

The Variogram

Introduction

- **Spatial Interpolation** is the domain of **geo-statistics**, which deals with a random variable that is observed at fixed sample point locations. It aims at **estimating a *spatial field* around the sample locations**.
In contrast, spatial statistical analysis deals with lattice/areal data (aggregation of points information within an area).
- The variogram focuses on **estimating the *second order variation* of a stationary surface based on the observed values at the sampling points**.
relationships between individuals
- A spatial field $Y(s_i)$ can be decomposed into three components:

$$Y(s_i) = \underbrace{\mu(s_i)}_{\text{general trend}} + \underbrace{\varepsilon'(s_i)}_{\text{error term structured model it through individuals}} + \underbrace{\varepsilon''(s_i)}_{\text{totally random}}$$

first order = exogenous trend	second order = stochastic inter-dependence	white noise = nugget effect
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- A **regionalized random variable $\varepsilon'(s_i)$** is continuous from point to point, but the **spatial variability with respect to the locations s_i** is too complex that it cannot be described with any tractable deterministic first order function (such as a trend surface).
Thus, its spatial variation of the stochastic variable $\varepsilon'(s_i)$ is best described by a covariance structure (function).
- The **first order** component can be generalized to any expected value of the random variable $E[Y(s_i)] = \mu(s_i)$ at an arbitrary location s_i . For trend surfaces this expectation simply becomes **$\mu(s_i) = \mathbf{x}^T(s_i) \cdot \boldsymbol{\beta}$**

- The concept of stationarity applies to the **covariance** among the pair of **random components** $\varepsilon'(s_i) + \varepsilon''(s_i)$:
 - $E[\varepsilon'(s_i) + \varepsilon''(s_i)] = 0$ and $Var[\varepsilon'(s_i) + \varepsilon''(s_i)] = const$
 - $Cov\left[\left(\varepsilon'(s_i) + \varepsilon''(s_i)\right), \left(\varepsilon'(s_j) + \varepsilon''(s_j)\right)\right]$ does not depend on the absolute locations of s_i and s_j , but only on their distance (and perhaps under *anisotropy* on the direction).
- The spatial arrangement of the sample points **constitutes the support of the regionalized variable**. Predictions outside the support become extrapolations. means applied area, once beyond this area, the discovered rule are not

Historical Background.

- South African mining engineer D.G. Krige. Early 1950 developed optimal core grade estimation from sample points.
- Meaning of "optimal"**: Interpolation weights $\lambda_i(s)$ are chosen to optimize the interpolation function $\hat{Y}(s) = \sum_{i=1}^n \lambda_i(s) \cdot Y(s_i)$ for any prediction location s given the sampling locations s_i , variance of prediction is smallest means best i.e., to provide a **Best Linear Unbiased Estimate (BLUE)** of the random field $\hat{Y}(s)$ at any location s .
- Geo-statistics was established by Georges Matheron of the Centre of Morphologie Mathematique in Fontainebleu, France. (early 1960).
- However, similar aspatial estimation approaches are known for a long time in other branches of sciences.
- Due to this origin in the mining industry much of the lingo of geo-statistics is in terms of mining engineering (e.g., **drift**, **sill** and **nugget**)

Basic Kriging models

first order component equal to 0

- For **simply kriging** it is assumed that $E[Y(s_i)] = 0 \quad \forall s_i \in \mathfrak{R}$, that is $\mu(s_i) = 0$ is given externally
- For **ordinary kriging** we assume that $E[Y(s_i)] = c \quad \forall s_i \in \mathfrak{R}$ has a constant expectation $\mu(s_i) = c$ unequal to zero. Thus the expectation does not depend on the location
- For **universal kriging** we assume that $E[Y(s_i)] = \mu(s_i) \quad \forall s_i \in \mathfrak{R}$ has spatially varying expectation, e.g.,
 $\mu(s_i) = \mathbf{x}^T(s_i) \cdot \boldsymbol{\beta}$.
 for instance, trend surface value
- The ordinary and universal kriging models can always be reduced to the simple kriging model by first estimating the expectation component $\hat{\mu}(s_i)$ and then transforming the observed spatial field $Y(s_i)$ by $[Y(s_i) - \hat{\mu}(s_i)]$ because $E[Y(s_i) - \hat{\mu}(s_i)] = 0$.

The expectation component $\hat{\mu}(s_i)$ must be estimated by **feasible generalized least squares** with an estimated covariance structure because the observed spatial field $Y(s_i)$ is autocorrelated, which violates the **OLS independence assumption**.

(Recall *ordinary least squares* requires that the individual observations are **independent** of each other, i.e.,
 $Cov[Y(s_i), Y(s_j)] = 0$ for all $s_i \neq s_j$)

Deterministic Inverse Distance Weighted Interpolator

- Inverse distance weighted interpolation is used frequently to obtain a first idea about the spatial variability of a spatial field $Y(s_i)$.
- Its spatial dependence (i.e., covariation) among the sample points is determined **deterministically**.

- It is defined as

$$\hat{Y}(\mathbf{s}) = \frac{\sum_{i=1}^n [h(\mathbf{s}_i, \mathbf{s})]^{-\alpha} \cdot Y(\mathbf{s}_i)}{\sum_{i=1}^n [h(\mathbf{s}_i, \mathbf{s})]^{-\alpha}}$$

with $h(\mathbf{s}_i, \mathbf{s})$ being the distance between the sample location \mathbf{s}_i and the prediction location \mathbf{s} and $\alpha \geq 0$ a smoothing parameter.

- By convention the smoothing value $\alpha = 2$ is used:
 - For $\alpha = 0$ extreme smoothing is achieved with $\hat{Y}(\mathbf{s}) = \frac{\sum_{i=1}^n Y(\mathbf{s}_i)}{n}$ at all prediction locations \mathbf{s} .
 - For $\alpha = \infty$ the smoothed pattern resembles that of a Voronoi polygon with each tile having a value equal to $\hat{Y}(\mathbf{s}) = Y(\mathbf{s}_i)$.
- For prediction locations $\mathbf{s}_i = \mathbf{s}$ the predicted value is $\hat{Y}(\mathbf{s}) = Y(\mathbf{s}_i)$.
- In deterministic smoothing the standard error of the prediction $\hat{Y}(\mathbf{s})$ cannot be evaluated, because no distribution assumptions are made.

Covariance and Semi-variogram

- Most geo-spatial surfaces exhibit some degree of *spatial persistence*. I.e. they exhibit *positive spatial autocorrelation*.
- The auto-covariance $C(\mathbf{s}_i, \mathbf{s}_j)$ between two points is

$$C(\mathbf{s}_i, \mathbf{s}_j) = E \left[\left(Y(\mathbf{s}_i) - \underbrace{\mu(\mathbf{s}_i)}_{\text{residual}} \right) \cdot \left(Y(\mathbf{s}_j) - \mu(\mathbf{s}_j) \right) \right] \text{ or in terms of the autocorrelation } \rho(\mathbf{s}_i, \mathbf{s}_j) = \frac{C(\mathbf{s}_i, \mathbf{s}_j)}{\sigma(\mathbf{s}_i) \cdot \sigma(\mathbf{s}_j)}$$

- For an **isotropic stationary** spatial process the covariance $C(\mathbf{s}_i, \mathbf{s}_j)$ reduces to $C(\mathbf{s}_i, \mathbf{s}_j) \Rightarrow C(\mathbf{s}_i - \mathbf{s}_j) \Rightarrow C(h_{ij})$, which depends only on the distance h_{ij} between points $\mathbf{s}_i = (x_i, y_i)^T$ and $\mathbf{s}_j = (x_j, y_j)^T$.

Note: **spherical** distances are better suited for small-scale maps to accommodate the curvature of the earth.

- For an **isotropic stationary** spatial process we get for the dissimilarity $Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})$:
 - $E[Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})] = E[Y(\mathbf{s} + \mathbf{h})] - E[Y(\mathbf{s})] = 0$
 for the individual expectations being **zero** or **constant** over space
 => Works for **simple and ordinary kriging** because any constant mean surface in the difference cancels out.
 => Consequently, we **do not need to know the mean** as long as it is constant at all locations
 - Otherwise we need to work with regression residuals, because their expectations are supposed to be identical within the study area: $E\left[\left(\varepsilon'(\mathbf{s}_i) + \varepsilon''(\mathbf{s}_i)\right) - \left(\varepsilon'(\mathbf{s}_j) + \varepsilon''(\mathbf{s}_j)\right)\right] = 0$
 => Works for **universal** kriging.
 - The variance of the distance at a distance $h = |\mathbf{h}|$ between two locations $Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})$ can be expressed as **semi-variogram** $\gamma(h)$ for simple and ordinary kriging:

$$\underbrace{Var[Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})]}_{=2\gamma(h)} = \underbrace{Var[Y(\mathbf{s} + \mathbf{h})]}_{=\sigma^2} + \underbrace{Var[Y(\mathbf{s})]}_{=\sigma^2} - \underbrace{2 \cdot C[Y(\mathbf{s} + \mathbf{h}), Y(\mathbf{s})]}_{=C(h)}$$

$$2 \cdot \gamma(h) = 2 \cdot \sigma^2 - 2 \cdot C(h)$$

$$\gamma(h) = \sigma^2 - C(h)$$
 because $Var(X \pm Z) = Var(X) + Var(Z) \pm 2 \cdot C(X, Z)$.

- The term **intrinsic stationarity** refers to the stationarity of **spatial difference** $[Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})]$ of regionalized random variable $Y(\mathbf{s})$ at two locations and not to the random variable $Y(\mathbf{s})$ itself. If the random field $Y(\mathbf{s})$ is weakly stationary then $[Y(\mathbf{s} + \mathbf{h}) - Y(\mathbf{s})]$ must be also intrinsic stationary.
- The semi-variogram is defined as $\gamma(h) = \sigma^2 - C(h) \Leftrightarrow C(h) = \sigma^2 - \gamma(h)$, thus the covariance can be directly calculated from the semi-variogram. It only depends on the distance h between the locations $\mathbf{s} + \mathbf{h}$ and \mathbf{s} .
- Relationship between the **covariance**, **correlation** and the semi-variogram

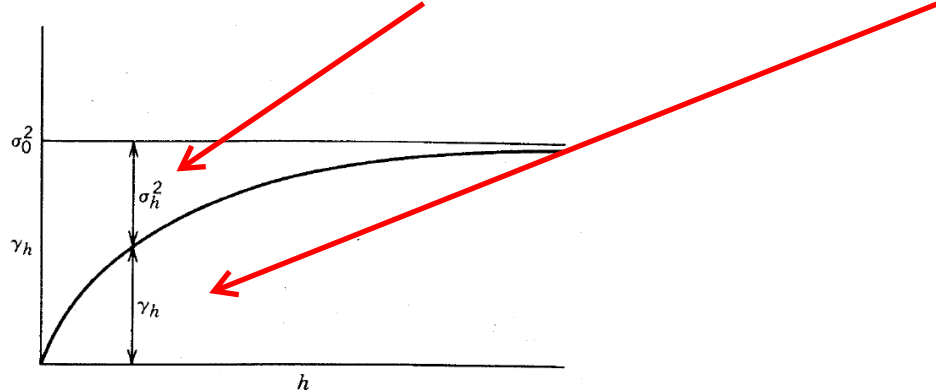


FIGURE 4.43 Relationship between semivariance γ and autocovariance σ^2 for a stationary regionalized variable. σ_0^2 is the variance of the observations, or the autocovariance at lag 0. For values of h beyond the range, $\gamma_h = \sigma_0^2$.

- Theoretical **properties** of the semi-variogram for intrinsic stationarity fields:
 - At $h = 0 \Rightarrow \gamma(0) = \sigma^2 - C(0) = 0$
 $\quad \quad \quad = \sigma^2$

- At $h \rightarrow \infty \Rightarrow \gamma(\infty) = \sigma^2 - C(\infty) = \sigma^2$. This is the **sill**.

$$\text{Var}(Y(s)) = 0$$
- The **threshold distance h** where the sill is reached is called **range**.
- If at some distance range the sample data exhibit **negative** spatial autocorrelation then the semi-variogram overshoots the sill.

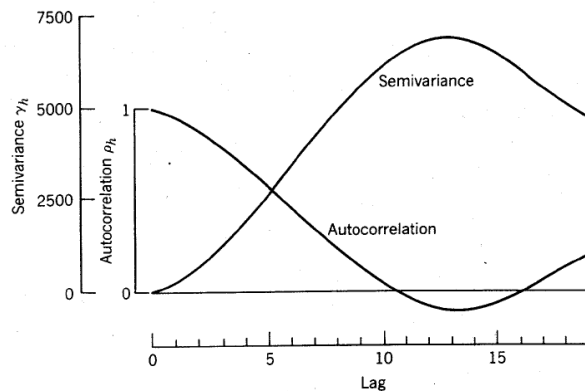


FIGURE 4.44 Relationship between semivariance γ_h and autocorrelation ρ_h for a stationary regionalized variable.

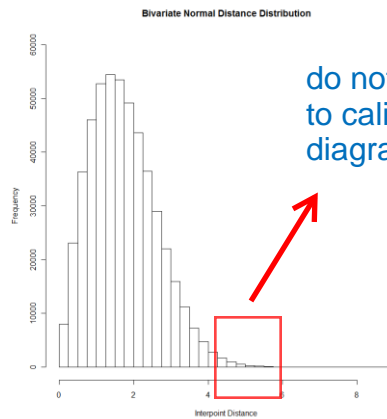
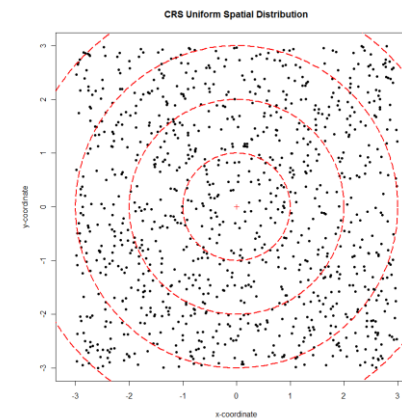
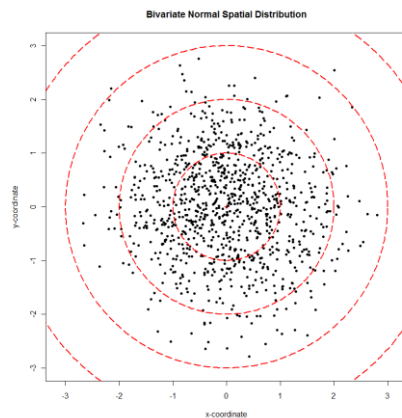
if the semi variance keep increasing, then would violate the stationary assumption (sill is variance.)

- If a semi-variogram is equal to the sill at all distances $h \geq 0$ then there is
 - **no spatial autocorrelation in** the underlying stochastic component $\varepsilon'(\mathbf{s}_i)$ and
 - $\gamma(h) = \text{Var}(\varepsilon''(\mathbf{s}_i))$ for all distances h .
 Consequently, spatial interpolation becomes meaningless because $\varepsilon''(\mathbf{s}_i) \sim i.i.d.$, thus it does not exhibit a spatial relationship.
- How fast the semi-variogram approaches the sill, if at all, for large h s is an indication of the presence of non-stationarities.

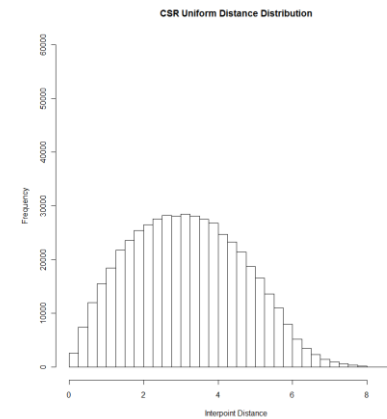
- One expects that spatial dependencies are fading out at some threshold distance h .
 - If this is not the case then non-stationarities are present.
 - For non-stationary spatial fields the **ergodasticity** assumption is violated because effects of a regionalized random variable will not be forgotten even at very high distances.
 - In theory, $\gamma(0 + \delta)$ must almost equal zero for any incremental small distance value δ .
If this is not the case we observe a **nugget effect**.
 - The **nugget effect** arises because the regionalized variable is
 - [a] so erratic over a short distance that the semi-variogram goes from zero to the level of the nugget effect in a distance less than the smallest sampling interval h_{min} between two points $\mathbf{s} + \mathbf{h}$ and \mathbf{s} .
 - [b] or the spatial process exhibits a discontinuity over a short distances (e.g. moving from inside a nugget to the surrounding environment). measurement error
 - [c] or our measurements $Y(\mathbf{s})$ at locations \mathbf{s} are impacted by some error $Var[\varepsilon''(\mathbf{s}_i)] \neq 0$.
- => Implication: A **predicted surface** $\hat{Y}(\mathbf{s})$ does not need to go exactly through the observed value $Y(\mathbf{s}_i)$ at the sampling locations \mathbf{s}_i .

Estimation of the semi-variogram

- Distribution of inter-sample point distances within a bound study area



do not use those distance
to calibrate the semi-variance
diagram



Notes:

- for higher distances the number of point pairs is decreasing again.
- For the normal distributed locations from the map center the squared distances h^2 are $h^2 \sim \chi^2_{df=2}$.

- Steps of variogram estimation:

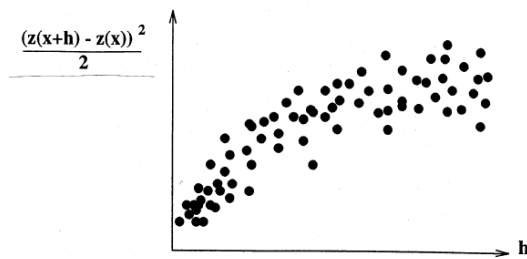


Figure 4.2: Plot of the dissimilarities γ^* against the spatial separation h of sample pairs: a variogram cloud.

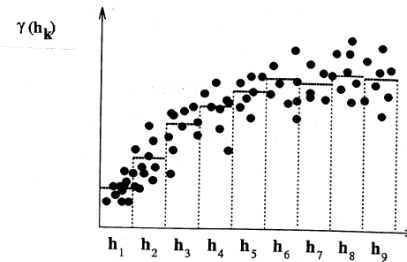


Figure 4.3: The experimental variogram is obtained by averaging the dissimilarities γ^* for given classes h_k .

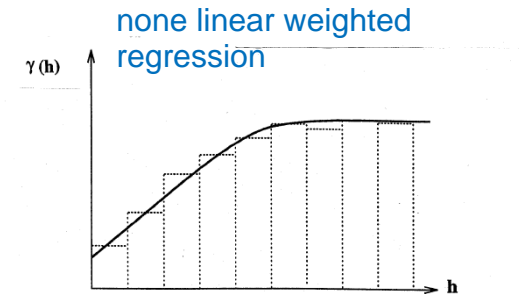


Figure 4.4: The sequence of average dissimilarities is fitted with a theoretical variogram function.

1. Plot pairwise squared semi-dissimilarities $(Y(s+h) - Y(s))^2 / 2$ for distances from zero to approx. half the **diameter** (spatial extend in either direction) of the study region.
 2. Take the averages of $(Y(s+h) - Y(s))^2 / 2$ in **distance bands**, which are also called **bins**.
 3. Fit a **theoretical semi-variogram functional form** to the mean levels in each bin:
 - => Elementary textbooks suggest fitting it by visual inspection of the sill, range and nugget effect
 - => Advanced methods use non-linear weighted statistical curve fitting techniques
 - The number of squared dissimilarities in each bin determines the weights.
 - A large number of pair-wise dissimilarities in a bin indicate higher precision of the estimate mean of the particular bin.
- => There are also approaches which do not require binning and use the pairwise squared semi-dissimilarities directly (see Bivand et al. p 228)

- What do we gain by estimating the semi-variogram instead of the covariance structure directly?

- Discrete intervals $[h_k - h_{k+1})$ are replaced by a function in the continuous distances h between two points.
- The semi-variogram is a simple function in just a few parameters.
- We do not need to know the mean of the spatial field as long as it is constant.
- We are not bound to the set of observed distances among observed **sample point** pairs.
This gives us the flexibility to estimate the co-variation between any two points as long as we know their distance.
Thus we can add **prediction points** and immediately calculate the covariation of these points with the given sample points.
- The semi-variogram is a smooth function and eliminates sampling variations from an empirically observed covariance matrix.
- The covariance matrix between the spatial objects can be derived from the theoretical forms of semi-variogram function $\gamma(h)$: $C(h) = \sigma^2 - \gamma(h)$
- However, only **specific functional specifications** for a semi-variogram function are permitted because the covariance matrix $C(h)$ has to satisfy the specific conditions:
 - [a] **symmetry** and all its eigenvalues are positive
 - [b] **positive definiteness** for any value of h .
- This leads to a selected number of **feasible** functional semi-variogram model specifications. For instance:

○ Spherical model:

$$\gamma(h) = \begin{cases} 0 & h = 0 \\ a + (\sigma^2 - a) \cdot \left(\frac{3 \cdot h}{2 \cdot r} - \frac{h^3}{2 \cdot r^3} \right) & 0 < h \leq r \\ \sigma^2 & \text{otherwise} \end{cases}$$

a is the nugget effect, σ^2 the sill and r is the range.

Discuss function's behavior at $h = 0$, $h = r$ and beyond.

○ Exponential model:

$$\gamma(h) = \begin{cases} 0 & h = 0 \\ a + (\sigma^2 - a) \cdot [1 - \exp(-3h/r)] & h > 0 \end{cases}$$

The exponential model will never exactly reach the sill, however, for large h it will come extremely close to it. **It can model mild non-stationarities.**

- For more models and a discussion of their properties see Waller & Gotway pp277-280.

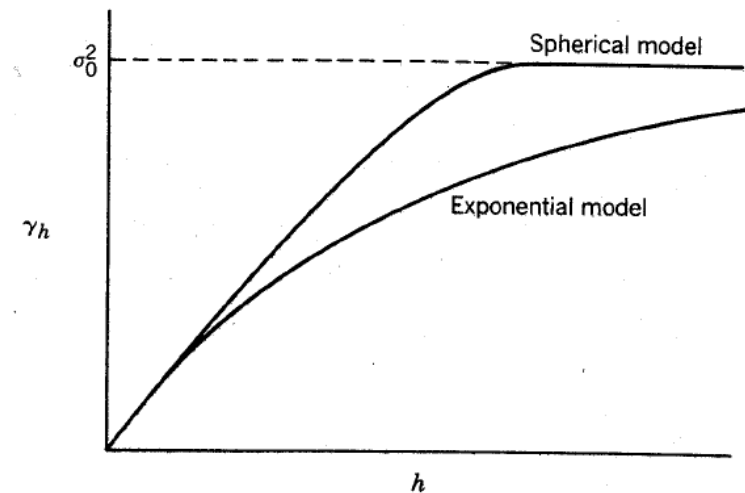


FIGURE 4.47 Exponential and spherical models of semivariogram. Both models have same initial slope and sill. After Clark (1979).

Extensions:

- The semi-variogram model can be relaxed to **accommodate simple forms anisotropy** (beyond the scope of this lecture). Here the sill or the range may differ depending on the direction of the distance measurement (see the Geo-Statistical Analyst in ArcGIS)

- The semi-variogram can be **generalized to a linear combination of several semi-variogram functions** such as one for shorter ranges and one for longer ranges, e.g., $\gamma(h) = \lambda \cdot \gamma_1(h) + (1 - \lambda) \cdot \gamma_2(h)$ (see Bivand et al. p 224 and `gstat::vgm(..., add.to=vgmComponent)` *divided data into two sets(short distance and long distance)*)
- One needs to question, however, if the more complex semi-variogram models add sufficiently new information and what their underlying theoretical justification is?