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clear; clc; close all
% Solution Output:
% Spencer Freeman, 4/12/2024
응
% 2a: 4247.345825, m/s
% 2b: 11669.486746, m/s
% 2c: 103.039085, m
% 2d: 0.055479, s
% 2e: 0.104177, m/s
% 2f: 5322.629358, nT
% 3a: 201261.842171, ions/something...
% 3b: 42.185130, 42.204214, 42.185249, ions/something... / s
% 4: -3.781406e-12, Pa/km
%% 2) A He+ ion is in near Earth space, at the geomagnetic equator. It has
% = 1.60 \times 10^{-2} and a charge of +1.60 \times 10^{-1}. It has a kinetic
% energy of 3.2 eV (where 1eV = 1.6 \times 10 - 19 J). The particle's motion is
% directed at an angle of 70 degrees away from the local geomagnetic field
% vector. The local geomagnetic field has a magnitude of 4700 nT, and a
% radius of curvature of 7300 km. Compute the following:
fprintf('Solution Output:\n')
fprintf('Spencer Freeman, 4/12/2024\n\n')
r2cyc = 1/(2*pi); % cycles/radian
eV2J = 1.6e-19; % J/eV
He mass = 6.64e-27; % kg
He c = 1.60e-19; % C
He u = 3.2 * eV2J; % J
dpsi = 70 * pi/180; % The particles motion directed away from the local geomagnetic ✓
field vector, rad
B mag = 4700e-9; % local geomagnetic field, T
Rc mag = 7300e3; % radius of curvature, m
% a) The velocity of the particle parallel to the magnetic field, in m/s
% (hint, use its kinetic energy to find its total speed first)
v_mag = sqrt(2*He_u/He_mass); % m/s
v B para = v mag * cos(dpsi); % m/s
fprintf('2a: f, m/s\n', v B para)
% b) The velocity of the particle perpendicular to the magnetic field, in m/s.
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v B perp = v mag * sin(dpsi); % m/s
fprintf('2b: %f, m/s\n', v B perp)
% c) The gyroradius of the particle, in m.
r_l = He_mass * v_B_perp / He_c / B_mag; % m
fprintf('2c: %f, m\n', r 1)
% d) The gyroperiod of the particle, in s. (hint, don't forget 2-pi)
w c = v B perp / r l * r2cyc; % Hz
T = 1 / w c; % s
fprintf('2d: %f, s\n', T)
% e) The magnitude of the gradient plus curvature drift of the particle, in m/s.
% V \nabla B + V (\nabla \times B) = m/q (R \nabla \times B) \rightarrow (R \nabla \times B) \rightarrow (V \| ^2 + 1/2 V \| ^2)
B = B mag * [0; 1; 0]; % B field vector (at geomagnetic equator), T
Rc = Rc mag * [1; 0; 0]; % m
v_{grad} = norm(He_{mass} / He_{c} * cross(Rc, B) / (Rc mag^2 * B mag^2) * (v B para^2 + \checkmark)
.5*v B perp^2)); % m/s
fprintf('2e: %f, m/s\n', v grad curv)
% f) The magnetic field strength, in nT, at which this particle will mirror.
B_m = B_mag * (v_mag / v_B_perp)^2; % T
fprintf('2f: %f, nT\n', B m * 10^9)
%% 3) In the E-region ionosphere, intense auroral electrons are creating a
% Chapman production layer. The altitude of peak production is at 125 km,
% with a production rate of 3.2*10^8 ions m-3 s-1. The recombination coefficient
% (alpha) in this region is 0.0079.
fprintf('\n')
clear
q = 3.2*10^8; % ions / m^3 / s
alpha = 0.0079;
\$ a) Estimate the ion density in this region at the time of the intense auroral oldsymbolarksim
electrons.
% n e (z,t) \approx \sqrt{(q(z,t)/\alpha)}
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n e0 = sqrt(q / alpha); % ions /
fprintf('3a: %f, ions/something...\n', n e0)
% b) A few moments later, the auroral electrons move away from this region.
% If it is currently nighttime (i.e. sun has set), determine the rate of
% change of ion density 3 seconds after the auroral electrons disappear.
% (hint, if you solve this numerically, use a small time-step to ensure
% dt * dn/dt is << n).
% (\partial n \ e) / \partial t = q(z, t) - \alpha n \ e^2
n e = n e0;
tend = 3; % s
dt = 1e-6; % s
ts = linspace(0, tend, round(tend/dt)); % s
for i = 1:length(ts)
    n = dot = -alpha * n e^2;
    n = n = + n = dot * dt;
   n es(i) = n e;
[t,nes] = ode45(@(t, ne) -alpha * ne^2, [0 tend], n e0);
n_e_ode = nes(end);
n \in tend = 1 / (alpha * tend + 1 / n e0);
% figure
% plot(ts, n es)
% grid on
fprintf('3b: %f, %f, ions/something... / s\n', n e, n e ode, n e tend)
%% 4) You have obtained observations of the O+ density in the upper F-region
% near the magnetic pole. At this altitude, the ion temperature is
% approximately constant with height, and the electron temperature is 1.5x
% the ion temperature.
% Find the vertical gradient in the pressure associated with the O+ plasma
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% in this region. You may ignore any other sources of pressure, and you may
% assume that O+ acts as an ideal gas when computing pressure (see Part 1,
% Lesson 1 slides if needed).
fprintf('\n')
clear
n_0 = 17312751370; % O+ density, m^-3
T O = 1606; % Ion temperature, K
T e = T O * 1.5; % electron temperature, K
h = 584; % Altitude of observation, km
Re = 6378; % Radius of Earth, km
mass O = 2.66e-26; % Mass of O+, kg
G = 6.67e-11; % Gravitational constant G
mass earth = 5.97e24; % Mass of Earth, kg
kb = 1.38e-23; % Boltzmann's constant
% use iterative method to calculate density
km2m = 1000; % m/km
m2km = 1/km2m; % km/m
delz = .1*km2m; % m
z2 = h: (delz*m2km): (h + 100); % km
rho0 = n O * mass O; % kg/m^3
rho = rho0;
for i = 2:length(z2)
    zi = z2(i);
    r = Re + zi; % km
    g = G*mass earth/(r*km2m)^2; % m/s^2
    H = kb*T O/mass O/g;
    rho(i) = rho(i - 1) *exp(-delz/H); % kg/m^3
N = rho/mass O; % number of ions / m^3
% ideal gas law
p = N*kb*T O; % Pa
% forward difference
dpdz = diff(p(1:2)) / diff(z2(1:2)); % Pa/km
figure
plot(p, z2)
grid on
xlabel('Pressure (Pa)')
ylabel('alt (km)')
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fprintf('4: %e, Pa/km\n', dpdz)