



Geophysical Research Letters*

RESEARCH LETTER

10.1029/2022GL100092

Key Points:

- Multiple satellite observations, modeling, and reanalysis data were used to reveal groundwater storage changes in the Tibetan Plateau
- Groundwater storage increased (5.59 ± 1.44 Gt/yr) in the entire Tibetan Plateau from 2003 to 2016
- Declining solid water
 (-17.72 ± 1.53 Gt/yr) including
 glaciers, snow, and permafrost
 dominates increasing groundwater
 storage

Supporting Information:

Supporting Information may be found in the online version of this article.

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Citation:

Zou, Y., Kuang, X., Feng, Y., Jiao, J. J., Liu, J., Wang, C., et al. (2022). Solid water melt dominates the increase of total groundwater storage in the Tibetan Plateau. *Geophysical Research Letters*, 49, e2022GL100092. https://doi. org/10.1029/2022GL100092

Received 16 JUN 2022 Accepted 7 SEP 2022

Solid Water Melt Dominates the Increase of Total Groundwater Storage in the Tibetan Plateau

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Abstract Understanding how groundwater storage (GWS) responds to climate change is essential for water resources management and future water availability in the Tibetan Plateau (TP). However, the dominant factor controlling long-term GWS changes remains unclear and its responses to climate change are not well understood. Here we combined multi-source datasets including in-situ measurements, satellite observations, global models, and reanalysis products to reveal that GWS increased at 5.59 ± 1.44 Gt/yr during 2003-2016 while showing spatial heterogeneities with increasing trends in northern TP and glacial regions and declining trends in central and southern TP. The accelerated transformation from solid water (glaciers, snow, and permafrost; -17.72 ± 1.53 Gt/yr) into liquid water provide more recharge to groundwater, dominating the total GWS increase. This study contributes to a better understanding of the hydrological cycle under climate change and provides key information for projecting water availability under different future scenarios in the TP.

Plain Language Summary Long-term groundwater storage (GWS; the amount of water stored in the pores, fissures, and caves of soil and rocks within the underground saturated zone) changes in response to climate change in the Tibetan Plateau (TP), and its dominant controlling factor has not been fully understood. There are very limited observation wells in the TP to monitor the groundwater level changes. This study used satellite observations, modeling, and reanalysis data to investigate GWS changes in the TP during 2003–2016. Results show that the GWS increased by 5.59 ± 1.44 Gt/yr during this period. The accelerated climate warming on the TP has led to significant solid water melt, including glacier melt, snow melt, and permafrost thaw. The solid water storage showed a decreasing trend (-17.72 ± 1.53 Gt/yr) and provided more water to infiltrate to recharge groundwater. The solid water melt has been identified to be the dominant source of GWS increase in the TP. The results of this study will help us to better understand the changes in the water cycle on the TP caused by climate change.

1. Introduction

Known as the Asian water tower, the Tibetan Plateau (TP) plays a critical role in providing water resources for over two billion people living downstream (Immerzeel et al., 2010). A better understanding of water storage changes is essential for water resources management to ensure water security and food production (Immerzeel & Bierkens, 2012). Groundwater sustains rivers in winter and buffers droughts to maintain fragile ecosystems, especially in arid or semi-arid regions (de Graaf et al., 2019; Famiglietti, 2014; Fan et al., 2013; Gleeson et al., 2016; Rodell et al., 2009; Taylor et al., 2013; Yao et al., 2021). Over the past decades, TP has experienced dramatic glacier retreat, permafrost degradation, and lake expansions caused by global warming (Brun et al., 2017; Cheng & Wu, 2007; Feng et al., 2019; Ji et al., 2020; Kuang & Jiao, 2016; Yao et al., 2012, 2022; Yi et al., 2021; Zhang et al., 2020). The accelerated transformation from solid water into liquid water may alter groundwater recharge, storage, and discharge to varying degrees, resulting in substantial changes in the groundwater system (Cheng & Jin, 2013; Gao et al., 2018; Huss & Hock, 2018; Lin et al., 2020; Yin et al., 2021). However, how groundwater storage (GWS) responds to climate change is poorly known because of limited observations caused by the harsh environment throughout the TP (Andermann et al., 2012; Feng et al., 2018). The Gravity Recovery and Climate Experiment (GRACE) mission provided an unprecedented opportunity to track water moving around the world (Tapley et al., 2004). By removing other hydrological contributions including soil moisture,

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ice, and surface water from Terrestrial Water Storage (TWS) changes, GWS changes can be obtained (Rodell & Famiglietti, 2002). GRACE is an effective tool to evaluate large-scale GWS changes in data-scarce or ungauged basins (Chen et al., 2016), especially the TP.

Recently, significant progress has been made in monitoring water storage changes in lakes, glaciers, snow, permafrost, soil moisture, and groundwater over the TP. Lake volume changes in 1,257 lakes (>1 km²) were monitored by multiple satellites (Li, Long, et al., 2019; Wang, Wang, Li, et al., 2021; Zhang, Bolch, et al., 2021). Hugonnet et al. (2021) estimated global glacier mass changes with high spatio-temporal resolution and precision. Kraaijenbrink et al. (2021) combined the snow model with satellite observations to estimate snowpack variations during the past four decades in High Mountain Asia. Other researchers also simulated active layer thickness (ALT) changes in the TP (Ji et al., 2022; Ran et al., 2021). Soil moisture in-situ observation networks were constructed to validate satellite retrievals and improve the modeling of soil moisture (Dente et al., 2012; Su et al., 2011; Yang et al., 2013). Previous studies found an increasing GWS in the subregions of the TP from the early 2000s (Bibi et al., 2019; Chao et al., 2019; Jiao et al., 2015; Xiang et al., 2016; Zhang et al., 2017) and attributed the GWS changes to various factors including precipitation, evapotranspiration, glacier retreat, permafrost degradation, snow melt, ecological project, etc. However, the dominant factors controlling the long-term changes in GWS remain unclear and its response to climate change is not well understood, thereby calling for improved estimations of the water storage components and a holistic analysis of GWS changes as well as their driving factors.

This study integrates multi-source datasets including in-situ measurements, satellite observations, global models, and reanalysis products to study changes in each water storage component and quantified their contributions to the TWS variations in 2003–2016 over the TP, particularly analyzing GWS changes and their controlling factors. In addition, we compared GRACE-derived GWS (G-GWS) changes with other independent estimations of GWS changes from baseflow recession, wells, Global Hydrological Models (GHMs), and Land Surface Models (LSMs). The value of this work is in providing a better understanding of the hydrological cycle under climate change and fundamental information for projecting water availability under different future scenarios in the TP.

2. Materials and Methods

The TP extends from the northern foothill of the Western Kunlun-Qilian Mountains in the north to the southern foothill of the Himalayas and other southern mountains in the south; from the Hindu Kush Mountains and the western edge of the Pamir Plateau in the west to the eastern edge of the Hengduan Mountains and other eastern mountains in the east with a total area over three million square kilometers and an average elevation of about 4,320 m (Zhang, Li, et al., 2021). It can be divided into the Endorheic Basins (including Amu Dayra, Tarim, Hexi Corridor, Qaidam, and Inner Plateau) and Exorheic Basins (including Indus, Ganges, Brahmaputra, Salween, Mekong, Yangtze, and Yellow (Figure S1 in Supporting Information S1; Zhang et al., 2013). We combined the Amu Dayra, Tarim, and Hexi Corridor (Amu-Tarim-Hexi); the Indus and Ganges (Indus-Ganges); the Mekong and Salween (Mekong-Salween) together, respectively, in analysis to ensure that the area requirement (200,000 km²) of GRACE's resolution accuracy is met (Rodell et al., 2018).

TWS changes are mainly the summation of water changes in lakes (L), glaciers (G), snow water equivalent (SWE), permafrost (PM), soil moisture (SM), and GWS in the TP (Zhang et al., 2017). Mass change signals caused by tectonic uplift, soil erosion, plant canopy, biomass changes, and river water storage changes were ignored as suggested by Xiang et al. (2016). GWS changes can be calculated by removing contributions of other components from TWS changes using the following water balance equation (Zhang et al., 2017):

$$\Delta GWS = \Delta TWS - \Delta L - \Delta G - \Delta SWE - \Delta PM - \Delta SM \tag{1}$$

Changes in solid water storage (including G, SWE, and PM) and liquid water storage (including L, SM, and GWS) were also analyzed.

TWS changes were from the Jet Propulsion Laboratory mascon products (JPL-M; Watkins et al., 2015) and the Center for Space Research mascon products (CSR-M; Save et al., 2016; Text S1 and Table S1 in Supporting Information S1). We compiled lake volume changes of 1,257 lakes from Li, Long, et al. (2019), Wang, Wang, Li, et al. (2021), and Zhang, Bolch, et al. (2021) (Text S2, Figure S2, and Table S2 in Supporting Information S1).

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Glacier mass changes were from Hugonnet et al. (2021) that used extensive Digital Elevation Models (DEMs) data to estimate global glacier mass changes with high accuracy and fine spatio-temporal resolution and validated against independent measurements (Table S3 in Supporting Information S1). SWE was retrieved from Kraaijenbrink et al. (2021), which combined the regional snow model with satellite observations to model snowpack changes and validated it against in-situ measurements and spaceborne satellite radar data (Text S3, Figure S3, and Table S4 in Supporting Information S1). For the permafrost, we retrieved ALT data from the Community Land Model Version 5 (Lawrence et al., 2019) and converted the ALT changes into mass changes by the method in Xiang et al. (2016) (Text S4, Figure S4-5, and Table S5 in Supporting Information S1). We evaluated 11 SM from LSMs, GHMs, and reanalysis products, and selected the ensemble mean of the Land component of the fifth generation of European ReAnalysis (ERA5-Land; Muñoz-Sabater et al., 2021), Modern-Era Retrospective Analysis for Research and Applications version 2 (MERRA2; Gelaro et al., 2017), and Mosaic from Global Land Data Assimilation System (GLDAS; Rodell et al., 2004) because of their superior performances over the TP (Text S5, Figure S6–S8, and Table S4 and S7 in Supporting Information S1). Three other independent estimations including baseflow-derived GWS (B-GWS), well-derived GWS (W-GWS), and GWS from GHMs and LSMs, were compared with G-GWS (Text S7, Figure S9-10, and Table S9-10 in Supporting Information S1). Meteorological data including precipitation and air temperature were used to analyze their effects on GWS changes (Text S8 and Table S8 in Supporting Information S1). Detailed information on processing procedures of data format (Text S8 in Supporting Information S1), trend estimation (Text S9 in Supporting Information S1), uncertainties, and component contribution ratio (Text S10-S11 in Supporting Information S1) are in the Supporting Information.

3. Results

3.1. Water Storage Components Trends and Their Contributions to TWS Variations

In the whole TP, glaciers (-16.61 ± 0.92 Gt/yr) dominated TWS trends, followed by lakes (7.28 ± 0.32 Gt/yr), groundwater (5.59 ± 1.44 Gt/yr), while permafrost (-1.10 ± 0.39 Gt/yr), soil moisture (0.66 ± 1.65 Gt/yr), and snow (-0.27 ± 0.86 Gt/yr) were minor (Figures 1a-1k; Table S11 in Supporting Information S1). Lakes (6.78 ± 0.32 Gt/yr) and glaciers (-14.91 ± 0.72 Gt/yr) played a major role in TWS trends in the Endorheic and Exorheic Basins, respectively. Among the sub-basins, lakes played a predominant role in TWS trends in the Inner Plateau (6.26 ± 0.28 Gt/yr) and Yellow (0.62 ± 0.02 Gt/yr). Glaciers dominated TWS trends in the Indus-Ganges (-7.28 ± 0.41 Gt/yr), Mekong-Salween (-1.43 ± 0.06 Gt/yr), and Brahmaputra (-5.49 ± 0.22 Gt/yr). Soil moisture dominated TWS trends only in the Amu-Tarim-Hexi (-0.74 ± 0.22 Gt/yr). Groundwater dominated TWS trends in the Qaidam (0.66 ± 0.13 Gt/yr) and Yangtze (2.16 ± 0.34 Gt/yr). Snow and permafrost had no prevailing trends in all regions. Besides, solid and liquid water showed distinct divergent changes with declining and increasing trends over the TP, respectively, except for the Mekong-Salween (Figure 11; Table S13 in Supporting Information S1).

Glaciers (28.70 \pm 0.84%) contributed to the largest TWS variations in the whole TP, followed by soil moisture (25.23 \pm 1.06%), groundwater (19.28 \pm 0.88%), lakes (12.08 \pm 0.29%), while both snow (8.53 \pm 0.70%) and permafrost (6.19 \pm 0.35%) were less than 10% (Figure 1m; Table S12 in Supporting Information S1). Lakes (32.20 \pm 0.64%) and glaciers (36.41 \pm 1.05%) accounted for most of the TWS variability in the Endorheic and Exorheic Basins, respectively. In the sub-basins, lakes were the major contributor to TWS variations in the Inner Plateau (37.69 \pm 0.76%). Glaciers contributed the most TWS variations in the Indus-Ganges (45.24 \pm 1.20%) and Brahmaputra (36.55 \pm 0.96%). Soil moisture was the major contributor to TWS variability only in the Yangtze (49.33 \pm 1.49%). Groundwater has the leading contribution in the Qaidam (45.08 \pm 1.60%), Amu-Tarim-Hexi (41.50 \pm 1.36%), Mekong-Salween (38.13 \pm 1.49%), and Yellow (29.91 \pm 1.60%). Snow and permafrost had no predominant contribution in all regions. Contributions of solid and liquid water to TWS variations were 43.87 \pm 1.68% and 56.13 \pm 1.60% in the whole TP, respectively (Table S13 in Supporting Information S1). Solid water had the dominant contribution only in the Indus-Ganges (58.59% \pm 1.74%), while liquid water dominated in other regions.

3.2. Spatio-Temporal Variations of GWS Changes in the TP

GWS showed a statistically significant increasing trend (5.59 \pm 1.44 Gt/yr, p < 0.01) in the entire TP from 2003 to 2016 (Figure 2c). This means that the total GWS increased by 78.26 \pm 20.16 Gt during this period, nearly

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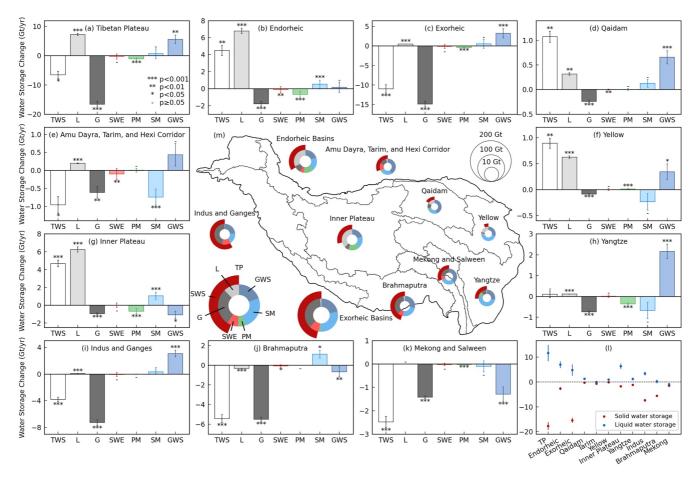


Figure 1. Trend of each water storage component in the (a) Tibetan Plateau (TP), (b) Endorheic Basins, (c) Exorheic Basins, (d) Qaidam, (e) Amu Dayra, Tarim, and Hexi Corridor, (f) Yellow, (g) Inner Plateau, (h) Yangtze, (i) Indus and Ganges (j) Brahmaputra, and (k) Mekong and Salween (k) from 2003 to 2016. (l) Trends of solid and liquid water storage in the TP and its sub-basins in 2003–2016 (Error bars represent the corresponding uncertainty of the GWS trends; *** indicates the significance level p < 0.001; ** indicates the significance level p < 0.05; - indicates the significance level p < 0.05; Tarim represents Amu Dayra, Tarim, and Hexi Corridor; Indus represents Indus and Ganges; Mekong represents Mekong and Salween). (m) Component contribution ratio of each water storage component to TWS variations in the TP and its sub-basins during 2003–2016 (The inner-circle diagram represents the magnitude of TWS variations). The basin boundaries were from Zhang (2019).

twice as much as the maximum capacity of the Three Georges Reservoir (Jiao et al., 2015). GWS increased both in the Endorheic (0.17 ± 0.60 Gt/yr) and Exorheic (3.23 ± 1.17 Gt/yr) Basins (Figure S11 in Supporting Information S1). However, GWS trends showed substantially spatial heterogeneities with increasing trends in northern TP and glacial areas while declining trends in central and southern TP (Figure 2a). GWS in nearly half of the domain showed declining trends, while the total GWS had an increasing trend. The glacial regions with increasing GWS accounted for small areas but their magnitudes are large than those in areas with declining GWS trends, thus having substantial effects on the whole TP. In sub-basins, GWS showed increasing trends in the Amu-Tarim-Hexi, Yellow, Yangtze, Indus-Ganges, and Qaidam ($0.34 \pm 0.15 \sim 3.06 \pm 0.50$ Gt/yr) while declining trends in other sub-basins ($-1.29 \pm 0.31 \sim -0.67 \pm 0.52$ Gt/yr; Figures 2c and 2d–2k).

4. Discussions

4.1. Solid Water Melt Dominates GWS Increase in the TP

From the perspective of sources and sinks of the groundwater budget, the substantial increase in the GWS was empirically linked to more recharge, while we found the declining precipitation from three different sources $(-0.58 \sim -2.00 \text{ mm/yr}, p > 0.05)$ during 2003–2016 (Figure S12 in Supporting Information S1). Ma and Zhang (2022) reported the overall increasing evapotranspiration in 1982–2016. Declining precipitation indicated

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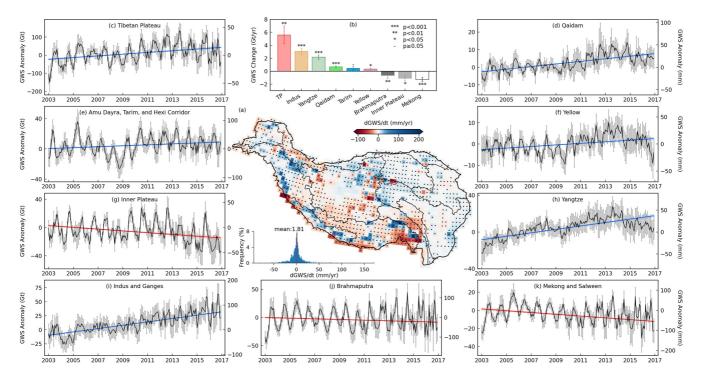


Figure 2. The spatial and temporal variations of groundwater storage (GWS) in the Tibetan Plateau (TP) and its sub-basins. (a) Spatial distribution of GWS trends from 2003 to 2016 in the TP (Grids with the significant trend (p < 0.05) are marked with black dots; the frequency histogram is GWS trends at all grids and the red line represents the mean value of all grids). (b) Ranking of GWS trends in the TP and its sub-basins during 2003–2016 (Error bars represent the corresponding uncertainty of GWS trends; *** indicates the significance level p < 0.001; ** indicates the significance level p < 0.05; - indicates the significance level p < 0.05; Tarim represents Amu Dayra, Tarim, and Hexi Corridor; Indus represents Indus and Ganges; Mekong represents Mekong and Salween). Monthly time series of GWS anomalies in the (c) TP, (d) Qaidam, (e) Amu Dayra, Tarim, and Hexi Corridor, (f) Yellow, (g) Inner Plateau, (h) Yangtze, (i) Indus and Ganges (j) Brahmaputra, and (k) Mekong and Salween during 2003–2016 (Linear trends are shown in red (declining) or blue (increasing) lines and their 95% confidence intervals are illustrated by transparent shades; error bars represent the corresponding uncertainty).

no more groundwater recharge from precipitation, so the increasing GWS should be recharged by other potential sources. Other potential groundwater recharge sources have not significantly changed except for solid water. Hence, the main reason for the increasing GWS was likely that groundwater gained more recharge from solid water melt (glaciers, snow, and permafrost) due to warming (air temperature trends range from -0.01 to 0.07° C/yr over the TP; Figure S13 in Supporting Information S1). We attribute the increasing GWS to declining solid water rather than the specific components because glacier mass changes from Hugonnet et al. (2021) include snow (Wang et al., 2018) that are difficult to separate.

The dominant controlling factors of GWS changes varied among the sub-basins. In the sub-basins including the Amu-Tarim-Hexi, Yellow, and Yangtze with increasing GWS, solid water melt was the dominant factor because of the declining solid water and precipitation (Figures 2e, 2f, and 2h; Figures S13e, S13f, and S13h in Supporting Information S1). Over the Qaidam and Indus-Ganges, solid water melt was still the predominant factor in the upward GWS as the increase in precipitation was not statistically significant enough to explain the increasing GWS (Figures 2d and 2i; Figures S13d and S13i in Supporting Information S1). In other sub-basins with declining GWS including the Inner Plateau, Brahmaputra, and Mekong-Salween, declined precipitation likely dominated the decreasing GWS since solid water was melting (Figures 2g, 2j and 2k; Figures S13g, S13j, and S13k in Supporting Information S1). Although declining GWS seemed to contradict lake expansions in the Inner Plateau, the increasing GWS was consistent with upward lake volume in the northern Inner Plateau (Wang, Wang, Li, et al., 2021). Besides, our results have the same spatial pattern of GWS trends as Qiao et al. (2021) with declining and increasing trends in the northern and southern Inner Plateau (Figure 2a), respectively. The downward GWS in the southern Inner Plateau may leak out by cross-basin groundwater flow through tensional fractures and deep normal faults (Yong et al., 2021). Interestingly, Wang, Wang, Zhou, et al. (2021) found that glacier mass loss in the southeastern TP did not offset the declined runoff in the upper Brahmaputra during 1998–2019. Our results demonstrated that the glacier retreat in the southeastern TP also did not offset the decreasing GWS in the upper

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Brahmaputra (Figures 2a and 2j). Although research subjects are different between our study and Wang, Wang, Zhou, et al. (2021), the spatio-temporal changes of runoff and GWS are similar, reflecting the same effects of glacier melt on the long-term changes in surface water and groundwater. The weakened Indian monsoon reduced water vapor transportation (Ouyang et al., 2020; Turner & Annamalai, 2012), resulting in less precipitation after 1998 Wang, Wang, Zhou, et al., (2021) and ultimately declining GWS in the Brahmaputra, while glacier melt did not reverse this downward trend.

Our results can also be supported by previous isotopic studies that demonstrated the dominant role of meltwater in groundwater recharge in glacial regions (Kong et al., 2019; Lone et al., 2021). Kong et al. (2019) found that non-monsoon precipitation (meltwater mainly coming from non-monsoon precipitation) accounts for $87 \pm 28\%$ of groundwater recharge at the Mingyong glacier in Hengduan Mountains, southeastern TP, even though this region is affected by the Indian monsoon. The mechanism of the high recharge ratio of groundwater from meltwater in Kong et al. (2019) is that soil water was nearly saturated at the beginning of monsoon seasons and the monsoon precipitation rarely infiltrated because meltwater from glaciers and snow continually infiltrated during non-monsoon seasons and rainfall is too concentrated to infiltrate in monsoon seasons. Besides, Lone et al. (2021) showed that meltwater supplies 83% of groundwater recharge in the upper Indus, highlighting the significance of the cryosphere in maintaining groundwater resources.

Increasing GWS caused by glacier melt was also found in other alpine regions. For example, Castellazzi et al. (2019) found that declining glacier mass was compensated by increased GWS in the Canadian Rocky Mountains and explained that part of the meltwater did not immediately flow into rivers but rather was stored in local aquifers, further supporting our results and indicating that this phenomenon may occur in global alpine regions.

4.2. Comparison of GWS Changes With Previous Studies

Our results showed consistency with previous studies in directions of GWS trends and an improvement in seasonal variations (Table S14 in Supporting Information S1). The directions of GWS trends in our study were in line with Xiang et al. (2016) except for the two regions. The seasonal GWS variations in Xiang et al. (2016) were much larger than our estimates because they only used linear trends of glaciers, lakes, and permafrost in calculations (Figures 1c–1k). In the Yangtze, we estimated an increasing GWS (2.16 ± 0.34 Gt/yr), which was comparable to Chao et al. (2019) (3.80 ± 0.50 Gt/yr). In the Qaidam, our results showed an increasing GWS (0.66 ± 0.13 Gt/yr) in 2003-2016, which was between the estimations in Bibi et al. (2019) (0.45 Gt/yr) and Jiao et al. (2015) (1.62 ± 0.14 Gt/yr). In the Inner Plateau, our results indicated an decreasing GWS (-1.09 ± 0.41 Gt/yr), which were comparable to Qiao et al. (2021) (-0.83 Gt/yr) but quite different from Zhang et al. (2017) (5.01 ± 1.59 Gt/yr). Zhang et al. (2017) concluded that GWS increases were comparable to lake volume increases (7.72 ± 0.63 Gt/yr) during 2003–2009. However, recent studies confirmed that TWS changes in the Inner Plateau were dominated by lake volume changes (Jiang et al., 2020; Xu et al., 2022), which are in agreement with our results. Leaving aside other factors, GWS changes in our study and Qiao et al. (2021) should be more reasonable as current studies use the updated GRACE products.

4.3. Comparison With GWS Estimations From Baseflow Recession and Wells

Figure S14 in Supporting Information S1 shows that the direction of trends in B-GWS and G-GWS are consistent in the eight headwater basins of the TP except for the headwater of Yangtze. Previous studies have confirmed that the magnitude of trends and variations in B-GWS is far lower than that of G-GWS (Cao et al., 2019; Schmidt et al., 2020). However, B-GWS still provides valuable independent information (Liu et al., 2020), and the consistencies between these two independent methods also indicate the robustness of our results. G-GWS has the same direction of GWS trends as W-GWS except in Golmud; G-GWS and W-GWS have the same magnitude of variations and trends In Lhasa, but not in Golmud and Huangzhong (Figure S15 in Supporting Information S1). These differences can be attributed to the unrepresentative single well, mismatched resolutions, or unreasonable specific yield estimations (Shamsudduha et al., 2012).

4.4. Inconsistencies in GWS Changes Between Our Study and Global Models

Figure 3a shows that there are varying degrees of differences in trends between G-GWS and other GHMs or LSMs in the TP during 2003–2012. We found that almost all existing GHMs and LSMs could not simulate the

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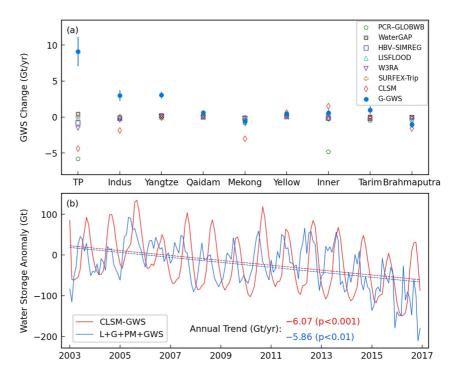


Figure 3. s(a) Comparisons of trends in Groundwater Storage (GWS) with five Global Hydrological Models (GHMs) and two Land Surface Models (LSMs) in the TP and its sub-basins during 2003–2012 (Error bars represent the corresponding uncertainty of GWS trends; information on GHMs and LSM are in Table S10; Inner represents Inner Plateau). (b) Comparisons of GWS anomalies from Catchment Land Surface Model (CLSM-GWS) with the summation of L, G, PM, and G-GWS in the entire TP during 2003–2016 (Red and blue dashed lines are the linear trend of CLSM-GWS and the summation of L, G, PM, and G-GWS, respectively; annual trends and significant levels are also shown).

substantially increasing GWS as G-GWS did (Figure 3a), particularly in the Yangtze and Indus-Ganges with large inconsistencies.

Many factors, including groundwater simulation schemes, incomplete water storage modules, groundwater depth, groundwater recharge formulations, forcing datasets, etc., could account for these inconsistencies. In general, these models used simplified water balance approaches to model GWS changes but did not simulate groundwater fluxes or heads directly, thus may result in inaccurate GWS simulations (Condon et al., 2021; Gleeson et al., 2021). Besides, these inconsistencies may be related to groundwater depths in the TP. Groundwater flow systems are characterized as deep circulation driven by topographic gradients that can reach deeper than 1 ~ 2 km in the TP (Ge et al., 2008), while these models only model shallow groundwater. Groundwater recharge schemes in these models are also important for interpreting these inconsistencies because most models only incorporate diffuse groundwater recharge formulation (Condon et al., 2021). Water-Global Assessment and Prognosis version 2.2 d (WaterGAP v2.2 d) and PCRaster Global Water Balance 2 (PCR-GLOBWB 2) improve it by considering recharge from surface water (Müller Schmied et al., 2021; Sutanudjaja et al., 2018), however, recharge from glaciers, snow, or permafrost is still not represented well in them. Therefore, the absence of lake, glacier, or permafrost modules in these models may largely cause these inconsistencies.

The Catchment Land Surface Model (CLSM) from the GLDAS version 2.2 (GLDAS-2.2) products show quite improvements in groundwater simulation as GLDAS-2.2 particularly assimilates TWS anomalies from the GRACE missions (Li, Rodell, et al., 2019). However, GWS from CLSM (CLSM-GWS) could not capture the increasing GWS affected by climate change in the TP. CLSM computes GWS changes by deducting contributions of other components (SM, SWE, and Canopy Interception) from TWS changes (Li, Rodell, et al., 2019), while the contributions of glaciers, permafrost, and lakes are not separated. Consequently, CLSM-GWS implicitly contains water storage changes in glaciers, permafrost, and lakes over the TP as Figure 3b indicates almost identical trends between CLSM-GWS (-6.07 Gt/yr) and the summation of the lake, glacier, permafrost, and GWS (-5.86 Gt/yr) in 2003–2016.

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4.5. Limitations and Uncertainties

One major limitation of this study is the accumulation of intrinsic uncertainties of each water storage component in the water balance approach. Uncertainties from soil moisture exert the largest effects on GWS estimations as using different products may result in opposite GWS trends (Figure S8 in Supporting Information S1). Although we used available in-situ measurements to evaluate soil moisture products, simulations and observations on soil moisture need to be further improved in the TP. Another limitation is the residual closure errors in GWS due to ignoring mass change induced by potential factors including tectonic uplift, soil erosion, plant canopy, biomass changes, water storage changes in the river channel, etc., in the water balance equation. Previous studies concluded that the effects of these factors can be ignored, while Ke et al. (2022) revealed that water storage changes in river channels could account for around 10% of the seasonal GRACE signals in the Upper Brahmaputra River Basin. Although water storage changes in river channels have limited effects on the long-term TWS trends, including them could improve estimations of seasonal GWS variations. In addition, datasets with coarse spatio-temporal resolution as well as simple assumptions and aquifer parameters in the independent GWS estimations from baseflow recession and well can also lead to uncertainties. Moreover, we have to acknowledge that no field observations could validate increasing GWS in the TP although we found this phenomenon in indirect ways. More efforts on observations and simulations will be needed to better understand the impacts of climate change on groundwater in the TP.

5. Conclusions

We combined multi-source data including in-situ measurements, satellite observations, global models, and reanalysis products, finding a substantial increase in GWS (5.59 ± 1.44 Gt/yr) in the whole TP from 2003 to 2016. The GWS changes showed spatial heterogeneities among the sub-basins with declining trends in the Inner Plateau, Brahmaputra, and Mekong-Salween, and increasing trends in other basins. Given the decreasing precipitation and no substantial changes in other potential recharge sources except for declining solid water (-17.72 ± 1.53 Gt/yr), the total GWS increase was likely dominated by groundwater recharge from solid water melt. The GWS contributes $19.28 \pm 0.88\%$ of the TWS variations in the entire TP and its contributions range from 19.64 ± 0.96 to $45.08 \pm 1.60\%$ in its sub-basins. G-GWS shows consistencies in trends with independent estimations of B-GWS, while almost none of the existing LSMs and GHMs could simulate the substantially increasing GWS in the TP due to the lack of cryospheric or lake modules and simplified representations of groundwater. This study identifies the dominant factor controlling long-term GWS changes and provides a better understanding of GWS response to climate change in the data-scarce TP.

Data Availability Statement

Basin boundaries from Zhang (2019) were provided by the National Tibetan Plateau Data Center (https://doi.org/10.11888/BaseGeography.tpe.249465.file). JPL-M products were provided by the Physical Oceanography Distributed Active Archive Center (https://doi.org/10.5067/TEMSC-3JC62). CSR-M products were provided by the Center for Space Research (http://www2.csr.utexas.edu/grace). Lake volume changes from Li, Long, et al. (2019) were provided by PANGAEA (https://doi.org/10.1594/PANGAEA.898411). Lake volume changes from Wang, Wang, Li, et al. (2021) were provided by Zenodo (https://doi.org/10.5281/zenodo.5543615). Lake volume changes from Zhang, Bolch, et al. (2021) were provided by National Tibetan Plateau Data Center (http://doi.org/10.11888/Hydro.tpdc.271169). Glacier mass changes from Hugonnet et al. (2021) were provided by Theia (https://doi.org/10.6096/13). SWE from Kraaijenbrink et al. (2021) were provided by Zenodo (https://doi.org/10.5281/zenodo.4715786). ALT from CLM5 was provided by the Climate Data Gateway at National Center for Atmospheric Research (https://doi.org/10.5065/d6154fwh).

Acknowledgments

This research was supported by the National Natural Science Foundation of China (Grants No. 92047202 and No. 91747204) and High-level Special Funding of the Southern University of Science and Technology (Grants No. G02296302 and No. G02296402).

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