

Analyze of Schwarz algorithms for idealized ocean-atmosphere coupling

Tutor : Éric Blayo et Florian Lemarié

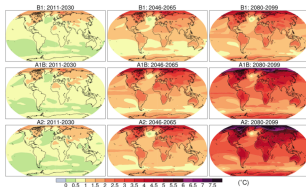
Sophie THERY

21/11/2017

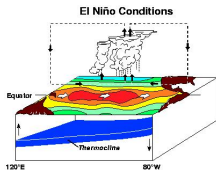
General framework: ocean-atmosphere coupling

Physical phenomena governed by the ocean atmosphere interactions:

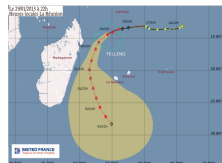
- climate
- El Niño
- tropical cyclones
- ...



climate modeling



seasonal forecasts

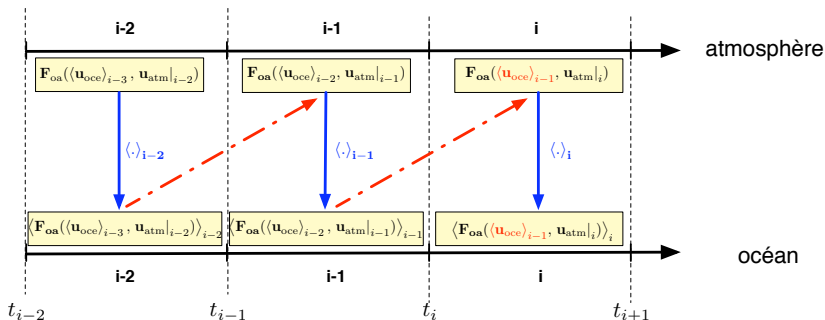


short term predictions

Motivations

The current methods of oceans-atmospheres coupling are unsatisfactory :

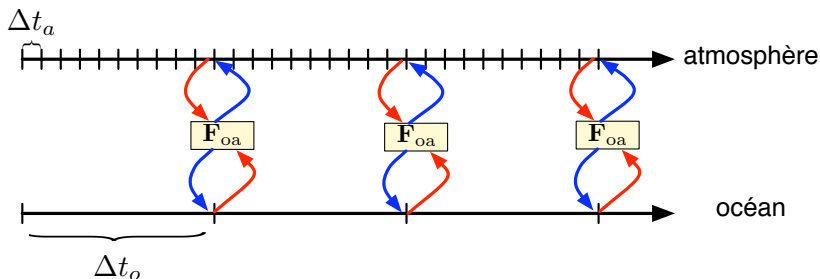
- Asynchrone coupling
 - Balance of the flows average on every window of time
 - problem of synchronisation



Motivations

The current methods of oceans-atmospheres coupling are unsatisfactory :

- Asynchrone coupling
- Synchrone coupling
 - ▶ lot of communication \Rightarrow inefficient implementation
 - ▶ problems of physical validity



Proposed solutions: Schwarz algorithms

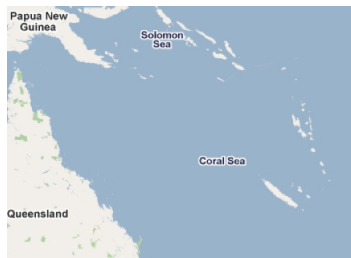
Proof of concept : Simulation of the tropical cyclone Erica (2003) :

- ROMS : Oceanic model

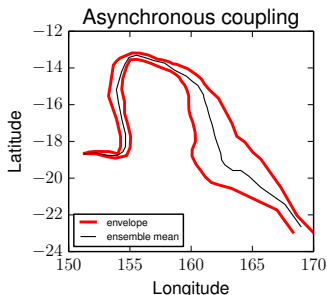
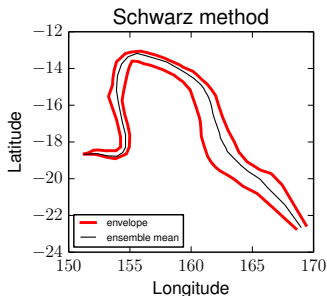
(Shchepetkin-McWilliams, 2005)

- WRF : Atmospheric model

(Skamarock-Klemp, 2007)



Impact



⇒ Reduction of the uncertainty on the trajectory and the intensity of the cyclone.

Lemarié et al. 2014

Aim of the internship: Study of the convergence of the Schwarz algorithms on simplified ocean-atmosphere coupling

1 Présentation du sujet

- Modelling of the ocean atmosphere interaction
- Schwarz algorithms
- Description of the method: example without Coriolis effect
- With Coriolis effect

2 The instationary case

- Convergence factor in the instationary case
- Constant diffusion coefficients
- Linear diffusion coefficients
- Numerical results and optimisation

3 What remains to be done

- Phd

Ocean atmosphere coupling model

Ocean atmosphere coupling model

- Navier-Stokes equations

Ocean atmosphere coupling model

- Navier-Stokes equations
- Dominant physics : vertical axis (1d)

Ocean atmosphere coupling model

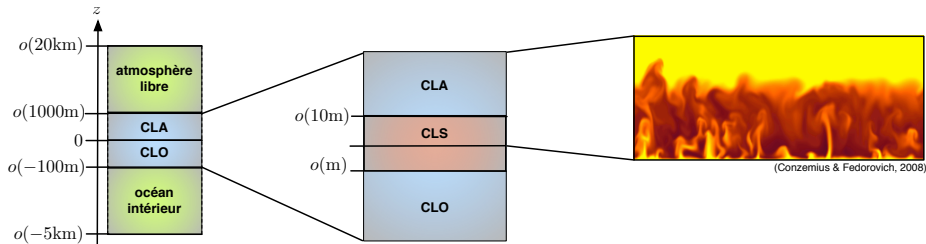
- Navier-Stokes equations
- Dominant physics : vertical axis (1d)
- Dominant terms

$$\begin{cases} \partial_t u - fv + \partial_z(wu) = \nu \partial_{zz}^2 u \\ \partial_t v + fu + \partial_z(wv) = \nu \partial_{zz}^2 v \end{cases}$$

- ▶ (u, v, w) : speed
- ▶ z : altitude
- ▶ ν : molecular diffusion
- ▶ f : Coriolis frequency

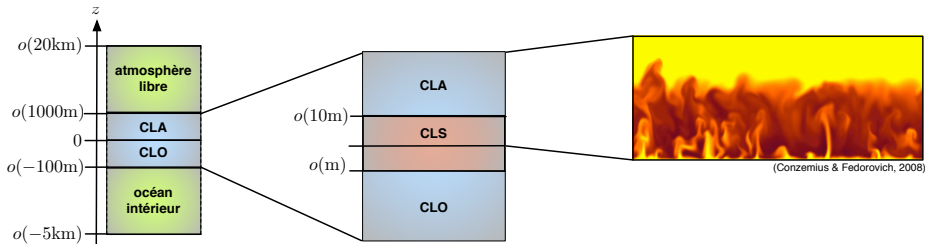
Simplified ocean atmosphere coupling model

Small-scale turbulence parametrisation



Simplified ocean atmosphere coupling model

Small-scale turbulence parametrisation

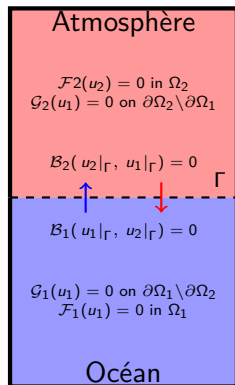


⇒ *Reynolds average*

$$\begin{cases} \partial_t u - fv - \partial_z(D(z)\partial_z u) = F^u \\ \partial_t v + fu - \partial_z(D(z)\partial_z v) = F^v \end{cases}$$

with $D(z) = az + b > 0$,

Our model of study



Sur $\Omega_2 \times [0, T]$:

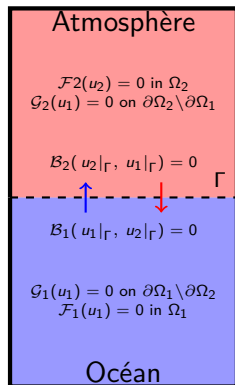
$$\begin{cases} \partial_t u_2 - f v_2 - \partial_z(D_2(z)\partial_z u_2) = F_2^u \\ \partial_t v_2 + f u_2 - \partial_z(D_2(z)\partial_z v_2) = F_2^v \end{cases}$$

Sur $\Omega_1 \times [0, T]$:

$$\begin{cases} \partial_t u_1 - f v_1 - \partial_z(D_1(z)\partial_z u_1) = F_1^u \\ \partial_t v_1 + f u_1 - \partial_z(D_1(z)\partial_z v_1) = F_1^v \end{cases}$$

- + initial conditions
- + outside conditions
- + interfaces conditions

Our model of study



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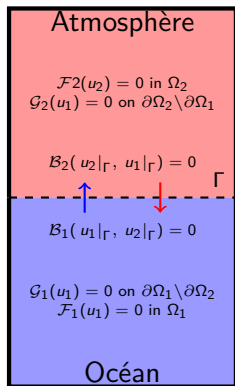
Sur $\Omega_1 \times [0, T]$:

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- + initial conditions
- + outside conditions
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\Rightarrow complexe values $U = u + iv$

Ocean-atmosphere coupling model



$$\left\{ \begin{array}{l} \text{Sur } \Omega_1 \times [0, T] : \\ \partial_t U_1 + ifU_1 - \partial_z(D_1(z)\partial_z U_1) = \tilde{F}_1 \\ \\ \text{Sur } \Omega_2 \times [0, T] : \\ \partial_t U_2 + ifU_2 - \partial_z(D_2(z)\partial_z U_2) = \tilde{F}_2 \\ + \text{Conditions initiales} \\ + \text{Conditions aux limites extérieures} \\ + \text{Conditions d'interfaces} \end{array} \right.$$

Théorie de couche limite :

$$D_j(z) = a_j z + b_j > 0$$

Schwarz algorithms

Algorithm 1 Schwarz algorithms

Require: u_2^0 sur Γ

$n = 0$

while non convergence ou $n < n_{max}$ **do**

solve

$$\begin{cases} \mathcal{L}_1 u_1^n = f_1 & \text{sur } \Omega_1, \\ \mathcal{G}_1 u_1^n = g_1 & \text{sur } \partial\Omega_1^{\text{ext}}, \\ \mathcal{B}_1 u_1^n = \mathcal{B}_1 u_2^{n-1} & \text{sur } \Gamma. \end{cases}$$

then solve

$$\begin{cases} \mathcal{L}_2 u_2^n = f_2 & \text{sur } \Omega_2, \\ \mathcal{G}_2 u_2^n = g_2 & \text{sur } \partial\Omega_2^{\text{ext}}, \\ \mathcal{B}_2 u_2^n = \mathcal{B}_2 u_1^n & \text{sur } \Gamma. \end{cases}$$

end while

State of the art

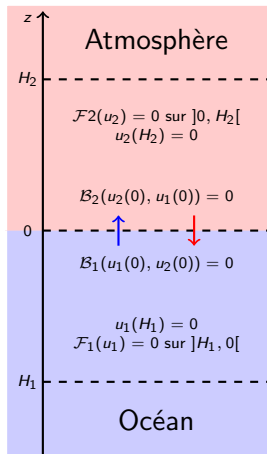
		Coefficients de diffusion constants	Coefficients de diffusion affines
Stationnaire	Sans effet de Coriolis	Lions J.L : <i>On the Schwarz . A variant for nonoverlapping</i>	Alternating method. III subdomains (1990)
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Aim of the internship: Study of the convergence of the Schwarz algorithms on simplified ocean-atmosphere coupling

We focus on two questions

- How the coriolis effect impact the convergence of the algorithm ?
- How "freeze" the diffusion coefficient by a constant impact the convergence ?

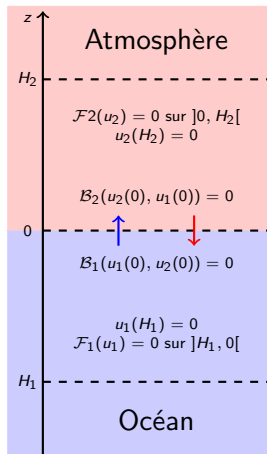
System verified by the errors



- On each area:

$$\begin{cases} -\partial_z(D_j(z)\partial_z u_j^n(z)) = F_j^u & \text{on } \Omega_j \\ u_j^n(H_j) = G_j \\ \mathcal{B}_j u_j^n(0) = \mathcal{B}_j u_k^{n-1}(0) \end{cases}$$

System verified by the errors



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- System verified by the errors :

$$e_j^n(z) = u_j^n(z) - u_j^*(z)$$

$$\begin{cases} -\partial_z(D_j(z)\partial_z e_j^n(z)) = 0 & \text{sur } \Omega_j \\ e_j^n(H_j) = 0 \\ \mathcal{B}_j e_j^n(0) = \mathcal{B}_j e_k^{n-1}(0) \end{cases}$$

Convergence factor

Definition of the convergence factor:

$$\rho = \frac{\|e_j^n\|}{\|e_j^{n-1}\|}$$

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Study of the convergence :

- If $\rho < 1$ then the algorithm converge.
- If $\rho \geq 1$ then the algorithm do not converge.

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Method for the stationary case :

- Solve probleme without interface conditions $\Rightarrow e_j^n(z) = C_j^n \tilde{e}_j(z)$

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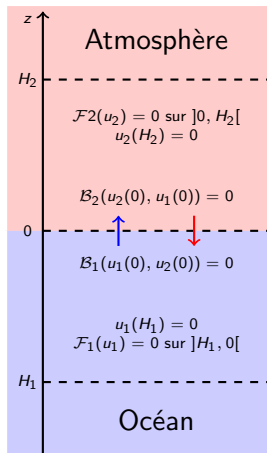
- If $\rho < 1$ then the algorithm converge.
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Method for the stationary case :

- Solve probleme without interface conditions $\Rightarrow e_j^n(z) = C_j^n \tilde{e}_j(z)$
- Convergence factor: $\rho = \left| \frac{C_j^n}{C_j^{n-1}} \right|$

Where C_j^n is determined by the interface conditions.

Dirichlet-Neumann conditions



Dirichlet-Neumann interface conditions :
exemple without Coriolis effect

$$\left\{ \begin{array}{ll} -\partial_z(D_j(z)\partial_z e_1^n(z)) & = 0 \\ & \text{sur }]H_1, 0[\\ e_1^n(H_1) & = 0 \\ e_1^n(0) & = e_2^{n-1}(0) \end{array} \right.$$

$$\left\{ \begin{array}{ll} -\partial_z(D_j(z)\partial_z e_2^n(z)) & = 0 \\ & \text{sur }]0, H_2[\\ e_2^n(H_2) & = 0 \\ D_2(0)\partial_z e_2^n(0) & = D_1(0)\partial_z e_1^n(0) \end{array} \right.$$

Convergence factor in stationary case without Coriolis effect, with Dirichlet-Neumann conditions:

$$\rho_{0,DN} = \frac{\int_{\Omega_2} (D_2(z))^{-1} dz}{\int_{\Omega_1} (D_1(z))^{-1} dz}$$

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- depends on D_j on all the area.
- $D_j(z)$ is not necessary constant.

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- depends on D_j on all the area.
- $D_j(z)$ is not necessary constant.
- particular case :

$$D_j(z) = \text{constant}$$

$$\rho_{0,DN}^{cst} = \frac{H_2 D_1}{H_1 D_2}$$

$$D_j(z) = a_j z + b_j :$$

$$\rho_{0,DN}^{var} = \frac{a_1}{a_2} \frac{\ln(1 + H_2 \frac{a_2}{b_2})}{\ln(1 + H_1 \frac{a_1}{b_1})}$$

Study of the convergence with Coriolis effect

		Coefficients de diffusion constants	Coefficients de diffusion affines
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With Coriolis effect

System to solve :

$$\left\{ \begin{array}{lcl} ife_1^n(z) - \partial_z(D_1(z)\partial_z e_1^n(z)) & = & 0 \quad \text{sur }]H_1, 0[\\ e_1^n(H_1) & = & 0 \\ e_1^n(0) & = & e_2^{n-1}(0) \end{array} \right.$$

$$\left\{ \begin{array}{lcl} ife_2^n(z) - \partial_z(D_2(z)\partial_z e_2^n(z)) & = & 0 \quad \text{sur }]0, H_2[\\ e_2^n(H_2) & = & 0 \\ D_2(0)\partial_z e_2^n(0) & = & D_1(0)\partial_z e_1^n(0) \end{array} \right.$$

- with $D_j = \text{constant} \Rightarrow$ well known solutions
- with $D_j = a_j z + b_j > 0$

With Coriolis effect

System to solve :

$$\left\{ \begin{array}{lcl} ife_1^n(z) - \partial_z(D_1(z)\partial_z e_1^n(z)) & = & 0 \quad \text{sur }]H_1, 0[\\ e_1^n(H_1) & = & 0 \\ e_1^n(0) & = & e_2^{n-1}(0) \end{array} \right.$$

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- with $D_j = \text{constant} \Rightarrow$ well known solutions
- with $D_j = a_j z + b_j > 0$

Bessel's equation

$$z\partial_{zz}^2 u + (2\alpha - 2\beta\nu + 1)z\partial_z u + (\beta^2\gamma^2 z^{2\beta} + \alpha(\alpha - 2\beta\nu))u = 0$$

Convergence factor in stationary case without Coriolis effect, with Dirichlet-Neumann conditions:

- $D_j = \text{constant}$:

$$\rho_{DN}^{cst} = \sqrt{\frac{D_1}{D_2}} \left| \frac{\tanh(H_2 \lambda_2)}{\tanh(H_1 \lambda_1)} \right|$$

with $\lambda_j = \sqrt{i \frac{f}{D_j}}$ and $f \neq 0$.

Convergence factor in stationary case without Coriolis effect, with Dirichlet-Neumann conditions:

- $D_j = \text{constant}$:

$$\rho_{DN}^{cst} = \sqrt{\frac{D_1}{D_2}} \left| \frac{\tanh(H_2 \lambda_2)}{\tanh(H_1 \lambda_1)} \right|$$

with $\lambda_j = \sqrt{i \frac{f}{D_j}}$ and $f \neq 0$.

- $D_j = a_j z + b_j$:

$$\begin{aligned} \rho_{DN}^{var} = & \sqrt{\frac{D_1(0)}{D_2(0)}} \left| \frac{l_0(\mu_2(H_2))K_0(\mu_2(0)) - K_0(\mu_2(H_2))l_0(\mu_2(0))}{l_0(\mu_1(H_1))K_0(\mu_1(0)) - K_0(\mu_1(H_1))l_0(\mu_1(0))} \right| \\ & \times \left| \frac{l_0(\mu_1(H_1))K_1(\mu_1(0)) + K_0(\mu_1(H_1))l_1(\mu_1(0))}{l_0(\mu_2(H_2))K_1(\mu_2(0)) + K_0(\mu_2(H_2))l_1(\mu_2(0))} \right| \end{aligned}$$

with $\mu_j(z) = 2 \sqrt{\frac{if}{a_j} \left(z + \frac{b_j}{a_j} \right)}$

Summary of the stationary case

		Coefficients de diffusion constants	Coefficients de diffusion affines
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Plan

1 Présentation du sujet

2 The instationary case

- Convergence factor in the instationary case
- Constant diffusion coefficients
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3 What remains to be done

Study of the convergence in the instationary case

$$\left\{ \begin{array}{ll} \partial_t e_j^n(t, z) + i f e_j^n(t, z) - \partial_z(D(z)\partial_z e_j^n(t, z)) & = 0 \quad \text{sur }]0, T[\times \Omega_j \\ e_j^n(t, H_j) & = 0 \quad \text{sur }]0, T[\\ \mathcal{B}e_j^n(t, 0) & = \mathcal{B}e_k^n(0) \quad \text{sur }]0, T[\\ e_j^n(0, z) & = 0 \quad \text{sur } \Omega_j \end{array} \right.$$

Study of the convergence in the instationary case

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Fourier transform on the time

$$\hat{u}(\omega, z) = \int_{-\infty}^{\infty} u(t, z) e^{-i\omega t} dt, \text{ with } \omega \in \mathbb{R}$$

Study of the convergence in the instationary case

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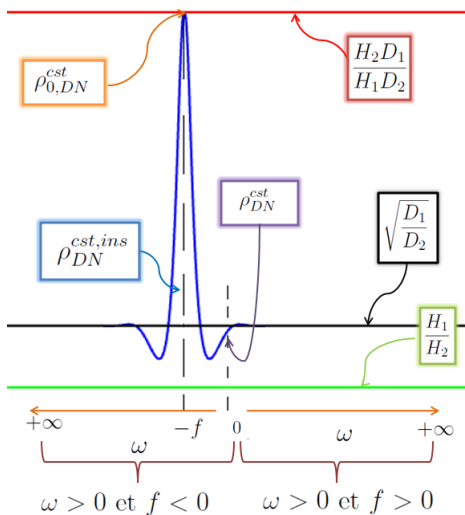
Fourier transform on the time

$$\hat{u}(\omega, z) = \int_{-\infty}^{\infty} u(t, z) e^{-i\omega t} dt, \text{ with } \omega \in \mathbb{R}$$

$$\left\{ \begin{array}{ll} i(f + \omega) \hat{e}_j^n(\omega, z) - \partial_z(D(z)\partial_z \hat{e}_j^n(\omega, z)) & = 0 \quad \text{sur } \Omega_j \\ \hat{e}_j^n(\omega, H_j) & = 0 \\ \mathcal{B} \hat{e}_j^n(\omega, 0) & = \mathcal{B}_j \hat{e}_k^n(\omega, 0) \end{array} \right.$$

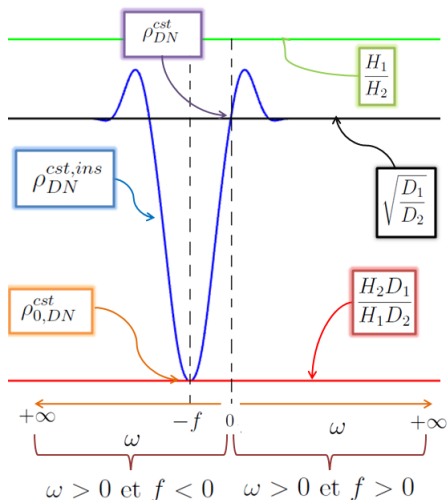
⇒ Stationary case equations with $f \rightarrow f + \omega$

Convergence factor in instationary case with Coriolis effect, with Dirichlet-Neumann conditions :



$$\text{If } \rho_{0,DN}^{var} \geq \sqrt{\frac{b_1}{b_2}}$$

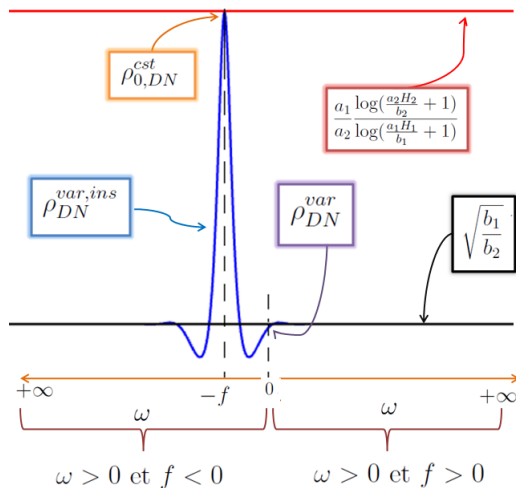
Facteur de convergence pour le cas instationnaire, des coefficients affines et conditions de Dirichlet-Neumann



If

$$\rho_{0,DN}^{var} \leq \sqrt{\frac{b_1}{b_2}}$$

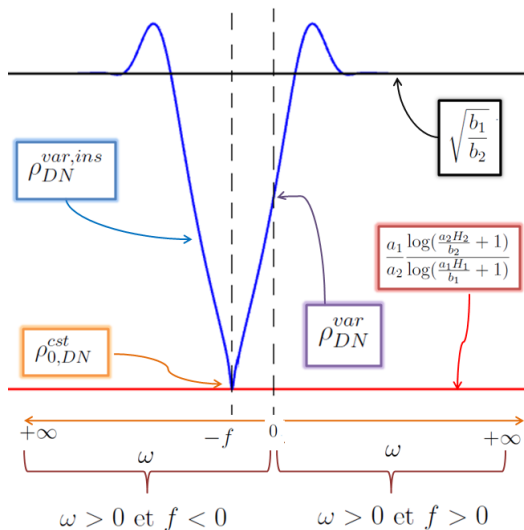
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Instationnaire	Sans effet de Coriolis	Martin, V. <i>Schwarz Waveform Relaxation Methods for oceanographic equations</i> (2003) Bennequin D, Gander M.J, Halpern L : <i>Optimized Schwarz Waveform Relaxation Methods for Convection Reaction Diffusion Problems</i> (2004)	Lemaire, F : <i>Toward an Optimized Global-in-Time Schwarz Algorithm for Diffusion Equations with Discontinuous and Spatially Variable Coefficient.</i> (2013)
	Avec effet de Coriolis	Gander M.L <i>Optimized schwarz methods for model problems with continuously variable coefficients</i> (2016)	

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- How can we optimize the convergence ?

Numerical results

- No disymetries between $f < 0$ et $f > 0$.
- When $\omega \ll |f|$ or $\omega \gg |f|$ theoretical results and numerical results are consistant.
- When $\omega \approx |f|$: dispersion of the numerical results.

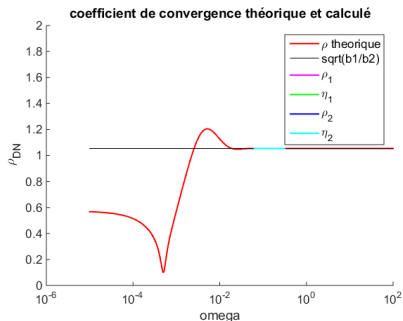


Figure: Theoretical and numerical convergence factor, $f < 0$

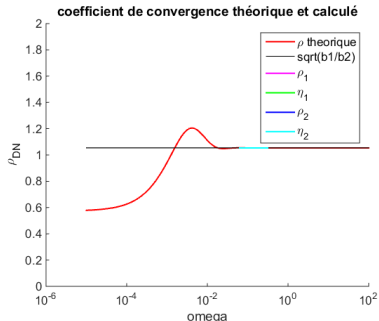


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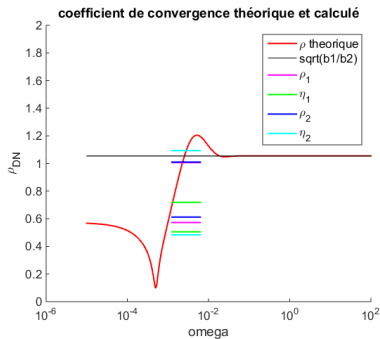


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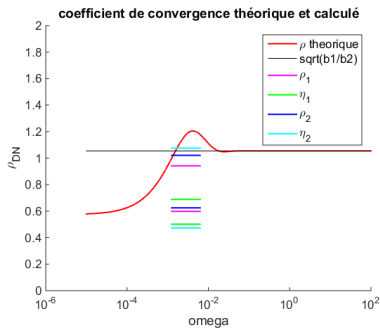


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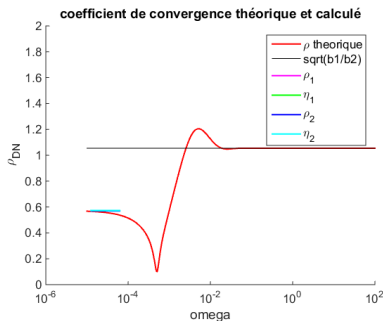


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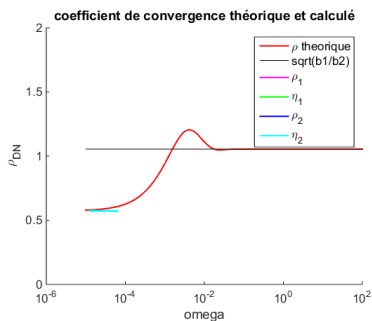


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- Easy to determinate in the stationary case.
- In the instationary case, a numerical method can calculate "good" Robin coefficients.
- Could we be help by the optimal coefficients from the stationary case to find better coefficient in the instationary case ?

During the Phd

- More realistic models.
⇒ more complicated interface conditions.
- Link with Charles's thesis.
- Realistic applications in collaboration with climatologists



Thanks