# Parallel Strategies for the GPS Radio Occultation Data

## Assimilation with a Nonlocal Operator in the Weather Research

## and Forecasting model

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#### ABSTRACT

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Several strategies are designed and implemented in the data assimilation system for the Weather Research and Forecasting model to parallel assimilate the global positioning system 11 (GPS) radio occultation (RO) sounding with the nonlocal excess phase delay operator, which 12 is computational expensive and has been proven to produce significantly better analysis for 13 numerical weather predictions compared to local refractivity operator. In particular, to solve 14 the parallel load imbalance problem due to the uneven geographic distribution of the GPS RO 15 observations, the round-robin scheduling is adopted to distribute GPS RO observations to 16 balance the workload among the processing cores. The wallclock time required to complete 5 17 iterations of minimization on a demonstration Antarctic case with 106 GPS RO observations, is reduced from more than 3.5 hours with single processor to 2.5 minutes with 106 processing cores. These strategies present the possibility of the application of the nonlocal GPS excess phase delay operator in the operational data assimilation systems with a cut-off time limit.

### $_{2}$ 1. Introduction

The observations from the global positioning system (GPS) radio occultation (RO) limb 23 sounding technique has been proven as a valuable observation of atmosphere for numerical 24 weather prediction (NWP) and climate research (Kuo et al. (1997); Zou et al. (1999); Zou 25 et al. (2000); Liu and Zou (2003); Healy et al. (2005); Huang et al. (2005); Cucurull et al. 26 (2006); Cucurull and Derber (2008); Healy and Thepaut (2006)). GPS RO data has several 27 advantages such as a high vertical resolution, no need for calibration, unaffected by cloud 28 cover and rainfall, and global coverage. In particular, in the middle of upper troposphere, the 29 GPS RO measurements have accuracy comparable with or better than that of radiosondes 30 (Kuo et al. (2005)). Since the launch of the Constellation Observing System for Meteorology, 31 Ionosphere, and Climate (COSMIC) mission in 2006, approximately ~1500-2500 globally 32 distributed GPS RO sounding are provided per day in near-real time. The COSMIC GPS 33 RO sounding are currently being used at several global operational NWP centers, including the National Centers for Environmental Prediction (NCEP; Cucurull and Derber (2008)), the European Centre for Medium-Range Weather Forecasts (ECMWF; Healy (2008)), the 36 Met Office (UKMO), and Météo France (Poli et al. (2009)). 37 Due to the success of COSMIC, U.S. agencies and Taiwan have decided to move forward 38 with a follow-on RO mission (called FORMOSAT-7/COSMIC-2) that will launch six satel-39 lites into low-inclination orbits in late 2015, and another six satellites into high-inclination 40 orbits in early 2018. The COSMIC-2 mission will provide nearly an order of magnitude more 41 RO data increase in the number of atmospheric and ionospheric observations that will greatly 42 benefit the research and operational communities (http://www.cosmic.ucar.edu/cosmic2/). 43 Depending on the level of data processing, various variables can be retrieved from GPS 44 RO observations for use in data assimilation, such as bending angel, refractivity and retrieved moisture/temperature profile (See Kuo et al. (2000), Kuo et al. (2004)). To account for the variations of the atmospheric states along the GPS ray paths, the nonlocal excess phase operator, introduced by Sokolovskiy et al. (2005), has proven to significantly improve the assimilation of GPS RO data (Sokolovskiy et al. (2005), Liu et al. (2008), Chen et al. (2009),

Ma et al. (2009) and Shao et al. (2009)). However, due to the high computational cost

associated with the nonlocal operator, it has been tested only in some research configurations

with affordable number of observations. The parallel implementation of the GPS RO data

assimilation with nonlocal operator is urgently needed to advance its applications in both

research and operational data assimilation systems.

Because both the local refractivity and nonlocal excess phase delay operators had been implemented in the data assimilation system for Weather Research and Forecasting model (WRFDA, Barker et al. (2012), Chen et al. (2009)), the three-dimensional variational (3D-Var) approach in WRFDA will be used throughout this paper to demonstrate the parallelization of GPS nonlocal operator. We believe that the parallel strategies for nonlocal operator are general and applicable for other parallel data assimilation systems (such as four-dimensional variational, and Ensemble Kalman Filter) employing the domain-decomposition approach.

This article is organized as follows. In Section 2, we briefly introduce both the local refractivity operator and nonlocal excess phase delay operators implemented in WRFDA and their computational costs. Section 3 provides the technical details of how the nonlocal GPS RO operator is paralleled in a domain-decomposition parallel context. The strategy to solve the load imbalance problem is presented in Section 4. Sec. 5 presents some further optimization to save the cost. The summary and discussion are given in Section 6.

# <sup>69</sup> 2. Local and Nonlocal GPS RO Operators

In both local and nonlocal GPS RO operators, the neutral atmospheric refractivity can be calculated from model variables via the following relationship

$$N = 77.6 \frac{P}{T} + 3.73 \times 10^5 \frac{PQ}{T^2(0.622 + 0.378Q)} \tag{1}$$

where P is the total atmospheric pressure in hPa; T is the atmospheric temperature in K; and Q is the specific humidity in kg/kg. The local GPS RO refractivity operator assumes that the observed refractivity is modeled as the local refractivity at the perigee point where the GPS ray is closest to the earth. The first guess fields of P, T and Q are interpolated horizontally and vertically to the perigee point of the RO observation and the Eqs. (1) is used to calculate the local refractivity. The local GPS RO refractivity operator is simple and low computational cost, and it is used by most of the data assimilation systems.

To account for the variations of the atmospheric states along the GPS ray paths, the nonlocal excess phase operator, introduced by Sokolovskiy et al. (2005), simulates the excess phase by integrating the local refractivity along the ray path, which is approximated by a straight line. The new observation S (excess phase, which is to be assimilated) is defined as

$$S = \int_{ray} N \, \mathrm{d}l \tag{2}$$

where l is the ray path and N is the refractivity. There are two steps associated with the implementation of the nonlocal operator (Chen et al. (2009), Ma et al. (2009), and Liu et al. (2008)): Firstly, the observed excess phase is calculated by integrating the refractivity from RO observations:

$$S_{obs} = \int_{ray} N_{RO}(r) \, \mathrm{d}l \tag{3}$$

where r is the radius vector derived from  $r = r_c + z$ ,  $r_c$  is the local curvature radius of the earth, and z is the height above the earth's surface,  $N_{RO}$  is the observed refractivity interpolated on the model mean heights of the tangent point position by a vertical average. The next step is to calculate the model counterpart  $S_{mod}$ , which is given by

$$S_{mod} = \int_{ray} N_{mod}(r) \, \mathrm{d}l \tag{4}$$

where  $N_{mod}$  is the simulated refractivity from first guess fields of T, P, and Q by Eqs. (1) and interpolated at the tangent point position.

Compared to the local GPS refractivity operator, the computational cost of the nonlocal excess phase operator is increased dramatically due to the integration of refractivity

along the GPS ray. Liu et al. (2008) reports that the cost of nonlocal to local operator is at least 100 times on a Linux cluster of NCAR. To justify the necessity to parallel the GPS RO nonlocal operator, an Antarctic domain with 30km horizontal resolution, shown 97 as Fig. 1, is chosen to demonstrate the computational cost of the GPS RO operators. This is a Advanced Research WRF (Skamarock et al. (2008)) model domain of 1800 UTC 11 December 2007 with  $401 \times 401$  mesh size and 55 vertical layers between surface and 10 hPa. Fig. 1 also shows the locations of the 106 GPS RO profiles within  $\pm 3$  hours 101 window centered at 1800UTC. The 106 GPS RO profiles are assimilated with WRFDA 102 3D-Var on National Center for Atmospheric Research (NCAR)'s supercomputer yellowstone (http://www2.cisl.ucar.edu/resources/yellowstone) and the wallclock time of 5 iterations of 104 minimization are recorded to demonstrate the computational costs. After excluding the pro-105 gram initialization and I/O, with single processing core, it takes only 146s to assimilate 106 106 GPS RO profiles by local operator. However, about 12,762s ( $\approx 3.5$  hours) are needed to run 107 5 iterations with nonlocal operator. The ratio of the computational cost of nonlocal to local 108 operator is about 87 for this case. In terms of the production 3D-Var run with 60 iterations 109 (assumes 2 outer loops) approximately, the wallclock time for serially running this case will 110 be more than 42 hours on yellowstone. Therefore, the cost of nonlocal operator with single 111 processing core is extremely unaffordable for either research or operational purposes. 112

# 3. Parallelization Strategy

Most of the atmospheric models and their data assimilation systems employ the horizontal domain-decomposition method for parallel processing. In data assimilation systems, the
observation is usually assigned to the subdomain where it is geographically located. However, for such as satellite radiance data assimilation, due to the uneven distribution of the
observations and the expensive radiative transfer model, the radiance operator is expensive
and the load imbalance issue has to be considered for operational practice. We may either

change the way of the horizontal domain-decomposition method to have each subdomain cover similar number of radiance data, or redistribute the radiance data among processing cores to have each core has similar workload. Please note that the load balance algorithm itself might be costly and complicated.

The difficulty to parallel the nonlocal operator under the domain-decomposition context 124 roots in its nonlocal integral nature of Eq. (4) along the ray-paths at all vertical levels 125 above the tangent point and below the model top. The ray-path might intercepts several subdomains located on different processing cores respectively and each processing core is 127 only aware of the atmospheric states of the local subdomain. Apparently, the most suitable strategy to parallel nonlocal GPS RO operator is the ray-path-wise distribution among pro-129 cessing cores (Zhang et al. (2004)), which distributes the workload among processing cores 130 based upon the number of the GPS ray-paths, instead of the geographic location of the ob-131 servations. However, in our parallelization strategy, we must consider the fact of the existed 132 domain-decomposition method to minimize the implementation cost. 133

Although Eq. (4) is an integration of the refractivity along the ray-path, which might 134 go through the whole model domain, we noticed that only one derived variable – model 135 simulated refractivity  $(N_{mod})$  is used for the integration. If each processing core (subdomain) 136 is aware of the global  $N_{mod}$ , the integration is able to be done on each whole ray-path parallel. 137 The calculate of  $N_{mod}$  from firstguess fields with Eq. (1) is trivial and each processing core 138 can calculate the  $N_{mod}$  of subdomain locally in advance. It turns out that we can use the parallel collecting-and-broadcasting operation to collect the local calculated  $N_{mod}$  from each 140 subdomain to a global  $N_{mod}$  array and to broadcast the global  $N_{mod}$  to each processing core. 141 The costs for this strategy is the additional memory storage for several global arrays and 142 the collecting-and-broadcasting operation in each iteration. For modern distributed memory 143 supercomputers, the costs are trivial. 144

With the implementation of above parallel strategy (experiment "Parallel"), Fig. 2(a) shows the parallel wallclock timing results with up to 512 processing cores on NCAR's

yellowstone (red bars represent experiment "Parallel"). The wallclock time of 5 iterations minimization reduced from around 3.5 hours with serial run to 279 ( $\approx 4.5$  minutes) seconds with 512 processing cores.

## 4. Load Balance

Fig. 3 shows the parallel speedups, which are the ratio of the wallclock times of parallel 151 runs against that of the serial run. The black line is the linear parallel speedup and represents 152 the ideal speedup or acceleration when multiple processing cores are used. The red line 153 represents experiment "Parallel". The speedup with 512 processing cores is 46 and one 154 may argue that the actual speedup is much lower than the ideal speedup (512) and the 155 parallelization strategy is not cost-effective. Therefore, the analysis of the parallel algorithm 156 of the GPS RO operator will be helpful to understand the unsatisfactory speedup and identify 157 the reason behind. 158

3, we emphasized that we have to consider the existed horizontal domain-159 decomposition method for the GPS RO data distribution, which indicates that each GPS 160 RO profile is assigned to the connected subdomain based on its geographic location. The location of the GPS RO profile is not fixed and changes for every occultation. The geographic distribution of GPS RO profiles is not even (see Fig. 1). Since the nonlocal excess phase 163 operator for GPS RO data is very expensive and it is very likely that some processing cores 164 or subdomains get more GPS RO profiles than others. One may deduce that the overall 165 performance is solely determined by the workload of the processing cores which being as-166 signed the most number of profiles to process. Fig. 2(b) shows the variation of the maximum 167 number of assigned observations per subdomain with the numbers of processing cores (red 168 bars represent experiment "Parallel"). Because the uneven geographic distribution of the 169 GPS RO profiles, even we used 512 processing cores, there is still one out of 512 subdo-170 mains covers two GPS RO profiles and most of the processing cores are idle. Therefore, 171

the theoretic speedup with 512 processing cores is 106/2 = 53 and the actual speedup of 46 should be considered as efficient enough for this strategy. It is impossible to achieve the ideal speedup before solving the load imbalance issue. Visual comparison of (a) and (b) in Fig. 2 for experiment "Parallel" suggests that the parallel performance has a very high correlation with the maximum number of the observation per processing core/subdomain. The load imbalance is the bottleneck of the overall parallel performance.

We have indicated in Sec. 3 that the most suitable parallel strategy for the nonlocal GPS 178 RO operator is the ray-path-wise distribution. Taking advantage of the parallel strategy 179 implemented in Sec. 3, each processing core is aware of the global  $N_{mod}$ , which means that 180 each processing core can calculate the integral of Eq. 4 along any ray-path of any profile. 181 Therefore, it is feasible to distribute the GPS RO data ray-by-ray among processing cores 182 in a round-robin fashion. However, taking into account the cost of implementation, it is 183 much easier to distribute the GPS RO data profile-by-profile among processing cores in 184 terms of the amount of code modification. Please note that since different GPS RO profile 185 may include different number of GPS ray-paths at vertical levels above the tangent point 186 and below the model top, the profile-by-profile distribution may sacrifice some of the load 187 balance compared to the ray-by-ray distribution. 188

The experiment "Loadbalance" in Fig. 2(a) shows the wallclock time spent with the 189 number of processing cores up to 128. Actually, 106 is the minimum number of processing 190 cores for this case to get the maximum theoretic parallel efficiency since we have 106 GPS RO profiles. With the round-robin distribution of the GPS RO profiles, each processing core 192 is assigned with approximately equal number of the profiles. With 128 processing cores, 193 the wallclock time is reduced to 162 ( $\approx 2.5$  minutes) seconds for 5 iterations minimization 194 including initialization and I/O. Compared to 492 seconds with 128 processing cores and 279 195 seconds with 512 processing cores in Sec. 3, the load balance strategy tremendously increases 196 the parallel efficiency of the nonlocal excess phase GPS RO operator. The experiment 197 "Loadbalance" in Fig. 3 shows that the speedup with 128 processing cores is 80. More 198

precisely speaking, the speedup with 106 processing core is 80. Without the load balance strategy, the speedup with 128 processing cores is 33. Again, the high correlation between the wallclock times and the maximum number of observations per processing cores of experiment "Loadbalance" in (a) and (b) of Fig. 2 confirms the analysis that the importance of the load balance strategy in parallel data assimilation of GPS RO data with nonlocal operator.

# 5. Further Optimization

In variational data assimilation methods, not only the observation operator is needed to 205 calculate the innovation, but also the corresponding tangent linear and adjoint observation 206 operators are required during the minimization. The tangent linear and adjoint operators 207 are used to evolve the perturbations forward and backward along the basic trajectory, re-208 spectively. Therefore, it is an economic way to save the computational cost if the basic 209 trajectory could be recorded, other than recomputed. In terms of the nonlocal GPS RO 210 operator implementation in the WRFDA system (Chen et al. (2009)), the locations of each 211 ray-path should be recored during the innovation calculation and the recorded location of 212 each ray-path can be used in tangent linear adjoint operators directly. The experiment "Op-213 timization" in Fig. 2(a) shows the wallclock timing results with this further optimization. Averaged  $4\% \sim 10\%$  further acceleration is observed.

# 216 6. Summary and Discussion

The nonlocal excess phase operator for GPS RO data has been demonstrated to be a solid and promising method to simulate the the observed GPS excess phase delay from the model states. However, due to its high computational cost and the nonlocal nature in the algorithm which includes the integration along the GPS ray-path across the model domain, it is not easy to implement this new operator in data assimilation systems parallelized based

on the horizontal domain-decomposition method. Therefore, it has been tested only in some research configurations with affordable number of observations.

To parallel the nonlocal excess phase GPS RO operator, the first strategy is to enable 224 each processing core to be aware of the global simulated model refractivity, which is the only 225 variable needed for the nonlocal operator and is calculated in advance. Thus, each processing 226 core can process the GPS observations geographically located within its connected subdo-227 main. However, the performance analysis reveals that the load imbalance associated with the default geographic observation distribution among processing cores seriously constrains the parallel efficiency. Leveraging the implementation of the first strategy, GPS RO profiles 230 can be alternatively distributed among processing cores in a round-robin fashion, which en-231 sures the best possible load balance with available computing resource. The demonstration 232 case with 106 GPS RO profiles over a Antarctic domain shows that the wallclock time for 233 5 iterations minimization with WRFDA reduced from about 3.5 hours with one processing 234 core to approximately 2.5 minutes with 106 processing cores. This is affordable in terms of 235 both research and operational practices. 236

Depends on the height of tangent point of the GPS occultation, different GPS RO profiles
may have different number of levels, therefore, different number of ray-paths. As mentioned
before, better load balance and further acceleration is still possible if the GPS RO data is
distributed among processing cores ray-by-ray. However, the ray-path has to be determined
before the data distribution and this may include some substantial code changes.

# <sup>242</sup> 7. Figures and tables

- 243 a. Figures
- 244 b. Tables
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## REFERENCES

- Barker, D. M., et al., 2012: The weather research and forecasting (WRF) model's community
- variational/ensemble data assimilation system: WRFDA. Bull. Amer. Meteor. Soc., 93,
- 252 831-843.
- <sup>253</sup> Chen, S. Y., C. Y. Huang, Y. H. Kuo, Y. R. Guo, and S. Sokolovskiy, 2009: Assimilation
- of GPS refractivity from FORMOSAT-3/COSMIC using a nonlocal operator with WRF
- 3DVAR and its impact on the prediction of a typhoon event. Terr. Atmos. Ocean Sci.,
- 256 **20**, 133–154, doi:10.3319/TAO.2007.11.29.01(F3C).
- <sup>257</sup> Cucurull, L. and J. C. Derber, 2008: Operational implementation of COSMIC observations
- into NCEP global data assimilation system. Wea. Forecasting, 23, 702–711.
- <sup>259</sup> Cucurull, L., Y. H. Kuo, D. Barker, and S. R. H. Rizvi, 2006: Assessing the impact of
- simulated cos mic gps radio occultation data on weather analysis over the antarctic: A
- case study. Mon. Wea. Rev., **134**, 3283–3296, doi:10.1175/MWR3241.1.
- Healy, S. B., 2008: Forecast impact experiment with a constellation of GPS radio occultation
- receivers. Atmos. Sci. Lett., 9, 111–118.
- Healy, S. B., A. M. Jupp, and C. Marquardt, 2005: Forecast impact experiment with
- GPS radio occultation measurements. Geophys. Res. Lett., 32, L03804, doi:10.1029/
- 266 2004GL020806.
- Healy, S. B. and J. N. Thepaut, 2006: Assimilation experiments with CHAMP GPS radio
- occultation measurements. Q. J. R. Meteorol. Soc., 132, 605–623, doi:10.1256/qj.04.182.
- Huang, C. Y., Y. H. Kuo, S. H. Chen, and F. Vandenberghe, 2005: Improvements on typhoon

- forecast with assimilated gps occultation refractivity. Weather Forecast., 20, 931–953, doi:
- 10.1175/WAF874.1.
- <sup>272</sup> Kuo, Y. H., W. Schreiner, J. Wang, D. Rossiter, and Y. Zhang, 2005: Comparison of GPS
- radio occultation soundings with radiosondes. Geophys. Res. Lett., 32, L05817, doi:10.
- 274 1029/2004GL021443.
- Kuo, Y. H., S. V. Sokolovskiy, R. A. Anthes, and F. Vandenberghe, 2000: Assimilation of
- GPS radio occultation data for numerical weather prediction. Terr. Atmos. Oceanic Sci.,
- **11**, 157–186.
- Kuo, Y. H., T.-K. Wee, S. Sokolovskiy, C. Rocken, W. Schreiner, D. Hunt, and R. A. Anthes,
- 2004: Inversion and error estimation of GPS radio occultation data. J. Meteor. Soc. Japan,
- 280 **82**, 507–531.
- Kuo, Y. H., X. Zou, and W. Huang, 1997: The impact of GPS data on the prediction of an
- extratropical cyclone: An observing system simulation experiment. Dyn. Atmos. Oceans,
- 283 27, 439–470, doi:10.1016/S0377-0265(97)00023-7.
- Liu, H., J. Anderson, Y. H. Kuo, C. Snyder, and A. Caya, 2008: Evaluation of a nonlocal
- quasi-phase observation operator in assimilation of CHAMP radio occultation refractivity
- with WRF. Mon. Wea. Rev., 136, 242–256, doi:10.1175/2007MWR2042.1.
- Liu, H. and X. Zou, 2003: Improvements to GPS radio occultation ray-tracing model and
- their impacts on assimilation of bending angle. J. Geophys. Res., 108, 4548, doi:10.1029/
- 289 2002JD003160.
- Ma, Z., Y. H. Kuo, B. Wang, W. S. Wu, and S. Sokolovskiy, 2009: Comparison of local and
- nonlocal observation operators for the assimilation of GPS RO data with the NCEP GSI
- system: An OSSE study. Mon. Wea. Rev., 137, 3575–3587, doi:10.1175/2009MWR2809.1.

- Poli, P., P. Moll, D. Puech, F. Rabier, and S. Healy, 2009: Quality control, error analysis
- and impact assessment of FORMOSAT-3/COSMIC in numerical weather prediction. Terr.
- 295 Atmos. Oceanic Sci., **20**, 101–113.
- Shao, H., X. Zou, and G. A. Hajj, 2009: Test of a non-local excess phase delay operator for
- GPS radio occultation data assimilation. J. Appl. Remote Sens., 3, 033508, doi:10.1117/
- 1.3094060.
- Skamarock, W. C., et al., 2008: A description of the advanced research WRF version 3.
- Technical report, NCAR Tech. Note NCAR/TN-475+STR, 113 pp. [Available from UCAR]
- communications, P. O. Box 3000, Boulder, Co 80307-3000.].
- Sokolovskiy, S., Y. H. Kuo, and W. Wang, 2005: Evaluation of a linear phase observation
- operator with CHAMP radio occultation data and high-resolution regional analysis. Mon.
- Wea. Rev., **133**, 3053–3059.
- Zhang, X., Y. Liu, B. Wang, and Z. Ji, 2004: Parallel computing of a variational data
- assimilation model for GPS/MET observation using the ray-tracing method. Adv. Atmos.
- sor Sci., 21, 220–226, doi:10.1007/BF02915708.
- Zou, X., B. Wang, H. Liu, R. A. Anthes, T. Matsumura, and Y. J. Zhu, 2000: A ray-tracing
- operator and its adjoint for the use of GPS/MET refraction angle measurements. Q. J.
- R. Meteorol. Soc., **126**, 3013–3040, doi:10.1002/qj.49712657003.
- Zou, X., et al., 1999: A ray-tracing operator and its adjoint for the use of GPS/MET
- refraction angle measurements. J. Geophys. Res., 104, 22301-22318, doi:10.1029/
- 1999JD900450.

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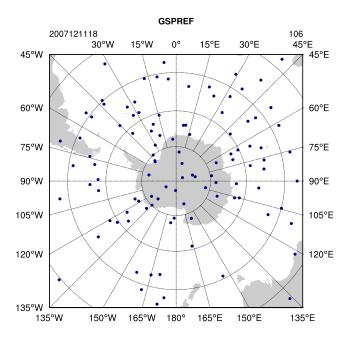
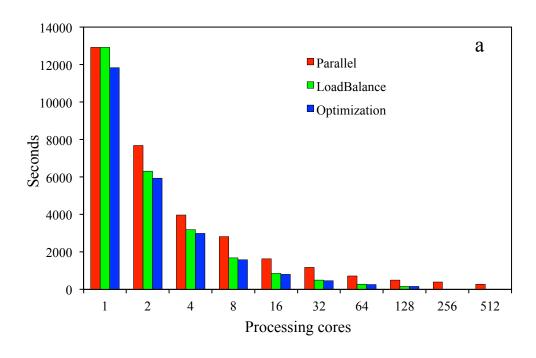


Fig. 1. Experiment domain and the locations of 106 GPS RO profiles (blue dots) within  $\pm 3$  hours of 1800 UTC 11 December 2007



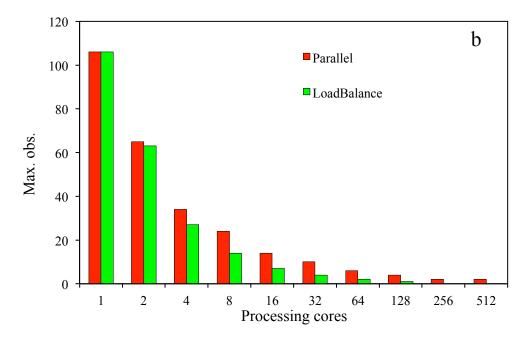


Fig. 2. The wallclock times (a) and maximum number of observation per processing core (b) for 5 iterations of minimization on NCAR yellowstone.

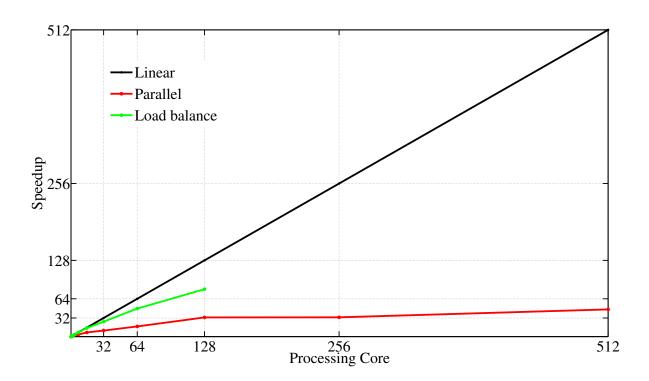


Fig. 3. The same as Fig. 2 , but for the parallel speedup