

## Rapid emplacement of massive Duluth Complex intrusions within the Midcontinent Rift

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**Abstract:**

The Duluth Complex is one of the largest mafic intrusive complexes in the world. It was emplaced as the Midcontinent Rift developed in Laurentia's interior during an interval of magmatism and extension from ca. 1109 to 1084 Ma. This duration of magmatic activity is more protracted than is typical for large igneous provinces interpreted to have formed from decompression melting of upwelling mantle plumes. While the overall duration of magmatic activity was protracted, there were intervals of more voluminous magmatism. New high-precision  $^{206}\text{Pb}/^{238}\text{U}$  dates for the anorthositic and layered series of the Duluth Complex constrain these units to have been emplaced ca. 1096 Ma in less than 1 million years. Comparison of paleomagnetic data from these units with the apparent polar wander path supports this interpretation. This timing corresponds with eruptions of the North Shore Volcanic Group. The rapid emplacement of Duluth Complex intrusions and these eruptions bear similarities to the geologically short duration of well-dated large igneous provinces. These data support hypotheses that call upon the co-location of lithospheric extension and anomalously hot upwelling mantle undergoing decompression melting. This rapid magmatic pulse occurred more than 10 million years after initial magmatic activity following more than 20° of latitudinal plate motion. A likely scenario is one in which upwelling mantle encountered the base of Laurentian lithosphere and flowed via "upside-down drainage" to the locally thinned lithosphere of the Midcontinent Rift.

1   **Rapid emplacement of massive Duluth Complex intrusions within the**  
2   **Midcontinent Rift**

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9   **ABSTRACT**

10   The Duluth Complex is one of the largest mafic intrusive complexes in the world. It was  
11   emplaced as the Midcontinent Rift developed in Laurentia's interior during an interval of  
12   magmatism and extension from *ca.* 1109 to 1084 Ma. This duration of magmatic activity is  
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14   from decompression melting of upwelling mantle plumes. While the overall duration of  
15   magmatic activity was protracted, there were intervals of more voluminous magmatism.  
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17   Complex constrain these units to have been emplaced *ca.* 1096 Ma in less than 1 million  
18   years. Comparison of paleomagnetic data from these units with the apparent polar wander  
19   path supports this interpretation. This timing corresponds with eruptions of the North  
20   Shore Volcanic Group. The rapid emplacement of Duluth Complex intrusions and these  
21   eruptions bear similarities to the geologically short duration of well-dated large igneous  
22   provinces. These data support hypotheses that call upon the co-location of lithospheric  
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24   rapid magmatic pulse occurred more than 10 million years after initial magmatic activity  
25   following more than 20° of latitudinal plate motion. A likely scenario is one in which  
26   upwelling mantle encountered the base of Laurentian lithosphere and flowed via "upside-  
27   down drainage" to the locally thinned lithosphere of the Midcontinent Rift.

28 **INTRODUCTION**

29 The Midcontinent Rift represents a protracted tectonomagmatic event in the interior of  
30 Laurentia (the North American craton). Voluminous outpouring of lava and emplacement  
31 of intrusions accompanied rift development (Fig. 1). Magmatic activity initiated *ca.* 1109  
32 Ma and continued until *ca.* 1084 Ma (Swanson-Hysell et al. 2019). Preserved thicknesses of  
33 the volcanic successions range from nearly 10 km for partial sections exposed on land, such  
34 as along the North Shore of Minnesota (Green et al., 2011), to ~25 km under Lake Superior  
35 (Cannon, 1992). These volcanics and associated intrusions are much more voluminous than  
36 is typical for a tectonic rifting event. Analysis of seismic data leads to an estimate that total  
37 eruptive volume exceeded  $2 \times 10^6 \text{ km}^3$  and that a much greater volume was added to the  
38 lithosphere as intrusions and an underplate (Cannon, 1992). The ~25 Myr duration of  
39 volcanism in the Midcontinent Rift is much longer than is typical for large igneous province  
40 emplacement associated with decompression melting of an upwelling mantle plume. Well-  
41 dated large igneous provinces, such as the Central Atlantic Magmatic Province (Blackburn  
42 et al., 2013), the Karoo-Ferrar (Burgess et al. 2015), and the Deccan Traps (Schoene et al.  
43 2019; Sprain et al. 2019) have durations of <1 Myr for the bulk of their magmatism. An  
44 explanation for prolonged volcanism in the Midcontinent Rift could attribute rift initiation  
45 and initial volcanism via plume arrival with continued volcanism resulting from rift-driven  
46 asthenospheric upwelling. However, the most voluminous period of magmatism occurred  
47 more than 10 million years after initial flood volcanism during an interval known as the  
48 “main magmatic stage” (Vervoort et al. 2007). Main stage magmatism has been attributed  
49 to an upwelling mantle plume based both on the large volume and geochemical signatures  
50 (Nicholson and Shirey 1990; White and McKenzie 1995).

51 Pioneering Midcontinent Rift geochronology utilized  $^{207}\text{Pb}/^{206}\text{Pb}$  dates on zircon from  
52 volcanics (Davis and Green 1997) and intrusions (Paces and Miller 1993) to illuminate the  
53 magmatic history. Subsequent advances in U-Pb geochronology enable higher precision  
54  $^{206}\text{Pb}/^{238}\text{U}$  dates to be used when chemical abrasion methods have mitigated Pb-loss  
55 (Mattinson 2005). An updated chronostratigraphic framework for Midcontinent Rift  
56 volcanics was recently published (Swanson-Hysell et al. 2019) that included new U-Pb  
57 dates developed using these methods (Fig. 2). With these higher precision constraints, the

58 timing and tempo of magmatic activity within the rift can be reevaluated. Of particular  
59 interest is whether magmatic activity was continuous or punctuated by pulses. Key to  
60 evaluating this question is the timing of emplacement of intrusive rocks throughout the  
61 Midcontinent Rift, particularly the largest intrusive suite – the Duluth Complex (Fig. 1).  
62 With its arcuate area of 5630 km<sup>2</sup>, the tholeiitic Duluth Complex is the second-largest  
63 exposed mafic intrusive complex on Earth (Miller et al. 2002). It was emplaced as sheet-like  
64 intrusions into the base of a comagmatic volcanic succession with the majority of its  
65 volume associated with the anorthositic series and the layered series of gabbroic and  
66 troctolitic cumulates (Miller et al. 2002; Fig. 1). We present <sup>206</sup>Pb/<sup>238</sup>U dates from the  
67 Duluth Complex, as well as the Beaver Bay Complex, to establish the duration of Duluth  
68 Complex magmatism and contextualize it with the chronology of volcanism.

69 **METHODS and RESULTS**

70 U-Pb geochronology methods for isotope dilution thermal ionization mass spectrometry  
71 (ID-TIMS) follow Schmitz (2012). Zircon crystals were chemically abraded prior to analysis  
72 in the Boise State Isotope Geology Laboratory. Weighted means were calculated from  
73 multiple single zircon dates with some dates being excluded due to Pb-loss (Fig. 2 and  
74 Table 1).<sup>1</sup> These <sup>206</sup>Pb/<sup>238</sup>U dates can be compared to one another, and to the volcanic  
75 dates of (Swanson-Hysell et al. 2019), at the level of analytical uncertainty (X error in Table  
76 1) given that they all have been developed using EARTHTIME tracer solutions (Condon et  
77 al., 2015). This analytical uncertainty will be referred to when dates are reported and  
78 discussed in the text.

79 The Duluth Complex anorthositic series comprises plagioclase-rich gabbroic cumulates  
80 varying from anorthositic gabbro to anorthosite. Samples FC1 and FC4b are from gabbroic

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<sup>1</sup> GSA Data Repository item 2020XXX, table of individual zircon dates is available online at <http://www.geosociety.org/datarerepository>. All paleomagnetic data and interpreted specimen directions are available to the measurement level in the MagIC database (<https://earthref.org/MagIC/doi/>). All code associated with statistical tests and data visualization is available within a Zenodo repository (currently available for reviewers in this Github repository: [https://github.com/Swanson-Hysell-Group/2020\\_Duluth\\_Complex](https://github.com/Swanson-Hysell-Group/2020_Duluth_Complex)).

81 anorthosite exposures near the former logging town of Forest Center. A weighted mean  
82  $^{206}\text{Pb}/^{238}\text{U}$  date for FC1 of  $1095.81 \pm 0.16$  Ma is calculated based on 10 single zircon dates  
83 (Table 1). The FC4b date is indistinguishable from FC1 with a weighted mean  $^{206}\text{Pb}/^{238}\text{U}$   
84 date of  $1095.71 \pm 0.17$  Ma based on dates from 8 zircons. These dates are indistinguishable  
85 from the weighted mean  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1095.86 \pm 0.19$  Ma developed from  
86 chemically-abraded zircons of gabbroic anorthosite sample AS3 collected from the  
87 anorthositic series in the vicinity of Duluth (Schoene et al., 2006; Fig. 2, Table 1).

88 The layered series of the Duluth Complex is a suite of stratiform troctolitic to gabbroic  
89 cumulates that were emplaced as discrete intrusions (Fig. 1). The PRI sample is a coarse-  
90 grained augite troctolite from the Partridge River intrusion which is at the base of the  
91 complex in contact with underlying Paleoproterozoic metasedimentary rocks (Fig. 1). Data  
92 from 6 zircons result in a weighted mean  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1096.19 \pm 0.19$  Ma (Fig. 2).  
93 The BEI sample is a coarse-grained olivine gabbro from the Bald Eagle intrusion. This  
94 intrusion has been interpreted as one of the youngest layered series units based on cross-  
95 cutting relationships inferred from aeromagnetic data (Miller et al. 2002). Dates from 6  
96 zircons of BEI result in a weighted mean  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1095.89 \pm 0.19$  Ma (Fig. 2) that  
97 is indistinguishable from the anorthositic series dates.

98 The Beaver Bay complex is a suite of dominantly hypabyssal intrusions that cross-cut the  
99 North Shore Volcanic Group (Fig. 1). Sample HCT is an augite troctolite from the  
100 Houghtaling Creek troctolite. The medium-grained olivine-plagioclase cumulates of this  
101 intrusion are interpreted to have been emplaced as a macrodike (Miller et al. 2001). While  
102 some zircons from HCT have Pb-loss that was not fully mitigated by chemical abrasion,  
103 dates from 4 concordant zircons result in a weighted mean  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1095.44 \pm$   
104  $0.26$  Ma (Fig. 2). A sample of coarse-grained ferrodiorite was collected as WLFG from the  
105 Wilson Lake ferrogabbro of the Beaver Bay Complex. This plug-shaped zoned intrusion was  
106 emplaced into the roof zone of the Duluth Complex. Dates from 5 zircons result in a  
107 weighted mean  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1091.63 \pm 0.35$  Ma (Fig. 2). This date overlaps within  
108 uncertainty with the  $^{206}\text{Pb}/^{238}\text{U}$  date of  $1091.61 \pm 0.14$  Ma from a Silver Bay intrusion of  
109 the Beaver Bay Complex (Fairchild et al., 2017; Figs. 1 and 2).

110 Paleomagnetic data from sites of the layered series (37 sites) and the anorthositic series  
111 (11 sites) in the vicinity of the city of Duluth were published in (Beck 1970) (Fig. 3).  
112 Statistical tests show site directions of the layered and anorthositic series to share a  
113 common mean, consistent with their overlapping U-Pb dates. In order to have  
114 paleomagnetic data directly paired with the geochronology, oriented cores were collected  
115 and analyzed from the sites of the FC1, FC4 and HCT samples. Magnetization was measured  
116 on a 2G DC-SQUID magnetometer at UC Berkeley. Samples underwent alternating field or  
117 thermal demagnetization steps and fits were made using the PmagPy software (Tauxe et al.  
118 2016). While (Beck 1970) did not discuss or implement tilt corrections, the Duluth  
119 Complex and overlying lava flows gently dip towards Lake Superior and the paleomagnetic  
120 data need to be corrected for this tilt. We compile abundant igneous layering orientation  
121 measurements, which are similar to the orientations of overlying lavas and interflow  
122 sediments, and use them for tilt-correction.

123 The rapid progression of poles within the Midcontinent Rift apparent polar wander path  
124 (APWP) enable these paleomagnetic data to give chronological insight. The positions of  
125 poles from the early stage of rift magmatism are quite different from main stage lavas (Fig.  
126 3). The similar position of virtual geomagnetic poles (VGPs) from across the Duluth  
127 Complex layered and anorthositic series, including the FC sites, is consistent with  
128 contemporaneous emplacement and they can be combined into a mean pole (Fig. 3). This  
129 paleomagnetic pole can be compared to a synthesized Midcontinent Rift APWP developed  
130 using an Euler pole inversion of the chronostratigraphically-constrained volcanic poles  
131 (Swanson-Hysell et al. 2019). The Duluth Complex pole lies between the 1100 Ma and 1095  
132 Ma path positions with the  $A_{95}$  uncertainty of the pole overlapping with the two angular  
133 standard deviations ellipse of the 1095 Ma path position. This result is consistent with a *ca.*  
134 1096 Ma age for the layered and anorthositic series and strengthens the correlation with  
135 the volcanics.

## 136 **DISCUSSION**

137 The new U-Pb dates, together with the paleomagnetic data, imply that the bulk of both the  
138 layered series and the anorthositic series of the Duluth Complex were emplaced in less

than 1 million years. With the oldest date of  $1096.19 \pm 0.19$  Ma and the youngest of  $1095.71 \pm 0.17$  Ma, the five dates from these series are within 500,000 years of one another and within 850,000 years if the  $2\sigma$  errors are considered (Fig. 2). This emplacement was coeval with eruption of the upper southeast sequence of the North Shore Volcanic Group (NSVG) which comprises ~7900 meters of lavas and is the thickest exposed Midcontinent Rift volcanic succession. The indistinguishable ages of the anorthositic and layered series, together with coeval NSVG eruptions, indicate that *ca.* 1096 Ma there was a large pulse of melt generation.

The  $1095.44 \pm 0.26$  Ma age of the Houghtaling Creek troctolite is indistinguishable from the younger Duluth Complex dates. This result indicates that this pulse of voluminous magmatic activity is represented in some intrusions within the Beaver Bay Complex. A younger pulse of Beaver Bay Complex magmatism postdates NSVG eruptions as units such as the Beaver River diabase and the Silver Bay intrusions penetrate the youngest NSVG lavas, including the  $1093.94 \pm 0.28$  Ma Palisade Rhyolite (Miller et al. 2001; Swanson-Hysell et al. 2019; Fig. 1). The age of this magmatism is represented by the indistinguishable dates of  $1091.63 \pm 0.35$  Ma for the Wilson Lake ferrogabbro and  $1091.61 \pm 0.14$  Ma from the Silver Bay intrusions (Fig. 2; Table 1). This younger Beaver Bay Complex magmatism is coeval with the eruption of the >5 km thick Portage Lake Volcanics that are exposed on the Keweenaw Peninsula and Isle Royale (Fig. 2).

Rapid emplacement of the voluminous layered and anorthositic series of the Duluth Complex bears similarities to the geologically short duration (<1 Myr) of well-dated continental flood basalt provinces (Burgess et al. 2015; Schoene et al. 2019). This similarity supports the hypothesis put forward by (Green 1983), and advanced by others including (Cannon and Hinze 1992) and (Stein et al., 2015), that the co-location of massive magmatism and rifting is the result of lithospheric extension atop decompression melting of an upwelling mantle plume. Contemporaneous heating of Laurentia lithosphere 600 km to the north of the rift is indicated by thermochronologic data from middle to lower crustal xenoliths (Edwards and Blackburn 2018). Basaltic magma was also emplaced throughout the Southwest large igneous province coeval with rift magmatism, including sills more than 2300 km from Duluth (Bright et al., 2014). That such a broad region of Laurentia

169 lithosphere experienced heating and magmatism supports the hypothesized large-scale  
170 mantle upwelling.

171 Both the *ca.* 1108 early stage and *ca.* 1096 Ma main stage volcanism within the  
172 Midcontinent Rift was voluminous and interpreted to be the result of a plume-related  
173 thermal anomaly. The interpretation that this volcanism is associated with a deep-seated  
174 mantle plume needs to be reconciled with the long duration of magmatism and the rapid  
175 equatorward motion of North America from a latitude of ~54° N *ca.* 1108 Ma during early  
176 stage flood basalt eruptions to ~32° N by the *ca.* 1096 Ma main stage (paleolatitudes for the  
177 location of Duluth, MN). While some of this motion could be associated with true polar  
178 wander, in which the mesosphere and asthenosphere rotated in conjunction with the  
179 lithosphere, (Swanson-Hysell et al. 2019) showed that the record of paleomagnetic poles  
180 requires a substantial component of differential plate tectonic motion. The pulsed nature of  
181 magmatic activity could support a model wherein there were multiple upwelling pulses. As  
182 postulated by (Cannon and Hinze 1992), the initial pulse expressed by *ca.* 1108 early stage  
183 flood basalt volcanism initiated lithospheric thinning. Given the significantly thinned  
184 lithosphere in the Midcontinent Rift region, subsequent positively-buoyant plume material  
185 that encountered Laurentia lithosphere would have experienced “upside-down” drainage  
186 wherein relief at the base of the lithosphere resulted in lateral and upward flow into the  
187 Midcontinent Rift (Sleep 1997; Swanson-Hysell et al. 2014). Flow of upwelling mantle to  
188 where the lithosphere was locally thin would have led to ponding and concentrated  
189 decompression melting within the rift axis region. One scenario is that Laurentia was  
190 migrating over a plume generation zone (Burke et al., 2008) from which multiple deep-  
191 seated mantle plumes upwelled to the lithosphere over that time interval. The first could  
192 have been centered on the present-day Lake Superior region with the second encountering  
193 Laurentian lithosphere and being directed to the rift by upside-down drainage in addition  
194 to driving magmatism in southwest Laurentia. Another scenario is that *ca.* 1096 Ma  
195 magmatism was invigorated by upwelling return flow enhanced by slab avalanche induced  
196 downwelling connected to the rapid plate motion of Laurentia that initiated in the early  
197 stage (Swanson-Hysell et al. 2019). Overall, the constraint that both the anorthositic and  
198 layered series of the Duluth Complex were emplaced in less than 1 million years requires

199 an exceptional thermal anomaly that lead to voluminous rapid melt generation during the  
200 main stage of Midcontinent Rift development.

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207

208 **FIGURE CAPTIONS**

209 Figure 1: Geologic map of NE Minnesota (simplified from (Jirsa et al. 2011)) highlighting  
210 major intrusive complexes of the Midcontinent Rift and showing geochronology sample  
211 locations. U-Pb dates from the anorthositic and layered series of the Duluth Complex  
212 (shown in light and dark blue) indicate rapid emplacement in less than 1 million years

213 Figure 2: Date bar plot of CA-ID-TIMS  $^{206}\text{Pb}/^{238}\text{U}$  dates for Midcontinent Rift volcanics and  
214 intrusives. Each vertical bar represents the date for an individual zircon while the  
215 horizontal lines and grey boxes represent the weighted means and their uncertainty.

216 Figure 3: Left panel: tilt-corrected site mean paleomagnetic directions from anorthositic  
217 and layered series sites of (Beck 1970) in the vicinity of Duluth and from the FC1, FC4, and  
218 HCT sites. Center panel: Virtual geomagnetic poles (VGPs) for sites with  $\alpha_{95} < 15^\circ$  give a  
219 mean pole of:  $188.7^\circ \text{ E}, 35.6^\circ \text{ N}, N=24, A_{95}=3.1, k=92$ . Right panel: Duluth Complex  
220 paleomagnetic pole shown with a synthesized pole path developed using an Euler pole  
221 inversion of Midcontinent Rift volcanic poles.

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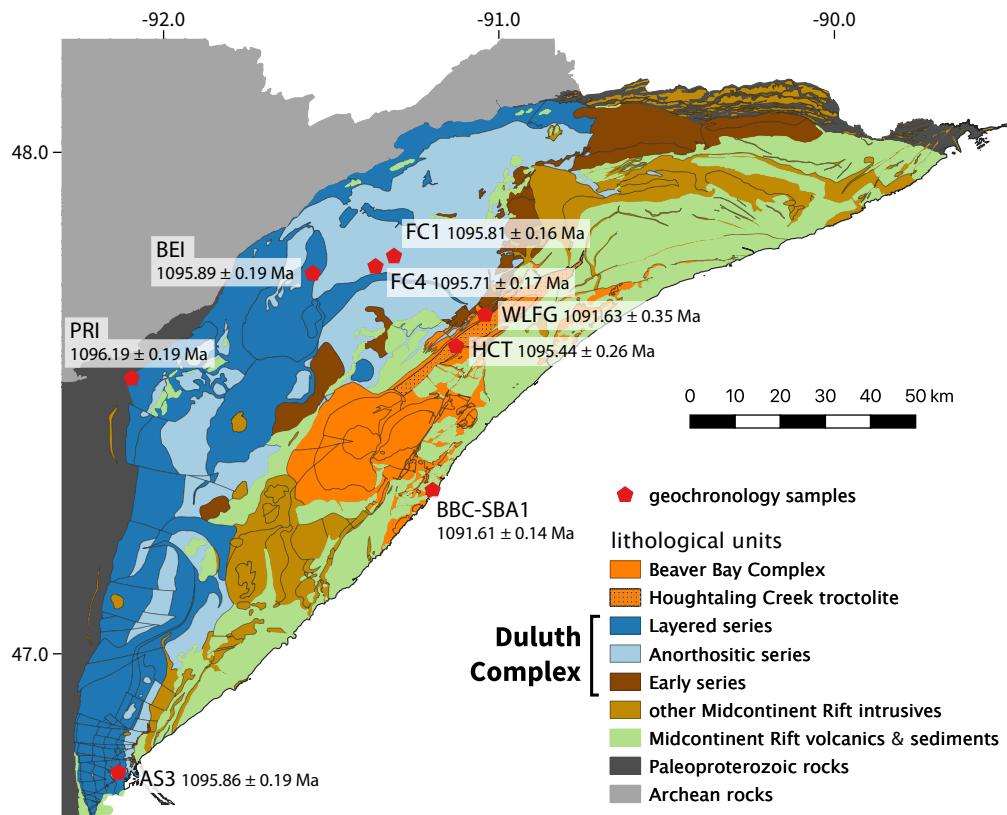
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# Figures and Table

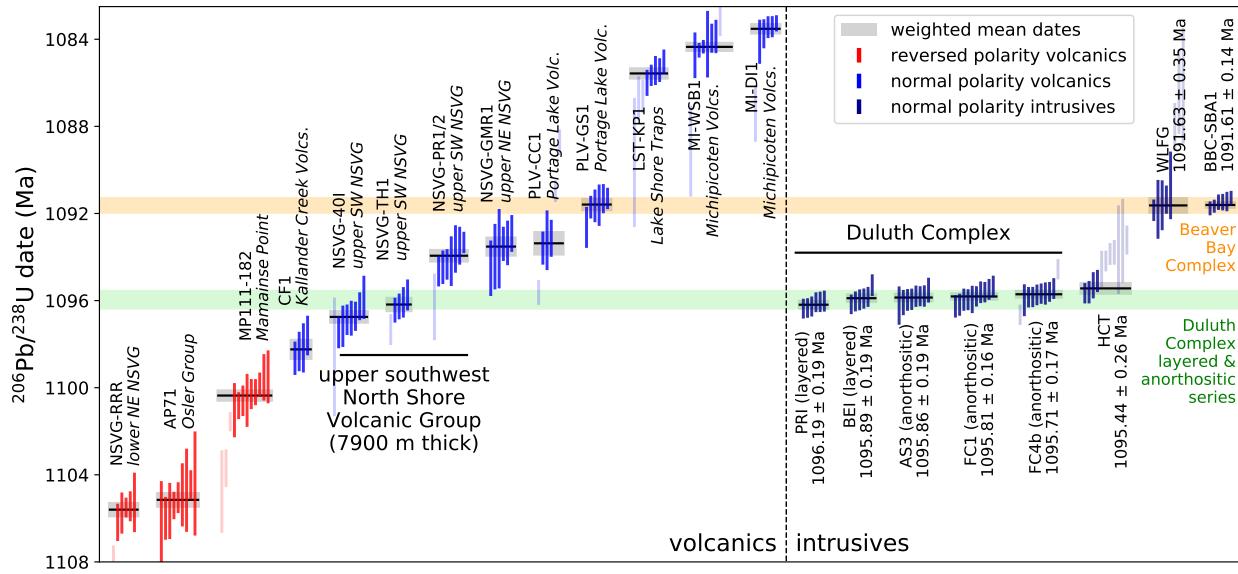
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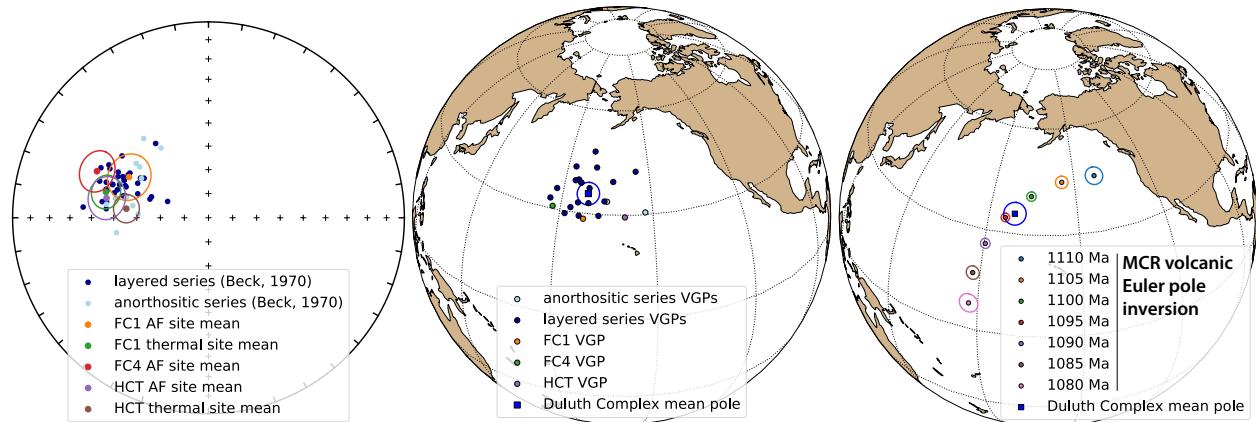
**Figure 1.** Geologic map of NE Minnesota (simplified from Jirsa et al., 2011) highlighting major intrusive complexes of the Midcontinent Rift and showing geochronology sample locations. U-Pb dates from the anorthositic and layered series of the Duluth Complex (shown in light and dark blue) indicate rapid emplacement in less than 1 million years.

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**Figure 2.** Date bar plot of CA-ID-TIMS  $^{206}\text{Pb}/^{238}\text{U}$  dates for Midcontinent Rift volcanics and intrusives. Each vertical bar represents the date for an individual zircon while the horizontal lines and grey boxes represent the weighted means and their uncertainty.



**Figure 3.** Left panel: tilt-corrected site mean paleomagnetic directions from anorthositic and layered series sites of Beck (1970) in the vicinity of Duluth and from the FC1, FC4, and HCT sites. Center panel: Virtual geomagnetic poles (VGPs) for sites with  $\alpha_{95} < 15^\circ$  give a mean pole of:  $188.7^\circ \text{ E}, 35.6^\circ \text{ N}, N=24$ ,  $A_{95}=3.1$ ,  $k=92$ . Right panel: Duluth Complex paleomagnetic pole shown with a synthesized pole path developed using an Euler pole inversion of Midcontinent Rift volcanic poles.

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**Table 1.** Summary of CA-ID-TIMS  $^{206}\text{Pb}/^{238}\text{U}$  dates from Midcontinent Rift intrusions

Sample	Group	Latitude Longitude	$^{206}\text{Pb}/^{238}\text{U}$ date (Ma)	Error ( $2\sigma$ )			MSWD	n
				X	Y	Z		
PRI <i>Partridge River intrusion</i>	Duluth Complex (layered series)	47.5480° N 92.1074° W	1096.19	0.19	0.36	1.15	0.45	6
BEI <i>Bald Eagle intrusion</i>	Duluth Complex (layered series)	47.7516° N 91.5680° W	1095.89	0.19	0.36	1.15	1.59	6
AS3 <i>Duluth anorthosite</i>	Duluth Complex (anorthositic series)	46.7621° N 92.1590° W	1095.86	0.19	0.36	1.15	0.43	8
FC1 <i>Forest Center anorthosite</i>	Duluth Complex (anorthositic series)	47.7827° N 91.3266° W	1095.81	0.16	0.34	1.14	1.44	10
FC4b <i>Forest Center anorthosite</i>	Duluth Complex (anorthositic series)	47.7677° N 91.3753° W	1095.71	0.17	0.35	1.14	0.38	8
HCT <i>Houghtaling Creek troctolite</i>	Beaver Bay Complex	47.6009° N 91.1497° W	1095.44	0.26	0.40	1.16	1.13	4
WLFG <i>Wilson Lake ferrogabbro</i>	Beaver Bay Complex	47.6620° N 91.0619° W	1091.63	0.35	0.46	1.18	0.74	5
BBC-SBA1 <i>Silver Bay aplite</i>	Beaver Bay Complex	47.6620° N 91.0619° W	1091.61	0.14	0.30	1.2	1.0	6

Notes: X=internal (analytical) uncertainty in the absence of external or systematic errors; Y=uncertainty incorporating the U-Pb tracer calibration error; Z=uncertainty including X and Y, as well as  $^{238}\text{U}$  decay constant uncertainty (0.108%; Jaffey et al., 1971). This Z error needs to be utilized when comparing to dates developed using other decay systems (e.g.,  $^{40}\text{Ar}/^{39}\text{Ar}$ ,  $^{187}\text{Re}-^{187}\text{Os}$ ); MSWD=mean square of weighted deviates; n=number of individual zircon dates included in the calculated sample mean date. All dates are from this study with the exceptions of AS3 which was published in Schoene et al. (2006) and BBC-SBA1 which was published in Fairchild et al. (2017). Data for individual zircons are provided in the Data Repository.