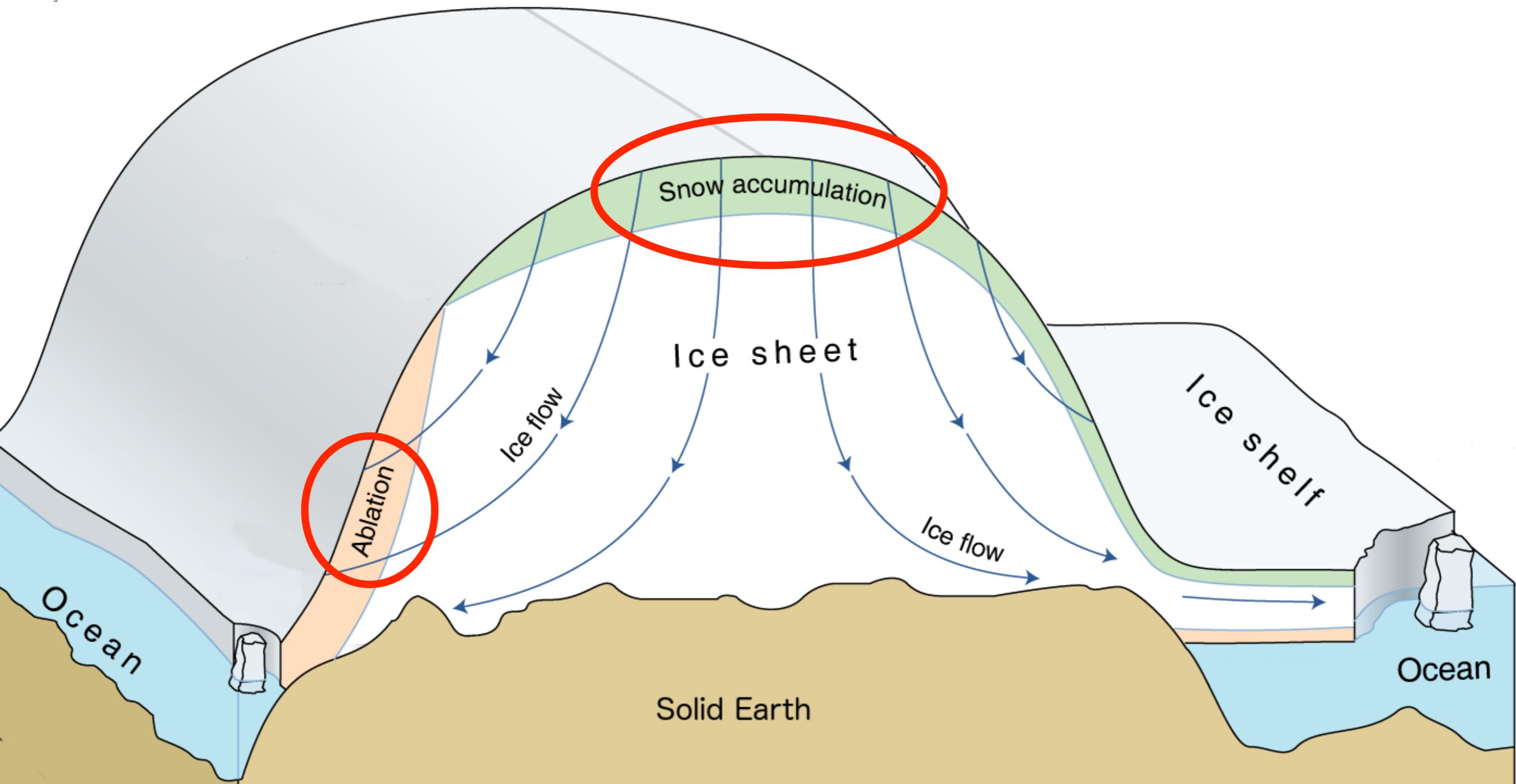


Surface mass balance

What

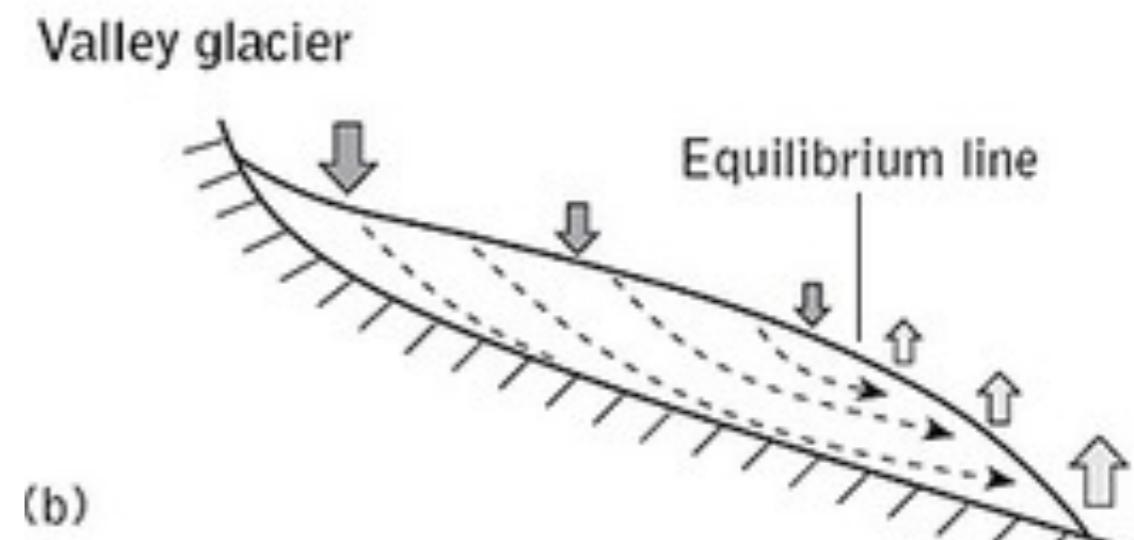
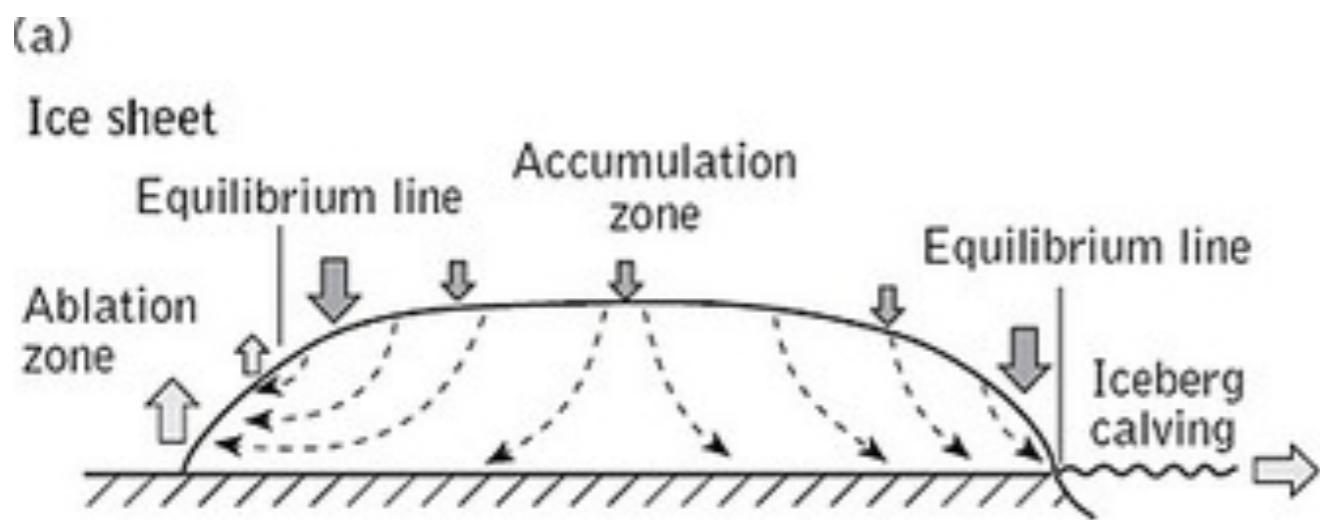


Mass balance

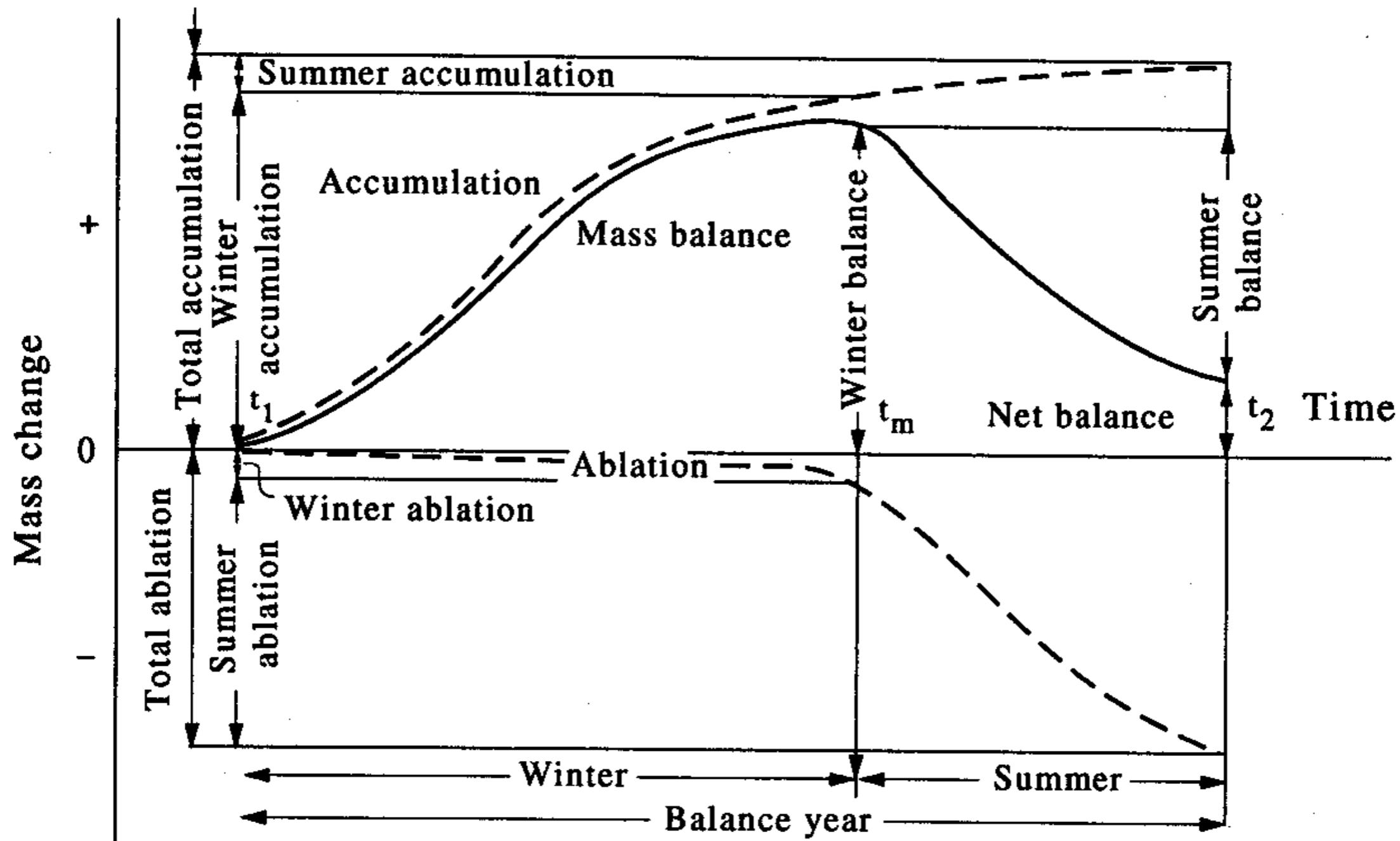
Mass balance: the sum of all sources and sinks of ice in a glacier/ice sheet

For a glacier that is at steady-state, there is typically:

1. An accumulation zone
2. An ablation/wastage zone
3. Ice flow from the accumulation zone to the ablation zone



Local Mass Balance

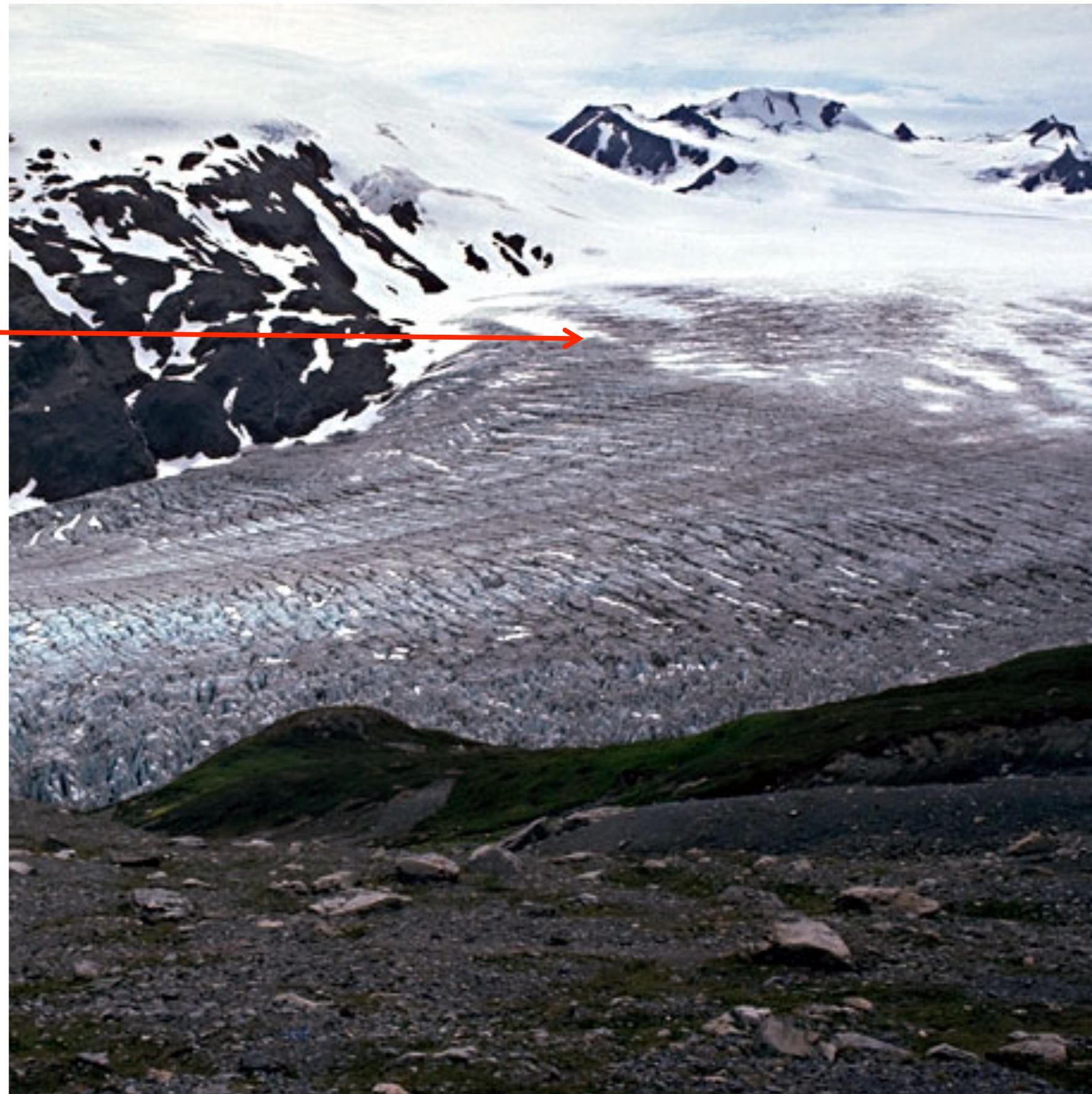


From Paterson (1994)

Equilibrium line

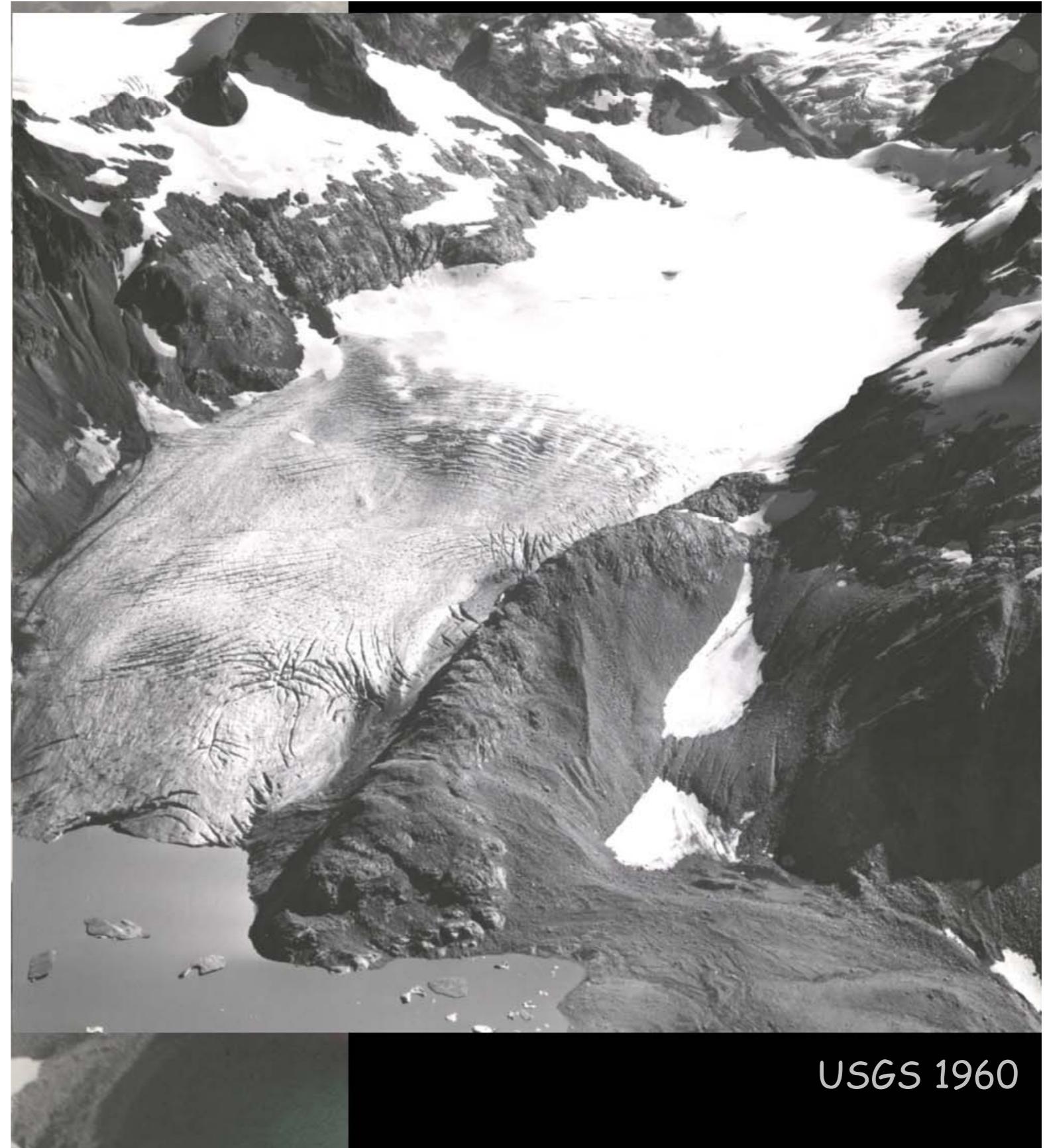
equilibrium
line/zone

Equilibrium line:
boundary between
accumulation zone
and ablation zone
where local surface
mass balance is
zero. Often used
as single-number
proxy for climate.



Equilibrium line

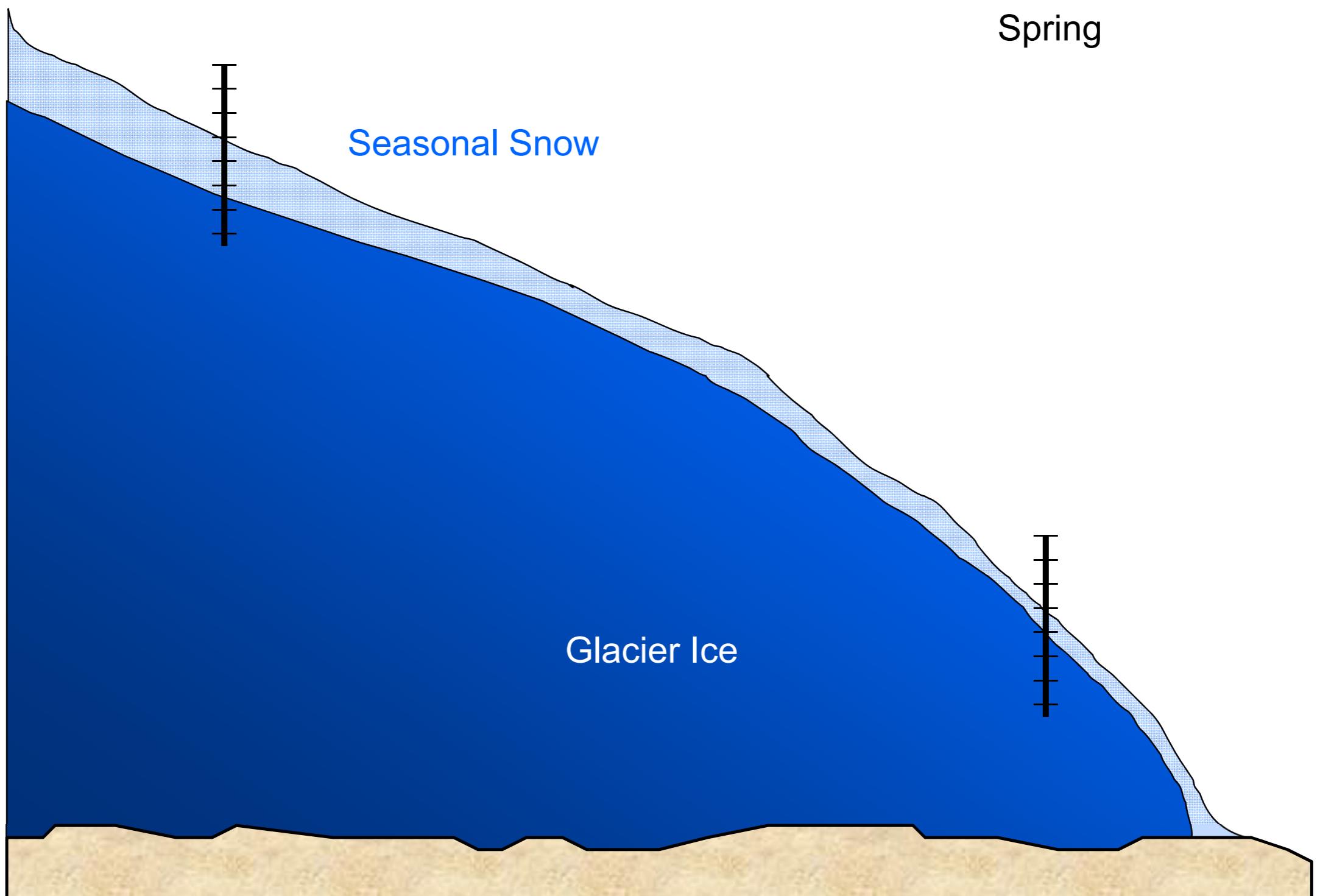
Equilibrium line:
boundary between
accumulation zone
and ablation zone
where local surface
mass balance is
zero. Often used
as single-number
proxy for climate.



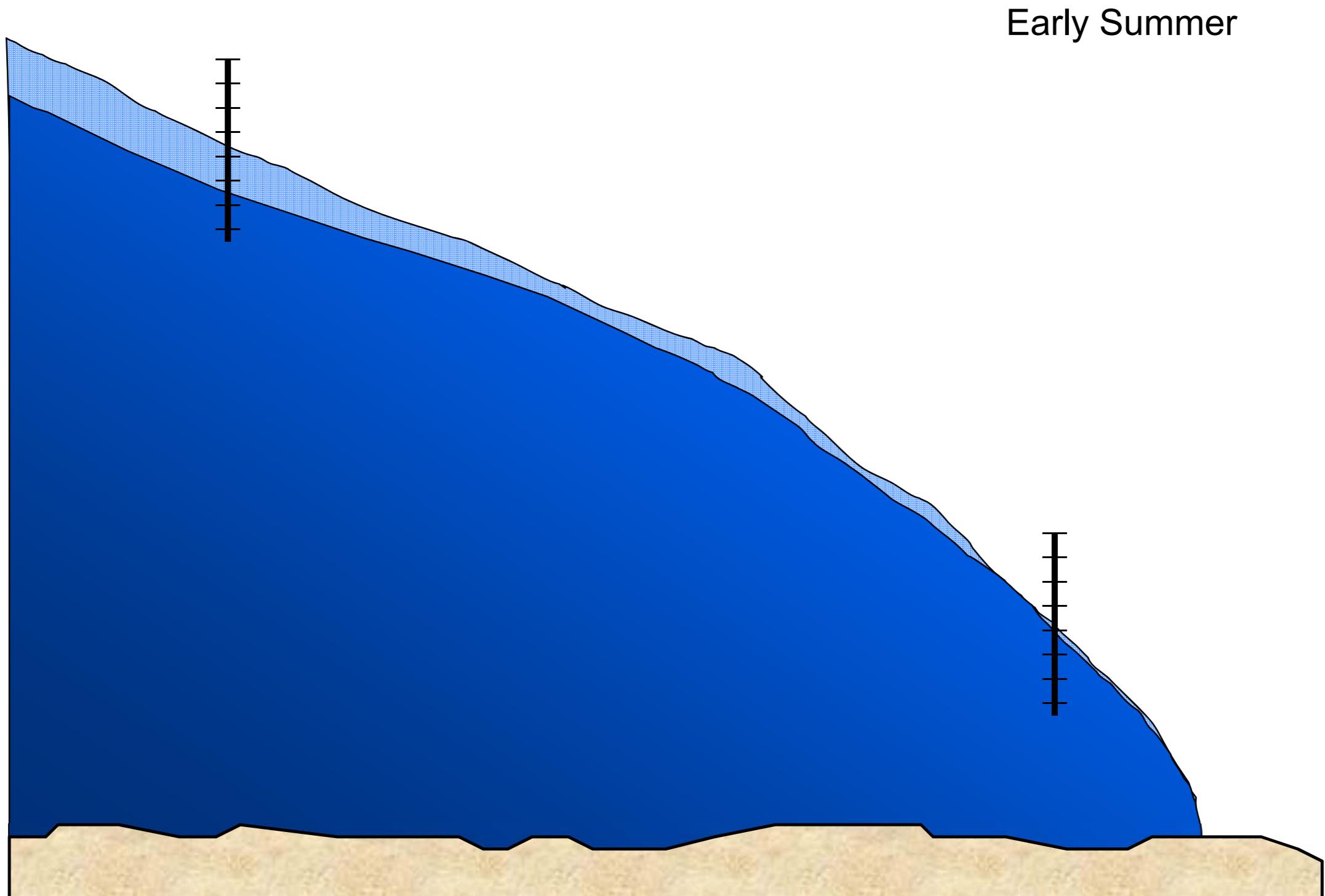
USGS 1960

Surface mass balance is strongly seasonal

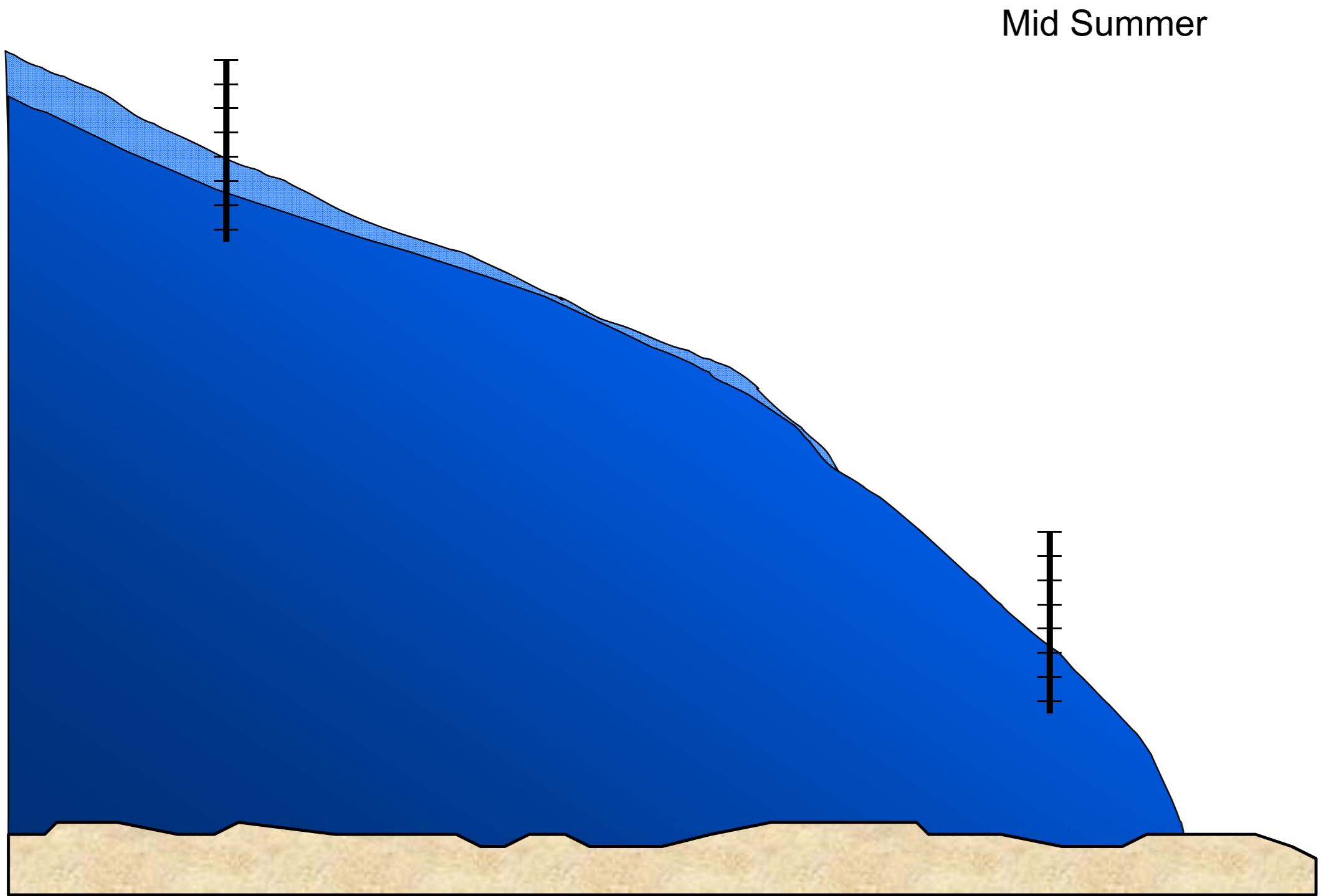
Migration of the Snow Line



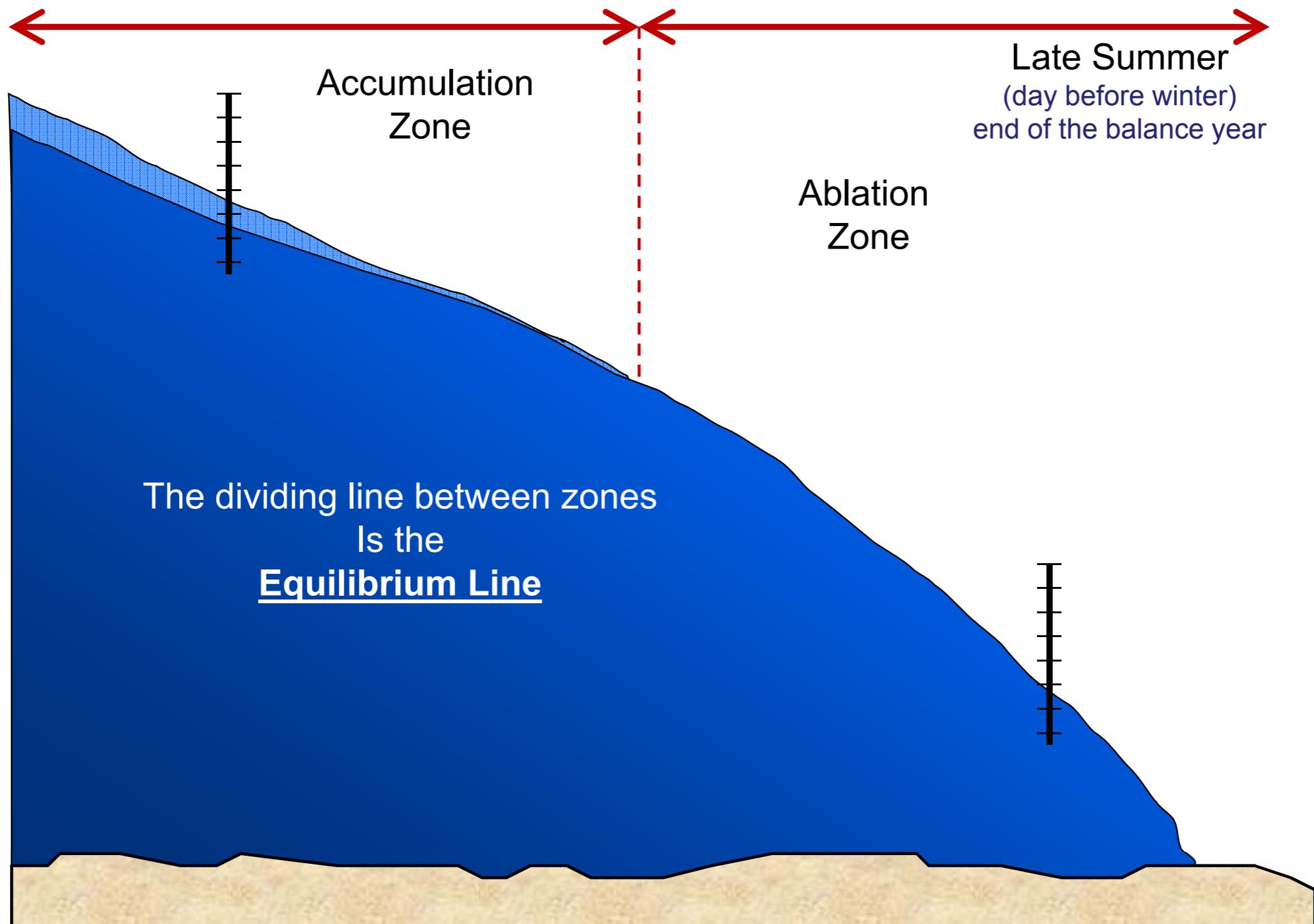
Surface mass balance is strongly seasonal



Surface mass balance is strongly seasonal



Surface mass balance is strongly seasonal



How mass gets in

- Snowfall
- Water freezing in the snowpack
- Avalanches
- Rime formation

Snow does not just fall

Wind transport mechanisms



1-10cm



ROLLING

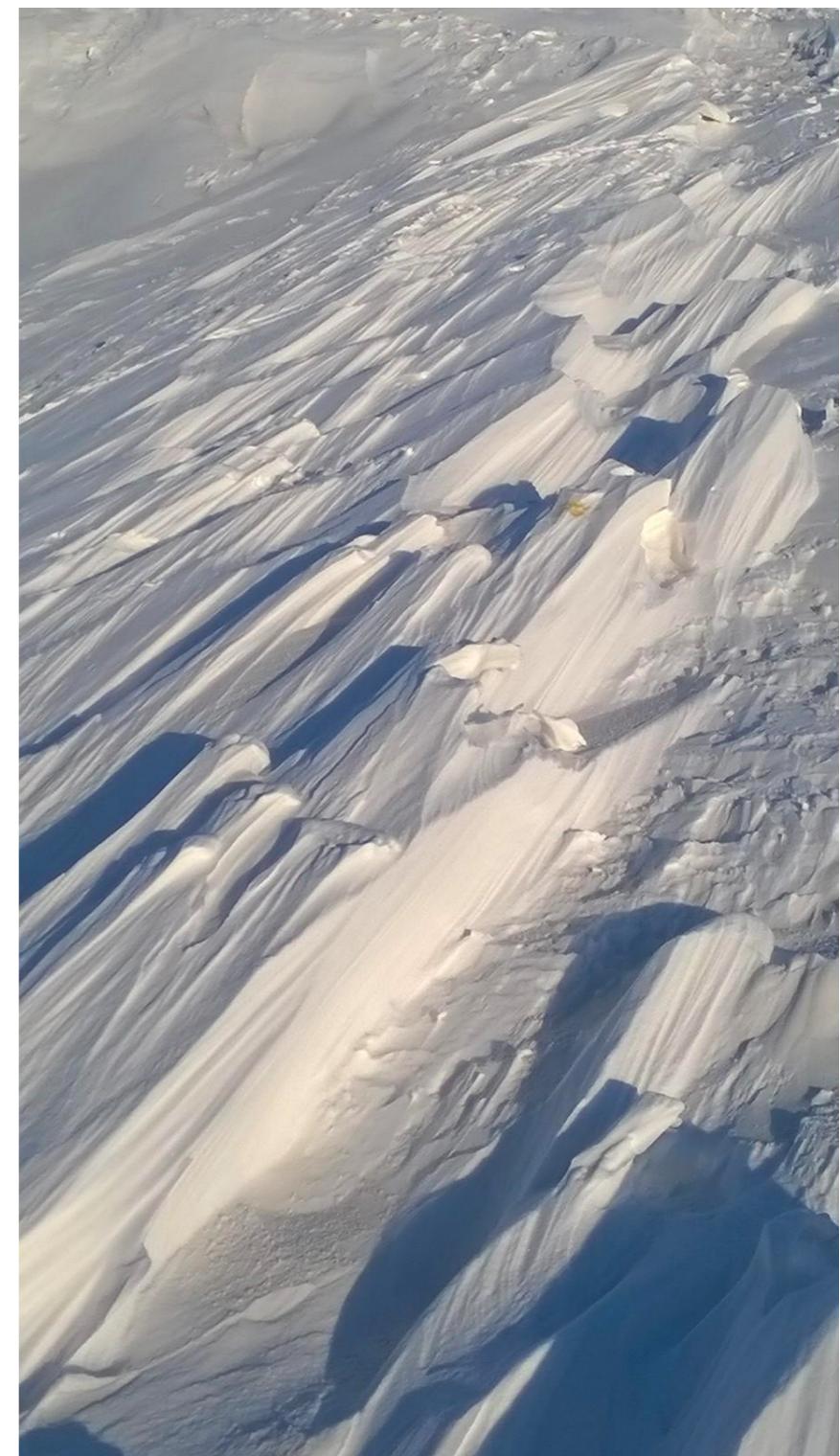
SALTATION



TURBULENT
SUSPENSION

Sastrugi

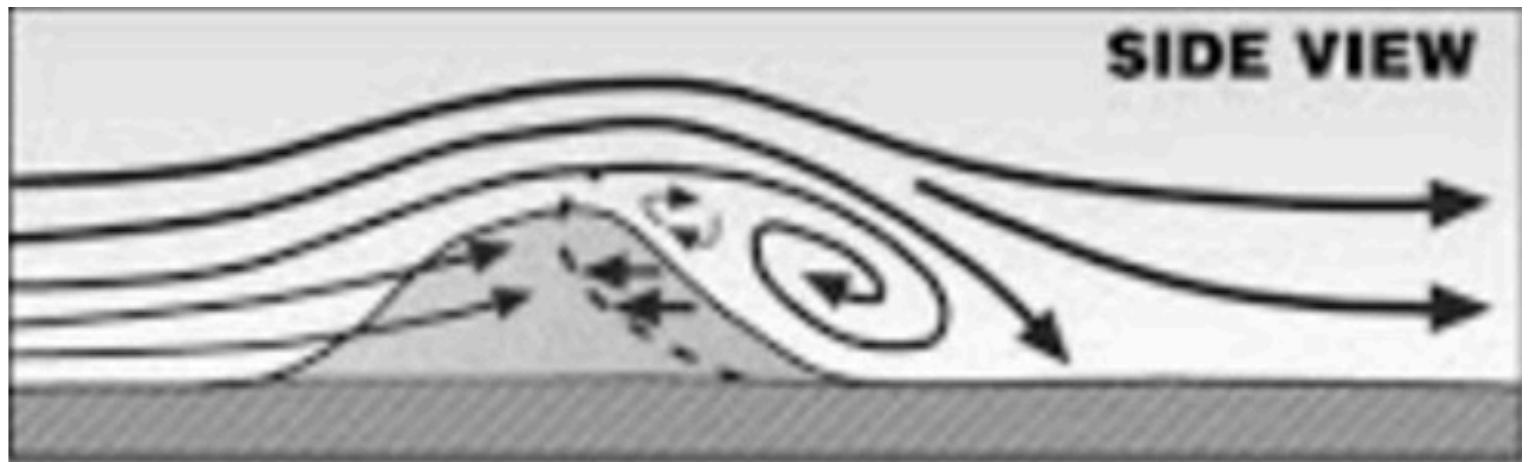
- Irregular grooves and ridges on snow surface
- Formed by wind erosion, saltation, and deposition



Snow dunes



- Very similar to aeolian sand dunes in formation, form
- Snow dunes consist entirely of eroded snow, ‘raised’ by blizzard winds
- Saltating snow forms moving dunes



Mountain wind and snow

Snow transport depends on wind speed

- Accelerating air can pick up more snow
- Decelerating air drops some transported snow

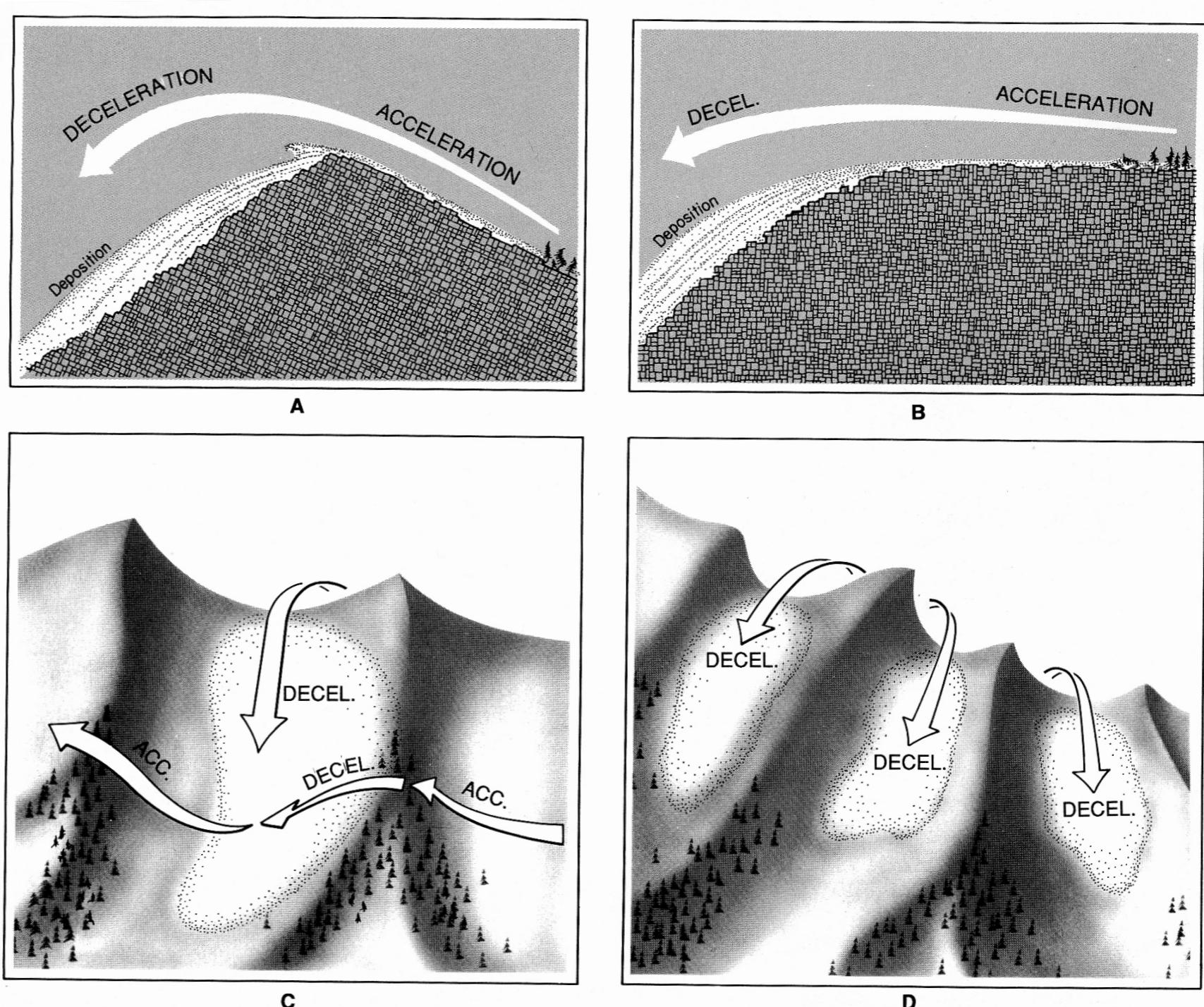


Figure 2.15. Snow is picked up in acceleration regions and deposited in deceleration regions (A, B). This produces lee zone deposition, cross-loading, and deposition in gullies and notches (C, D).

How mass gets out

- Melting and then run off
- Evaporation/Sublimation
- Wind scour, blowing snow
- Ocean heating at ice-water interface
 - Turbulent boundary layer determines rate of heat transfer (Jenkins 1991, and many others)
- Iceberg calving
 - We will talk about more later in the semester

Energy balance pieces

- Transfer through phase change
 - Latent heat

Typical Latent Heat Fluxes

Air	dry	humid (warm)
	$LH \uparrow$ sublimation evaporation	$LH \downarrow$ deposition condensation
Snow	humid	humid (cold)
Important on wet slopes		Important on high plateaus

Energy balance pieces

- Transfer through phase change
 - Latent heat
- Transfer by conduction
 - The turbulent near-surface boundary layer has a temperature which transfers heat to the ice through sensible heating
- Transfer by radiation

Radiation balance contributes to energy exchange between atmosphere and ice/snow surface

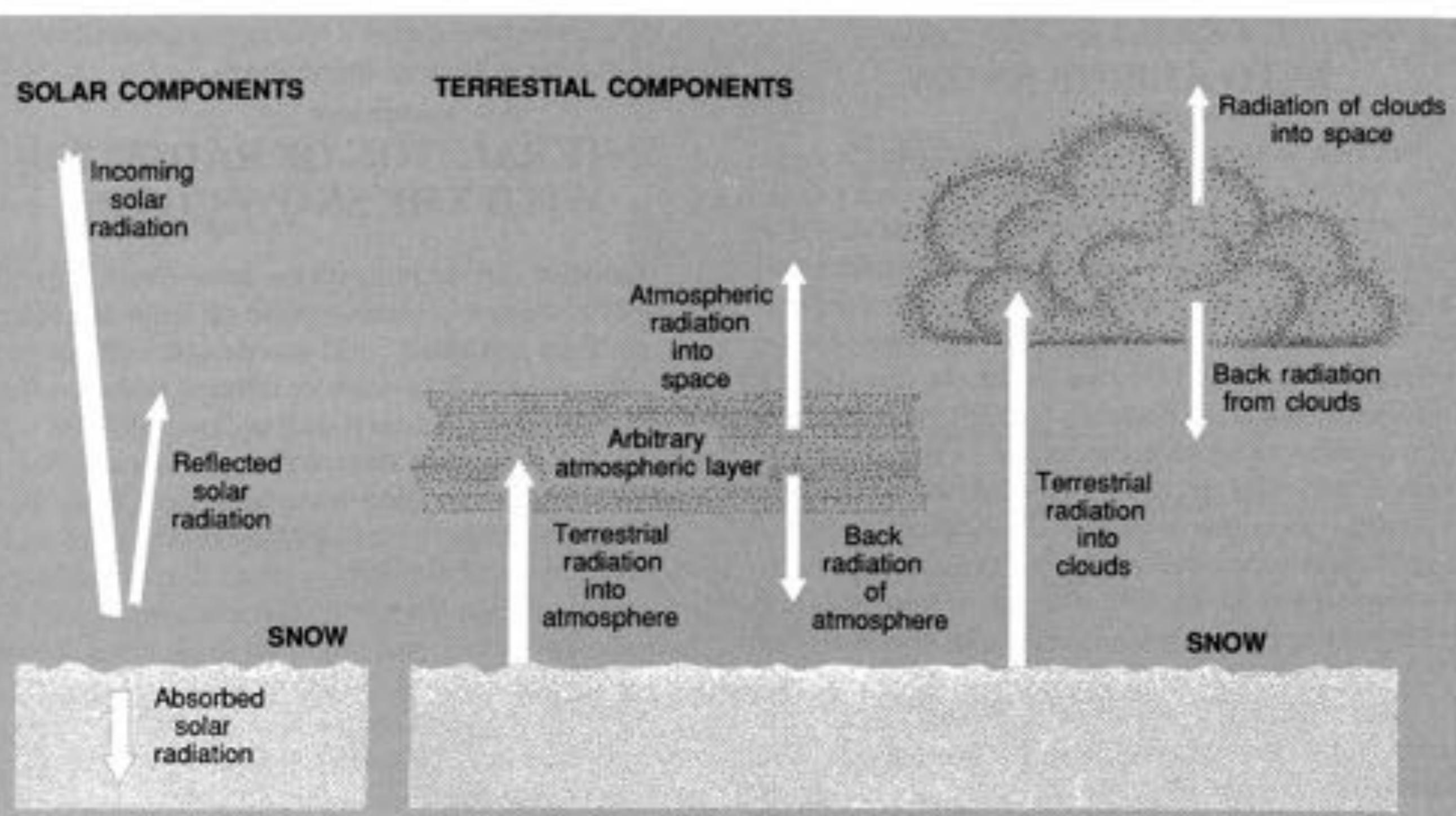


Figure 2.23. Radiation balance at snow surface. Incoming, absorbed, and reflected solar radiation is short wave (visible and ultraviolet). The other components from terrestrial sources are long wave (infrared).

Energy in Melt Season

Source	Typical Importance	Blue Glacier (Mt. Olympus):
Radiation	50-85%	57%
Sensible heat	}	34%
Latent heat		9%

Typically $R > SH > LH$

Surface Mass Balance with temperature

Ablation: warmer air temperature means more melt due to more rapid heat transfer across snow-air boundary layer

Accumulation: warmer air temperature mean more accumulation because warmer air holds more moisture (clausius-claperon), until precipitation starts falling as rain sufficiently above 0 C

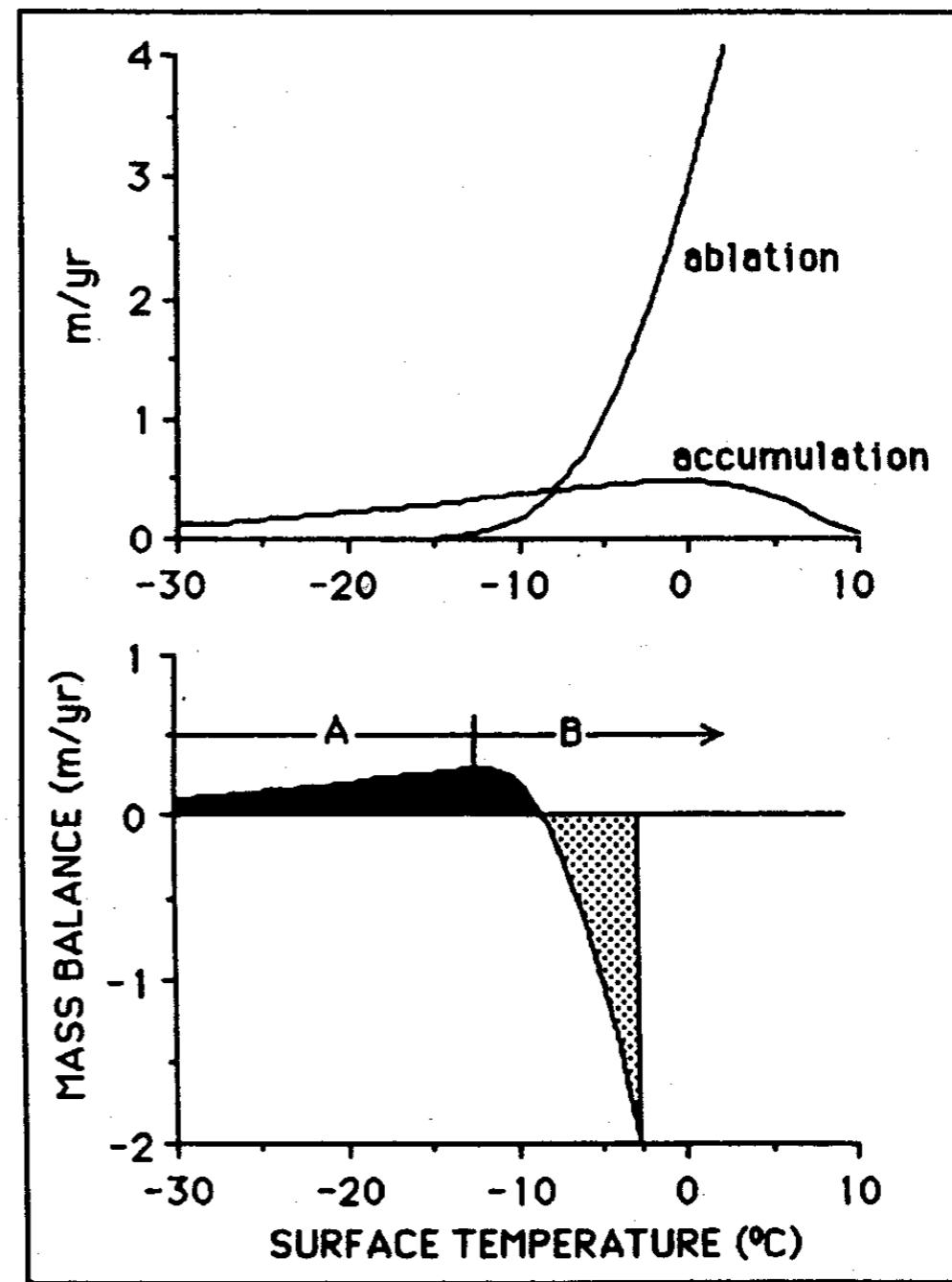
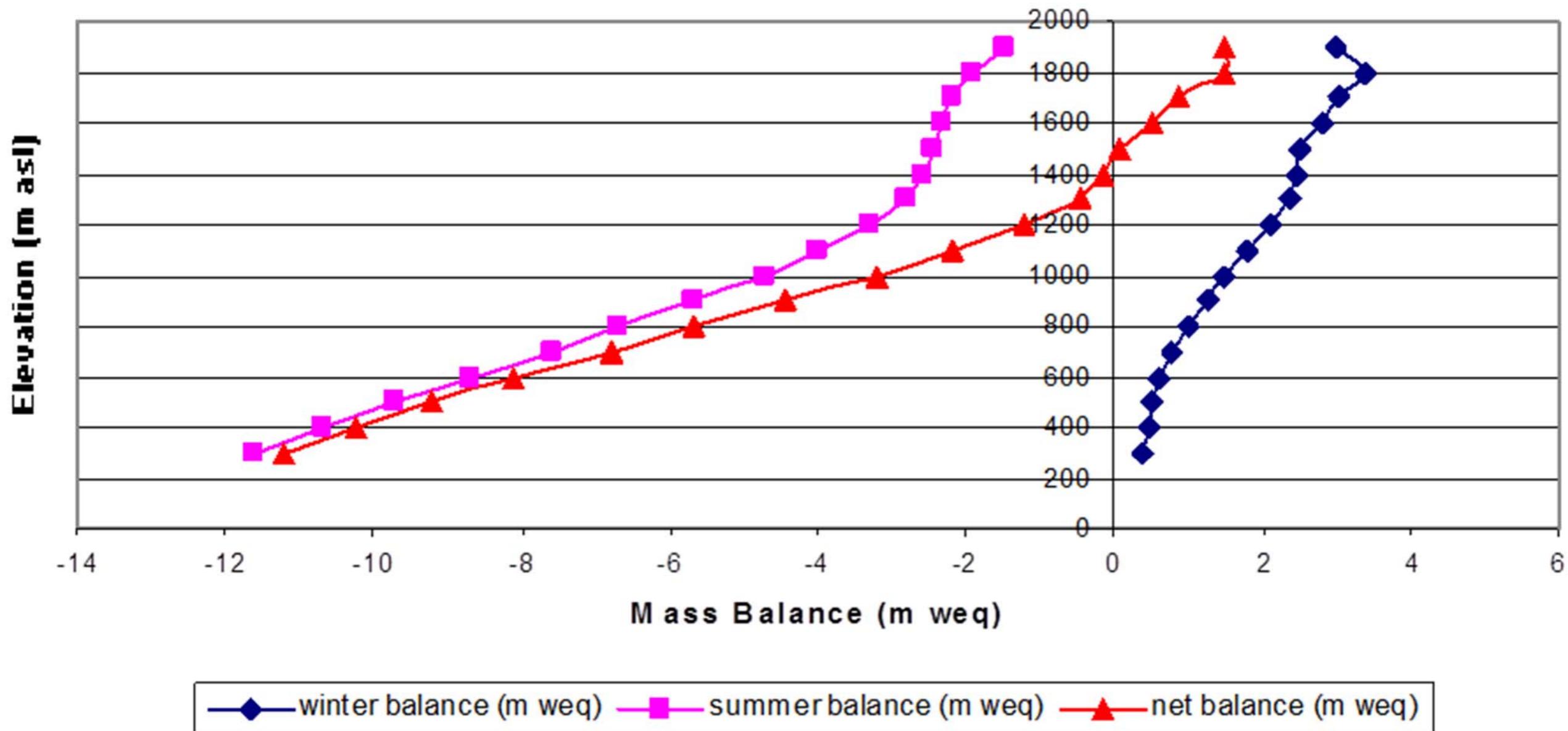


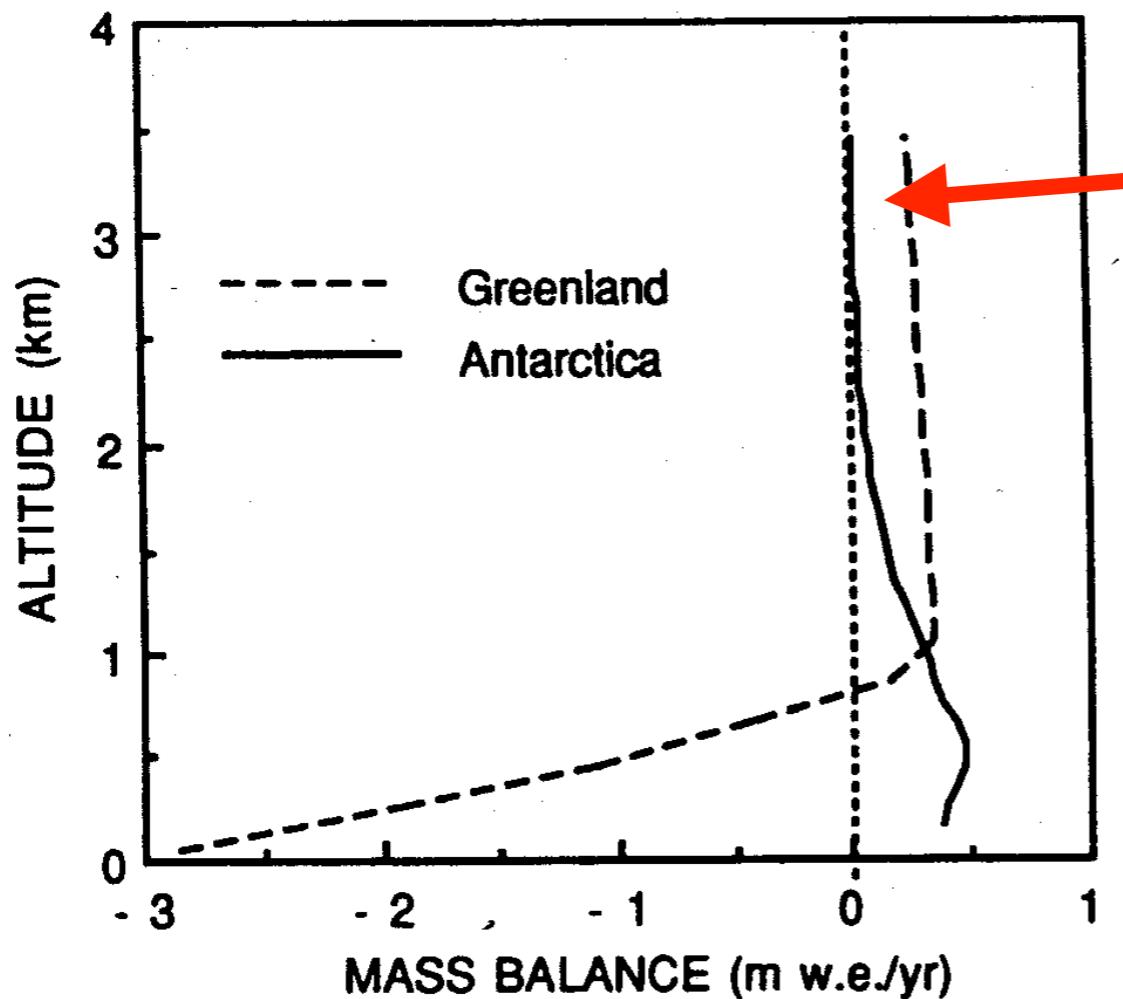
Fig. 3. Components of glacier mass balance at a fixed location in dependence of the annual mean surface temperature, with a given constant annual temperature range. The net balance (lower panel) shows a region where it *increases* with temperature (A) as well as a region where it *decreases* with temperature (B). The Antarctic ice sheet is in region A, the Greenland ice sheet mainly, but not entirely,

Surface Mass Balance with elevation

Mass Balance Nigardsbreen (1996-1997)



Surface Mass Balance with elevation



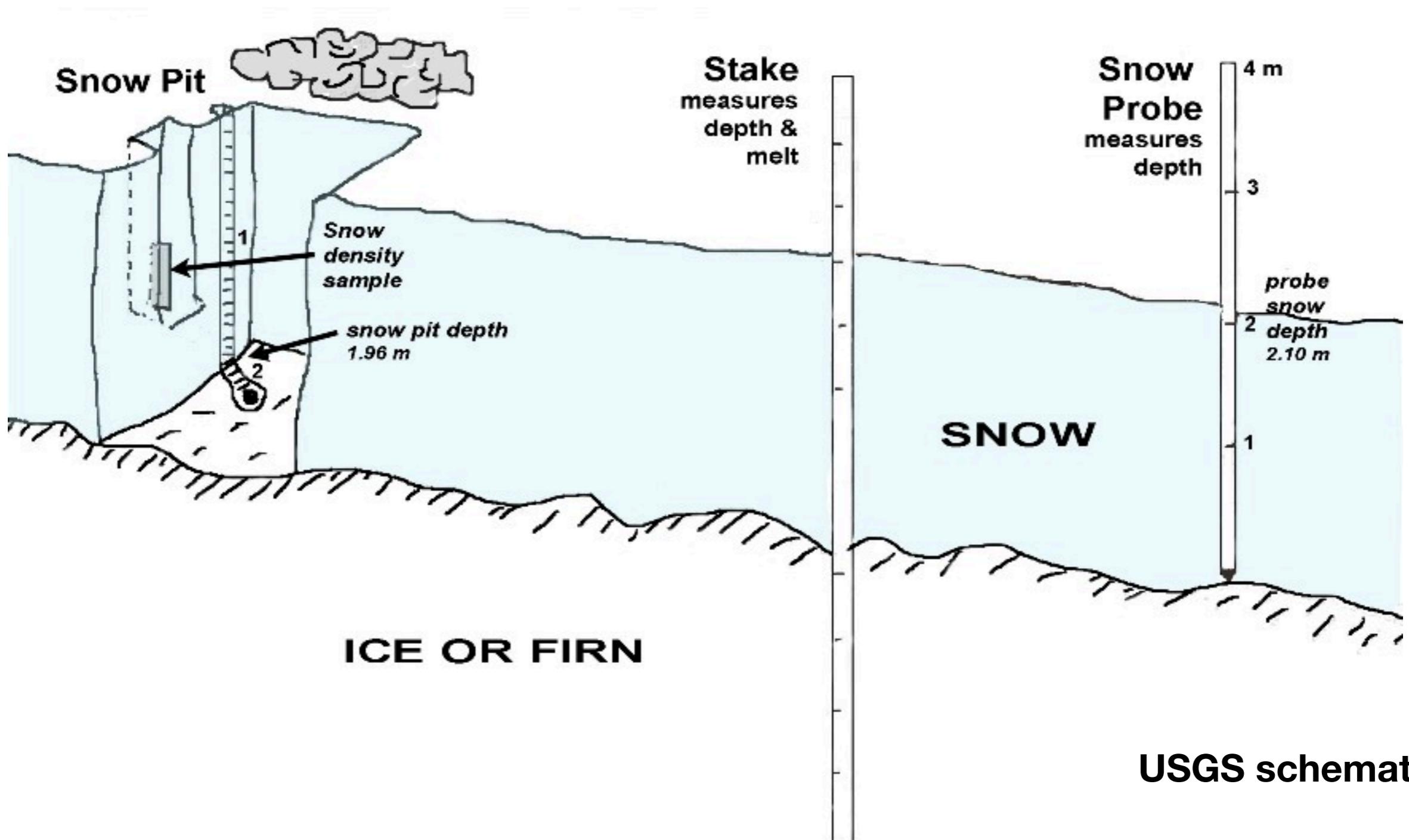
Elevation-Desert effect:
less moisture at higher
elevations (only
matters in thick polar
ice sheets)

Fig. 2. Mass balance as a function of elevation for the Antarctic and Greenland ice sheets. The curves represent mean values for the entire ice sheets, local conditions may be quite different. Based on Drewry (pers. comm. 1988) and Oerlemans et al. (1990).

Height-mass balance feedback

To the board! (and MATLAB)

Measuring surface mass balance



Measuring surface mass balance



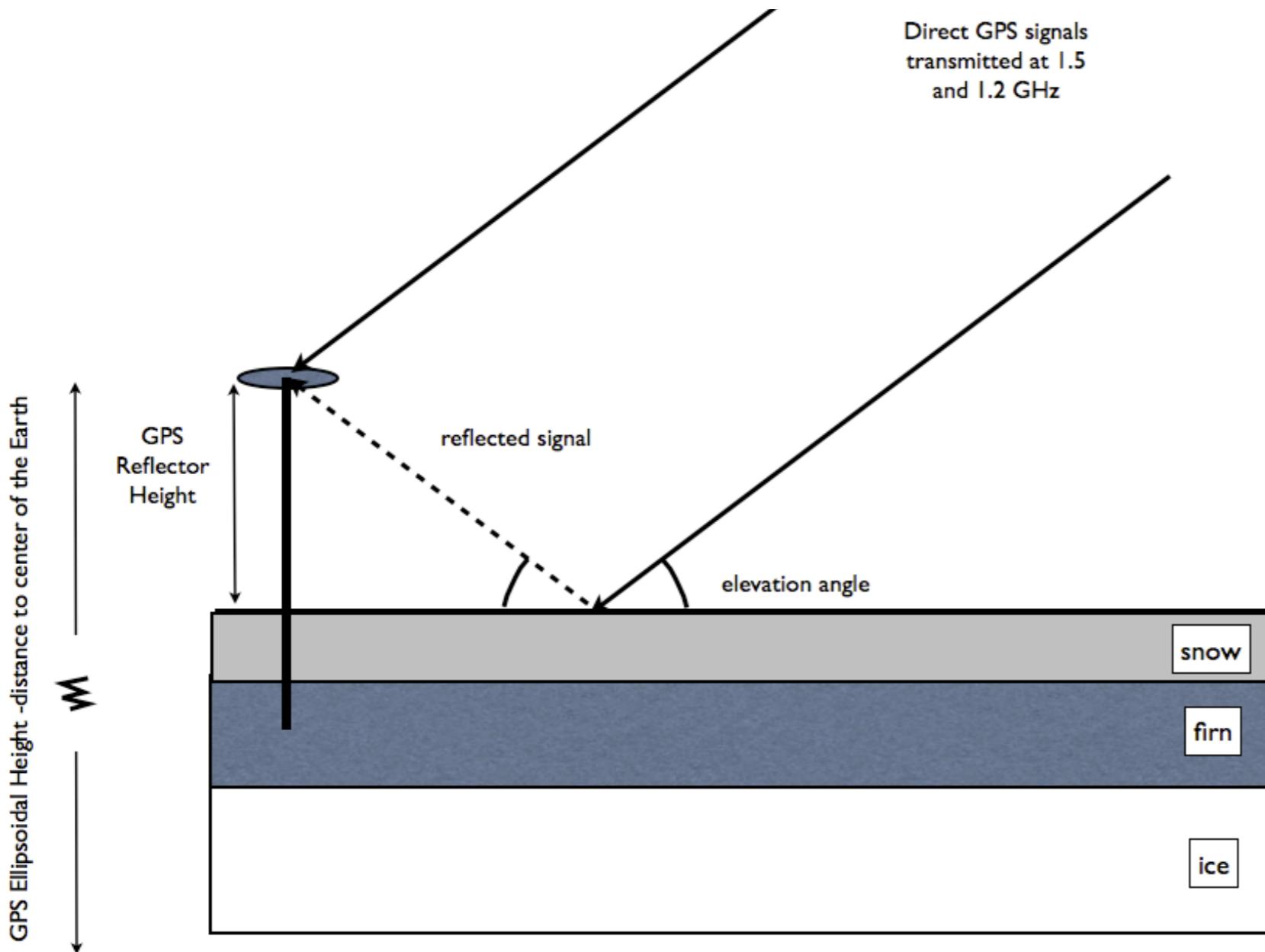
**Snow Pits
retrospectively
determine
accumulation since
last summer**

Measuring surface mass balance

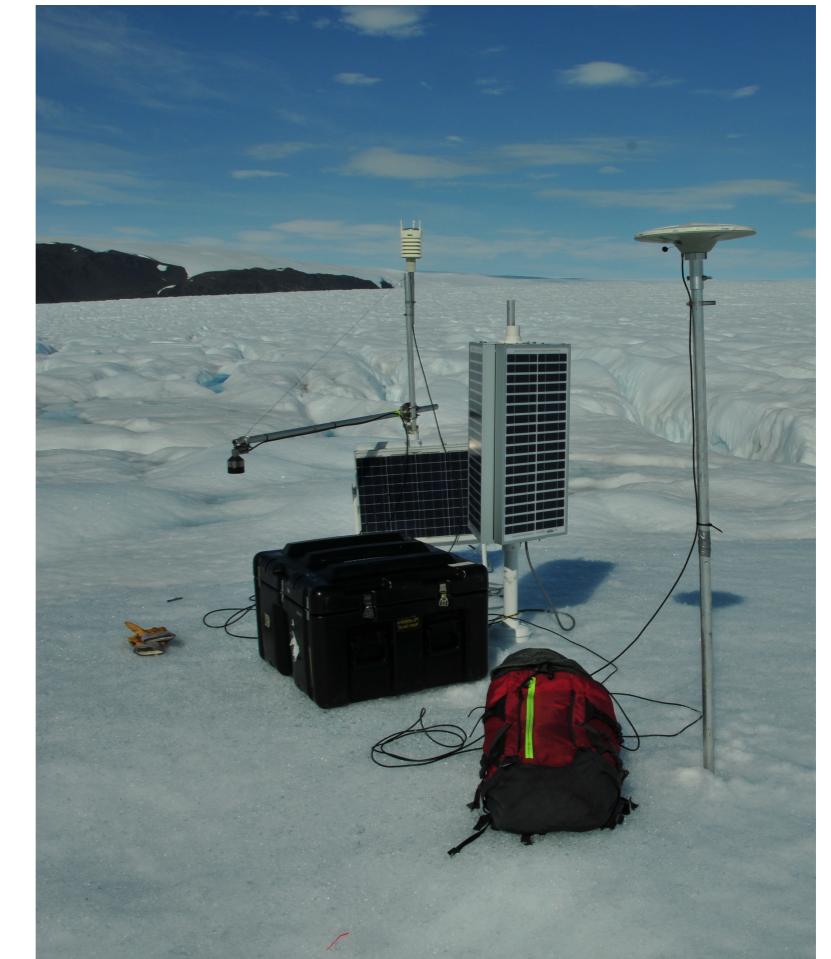


**Snow Stakes
less labor-
intensive**

Measuring surface mass balance

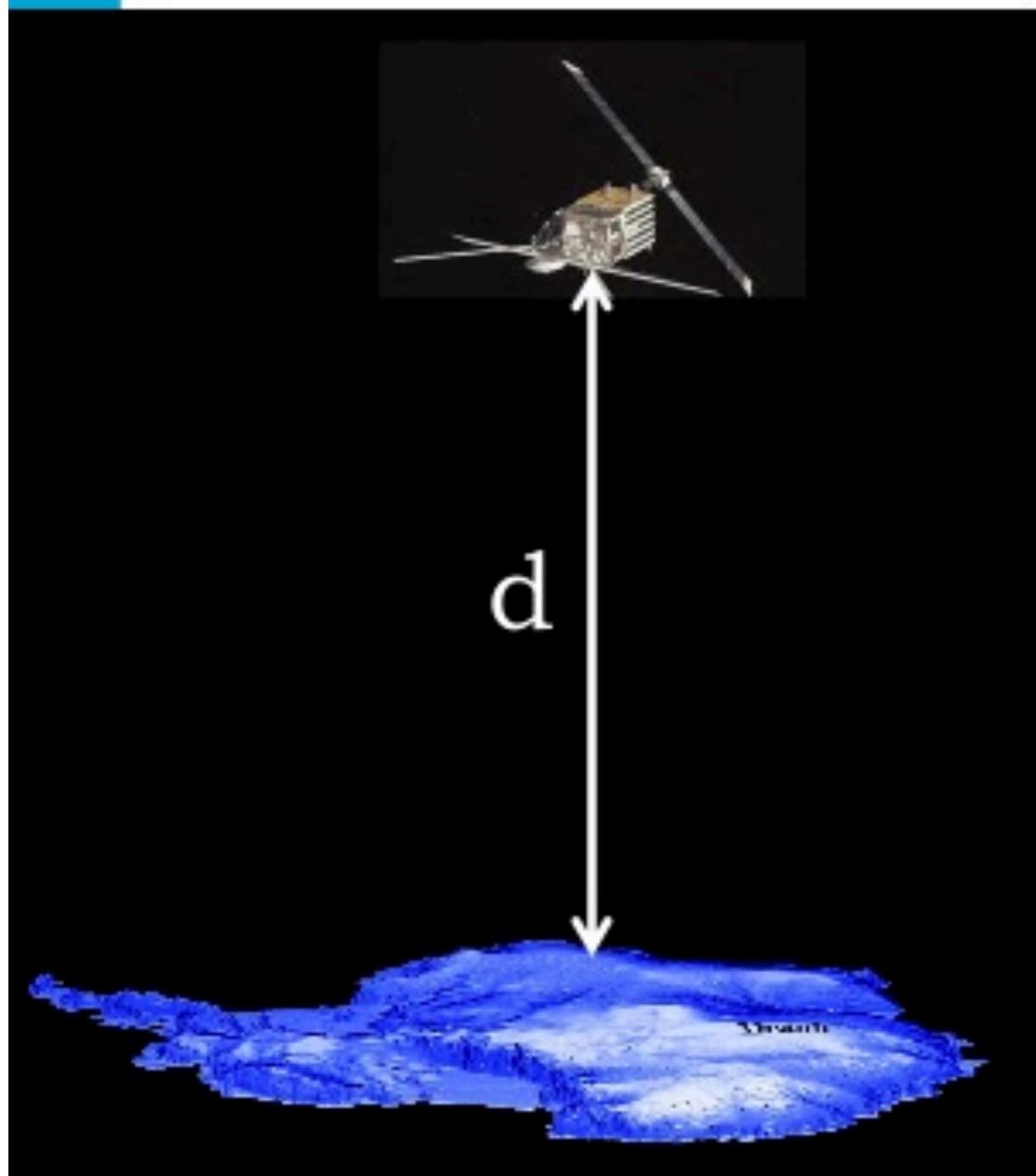


**GPS reflectometry:
more accurate, but
requires power**



Satellite Laser Altimetry

Satellite altimetry



$$h = \text{Alt} - d$$

Elevation Satellite altitude

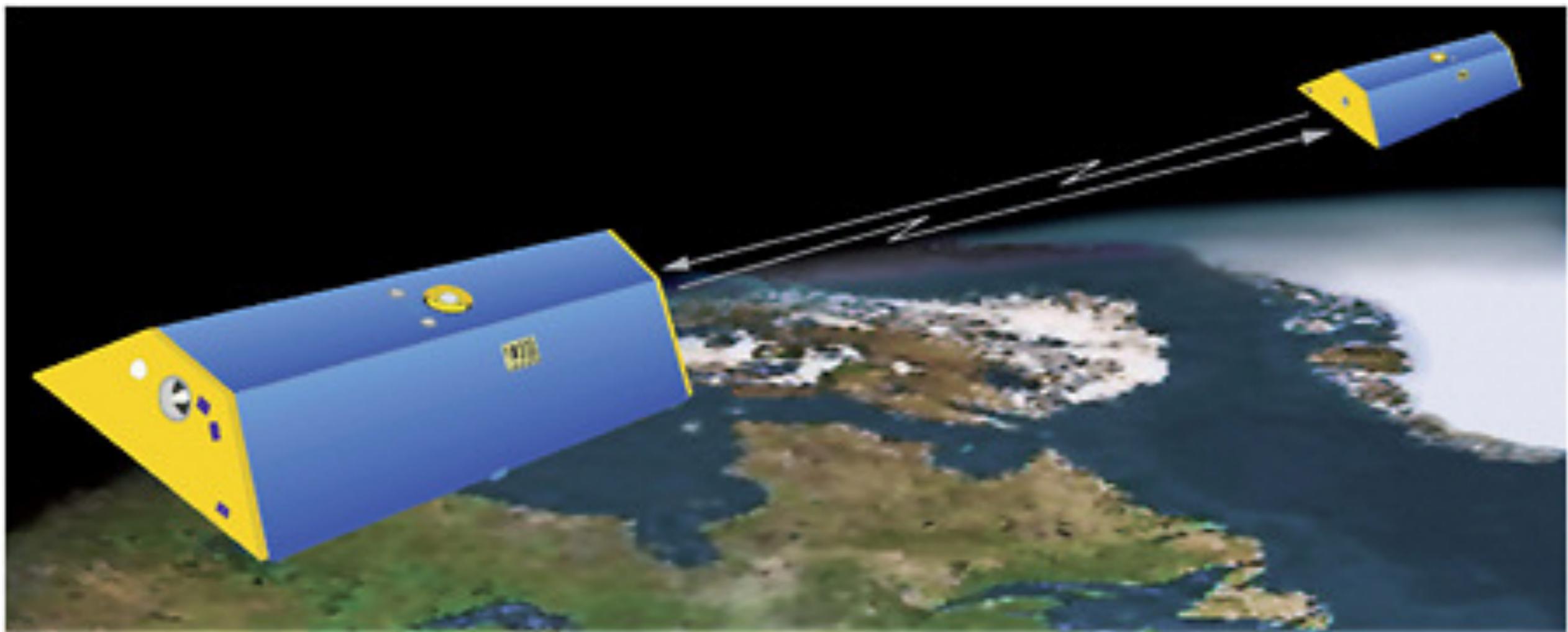
$$\Delta V = \Delta h * S$$

Volume change Elevation change Area

The diagram includes two mathematical equations. The first equation, $h = \text{Alt} - d$, has 'Elevation' and 'Satellite altitude' circled in red with arrows pointing to their respective terms. The second equation, $\Delta V = \Delta h * S$, has 'Volume change', 'Elevation change', and 'Area' circled in red with arrows pointing to their respective terms.

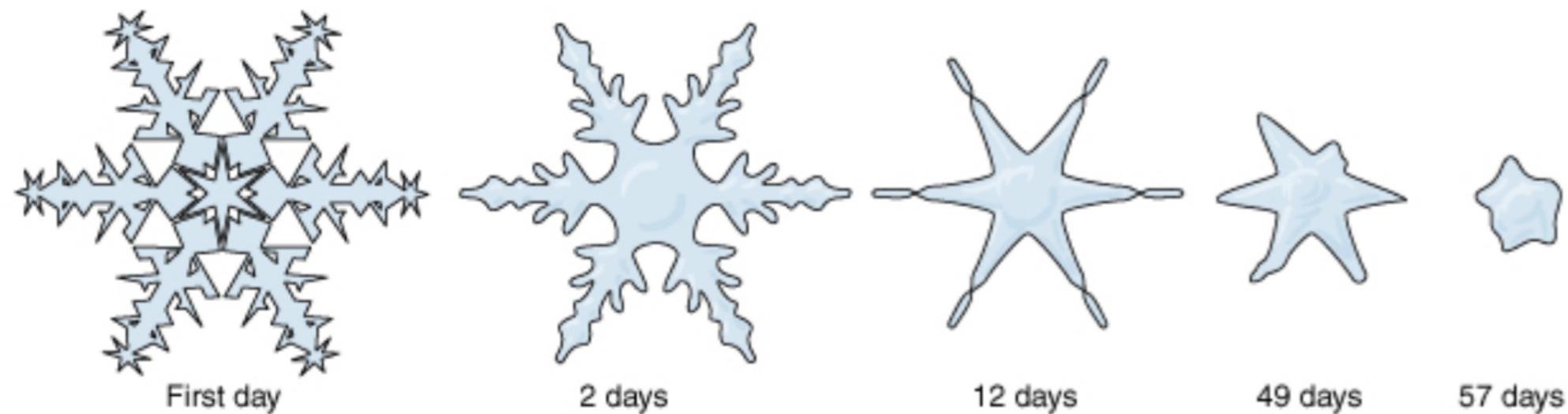
Gravity measurements

GRACE satellites measures small spatial variations in gravitational attraction by using very precise positioning



How snow becomes ice

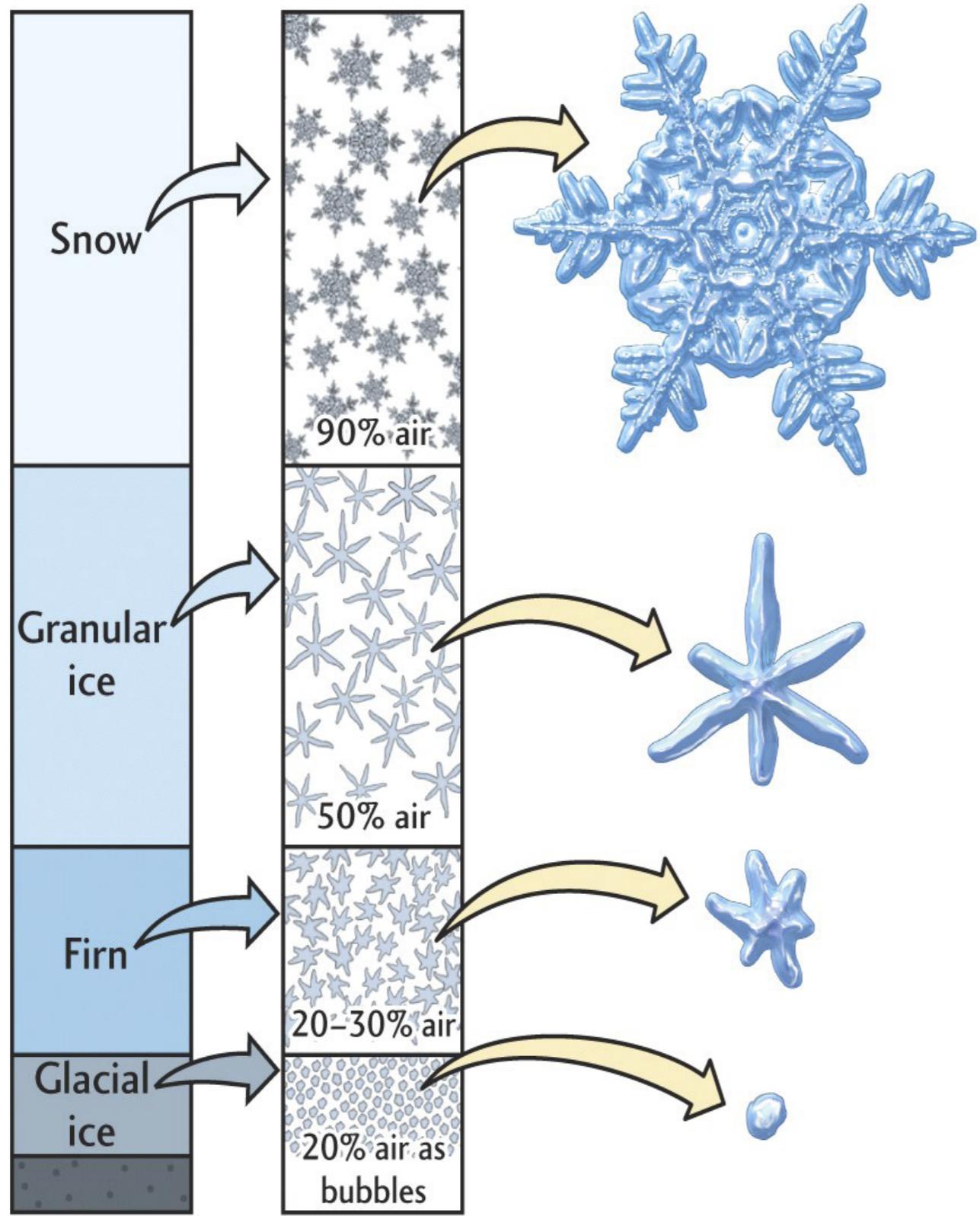
new snow —————→ old snow
(firn)



Copyright 1999 John Wiley and Sons, Inc. All rights reserved.

The Answer: Compression

Firn compaction



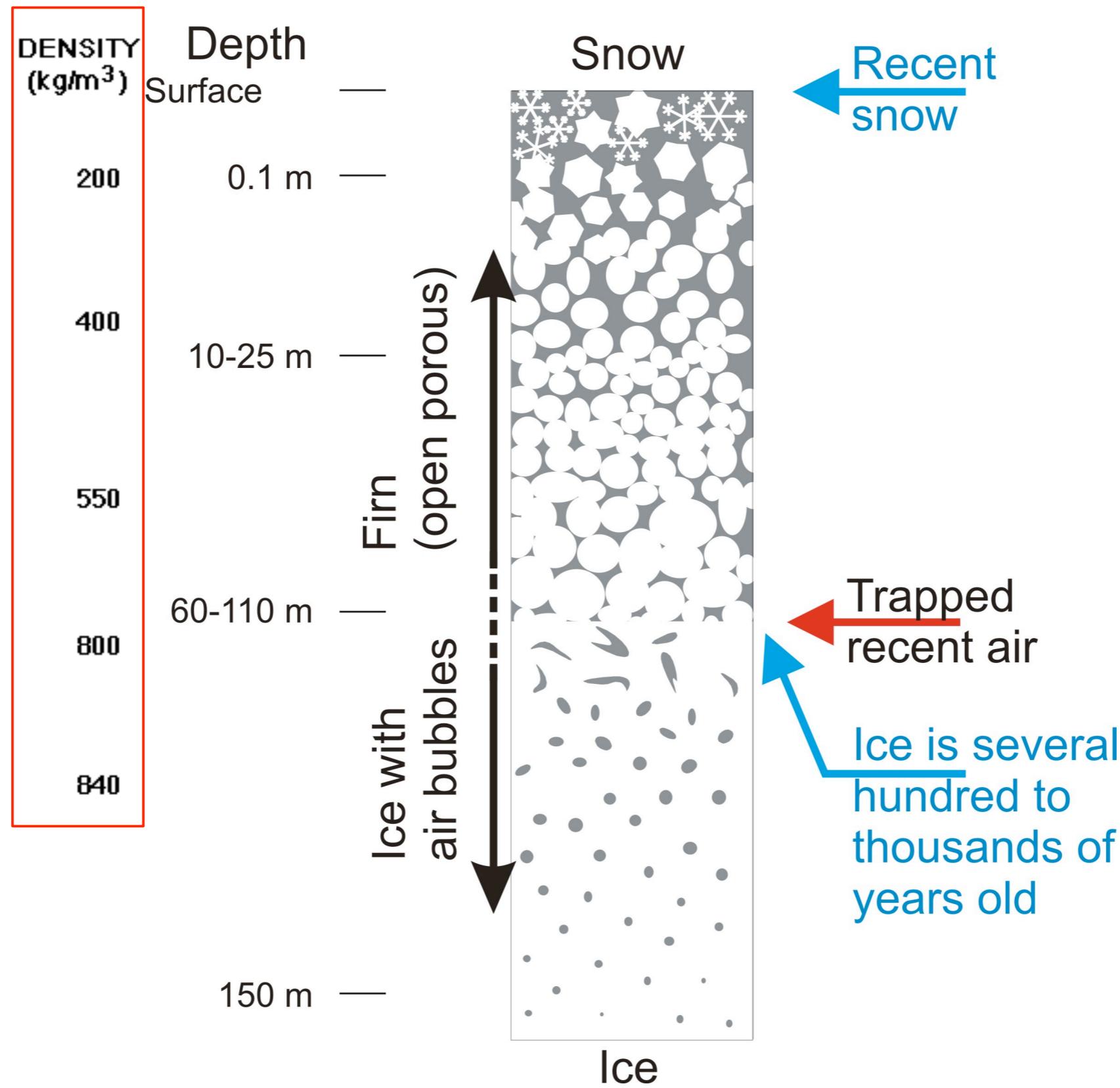
Snow
(50-70 kg/m³)

Firn: the intermediate stage between low-density snow crystals and high density glacial ice

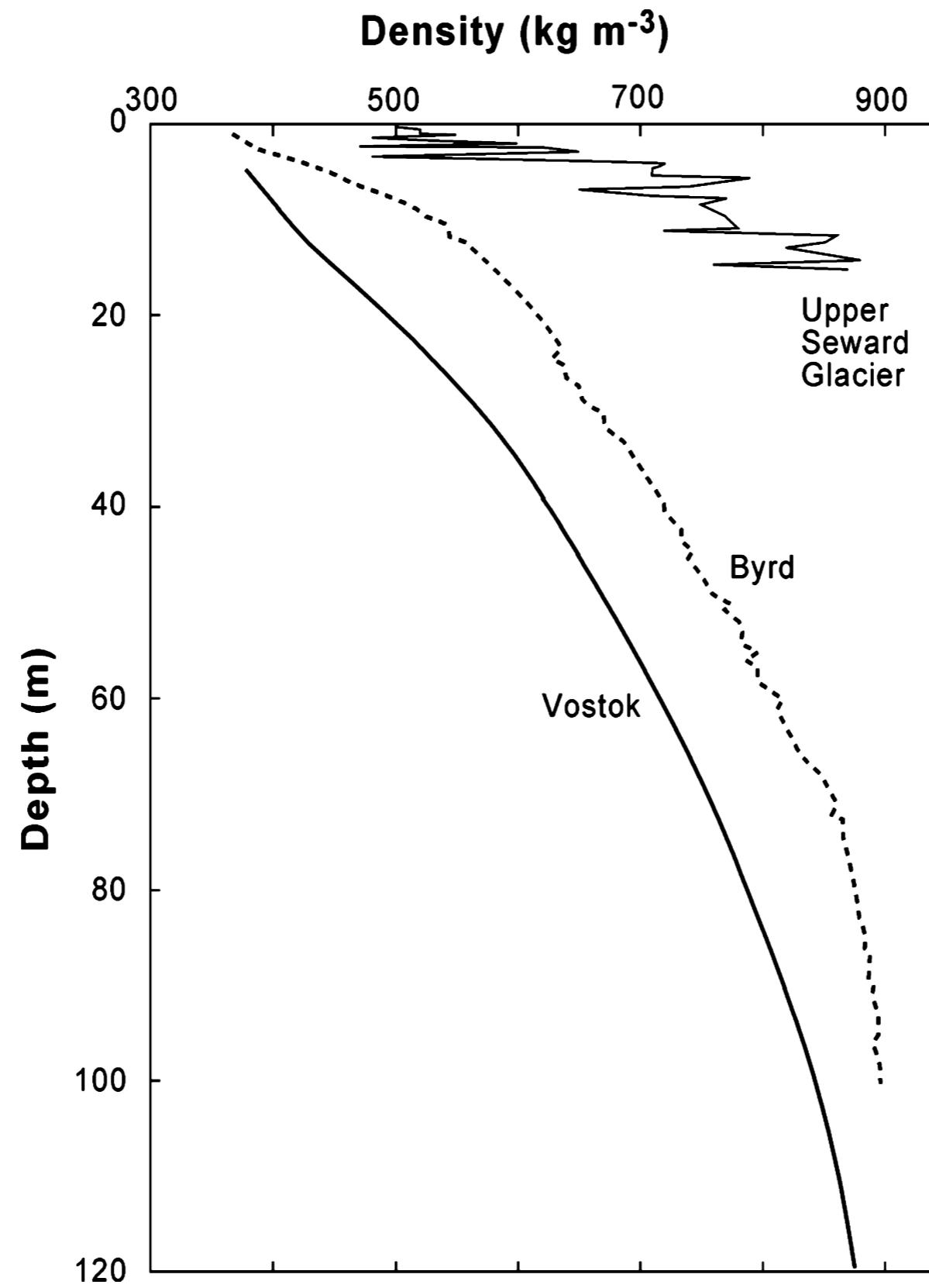
Firn
(400-800 kg/m³)

Ice
(800-930 kg/m³)

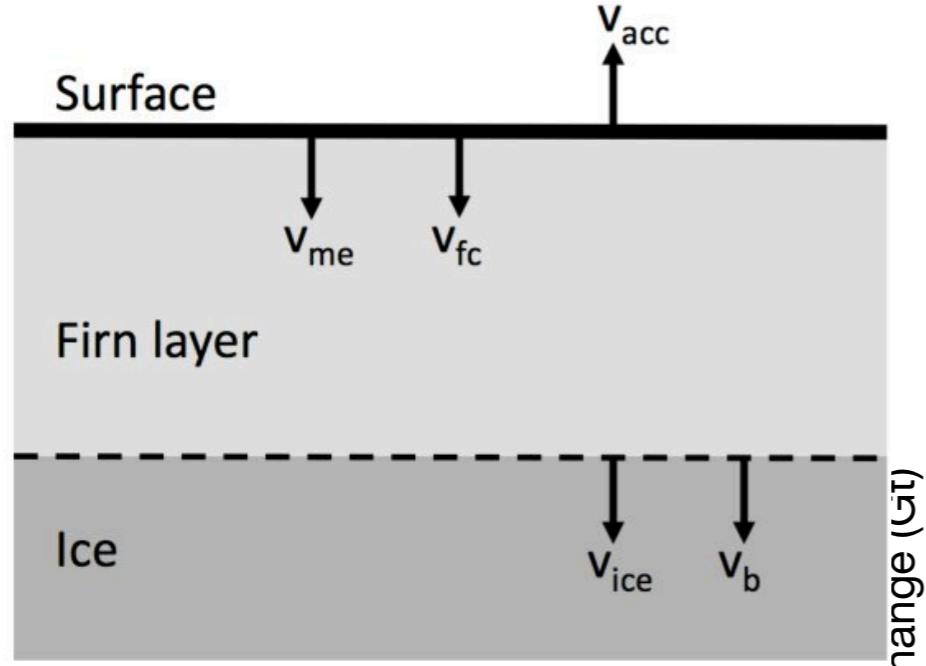
How snow becomes ice



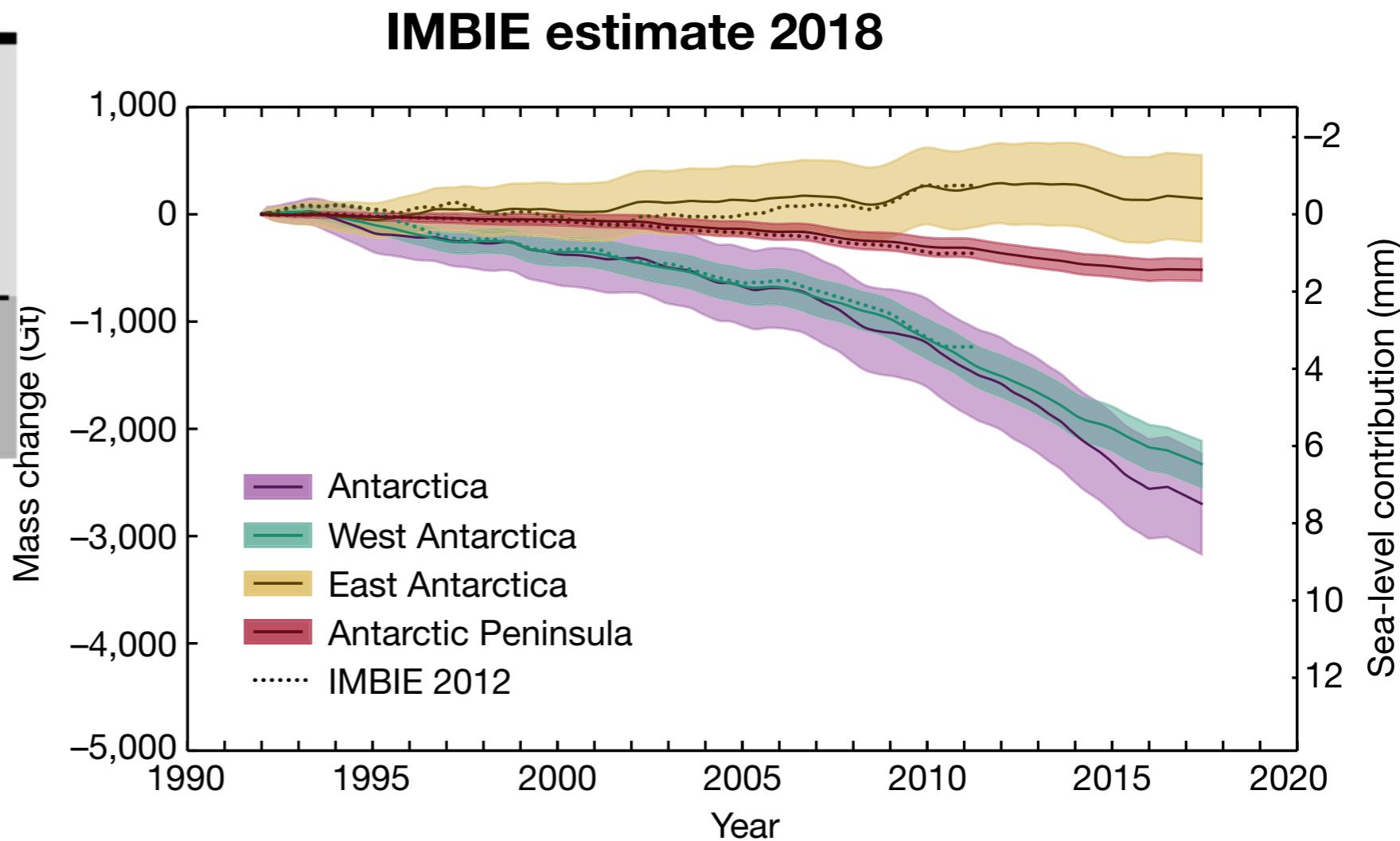
How snow becomes ice



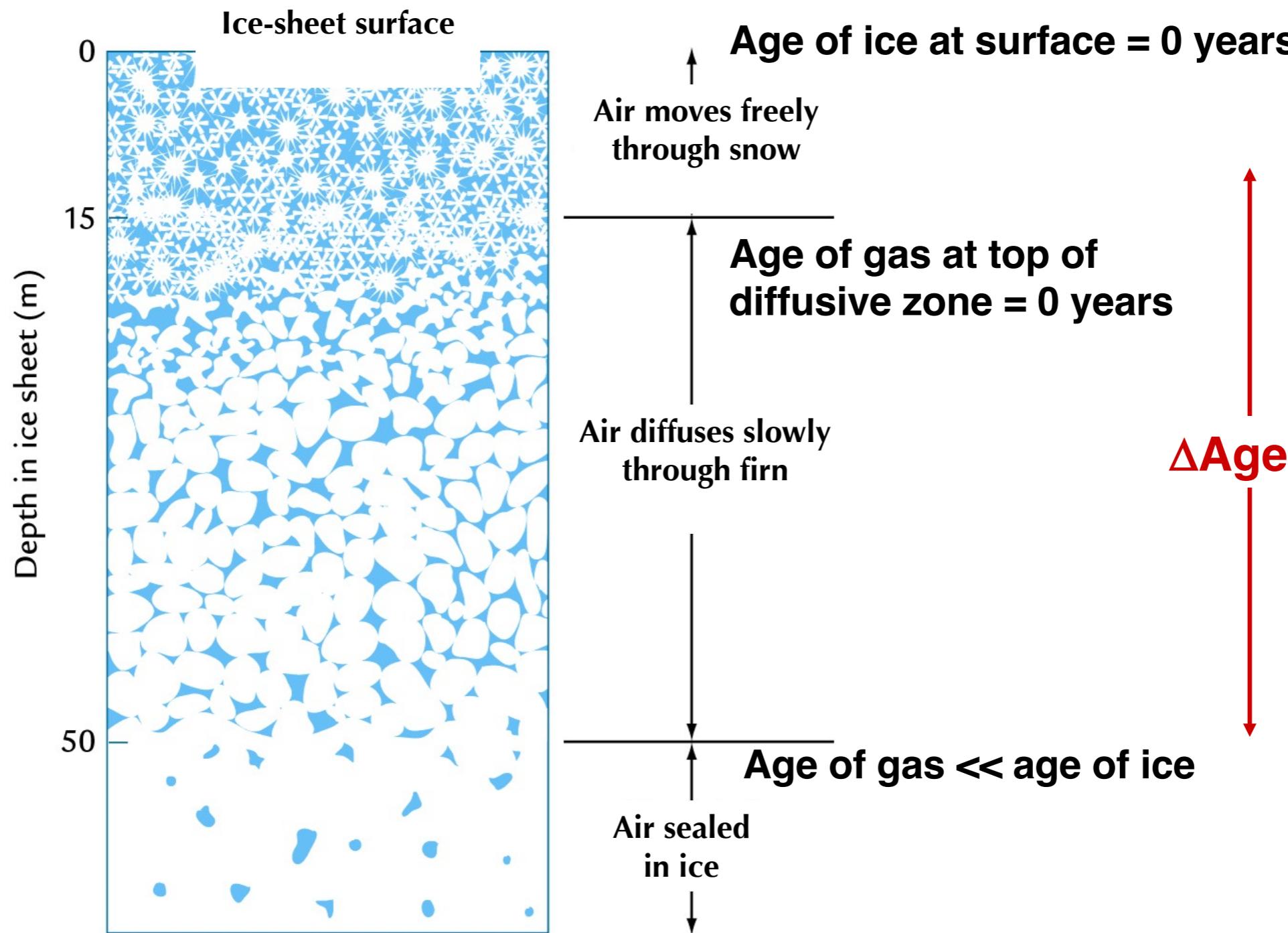
Understanding firn compaction is critical for interpreting elevation changes measured from satellites



Ligtenberg et al. 2011



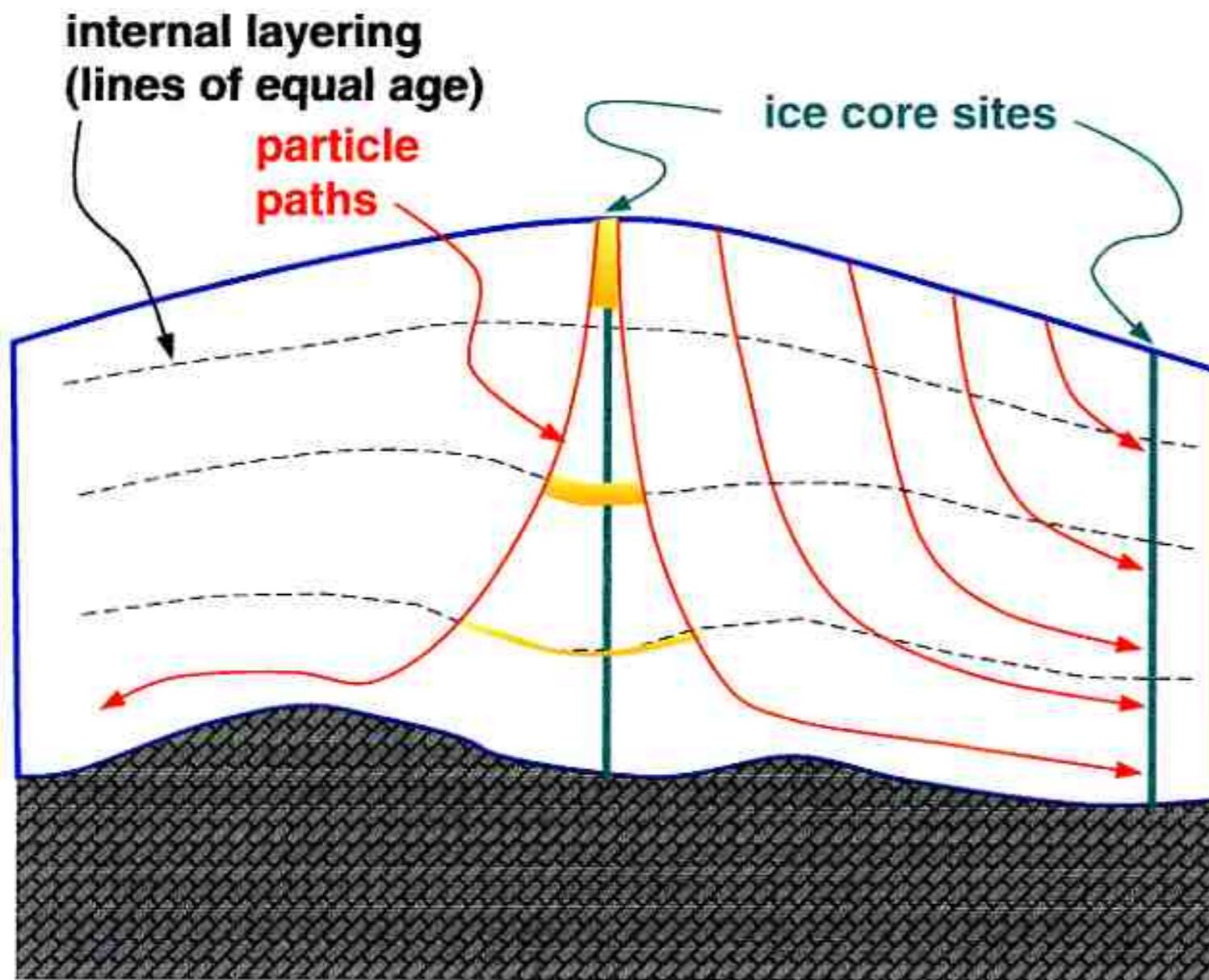
Firn traps atmospheric gases



Choosing an Ice-Core Site

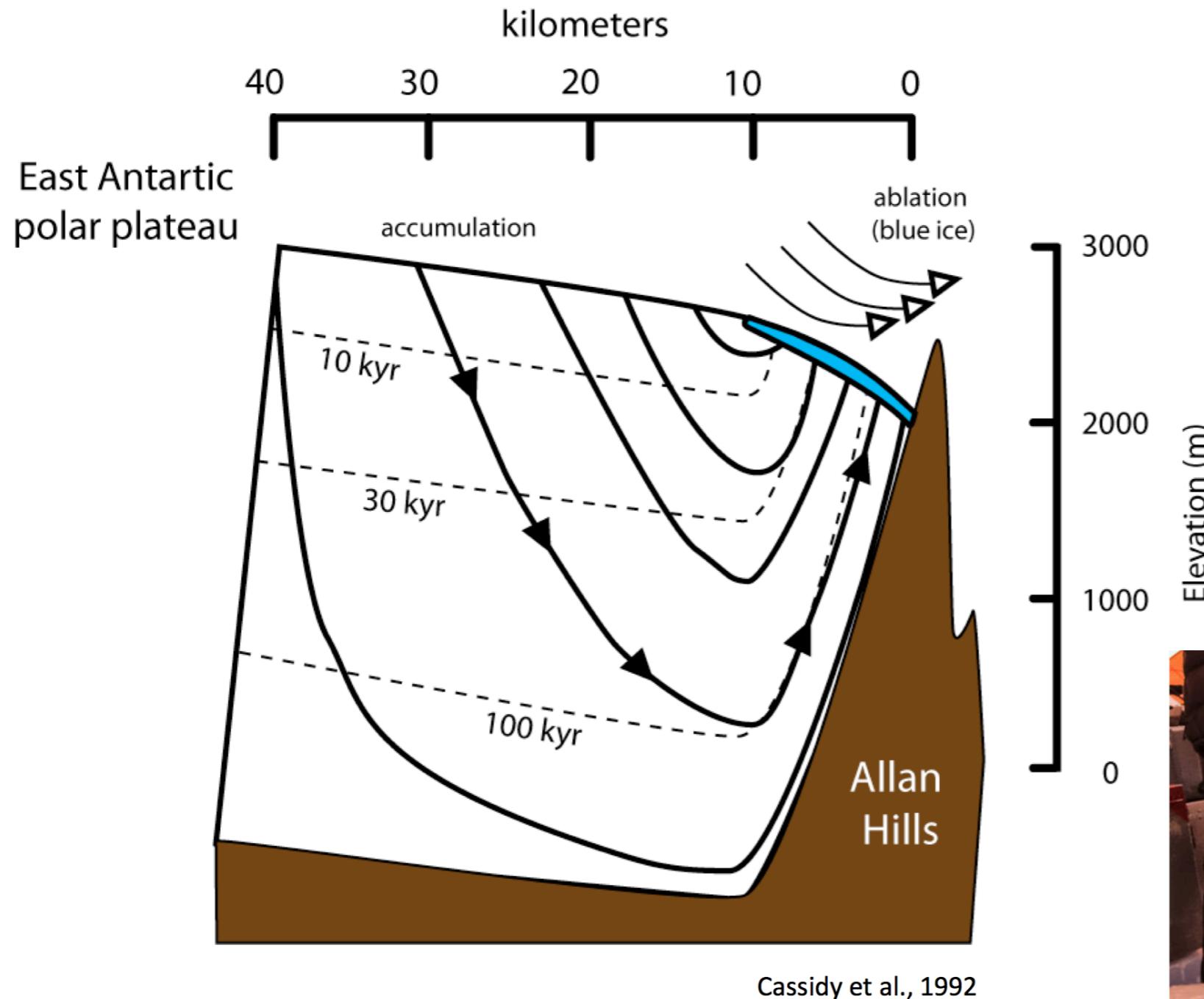
Near the summit is good.

- Climate signal relates to the same place for ice of all ages.

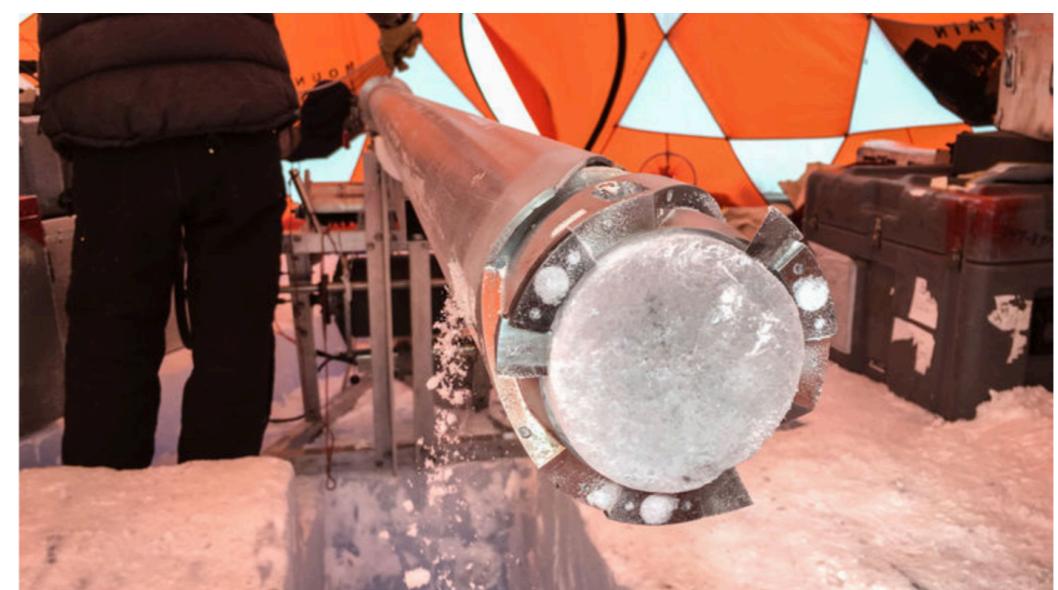


Layers get thinner over time due to flow.

But also, you can choose sites away from the divide to find very old ice



The closer you get to the bedrock in the hills, you get small chunks of older preserved ice



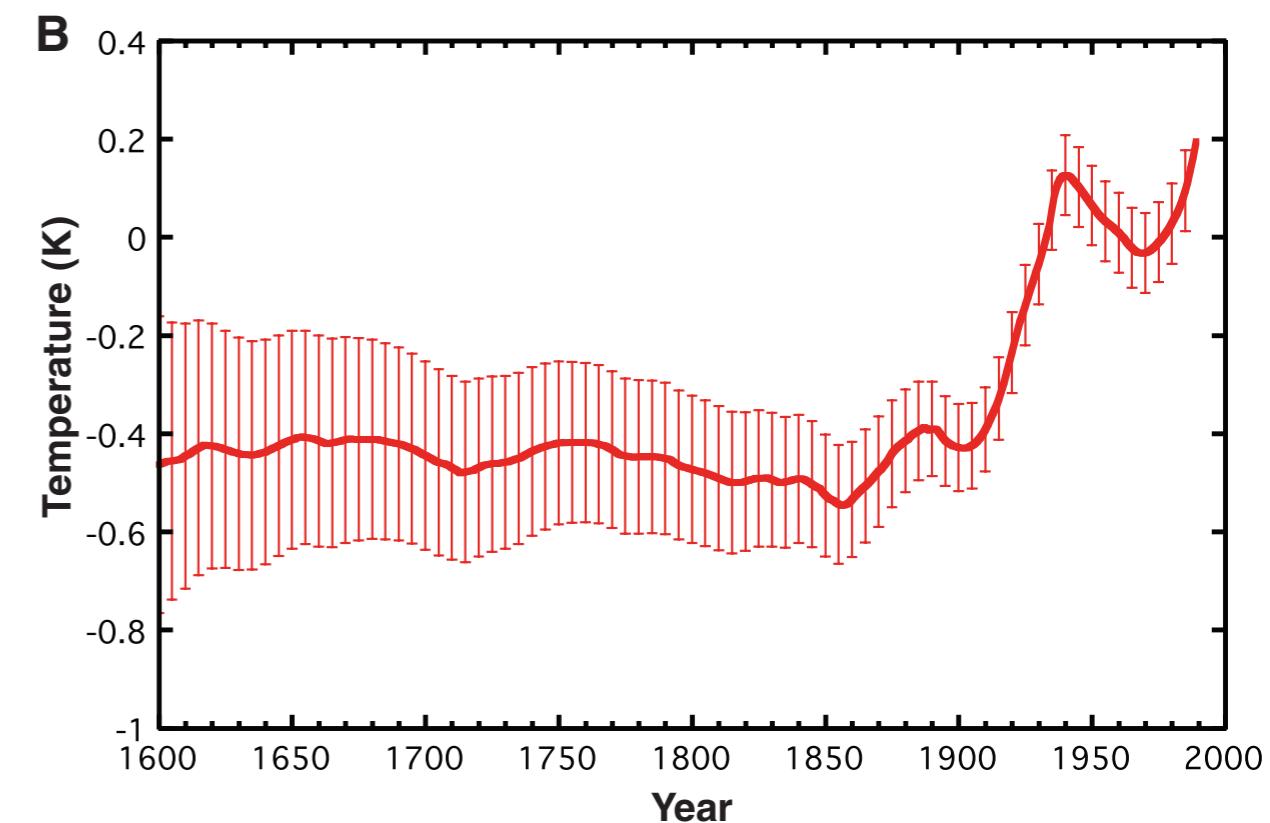
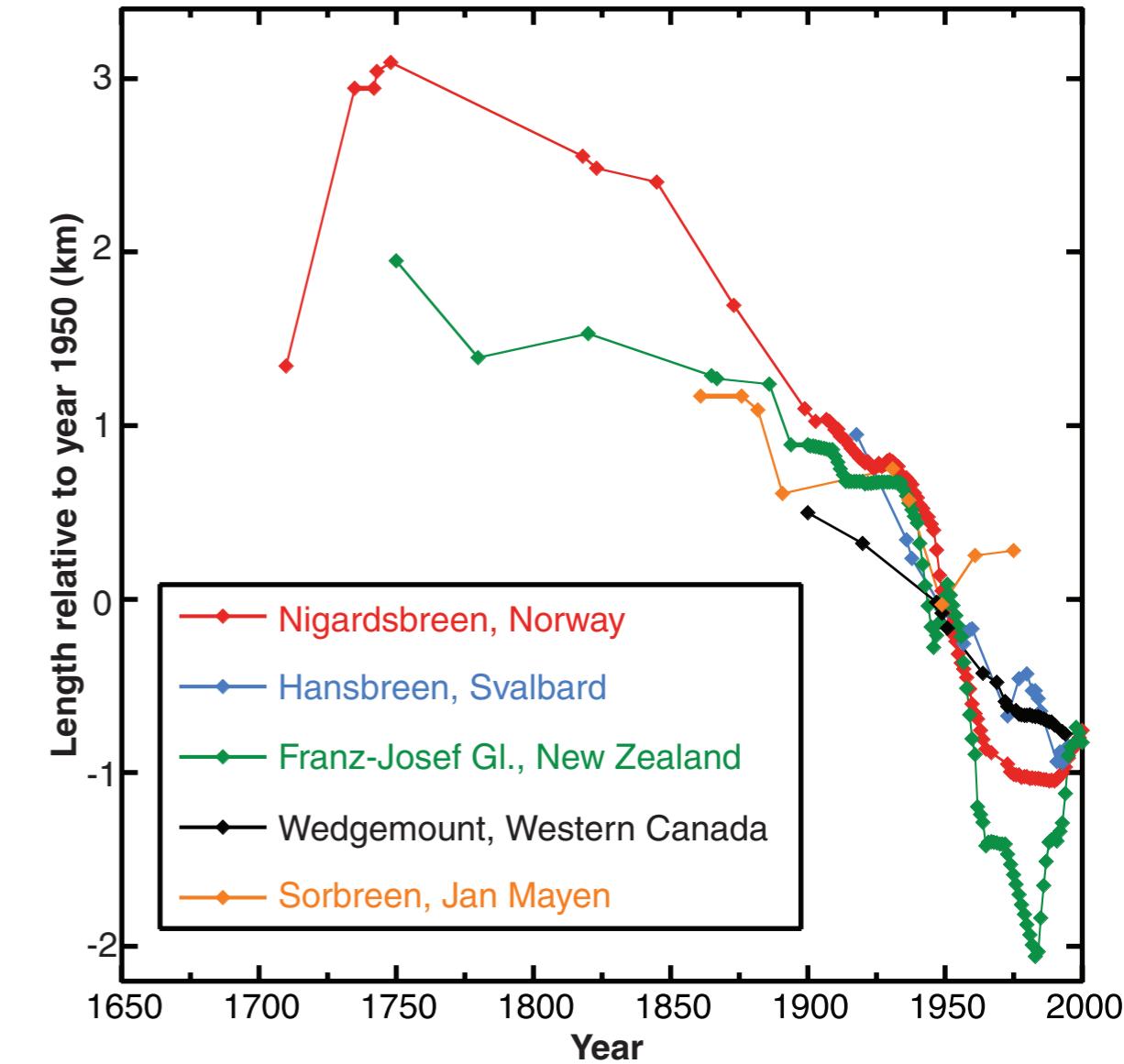
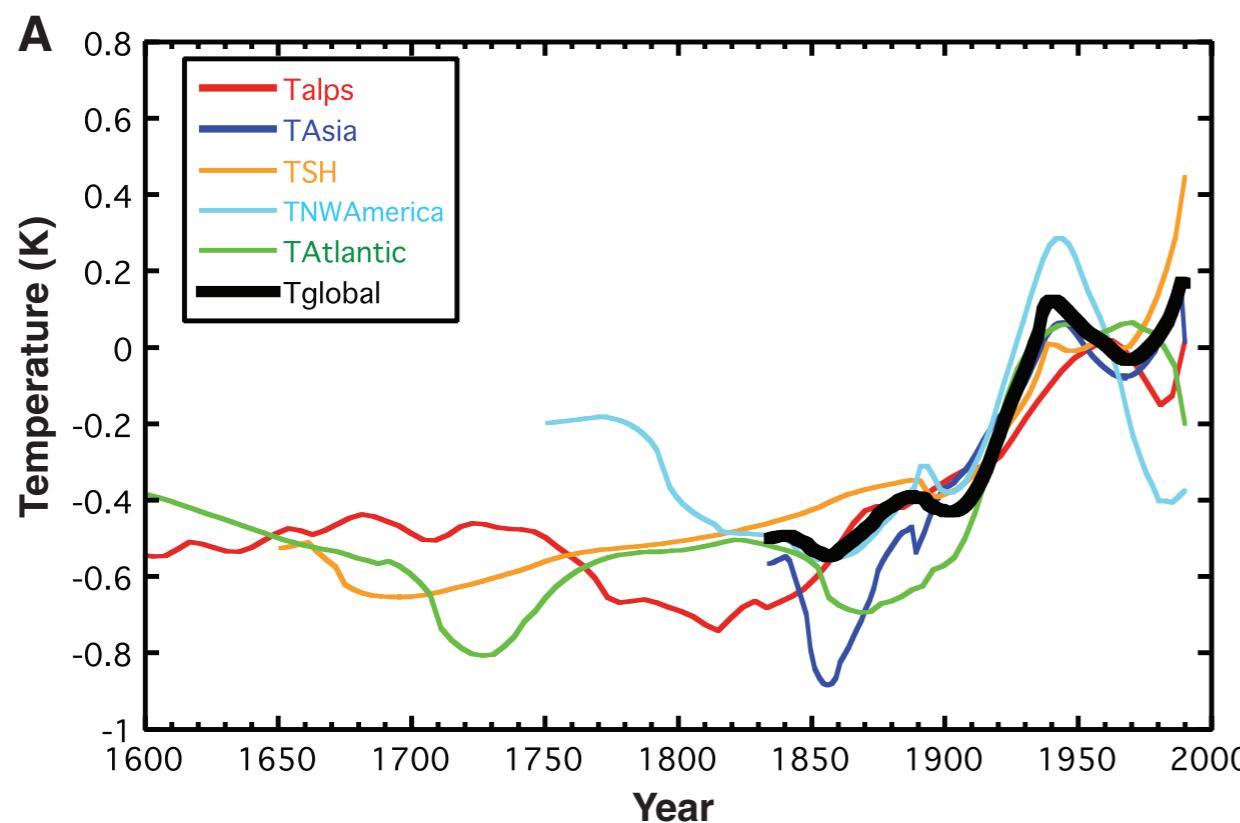
Scientists endured bitter winds to retrieve ancient ice from a blue ice field in the Allan Hills of Antarctica. YUZHENG YAN

Record-shattering 2.7-million-year-old ice core reveals start of the ice ages

How do mountain glaciers respond to climate forcing?

To the board!

Glacier state as a filter/recorder of climate (Oerlemans 2005)



Glacier retreat as “categorical” evidence of regional climate change (Roe et al. 2017)

