

Individual Project Summary 3/13

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Introduction

Rapidly progressing, anthropogenic climate change is not only real, but arguably the most dangerous threat facing human societies and global stability (Union of Concerned Scientists). In order to predict spatio-temporal changes in global climate under a variety of scenarios and assumptions, climate researchers developed computational models to emulate long-term atmosphere and ocean dynamics. There is hope that these models can provide policy and decision makers with predictive tools for climate change preparedness and mitigation (see, for example, [link](#)). A statistical hurdle, however, impedes this goal: model uncertainty. Much research has gone into understanding and limiting the sources of model uncertainty (Hawkins and Sutton 2009). Despite this progress, there remains a deeper concern regarding model independence and its effect on uncertainty quantification.

Because the models often share both code modules as well as the biases of their implementors, their outputs do not represent independent draws from a space of “all possible future climate trajectories”. Thus, agreement in model predictions does not grant greater certainty (Larose et al. 2005). To skirt this issue, Knutti et al describe a method to build consensus models from combinations of model prediction that takes into account their interdependence. By then placing a prior over all possible consensus models, independent samples of “possible future climates” can be drawn and used in analyses (Knutti 2010).

There are still a number of problems with this approach. For one, it doesn’t directly address the original central problem: the models are not independent, so model agreement doesn’t imply higher certainty. Beyond the lingering problems with prediction, we argue that there are shortcomings with the methods used in this result. In particular, each model is originally expressed - even before EOF (PCA) - in a severely limited form. Only 6 of the original climate variables are used in model comparison. This appears to be due to a computational constraint (at least, it certainly is on our machines). We would like to find a way to incorporate more variables into Knutti et al’s procedure. We would like to know if doing so will alter in any meaningful way their results, and if so, what variables are most to blame for the discrepancies.

Data Summary

We are using Coupled Model Intercomparison Project (CMIP5) data made available by ETH Institute for Atmospheric and Climate Science, as well as observational data gathered (according to (Knutti 2010) Table 1). We have access to the observational dataset, as well as data from a total of 46 models, 36 variables, under 10 different scenarios, combined using 47 different ensembles, available in daily/monthly/annual aggregations, and in two spatial coordinate grids (one provided by the model source, and another interpolated to a 2.5 by 2.5 degree grid).

To get an idea of the data in their most basic, time-series form, see Figure 1. This plot shows time series data for surface temperature taken from the ACCESS1-3 model, under the historicalGHG scenario, resampled with r1i1p1 ensemble, and grided to 2.5 by 2.5 degrees.

These data display the expected seasonal signature which can be seen in each time series’ auto-covariance plots (see figure 2). The degree to which this auto-covariance holds suggest that these data are suited for compression - a characteristic exploited by Knutti et al.

```
#knitr::include_graphics("/accounts/grad/yonid/Documents/Stat222/papers/boxpot.jpg")
```

In addition we can look at the entire spatial-map at an instance in time for a single model (see figure 3):

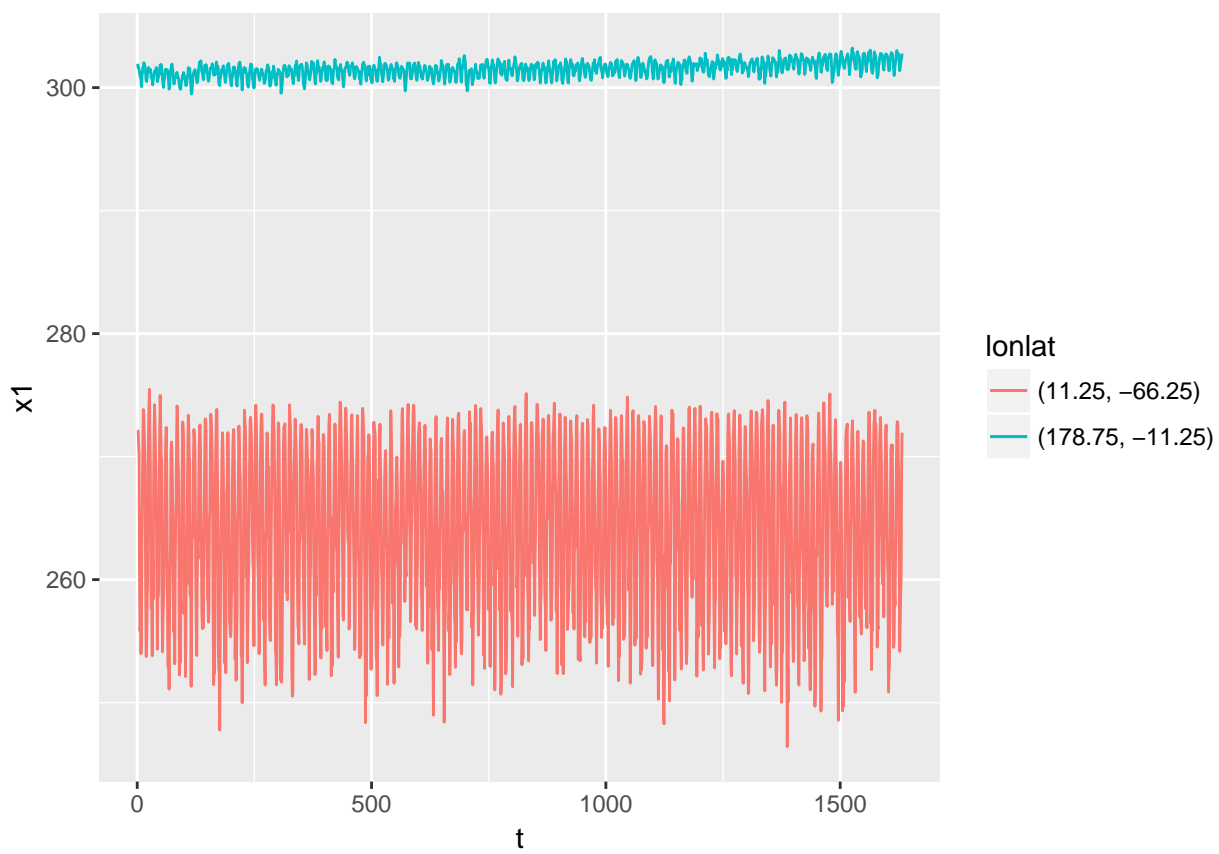


Figure 1: Surface Temperature at (178.75,-11.25) and (11.25, -66.25)

```
{r, echo = FALSE, fig.cap = "Geo-spatial variation in air surface temperature (K) 223  
years after simulation start"} knitr::include_graphics("presPic.png")
```

Here again we find strong auto-correlation, but this time of the spatial variety (see figure 4).

We can also look at the euclidean distance of all the timeseries to give us an idea of the amount of redundancy in the data (keeping in mind this is for air temperature at surface only):

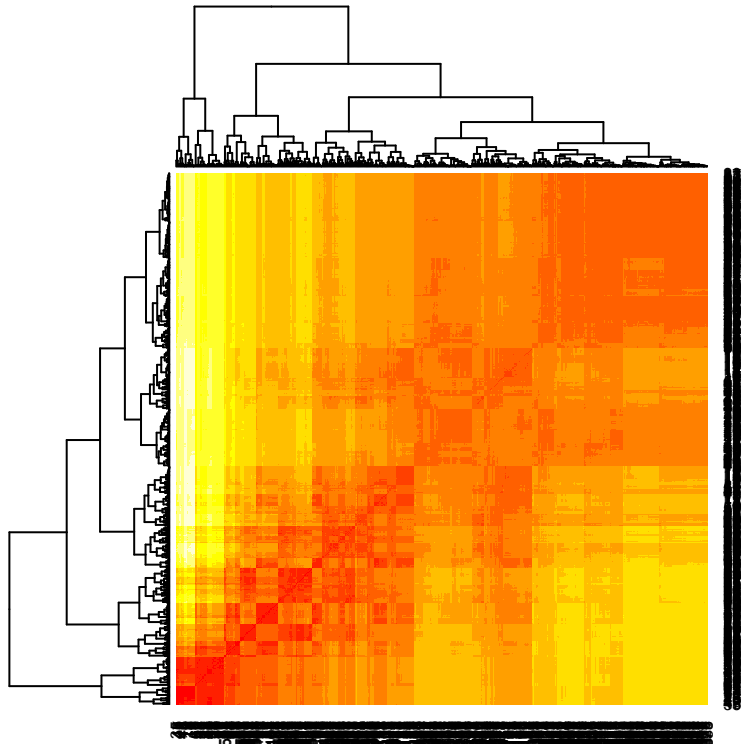


Figure 2: Heat map of spatial similarity of TAS time series

Despite the spatial autocorrelation demonstrated here, Knutti et al do not attempt to compress the data in the spatial domain.

Methods

We first attempted to reproduce (roughly) the pca/mds project methods performed in Knutti et al. which involve, in order: flattening the data for each variable; performing a PCA on the time-dimension to reduce dimensionality; reducing each model to a single vector by concatenating all compressed variable data; and finally performing an MDS on the matrix of all model rows (for mathematics, see below, excluding step A4).

We then attempted to incorporate our knowledge of spatial auto-correlation. After flattening the data we performed two PCA's: we first performed a PCA on the spatial dimension, selecting out the first N principle components, and then performed a PCA on the temporal dimension, represented by the transpose of the N principle components. We then selected K of these principle components and stretched them into a single vector representing the compressed data for that model variable. We then concatenate the vectors for all model variables involved to form a row representing the entire model in compressed form. With this row of data, we then proceeded with the Knutti et al methods, assembling a model matrix and performing an MDS on the model distances.

Our algorithm

- A) for mod^j in $Jmodelsandforvar_i$ in $Ivariables$:
 - 1) $M_{var_i} = mod^j_{.,.,.,var_i}$
 - 2) $T = (flatten(M_{var_i}))^T$
 - 3) Let P such that $T^T T = P^T \Lambda_{spatial} P$, then: $\tilde{M} = ((TP)_{.,1-N})^T$
 - 4) Let Q such that $\tilde{M}^T \tilde{M} = Q^T \Lambda_{time} Q$, then: $\hat{M} = (\tilde{M}Q)_{.,1-K}$
 - 5) $R_{var_i} = [\hat{M}_{1,.}, ..., \hat{M}_{N,.}]$
- Let $mod^j_{row} = [R_{var_1}, ..., R_{var_I}]$
- B) Let $\mathcal{M} = [(mod^1_{row})^T, ..., (mod^J_{row})^T]^T$
- C) Perform Classical MDS:
- Let $\Delta = [d^2_{uv}]$ where $d_{uv} = \sqrt{||mod^u_{row} - mod^v_{row}||}$
- Let $B = -\frac{1}{2}J\Delta J$ where $J = I - \frac{1}{J}\mathbb{1}\mathbb{1}^T$
- Let Q^B such that $B = Q^{BT} \Lambda_{dist} Q^B$, then: $\hat{\mathcal{M}} = Q^B_{.,1-2} diag \lambda_1, \lambda_2^{frac{1}{2}}$, where λ_1 and λ_2 are the two largest eigenvalues of B and $Q^B_{.,1-2}$ are the corresponding eigenvectors.

We chose N and K by keeping the eigenvectors for each dimension that contain 90 percent of the variation.

Results

Conclusions

There are multiple directions we could go. The first is to proceed with our initial goals: improve our methods by optimizing N and M and then incorporate all model variables as planned. Should this task prove too trivial, another more technically challenging option exists: in a sense, the most natural way to think about these models is not as matrices in $\mathbb{R}^{n \times n}$ but as multidimensional arrays in $\mathbb{R}^{L1 \times L2 \times T \times V}$ [as in citation]. We can then ask what multidimensional array in $\mathbb{R}^{l1 \times l2 \times t \times v}$, with $l1 < L1$, $l2 < L2$, $t < T$ and $v < V$, is closest to the original. Doing so would compress all dimensions at the same time, capturing correlations between and within dimensions that our flat PCA methods cannot. The caveats to this direction are the mathematics, the implementation, and the lack of guaranteed improvement in the results (given the effort needed for caveats one and two). Nonetheless, it is something to keep in mind.

Alternatively, we could switch gears all together. Instead of trying to improve on methods that sweep Independence under the rug, we could try to quantify the degree of non-independence. So far, we have found no indications in the literature that this has been done.

References

- Hawkins, Ed, and Rowan Sutton. 2009. "The Potential to Narrow Uncertainty in Regional Climate Predictions." *Bulletin of the American Meteorological Society* 90 (8): 1095–1107. doi:[10.1175/2009BAMS2607.1](https://doi.org/10.1175/2009BAMS2607.1).
- Knutti, Reto. 2010. "The End of Model Democracy?" *Climatic Change* 102 (3): 395–404. doi:[10.1007/s10584-010-9800-2](https://doi.org/10.1007/s10584-010-9800-2).
- Larose, Simon, Catherine F Ratelle, Frederic Guay, Marylou Harvey, and Evelyne Drouin. 2005. "in Science and Technology :". *Structure*: 171–192.