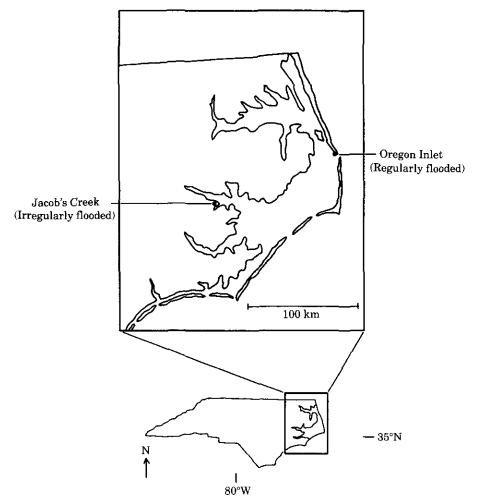


Figure 1. Location of the regularly and irregularly flooded estuarine marshes in North Carolina

accretion in estuarine wetlands (DeLaune et al., 1978; Smith et al., 1983; Sharma et al., 1987). Cesium-137 activity was measured for 8–24 h using a high purity germanium gamma detector (EG&G Ortec, Oak Ridge, Tennessee). Counting efficiency (2·08%) was determined by counting a 137 Cs standard of the same geometry. Depth increments also were analysed for bulk density, organic C and total N. Soils were analysed for organic C and total N using a Perkin-Elmer CHN analyser. Analysis (n=40) of bituminous coal (NBS standard no. 1632b; $78\cdot11\%$ C, $1\cdot56\%$ N) yielded values of $77\cdot59\pm0\cdot23\%$ C and $1\cdot55\pm0\cdot01\%$ N. Previous analysis of soils from these locations indicated no measurable carbonates in the samples (Craft et al., 1991). For this reason, it was assumed that all of the C in the samples was organic. Soil organic matter was calculated using the organic C content as described by Craft et al. (1991).

After analysis for ¹³⁷Cs, bulk density, organic C and total N, increments in the 0-10 cm depth of each core were combined for determination of particle density. Particle density was measured by the pycnometer method (Blake & Hartge, 1986). The particle density

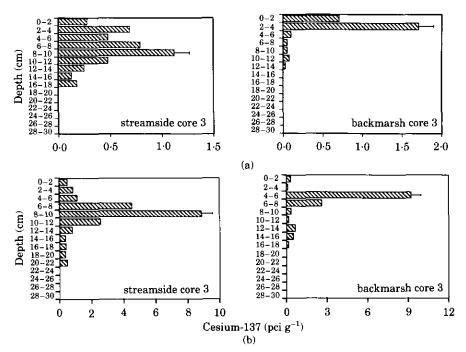


Figure 2. Depth distribution of 137 Cs in representative soil cores collected from (a) regularly and (b) irregularly flooded estuarine marshes. The 137 Cs data for all 20 cores are presented in Tables 1 and 2.

data was used to estimate the relative contribution of organic and mineral material to vertical accretion. All analyses were expressed on a dry weight basis by drying $1.0\,\mathrm{g}$ subsamples overnight at $105\,^{\circ}\mathrm{C}$.

Results and discussion

Accretion rates

There were well-defined 137 Cs maxima in the soils of both regularly and irregularly flooded marshes (Figure 2), corresponding to the period of maximum deposition of 137 Cs (approx. 1964). The depth of the 137 Cs peak occurred at a shallower depth in the regularly flooded marsh (0–10 cm) as compared to the irregularly flooded marsh (5–13 cm) (Tables 1 and 2). Using the 137 Cs maxima in the soil profile as a marker of the marsh surface in 1964, we calculated vertical accretion rates of 0·4–3·7 mm year $^{-1}$ in the regularly flooded marsh and 2·1–5·4 mm year $^{-1}$ in the irregularly flooded marsh.

Vertical accretion rates were significantly higher in the irregularly flooded marsh as compared to the regularly flooded marsh (Student's t-test, P=0·05; Figure 3). The average rate of accretion in the irregularly flooded marsh was 3·00 mm year⁻¹ while the regularly flooded marsh grew at a rate of 1·81 mm year⁻¹. The higher accretion rates in the irregularly flooded marsh were attributed to reduced tidal flushing in this low energy, brackish-water marsh. Irregularly flooded marshes in North Carolina are inundated only during spring and storm tides and during extended periods when north-easterly winds push water into the Albemarle and Pamlico Sounds. As a result, only a small amount of the net primary production (NPP) of these marshes is removed by the tides as most of the plant material remains on the marsh (Craft et al., 1988). In contrast, a larger proportion of

TABLE 1. Cesium-137 activity (pci g⁻¹) in soil cores collected from a regularly flooded estuarine marsh

Depth (cm)	Streamside core				Backmarsh core					
	1	2	3	4	5	1	2	3	4	5
0–2	0.08	0.43	0.28	0.38	0.28	1.04	0-20	0-71	0.46	0.15
2-4	0.13	0.56	0.69	0.43	0.69	0.66	0.51	1-71	0.31	0.59
4-6	0.36	0.48	0.48	0.74	0.69	0.23	$\overline{0.28}$	$\overline{0.10}$	0.13	0.41
6–8	0.66	0.79	0.79	0.69	0.56	0.13	0.08	0.05	0.08	0.13
8-10	0.56	$\overline{0.74}$	1.12	0.48	0.56	0.05	0.05	0.05	0.08	0.08
10-12	0.20	0.69	$\overline{0.48}$	0.51	0.15	0.08	0.08	0.08	0.10	0.03
12-14	0.08	0.25	0.25	0.13	0.13	0.08	0.03	0.03	0.05	0.00
14–16	0.05	0.13	0.13	0.13	0.10	0.03	a	a	a	
16–18	0.03	0.15	0.18	0.03	0.08	0.03		_	_	_
18-20	0.15	<u> </u>	<u> </u>	a	a	0.03	_	_		_
20–22	0.08	_	_	_	_	0.03	_	_	_	_
22-24	0.00	_	_	_	_	0.00	_	_	_	_
24-26	0.00	_		_	-	0.00	_	_	_	_
26–28	0.00	_	_	_	_	0.05	_	_	_	
28-30	0.00	_	_	_	_	0.03	_	_	_	_
Counting										
error ^b	0.14	0.13	0.15	0.19	0.15	0.11	0.08	0.19	0.10	0.06

The ¹³⁷Cs maximum in each core is underlined.

Table 2. Cesium-137 activity (pci g 1) in soil cores collected from an irrregularly flooded estuarine marsh

Depth (cm)	Streamside core				Backmarsh core					
	1	2	3	4	5	1	2	3	4	5
0–2	0.48	0.53	0-51	0.71	0.56	0.48	0.63	0.33	0.48	0.59
2-4	0.89	0.97	0.86	1.04	0.89	0.92	1.02	1.35	0.99	0.61
4-6	1.40	1.73	1.12	1.98	1.58	4.91	5.73	<u>9·24</u>	4.40	0.66
6-8	1.55	6.59	4.52	6.01	5.85	$\overline{4.78}$	1.96	2.65	4.61	3.66
8-10	2.16	4.78	8.88	1.25	5-06	0.28	0.33	0.36	0.56	0.87
10-12	4.73	1.17	2.54	0.38	1.04	0.36	0.18	0.15	0.20	0.28
12-14	<u>5·22</u>	0.38	0.79	0.15	0.71	0.51	0.18	0.66	0.25	0.10
14–16	3.69	0.13	0.36	0.31	0.38	0.31	0.48	0.51	0.31	0.41
16–18	2 19	0.23	0.41	a	a	a	a	0.15	a	0.25
18–20	1.55	0.23	0.38	_		_	_	a	_	0.36
20–22	1.25	0.23	0.48	_		_	_	_	_	0.18
22-24	1.22	0.08	a	_	_	_	_	_	_	0.28
24–26	0.71	a	_	_	_	_	_	_	_	0.31
26–28	0.56	_	_	_	_	_	_	_	_	0.28
28-30	0.00	_	_	_	_	_	_	_	_	0.25
Counting										
error ^b	0.47	0.40	0.71	0.36	0.41	0.00	0.46	0.74	0.46	0.37

The ¹³⁷Cs maximum in each core is underlined.

^{a 137}Cs was not measured below this depth, ^b counting error of the ¹³⁷Cs maximum.

^{a 137}Cs was not measured below this depth, ^b counting error of the ¹³⁷Cs maximum.

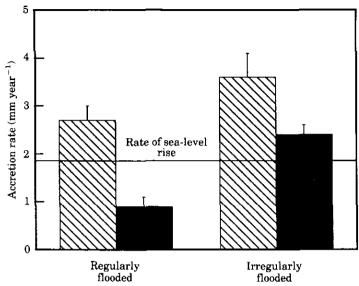


Figure 3. Mean vertical accretion rates in regularly and irregularly flooded marsh soils. Each value represents the mean $(\pm 1 \text{ SE})$ of five cores. \blacksquare , backmarsh; \square , streamside.

the NPP in regularly flooded marshes is removed by the astronomical tides so that less organic matter is available to contribute to vertical accretion. Craft *et al.* (1988) observed that 20% of the NPP accumulated in the soil of an irregularly flooded estuarine marsh annually as compared to 11-12% in regularly flooded marshes.

At both locations, accretion rates were higher in the streamside marsh than in the backmarsh (Figure 3). Accretion rates in the irregularly flooded marsh averaged 3·6 mm year⁻¹ along the streamside and 2·4 mm year⁻¹ in the backmarsh while the mean accretion rate in the streamside and backmarsh of the regularly flooded marsh was 2·7 and 0·9 mm year⁻¹, respectively. Several researchers (DeLaune et al., 1978; Richard, 1978; DeLaune & Patrick, 1980; Hatton et al., 1983) also observed higher rates of accretion in streamside marshes and attributed it to increased inputs of sediment from the water column. Sediment deposition may also enhance plant growth by adding particulate nutrients to the marsh (DeLaune et al., 1981; Bricker-Urso et al., 1989). DeLaune and Patrick (1980) estimated that sedimentation processes contributed 23·1, 2·31 and 99·1 g m⁻² year⁻¹ of N, phosphorus (P) and potassium (K) respectively to a Louisiana saltmarsh. The addition of particulate nutrients to the marsh may act as a positive-feedback to vertical accretion by stimulating plant growth and organic matter accumulation (Bricker-Urso et al., 1989).

With the exception of the backmarsh location at the regularly flooded marsh, accretion rates equalled or exceeded the rate of sea-level rise (Figure 3). The rate of sea-level rise along the North Carolina coast (1940–80) has been estimated at 1.9 mm year⁻¹ (Stevenson *et al.*, 1986) which indicates that the vertical growth of these marshes generally, is keeping pace with the current rate of sea-level rise.

Stevenson *et al.* (1986) postulated that the amount of tidal energy (expressed as mean tidal range) is a major factor controlling marsh accretion under conditions of both low and high sediment inputs. These researchers found a positive correlation (r=0.83) between accretionary balance (vertical accretion rate minus local apparent sea-level rise) and mean tidal range. According to Stevenson *et al.* (1986), marshes of microtidal environments

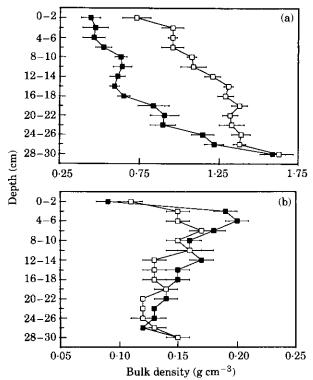


Figure 4. Mean bulk density by depth in (a) regularly and (b) irregularly flooded estuarine marsh soils. Each data point represents the mean $(\pm 1 \text{ SE})$ of five cores. $-\Box$ —, backmarsh; $-\blacksquare$ —, streamside.

(mean tidal amplitude of 0.5 m or less) have a negative accretionary balance and are not keeping pace with sea-level rise. However, both our regularly and irregularly flooded marshes are keeping pace with sea-level rise in spite of low tidal amplitude. The irregularly flooded marsh, Jacob's Creek (tidal range less than 10 cm), has an accretionary balance of 0.5–1.7 mm year⁻¹. The regularly flooded marsh, Oregon Inlet (tidal range of 0.3 m), exhibited a positive accretionary balance in the streamside zone (0.8 mm year⁻¹) but the backmarsh zone had a deficit of 1.0 mm year⁻¹. Our results suggest that regularly and irregularly flooded microtidal marshes in North Carolina are able to maintain elevation in response to the current rate of sea-level rise. However, the mechanisms controlling vertical accretion differ between the two marsh types with *in situ* organic matter accumulation being the primary mechanism in irregularly flooded marshes and mineral sediment deposition as the dominant mechanism in the regularly flooded marsh (see section on Accumulation rates of organic matter and mineral sediment).

Soil bulk density, organic carbon and total nitrogen

Soil bulk density varied inversely with the organic matter content of the soil. Bulk density was much higher in the low organic matter regularly flooded marsh $(0.44-1.63 \text{ g cm}^{-3})$ as compared to the organic soils of the irregularly flooded marsh $(0.09-0.20 \text{ g cm}^{-3})$ [Figure 4(a,b)]. In the regularly flooded marsh soil, bulk density was lowest at the surface and increased with depth [Figure 4(a)]. Soil bulk density was lower at the streamside location than the backmarsh [Figure 4(a)] as a result of the higher organic C content [see Figure

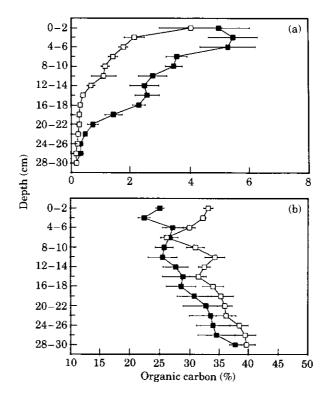


Figure 5. Mean organic carbon content by depth in (a) regularly and (b) irregularly flooded estuarine marsh soils. Each data point represents the mean $(\pm 1 \text{ SE})$ of five cores. $-\Box$, backmarsh; $-\blacksquare$, streamside.

5(a)]. In contrast to the regularly flooded marsh, bulk density in the irregularly flooded marsh soils exhibited no clear differences with regard to location or depth [Figure 4(b)].

The regularly flooded marsh soils contained much lower concentrations of organic C (<6%) than the soils of the irregularly flooded marsh (22–40%) [Figure 5(a,b)]. In the regularly flooded marsh, soil organic C decreased with depth in both the streamside and backmarsh locations [Figure 5(a)] while organic C increased with depth in the irregularly flooded marsh [Figure 5(b)].

The difference in concentration and distribution of soil organic C in the regularly and irregularly flooded marshes reflects the age difference of the two marshes. The regularly flooded marsh is located on the sound side of the North Carolina barrier islands. During the past 5000 years, as sea-level has risen, the barrier islands have migrated landward with the rising sea (Pilkey et al., 1982). As a result, these barrier island marshes are an ephemeral part of the landscape as the barrier islands 'roll over' in their landward migration (Leatherman, 1982). In contrast, the irregularly flooded marsh is located on the Pamlico terrace; a much older (approx. 50 000 years) landform (Daniels et al., 1984). The thickness of the organic layer (2–3 m) and uniform depth distribution of organic C suggests that this marsh has existed for several thousand years. In fact, ¹⁴C dating from the basal peat of a nearby marsh yielded a date of 2600 years BP (before present) (Bellis & Gaither, 1985).

Soil organic C was much lower at the backmarsh location of the regularly flooded marsh as compared to the streamside marsh [Figure 5(a)]. Again, the difference in soil organic C

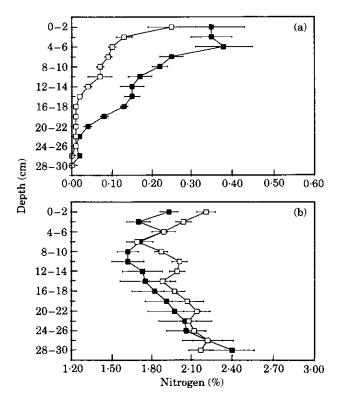


Figure 6. Mean total nitrogen content by depth in (a) regularly and (b) irregularly flooded estuarine marsh soils. Each data point represents the mean (± 1 SE) of five cores. $-\Box$ —, backmarsh; $-\blacksquare$ —, streamside.

between the streamside and backmarsh is due to the difference in age. As the sea-level rose, the streamside marsh was initially inundated and colonized by emergent vegetation. During this time the backmarsh area probably was dominated by maritime shrub vegetation such as red cedar (funiperus), wax myrtle (Myrica) and marsh elder (Iva) (Leatherman, 1982). Only recently, as the sea-level continued to rise, has this area been flooded and colonized by S. alterniflora.

In contrast to the regularly flooded marsh, soil organic C was lower along the streamside of the irregularly flooded marsh [Figure 5(b)]. The lower organic C content of the streamside marsh soil probably reflects higher sedimentation rates adjacent to the stream than in the interior of the marsh.

The distribution of total N in the marshes was similar to that of organic C. Soil N was much higher in the organic soils of the irregularly flooded marsh (1.62-2.40%) than in the sandy soils of the regularly flooded marsh [0-0.38%]; Figure 6(a,b)]. Total N decreased with depth in soils of the regularly flooded marsh [Figure 6(a)] and increased with depth in the irregularly flooded marsh [Figure 6(b)]. Like organic C, soil N was higher in the streamside of the regularly flooded marsh and in the backmarsh of the irregularly flooded marsh [Figure 6(a,b)].

Accumulation rates of organic carbon and total nitrogen

Rates of organic C and total N accumulation were much higher in the irregularly flooded marsh as compared to the regularly flooded marsh (Table 3) and were attributed mostly to

Marsh type	Location	Organic carbon (g m -2 year -1)	Total nitrogen (g m ² year ¹)
Regularly flooded	Streamside	58·9±7·3	4·11 ± 0·56
(Oregon Inlet)	Backmarsh	21.3 ± 4.7	1.32 ± 0.28
Irregularly flooded	Streamside	145·5 ± 19·4	10.28 ± 1.40
(Jacob's Creek)	Backmarsh	106.5 ± 6.8	6.89 ± 0.57

Table 3. Mean accumulation rates of organic carbon and total nitrogen in regularly and irregularly flooded estuarine marsh soils

Each value represents the mean (± 1 SE) of five cores.

much higher concentrations of organic C and total N (Figures 5 and 6) and, to a lesser extent, to higher accretion rates (Figure 3). Rates of organic C and total N accumulation in our marshes were much lower than reported for coastal marshes in Louisiana (DeLaune et al., 1981; Hatton et al., 1982; Smith et al., 1983). The difference was attributed to much higher accretion rates (5·2–14·0 mm year⁻¹) in the Louisana marshes (DeLaune et al., 1981; Hatton et al., 1982; Smith et al., 1983).

The higher rates of C and N accumulation in the irregularly flooded marsh may be attributed to: (1) reduced rates of decomposition caused by waterlogging, (2) higher net primary production (and organic matter deposition), (3) emergent vegetation (e.g. Juncus) which is inherently more resistant to decomposition or (4) reduced tidal flushing and export of detritus. The weak and irregular inundation in marshes such as Jacob's Creek may result in waterlogging of the sediments, slowing decomposition and enhancing organic matter accumulation as compared to coastal wetlands that are vigorously flooded by astronomical tides (Stevenson et al., 1986). However, Jernigan (1990) surveyed 12 marshes in North Carolina and found that the sediments underlying irregularly flooded Juncus marshes were less reduced (55 mV) than sediments underlying S. alterniflora (-112 mV). Differences in net primary productivity or decomposition rates of emergent vegetation also cannot account for the higher rates of C and N accumulation in the irregularly flooded marsh. Net primary productivity of emergent vegetation was similar in the regularly (1046 g C m⁻² year⁻¹) and irregularly flooded (900 g C m⁻² year⁻¹) marshes (S. W. Broome, unpubl. data). Decomposition rates of Juncus and Spartina also appear to be similar. Hackney and Cruz (1980) observed that, after 1 year, litter bags containing Juncus and Spartina roots lost 17 and 20% of their mass, respectively. Likewise, McKee and Seneca (1982) found that litter bags containing above-ground Spartina and Juncus tissue exhibited similar decomposition rates (approximately 80% weight loss after 13 months).

It is likely that rates of organic matter and nutrient accumulation are often higher in irregularly flooded marshes because less of the NPP is removed from the marsh surface by tidal action (Hackney & Cruz, 1982; Craft et al., 1988). The percentage of NPP buried annually was nearly four times higher in the irregularly flooded marsh (14%, 126 g C m⁻² year⁻¹) as compared to the regularly flooded marsh (3·4%, 40 g C m⁻² year⁻¹). Craft et al. (1988) also observed that only 11–12% of the NPP in a regularly flooded transplanted marsh accumulated in the sediments annually as compared to 20% in an irregularly flooded transplanted marsh.

Marsh type	Location	Organic matter (g m ⁻² year ⁻¹)	Inorganic sediment (g m ⁻² year ⁻¹)
Regularly flooded	Streamside	137±17(11%)	1139+210
(Oregon Inlet)	Backmarsh	$51\pm11(7\%)$	677 ± 197
Irregularly flooded	Streamside	$280 \pm 39 (47\%)$	311 ± 63
(Jacob's Creek)	Backmarsh	$196 \pm 13 (57\%)$	147 ± 14

Table 4. Mean accumulation rates of organic matter and mineral sediment in regularly and irregularly flooded estuarine marshes

Each value represents the mean $(\pm 1 \text{ SE})$ of five cores. Numbers in parentheses represent the mass contribution of organic matter to the vertical accretion rate

Accumulation rates of organic matter and mineral sediment

Like organic C, rates of organic matter accumulation were much higher in the irregularly flooded marsh (196–280 g m⁻² year⁻¹) than in the regularly flooded marsh (51–137 g m⁻² year⁻¹). At both sites, accumulation of organic matter was greater in the streamside marsh as compared to the backmarsh (Table 4). In contrast to organic matter, deposition of inorganic sediment was much higher in the regularly flooded marsh (677–1139 g m⁻² year⁻¹) as compared to the irregularly flooded marsh (147–311 g m⁻² year⁻¹). On a mass basis, organic matter accounted for 47–57% of the accretion of solids in the irregularly flooded and only 7–11% in the regularly flooded marsh (Table 4).

Accumulation rates of organic matter and mineral sediment were two to three times greater along the streamside as compared to the backmarsh of the regularly flooded marsh (Table 4). As a result, the streamside location is able to maintain its elevation (2·7 mm year⁻¹) against rising sea-level (1·9 mm year⁻¹). In contrast, accretion rates in the backmarsh location (0·9 mm year⁻¹) indicate that there is insufficient organic matter and mineral sediment accumulation to sustain vertical growth in the face of rising sea-level. It is likely that the accretionary deficit in the backmarsh location at Oregon Inlet is due to reduced mineral sediment inputs (Table 4) caused by a reduction in tidal inundation (as compared to the streamside location). There also appears to be little if any contribution by overwash at this site.

Soil bulk density and particle density (upper 10 cm) was much lower in the organic soils (>50% organic matter) of the irregularly flooded marsh than in the sandy soils (7–11% organic matter) of the regularly flooded marsh (Table 5). Particle density of the regularly flooded marsh soil was somewhat less $(2\cdot40-2\cdot50\,\mathrm{g\,cm^{-3}})$ than the particle density reported for mineral soils $(2\cdot60-2\cdot75\,\mathrm{g\,cm^{-3}};\mathrm{Brady},1984)$. Assuming a particle density of $2\cdot65\,\mathrm{g\,cm^{-3}}$ for inorganic sediment, we calculated that the particle density of organic matter in the irregularly flooded marsh was $1\cdot24-1\cdot34\,\mathrm{g\,cm^{-3}}$. Our values are within the range $(1\cdot10-1\cdot40\,\mathrm{g\,cm^{-3}})$ given by Brady (1984) for organic matter.

At both marshes, most of the volume in the upper 10 cm of the soil was occupied by pores (Table 5). Voids accounted for 65–80% of the volume in the regularly flooded marsh soil and 91–93% in the irregularly flooded marsh soil. Organic matter accounted for 4–5% of the total volume in the regularly flooded marsh compared to 16–30% for inorganic sediment (Table 5). In the irregularly flooded marsh, organic matter and inorganic sediment contributed 5–6 and 2–3% of the volume, respectively (Table 5). If the data are expressed as the percentage contribution to the solid phase, then soil organic matter

and	 properties (bulk density, pa pper 10 cm and percentage volume accretion	

	Regularly flooded (Oregon Inlet)		Irregularly flooded (Jacob's Creek)		
	Streamside	Backmarsh	Streamside	Backmarsh	
Bulk density				11	
(g dry weight cm ⁻³)	0.48 ± 0.01	0.85 ± 0.09	0.16 ± 0.02	0.13 ± 0.01	
Particle density					
(g dry weight cm ⁻³)	2.40 ± 0.07	2.50 ± 0.07	1.95 ± 0.06	1.87 ± 0.07	
Organic matter					
(% dry weight)	11 ± 1	7 ± 2	53 ± 2	55 <u>+</u> 2	
Particle density of					
organic matter (g cm -3)a	_	_	1.34 ± 0.08	1.24 ± 0.11	
Porosity					
(%) ^b	80 ± 1	65 <u>±</u> 3	91 <u>+</u> 1	93 ± 1	
% contribution of					
organic matter	4(20)	5 (14)	6 (67)	5 (71)	
% contribution of					
inorganic sediment	16 (80)	30 (86)	3 (33)	2 (29)	

Each number represents the mean (± 1 SE) of the averaged depth integrated (0–10 cm) values of the five cores.

values of the five cores.

$$^{a}D_{p}(OM) = \frac{D_{p}(soil) - (\%IS/100) (D_{p}(IS))}{(\%OM/100)}$$

accounts for 14-20 and 67-71% of the volume of solids in the regularly (7-11% organic matter by weight) and irregularly (53-55% organic matter by weight) flooded marshes, respectively.

Other studies also have documented the importance of soil organic matter in vertical accretion. DeLaune et al. (1983) observed that coastal marshes in Louisiana containing 20–30% soil organic matter by weight contributed approximately 50% of the volume of soil solids. In Rhode Island saltmarshes (30–60% organic matter), organic matter accounted for 50 and 73% of the volume of soil solids in the low and high marsh, respectively (Bricker-Urso et al., 1989). Our results and those of other researchers demonstrate the importance of organic matter to vertical accretion in both organic and mineral marshes by enhancing soil volume. Other studies also have documented the importance of soil organic matter to vertical accretion by providing structural support and pore space (DeLaune & Patrick, 1980; DeLaune et al., 1983; Hatton et al., 1983; Bricker-Urso et al., 1989).

Although both regularly and irregularly flooded microtidal marshes are able to grow vertically in response to the current rate of sea-level rise, the mechanisms controlling accretion differ between the two marsh types. Irregularly flooded marshes grow vertically through accumulation of organic matter from emergent vegetation (Tables 4 and 5). The weak and irregular inundation of marshes such as Jacob's Creek results in limited export

 $^{^{}b}$ Porosity = 1 - (D_{b}/D_{p}) .

Volume (OM or IS) = Mass (OM or IS)/ D_p (OM or IS). Calculated using the mean D_p for each location and the D_p (OM) and D_p (IS) of 1·29 and 2·65 g cm⁻³, respectively. Numbers in parentheses represent the percentage contribution of organic matter or mineral sediment to accretion of solids.

D_b, bulk density; D_n, particle density; IS, inorganic sediment; OM, organic matter.

of detritus so that most of the NPP accumulates on the marsh surface. Likewise, the low tidal energy associated with irregularly flooded marshes limits allochthonous inputs of mineral sediment, reinforcing the importance of organic matter to vertical growth in these marshes. In contrast, much of the NPP in regularly flooded marshes such as Oregon Inlet is exported by the astronomical tides. However, the increased tidal energy associated with regularly flooded marshes contributes substantial amounts of inorganic sediment to the surface of these marshes. As a result, the vertical growth of these microtidal regularly flooded marshes is sustained primarily by allochthonous inputs of mineral material (Tables 4 and 5).

Conclusions

Vertical accretion rates in an irregularly flooded estuarine marsh (2·4–3·6 mm year⁻¹) indicate that this coastal wetland is keeping pace with sea-level rise (1·9 mm year⁻¹). Accretion rates in a regularly flooded marsh suggest that the streamside marsh (2·7 mm year⁻¹) is maintaining its elevation relative to sea-level rise but the backmarsh (0·9 mm year⁻¹) is not keeping pace with rising sea-level. At both marshes, vertical accretion was greater in streamside than backmarsh locations as a result of greater deposition of organic matter and inorganic sediment.

The mechanisms controlling vertical accretion are different in the two marshes. In the irregularly flooded marsh, accretion occurs via *in situ* production and accumulation of organic matter. However, allochthonous inputs of mineral sediment are responsible for the vertical growth of the regularly flooded marsh. The irregularly flooded marsh exhibited a positive accretionary balance (vertical accretion rate minus local apparent sealevel rise) of 0·5–1·7 mm year⁻¹, suggesting that this marsh is capable of maintaining itself in response to an increase in the rate of sea-level rise. In the regularly flooded marsh, the streamside zone had a positive accretionary balance (0·8 mm year⁻¹) but the backmarsh exhibited an accretionary deficit of 1·0 mm year⁻¹. Our results suggest that regularly flooded microtidal (tidal amplitude less than 0·5 m) marshes would be susceptible to submergence if exposed to increased rates of sea-level rise. Vertical growth of these marshes in response to rising sea-level will be dependent on a concomitant increase in allochthonous mineral sediment inputs.

Acknowledgements

We wish to thank C. Campbell and L. Hobbs for assisting with sample collection, B. Crabtree and Dr Wayne Robarge for advice with the ¹³⁷Cs analysis and B. Taylor for analysing the samples for C and N.

References

Armentano, T. V. & Menges, E. S. 1986 Patterns of change in the carbon balance of organic soil-wetlands of the temperate zone. Journal of Ecology 74, 755-774.

Armentano, T. V. & Woodwell, G. M. 1975 Sedimentation rates in a Long Island marsh determined by Pb-210 dating. *Limnology and Oceanography* 20, 452-456.

Barnett, T. P. 1983 Global sea level: estimating and explaining apparent changes. In *Coastal Zone '83* (Magoon, O. T., ed.). American Society of Civil Engineers, New York, pp. 2777–2795.

- Bastian, R. K. & Benforado, J. 1988 Water quality functions of wetlands: natural and managed systems. In *The Ecology and Management of Wetlands. Vol. 1: Ecology of Wetlands* (Hook, D. D. et al., eds). Croom Helm, London, pp. 87-97.
- Bellis, V. J. & Gaither, A. C. 1985 Seasonality of aboveground and belowground biomass for six salt marsh plant species. *Journal of the Elisha Mitchell Society* 10, 95–109.
- Bellrose, F. C. & Trudeau, N. M. 1988 Wetlands and their relationship to migrating and winter populations of waterfowl. In *The Ecology and Management of Wetlands*. Vol. 1: Ecology of Wetlands (Hook, D. D. et al., eds). Croom Helm, London, pp. 183–194.
- Benninger, L. K. & Chanton, J. P. 1985 Fallout ^{239,240}Pu and natural ²³⁸U and ²¹⁰Pb in sediments of the North River marsh, North Carolina. *Eos* **66**, 276.
- Blake, G. R. & Hartge, K. H. 1986 Particle density. Agronomy 9, 377-382.
- Brady, N. C. 1984 The Nature and Properties of Soils. Macmillan, New York.
- Bricker-Urso, S., Nixon, S. W., Cochran, J. K., Hirschberg, D. J. & Hunt, C. 1989 Accretion rates and sediment accumulation in Rhode Island salt marshes. *Estuaries* 12, 300-317.
- Craft, C. B., Broome, S. W. & Seneca, E. D. 1988 Nitrogen, phosphorus and organic carbon pools in natural and transplanted marsh soils. *Estuaries* 11, 272–280.
- Craft, C. B., Seneca, E. D. & Broome, S. W. 1991 Loss on ignition and Kjeldahl digestion for estimating organic carbon and total nitrogen in estuarine marsh soils: calibration with dry combustion. Estuaries 14, 175-179.
- Daniels, R. B., Kleiss, H. J., Buol, S. W., Byrd, H. J. & Phillips, J. A. 1984 Soil systems in North Carolina. Bulletin No. 467. North Carolina Agricultural Research Service, Raleigh, North Carolina.
- DeLaune, R. D., Baumann, R. H. & Gosselink, J. G. 1983 Relationships among vertical accretion, coastal submergence and erosion in a Louisiana gulf coast marsh. *Journal of Sedimentary Petrology* 53, 147-157.
- DeLaune, R. D., Patrick, W. H. Jr & Buresh, R. J. 1978 Sedimentation rates determined by ¹³⁷Cs dating in a rapidly accreting salt marsh. *Nature* 275, 532-533.
- DeLaune, R. D. & Patrick, W. H. Jr 1980 Rate of sedimentation and its role in nutrient cycling in a Louisiana salt marsh. In *Estuarine and Wetland Processes with Emphasis on Modeling* (Hamilton, P. & Macdonald, K. B., eds). Plenum, New York, pp. 401-412.
- DeLaune, R. D., Reddy, C. N. & Patrick, W. H. Jr 1981 Accumulation of plant nutrients and heavy metals through sedimentation processes and accretion in a Louisiana salt marsh. *Estuaries* 4, 328–334.
- Frey, R. W. & Basan, P. B. 1978 Coastal salt marshes. In *Coastal Sedimentary Environments* (Davis, R. A., ed.). Springer-Verlag, New York, pp. 101-168.
- Gornitz, V., Lebedeff, S. & Hanson, J. 1982 Global sea level trend in the past century. Science 215, 1611-1614.
- Gosselink, J. G. & Turner, R. E. 1978 The role of hydrology in freshwater wetland ecosystems. In Freshwater Wetlands: Ecological Processes and Management Potential (Good, R. E., Whigham, D. F. & Simpson, R. L., eds). Academic Press, New York, pp. 63-77.
- Griffin, T. M. & Rabenhorst, M. C. 1989 Processes and rates of pedogenesis in some Maryland tidal marsh soils. Soil Science Society of America Journal 53, 862-870.
- Hackney, C. T. & de la Cruz, A. A. 1980 In situ decomposition of roots and rhizomes of two tidal marsh plants. *Ecology* **61**, 225-231.
- Hackney, C. T. & de la Cruz, A. A. 1982 The structure and function of brackish marshes in the north central Gulf of Mexico: a ten year case study. In Wetlands Ecology and Management (Gopal, B., Turner, R. E., Wetzel, R. G. & Whigham, D. F., eds). National Institute of Ecology and International Scientific Publications, Jaipur, pp. 89-107.
- Hatton, R. S., DeLaune, R. D. & Patrick, W. H. Jr 1983 Sedimentation, accretion and subsidence in marshes of Barataria Basin, Louisiana. *Limnology and Oceanography* 28, 494–502.
- Hatton, R. S., Patrick, W. H. Jr & DeLaune, R. D. 1982 Sedimentation, nutrient accumulation and early diagenesis in Louisiana Barataria Basin coastal marshes. In *Estuarine Comparisons* (Kennedy, V. S., ed.). Academic Press, New York, pp. 255–267.
- Jernigan, L. S. Jr 1990 Factors influencing the distributional patterns of *Juncus roemerianus* in two North Carolina salt marshes. Ph.D. Thesis, North Carolina State University, Raleigh, North Carolina.
- Jordan, T. E., Correll, D. L. & Whigham, D. F. 1983 Nutrient flux in the Rhode river: tidal exchange of nutrients by brackish marshes. Estuarine, Coastal and Shelf Science 17, 651-667.
- Kearney, M. S. & Ward, L. G. 1986 Accretion rates in brackish marshes of a Chesapeake Bay estuarine tributary. *Geo-Marine Letters* **6**, **41**–49.
- Knutson, P. L. 1988 Role of coastal marshes in energy dissipation and shore protection. In The Ecology and Management of Wetlands. Vol. 1: Ecology of Wetlands (Hook, D. D. et al., eds). Croom Helm, London, pp. 161-175.
- Leatherman, S. P. 1982 Barrier Island Handbook. University of Maryland, College Park, Maryland.
- Marinucci, A. C. 1982 Trophic importance of Spartina alterniflora production and decomposition to the marsh-estuarine ecosystem. Biological Conservation 22, 35-58.
- McKee, K. L. & Seneca, E. D. 1982 The influence of morphology in determining the decomposition of two salt marsh macrophytes. *Estuaries* 5, 302–309.

- Nixon, S. W. 1980 Between coastal marshes and coastal waters: a review of twenty years of speculation and research on the role of salt marshes in estuarine productivity and water chemistry. In Estuarine and Wetland Processes with Emphasis on Modeling (Hamilton, P. & Macdonald, K. B., eds). Plenum, New York, pp. 437-525.
- Odum, W. E., Rosas, L. P. & McIvor, C. C. 1988 A comparison of fish and invertebrate community composition in tidal freshwater and oligohaline marsh systems. In *The Ecology and Management of Wetlands*. *Vol. 1: Ecology of Wetlands* (Hook, D. D. *et al.*, eds). Croom Helm, London, pp. 561–569.
- Pilkey, O. H. Jr, Neal, W. J., Pilkey, O. H. Sr & Riggs, S. R. 1982 From Currituck to Calabash: Living with North Carolina's Barrier Islands. Duke University Press, Durham.
- Redfield, A. C. 1972 Development of a New England salt marsh. Ecological Monographs 42, 201-237.
- Revelle, R. R. 1983 Probable future changes in sea level resulting from increased atmospheric carbon dioxide. In Changing Climate: Report of the Carbon Dioxide Assessment Committee. National Academy Press, Washington, pp. 433-448.
- Richard, G. A. 1978 Seasonal and environmental variations in sediment accretion in a Long Island salt marsh. *Estuaries* 1, 29–35.
- Rosen, P. S. 1980 Erosion susceptibility of the Virginia Chesapeake Bay shoreline. Marine Geology 34, 45-59.
- Sharma, P., Gardner, L. R., Moore, W. S. & Bollinger, M. S. 1987 Sedimentation and bioturbation in a salt marsh as revealed by ²¹⁰Pb, ¹³⁷Cs and ⁷Be studies. *Limnology and Oceanography* **32**, 313–326.
- Smith, C. J., DeLaune, R. D. & Patrick, W. H. Jr 1983 Carbon dioxide emission and carbon accumulation in coastal wetlands. Estuarine, Coastal and Shelf Science 17, 21-29.
- Stevenson, J. C., Ward, L. G. & Kearney, M. S. 1986 Vertical accretion in marshes with varying rates of sea level rise. In *Estuarine Variability* (Wolfe, D. A., ed.). Academic Press, New York, pp. 241–259.
- Titus, J. G. & Barth, M. C. 1984 An overview of the causes and effects of sea level rise. In *Greenhouse Effect* and Sea Level Rise (Barth, M. C. & Titus, J. G., eds). Van Nostrand Reinhold, New York, pp. 1-56.