# **Cross-evaluation of Uncertainties in Extreme Precipitation Events using Multiplicative Triple Collocation**

## ABSTRACT

## INTRODUCTION

Extreme rainfall events associated with flash flood and landslides lead to tremendous damages including properties, fatalities, etc. According to Mazzoglio and Laio et al. (2019), extreme weather condition tends to intensify and become more frequent. It is thus a great concern to face these calamities. Among them, tropical cyclone is one of the most excessive rainfall producers because it carries great amount of water moisture from costal towards inland. Hurricane Harvey is termed as a category 4 hurricane that made landfall on Texas and impacted Louisiana, Oklahoma with devastating urban flooding and fatalities in August 2017. With data delivered by Texas state (Emanuel 2017), Hurricane Harvey produced the largest rainfall of hurricane in the recorded history of the United States, and caused at least 70 casualties, and economic lost beyond 150 million (Emanuel 2017, Omranian, Sharif et al. 2018). Harvey brings the maximum wind of 115 to 130 mph which damaged significant amount of trees, fences and power poles in the Metropolitan area. In Fall 2019, tropical cyclone Imelda made a landfall in Texas on 17, with similar impacted regions as Harvey. It produced 1096 mm total amount of rainfall in Texas, which is ranked as 4th in the history (contributors). Damages in Texas were reported to be beyond $3 million in this disaster while $1 billion over whole country (News 2019). The other two tropical cyclones Bill and Cindy, influenced the same region and bringing in copious rainfall in 2015 and 2017 respectively.

Traditionally, rainfall rate was captured by gauges as a direct measurement but only at a point. Even though gauge data are often treated as reference, it is still not impeccable as it suffers from splash-out with heavy rainfall, lack of sensitivity to light rain, under-catching by wind effect, and evaporations (Molini, Lanza et al. 2005, Hong and Gourley 2014). Especially in heavy rain events multiple studies demonstrated the error caused by these factors is not trivial, (Berlamont 2001, Molini, Lanza et al. 2005). Berlamont (2001) investigated the disadvantages of tipping bucket rain gauges during extreme weather conditions and found there was underestimation of rainfall volumes due to loss of water during the tipping action of the device. When merging point samplings to spatial coverage, it further introduces uncertainties coming from interpolation that may not be representative of rainfall variability depending upon the gauge quality and density. Stampoulis and Anagnostou (2012) discovered the convective nature of rainfall could increase gauge-interpolation uncertainties for the increased standard deviation values. The Wind from tropical cyclones would substantially affect the performance of rain gauges, with the relative bias ranging from 5 percent to 80 percent according to the speed (Pollock, Dutton et al. 2010), but no clear consensus on the percentage of undercatching under various wind speed categories (Medlin, Kimball et al. 2007). This is problematic when one applies gauge correction for other products e.g. radar, satellite in that event as gauge itself underestimates rainfall with up to 80 percent of bias.

In the past few decades, the emerging radar technology has been applied in metrology to measure rain rate by emitting and receiving electromagnetic signals. The most prominent advantage of radar over rain gauges is that radar provides a more refined spatial temporal scale and larger area coverage (He, Sonnenborg et al. 2013). But since it is an indirect measurement of rain rate, radar technology itself produces and propagates errors in the end products. The uncertainties of radar-based precipitation estimations can be categorized as incorrect calibration, sampling representativeness, non-weather echo and errors in Z-R relations (Medlin, Kimball et al. 2007, Ryzhkov, Diederich et al. 2014, Kirstetter, Gourley et al. 2015). These intrinsic systematic errors are challenging to mitigate by improving radar technology alone. Thus, many researchers tend to couple radar with rain gauges and satellites to improve the radar performance e.g. Kriging with External Drift (KED) (Jewell and Gaussiat 2015, Cecinati, Moreno-Ródenas et al. 2018), Mean Field Bias Correction (MFB) (Yoo, Park et al. 2014). Kidd et al. (2011) compared ground radar QPE with gauge data in Germany, and the results indicated radar overestimation in convective rainfall regimes. Medlin et al. (2007) has evaluated National Weather Service (NWS) Weather Surveillance Radar-1988 Doppler (WSR-88D) during Hurricane Danny. They concluded that both radar and rain gauge seriously underestimated event rainfall.

Another commonly used data source for retrieving precipitation is satellite attributing to the advantage of large spatial coverage and the top view. Satellite-based precipitation products utilize the information provided by Infrared/visible channels from geostationary (GEO) satellite, passive microwave sensor and spaceborne radar from low orbiting (LEO) satellite to provide accurate estimates. Just like weather radar QPE, satellite rainfall retrieval isan indirect measurement of rainfall. It is certain that the relationship between the satellite radiances retrieval and rain rate is less solid as compared to radar reflectivity vs. rain rate (Scofield and Kuligowski 2003). Many dedicated studies (Omranian, Sharif et al. 2018, Chaoying Huang 2019) stated that satellite data underestimated the magnitude of rain rate. (Hong, Hsu et al. 2006) demonstrated satellite performance decreased with the increase of rain rate but higher normalized bias in the light rain rates. (Chen, Hong et al. 2013) compared four satellite QPEs and ground based radar with gauge references for Typhoon Morakot, and they found satellite QPEs underestimate extreme rainfall possibly due to gauge smoother, missed precipitation outside of PMW overpasses, ice areas. (Omranian, Sharif et al. 2018) evaluated IMERG V05 final gauge adjusted products with The National Weather Service/National Centers for Environmental Prediction (NWS/NCEP) River Forecast Center (RFC) Stage-IV Quantitative Precipitation Estimates (QPEs) as reference in hurricane Harvey event. It revealed that IMERG is able to detect the spatial variability of the rainfall field and overestimated rain rainfall to some degree but it remains unclear whether NCEP stage IV radar QPE is able to be representative as the reference during Harvey event.

Given this context that no either source of product can be trustable, researchers performed stochastic approach to analyze these uncertainties.(Tian and Peters-Lidard 2010). Triple Collocation (TC) has been proven to be a statistical approach to estimate uncertainties within each of three independent products (Stoffelen 1998, McColl, Vogelzang et al. 2014, Massari, Crow et al. 2017, Li, Tang et al. 2018). TC was firstly applied to evaluate ocean surface wind variability by inputting different wind products (Stoffelen 1998). Thereafter, it has been extended to determine errors of sea surface temperature (Gentemann 2014), sea surface salinity (Ratheesh, Mankad et al. 2013), wave height (Caires 2003), leaf area index and soil moisture. (Roebeling, Wolters et al. 2012) was the first to apply TC in hydrometeorology to intercompare remote sensing, weather radar and rain gauges in Europe. (Massari, Crow et al. 2017) compared performance of five satellite QPEs over US, and deduced the results of correlation coefficient towards globe. (Alemohammad, McColl et al. 2015) introduced multiplicative triple collocation method (MTC), suggesting its appropriateness in rainfall error evaluation and then decomposed the error term in order to investigate the violation of assumptions. (Li, Tang et al. 2018) used TC method to perform uncertainty analysis over ungauged regions in Tibetan in China after validated TC with traditional intercomparing statistics. To the best of our knowledge, TC has not been utilized to evaluate the uncertainties in extreme weather conditions in which they contain more uncertainties than normal cases. To investigate the validity of TC and interpret the uncertainties with three common QPEs −rain gauge interpolated XXX, radar estimated XXX, satellite estimated XXX, the objectives of this study are designed as following: 1. Compare differences and performances of three independent products during multiple events; 2. Perform uncertainty estimation over multiple extreme events with special emphasis on hurricane Harvey; 3. Evaluate the applicability of TC during extreme events.

We organize this article into four sections: section 2 will introduce the study domain and briefly review the three datasets been used in this study. Section 3 will describe in detail the formula to derive Root Mean Squared Error (RMSE) and Correlation Coefficient (CC) from TC method. Section 4 and 5 will follow up with results and conclusions from this study. The structure of this article will start from a broad overview (multiple events) and then dive into specific event (Harvey) to scrutinize the differences.

(This is the experience from getting edits from Prof. Hong, you need to cite more paper from the journal you are going to submit to. There are only three papers from J of Hydrometeorology.)

## STUDY AREA and DATASETS

### ***Study domain***

The area we are of interest in the Gulf Coast of North America, which had multiple hurricane landings and set many extreme precipitation records throughout the US history. It is one of the most frequently impacted areas by tropical cyclones, such as Hurricane Harvey, and Tropical Storms Imeda, Bill, and Cindy. **Figure 1** illustrates the storm track and total accumulative precipitation amount of Tropical Storm Bill (2015), Cindy (2017) and Imelda (2019) as well as Hurricane Harvey from 2017. The impacted area of the aforementioned events contains the states of Texas, Oklahoma, Louisiana, Arkansas, Tennessee, Mississippi and Alabama, which accounts for almost 10 percent of the conterminous United States (CONUS). These events had similar moving patterns that they approached inland from Gulf of Mexico and bending towards Northwest after landing, except for Imelda which dissipated shortly after making the landfall. **Table 1** here listed out details about four events including their durations, amount of rainfall falling in that period. To be comparative, we concatenated Bill, Cindy and Imelda so that the total duration and also rainfall amount are similar to Harvey. After doing so, hurricane Harvey even observed slightly higher amount than the other three added together which illustrates the severity of that case. Due to the unprecedented nature of Hurricane Harvey, three other tropical storms were selected for analysis to improve the genericity of the study findings. All three tropical storms in total would match the duration of Hurricane Harvey, which gave the identical data sample size for two study groups: Harvey and non-Harvey.

### ***Datasets description***

NCEP gridded gauge only hourly precipitation product (NCEP) is an operational product (Lin 2011) covering the CONUS and parts of Puerto Rico. It is automatically derived from approximately 3000 operational hourly rain gauge observations across 48 states to produce 4km/hour rainfall field, using the interpolation method from (Seo 1998) , which introduced Double Optimal estimation (DO) and Single Optimal estimation (SO) to gain conditional expectation of rainfall estimation. This technique accounts for fractional coverage of rainfall due to sparse gauge networks. (Gourley, Hong et al. 2009) performed inter-comparisons of NCEP gauge only QPE, NCEP stage IV radar QPE and PERSIANN-CSS satellite QPE, and demonstrated that NCEP gave the best performance at longer time scale e.g. seasonal, daily. However, it encounters underperformance at shorter time scale (1 hour) especially for storm. NCEP gridded gauge only QPE were obtained from National Center for Atmospheric Research/Earth Observing Laboratory (NCAR/EOL):<https://data.eol.ucar.edu/dataset/21.088>.

MRMS has around 180 integrated operational radars including 146 S-band and 30 C-band radars, creating seamless 3D radar mosaic across the CONUS and Southern Canada at 1km/2min resolution (Zhang, Howard et al. 2016). It produces both radar based QPE and radar gauge corrected QPE to improve performance. The radar-based QPE was used in this study as an independent precipitation observation from radar. (Gao, Zhang et al. 2018) evaluated the performance of MRMS QPE during hurricane Harvey, they envisioned MRMS QPE underestimated total accumulated rainfall by a small factor and overestimated very light precipitation. For this study, 1km/2min radar based QPE were retrieved and aggregated by averaging to 4km/hourly, in order to be compatible with NCEPdata. Historical MRMS radar only QPE was downloaded from Iowa State Mesonet MRMS archive: <http://mtarchive.geol.iastate.edu/>.

IMERG satellite precipitation product V06 (Huffman 2019) is integrated from its core satellite (GPM Core Observatory), microwave constellations, Infrared and additional constellations, aiming at providing global coverage of rainfall field (90N-S from V06 onward) beyond its predecessor Tropical Rainfall Measurement Mission (TRMM). GPM Core Observatory has additional channels of dual-frequency precipitation radar (DPR) and GPM Microwave Imager (GMI) which are capable of detecting very light precipitation and falling snow (Skofronick-Jackson, Petersen et al. 2017). It produces three stages: early run, late run and final run with 4 hours latency, 12 hours latency and 3.5 months latency respectively at a half-hour and 0.1 arc-degree scale. Early run provides near real-time brief observations with inter-calibrated satellite products, and late run adds up the late coming high quality PMW data and climatological calibration. Final run compares late run product with Global Precipitation Climatology Project (GPCP), and adjust the factor to compensate for under/over-estimation (Huffman 2019). To account for independency, final run without calibration/late run with calibration is utilized in this study. In order to perform pixel-wise analysis, IMERG data needs to be downscaled by nearest neighbor and then accumulated to 1 hour. Current IMERG final run product V06 can be accessed at <https://disc.gsfc.nasa.gov/datasets/GPM_3IMERGHH_06/summary?keywords=IMERG>.

## METHODOLOGY

### ***Assumptions***

For TC method, a few assumptions are needed: 1) the three forcing data should be independent, for instance, they are derived from different instruments in our dataset; 2) The errors of three independent products should be independent or unrelated (,) which refers as zero cross-correlation. 3) The expectation of error is treated as zero known as the unbiasedness assumption which is often described in geostatistical analysis e.g. kriging.

The three precipitation products we selected based on gauge only, radar only, and satellite only to meet above criteria well because they are not able to interact with each other. In fact, three sensors are the natural testbed for TC in rainfall uncertainty estimation. The detailed TC functions and explanations are listed in Appendix A.

### ***Expressions***

The basics of TC method is to treat three independent products as equally important, and thus no bias is produced in between. Since no ground truth value assumed, it then uses linear combination of three products and affine transformed error model to derive root mean square error and correlation coefficient (Zwieback, Scipal et al. 2012).

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Where indicates each of the independent source data, is the “relative truth”, is the weights and biases to adjust, and represents error for each product.

(Tian, Huffman et al. 2013) then proposed a way to transform the additive error model to multiplicative by logarithmic transformation, and it is proved to be more appropriate in rainfall error estimation (Tian, Huffman et al. 2013, Alemohammad, McColl et al. 2015, Li, Tang et al. 2018). Hence, the error model can be reformed as:

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From that, we can derive rain rate and error model by transforming into linear combination so that it fits into TC method.

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Where demonstrating the multiplicative error, indicating the residual error, and as the deformation error. Hence, after we transform the model to be multiplicative form, these parameters along with error should be also in logarithmic form. In the analysis of (Alemohammad, McColl et al. 2015), they later transformed back error to linear scale by Taylor series expansion as first order approximation.

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here represents the mean of field. In doing so, the error field could be identified as linear scale, meaning the same unit of mm/hour as rain rate. From linear equation (3), we are able to derive RMSE in the following set of equations based on the covariance of triples (McColl, Vogelzang et al. 2014):

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(McColl, Vogelzang et al. 2014) introduced a way to evaluate correlation coefficient (CC) from manipulating covariance matrices with called “ETC” method in which CC is formed shown below as a set of equations.

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Since both RMSE and CC are derived from covariance between the three products, they infer the relative error which are treated as uncertainties. Therefore, the less uncertain product or the best performance in the following sections is the one that has lowest RMSE and highest CC. Likewise, the most uncertain is associated with highest RMSE and lowest CC.

### ***Data preparation and model setup***

The rainfall data retrieved from each source are preprocessed products, but containing different extent. To make them comparable, IMERG and MRMS data are scaled to the same degree of NCEP gauge only data with the same array dimensions, with 4 km spatial resolution and 1 hour temporal resolution. Because we need to transform data into logarithmic scale, it is important to filter out zero-rain and NAN data. The treatment of zero values is varying, either by simply removing zero values (Alemohammad, McColl et al. 2015, Massari, Crow et al. 2017) or replacing with near-zero values(Roebeling, Wolters et al. 2012). (Li, Tang et al. 2018) tested the sensitivity by reassigning , to non-rain samples. Removing zero values may shorten the samples especially for events based evaluation. Thus, we dropped NAN values and treat zero to be . To obtain more robust results and filter out noise, we utilized bootstrap with 500 trials for evaluation at each pixel, and stored the mean value of RMSE and CC.

### ***Statistical evaluation metrics***

A list of evaluation indicators are summarized in **(Table 2)**. As we assumed there was no ground truth in the extreme events, we will term “error” as “difference” between products e.g. Root-Mean-Square-Difference. Quantitative differences including Correlation Coefficient (CC), Root Mean Squared (RMS) difference are the 1st and 2nd order evaluations respectively. They are computed in a domain at each pixel (4km) for each pair. For categorical difference, we considered Probability of Detection (POD), False Alarm Rate (FAR), Critical Success Index (CSI). The “reference” data in the denominator is chosen by the least uncertain product calculated by TC method. For instance, TC predicted radar product to be less uncertain which is quantified by RMSE compared to satellite. Then we put radar as the relative reference data for the categorical difference.

Since it is more insightful to compute these metrics in a conditional form during extreme events (Sukovich, Ralph et al. 2014), We further condition these metrics within the core of Hurricane Harvey to validate the consistency between TC metrics and traditional statistical methods e.g. . This condition is based on different accumulated rainfall observed for each pixel but still adopts the similar formula in **(Table 2)** . We consider conditions when hourly rainfall exceeds 50 percentile, 75 percentile and 99 percentile based on the estimate itself. For instance, when evaluating the NCEP gauge performance at Harvey core, we filter out pixel time series of rain rate to be less than 50 percentile, 75 percentile and 95 percentile of gauge rain rate at collocated point, and use the remaining to compute conditional metrics.

### ***Harvey core separation***

In order to be specifically focus on extreme rainfall, then the core, non-core regions are identified. According to the MRMS radar observation, 95 percentiles of rainfall falling outside of the core region, which corresponds to 400 mm accumulative rainfall isoline, depicted in **(Figure 8)**. We further used that line as a separation between core and non-core region. In core regions, we computed the conditional metrics described above, and selected several representative pixels to plot the time series. (You need to explain more what’s the purpose of separating the core and non-core)

(During the method section, you didn’t explain anything about other non-Harvey events, you need to find a place to do that and state the reasons)

## RESULTS

### ***Overall behaviors***

The accumulative rainfall for all events, Harvey and non-Harvey in **Figure 2** shows that MRMS product always estimated the highest accumulative rainfall amount at south eastern Texas among all tropical cyclone events following with IMERG and NCEP.. We can observe the lumped patches-like pattern for NCEP data, meaning the quality of gauge during extreme events is varying and will be investigated further iwenshenn the Harvey events. But for MRMS and IMERG, they are stratified well which show the spatial continuity of rain field.

The first hand overview of applying TC to four events concatenated together, denoted as “All”, Harvey-only and non-Harvey that denoted as “other” are illustrated in **Figure 3** for CC**, Figure 4** for RMSE. Firstly, both CC and RMSE demonstrate consistent behaviour in which higher CC and lower RMSE. All three cases identified the same result that the ordered product uncertainties (from high to low) are MRMS, IMERG, NCEP according to the median value and metrics distribution at the whole domain. The median value is more insightful than mean value because there are noises that tend to shift up mean values especially in places where rainfall samples are limited. NCEP gauge data tends to have higher CC value in the western regions in which smaller amount of rainfall observed compared to core regions. MRMS is behaving more uniform that higher CC is wide spread over the domain. IMERG concentrates its higher value in the western region as well as NCEP, but cover more regions than NCEP. The violin plot inside each axes indicates the distribution of CC. They are more skewed to higher values (above 0.5) but still have some samples that close to zero, especially for NCEP data. These values normally appear outside of the core where we set non-rain samples to be . When the covariance of the estimated product e.g. NCEP with either other two products e.g. MRMS and NCEP becomes 0, it will impose estimated CC to be zero, meaning no correlation with the other two products. These low values will be eliminated when we only consider core regions. We would like to point out that, in event Harvey, CCs of three product is more divergent compared to either All or Other. This indicates the three products agree more with other for non-Harvey case, and also push us to investigate Harvey with special attention.

**Figure 4** depicted the RMSE value from TC results which reveals the other side of CC value. In **Figure 4**, we emphasized two remarkable consistent behaviours. One is for NCEP data that All and Harvey both experience higher RMSE value in Houston region. This needs our further scrutinization by inspecting its time series in the detailed Harvey analysis. The other for MRMS and IMERG shows anomalous signals at the boundary of Louisiana and Mississippi. The corresponding time series plot of rainfall at one location is selected and shown up in the small axes. The overall trend is well followed by three products except that MRMS radar has some anomalous impulse that disagrees with the other two data sources. This anomalous echo may not be coming from hydrometeors. It is probably caused by debris rolled up by strong wind that was recorded of 108 mph.

### ***Detailed Harvey Analysis***

**Table 3** listed both quantitative intensity difference and categorical difference in Harvey with three different percentiles. NCEP vs. MRMS, NCEP vs. IMERG, IMERG vs. MRMS are taken into account and the latter product is considered as “reference” because of lower uncertainty in the previous TC results. What this result conveys is with analogy to the TC results. IMERG vs. MRMS comparison has highest correlation coefficient for three percentiles, even though they are slightly lower than NCEP vs. IMERG in terms of RMS difference in lower percentiles. For categorical difference, all results suggest IMERG and MRMS behave relatively close while NCEP deviates from both MRMS and IMERG. It again validates the results that TC is capable to convey some consistent results with traditional methods.

The CC and RMSE from TC are further split based on accumulative rainfall by MRMS data at each interval of 50 mm in **Figure 5.** In doing so, we could identify the range that is suitable for each product to be having less uncertainty. As an overview, MRMS data is substantially robust and stable as rainfall increases, but NCEP gauge data generally acquire more uncertainties than IMERG and MRMS at all ranges. Besides, NCEP obviously aggravate its behavior at higher range of the rainfall i.e. above 1200 mm. IMERG data is behaving worse at lower tail i.e. below 150 mm. That indicates satellite data under-performs at lower tail probably due to the sensor sensitivity, signal attenuation and smoothing effect of large size of the footprint. It is been reported in literatures that IMERG data typically tends to overestimate light rain and underestimate heavy rainfall (Guo, Chen et al. 2016, Sharifi, Steinacker et al. 2016, O, Foelsche et al. 2017). (Omranian, Sharif et al. 2018) also concluded that IMERG final product has generally better performance with higher precipitation rates compared to lower rates in the case of hurricane Harvey.

**Figure 6** depicts TC metrics CC and RMSE for three grouped regions – whole, core and non-core. After thresholding, we clearly witnessed the improved performance for both MRMS and IMERG – higher CC and lower RMSE – in terms of median value and uncertainty bound. However, the RMSE for NCEP even remained unchanged. This points out the noise removal for NCEP data inside the core is superseded by the degradation of performance probably due to the impact of intensive wind or out-of-splash. While for the performance of each product, the results are still remained, which again validate the consistency of TC method. The whole region performance sits in between core and non-core because it neutralize the two tails. Looking at the distribution of rainfall for each product in **Figure 7**, It associates with the characteristics of the product itself. For NCEP gauges, It is likely to underestimate total rainfall because of splash-out-of-water, wind under-catching and not representative of rainfall variability. It is thus obvious for NCEP data to be concentrated in the range of 400 to 600 mm. Even though IMERG data is more wide spread than NCEP, it still cuts off at 1100 mm. The reasons behind are myriad e.g. the sensitivity of sensors, type of sensors (IR, PMW, DPR) but most importantly, since IMERG has resolution of 0.1 degree (around 10 km), it acts like a smoother that take the average of the cell. Hence, it is difficult to capture fine scale rainfall field. To be noted, MRMS data could also suffer from overestimation since we observed some anomalous impulses during this event, but we tend to believe that MRMS radar QPE is more or less close to “reference” as TC results suggested.

By inspecting the time series at selected representative points in **Figure 8**, we are able to unravel the reason that NCEP data behaved relatively worse. Point 1,3, 4 and 5 are picked based on the RMSE spatial map in which NCEP data had highest value or maximum rainfall are captured. In the corresponding time series plots, the grey window mentioned that in that certain period, NCEP data either recorded as zero value but the other two captured intensive rainfall (point 1,3) or stopped recording any data (point 4,5). This anomaly could be caused by wind that blows rain drops out of the range, or mechanical misfunctioning but more likely due to uncertainties introduced by interpolation. Because of the variability of rainfall field, sparse gauge network (e.g. 3000 in total around the U.S.) is not able to capture this variability. Point 2 is selected as IMERG data recorded maximum amount of rainfall in this event. The horizontal blue line marked the maximum rain rate that GMI (GPM Microwave Imager) can record due to sensor sensitivity (Skofronick-Jackson, Petersen et al. 2017). In other words, IMERG product will cut off any rain rate larger than 60 mm/h except places swept by DPR (dual-frequency precipitation radar). Nevertheless, DPR has limited swath and coverage. Point 4 is where MRMS has relatively low CC value inside the core. We indeed found that MRMS sometimes jumped up instantaneously while NCEP and IMERG showed some agreements such as highlighted red window. It probably attributes to non-weather echoes or applying incorrect Z-R relations.

**Figure 9** has depicted the conditional inter-comparison results for each pair. As expected, those metrics are getting worse when larger percentile of data are exceeded because larger rain rate is associated with more difference. Except for RMS difference, all other statistics suggest IMERG and MRMS are more or less similar in whatever kinds of extreme rainfall conditions which is again the same signature in **Table 3**. It once again proved that NCEP data inside Harvey core has certain degree of degradation. Researchers need to be cautious when deciding to use this data set to evaluate extreme weather conditions, and additional justification should be provided.

## CONCLUSIONS AND FUTURE WORKS

In this experiment, we tried to explore the applicability of TC method in extreme events, and interpret the results from it. It is an ill-posed statement because we cannot find valuable reference data to absolutely validate the performance of TC method even though a great amount of previous works has proven that TC is a powerful statistical tool to analyse the uncertainties of three independent precipitation products. However, instead of directly compare TC with traditional methods for given reference data, we could evaluate the consistency of TC in the following ways:

1. Applying TC in the overall extreme events combined, and separately perform it with individuals. The results in **(Figure 3 and 4)** have suggested that MRMS data are always providing lowest uncertainties, and then follows IMERG satellite QPE. NCEP data are associated with largest amount of uncertainties, which may not be an appropriate product in evaluating rainfall at extremes.
2. The traditional evaluation metrics e.g. RMS difference, Correlation Coefficient, POD, FAR, CSI and also their conditioned values are well in line with the TC results. They indicate that NCEP data have larger difference with MRMS and IMERG.
3. The results to separate Harvey core and non-core **(Figure 5)** also demonstrated NCEP data performance is not substantially improved inside the core while the other two get different amount of increment. This points out that NCEP data may be subject to degradation in the Harvey core.
4. The time series for collocated pixels where gauge have more uncertainties inside the core showed that a large portion of NCEP data didn’t record rainfall while the other two products indeed observed some amount. We believe that gauge sparsity plays the first order role, and then it may also subject to other effects e.g. wind under-catching, splash-out-of-water, mal-functioning.
5. IMERG data has the second largest uncertainty because: 1. Its large resolution smooths out rainfall variability so that it does not capture enough rainfall **(Figure 8)**; 2. The sensitivity for IMERG data is not trivial in extreme conditions **(Figure 7)** as it cuts off rain rate beyond certain threshold; 3. Signal attenuation may also lower the rain rate captured by satellite QPEs; 4. Gauge may also smooth out extreme values for gauge adjusted satellite QPEs.

The TC results also provide us insightful considerations. Firstly, it can assess the quality of each product. In our extreme rainfall cases, we found NCEP gauge based QPE encountered some problems that may be due to the uncertainty introduced by interpolation, or systematic error e.g. wind under-catching, splash-out-of-water etc. IMERG satellite QPE sits in between NCEP and MRMS. MRMS radar QPE is proved to be the most reliable dataset in capturing extreme rainfall. Secondly, it tells the story about limiting boundaries of each product. NCEP gauge data increase its uncertainties as accumulative rainfall amount increases. One can infer that gauge-like QPE are more susceptible to higher rainfall rate due to its own deficiencies mentioned above. IMERG data have higher uncertainties at lower tail meaning light rainfall could cumber its performance. MRMS radar data are behaving more robust and stable at either high rain amount or low rain amount. It suggests that MRMS radar QPE is more appropriate to evaluate precipitation during extreme events.

This paper serves as the first order overview of the quality of each product during extreme events as each suffers from its own deficiencies. However, it is still unclear which systematic error plays the first role for radar and satellite QPE. It motivates us to explore more detailed error decomposition in the future researches. Beyond that, we believe that the gauge corrected products i.e. radar and satellite will not get improved performance in extreme events. Further comparisons regarding to this topic could also help developers to adopt more reasonable algorithms for quantitative precipitation estimation.

## ACKNOWLEDGEMENTS

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