

$^{40}\text{Ar}/^{39}\text{Ar}$ Thermochronology: Diffusion and Dodson

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Learning Outcomes

- You will gain an understanding of:
 - Diffusion in minerals and rocks
 - Closure temperature
 - The assumptions behind the Dodson T_C formulation
- You will be able to:
 - Calculate closure temperatures

Diffusion

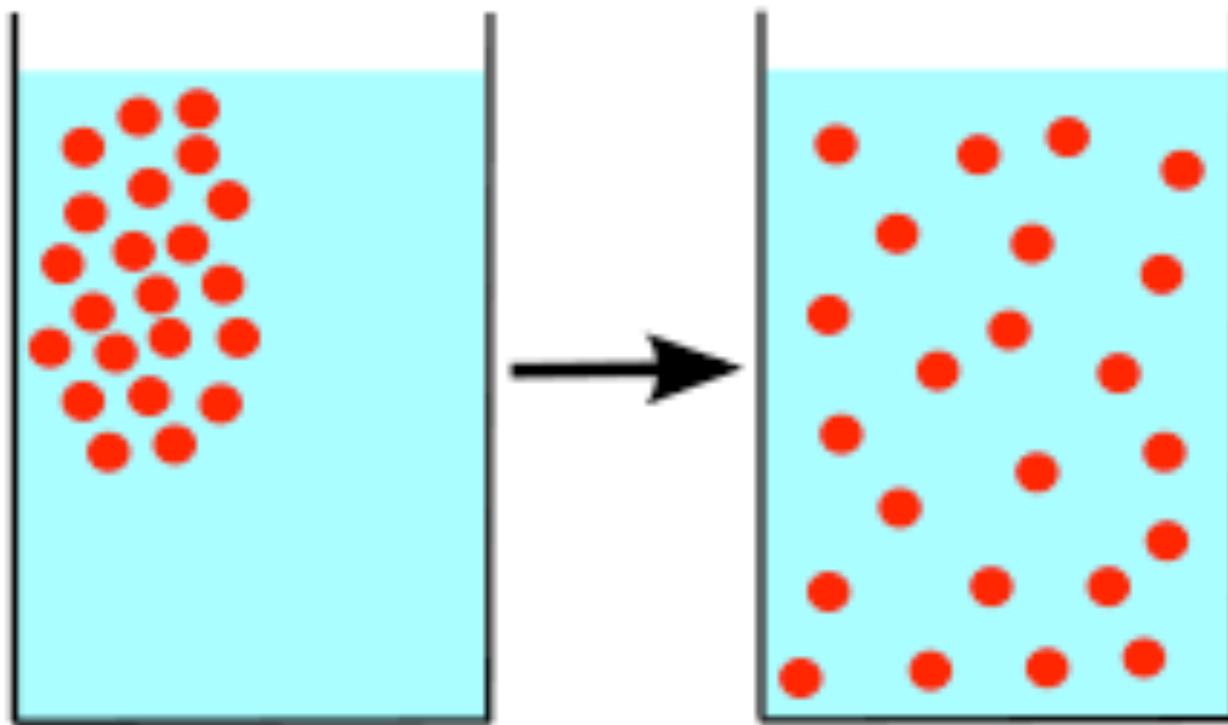


Image: Wikipedia

Diffusion in Minerals

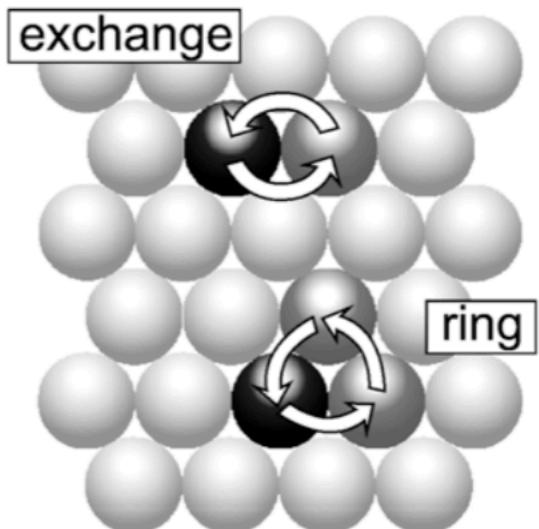
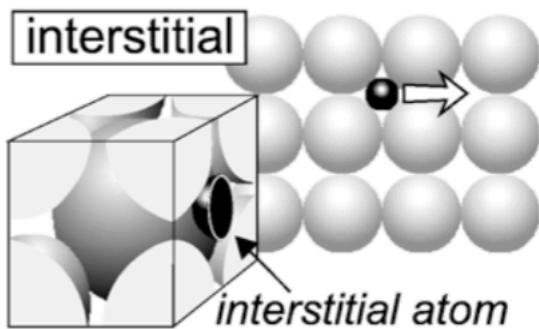
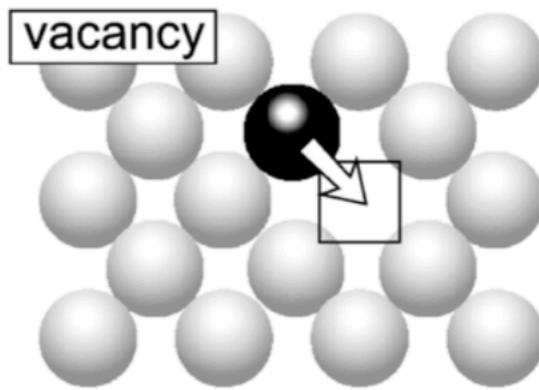
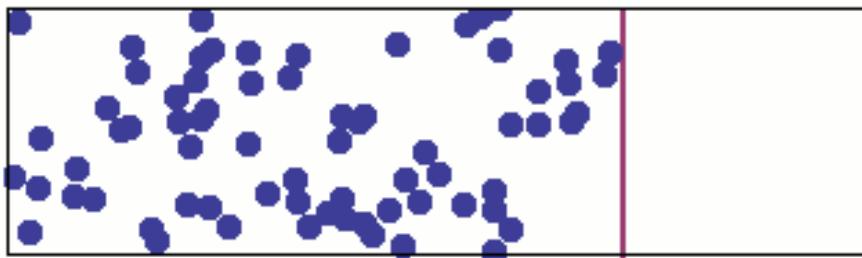


Image: Watson + Baxter 2010



Low concentrations: random walk



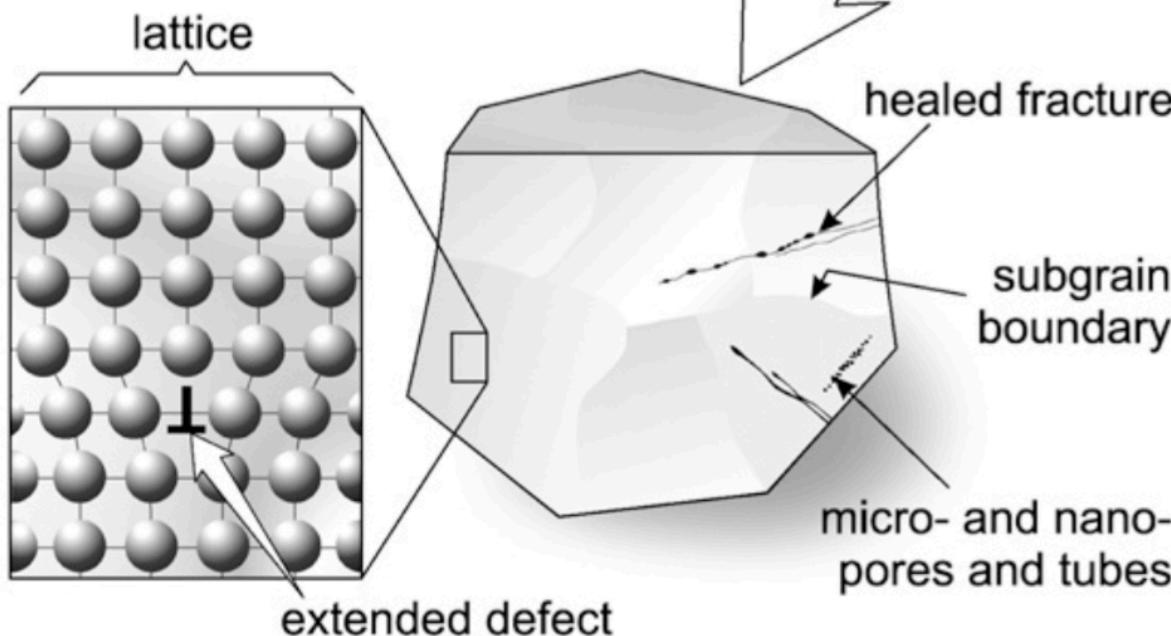
High concentrations: smooth and deterministic

Bulk-Rock Diffusion Pathways

- a. intragrain
- b. grain boundary
- c. grain-edge fluid

$$b + c = \text{ITM}$$

Intragrain Pathways



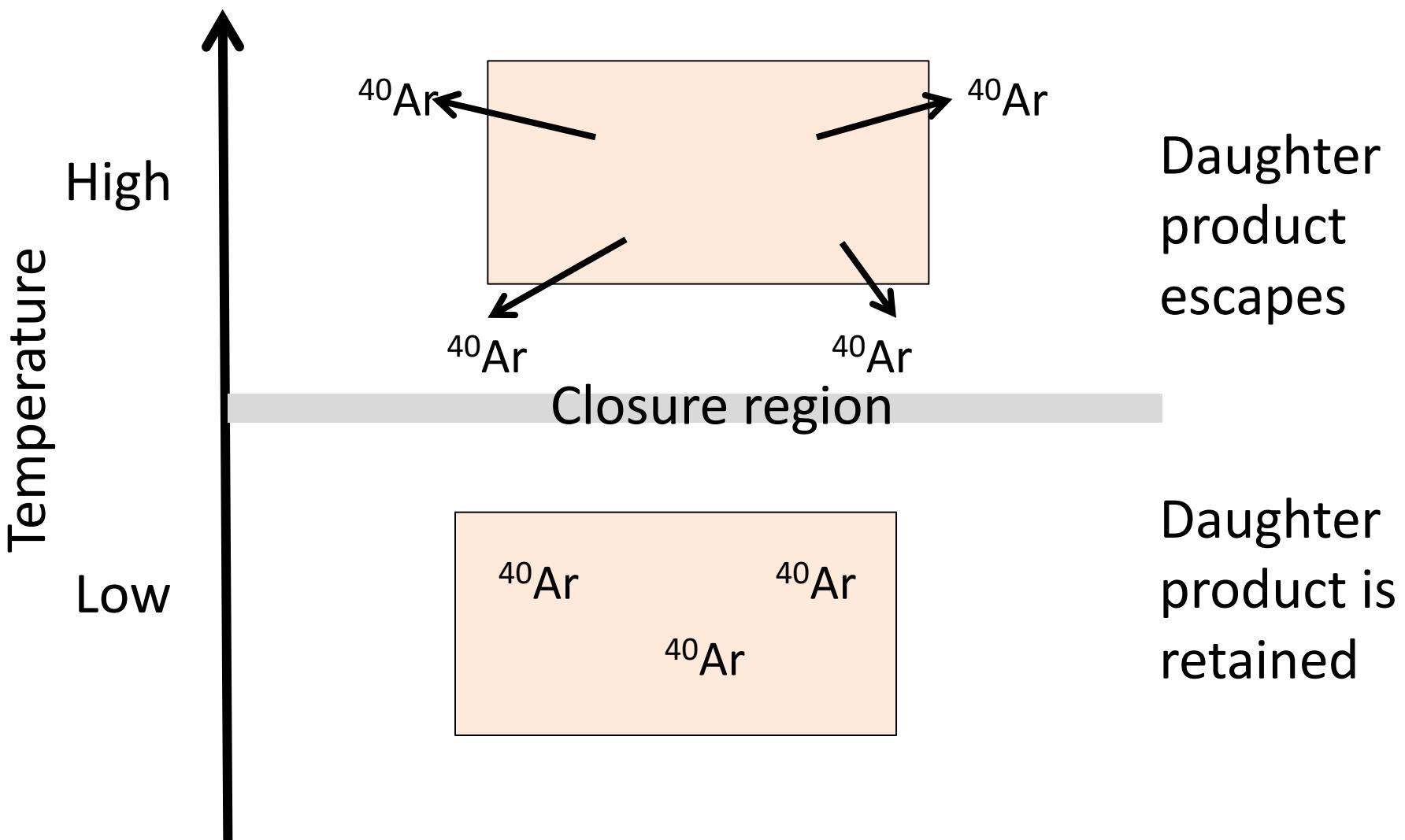
Diffusion in Rocks

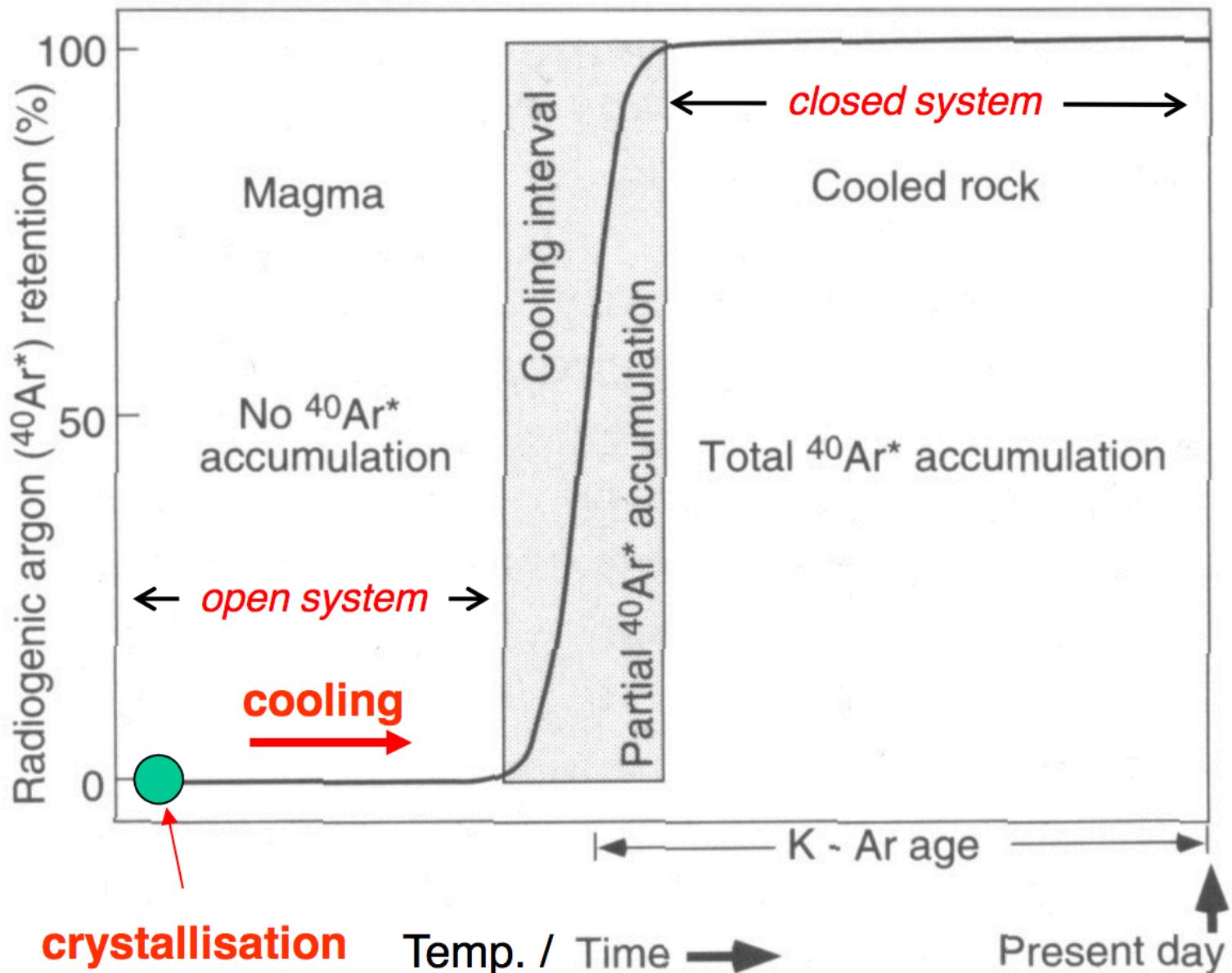
Not volume diffusion

Could be barriers or pathways

Image: Watson + Baxter 2010

Closure Temperature





Fick's 2nd law of diffusion

$$\frac{\partial c}{\partial t} = D \nabla^2 c + S$$

This equation describes the concentration (of argon) in 3 dimensions as a function of space and time

C= concentration

D = diffusion coefficient

∇^2 = (x,y,z directions)

S = source term (Ar is being created over time)

Linking age to T: Dodson T_c

$$\frac{\partial c}{\partial t} = D \nabla^2 c + S$$

Note: Mistake in abstract of Dodson paper: E_a and R transposed!

$$T_c = E_a / [R \ln(A\tau D_0 / a^2)]$$

R = gas constant

E_a = activation energy

A = shape factor (line, plane, sphere)

τ = time constant with which D diminishes with T

D_0 = diffusion coefficient

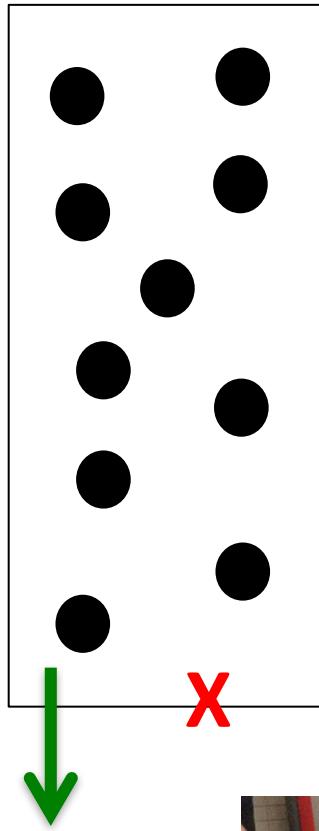
a = diffusion radius

$$\tau = RT^2/(E dT/dt)$$

Note: dT/dt must be decreasing

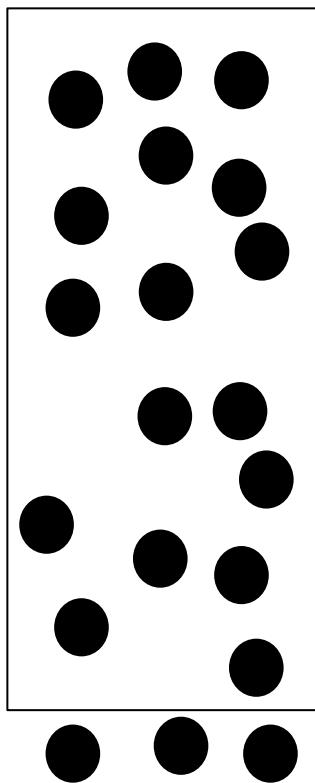
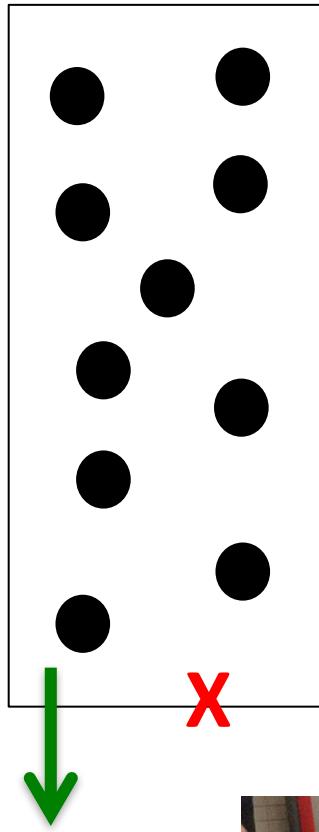
Dodson, CMP, 1973

10



10

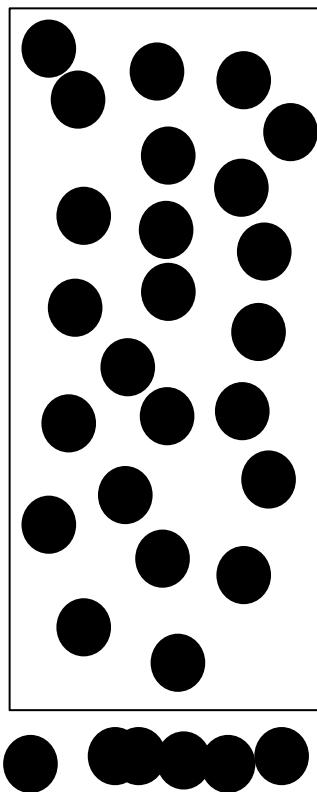
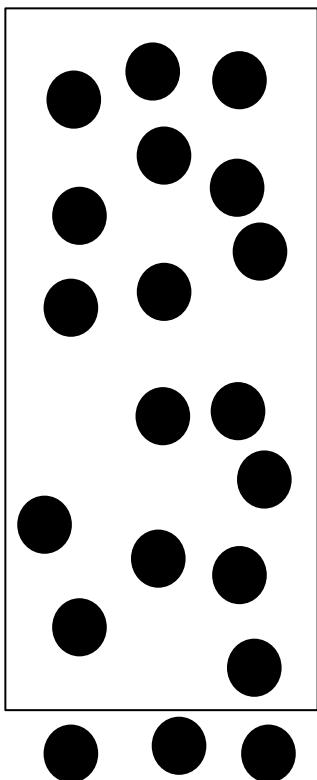
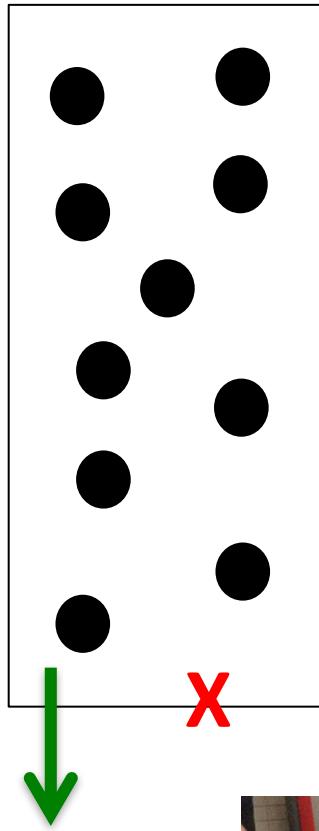
10+10
-3



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10+10
-3

10+10+10
-3-3

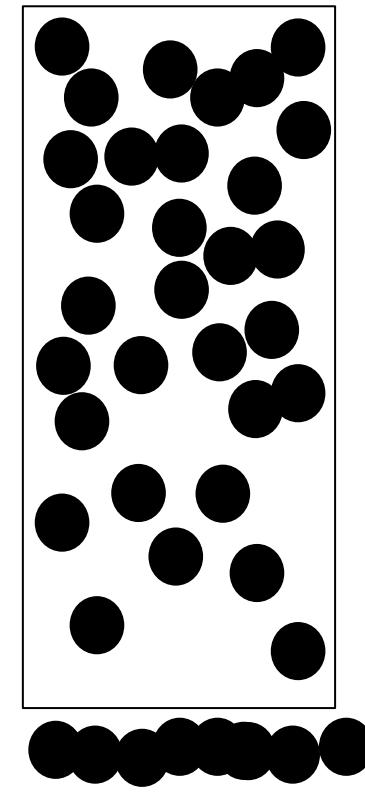
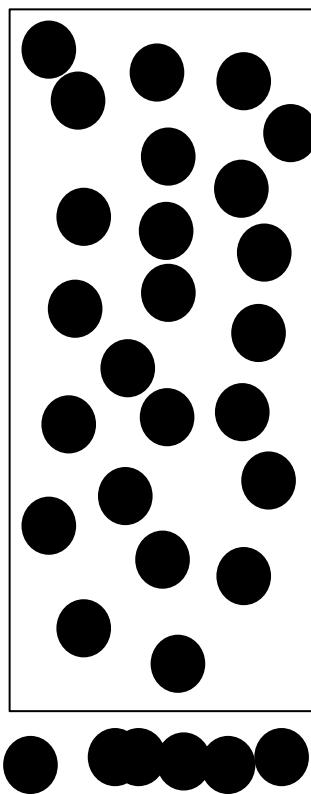
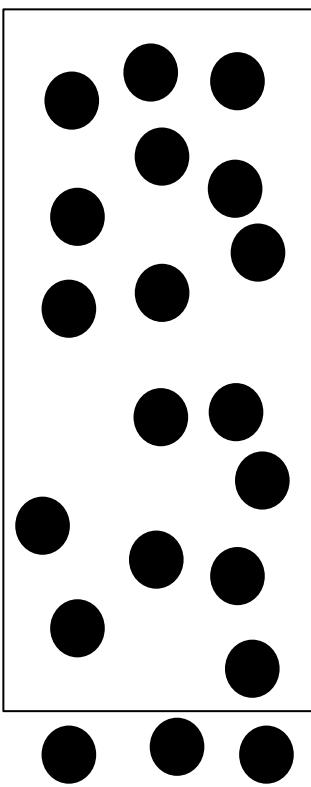
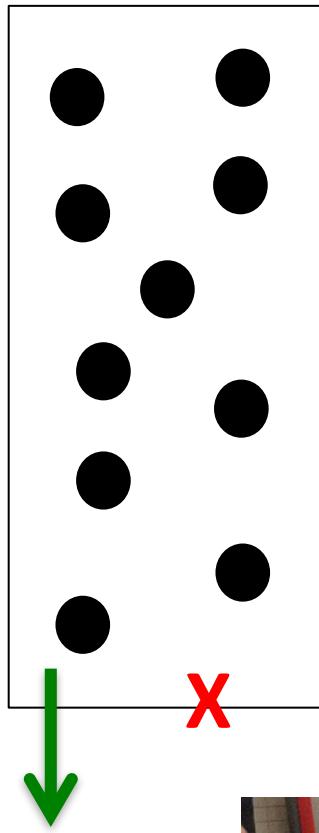


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10+10
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10+10+10
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10+10+10+10
-3-3-3



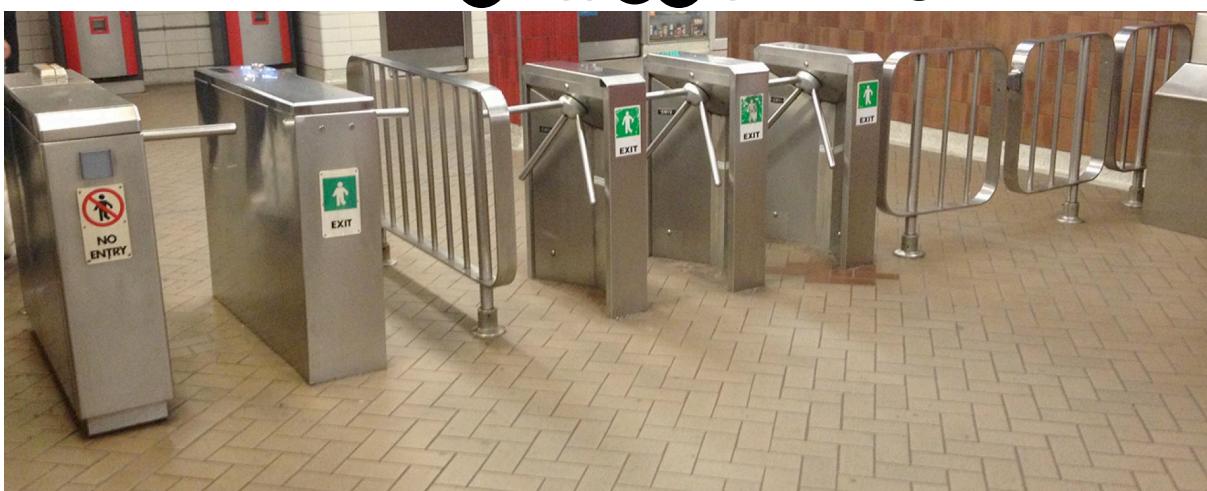
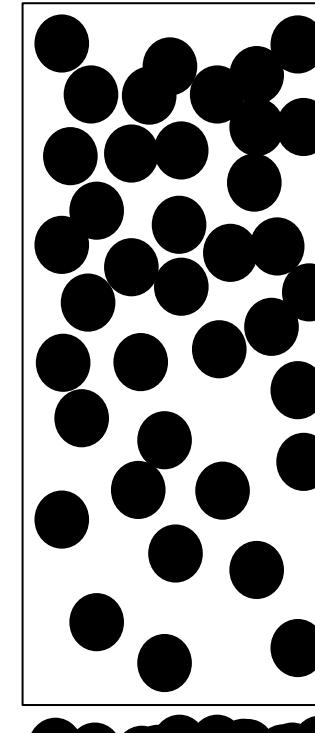
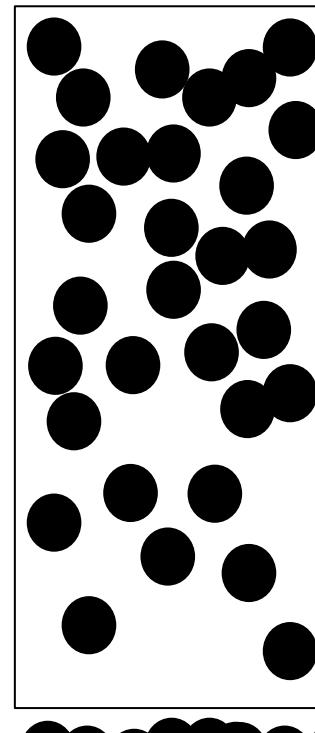
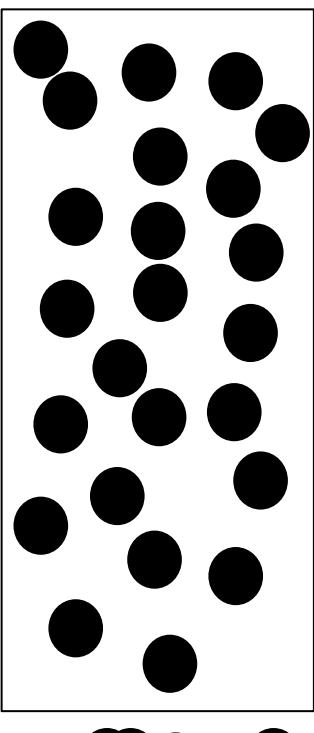
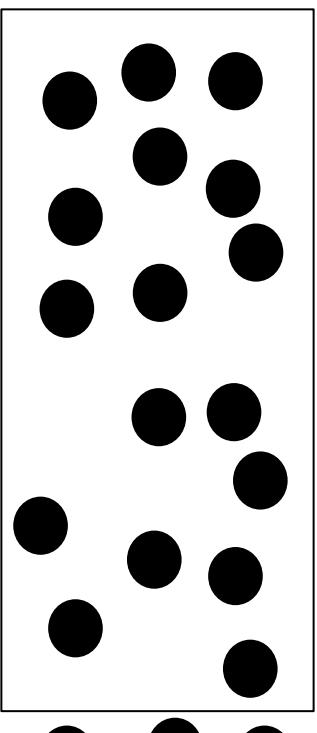
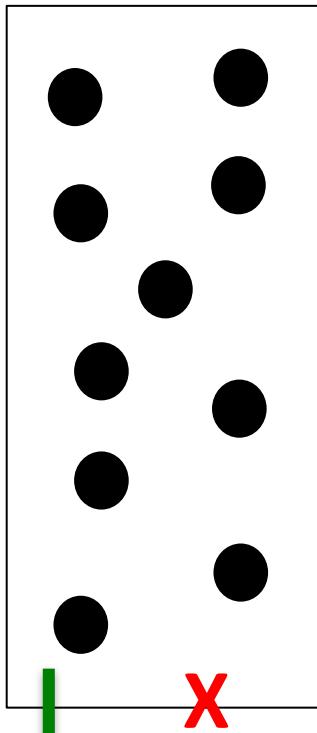
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10+10
-3

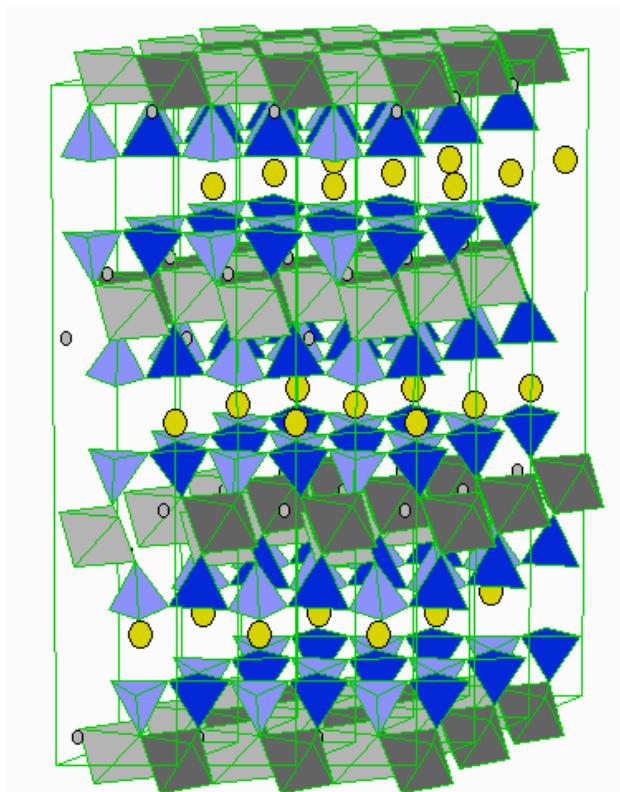
10+10+10
-3-3

10+10+10+10
-3-3-3

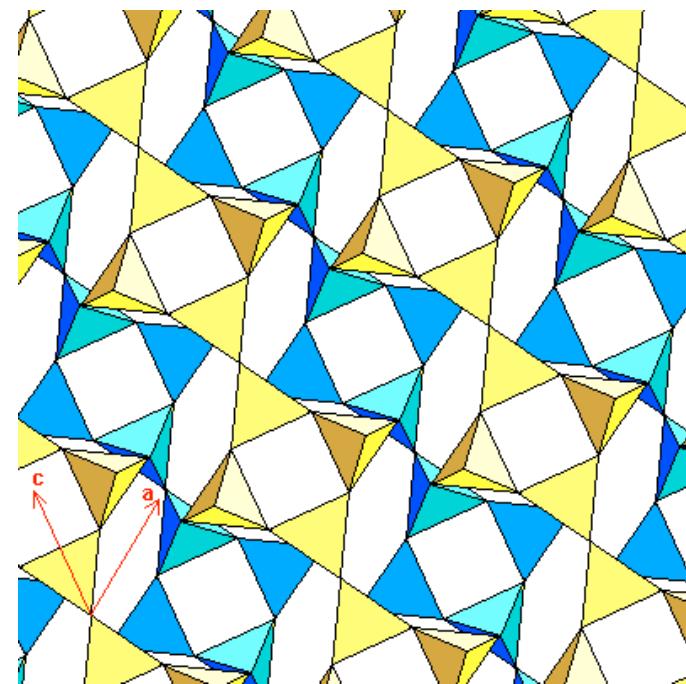
10+10+10+10+10
-3-3-3-3



Remember minerals have different lattice properties



Mica: cylinder (2D)



Feldspar: sphere (3D)

Create a T_c Spreadsheet for muscovite

activity
TIME

Constants:

$$R = 1.986 \times 10^{-3} \text{ kcal/K/mol}$$

$$D_0 = 2.3 \text{ cm}^2 \text{ s}^{-1} \text{ (Harrison et al., 2009)}$$

$$E_A = 63 \text{ kcal mol}^{-1}$$

$$A = 27 \text{ for cylinder}$$

$$T_c = E_a / [R \ln(A\tau D_0 / a^2)]$$
$$\tau = RT^2 / (E_a dT/dt)$$

Let's use a grain radius of 0.05 cm

And a cooling rate of 10 K Ma⁻¹ (convert to K s⁻¹)

What is T_c ?

Hint: you will need to iterate. Choose a starting temperature to calculate τ , then iterate from that value

Now calculate and plot a graph of:

activity
TIME

- Closure T for muscovite for cooling rates from 0.01 to 100 K Ma⁻¹
- For grain radii of 0.001 to 1 cm

Well done!

- You can use this spreadsheet to calculate Tc of:
- Ar in other minerals
- He in other minerals
- Pb in other minerals....
- Basically diffusion of any radiogenically-produced element in any mineral.

Mineral	D_0 (cm 2 s $^{-1}$)	E_a (kJ mol $^{-1}$)	Reference
Phlogopite (Ann ₄)	$0.75^{+1.7}_{-0.52}$	242 ± 11	Giletti (1974)
Biotite (Ann ₅₆)	$0.077^{+0.21}_{-0.06}$	196 ± 9	Harrison et al. (1985)
Biotite (Ann ₅₆)	$0.015^{+0.022}_{-0.005}$	188 ± 12	Grove and Harrison (1996)
Biotite (Ann ₅₆)	$0.075^{+0.049}_{-0.021}$	197 ± 6	Combined data of Harrison et al. (1985) and Grove and Harri- son (1996)
Biotite (Ann ₅₆)	$0.40^{+0.96}_{-0.28}$	211 ± 9	Grove and Harrison (1996)
Muscovite *	$0.033^{+0.213}_{-0.029}$	183 ± 38	Hames and Bowring (1994)
Hornblende	$0.06^{+0.4}_{-0.01}$	276 ± 17	Harrison (1981)

Now been updated by Harrison et al., 2009

activity
TIME

What approximations are involved in using T_c to link age to temperature?

Linking $^{40}\text{Ar}/^{39}\text{Ar}$ ages to temperature: assumptions

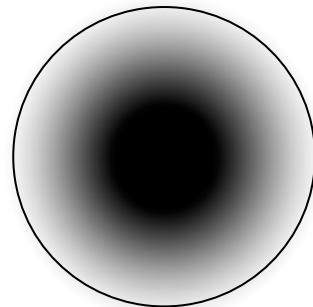
- No initial Ar in grain

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Linking $^{40}\text{Ar}/^{39}\text{Ar}$ ages to temperature: assumptions

- No initial Ar in grain
- Thermally-activated
volume diffusion

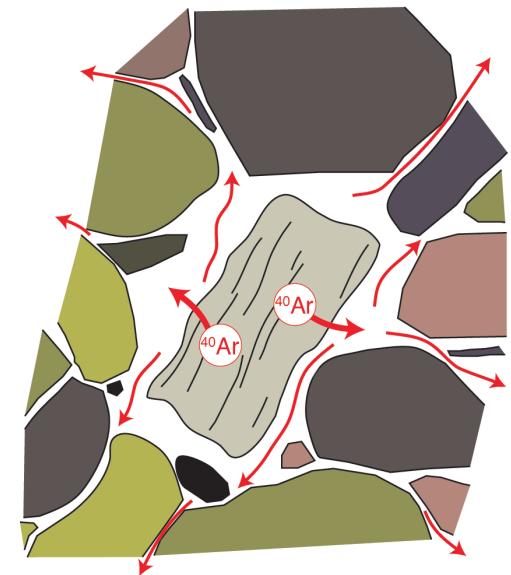
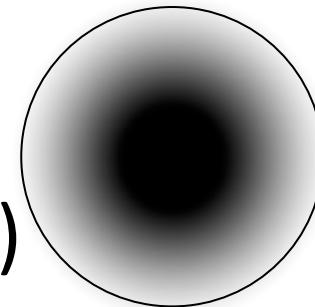
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Linking $^{40}\text{Ar}/^{39}\text{Ar}$ ages to temperature: assumptions

- No initial Ar in grain
- Thermally-activated volume diffusion
- Infinite grain boundary reservoir (open system)

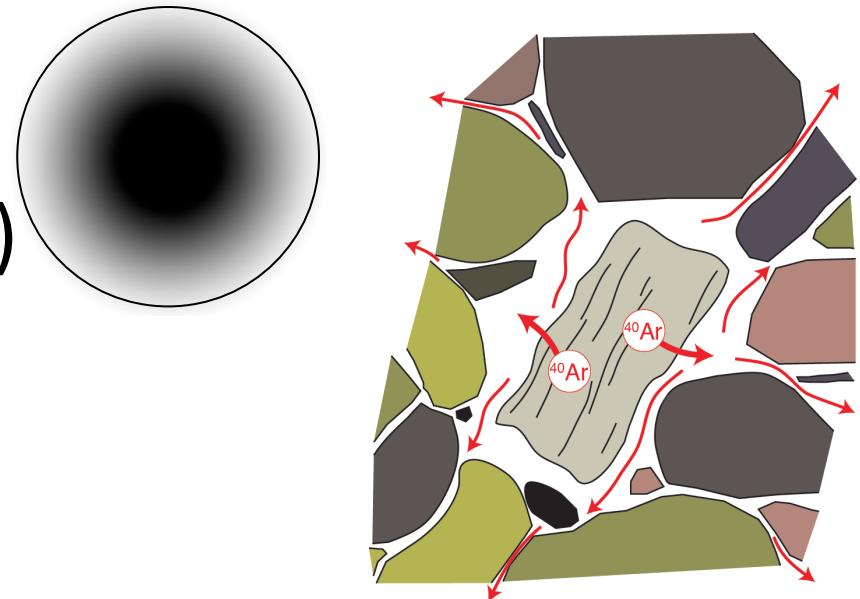
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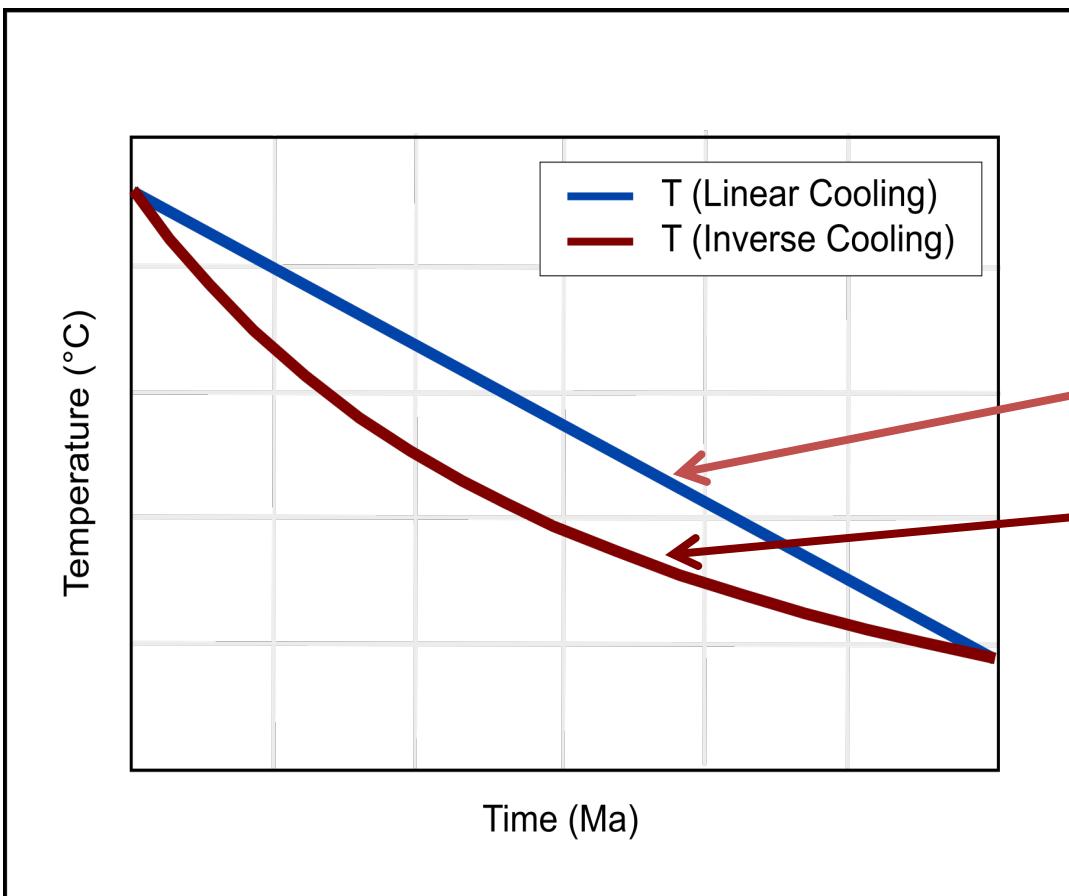
Linking $^{40}\text{Ar}/^{39}\text{Ar}$ ages to temperature: assumptions

- No initial Ar in grain
- Thermally-activated volume diffusion
- Infinite grain boundary reservoir (open system)
- $T_{\text{crystallisation}} \gg T_{\text{closure}}$

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Dodson: Specific $1/T$ cooling path



Our normal
assumption

The Dodson
solution
requirement

Conditions for the Dodson solution

- Initially, concentration = 0

$$\frac{\partial c}{\partial t} = D \nabla^2 c + S$$

- Where does the time dependence sit?

Conditions for the Dodson solution

- Initially, concentration = 0

$$\frac{\partial c}{\partial t} = D \nabla^2 c + S$$

- Where does the time dependence sit?
 - In S (source term)
 - In $D = D_0 \exp(-E_a / RT)$ (temperature linked to time)
 - Analytical solution possible if E_a/RT is linear with time
 - i.e a 1/time temperature dependence (messy!)

Linking $^{40}\text{Ar}/^{39}\text{Ar}$ ages to temperature: assumptions

- No initial Ar in grain
- Thermally-activated

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**Shouldn't use Dodson's formulation
unless these assumptions are satisfied**

How can we tell? That's up next...



Learning Outcomes

- You will gain an understanding of:
 - Diffusion in minerals and rocks
 - Closure temperature
 - The assumptions behind the Dodson T_C formulation
- You will be able to:
 - Calculate closure temperatures