## Notes on the function gsw\_pot\_enthalpy\_ice\_freezing\_first\_derivatives(SA,p)

This function,  $gsw_pot_enthalpy_ice_freezing_first_derivatives(SA,p)$ , evaluates the first derivatives of the potential enthalpy (reference pressure of 0 dbar) of ice Ih at the temperature at which it melts into seawater of Absolute Salinity  $S_A$  at pressure p.

The specific enthalpy of ice  $h^{lh}$  is given in terms of the Gibbs function of ice whose independent variables are *in situ* temperature t and pressure p, so that the native functional form of specific enthalpy is  $h^{lh} = h^{lh}(t,p)$ . The potential enthalpy of ice,  $\hbar^{lh}$ , and the potential temperature of ice will always be considered with respect to the reference pressure of 0 dbar, so that  $\theta^{lh} = \theta^{lh}(t,p)$  and

$$\theta^{\mathrm{lh}} = \theta^{\mathrm{lh}}(t, p)$$
 and  $\hbar^{\mathrm{lh}} = h^{\mathrm{lh}}(\theta^{\mathrm{lh}}, 0) = \hbar^{\mathrm{lh}}(\theta^{\mathrm{lh}}).$  (1a,b)

The *in situ* temperature at which seawater is in thermodynamic equilibrium with ice (the freezing temperature) can be expressed as a function of the Absolute Salinity  $S_A$  of the interstitial seawater and the ambient pressure p as

$$t_{\text{freezing}} = t_{\text{freezing}}(S_{\text{A}}, p).$$
 (2)

The potential enthalpy of ice at the freezing temperature is then given by combining the above functional forms as

$$\hbar_{\text{freezing}}^{\text{lh}} = h^{\text{lh}} \left( \theta_{\text{freezing}}^{\text{lh}}, 0 \right) = h^{\text{lh}} \left( \theta^{\text{lh}} \left[ t_{\text{freezing}} \left( S_{\text{A}}, p \right), p \right], 0 \right). \tag{3}$$

From this equation we can derive expressions for the desired partial derivatives of the potential enthalpy of ice with respect to (i) Absolute Salinity and (ii) pressure as

$$\frac{\partial h_{\text{freezing}}^{\text{lh}}}{\partial S_{\text{A}}} \bigg|_{p} = \frac{dh^{\text{lh}}}{d\theta^{\text{lh}}} \frac{\partial \theta^{\text{lh}}}{\partial T} \bigg|_{p} \frac{\partial t_{\text{freezing}}}{\partial S_{\text{A}}} \bigg|_{p}, \tag{4}$$

and

$$\frac{\partial h_{\text{freezing}}^{\text{lh}}}{\partial P}\bigg|_{S_{\text{A}}} = \frac{dh_{\text{lh}}^{\text{lh}}}{d\theta_{\text{lh}}^{\text{lh}}} \left( \frac{\partial \theta_{\text{lh}}^{\text{lh}}}{\partial T} \bigg|_{P} \frac{\partial t_{\text{freezing}}}{\partial P} \bigg|_{S_{\text{A}}} + \left. \frac{\partial \theta_{\text{lh}}^{\text{lh}}}{\partial P} \bigg|_{T} \right).$$
(5)

We now need to evaluate the right-hand sides of these two equations. Expressions for the two partial derivatives of the freezing *in situ* temperature are well-known in the TEOS-10 literature, and here we repeat these expressions from Eqns. (C3a) and (C3b) of McDougall et al. (2014), (here the Gibbs function of ice is  $g^{\text{lh}} = g^{\text{lh}}(t,p)$  and the Gibbs function of seawater is  $g = g(S_{\text{A}},t,p)$ )

$$\frac{\partial t_{\text{freezing}}}{\partial S_{\text{A}}} \bigg|_{\text{P}} = \frac{S_{\text{A}} g_{S_{\text{A}}} S_{\text{A}}}{g_{T} - S_{\text{A}} g_{TS_{\text{A}}} - g_{T}^{\text{lh}}} , \tag{6}$$

and

$$\frac{\partial t_{\text{freezing}}}{\partial P}\bigg|_{S_{\text{A}}} = -\frac{g_P - S_{\text{A}} g_{PS_{\text{A}}} - g_P^{\text{lh}}}{g_T - S_{\text{A}} g_{TS_{\text{A}}} - g_T^{\text{lh}}} .$$
(7)

These partial derivatives are available in the GSW Oceanographic Toolbox from the function <code>gsw\_t\_freezing\_first\_derivatives</code>(SA,p,saturation\_fraction), and in this work on frazil ice and seawater mixtures we will always take the saturation\_fraction of air in the seawater to be zero in order to obtain thermodynamic consistency between the Gibbs functions

The first term on the right-hand sides of both Eqns. (4) and (5) can be readily calculated from the Gibbs function of ice Ih as the isobaric specific heat capacity evaluated at a temperature of  $\theta^{lh}$  and at 0 dbar, namely

$$\frac{d\hbar^{\text{lh}}}{d\theta^{\text{lh}}} = c_p^{\text{lh}} \Big( \theta^{\text{lh}}, 0 \Big), \tag{8}$$

and this derivative can be evaluated using the GSW function **gsw\_cp\_ice**(t,p).

In order to evaluate Eqns. (4) and (5) we also need

$$\frac{\partial \theta^{\text{lh}}}{\partial T}\Big|_{p}$$
 and  $\frac{\partial \theta^{\text{lh}}}{\partial P}\Big|_{T}$ , (9)

and we find these by differentiating the entropy equality

$$\eta^{\text{Ih}}(t,p) = \eta^{\text{Ih}}(\theta^{\text{Ih}}[t,p,0],0), \tag{10}$$

first with respect to *in situ* temperature and then with respect to pressure (see appendix A.15 of IOC et al. (2010) for motivation based on the case of seawater rather than of ice Ih). This procedure leads to

$$\frac{\partial \theta^{lh}}{\partial T}\bigg|_{p} = \frac{g_{TT}^{lh}(t,p)}{g_{TT}^{lh}(\theta^{lh},0)} = \frac{\left(T_{0} + \theta^{lh}\right)c_{p}^{lh}(t,p)}{\left(T_{0} + t\right)c_{p}^{lh}(\theta,0)}, \tag{11}$$

and

$$\frac{\partial \theta^{\text{lh}}}{\partial P}\bigg|_{T} = \frac{g_{TP}^{\text{lh}}(t, p)}{g_{TT}^{\text{lh}}(\theta^{\text{lh}}, 0)} . \tag{12}$$

We are done. Now we can combine the above expressions, and achieve some benefit of cancellation, so that the desired expressions of Eqns. (4) and (5) become

$$\left. \frac{\partial h_{\text{freezing}}^{\text{lh}}}{\partial S_{\text{A}}} \right|_{p} = \frac{\left(T_{0} + \theta^{\text{lh}}\right)}{\left(T_{0} + t\right)} c_{p}^{\text{lh}}\left(t, p\right) \frac{\partial t_{\text{freezing}}}{\partial S_{\text{A}}} \right|_{p}, \tag{13}$$

and

$$\left| \frac{\partial h_{\text{freezing}}^{\text{lh}}}{\partial P} \right|_{S_{A}} = \left| \frac{\left( T_{0} + \theta^{\text{lh}} \right)}{\left( T_{0} + t \right)} c_{p}^{\text{lh}} \left( t, p \right) \frac{\partial t_{\text{freezing}}}{\partial P} \right|_{S_{A}} - \left( T_{0} + \theta^{\text{lh}} \right) g_{TP}^{\text{lh}} \left( t, p \right) \right|. \tag{14}$$

These Eqns. (13) and (14) are evaluated in the present GSW code,

**gsw\_pot\_enthalpy\_ice\_freezing\_first\_derivatives**(SA,p), with every occurrence of *in situ* temperature t being the freezing the *in situ* temperature  $t_{\text{freezing}} = t_{\text{freezing}} \left( S_{\text{A}}, p \right)$  and every occurrence of the potential temperature of ice  $\theta^{\text{lh}}$  being  $\theta^{\text{lh}}_{\text{freezing}} = \theta^{\text{lh}}_{\text{freezing}} \left( S_{\text{A}}, p \right)$ .

## References

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