Generalising the impact of forest cover on streamflow from experimental data: it is not that simple.

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4 Abstract

Three recent papers review and analyze large global datasets related to impacts of forest cover on streamflow. Using three different approaches, they all find a strong relationship between forestation, de-forestation and streamflow. However, the results are problematic, the underlying data set is unbalanced, and there are correlations in the data that warrant further investigation as this would influence the results. For example, the area of the catchment is strongly related to the assessment technique and the variability in the response data. For this study, the data for the recent three papers were reviewed, combined, and supplemented with new studies. Subsequently, the data were re-analyzed using generalised additive modelling. The results highlight that there are four interlinked reasons that make the general outcomes from the previous papers problematic: 1) The existence of latent variables in the data that create the appearance of a relationship that really does not exist; 2) The difficulty in fully interpreting the specifics of different studies; 3) The difficulty of integrating data from seemingly similar studies, but with quite different objectives; and 4) The chance of transcription errors influencing the data. Overall this indicates that while valuable data can be extracted from past studies, the above problems need

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to be considered before results are generalised and extrapolated to continental and global scales.

5 Keywords: data analysis, forest cover change, global scales, literature review

6 1. Introduction

There is a urgent need to identify the impacts of human intervention on streamflow at a global scale and to separate this from climate effects [35, 18]. More specifically, the impacts of global deforestation and reforestation are important through their perceived influence on streamflow and blue and green water availability [18, 28]. The past work reviewing these impacts [2, 19, 41, 9, 11 10, 16] highlights a general consensus that if forest areas increase, streamflow decreases and vice-versa. The most dramatic example of this is Figure 5 in Zhang 13 et al. [39] indicating (for Australian catchments) a 100% decrease in streamflow for catchments with 100% forest cover. However, on the other end of the 15 spectrum, for three French catchments [13], there was no change in streamflow characteristics in two of the catchments after deforestation. For reforestation, a modelling study across the 1 millon km² Murray Darling Basin also found no major effect, especially in larger catchments [32], but a recent study [18] 19 found an 8% change in streamflow as a result of reforestation. Similarly a mod-20 elling study by Beck et al. [3] found no significant change in streamflow in 12 catchments in Puerto Rico as a result of deforestation. In contrast, in a recent study in Brazil across 324 catchments, Levy et al. [22] found a significant increase in streamflow, particular in the dry season, as a result of deforestation. 24 This suggests that there can be significant variation across the different studies, methodologies and geographical regions. For the purpose of this paper, watershed and catchment are interchangeable terms. Many of the US studies use watershed, while European and Australian studies use *catchment*. In particular, we retained the term "paired watershed studies" and "quasi-paired watershed studies" as this is the most common terminology, but further mostly use the term catchment.

There has been a recent push in the hydrological community [14] to use 'meta-analysis' to summarise past studies. The suggestion is that, because meta-analyses use clearly defined search terms and statistical methods to analyze the results, this will lead to more reliable summaries of past research. As a result, several review papers have summarized the plethora of forestation and deforestation studies across the globe, in relation to paired watershed studies [9, 8], related to reforestation in particular [16], and more generally [19, 41]. These studies aim to generalize the individual experimental and research findings and to identify if there are global trends or relationships. Others have used the understanding from a global analysis to extrapolate to global scales [18].

The recent paper by Filoso et al. [16] is a clear meta-analysis, but most others [41, 18, 44] are not. However, n impressive global database of catchment studies 43 with changes in streamflow due to changes in forest cover has been developed [41, 16] and statistical approaches are used to analyze the resulting data. The Zhang et al. [41] dataset, which covers over 312 studies, is described in terms of the change in streamflow as a result of the change in forest cover, where studies related to both forestation (increase in forest cover) and deforestation (decrease 48 in forest cover) were included. In contrast, the paper by Filoso et al. [16] focused primarily on reforestation, and covered an equally impressive database of 167 studies using a systematic review. In this case the collected data is mostly coded as count data and only a subset of 37 studies was analyzed for actual 52 water yield change. There is some overlap between the two data sets, but there are also some studies unique to both sets. The more regionally concentrated and detailed study by Levy et al. [22] is a further independent dataset with no overlap with the other studies. However, for this study only the flow and rainfall data is available for the catchments, and the change in landcover was derived from satellite data and was not made available.

The conclusions of the first mentioned major review paper [41] indicates that

there is a distinct difference in the change in flow as a result of forestation or deforestation between small watersheds (catchments), defined as $< 1000 \text{ km}^2$ 61 and large watersheds (catchments) > 1000 km². While for small catchments 62 there was no real change in runoff with changes in cover, for large catchments there was a clear trend showing a decrease in runoff with increases in forest cover. The main conclusion was that the response in annual runoff to forest cover was scale dependent and appeared to be more sensitive to forest cover change in water limited catchments relative to energy limited catchments [41]. The second study [16] is a systematic review of reforestation studies (only studies in which forest cover increased). This study classified the historical research and highlighted gaps in the spatial distribution, the types of studies and 70 the types of analysis. Their main conclusion was also that reforestation decreases streamflow, but that there were many interacting factors. For a subset of the data (37 data points) they also indicated decreasing impacts of reforestation with increasing catchment size (agreeing with Zhang et al. [41]), but they did not identify a distinct threshold and fitted a log-linear relationship. In addition, 75 they identified that studies with shorter periods of data collection resulted in larger declines in streamflow. 77

An earlier paper, that includes much of the same data as Zhang et al. [41] and Filoso et al. [16], is Zhou et al. [44], which has one author in common with Zhang et al. [41]. However, this paper aims to explain the variation in the data using the elasticity approach in the Fuh model, which is similar to well-known Budyko approaches [40]. In particular, it aims to link the variation in the observed data to variations in the exponent m in the Fuh model. A key

observation is that in drier environments, the effects of removing forest cover are much greater than in wetter environments, which is also suggested by Figure 4 in Zhang et al. [41]. The Fuh model and the related variations of the Budyko equilibrium modelling approach was also used by Hoek van Dijke et al. [18] to interpret the global impact of reforestation.

However, concerning is that there are some clear limitations in these studies, and some of this applies to meta-analyses in general. The main method in the work by Zhang et al. [41] is a single covariate linear regression. In contrast, the systematic review from Filoso et al. [16] mainly emphasises the classification and distributions of the study. Zhang et al. [41] points out that a main assumption in their work is that the catchment size threshold at 1000 km² is a distinct separation between "small" and "large" catchments. However, a subset of 37 data points in Filoso et al. [16] (their Figure 9) does not appear to support this, suggesting a continuum. And while the work Filoso et al. [16] provides important insights in study types, analysis types, forest types and broad classification, there is limited quantification of actual impact.

In contrast to the single covariate linear regression in the earlier studies [41, 16] and the top-down Budyko modelling [44, 18], the regional Brazilian Cerrado study [22] provides an example of an carefully designed statistical approach using mixed effects modelling and Differences-in-Differences modelling focusing specifically on the effect of deforestation. The analysis specifically accounted for differences between catchments and differences due to variations in climate.

Not all datasets are however suitable for this kind of in-depth analysis.

Given all these previous reviews and the seemingly clear conclusions about the impact of forest cover change on streamflow, the question is why another review paper on this topic? There is a real attraction in the concept of statistical analysis of past studies encapsulated in meta-analysis to be able to extrapo-

- late findings to larger scales, and to identify factors across global scales [14].

 However, there are also some hidden complications in this that can invalidate
 results, which this paper aims to highlight. There are four potential errors (or
 limitations) in such global meta-analyses:
- Impact of latent variables that are not included in the typical single covariate analysis;
- Interpretation errors due to incomplete descriptions of the experiments in
 the original papers;
- Aggregation of data that originates from different experiments with different objectives across a wide time period, but have similar keywords; and, finally

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 Transcription errors in the data, especially if data is collected from other review papers as some of the original papers are difficult to locate.

The aim of this paper is to first reanalyze the global dataset [41, 16] using 124 some more detailed statistical modelling and to use this to highlight examples 125 of each of these limitations. This will show how they have influenced the out-126 comes of the past work, and provide suggestions of how we can overcome these 127 limitations. In addition, by applying more complex statistical models, we will 128 highlight the conclusions that can be drawn from the data. Finally, we will 129 highlight future research needs in the area of forest cover change impact on 130 streamflow. 131

We are taking advantage of the earlier work by Zhang et al. [41], Filoso et al. [16] and Zhou et al. [44] and the large database of studies these authors have shared.

35 2. Methods

in the models.

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2.1. The original data set

As indicated, the starting point of this paper is the data base of studies which 137 were included in Zhang et al. [41] as supplementary material. The columns 138 in this data set are the catchment number, the catchment name, the Area in 139 km², the annual average precipitation (Pa) in mm, the forest type, hydrological 140 regime, and climate type, the change in forest cover in % ($\Delta F\%$) and the change 141 in streamflow in % ($\Delta Qf\%$), the precipitation data type, the assessment tech-142 nique, and the source of the info, which is a citation. The change in streamflow 143 $(\Delta Qf\%)$ is based on equation 1 in Zhang et al. [41]. Several of these columns contain abbreviations to describe the different vari-145 ables, which are summarised in Table 1. These abbreviations will later be used

Table 1: Summary of abbreviations of factors used in the Zhang et al. (2017) data set

Factor	Abbreviation	Definition
forest type	CF	coniferous forest
	BF	broadleaf forest
	MF	mixed forest
hydrological regime	RD	rain dominated
	SD	snow dominated
climate type	EL	energy limited
	WL	water limited
	EQ	equitant
precipitation data type	ОВ	observed

Factor	Abbreviation	Definition
	SG	spatial gridded
	MD	modelled
assessment technique	PWE	paired watershed experiment
	QPW	quasi-paired watershed
		experiment
	$_{ m HM}$	hydrological modelling
	$\mathbf{E}\mathbf{A}$	elasticity analysis
	SH	statistical modelling and
		hydrographs

The paper by Zhang et al. [41] also uses the dryness index, which is the annual rainfall (Pa) divided by the potential or reference evapotranspiration (ET₀ or E₀) in their analysis, and have used the dryness index to identify the climate type. However, the potential or reference ET used for this calculation was originally not included in the published data set. We will discus below how we derived the dryness index in our data set. We combined the tables for small catchments ($< 1000 \text{ km}^2$) and large catchments ($>= 1000 \text{ km}^2$) from Zhang et al. [41] in our analysis.

2.2. Additional data collection

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To enhance the existing data set, this study added additional variables and cross-checked the studies with the data set from Filoso et al. [16]. In particular, we focused on the 37 data points related to the quantitative regression analysis used in Filoso et al. [16].

In addition, a few additional variables were included to enhance the data set. We added latitude and longitude for the center of the catchment as an approximation of its spatial location. Mostly the data reported by the authors
was used, but in some cases the variables had to be approximated from the
location of the centre of the catchment using Google MapsTM. In the dataset,
an additional column has been added to indicate the source of the location data
to indicate if this is directly from the paper or elsewhere.

As highlighted, Zhang et al. [41] did not provide values for evapotranspiration in the data base. Using the location information, reference evapotranspiration (E_0) was extracted from the Global Aridity Index and Potential Evapo-Transpiration (ET_0) Climate Databasev2 [31], if a value of E_0 was not available from the original papers. For large catchments, this value (and the associated coordinates), similar to annual average rainfall, is only an approximation of the climate at the location.

Similar to Zhang et al. [41], the Dryness index was calculated from the catchment estimate of reference evapotranspiration and the catchment estimate of annual average rainfall (Pa) as:

$$Dryness = \frac{E_0}{Pa} \tag{1}$$

The length of the study can be a variable influencing the change in flow [e.g. 178 19, 16], as for example, more mature plantations are thought to have smaller 179 impacts on flow or regrowth might follow a "Kuczera curve" [21]. It is not clear 180 if this is an effect of increased water use in growth [33] or due to changes in 181 interception [29]. Therefore, the length of the study calculate as the difference 182 between the starting data and completion date of the different studies was ex-183 tracted from the references provided by Zhang et al. [41]. The length of the 184 study was already included in the data from Filoso et al. [16], but these were 185 checked against the original publications.

Several additional data points from catchment studies were extracted from

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Almeida et al. [1], Ferreto et al. [15], Zhang et al. [39], Zhao et al. [42], Borg et al. [7], Thornton et al. [30], Zhou et al. [43], Rodriguez et al. [26], Ruprecht et al. [27] and Peña-Arancibia et al. [24], and these were checked against the existing studies to prevent overlap. In the citation column in the accompanying data set, the main reference for the calculated change in streamflow was generally used, because sometimes the original study did not provide the quantification of the change in streamflow [i.e. Table 6 in 39].

We conducted a thorough review of all the studies mentioned in the data base of Zhang et al. [41] and sourced all the original papers. As a result of this we made several changes to the data base, which are all recorded in Supplementary Data part 1. Overall 36 data points were changed and the most common problem was a change in the sign for the change in forest cover or the change in flow. We assume that these were transcription errors.

We also removed one data point from the data set, which corresponds to catchment #1 (Amazon) in Zhang et al. [41]. This is because the cited reference [25]
only relates to 1 and 1.5 ha paired catchment studies in French Guyana, and in
which the actual change in forest cover is not recorded. Finally, on review of all
the data in Zhang et al. [41] and Filoso et al. [16], 29 potential duplicates were
identified and flagged in the data, and not used in the analysis.

The final column in the improved data base is a "notes" column, which we added, but is not further used in the analysis. It gives context to some of the data for future research and highlights some of the discrepancies that we found between the original papers and the data in the tables from Zhang et al. [41]. This will allow future research to scrutinise our input for errors.

2.3. Statistical modelling

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The aim of the statistical analysis is to highlight the most important variables in the data set that explain the change in flow as a result of changes in forest cover. This first aim is similar to Zhang et al. [41], but the main difference is
that we start off with all variables in the data set in the model. Subsequently the
analysis will concentrate on how the individual variables in the dataset relate
to each other and how latent variables in the data set can be masked and result
in relationships that might not really exist. Finally, the analysis will highlight
how the results are conditional on the dataset.

In the statistical analysis we are not necessarily seeking the best "predictive" model, and as such do not perform a traditional variable selection process.

Rather, we focus on analyzing the predictor variables in the full model to identify how all the variables explain the variance in the dependent variable.

To estimate how the change in streamflow is affected by the change in forest cover, while considering the effects of the other variables, we applied generalised additive modelling (GAM) [38].

The general model tested is:

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$$\Delta Qf\% \sim \Delta\% forest\ cover +$$

$$\sum X_i + \sum s(Z_i) + \varepsilon \tag{2}$$

Here X_i are factorial variables, while Z_i are continuous variables. As a 229 first step, the model assumes no direct interactions and that all variables are 230 additive. A further assumption in the model is that all continuous variables 231 Z_i (such as annual precipitation (Pa)) can have either a linear or a non-linear 232 relationship with $\Delta Qf\%$. This means that a smooth function s() can be applied 233 to the Z_i variables. For the smoothing function we applied thin plate regression 234 splines with an additional shrinkage penalty. The result of this approach is 235 that for high enough smoothing parameters (i.e. if the data is very "wiggly") 236 the smooth term can be shrunk to 0 and thus will be no longer significant [38]. This is done because a highly flexible smooth term could always fit the data, but would not necessarily indicate a relevant relationship. In other words, the approach balances finding a smooth non-linear relationship for the variable against overfitting the data.

The changes in forest cover contain both positive (forestation) and negative values (deforestation). In Zhang et al. [41], these changes were jointly analyzed, assuming the effect on the change in flow was linear and the effect of removing forest cover was the same as an equivalent addition of forest cover.

However, the impact of an increase in forest cover can be different from the same fractional decrease in forest cover. The question becomes how best to 247 analyze this. One approach would be to allow a different slope and a different intercept for the decreases relative to the increases. This can be tested by con-249 verting all the change in forest cover data to positive values, and an additional binary column $(sign_{forestcover})$ can be included indicating whether it was a for-251 est cover increase or decrease. In the model, the parameter for $sign_{forestcover}$ 252 will indicate the difference in the changes in flow for increases in forest cover 253 compared to decreases in forest cover. The disadvantage of this approach is that 254 the relationship with forest cover becomes discontinuous at the origin (0 change 255 in forest cover). 256

A second approach is to test the change in forest cover as a non-linear relationship in the GAM model. Because a shrinkage penalty is used, this will
also test the non-linear assumption and allows the variable for forest cover to be
continuous. The disadvantage of this approach is that the relationship between
forest cover and change in flow is less easy to interpret, as the non-linear fit in
the GAM has no direct parametric form. All three approaches are tested in this
study.

The overarching test focuses on identifying the change streamflow as a result

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of a change in forest cover and how this is potentially affected by different other factors (as indicated by the previous research: Zhang et al. [41]; Filoso et al. [16]; Zhou et al. [44]): climate, size of catchment and length of study. In addition to these earlier identified factors, this study also tested for the factors listed in Table 1

As an initial approach we tested whether the additional catchments added to the original data from Zhang et al. [41] did not majorly influenced the results (This analysis is in supplementary material part 2). This analysis highlights that the newly added catchment and the changes to the dataset create minor differences when repeating the analysis from the original paper. However, this means that the results of the studies are still comparable.

To make all the data and code used for the analysis publicly available, all the final data and analysis for this paper are located on github:

https://github.com/WillemVervoort/Forest_and_water on the "publish" branch.

279 3. Results

3.1. Description of the data

The overall dataset contains 329 observations of changes in flow, which in-281 cludes the newly identified data sets and after removing identified duplicate data 282 and lines with missing data. In contrast, the original dataset from Zhang et al. 283 [41] contained 312 catchments and the Filoso et al. [16] study used 37 catch-284 ments (Table S2 in Filoso et al. [16]). The overall distribution of changes in flow 285 is highly skewed as is the distribution of changes in forest cover and $Area \ km^2$. 286 The values of changes in flow greater than 100% and smaller than -100% clearly 287 create long tails in the change in flow distribution. Note also the large number 288 of studies with 100% forest cover reduction. Clearly visible is also that smaller catchments dominate the database with 42% of the data from catchments < 1290

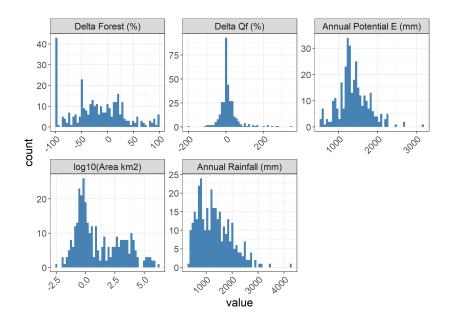


Figure 1: Overview of the distribution of the data set for five of the included variables. Note that the first panel (showing the distribution of the catchment areas) indicates the distribution of the $log_10_$ transformed Area (in km²).

 291 km² and 65% of the data for catchments < 10 km² (Figure 1). This high skew 292 in some of the data can create difficulties in the statistical modelling and this 293 will be discussed later.

3.1.1. Geospatial location of the catchments

Apart from looking at the distribution of the values, the spatial locations of the data can also be important, in particular when analysing the effect of climate. The catchments are spread across the world, and relative to Zhang et al. [41], this dataset has a very similar geospatial distribution. The major climate gradients are represented in the data, but there appears to be some bias in the spatial locations of the data. As the global map (Figure 2) shows, the distribution of case study catchments covers multiple continents. There is some spatial clustering in the studies in North America, Australia and East Asia.

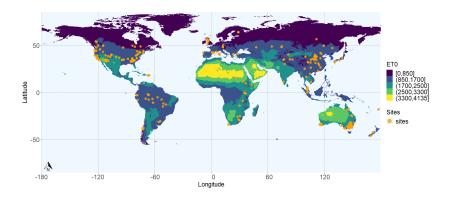


Figure 2: Distribution of included catchments across the globe based on reported or estimated latitude and longitude

3.1.2. Cross correlation between the different variables

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A final data exploration is to identify potential cross correlations in the data, which can point to possible interactions or potential biases. This analysis can also provide further insight for the statistical modelling, highlighting potential latent variables in the data set.

The correlation plot (Figure 3) highlights several correlations thate are worth investigating, even though in general cross correlations between variables are quite low. Some interesting relationships that appear in this graph are:

- the negative relationship between log10(Area) and change in forest area (DeltaF_perc), indicating that in the data set larger catchments tended to have (obviously) smaller areas of forest change.
- the weak positive relationship between log10(Area) and the assessment method using hydrological models. This highlights that paired catchment studies mostly concentrate on smaller scales.
 - A strong inverse relationship between log10(Area) and the paired watershed assessment method (simply the inverse from the last point), which

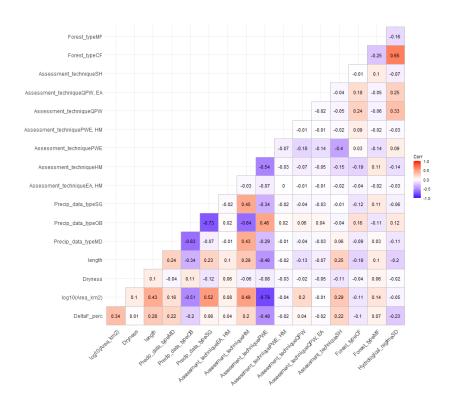


Figure 3: Correlation matrix for all variables

is also indicated by the negative relationship between the two assessment methods. This is further visible in the relationship between the change in forest cover and the paired watershed assessment method, showing the impact of the latent variable (log10(Area)). Smaller catchments used in paired watershed assessments are easier to fully clear or fully replant.

3.2. Statistical analysis

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The results of the overall statistical model that includes all the variables (but no interactions) reinforces some of the results from the correlation analysis.

This includes introducing non-linearity (Equation (2)) for the numerical variables in the model. While increasing non-linearity in the model can increase the flexibility if the model, the shrinkage splines assist with limiting overfitting. Following Wood [38], the number of degrees of freedom k in the non-linear variables was based on assessment of the effective degrees of freedom in the model output. If the effective degrees of freedom were close to k-1 then k was increased and the model rerun. By using shrinkage splines, this also results in the whole term being shrunk to zero if needed [38].

Table 2: Statistical summary for the linear terms the full model

	Estimate	Std. Error	t value	$\Pr(> t)$
(Intercept)	-4.81	16.31	-0.29	0.77
${\bf DeltaF_perc}$	-0.6	0.06	-10.71	0
${\bf Precip_data_typeOB}$	-21.4	13.23	-1.62	0.11
${\bf Precip_data_typeSG}$	9.36	15.17	0.62	0.54
${\bf Assessment_techniqueEA},$	20.64	42.73	0.48	0.63
$_{ m HM}$				
${\bf Assessment_techniqueHM}$	22.81	11.71	1.95	0.05

	Estimate	Std. Error	t value	$\Pr(> t)$
Assessment_techniquePW	E 30.63	11.94	2.57	0.01
${\bf Assessment_techniquePW}$	$(\mathbf{E}, 17.42)$	43.26	0.4	0.69
$_{ m HM}$				
Assessment_techniqueQP	W 39.52	20.15	1.96	0.05
Assessment_techniqueQP	W , 24.39	24.42	1	0.32
$\mathbf{E}\mathbf{A}$				
${\bf Assessment_techniqueSH}$	45.3	11.83	3.83	0
${\bf Forest_typeCF}$	-9.45	7.6	-1.24	0.21
${\bf Forest_typeMF}$	-8.05	7.56	-1.06	0.29
${\bf Hydrological_regimeSD}$	3.57	9.16	0.39	0.7

Table 3: Statistical summary for the smooth terms for the full model

	edf	Ref.df	F	p-value
s(log10(Area_km2))	0.79	4	0.99	0.02
s(Dryness)	4.64	9	2.26	0
s(Length)	4.45	34	0.22	0.12

The overall explaining power of the model can be interpreted from the adjusted r^2 (which is penalised for the number of parameters). This indicates an adjusted r^2 of 0.45 and deviance explained is 0.49, suggesting the model only explains about 50% of the variance in the data.

Inspecting the significance of the variables (Table 2 and Table 3) indicates
some interesting features. The overall partial slope of the change in forest cover
is -0.6, if all other variables are kept constant. This suggest quite strong change

in streamflow, moving from fully forested to fully cleared. Over the whole forest cover range, this is a change of -120 mm, with other variables held constant.

This change is highly significant, as indicated by the low p-value.

In addition, all the smoothed variables $log10(Area~(km^2))$ (p = 0.02)), Dry-log16 ness (p = 0)) and log16 (p = 0.12)) explain variation in the data. For log16 the p-value is not strictly smaller than 0.05, but still indicates some reasonable evidence that the variable explains some of the variation in the change in
streamflow.

Furthermore Table 2 indicates that several of the assessment methods explain variation in the change in streamflow, which was also indicated in the correlation analysis. In particular, the assessment methods Paired Watersheds Experiments (PWE), Hydrological Modelling (HM) and Statistical modelling and hydrographs (SH) are important explaining variables (p < 0.05).

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The remaining variables related to rainfall observation technique, forest type, or hydrological regime don't appear to have an influence on the change in flow.

Table 4: Comparison of alternative models for the relationship between the change in forest cover and the change in streamflow. (See Supplementary Material part 3)

Model for change in forest cover	Deviation explained	AIC
linear across range	0.49	3182
different for forestation and	0.45	3227
deforestation		
non-linear across the range	0.5	3182

As discussed in the methods, the overall linear response to the change in forest cover was compared to a transformation of the negative forest cover to positives and a check whether the relationship might be non-linear. This approach tests whether the impact on streamflow from removing forest cover is different from reforestation, as outlined in the methods. The detail of the comparison is highlighted in Supplementary material part 3. However, generally the results of the analysis showed two main points (Table 4):

1. The model assuming a simple linear relationship between change in forest cover (both positive and negative) and the change in flow explained the most variation in the data and indicated the best performance in terms of the Aikaike Information Criterium (AIC); and

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2. There is no need to assume a non-linear relationship, as a linear relationship provides a similar performance for the fit to the data.

The smoothed variables in the model can be inspected visually to identify if there are any issues with the fit. This is in addition to the earlier mentioned checks using gam.check() in the R package mgcv to test whether the number of degrees of freedom k is adequate.

Figure 4 highlights that the relationship between $log10(Area~km^2)$ and the change in flow is essentially linear. It indicates the negative slope that was also clear from Zhang et al. [41], indicating that in larger catchments changes in forest cover have less impact on streamflow than for smaller catchments.

Both the *Length* and *Dryness* variables show strong non-linearity, but the relationships do not show a clear trend due to the scatter and the distribution of the data. A further problem appears to be that *Length* and *Dryness* have several points with very high leverage that determine much of the non-linearity in the relationship.

As this is not always shown in papers discussing regression relationship, the residual distribution is provided in more detail (Figure 5). Visually, the residuals appear approximately normal, although there is a noticable skew in a limited

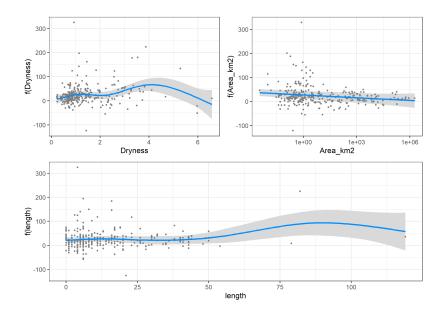


Figure 4: Visualisation of the smooth variables in the model, the shaded areas are the 95% confidence intervals associated with the fit of the smooth, the blue line is the mean smoothed relationship, with data plotted as individual points

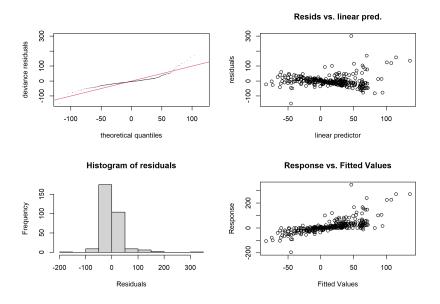


Figure 5: Residual plots for the regression model indicating a slightly fat-tailed residual distribution

number of the data in the upper part of the distribution (Figure 5, top left).

This is related to a limited number catchments that have very high changes in

streamflow in the data set. In other words, the distribution of the residuals is

somewhat fat-tailed.

One solution could be to transform the data, however this is not that simple.

As the data for the change in flow cover the domain \mathbb{R} , a simple log or Gamma

transformation is not a solution. More complex transformations make the re-

393 sults of the regression difficult to interpret, and at some point can be slightly

contrived.

Given the majority of the residuals indicate a relatively well behaved distribution, we simply note the behaviour at the extremes and will discuss this later in the paper, and explain how this relates to the characteristics of the dataset.

3.2.1. Removal of studies of great length and for very dry catchments

Table 5: Catchments for which the dryness index > 5

Number	Latitude	Longitude	Catchment name
76	34.67	-111.7	Beaver Creek, AZ #3-2
225	32.74	-111.5	Natural Drainages, Ariz.,
			U.S.A, A
226	32.74	-111.5	Natural Drainages, Ariz.,
			U.S.A, C
356	-25.75	28.23	Queens river

The flexible nature of the splines means that the Length variable highlights

substantial non-linearity in the data, but it is unclear what exactly is captured.

401 The shape of the conditional response (Figure 4) does not reflect a similar

response as indicated by Filoso et al. [16] and Jackson et al. [19]. One reason

could be that the relationship is dominated by the few data points with very long data series, which show highly variable responses (Figure 4).

The points related to catchments with very long studies (> 60 years) might
be questionable, as changes other than forest cover change could affect streamflow. In addition, a few of the catchments have Dryness values that are very
large (> 5) and these values have high leverage in the data, affecting the residual
distribution. These catchments are listed in Table 5, and are three catchments
in Arizona and 1 catchment in South Africa. It is possible that catchments in
these climate zones behave different from the rest of the catchments.

Table 6: Statistical summary for the linear terms the restricted model

	Estimate	Std. Error	t value	$\Pr(> t)$
(Intercept)	-10.17	17.98	-0.57	0.57
${\bf DeltaF_perc}$	-0.59	0.08	-7.45	0
Forest_SignIncrease	0.41	9.79	0.04	0.97
Precip_data_typeOB	-15.97	12.5	-1.28	0.2
${\bf Precip_data_typeSG}$	15.71	14.85	1.06	0.29
$Assessment_technique EA,$	20.38	41.03	0.5	0.62
$\mathbf{H}\mathbf{M}$				
${\bf Assessment_techniqueHM}$	26.42	11.4	2.32	0.02
Assessment_techniquePWI	E 28.51	12.15	2.35	0.02
Assessment_techniquePWI	E, 17.4	42.05	0.41	0.68
$\mathbf{H}\mathbf{M}$				
Assessment_techniqueQPV	V 41.49	19.53	2.12	0.03
Assessment_techniqueQPV	V , 24.81	23.32	1.06	0.29
TrΛ				

 $\mathbf{E}\mathbf{A}$

	Estimate	Std. Error	t value	$\Pr(> t)$
Assessment_techniqueSH	47.26	11.49	4.11	0
${\bf Forest_typeCF}$	-9.47	7.3	-1.3	0.2
${\bf Forest_typeMF}$	-6.01	7.35	-0.82	0.41
${\bf Hydrological_regimeSD}$	2.5	8.89	0.28	0.78

Table 7: Statistical summary of the smooth terms reducing dataset to studies with the study length shorter than 60 years and Dryness ≤ 5 .

	edf	Ref.df	F	p-value
s(Dryness)	4.02	9	2.16	0
$s(\log 10(Area_km2))$	0.87	4	1.53	0.01
s(Length)	0	9	0	0.98

Therefore it is worth investigating what effect removing these few data points has on the overall model and the significance of the variables. Data that have Dryness <= 5 and Length <= 60 years were removed from the dataset and the model based on a reduction of the data set from 329 to 310 catchments is run again.

This model, which excludes data with long studies and very dry catchments explains only slightly less of the variation with an adjusted r^2 of 0.44 and a deviance explained of 0.48.

Investigating the non-linear responses suggest that Dryness has a clear non-linear response, which is significant, where changes in forest cover in drier catchments having a greater impact on streamflow (Figure 6 and Table 7). Catchment area $(log10(Area~(km^2)))$ still has an impact on flow with p = 0.01, and the rela-

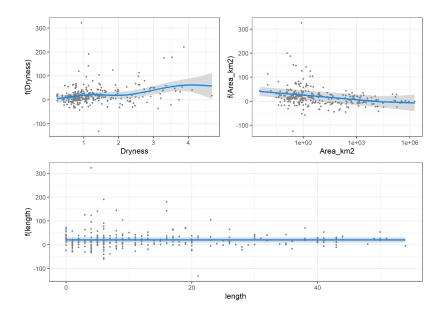


Figure 6: Visualisation of the smooth variables in the model with reduced data for dryness and length

- $_{424}$ tionship looks almost linear. More importantly, the variable Length is no longer
- significant, after removal of the two studies with very long lengths.

Table 8: Distribution of assessment techniques in the data set

Assessment_technique	n
PWE	185
$_{ m HM}$	57
SH	42
EA	32
QPW	7
QPW, EA	4
EA, HM	1
PWE, HM	1

One concern with the results presented so far is that there are a few assessment techniques in the data set with a very low number of observations and could influence the results of the analysis. This includes the category of Quasi paired watersheds and combinations of elasticity analysis and hydrological modelling (EA,HM) and paired watersheds and hydrological modelling (PWE,HM) (Table 8).

Therefore, the model was rerun excluding the combined assessment techniques (EA, HM), (PWE, HM) and (QPW, EA) and the assessment technique
QPW, which were all non-significant (Table 8). This resulted in a data set of
323 catchment studies.

The model based on assessment techniques that have more than 10 observations in the data set does not change much in the results (results not shown). It strengthens the significance of the different assessment techniques, but generally results in the same interpretation. Overall this suggests that although those observations have some impact on the overall relationships, they do not strongly bias the outcomes.

The overall model results clearly highlight that some of the assessment techniques (in particular paired watershed studies (PWE) and combined use of statistical methods and hydrographs (SH)), have a strong impact on the predicted
change in flow. Particularly, relative to EA (elasticity approaches) all other
assessment techniques have higher predicted changes in flow. In other words,
there is a distinct difference in the way the change in flow is assessed, and the
EA method (for example in Zhou et al. [44]) appears to suggest a much smaller
effect on the change in flow.

50 4. Discussion

- The generalised additive models appear to reach the same conclusions as the single variable regression in earlier papers [41, 16]. It appears that:
- 1. Larger catchments show lower impact of forest cover change on streamflow;
- 2. Drier catchments show a greater impact of forest cover change on streamflow; and
- 3. There is a general linear relationship between the change in forest cover and the change in streamflow.
- This might suggest that the simpler models have reached the correct conclusion. However, this is somewhat premature. given that the other major point coming out of the results is:
- 461 4. There is a clear relationship between size of catchments, area cleared and
 462 type of experiments, with particular Paired Watershed Experiments con463 taining the smallest catchments, the largest % forest cover change and the
 464 largest variability in the flow response.
- Figure 7 provides a clear overview of the whole data set, and in this figure
 the size of the catchments and the different assessment methods are highlighted.
 This figure clearly indicates that the data relating to high changes in forest cover
 are all small catchments and relate mostly to paired watershed experiments.
 In contrast, data related to large catchments are related to smaller changes in
 forest cover and different methods, such as hydrological modelling and elasticity
 analysis. This confirms the model results (Table 6) and the earlier correlation
 analysis (Figure 3).
- It is possible that one of the reasons why Zhang et al. [41] separated their analysis in large ($> 1000 \text{ km}^2$) and small ($< 1000 \text{ km}^2$) catchments, is that they realised this difference in assessment methods and wanted to account for

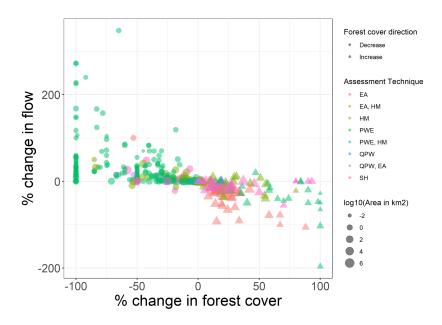


Figure 7: Overview of the data highlighting the dominance of small catchment studies which are fully forested or cleared and the scatter in the data

this. However, this is not explicitly identified, and there is no real physical explanation of the $1000~\rm{km}^2$ threshold.

The other interesting point in Figure 7 is that the variation in the data increases as the catchment size decreases and the change in forest cover increases.

This also means that the overall variation in the data for paired watershed experiments (PWE) is much greater than for any of the other methods.

4.1. Is there a problem with extending local experimental data to larger scales?

The overarching reason for combining past studies at a global scale is to infer relationships that can be used to make more general statements or develop more global scale modelling of impacts [i.e. 44, 19, 18]. Therefore, the results from the analysis could be seen as a confirmation of the earlier research [41, 16, 44, 19]. However, the explaining power of the developed model is quite low and a lot of variation in the data is unexplained. As is highlighted in the introduction there are four major issues with this type of analysis, and the results from this paper also highlight these issues. Here, these issues are further explained.

491 4.1.1. Issue 1: Latent variables are not included in the typical single covariate
492 analysis

The results show that it is simply impossible to analyze a single covariate 493 relationship, as there are several latent variables in the data. An example of this is the general relationship of the change in flow as a function of the change 495 in forest cover. Clearly the relationship is highly impacted by the fact that all 496 the small catchments have large changes in forest cover and are all associated 497 with paired watershed experiments. Without taking these factors into account, 498 a definite answer about the impact of forest cover on the change in flow cannot 499 be given. Furthermore, the large variability in the change in flow data for these 500 small catchments (Figure 7) indicates that there is a further (unknown) variable that explains the variation in the data. 502

If the remaining variation in the residuals is small relative to the trend, then there is little need to identify further latent variables, but if the variation is large, then it is unclear if it is the latent variable that determines the trend, or the actual relationship in the data.

Similarly, the data for the larger catchments containing smaller changes in 507 forest cover are dominated by hydrological modelling studies, resulting in a 508 further complication. If the reponse of the streamflow in the modelling studies is 509 the result of the conceptualised relationship between streamflow and forest cover 510 (possibly from a subset of the paired catchment studies), then it is impossible 511 to say if the change in streamflow is real, or simply a result of a pre-conceived 512 model relationship. Is the smaller variation in the data for smaller changed in 513 forest cover (Figure 7) a result of similar conceptualised model relationships, or 514 actual variation between catchments and climate types? Currently this question 515

516 cannot be answered.

This becomes problematic when extrapolated to larger scales. A clear example of this is the paper by Hoek van Dijke et al. [18] where the conceptualised
relationship between forest cover and streamflow pre-determines the outcomes
of the global modelling.

The only way to analyze changes in streamflow as a function of forest cover in larger catchments is to actually derive this from observed data of long term streamflow and forest cover (as was done in Levy et al. [22]).

One of these latent variables could be the total area of forest in a catchment, 524 as was analysed in Levy et al. [22]. In this case, the total % area of forest was 525 not included in the data. As a test, the total % area of forest for the larger catchments ($> 1000 \text{ km}^2$ in Zhang et al. [41]) were added to the dataset and the 527 model for just the large catchments was tested. This showed that the % area of forest was not significant explaining the change in flow for the larger catchments 529 (retaining all other variables in the model, results not shown). While this might 530 be an area of further research on the full dataset, it is complicated for two 531 reasons: 532

- 1. The area of forest is not always indicated in the original papers, or a range of values is given, complicating the data collection.
- 2. Many of the small catchments have 100% area covered in forest, introducing a strong skew in the data and complicating if total area of forest has an impact on the change in flow.

We are not arguing that there is no relationship between streamflow and forest cover, and there might indeed be a global relationship that can be discovered. But, this relationship can only be discovered if we are able to address some of the major other factors that explain the variability, and work with actual data and not model outputs. 4.1.2. Issue 2: Interpretation errors due to complex descriptions of the experiments in the original papers

The second major issue that became clear from reviewing many of the original papers is that some of the variability might be an interpretation problem. In many cases the original description in the paper is interpreted to extract the % change in streamflow from the % change in forest cover. This seems like a simple activity, but this is not always the case.

Two examples can be highlighted:

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- The papers from Almeida et al. [1] and Ferreto et al. [15] partly discuss 551 the same experiment and the same catchment. In Almeida et al. [1], 552 the methods discuss how two experimental catchments of approximately 80ha in size which were harvested. One catchment was 100% harvested 554 and the other 30% harvested. Throughout the paper the catchments are indicated as 100% harvested and 30% harvested. However, only after 556 reading Ferreto et al. [15], did we discover that in fact the 100% and 557 30% refer to the "eucalyptus plantation area", which was about 60% of 558 the total area. This is in fact mentioned in Table 1 in Almeida et al. 559 [1], but does not appear in the text. The question then becomes how to 560 interpret this in the data base for this paper. Clearly it was a 100% and 561 30% change in forest cover, but only for the 60% plantation cover, not for any of the other areas in the catchment, which included native vegetation 563 and riparian vegetation. There are several other examples like this in the 564 different papers [for example 6, 5]. 565
 - Another example is the paper by Waterloo et al. [36]. This modelling study
 in Fiji of the clearing of a catchment reports the changes in streamflow
 over parts of the year. For a period of 324 days the streamflow increased
 from 252 mm to 580 mm (a 230% increase if calculated as 580/252 * 100)

and for a second period of 309 days the streamflow increased from 90 mm to 194 mm (a 215 % increase). However, how we convert this to a change in annual flow (which most of the other data relate to) is difficult. The original data base listed a 50 % change in flow, but it is difficult to identify how this is calculated. We suspect that results from $252/580*100 \approx 50$ and $90/194 \approx 50$.

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Clearly, interpreting older papers can be difficult and this can result in variation in the data that is being analyzed. Similar to the last issue, if these errors only introduce small variation in the data, then it will not limit the interpolation to larger scales. At this point, it is not clear if this is indeed the case. The large variation in the experimental watershed data suggests that this might be a more serious problem.

4.1.3. Issue 3: Aggregation of data that originates from different experiments
 with different objectives across a wide time period

For many of the small catchment studies listed in the database, the assumption is that the original experimental design can be interpreted in terms of a binary "forestation" or "deforestation". However, the real situation is often much more complex and fuzzy.

Many of the paired watershed experiments included a harvesting and replanting or regrowth after harvesting or fire experiment [e.g. 11, 12, 37]. As a result, it becomes difficult to assess how we interpret the change in flow as a result of a change in cover. In many cases we would expect the flow to change over time as a function of the recovery [20] and therefore the timeseries of the flow needs to be assessed over a longer time.

Many of the papers in the database report early results (for example 1 or 3 years after harvesting), but some also report longer time periods. As earlier work [12, 20] has highlighted, we can always expect a larger effect directly

after harvesting, but this effect diminishes over time (even if it does not always return to the original state). Comparing studies reporting results directly after treatment to longer term studies therefore becomes problematic.

In our work, the variable *Length* was used in the model to test for some of these effects, but this was insignificant in the model (Table 7). Given the other variation in the data, this does not necessarily mean that there is no effect.

This is further complicated by the variation in different types of clearing and the different types of vegetation. In the original Zhang et al. [41] a variable to describe the *forest type* was included (Table 1), but in the model this is not significant (Table 2). This is probably because the broad classification used does not capture the actual variation in runoff response. In addition, as Figure 3 shows, there is a correlation between coniferous forests and snow dominated hydrological regimes, further complicating the analysis.

An additional complication related to combining studies related to wild fires or bush fires and logging studies is the differences in vegetation recovery. For example, Heath et al. [17] found that catchments with resprouting species around Sydney, Australia, indicated little change in the streamflow in comparison to species regrowing from seed further south on the continent [45].

As a result, it can be difficult to exactly pinpoint the change in flow as a result of the change in cover, as well as being difficult to assess what the exact change in cover actually was.

As indicated before, if the overall variation due to this issues is small, then
this would not be an issue for upscaling the results, but the large variation for
the smaller catchments suggest that effects could be considerable. As Jones et al.
[20] indicate, this really needs time series analysis of the different experiments.
However, some of the time series data might not be recoverable from the older
experiments, which will limit the opportunities for analysis. We will discuss this

further below.

625 4.1.4. Issue 4: Transcription errors in the data

This issue seems to mainly occur if data is collected from other review papers.

This might be because some of the original papers are difficult to locate and
therefore values from reporting papers are used. In supplementary data part 1,
several changes to the original data sets have been documented, and as can be
seen several of these are transcription errors.

This does influence the results in Zhang et al. [41], comparing the results in
Supplementary material 2 with the original paper. The main example is that in
this study the largest catchment (watershed #1 in Zhang et al. [41]) had to be
removed, as this study actually involved paired watershed experiments on very
small plots, for which the characteristics were not recoverable.

Clearly, this is a problem for all reviews that attempt to bring together large numbers of results from published papers, and where actual results are copied rather than using some sort of automated text analysis.

In the end, careful review of the data and the original papers can circumvent most of this issue. And, making the data available (as Zhang et al. [41], Zhou et al. [44] and Filoso et al. [16] have done) provides an opportunity for review by other researchers, and over time most of the transcription errors can be resolved.

4.2. General discussion

In this paper, a few studies have been singled out in the analysis. The choice
of focus was mainly driven by the data that was made available by the authors
of these papers [41, 16], which provide a rich case study for the current paper.
Field research is by nature limited in space and time, due to the high costs
involved of setting up experiments. This is particularly true for experiments in

hydrology and forest hydrology, where field sites need to cover sufficient spatial

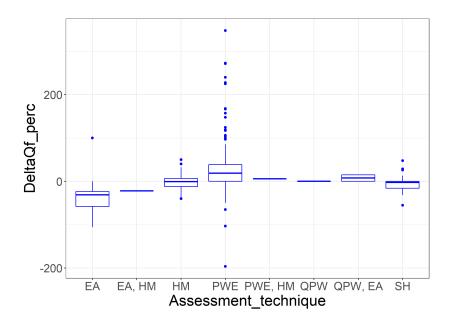


Figure 8: Boxplot of the variation in the change in flow for the different assessment techniques, showing the dominance of the variation and the outliers in the dataset in the paired watershed experiments

and temporal variability. This means there is a general need to extrapolate the local results to larger scales to inform decision making and policy.

However, as is demonstrated in this paper, there are multiple issues when
this local scale data is extrapolated to larger scales. It clearly demonstrates that
the results of any model (in this case a regression model) is highly dependent
on the data, but also on the assumptions in the model. From the perspective of
extrapolating local data to global scales for policy advice and decision making
[i.e. 18, 19], this is an important point.

4.2.1. Residuals of the model

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The residuals of the final model presented in this paper (Figure 5) indicate that the residual distribution remains fat-tailed, causing deviations from an assumed $\epsilon \sim N(0, \sigma^2)$. This once again highlights that there is unexplained variation at the extremes of the distribution, once again related to the paired

watershed experiments (Figure 8). Generally, in statistical models, the approach would be to further normalise the residuals through transformations. However, in this case this might be difficult and might not resolve all the issues due to the large variation in the data.

4.2.2. Interactions

The current modelling approach does not consider any interactions between
the variables, and this would offer another approach to understand the variation
in the data. As already indicated in Figure 3, there are interactions between
different variables. This further complicates the extrapolation of the local scale
experiment data to global scales and to extend historical data to current management and decisions.

In this case, interactions were not included because, as was shown, there are bigger problems with trying to extrapolate the existing data, and the data itself can be problematic. To be able to model the interactions well, the nature of the variables and interactions need to be understood and or clearly hypothesized.

Otherwise it becomes another case of correlation without causation.

4.2.3. Implications for other "meta-analysis" studies

There has been a recent push to develop more meta-analysis studies in 680 hydrology [35, 14], and we strongly believe that developing new insights by 681 combining historical data sets from reviewed papers is highly valuable. How-682 ever, this paper highlights that there is considerable chance that large histor-683 ical data sets include latent variables and are more complex than envisioned. 684 This is particularly true for more historical work, as methods of observation and 685 even apporaches to management have changed considerably. The same management description is not necessarily the same action on the ground. A carefully 687 designed and systematic approach can prevent some of bigger problems as is demonstrated in Wang et al. [35], where both the approach and the catchment area are investigated as latent variables. This is particularly relevant, where the results of meta-analyses are extrapolated to make global predictions without clearly quantified uncertainties (such as in Hoek van Dijke et al. [18] and Wang et al. [35]).

A second potential danger is the extrapolation of the local small catchment 694 results and conclusions to larger scales, but beyond the original scope of the 695 studies. For example, the current database is mainly related to forest harvest, 696 bush fire and reforestation/plantation management. It is tempting to use the result of a large scale analysis of this data to make inferences about overall 698 landuse change [23, 35], but this would not be valid, as the deforestation studies are generally not a transition to an agricultural landuse or pasture, but 700 regrowing into forest. Similarly, using the plantation studies to extrapolate to "reforestation" (as in Filoso et al. [16] and Hoek van Dijke et al. [18]) is also 702 tenuous. Plantation forests are generally fast growing hybrids that will have 703 quite different ecophysiology, particularly in South America [20, 4], while other 704 reforestation, for example for salinity control in Australia, might focus on a mix 705 of native species. Given the link between ecophysiology and water and carbon 706 budgets [19], care should be taken in extrapolation, introducing a further error. 707

A final factor is ignoring the effect of climate change [34] on runoff, even if the effects are still minor. Earlier papers [23, 35] have analyzed climate effects relative to management effects in the data, but these studies did not explicitly test for climate change. Given that the database of studies now captures almost 100 years of work, we cannot ignore a climate change trend that is potentially hidden in the data. A simple inclusion of the start date of the experiment (From) in the GAM model does suggests an increase in change in the percentage of flow over time. However, as the data distribution is uneven in time, and consists

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of multiple assessment techniques there could be multiple complicating factors, and drawing a firm conclusion would be premature.

718 4.2.4. Future research needs (implications for forest hydrology)

Beyond a more formal approach to investigating climate change effects in the
data, this study also points to several further opportunities and future research
needs.

A major focus of many of the papers related to forest hydrology has been 722 on the impact of plantation forest operations on the catchment, rather than 723 the transition of forestry to agriculture. As the paper by Jones et al. [20] 724 highlights this means there are opportunities to analyze the time evolution of 725 the catchment response to forestation. Given the large number of studies that 726 look at a time evolution of forest cover (i.e. either clearing and regrowth, or 727 burning and regrowth), this data can offer further insights into the dynamic response of catchments to changes in land cover. As highlighted, some of the 729 older data is not fully recoverable, but there is often a series of papers related to one experiment, which at least would provide individual time points. 731

More generally there is a clear need for a more in depth analysis of the data base of studies used here. In particular, more detailed data can potentially be extracted from many of the studies in terms of vegetation species, streamflow responses and responses of components of streamflow (slow flow, quick flow etc.), as well as a more in depth description of the management and actual experimental design.

There is also a clear need to understand the impact of the assessment methods with respect to scale. Extrapolating paired watershed experiment results
into models can possibly overlook landscape interactions that are visible at
larger scales, but do not occur on smaller scales. For example, this could be the
effects of lateral flow and groundwater connectivity and impacts of elevation on

landuse. A carefully designed simulation study that specifically investigates the change in stream flow repsonse with scale using local field data for verification can help solve this problem.

At the moment, providing answers to the impact of streamflow at larger scales should generally not be approached by simulation modelling. A better 747 approach is analyzing streamflow data at multiple spatial and temporal scales 748 for responses (rather than running simulations) and using satellite data to dy-749 namically include landuse changes. The highlighted paper by Levy et al. [22] is 750 currently the best example of a solid statistical approach to analyzing streamflow responses. Simulation modelling can be an approach to analyze different 752 scenarios, if there is clear recognition of the potential impact of the model structure (the algorithms and parameters that describe for example plantation tree 754 growth) on the simulation outcomes.

We envision that in the future more innovative approaches to analyzing data at different scales will be developed.

758 5. Conclusions

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This study demonstrates that analyzes of large databases of essentially "aggregated data" should be considered carefully and simple single variable regressions often present simplistic relationships that can be misleading.

While the analysis reveals similar conclusions in relation to the response of streamflow to forest cover, there are four major interlinked reasons why these results should be considered carefully. This subsequently has implications for meta-analyses in Environmental Science and Hydrology in general. The reasons higlighted in this paper are:

• The existence of latent variables in the data that create the appearance of a relationship that really does not exist;

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• The difficulty in fully interpreting the specifics of different studies;

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• The difficulty of integrating data from seemingly similar studies, but with
quite different objectives; and

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• The chance of transcription errors influencing the data.

Any statistical analysis, including the one in this paper, needs to be considered "conditional on the data", and given the issues indicated, extrapolation of the results of summary studies to larger scales and into global hydrological models has to be done with great care. Better would be to analyze observed data and explicitly include uncertainty in the extrapolation of the results.

This therefore has implications for the recent growth in meta-analysis review papers, which has been boosted by increased computational capacity and much better on-line accessible data bases with research data. Clearly, this requires careful definition of the search terms, and follow-up review of the harvested papers, as well as an understanding that the statistical relationships can be hiding other unknown factors. As the old adagium says: Correlation is not the same as causation.

788 6. Acknowledgements

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