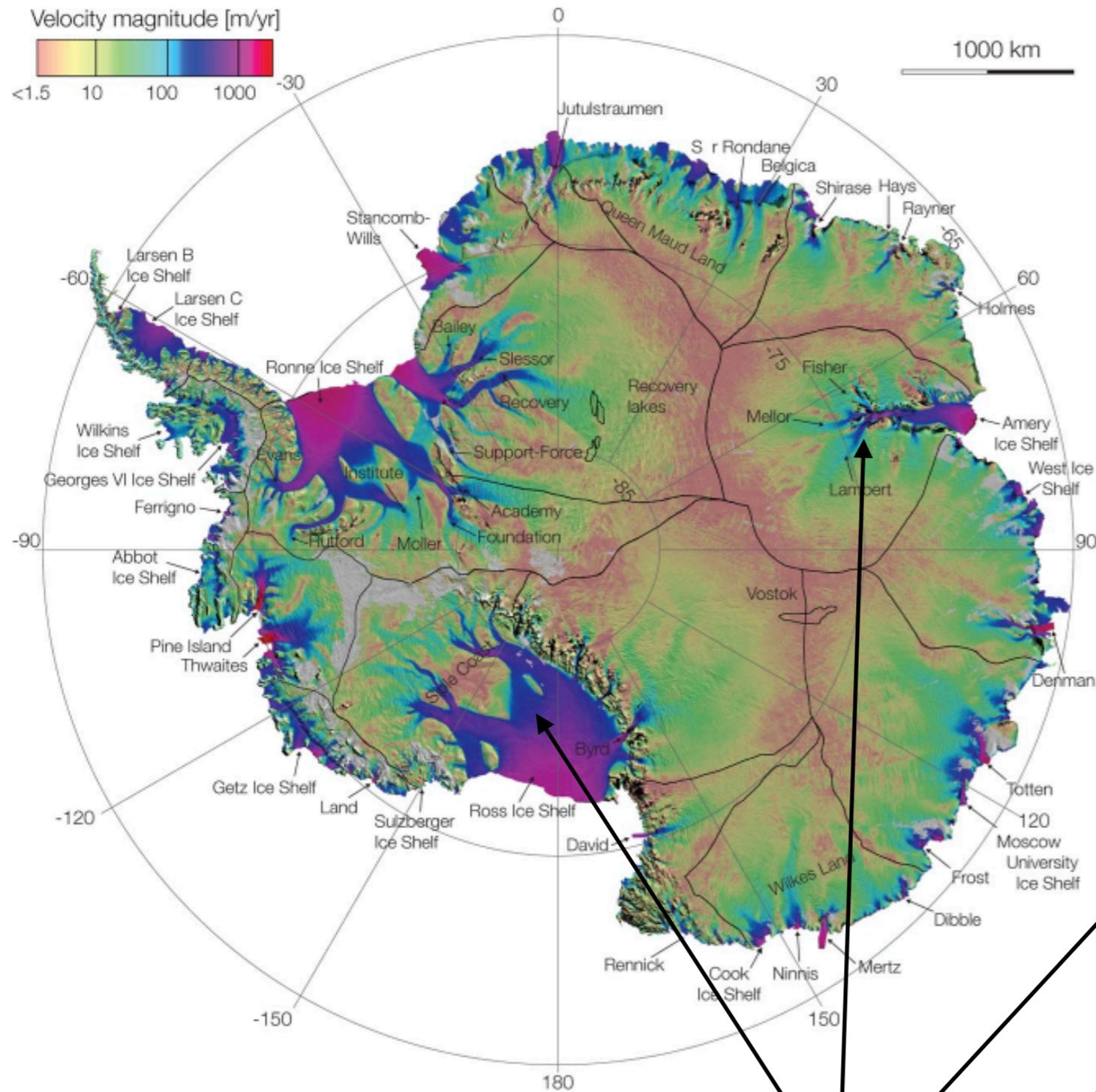


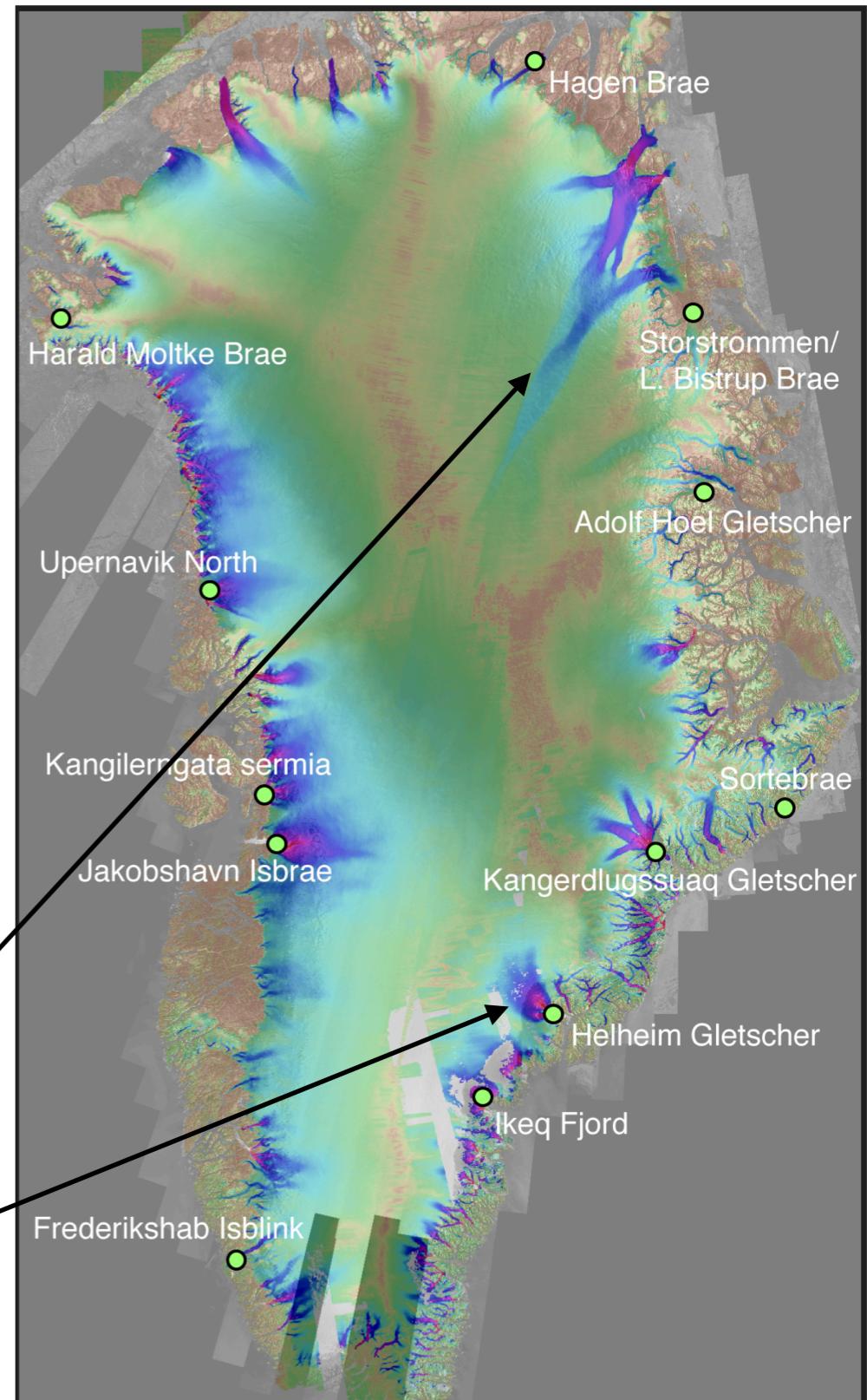
# **Ice Streams**

# Ice Sheet Velocities

Antarctic Ice Sheet

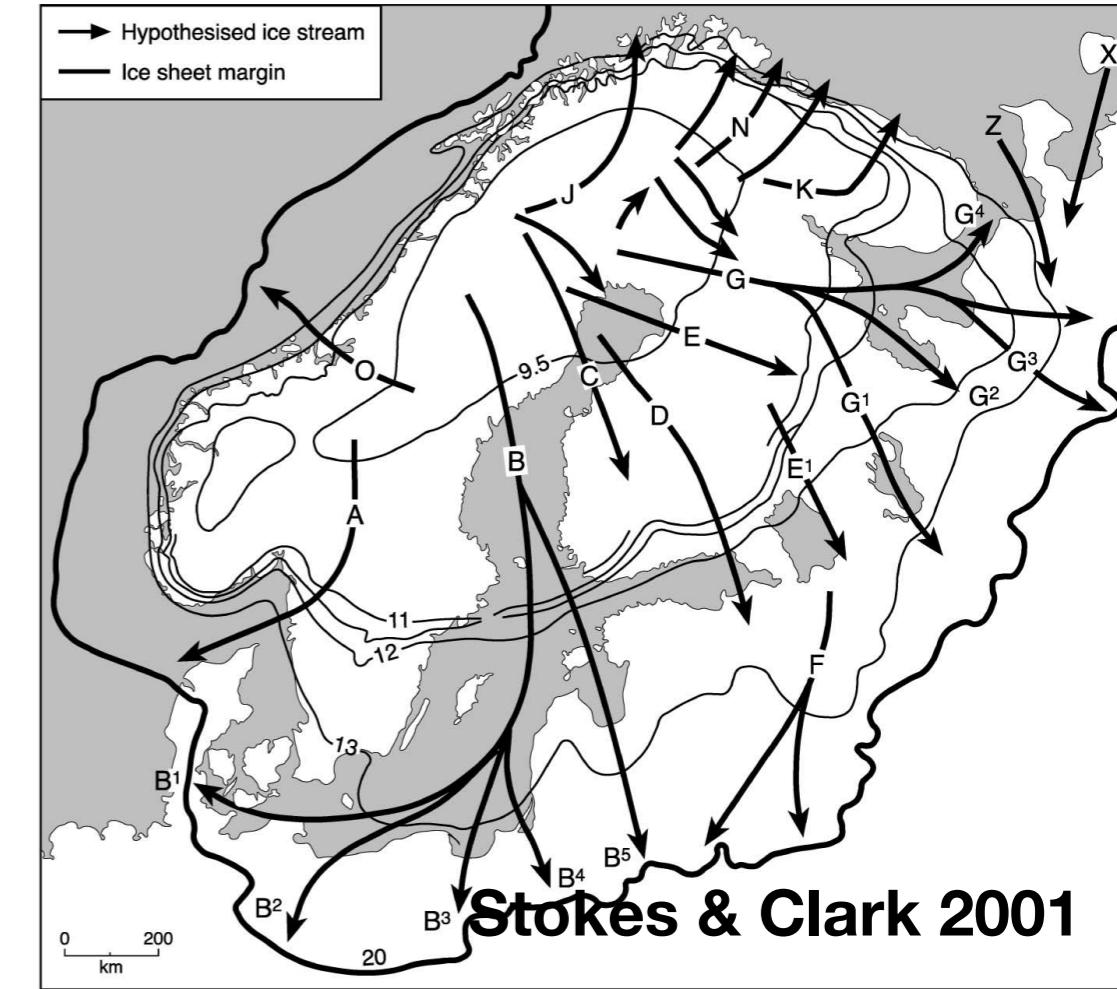
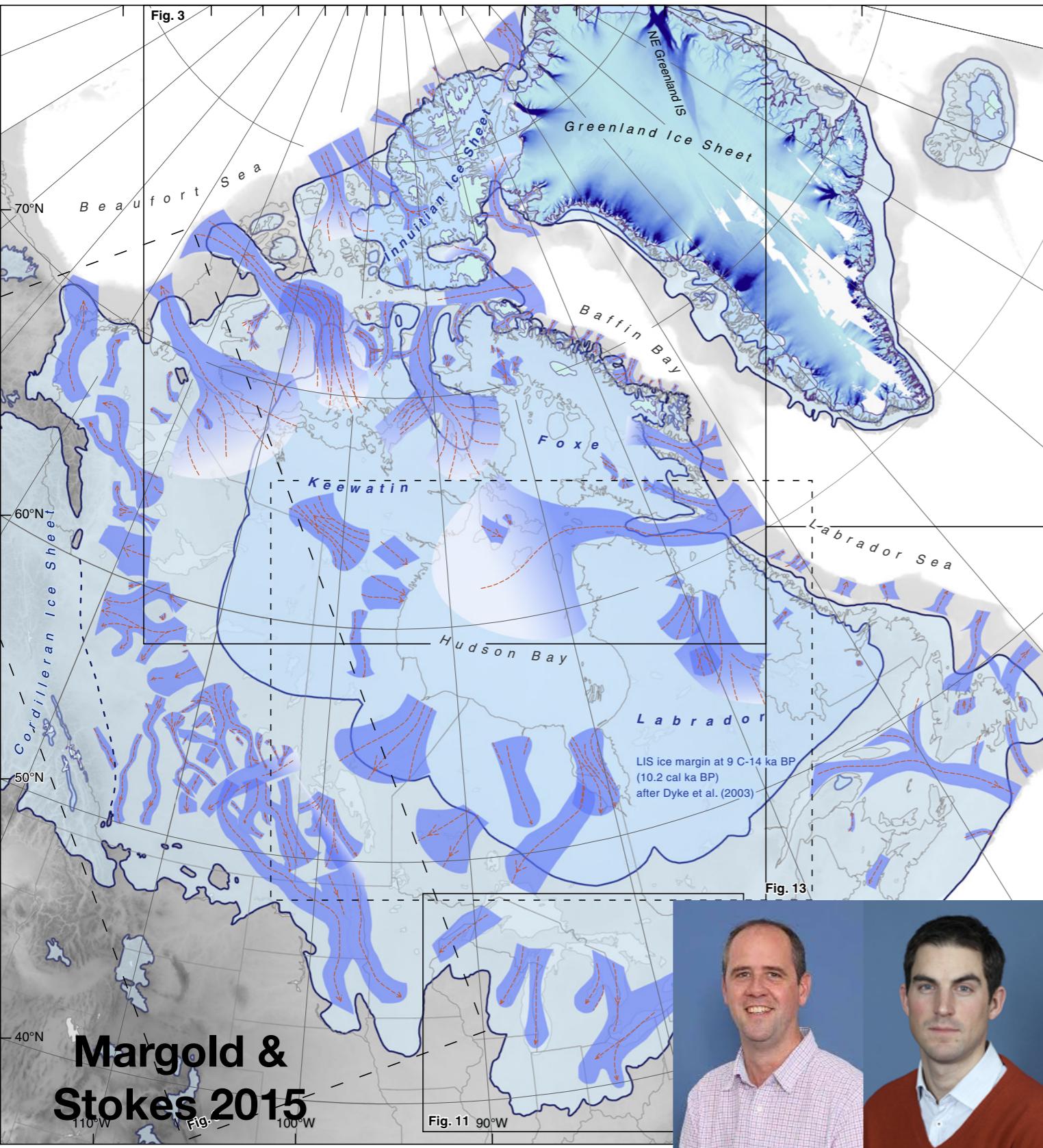


Greenland Ice Sheet



Most ice flow occurs through these.  
What are they? Why are they?

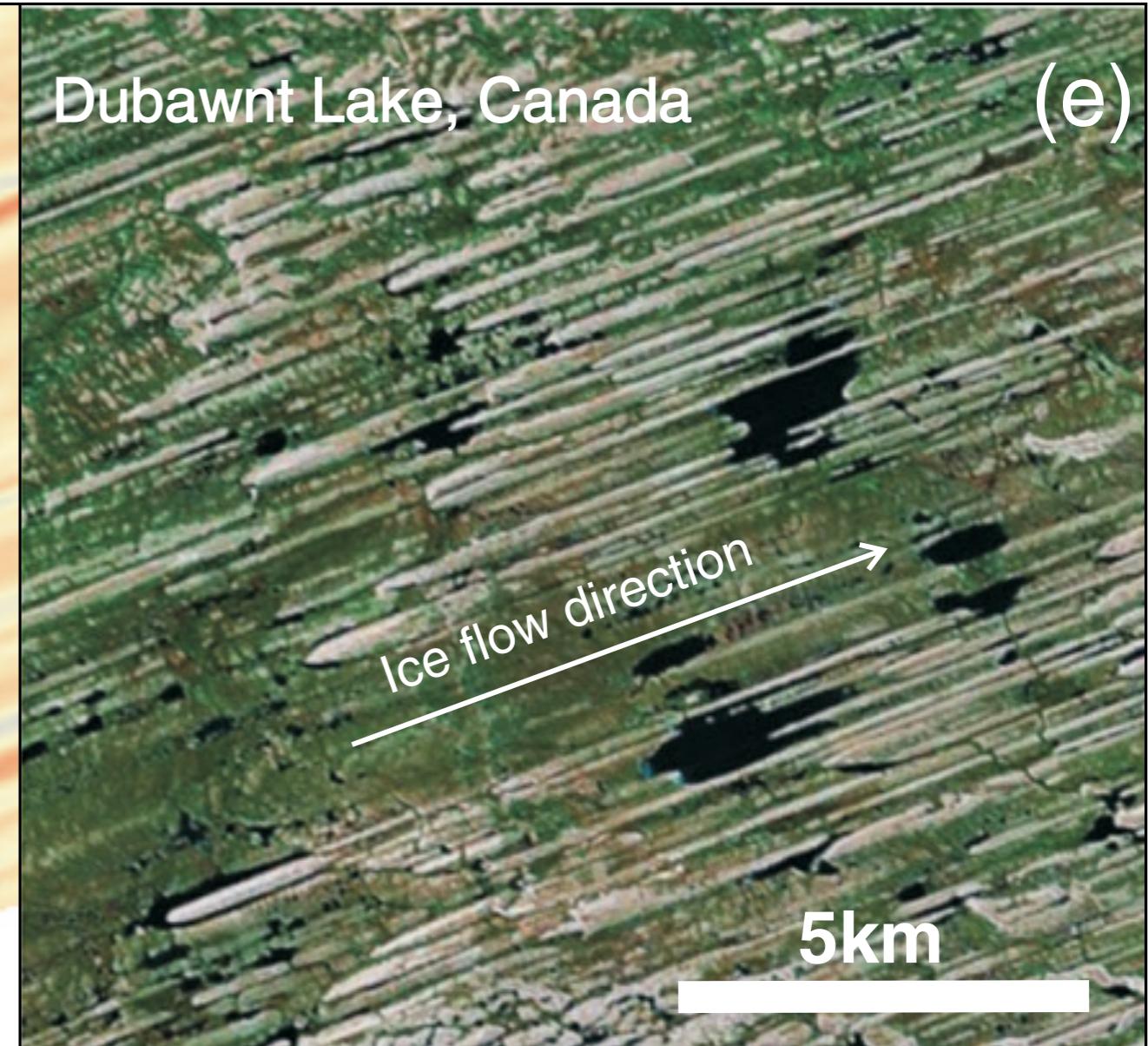
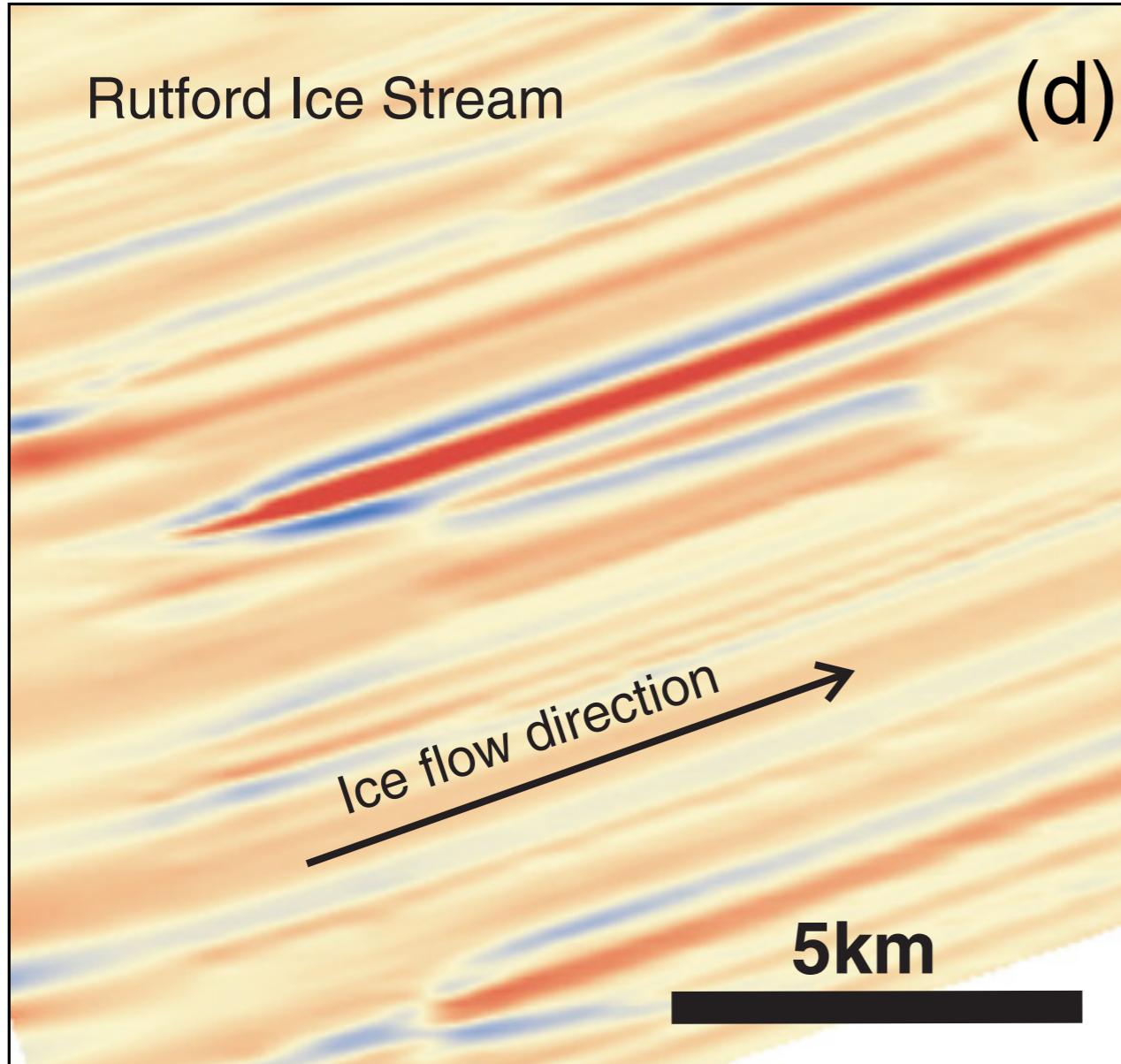
# Geological evidence for ice streaming in past ice sheets



There is also strong evidence that ice streams were pervasive in past ice sheets and were important to their overall mass transport as well



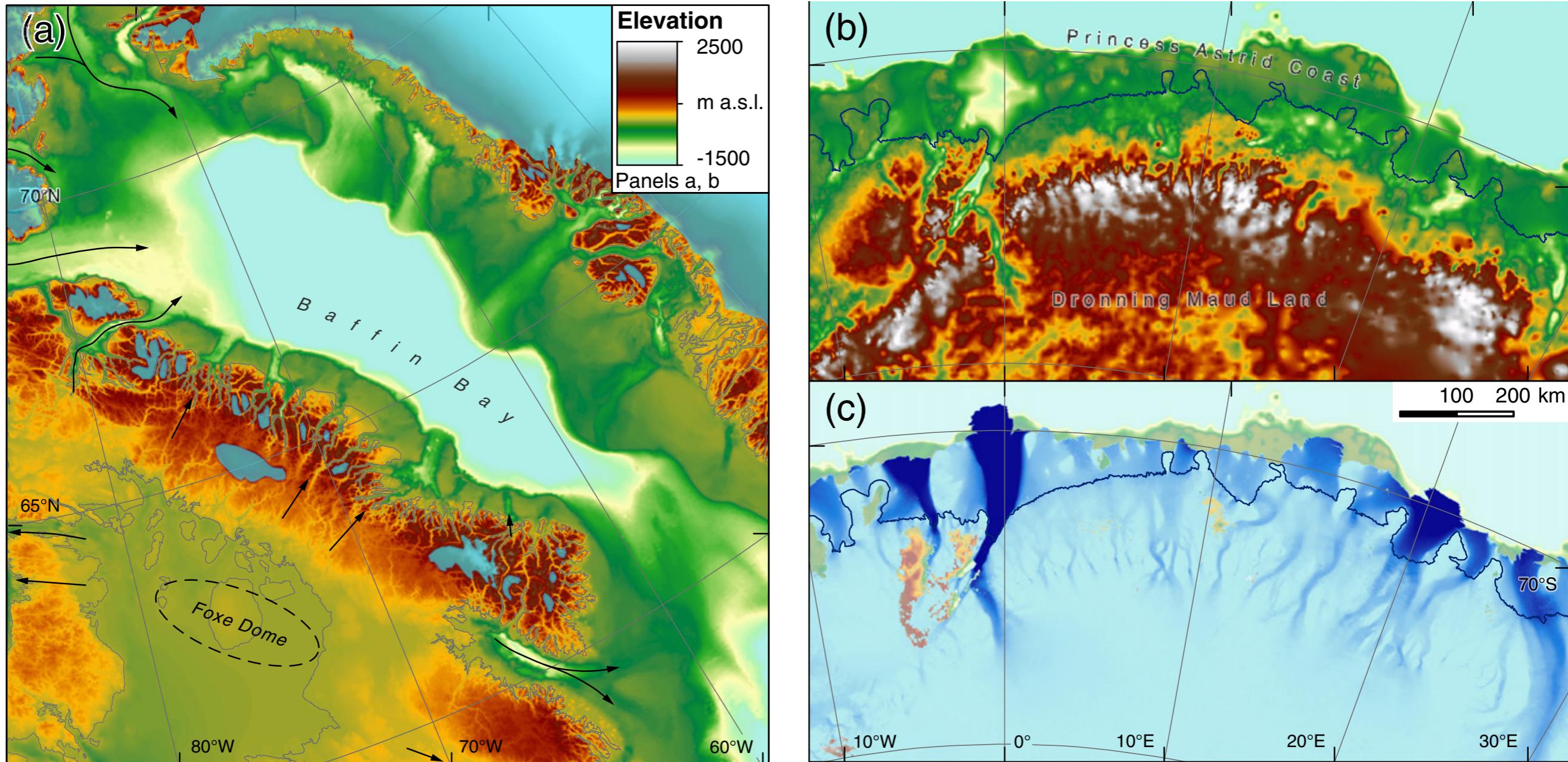
# Geological evidence for ice streaming in past ice sheets



Margold and Stokes 2015

**Megascale Glacial Lineations (MSG) indicate fast flow and direction of flow from past (and current) ice streams**

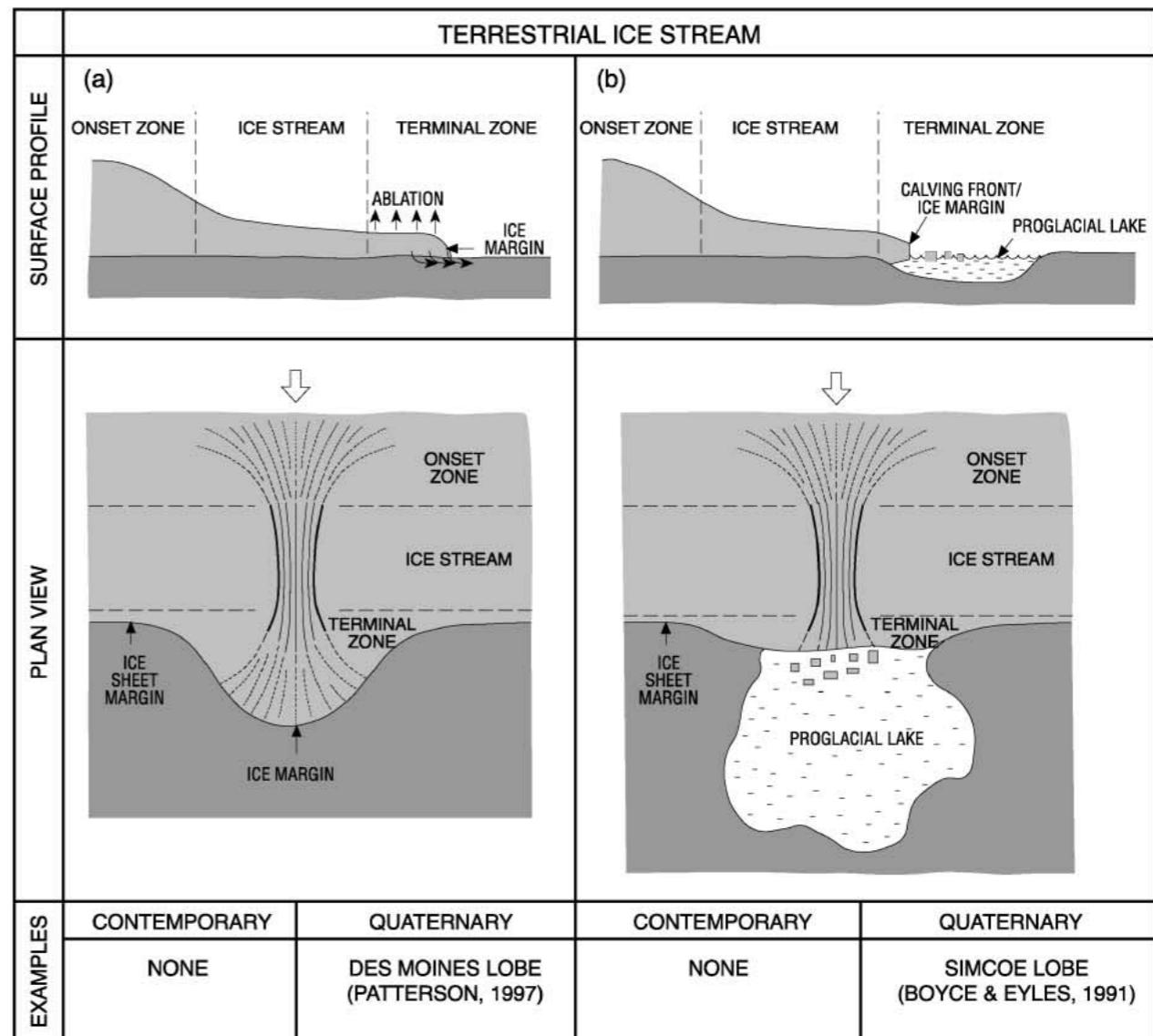
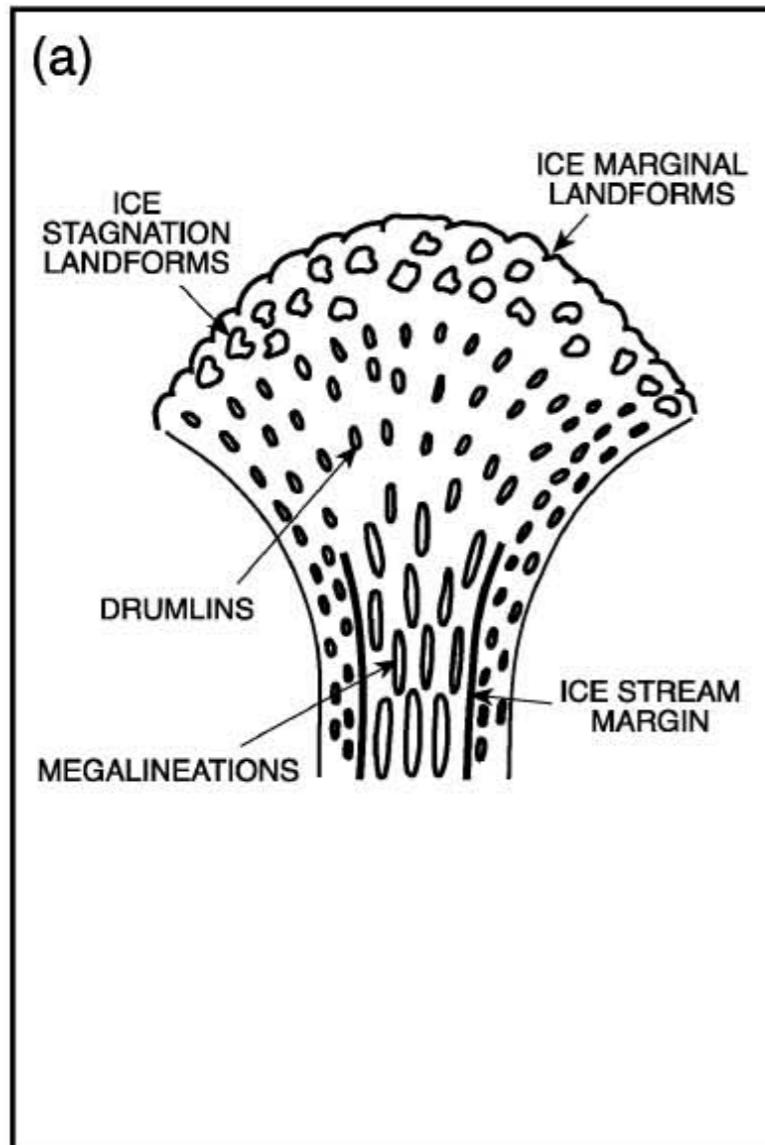
# Geological evidence for ice streaming in past ice sheets



Margold and Stokes 2015

Bathymetric troughs also indicate places where past ice streams have eroded into the bed

# Geological evidence for ice streaming in past ice sheets

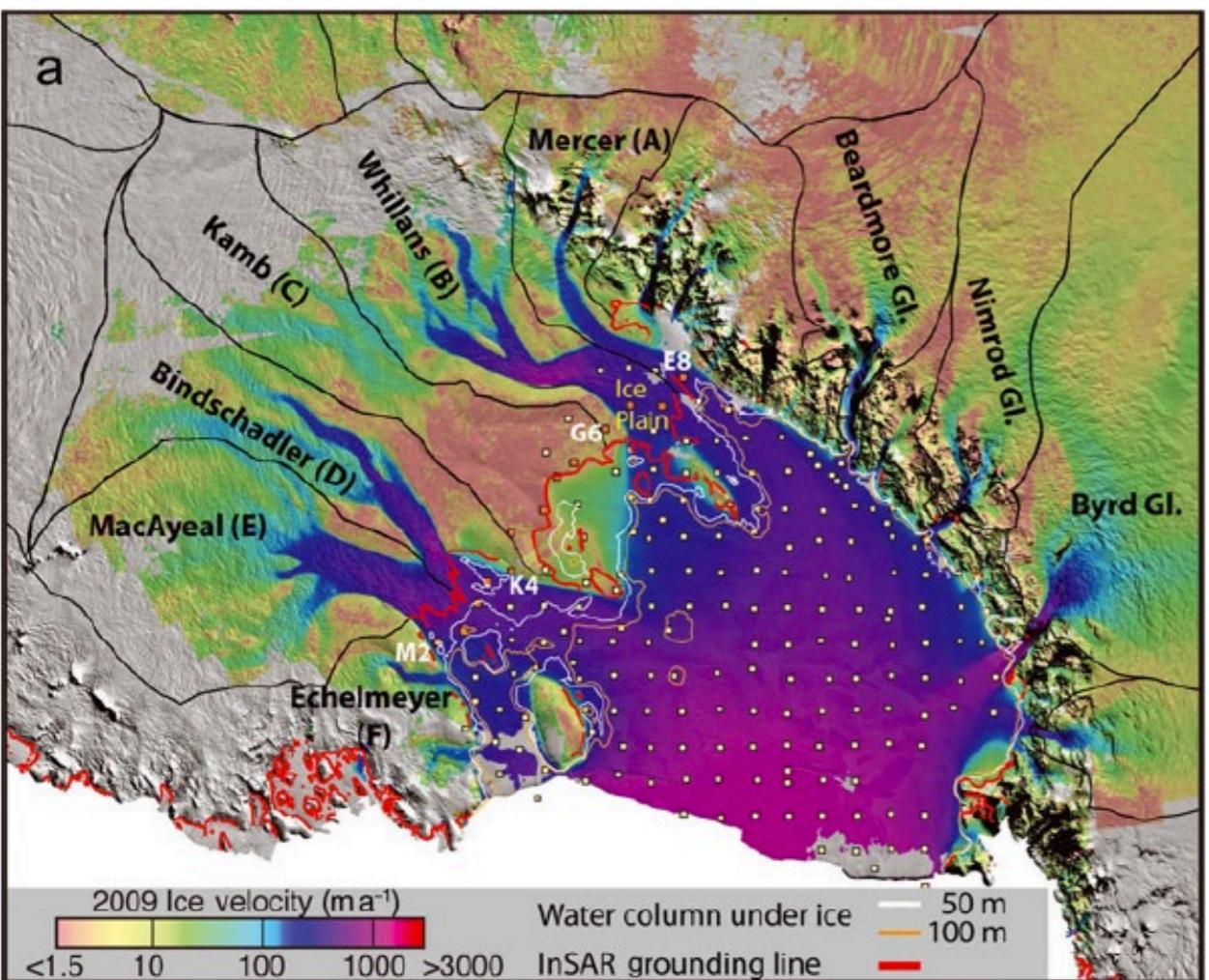


Stokes & Clark 2001

Moraines, drumlins, and other sedimentary deposits may indicate furthest extent of ice stream flow (since ice streams often terminate at different places than rest of margin)

# Terminology

“Ice stream” is used interchangeably with “outlet glacier” - refers to **any place in an ice sheet where high ice velocities (>100 m/yr) are embedded within slower ice velocities (<100 m/yr)**



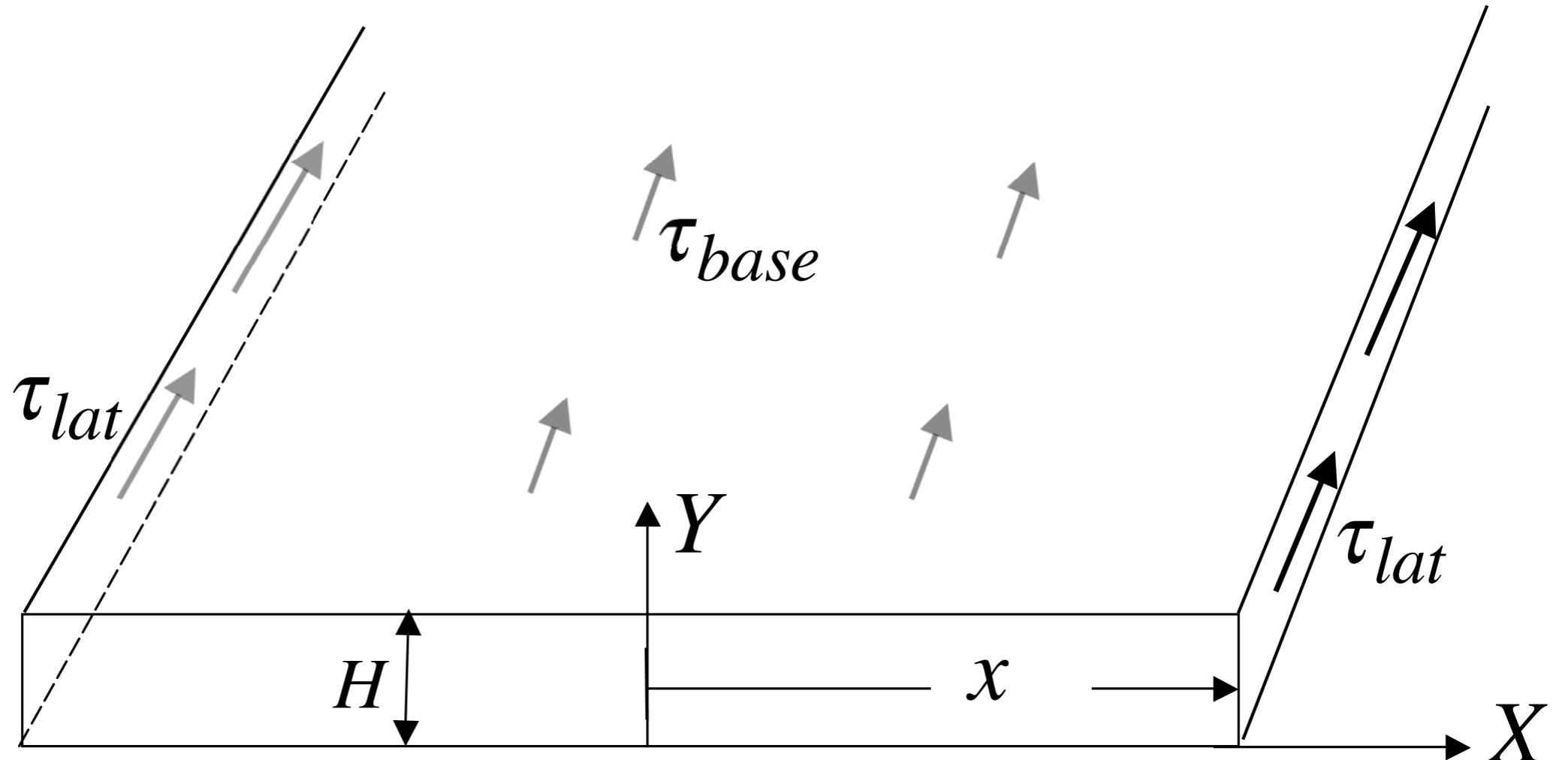
Ice streams are like the jet stream in the atmosphere or boundary currents in the ocean - they are “jets” of fast flow embedded within slow flow - typically indicates that something special is happening

# So...why do ice streams exist?

Group discussion (3 min): what processes can cause localized regions of fast ice flow?

(Think back to previous classes about what sets ice flow velocity - there is not a single right answer)

# The force balance of an ice stream



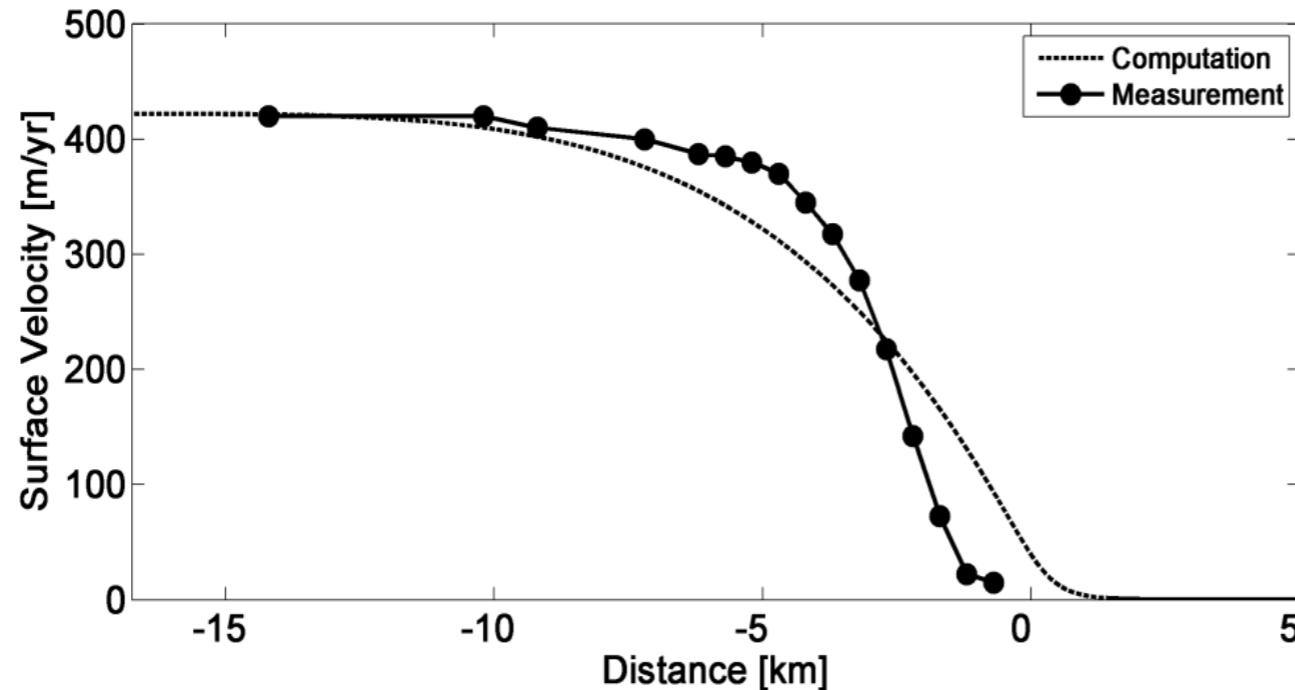
**Driving stress is balanced by friction at the bed and lateral friction on the ice stream sides from rock or slow ice**

# High ice stream velocities can be understood from a simple force balance

- To the board!

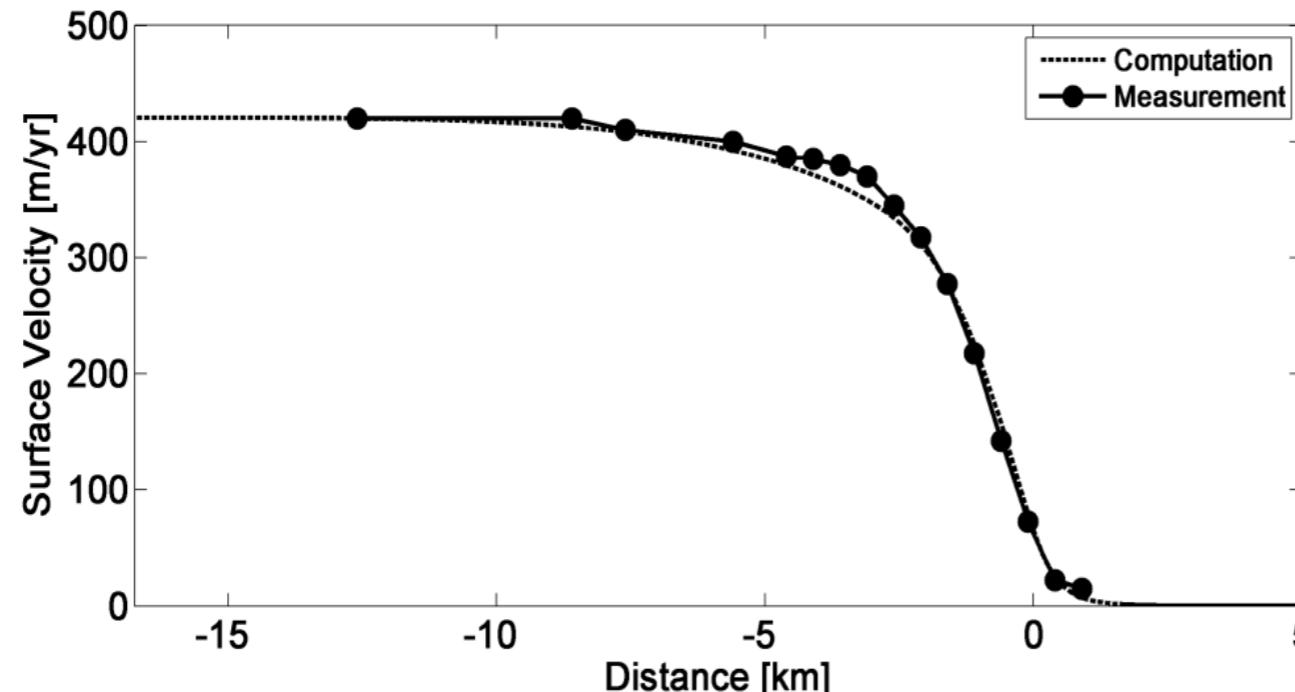
# Structure of shear margins

B2. Surface velocities (nonlinear, temperature-independent rheology)

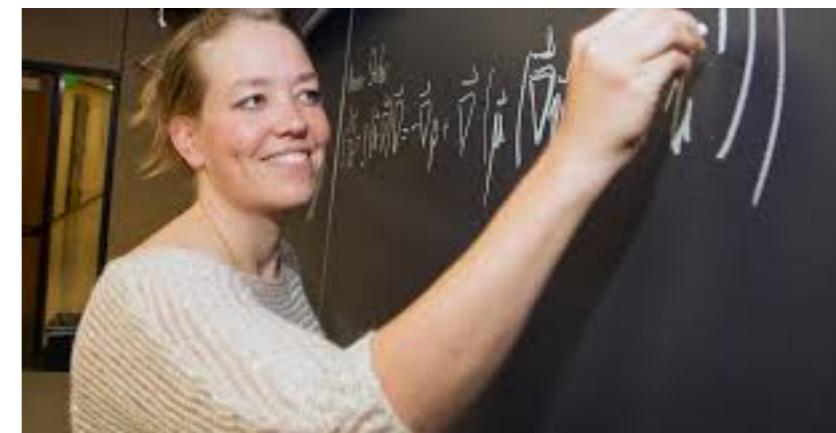


**Shear-dependent viscosity**

B3. Surface velocities (nonlinear, temperature-dependent rheology)

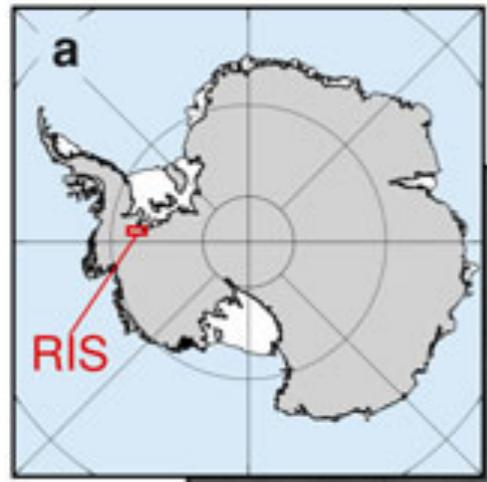


**Shear- and temperature-dependent viscosity  
(shear heating in the margin)**

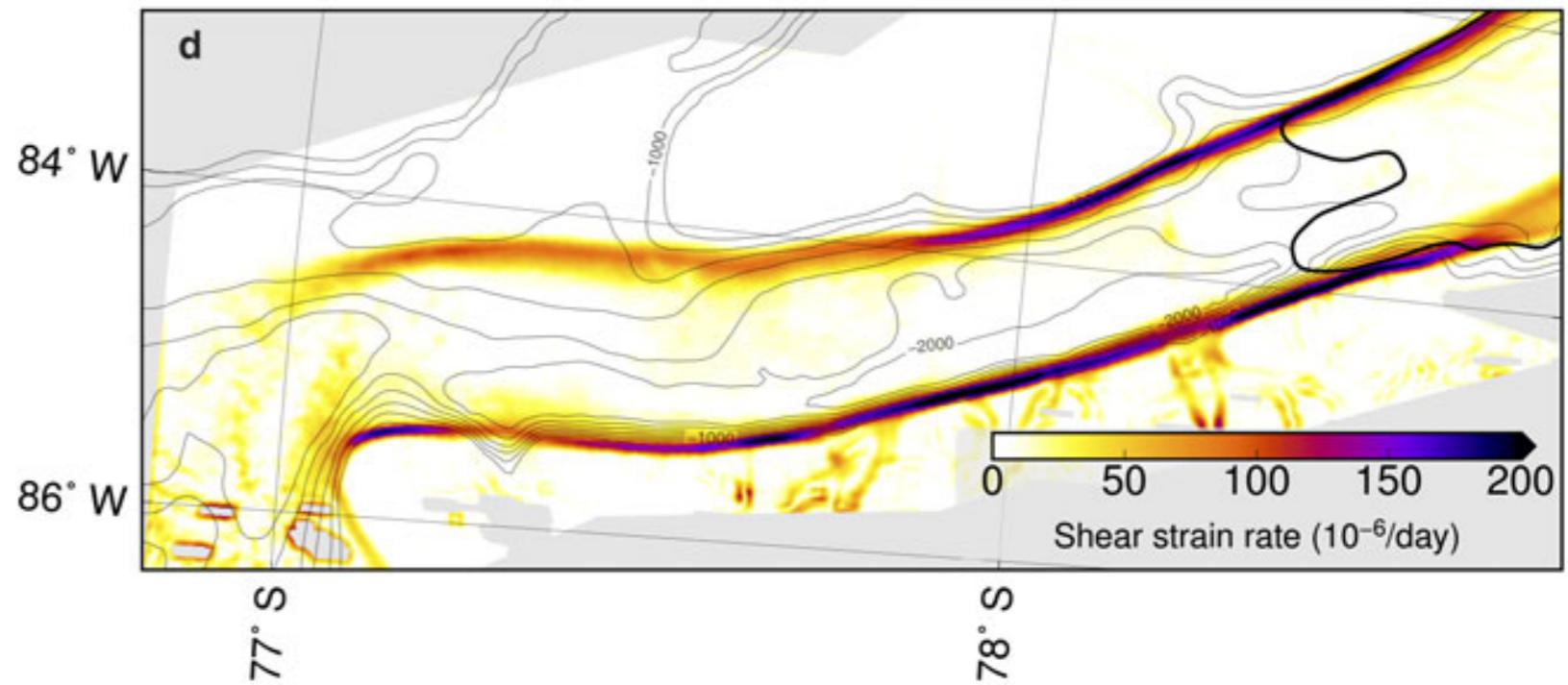
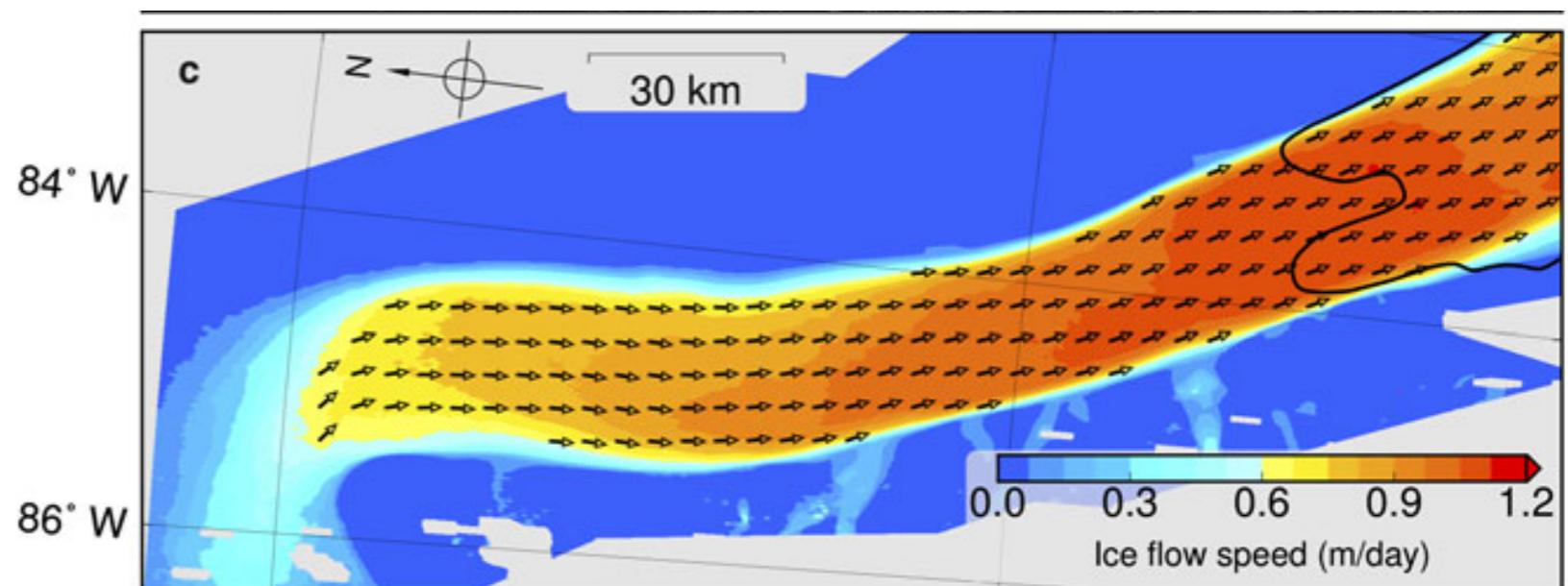


# Shear margins

Shear margins are the transition zones in between fast ice flow in the ice stream to slow ice flow outside the ice stream - very high shear stresses in shear margins

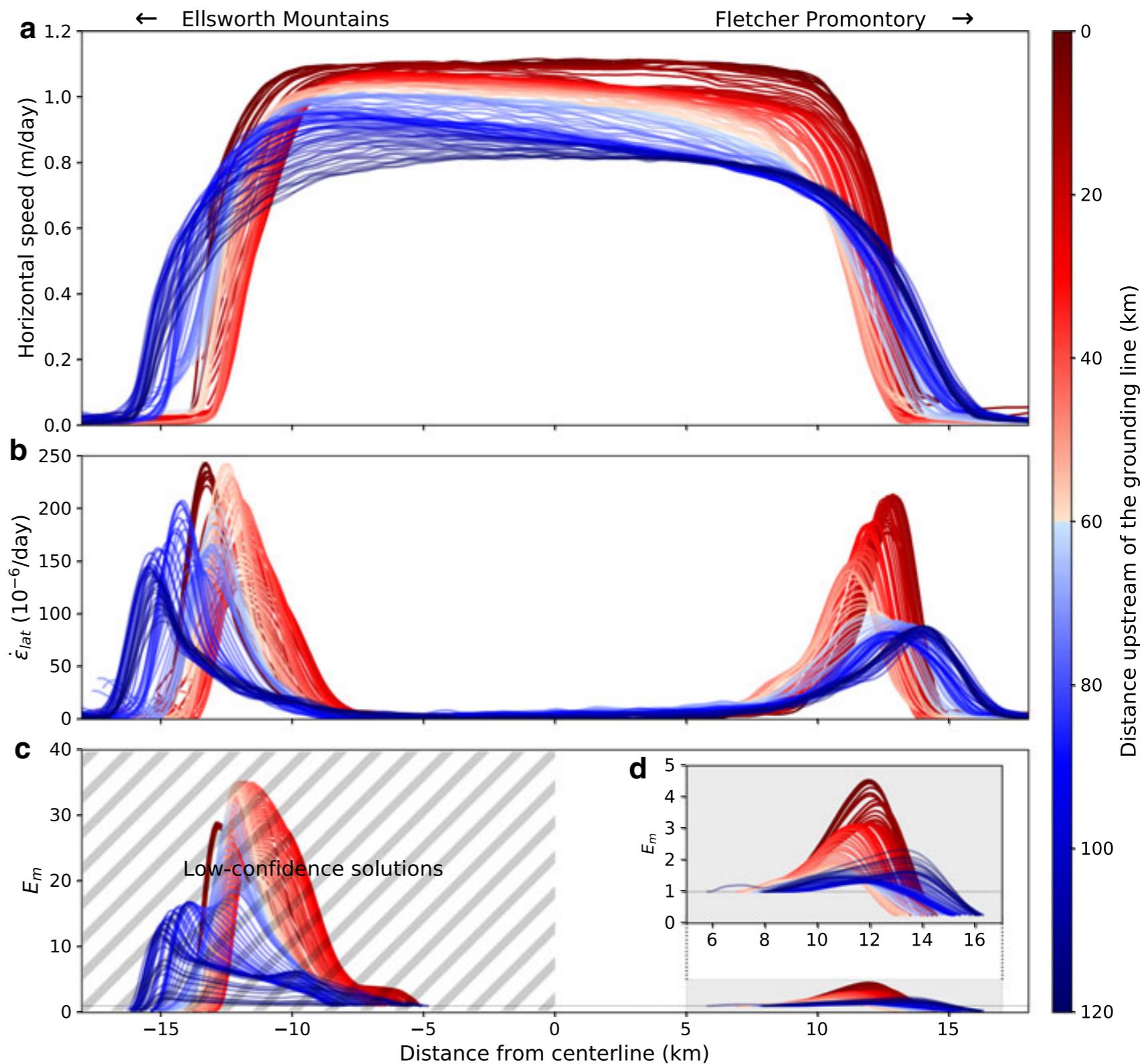


Minchew et al. 2018



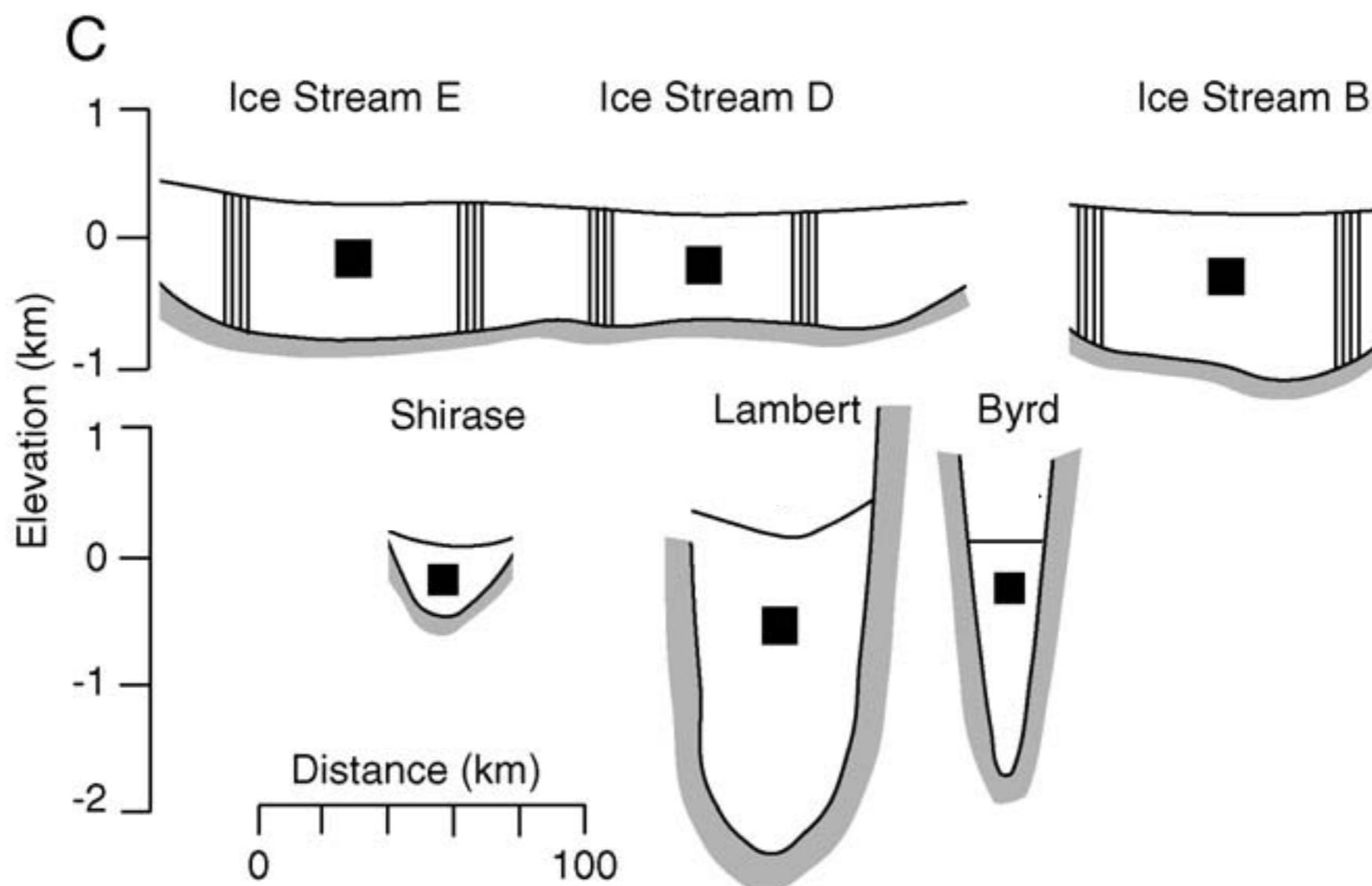
# Shear margins

Shear margins are the transition zones in between fast ice flow in the ice stream to slow ice flow outside the ice stream - very high shear stresses in shear margins - very low shear stresses in ice stream interior



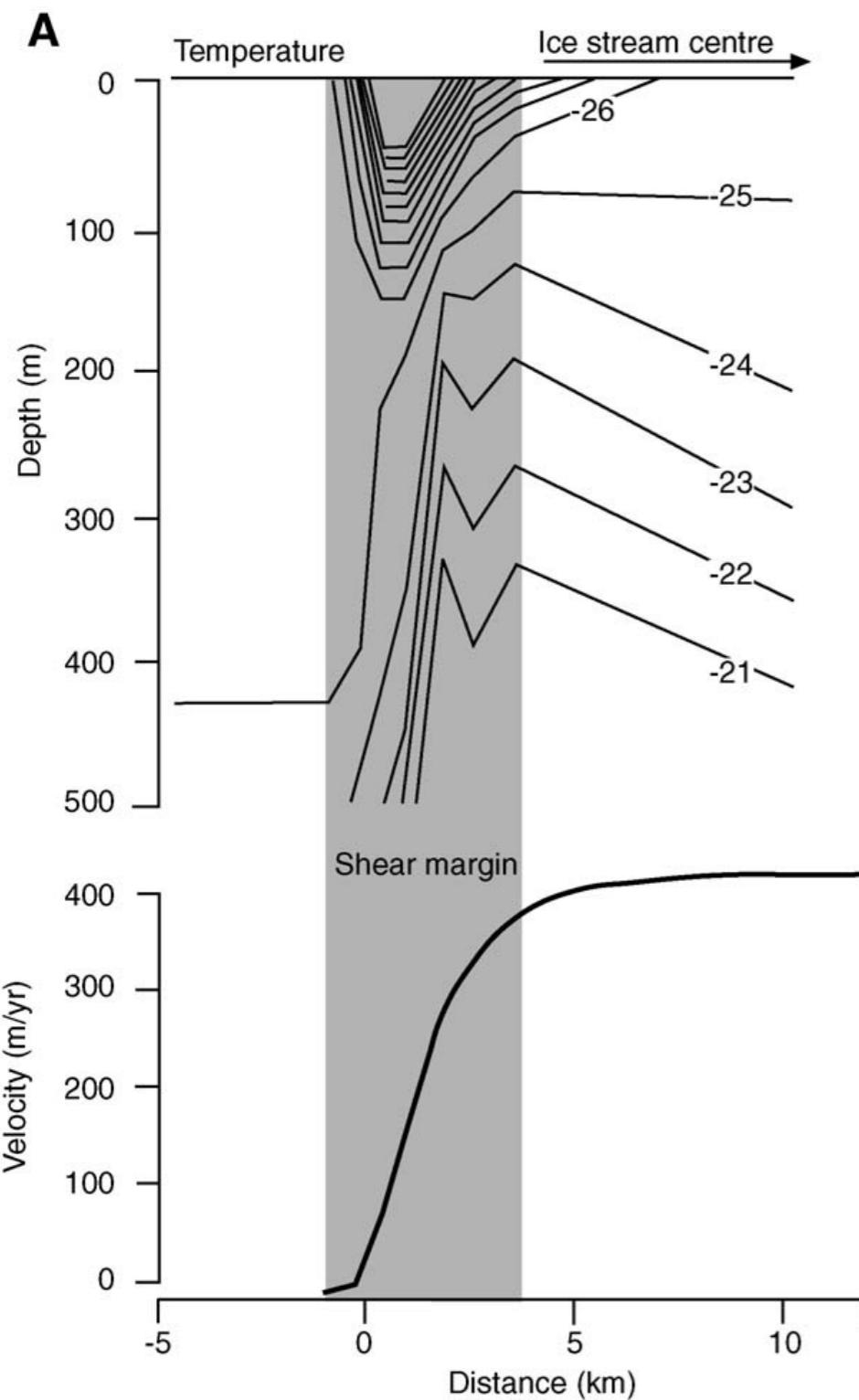
# Shear margins

Ice-ice or ice-rock



Bennett 2003

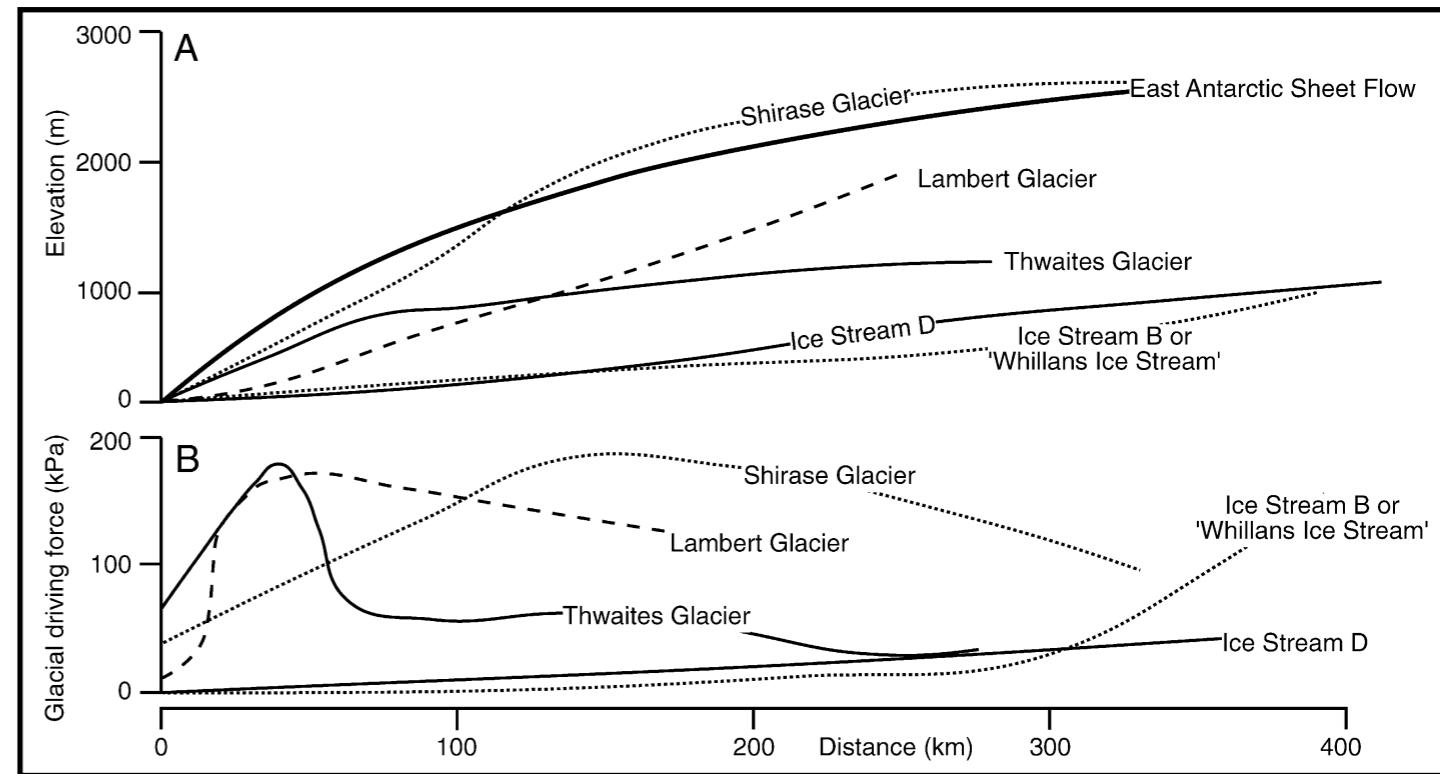
# Shear margins



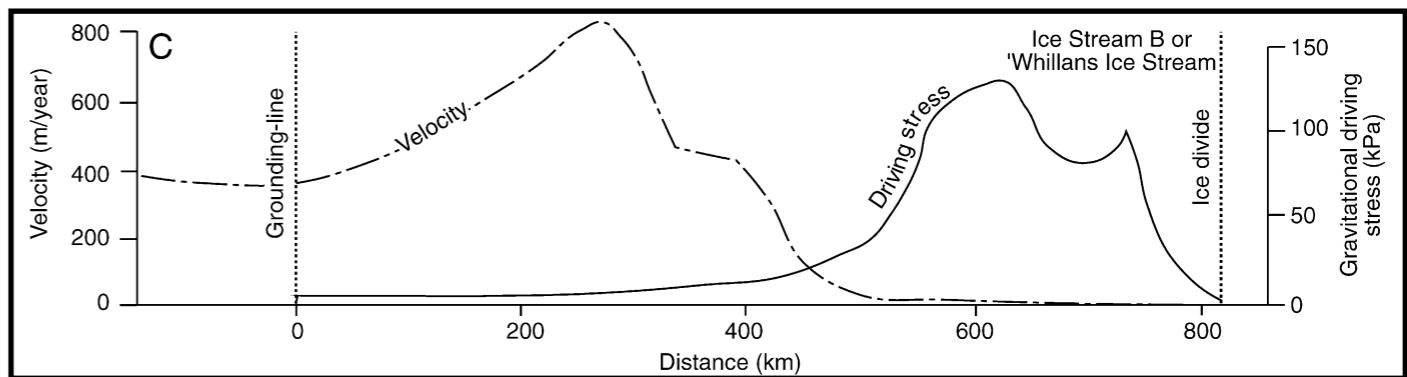
Shear margins also tend to be quite warm because shear causes heat production within ice (due to the rapid conversion of ice crystal bond dislocations to heat)

# Two types of ice streams

- **Topographic ice streams:** Gravitational driving stress is high (perhaps due to strongly sloping bed topography) - more common in Greenland and East Antarctica



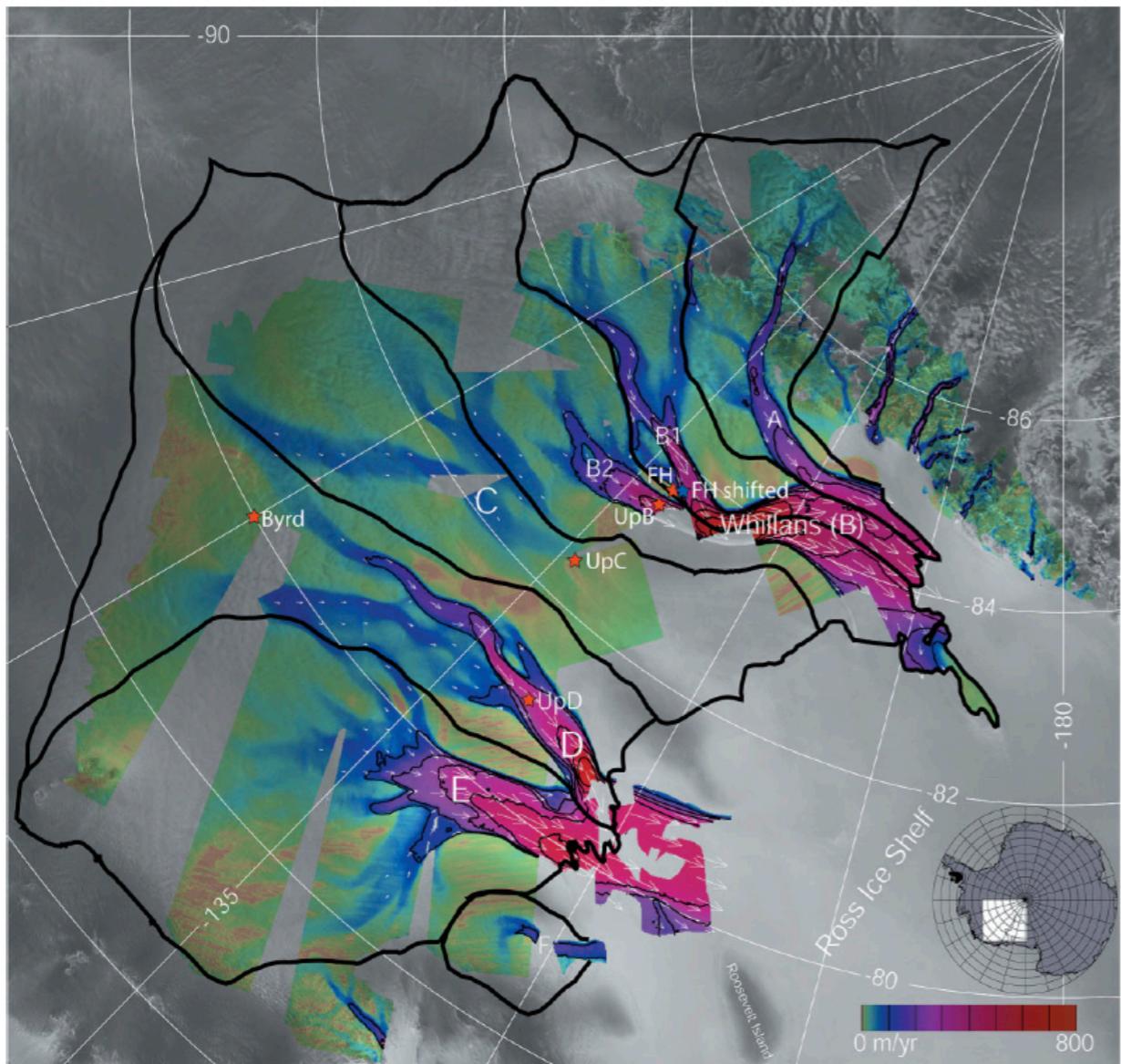
- **Pure ice streams:** the bed is very weak, driving stress is weak too - more common in West Antarctica



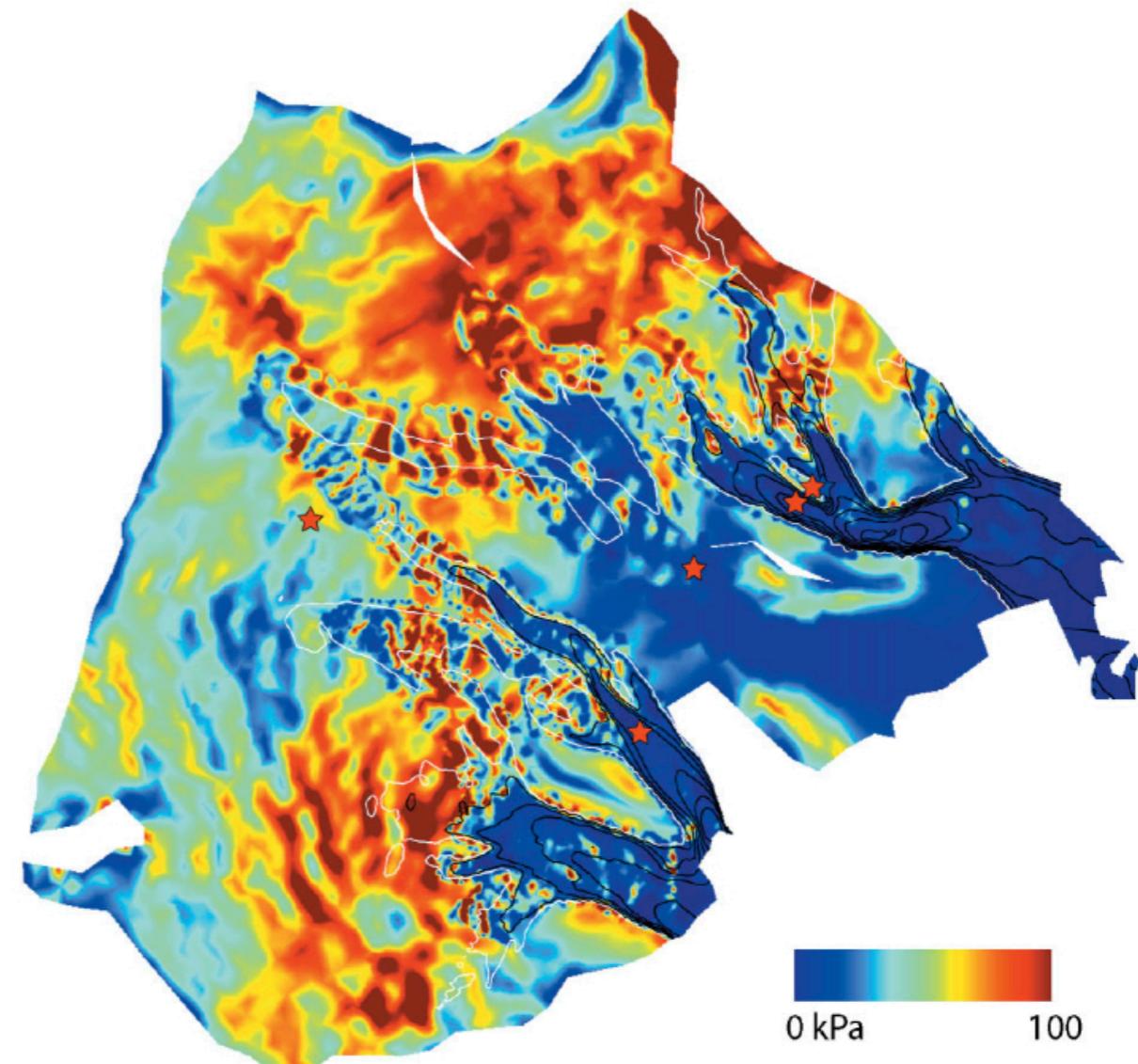
Bennett 2003

# So how does the bed get so weak in pure ice streams?

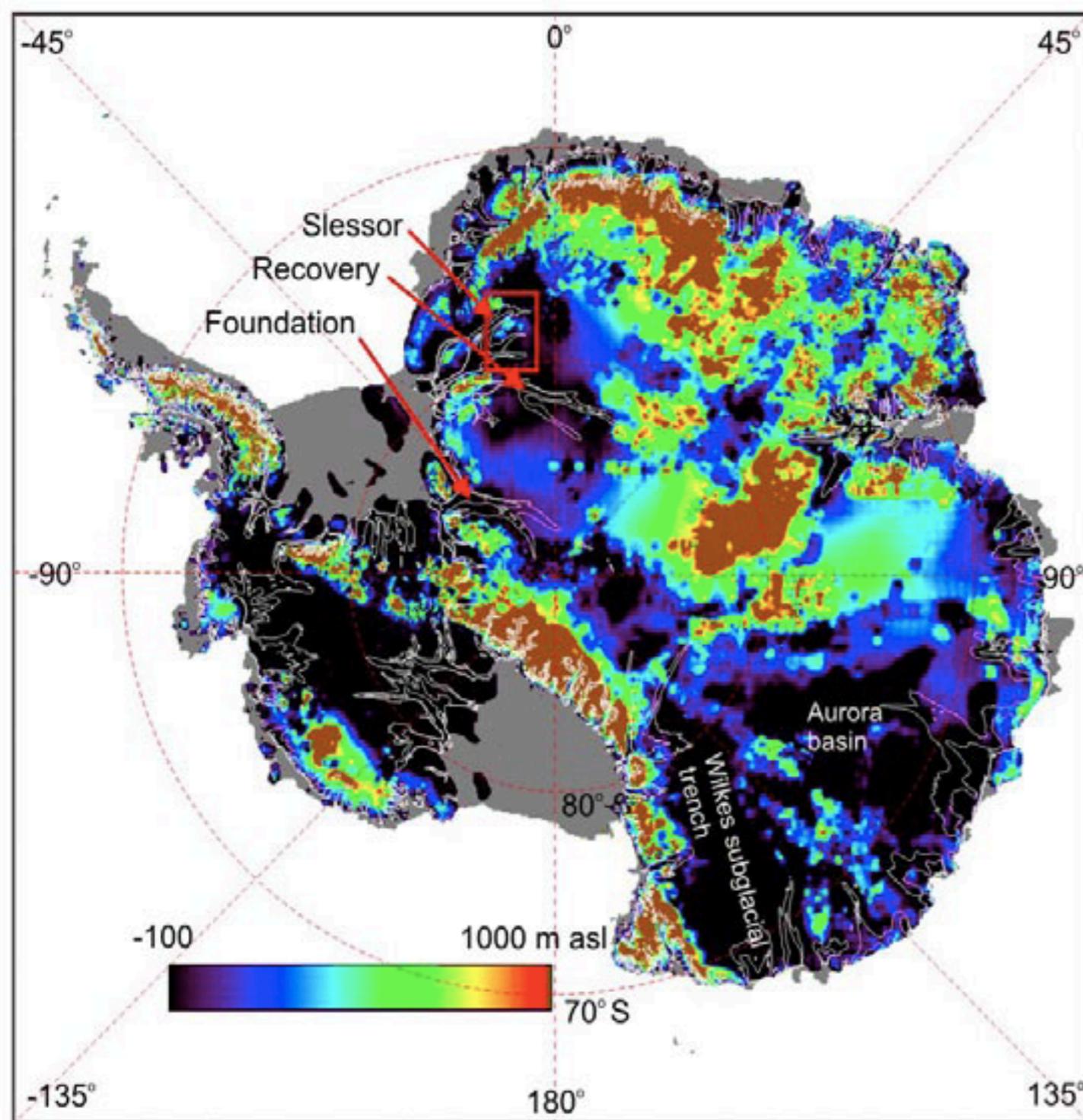
Siple Coast, W. Antarctica Velocities



Siple Coast, W. Antarctica Bed Friction

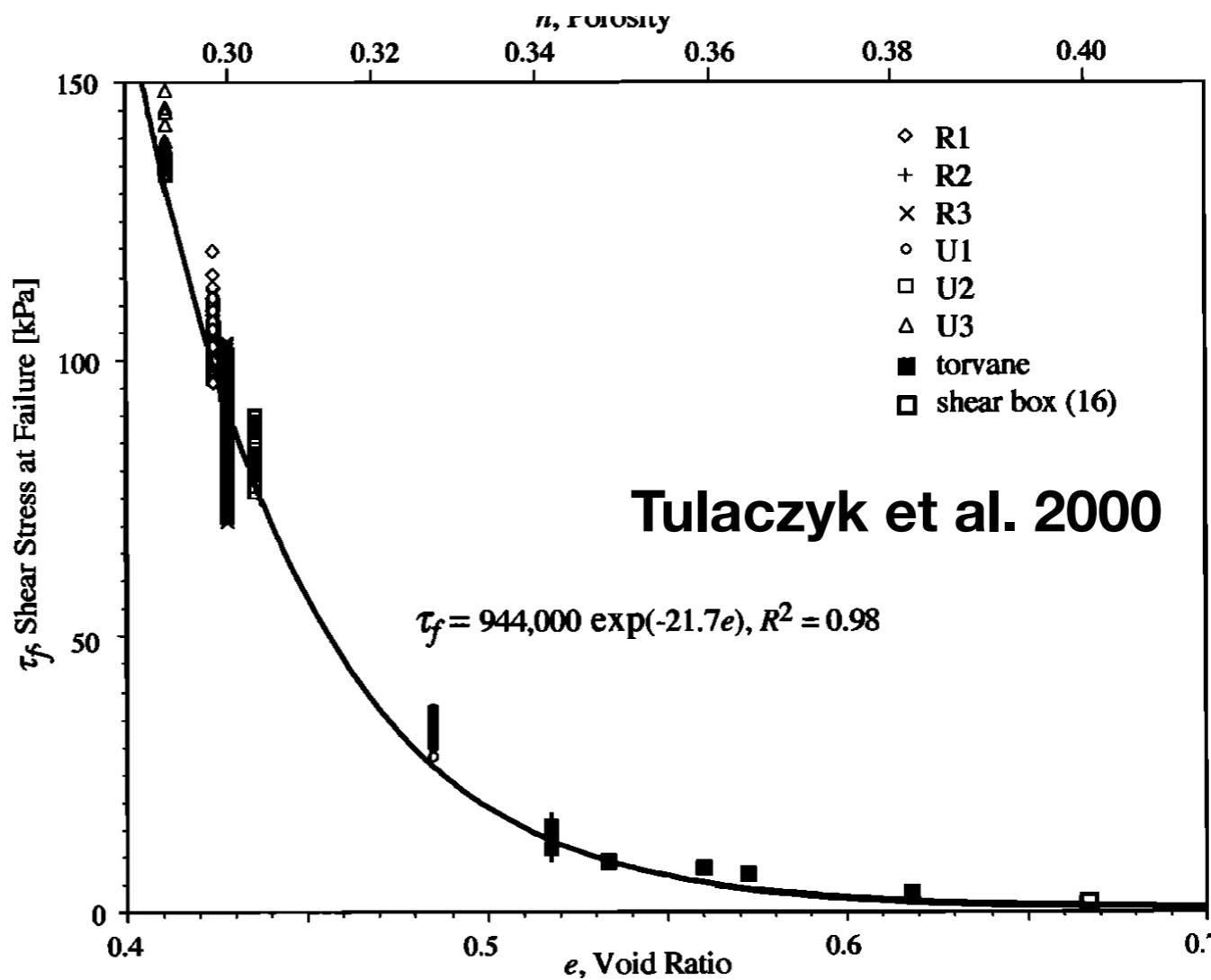
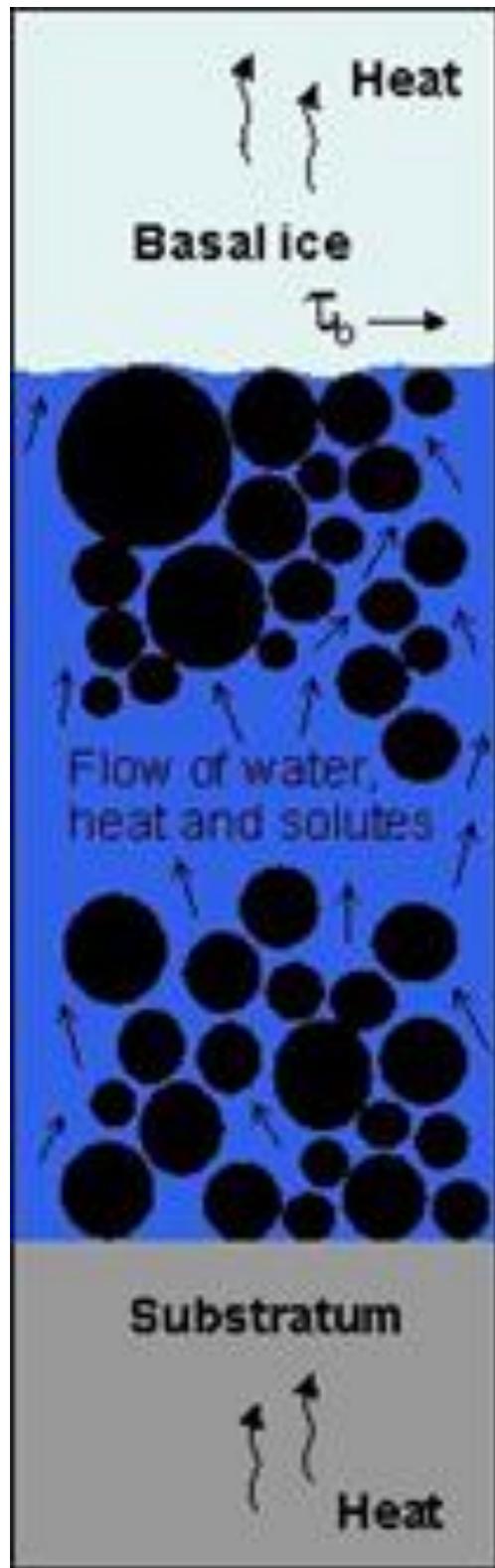


# So how does the bed get so weak in pure ice streams?



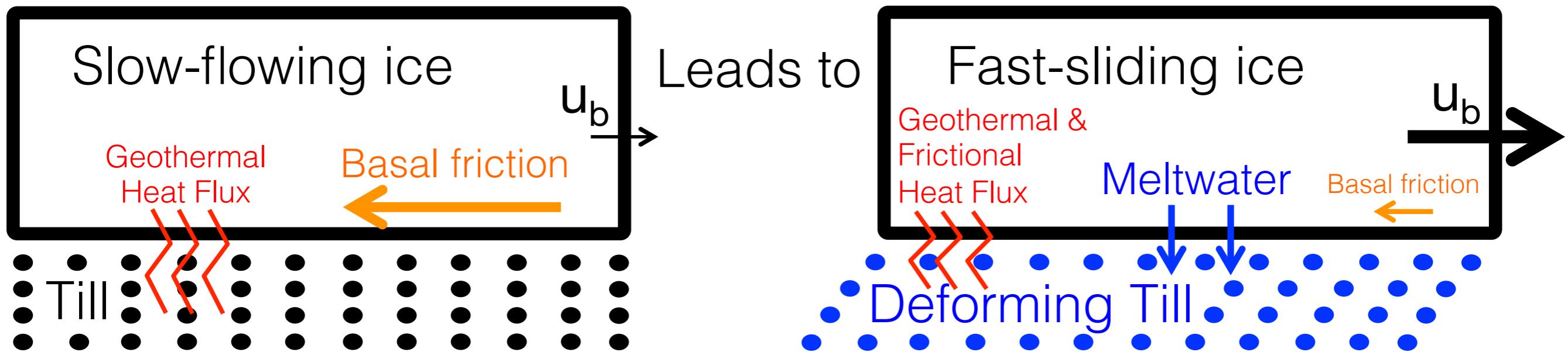
**Parts of Antarctic Ice Sheet that are below sea level (black on plot to the left) have been covered by ocean in the past when the ice sheet was much smaller and so there are lots of marine sediments**

# Till yield strength and water content



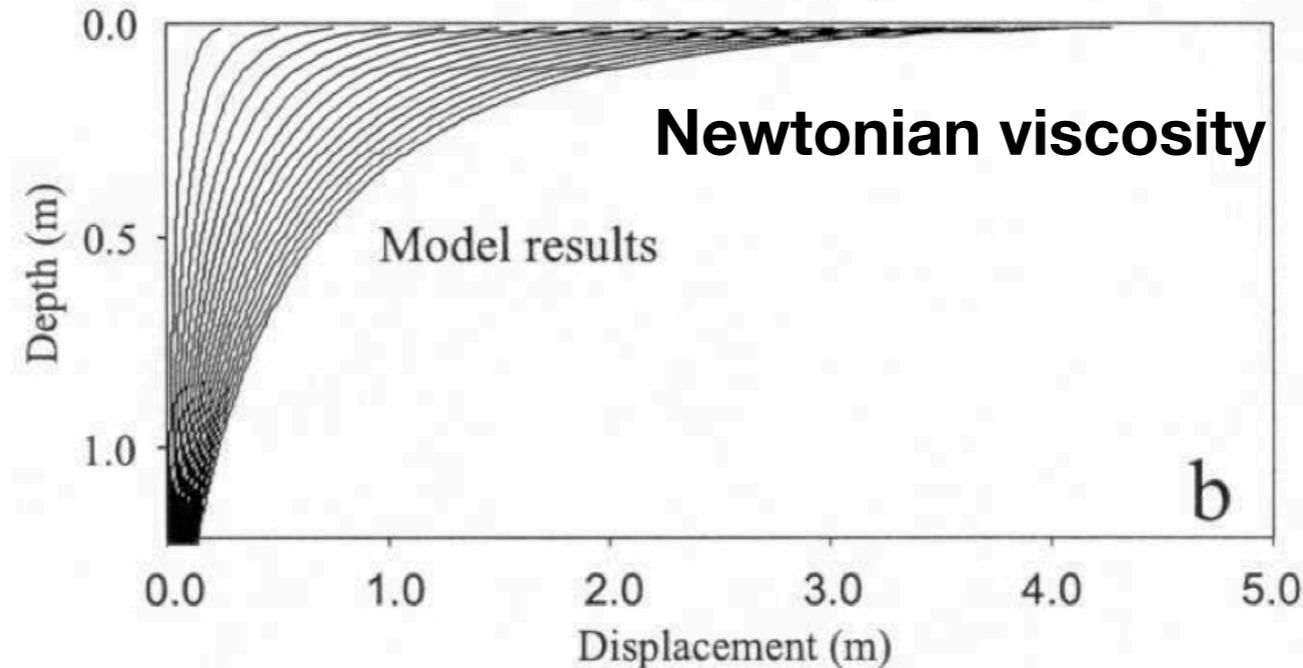
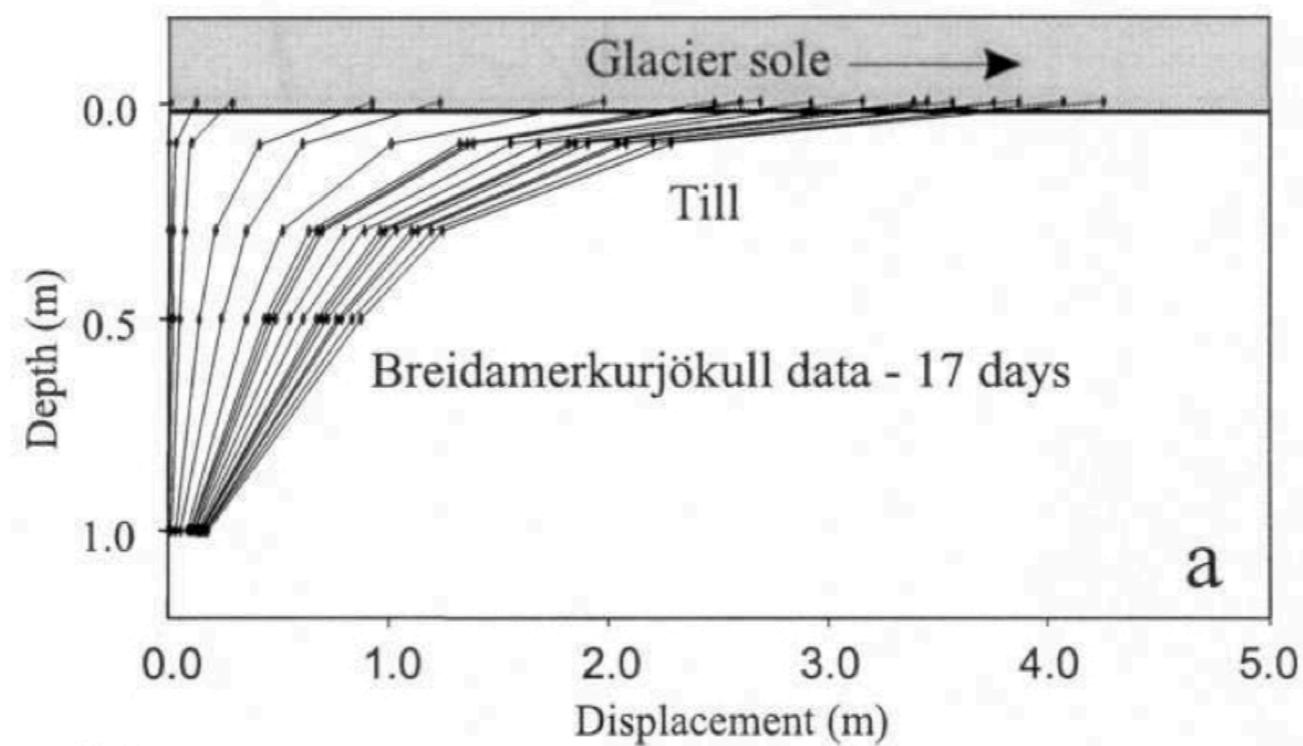
In water-rich tills, the addition of more water beyond the consolidation threshold of till causes a reduction in contact between grains within the till, lowering the basal shear stress required for plastic yielding of the till

# So how does the bed get so weak in pure ice streams?



# What happens where the bed is soft and can deform?

**Measurements from stakes underneath a glacier in Iceland (Boulton 1979)**



# What happens where the bed is soft and can deform?

Measurements from  
a ring shear  
apparatus with till  
extracted from  
glaciers

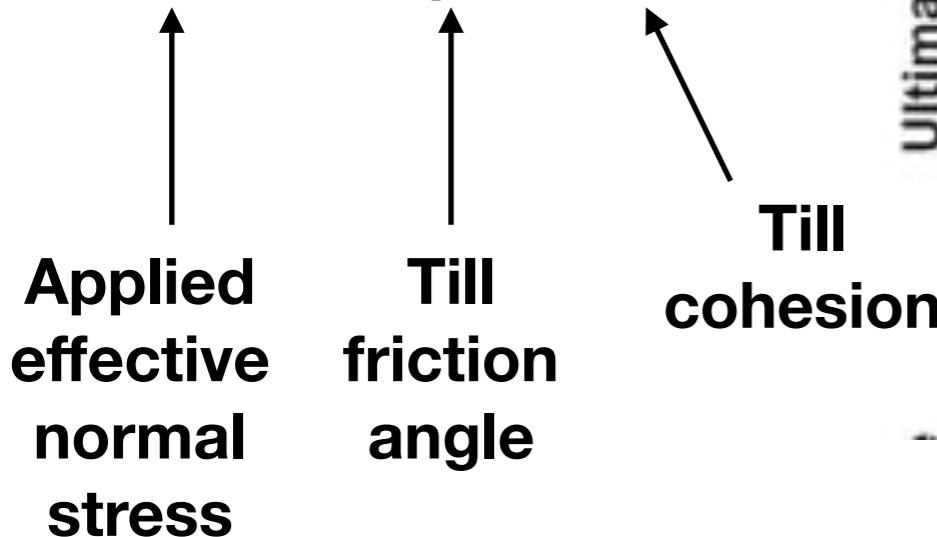


# What happens where the bed is soft and can deform?

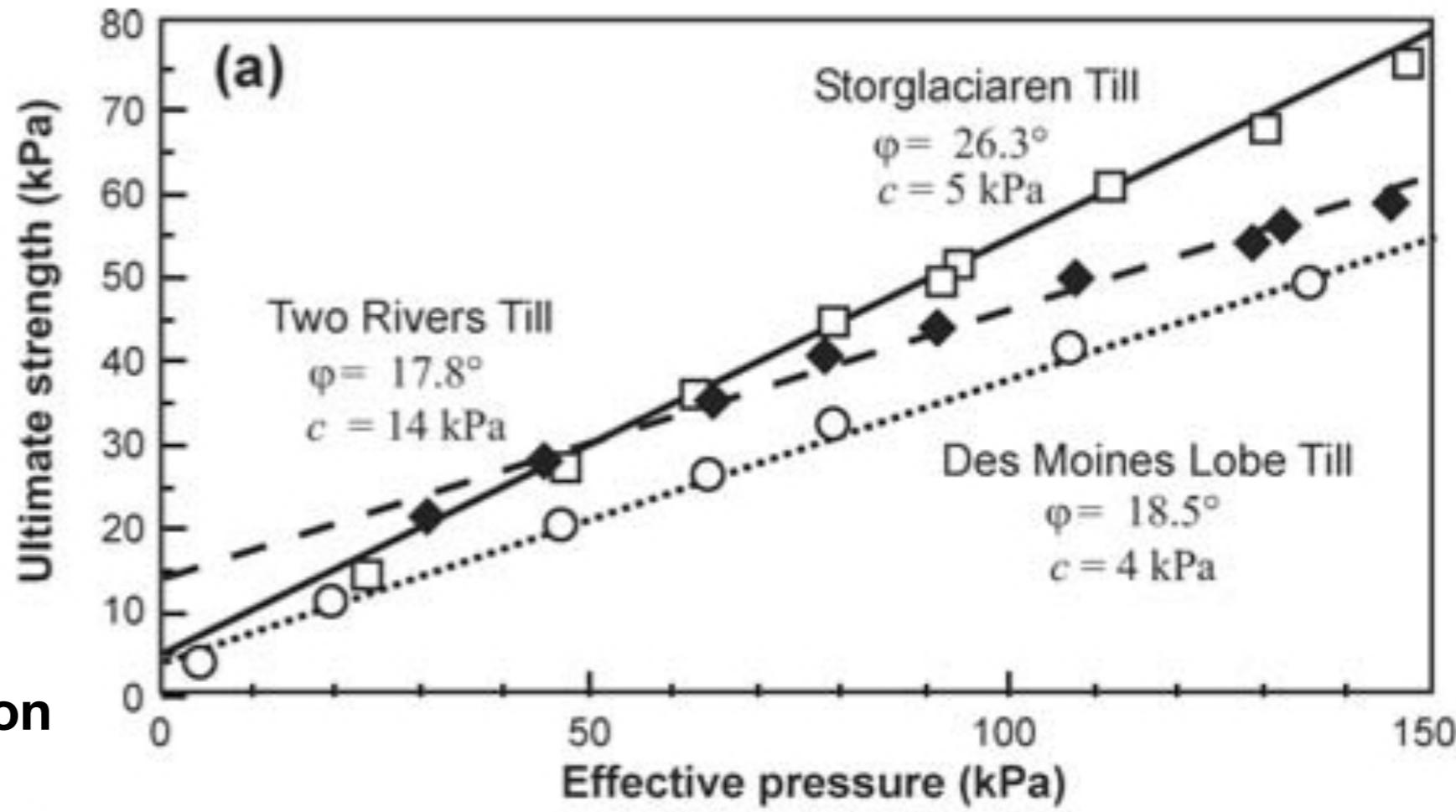
Till is plastic (not deformation below a certain yield stress) - that yield stress depends on effective pressure

Mohr-Coulomb failure criterion

$$\tau = N \tan(\phi) + c$$



Measurements from a ring shear apparatus with till extracted from glaciers



Effective pressure increases with depth in till - lowest yield stress nearest to the ice-till interface - still explains Iceland measurements!

# Cohesion

**Cohesion - the attraction of grains to each other - depends on the grain properties and what's in between the grain**

**Sand grains are not very cohesive.**  
**They will only stick to one another if there is some water in the sand, but not too much**



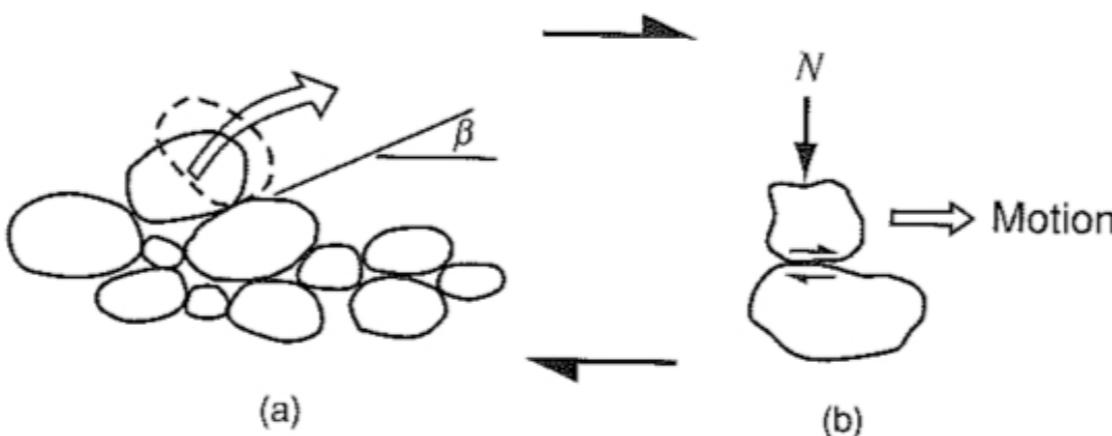
# Cohesion

**Cohesion - the attraction of grains to each other - depends on the grain properties and what's in between the grain**

**Clay grains  
are very small,  
and hence  
cohesive.  
Water makes  
them even  
more  
cohesive**

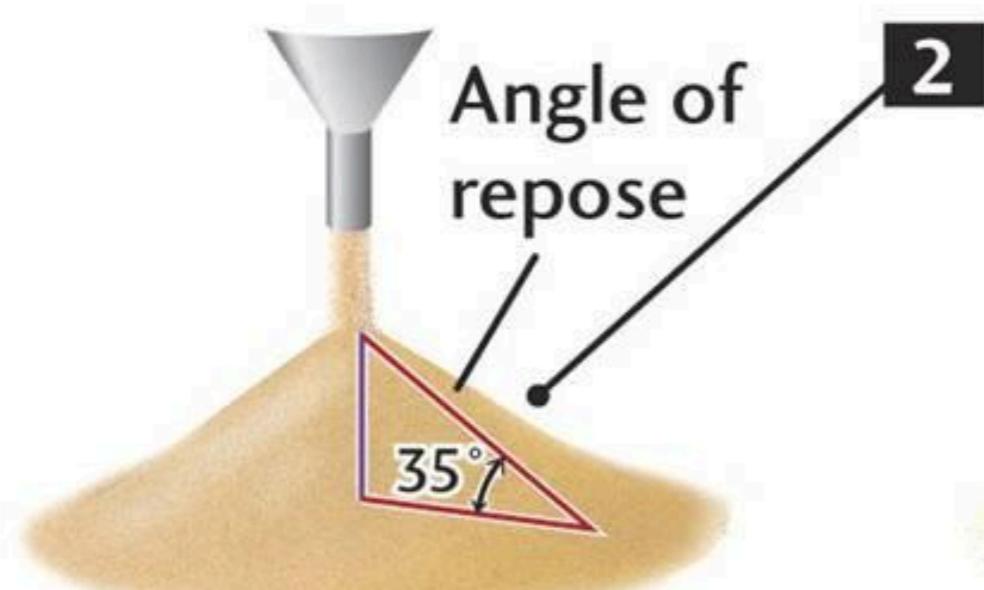


# Friction angle and the angle of repose

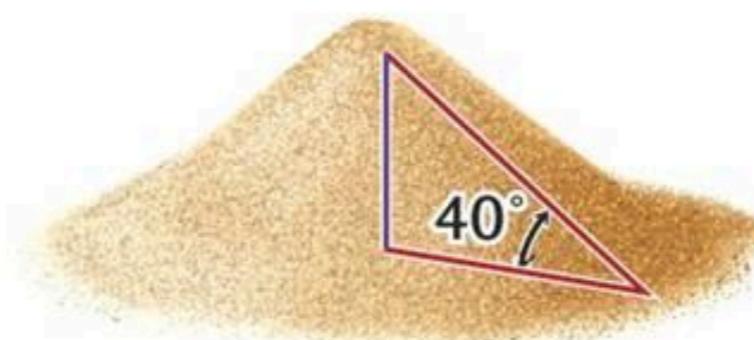


MASS MOVEMENT DEPENDS ON THE NATURE OF MATERIAL, WATER CONTENT, AND SLOPE STEEPNESS

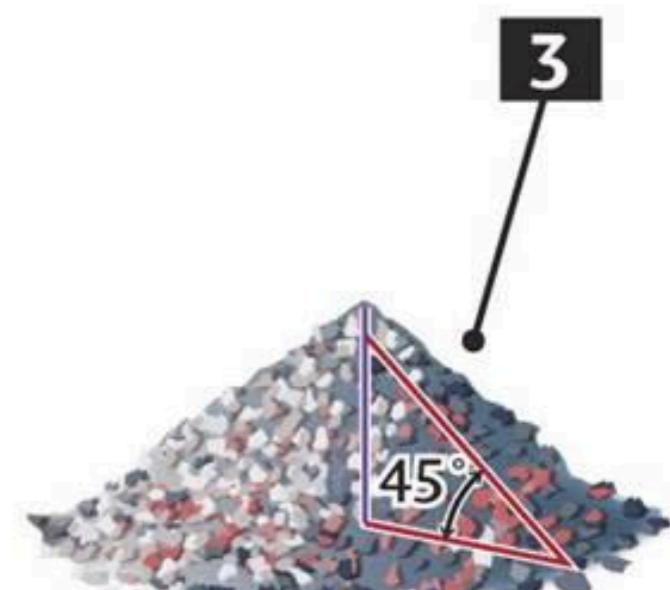
1



Fine sand

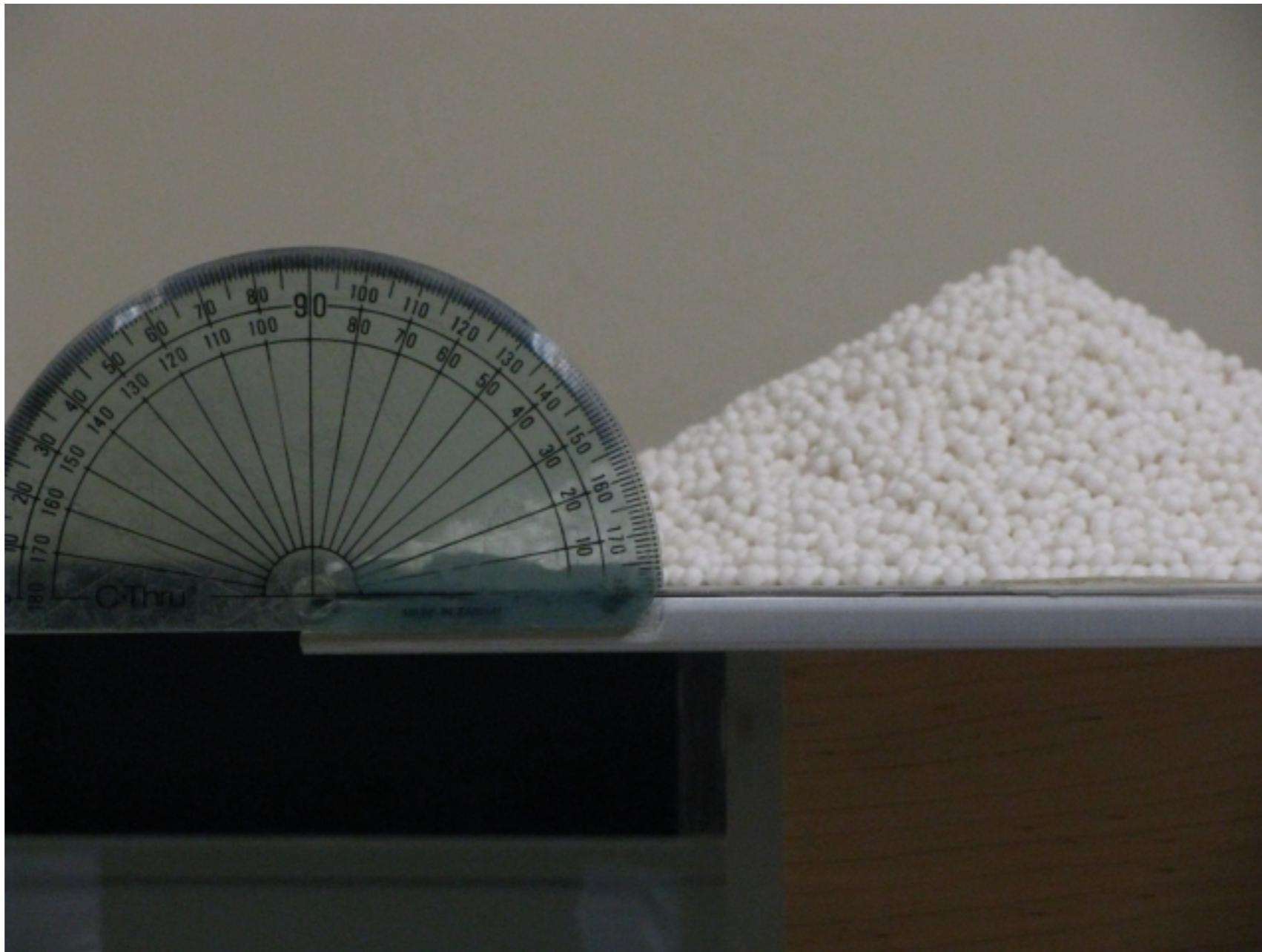


Coarse sand



Angular pebbles

# Friction angle and the angle of repose

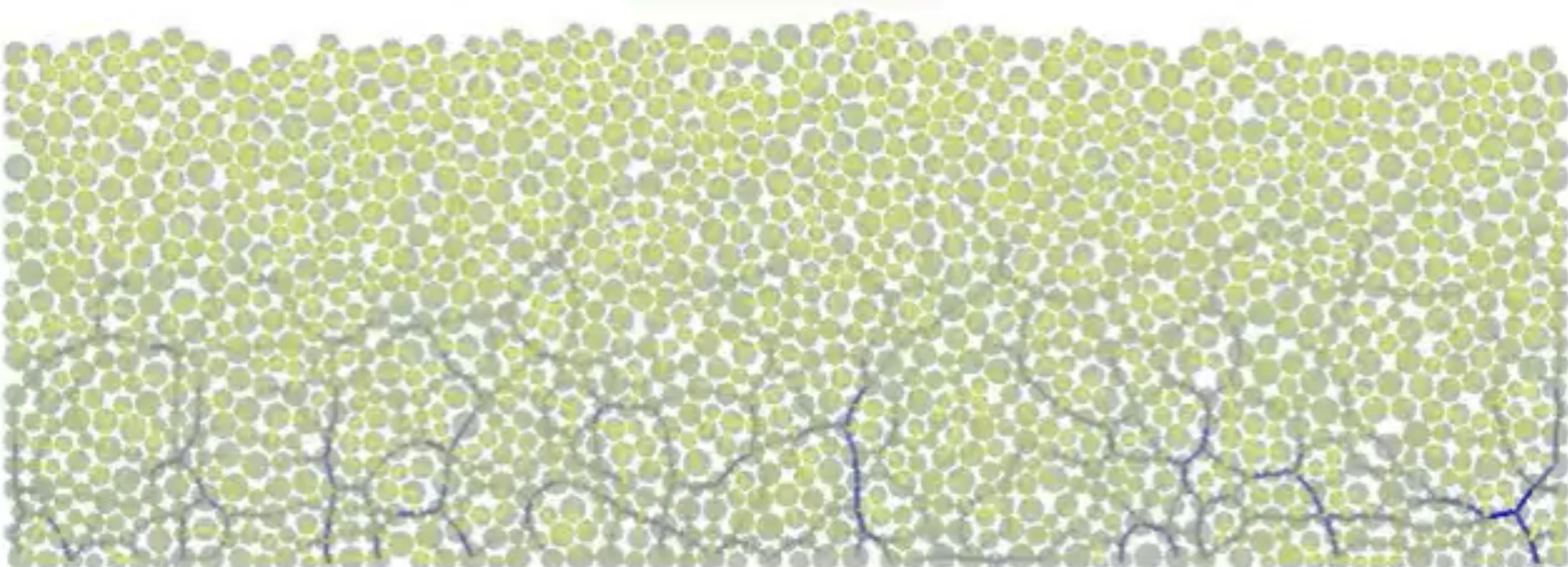
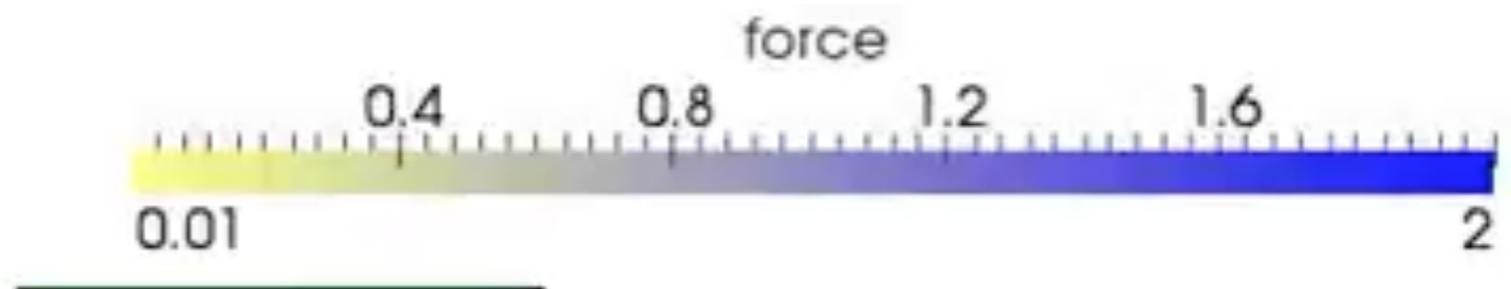


# Grain contacts transmit stress from boundaries over long distances

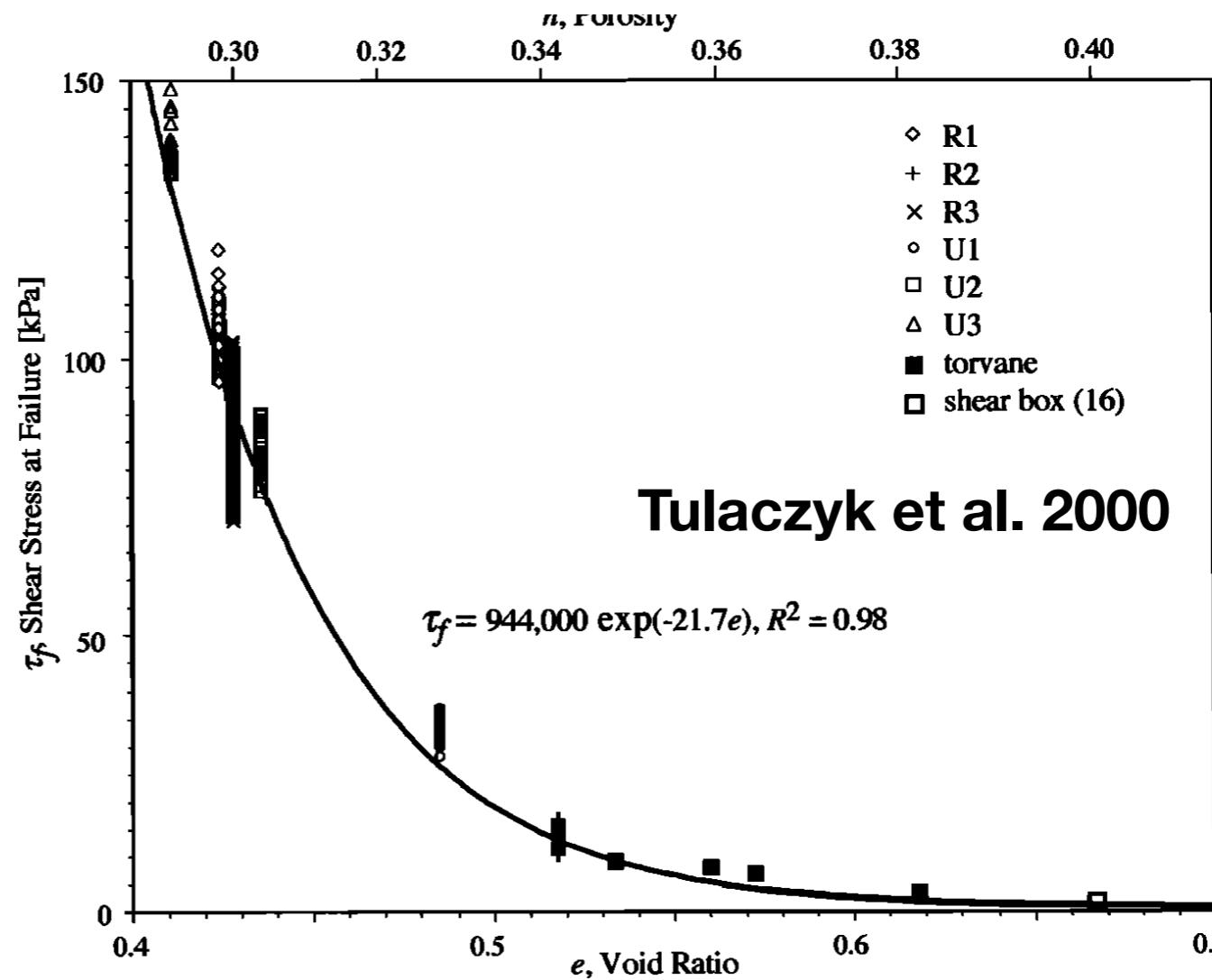
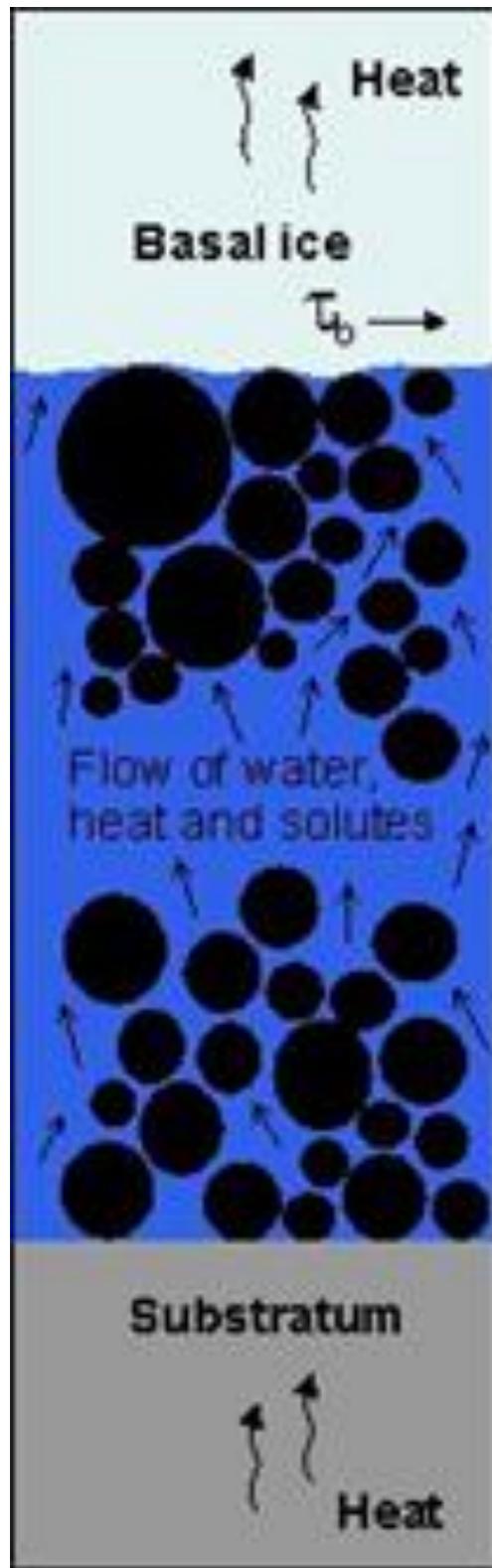


# Grain contacts transmit stress from boundaries over long distances

Time: 5.00 s

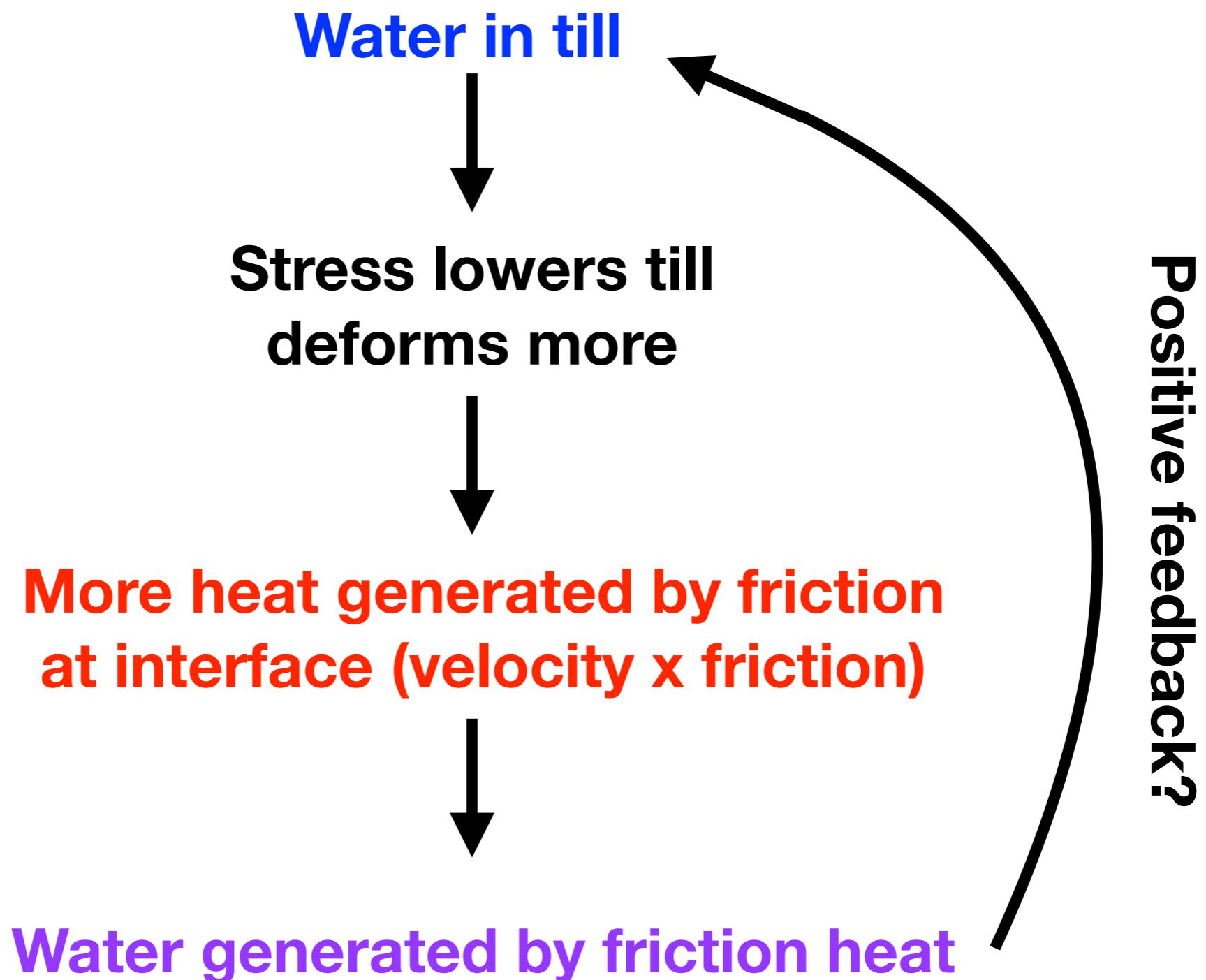


# Saturation of till

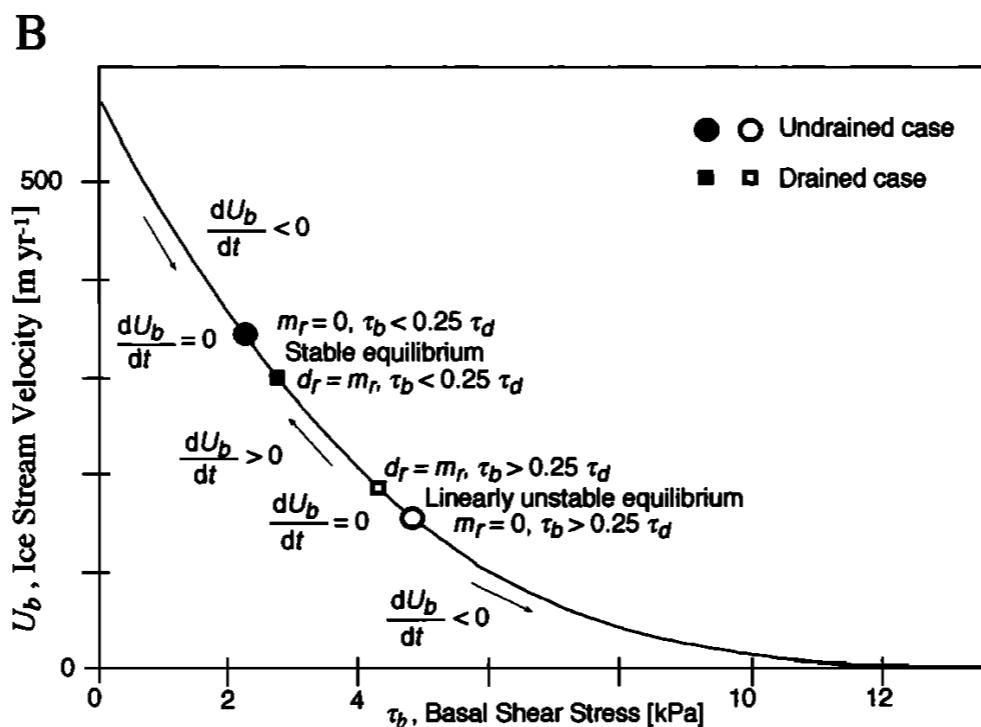
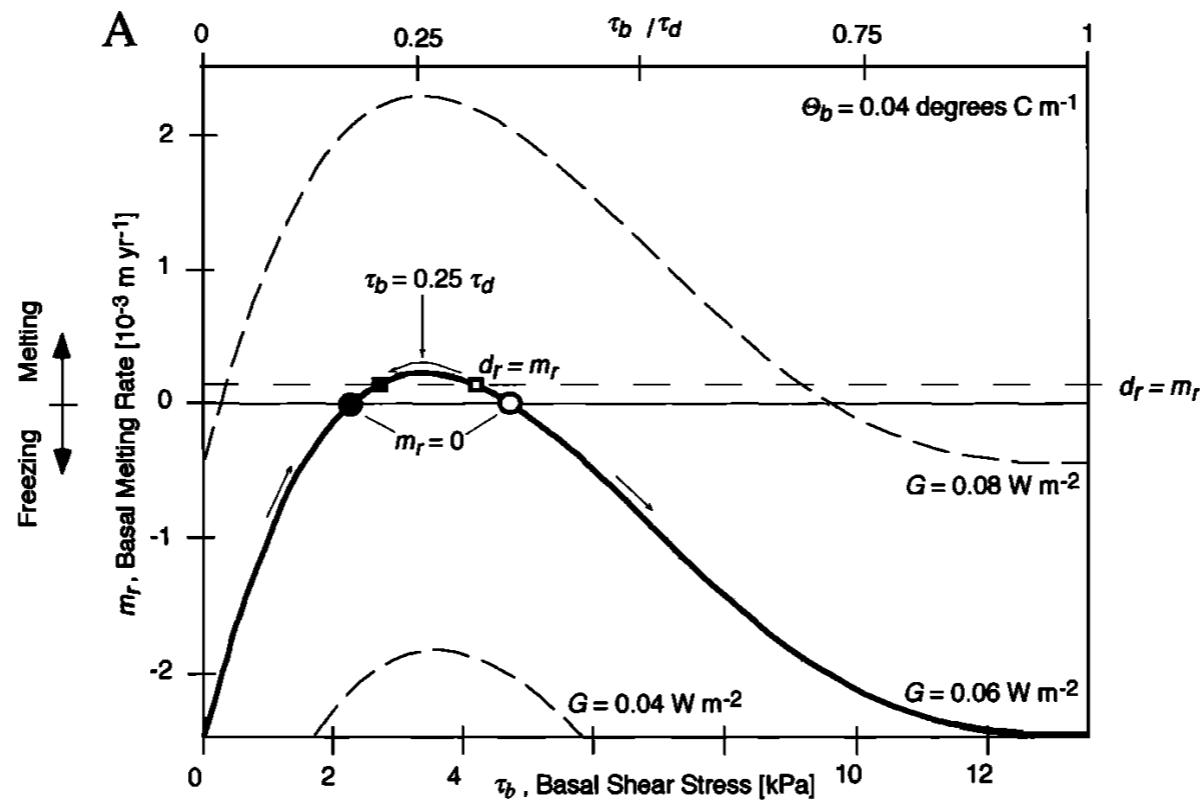


In water-rich tills, the addition of more water beyond the consolidation threshold of till causes a reduction in contact between grains within the till, lowering the basal shear stress required for plastic yielding of the till

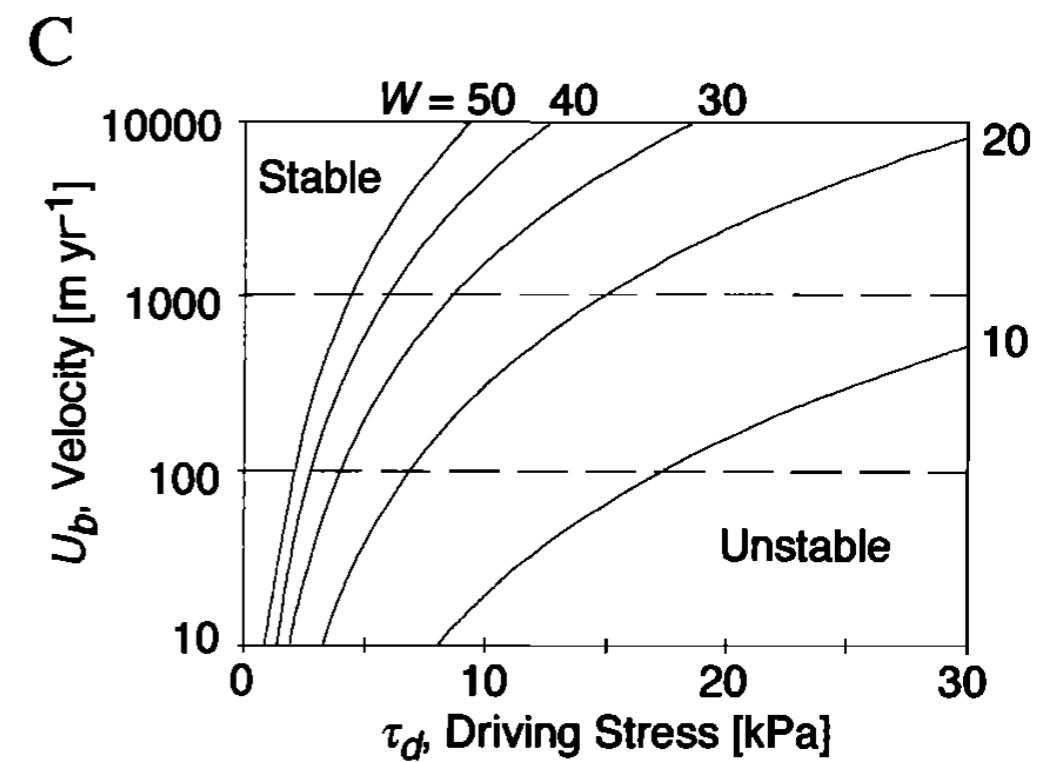
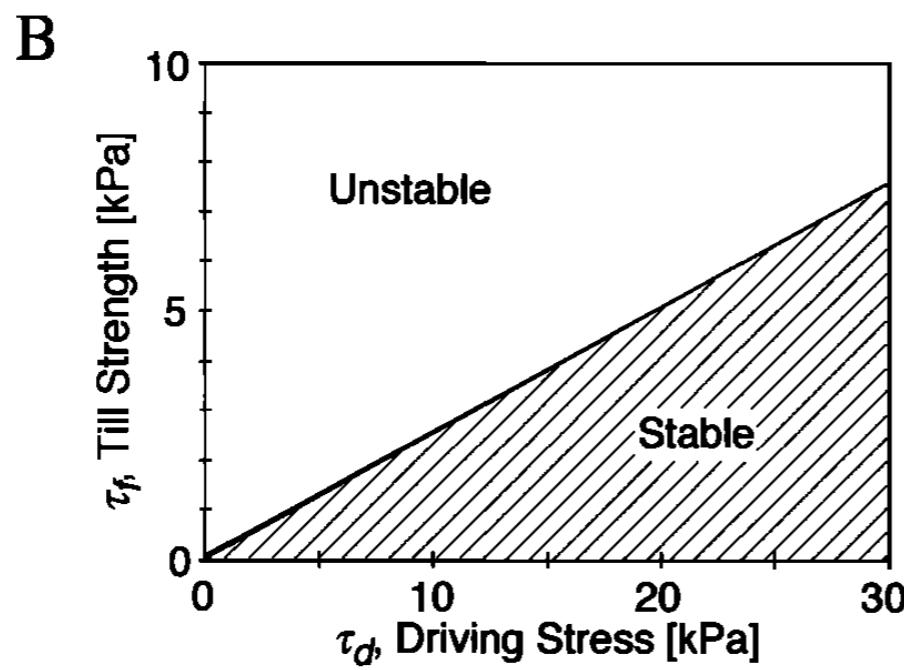
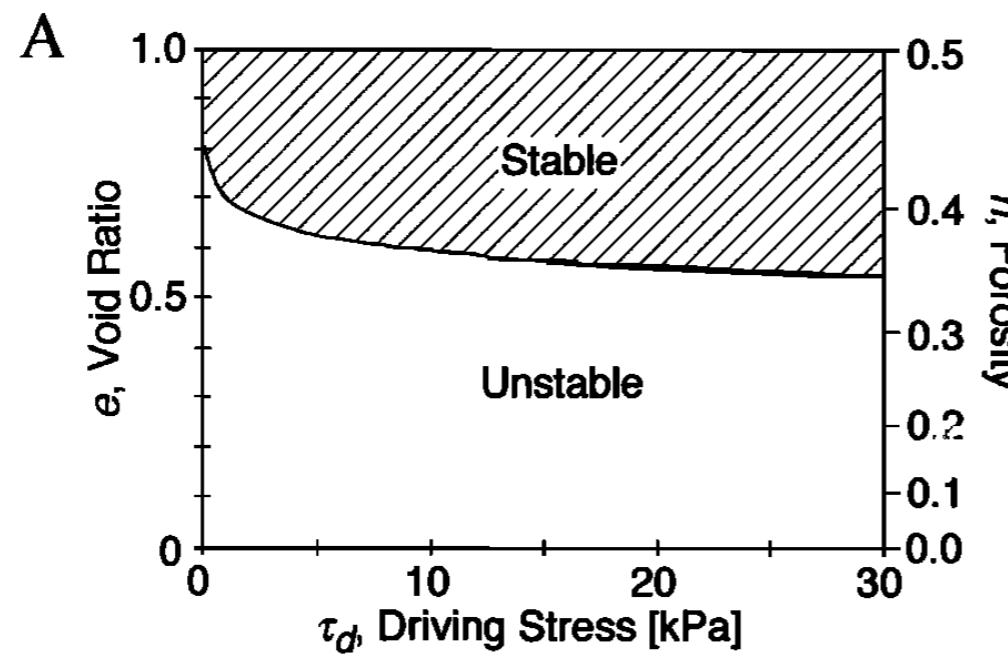
# Saturation of till



# Saturation of till

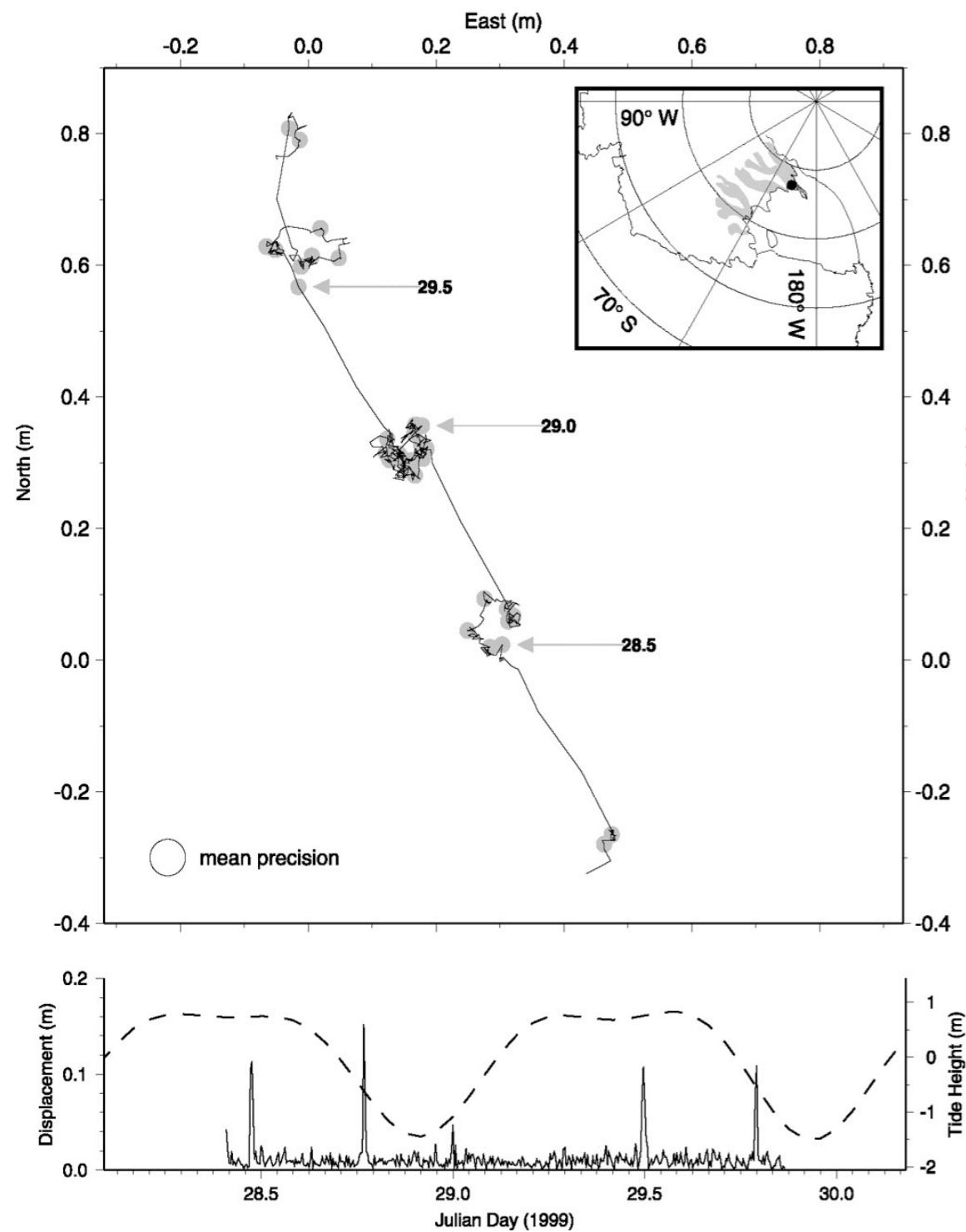


# Saturation of till



# Stick-slip ice stream variability

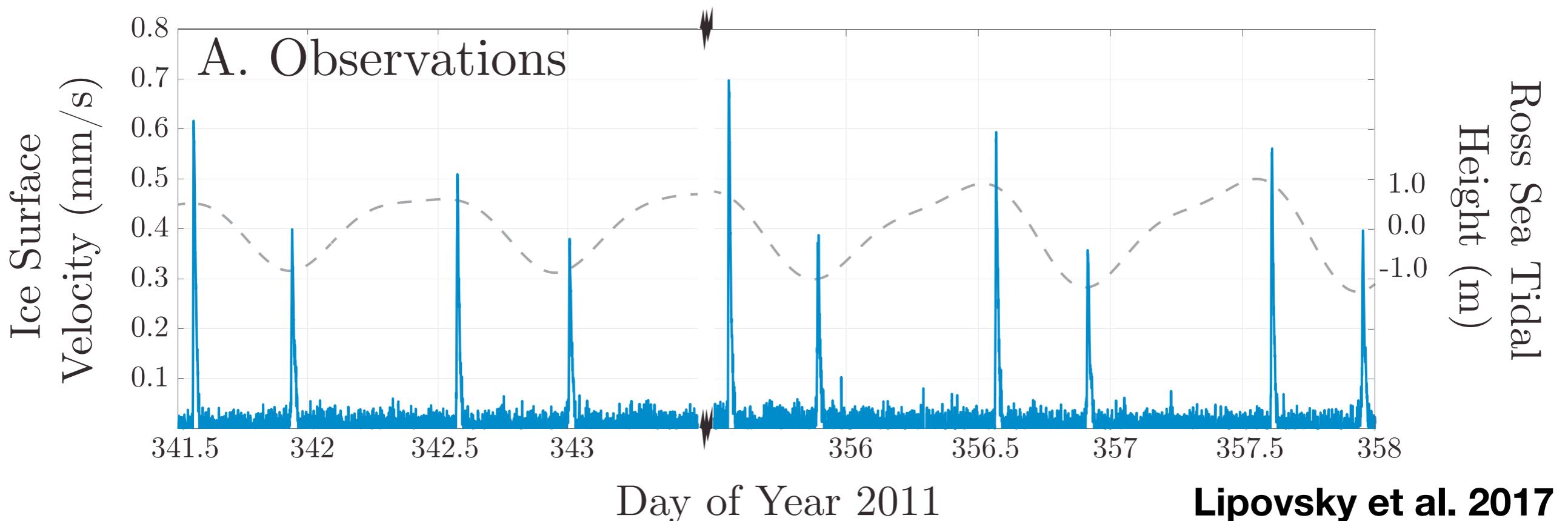
## Whillans Ice Stream, West Antarctica



**Stick-slip motion:** ice is mostly static (stick), but periodically it moves very quickly over a short duration (slip). Similar to how faults move during and in-between earthquakes

# Stick-slip ice stream variability

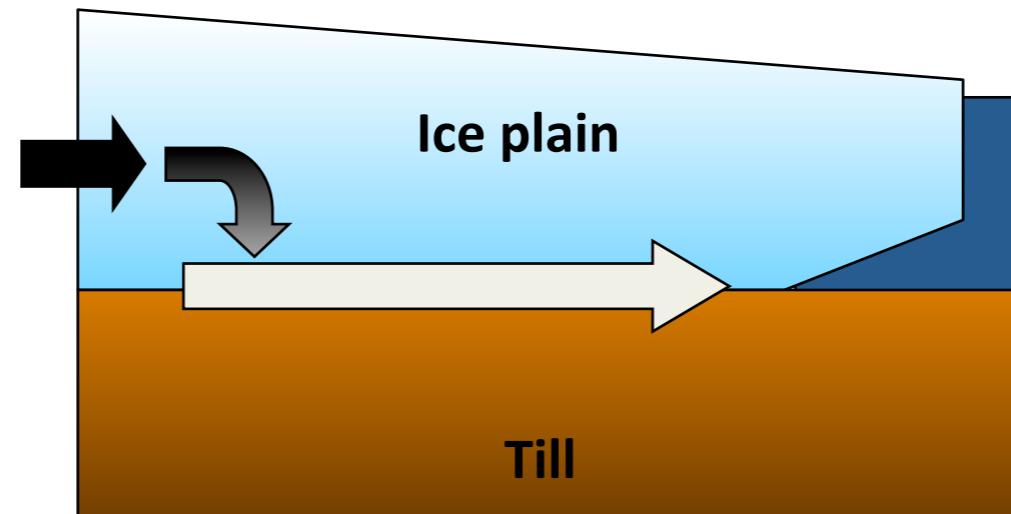
Whillans Ice Stream, West Antarctica



**The slip events at this one ice stream also seem to occur during high or low tide**

# Stick-slip ice stream variability

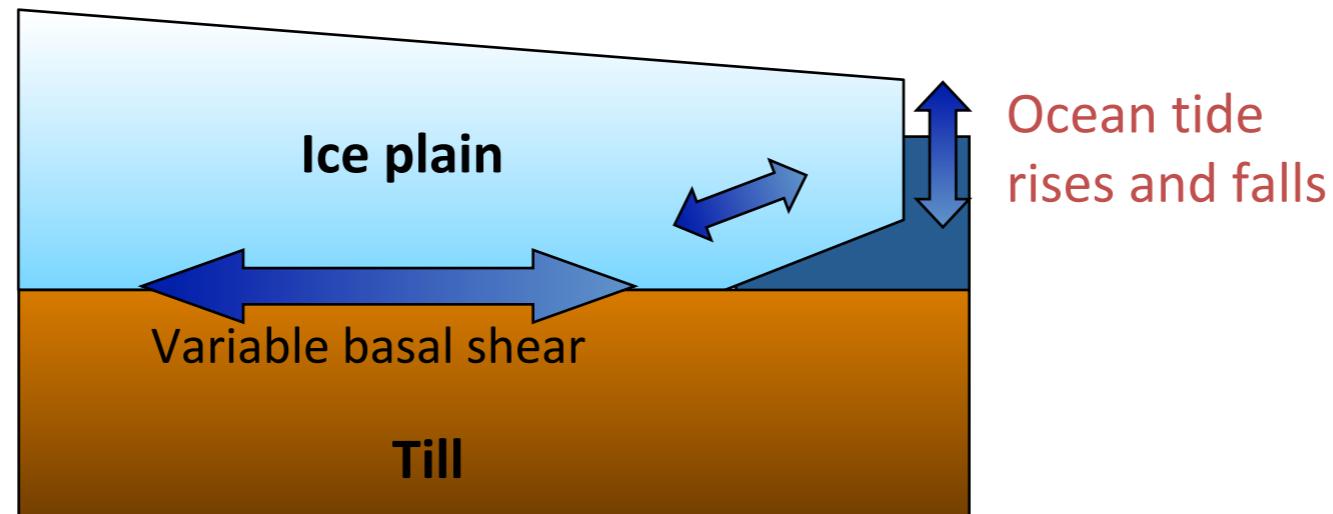
Whillans Ice Stream, West Antarctica



One idea: plastic till is below the “yield stress” that is required for deformation to occur in the till

# Stick-slip ice stream variability

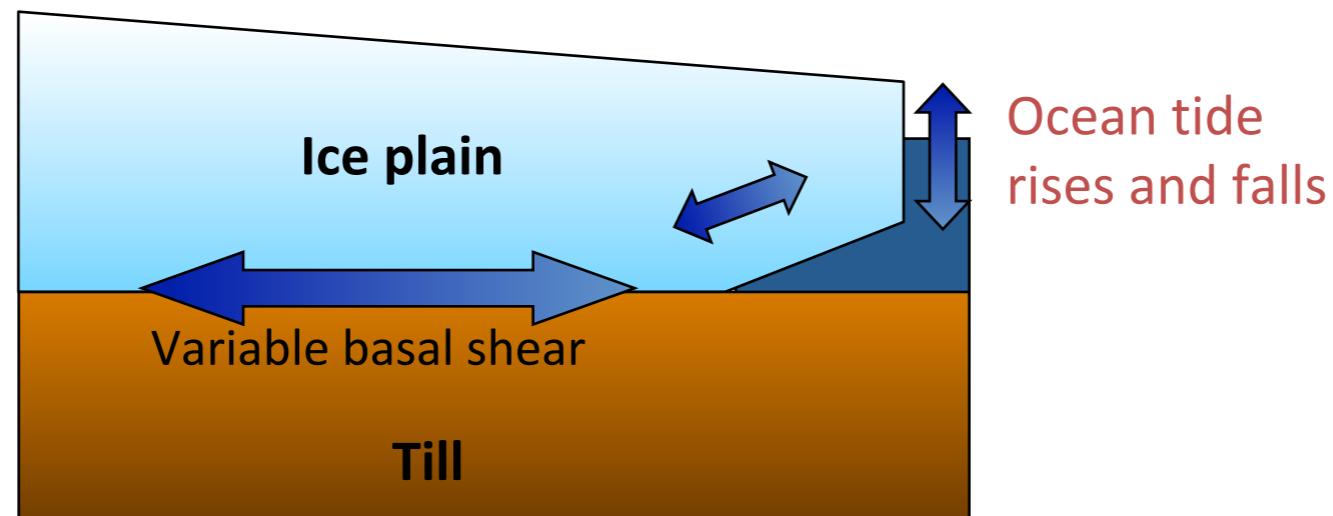
Whillans Ice Stream, West Antarctica



**One idea: when the yield stress is reached, the till slips arbitrarily fast until stress is relieved**

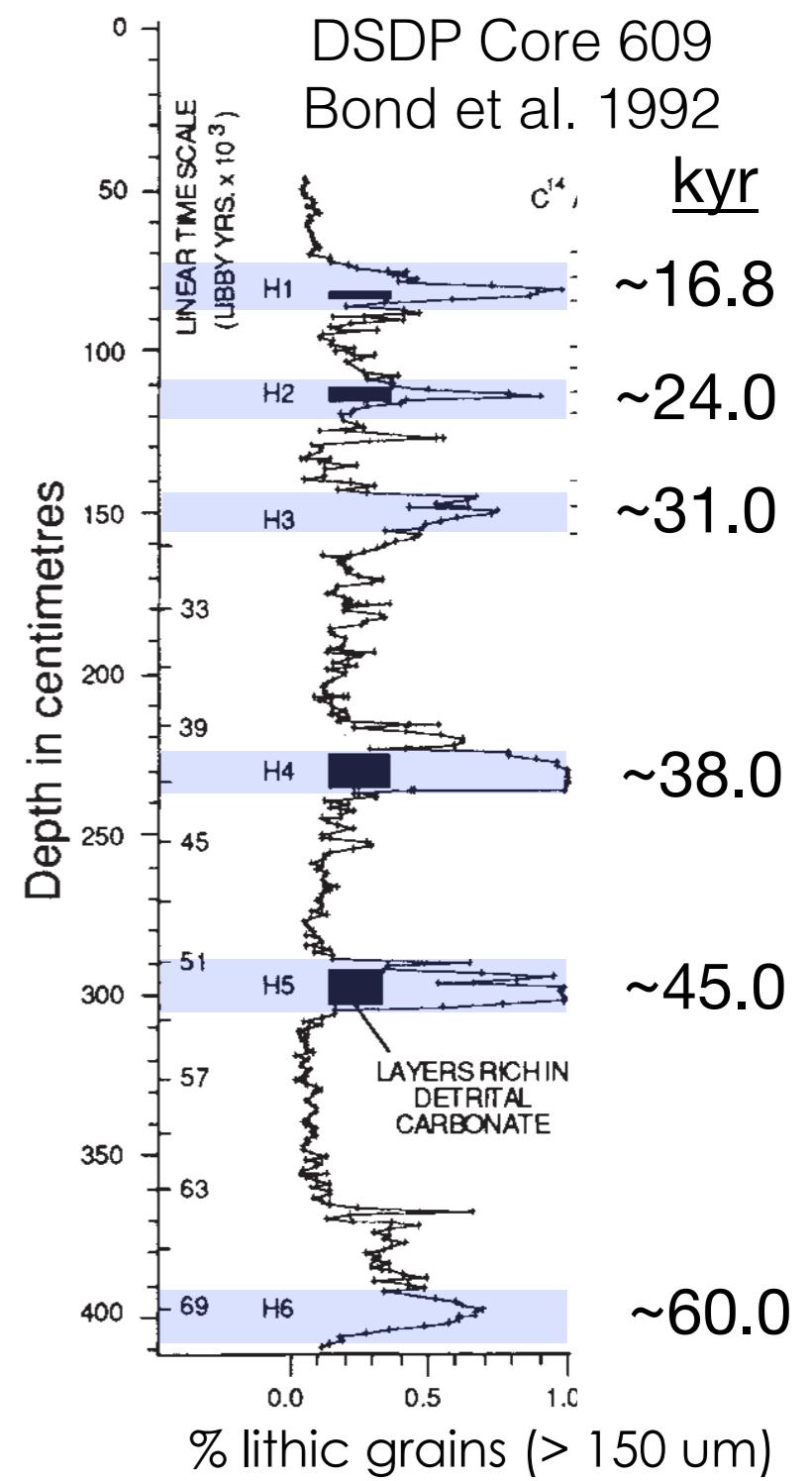
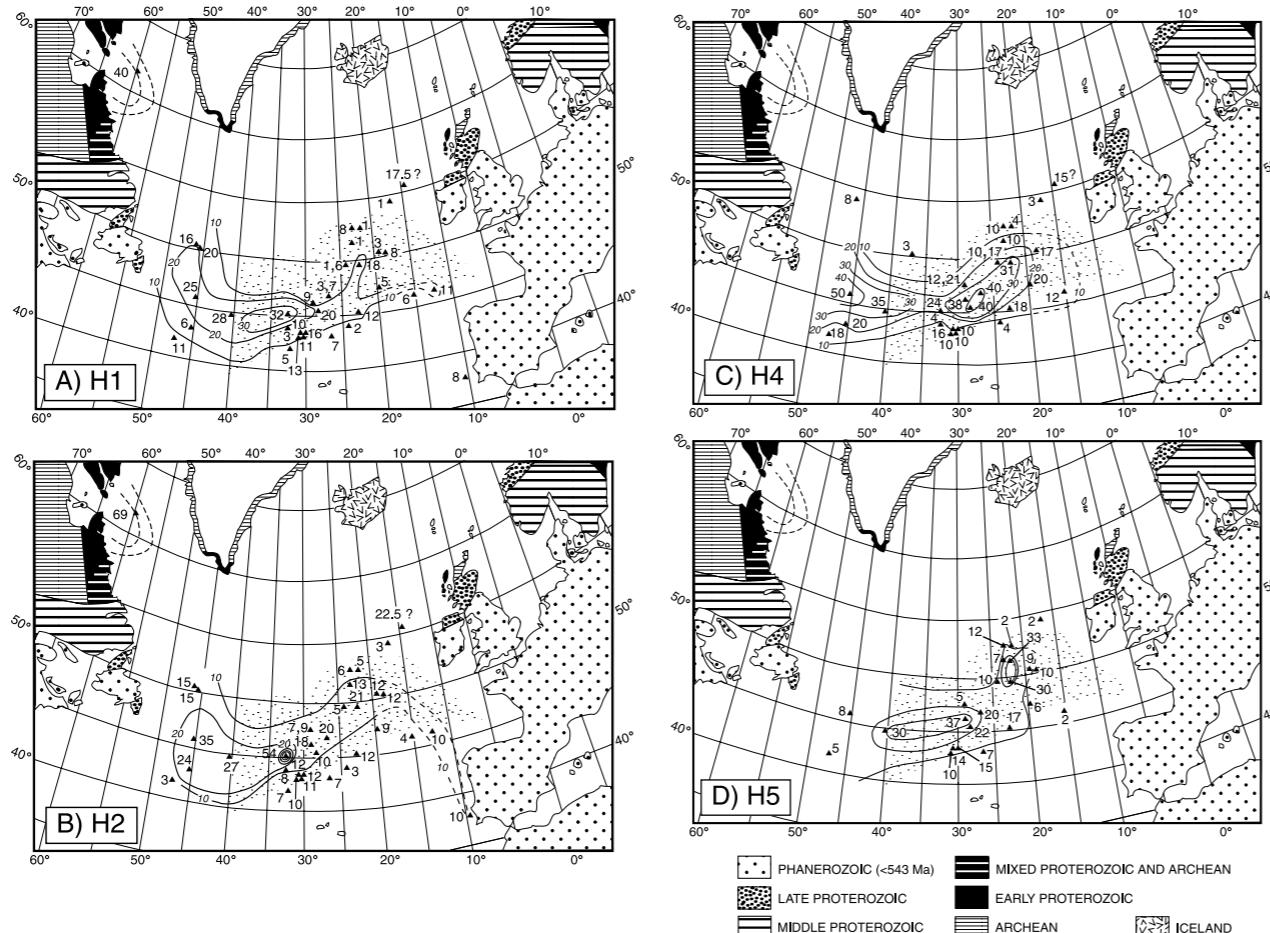
# Stick-slip ice stream variability

Whillans Ice Stream, West Antarctica

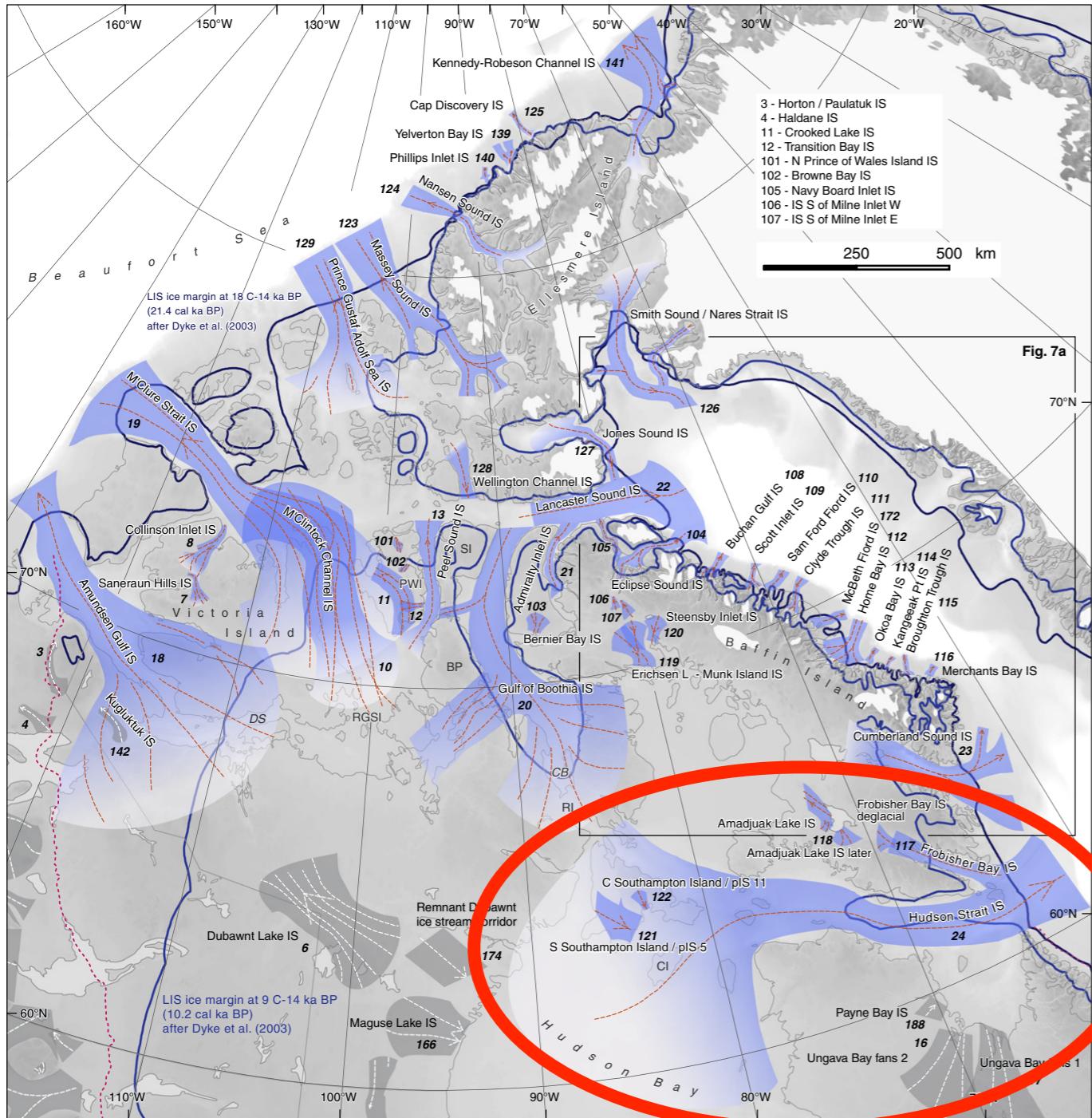


However, to explain the particular phasing and time scale of stick-slip behavior of ice streams (such as Whillans) requires more complicated models than just plastic till (i.e. rate and state behavior, see Lipovsky et al. 2017)

# Centennial/millennial scale ice stream variability

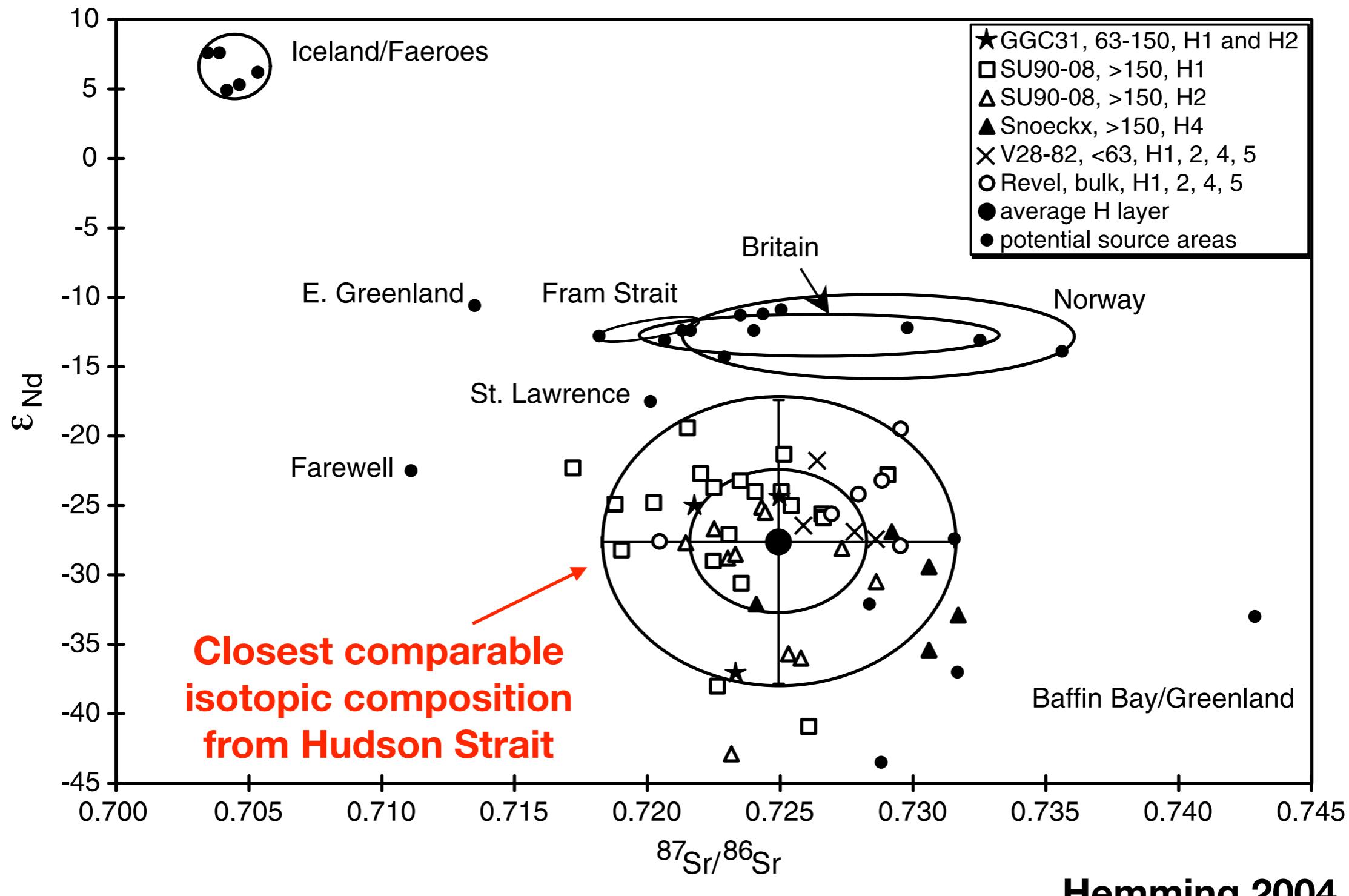


# Centennial/millennial scale ice stream variability

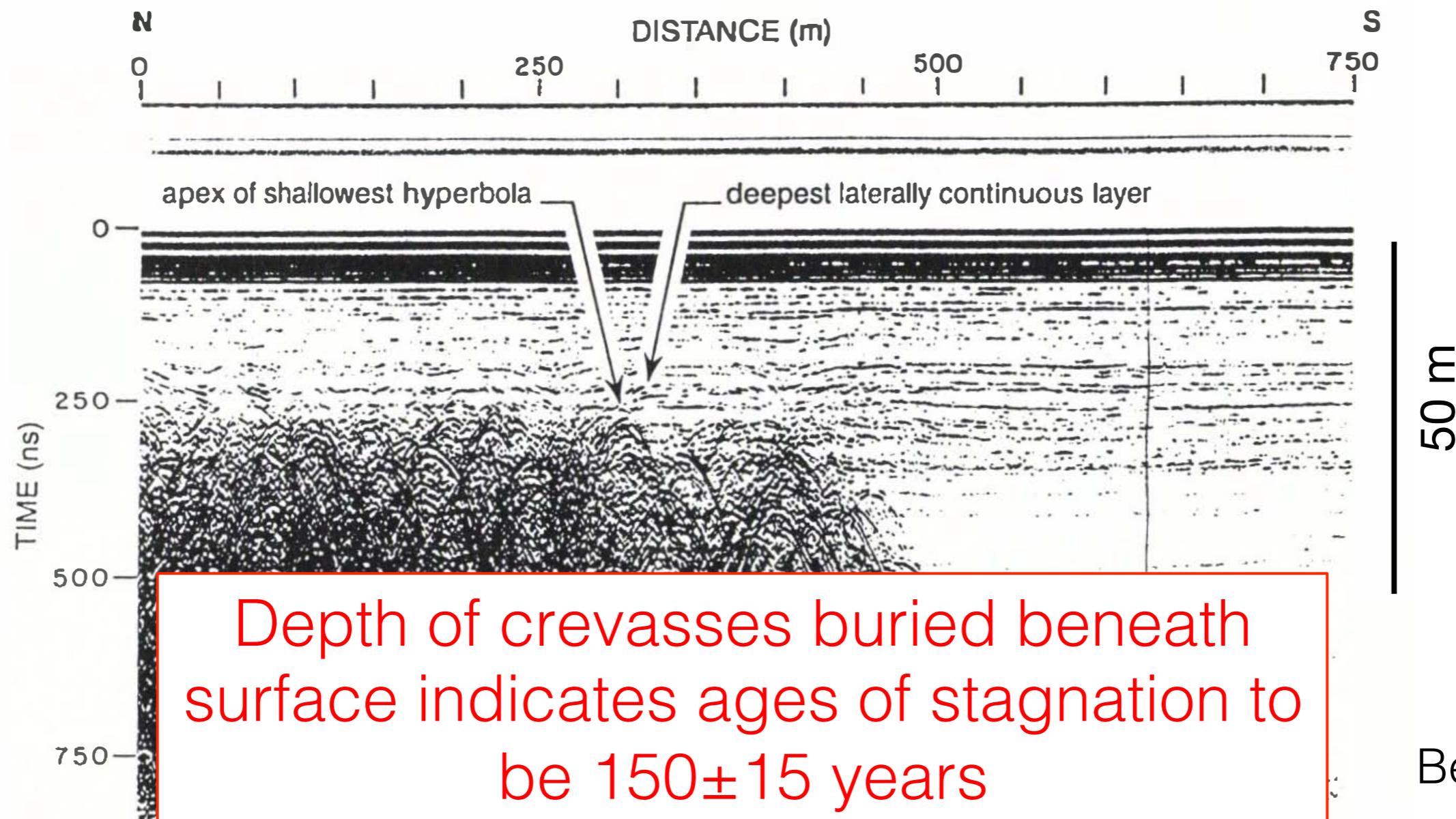


**Hudson Strait Ice Stream - potentially the largest ice stream existing in the geological record (about as long as the distance from Atlanta to Texas), drained the heart of the Laurentide Ice Sheet - variability in ice flow thought to play a role in generating Heinrich events through discharge of icebergs**

# Centennial/millennial scale ice stream variability



# Centennial/millennial scale ice stream variability



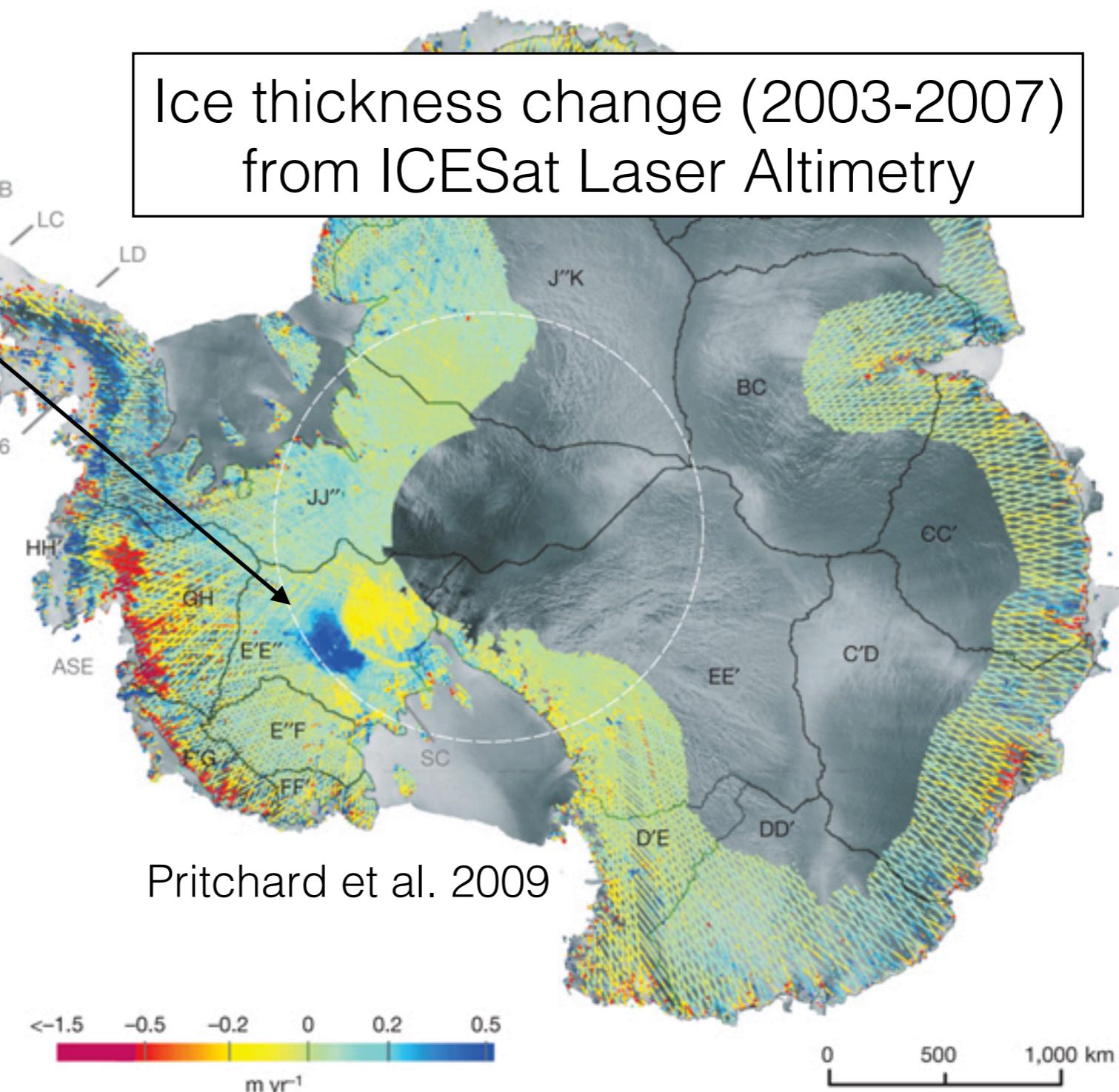
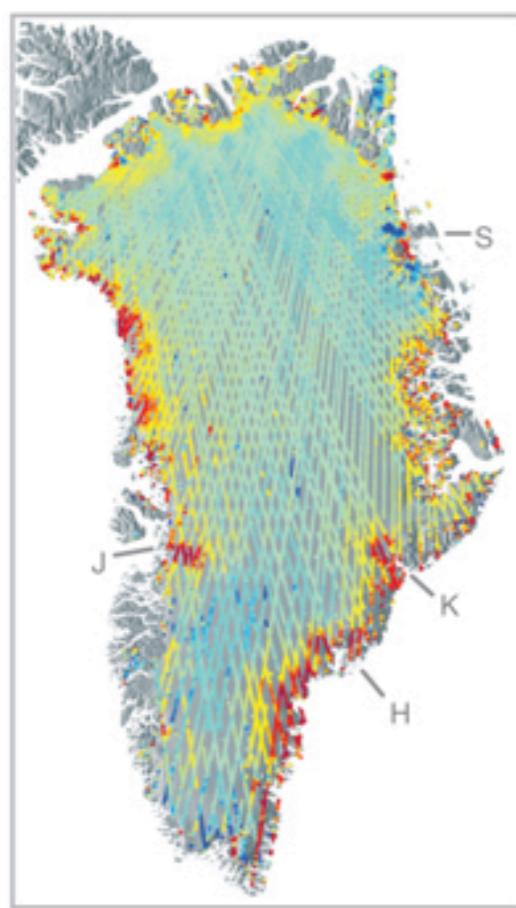
Ice streams in West Antarctica also seem to exhibit flow variability on time scales of centuries

# Centennial/millennial scale ice stream variability

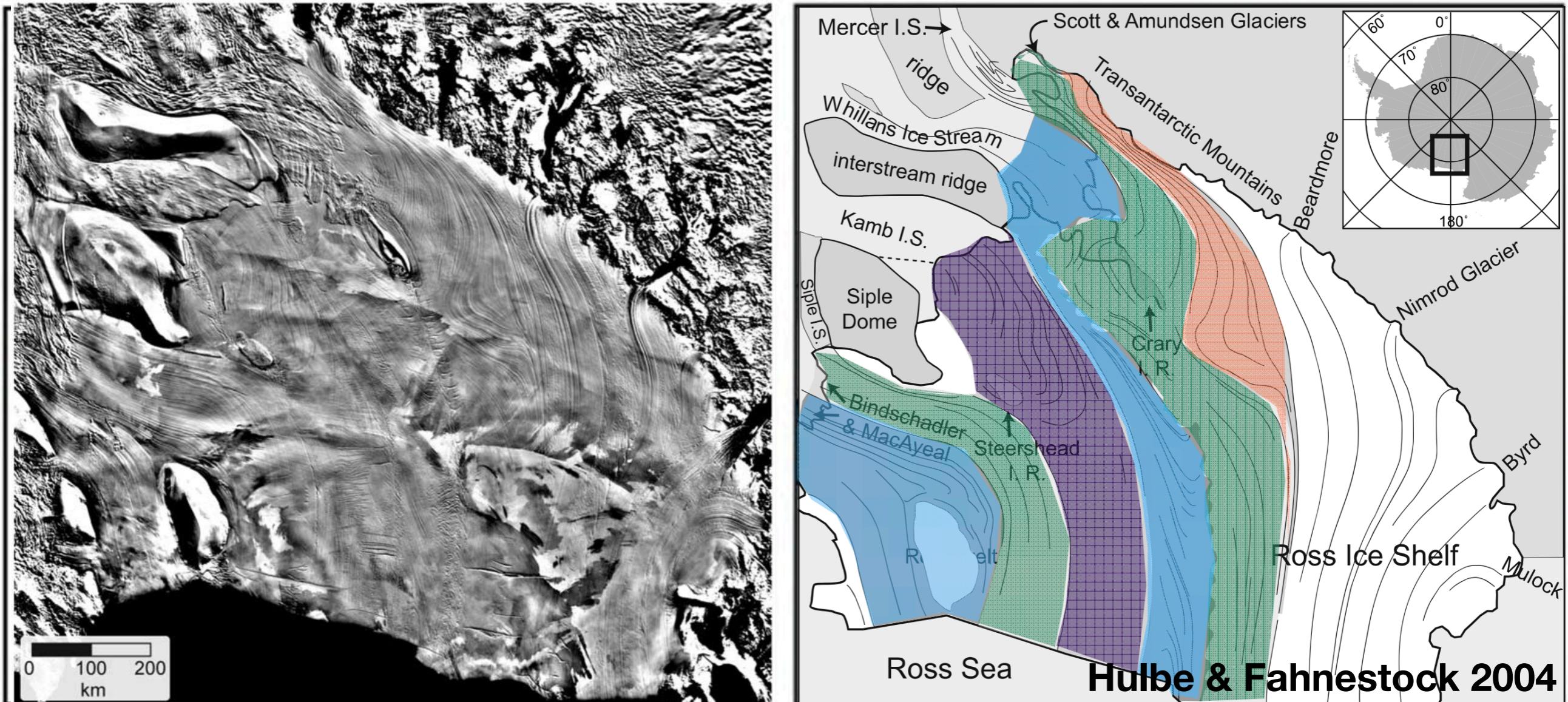
Kamb Ice Stream

shutdown is the cause  
of this odd “thickening  
anomaly” in Siple  
Coast

Ice thickness change (2003-2007)  
from ICESat Laser Altimetry

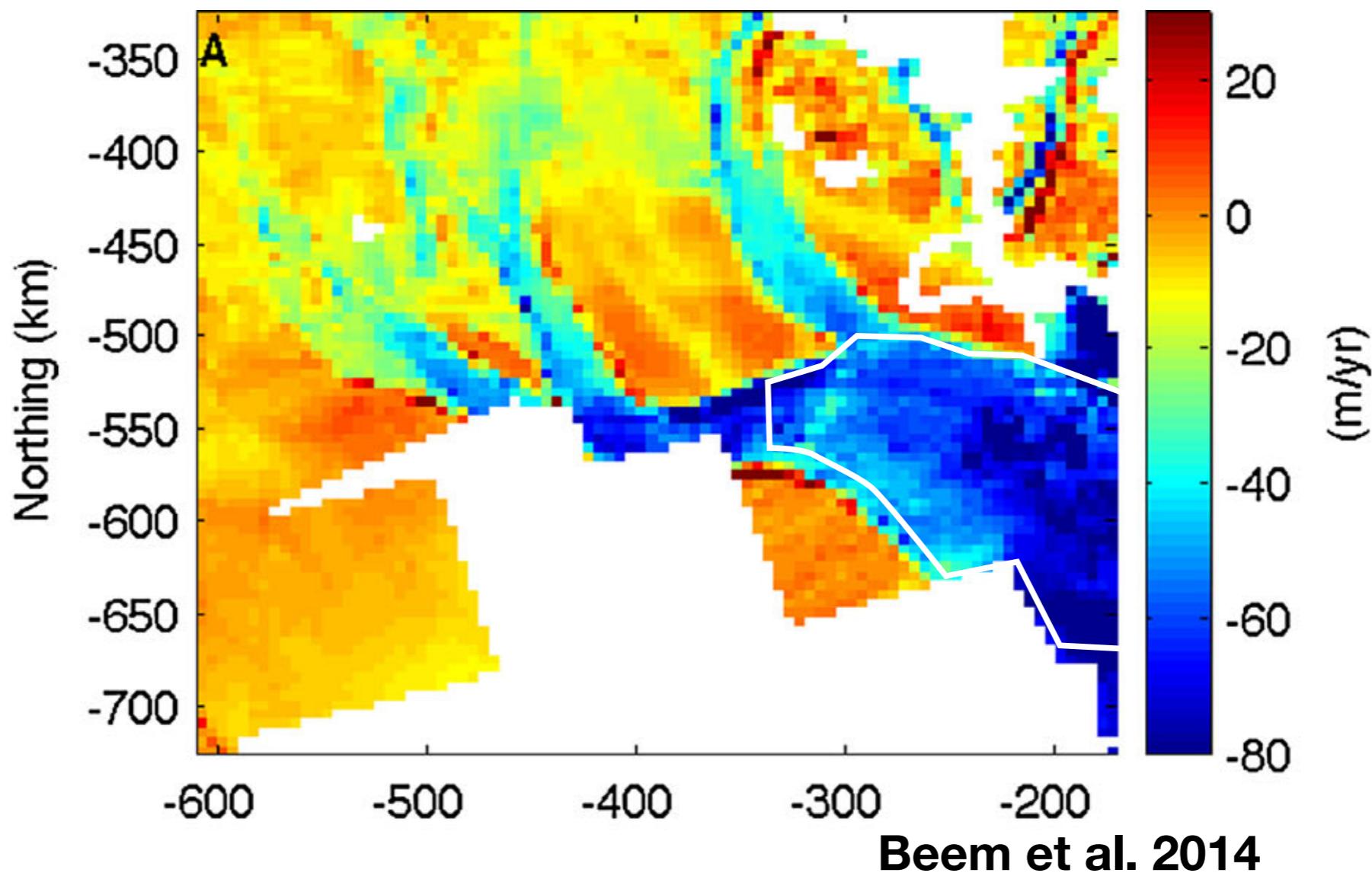


# Centennial/millennial scale ice stream variability



Folds/undulations in streamlines the Ross Ice Shelf (in the absence of other obstructions) indicate that the relative flow speed of adjacent ice streams has changed over time

# Centennial/millennial scale ice stream variability



Difference in Whillans ice stream velocity between  
1997 and 2009

# Ice stream variability

