

Modeling of Geophysical Flows – Analysis of Models and Modeling Assumptions Using UQ

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Abstract

Dense large scale granular avalanches are a complex class of flows with physics that has often been poorly captured by models that are computationally tractable. Sparsity of actual flow data (usually only a posteriori deposit information is available) and large uncertainty in the mechanisms of initiation and flow propagation make the modeling task challenging and a subject of much continuing interest. Models that appear to represent the physics well in certain flows turn out to be poorly behaved in others due to intrinsic mathematical or numerical issues. Nevertheless, given the large implications on life and property many models with different modeling assumptions have been proposed.

While, inverse problems can shed some light on parameter choices it is difficult to make firm judgments on the validity or appropriateness of any single or set of modeling assumptions for a particular target flow or potential flows that needs to be modeled for predictive use in hazard analysis. We will present here an uncertainty quantification based approach to carefully, analyze the effect of modeling assumptions on quantities of interest in simulations based on three established models (Mohr-Coulomb, Pouliquen-Forterre and Voellmy-Salm) and thereby derive a model (from a set of modeling assumptions) suitable for use in a particular context. We also illustrate that a simpler though more restrictive approach is to use a Bayesian modeling average approach based on the limited data to combine the outcomes of different models.

We juxtapose observation data to the simulation results, only with the purpose to maintain a link to the real occurrence of a geophysical flow representing a possible outcome of the case studies described. The fundamental focus of this paper is the exploration of the dynamics of the simulated flows, enabling a notion of the contribution of different mechanisms or models elements inside the simulation procedure, in a fully quantitative, predictive-use oriented and statistical framework.

1 Introduction

1.1 Geophysical mass flows

Geophysical mass flows include debris and mud flows, landslides, snow and rock avalanches, glacier flows, pyroclastic surges, block and ash flows, and pumice flows, lahars, jökulhaups and many other examples. These flows are sometimes tens of kilometers in length and may travel at speeds as fast as hundreds of meters per second. Their deposits can be as much as tens of meters deep and kilometers long. In other words, there is no single universal description of a “typical” geophysical mass flow. Different types of flows are the results of different types of mechanisms/processes and hence different types of flows will have significantly different physical characteristics.

The rheology complexity of the fluidized material, and the mathematical problem of modeling and computing, make the description of the dynamics of those flows really challenging. In many cases the flows are *shallow*, i.e. the horizontal dimension is significantly larger than the flow depth. This assumption allows to perform depth-integration on the governing 3D Navier-Stokes equations, resulting in the 2D formulation of Shallow Water Equations (SWE), also known as Saint Venant Equations (Batchelor, 2000; Luca et al., 2016). Pioneering work of Savage and Hutter (1989) that was followed by Hutter et al. (1993); Dade and Huppert (1998) developed a depth-averaged model for flow of granular materials. Their model was used to model a flow of granular material down an inclined plane, originally using the Mohr-Coulomb rheology (MC) (see also Jaeger et al. (1989); Fraccarollo and Toro (1995)). Afterwards, large number of studies focused on the modeling of granular flows, including geophysical mass flows, using SWE approach and many of them were carefully reviewed in Pudasaini and Hutter (2007).

Modeling flow of granular material down an inclined plane was explored in detail by several further studies, both theoretically and experimentally (Pouliquen, 1999; Ruyer-Quil and Manneville, 2000; Silbert et al., 2001; Bursik et al., 2005; DaCruz et al., 2005). Granular material slumping (Balmforth and Kerswell, 2005; Lajeunesse et al., 2005), rapid flow down smooth inclines (Greve et al., 1994; Wieland et al., 1999) and shock waves (Gray et al., 2003; Hokanardottir and Hogg, 2005) were modeled and tested. Terrain erosion effects were investigated (Pitman et al., 2003a; Edwards and Gray, 2015), and also material deposition and self-channeling effects (Mangeney-Castelnau et al., 2005, 2007).

In particular, the experiments on rough inclined planes led to the development of the Pouliquen-Forsterre rheology (PF), assuming a variable frictional behavior as a function of flow regime (i.e. Froude Number, Fr) and flow depth (Pouliquen, 1999; Forsterre and Pouliquen, 2002; Pouliquen and Forsterre, 2002; Forsterre and Pouliquen, 2003). They allowed for an improved modeling of front propagation, mass spreading, surface wave-propagation and vortices-evolution in the flows on a rough terrain (Forsterre, 2006; Jop et al., 2006; Forsterre and Pouliquen, 2008).

In parallel to and sometimes anticipating the earth avalanches experiments and models, study of snow avalanches led to the development of the Voellmy-Salm rheology (VS) (Voellmy, 1955; Salm et al., 1990; Salm, 1993; Bartelt et al., 1999). Dense snow or debris avalanches consist of mobilized, rapidly flowing ice-snow mixed to debris-rock granules (Bartelt and McArdell, 2009). The VS rheology assumes a velocity dependent resisting term in addition to the traditional basal friction, ideally capable of including an approximation of the turbulence-generated friction. Many experimental and theoretical studies were developed in this framework (Gruber and Bartelt, 2007; Kern et al., 2009; Christen et al., 2010; Fischer et al., 2012).

In Iverson (1997); Iverson and Denlinger (2001); Denlinger and Iverson (2001, 2004); Iverson et al. (2004), the depth-averaged model was further studied and applied in the simulation of geophysical flows over digital models of natural terrains. Moreover, Gray et al. (1999, 2003) tested modeling of fast avalanches, exploring the high-velocity effects of channelizing/chuting topographies. A compensation for the effect of earth pressure changes was implemented and explored in detail (Pirulli et al., 2007; Pirulli and Mangeney, 2008), as well as the implementation of a significant curvature in the terrain (Pudasaini and Hutter, 2003; Fischer et al., 2012). Several specific studies on the pore pressure effects on the flow initiation and fluidization were developed (Savage and Iverson, 2003; Iordanoff and Khonsari, 2004; Iverson and George, 2014). A two-phase model was also implemented to more accurately simulate heterogenous flows with a significant portion of interstitial fluid (Pitman and Le, 2005). All of these complex modeling choices correspond to alternative or additional physical assumptions, and they are sometimes very difficult to be evaluated, compared or reasonably combined together.

A particular interest is raised by the specific efforts devoted to the modeling of volcanic mass flows

with SWE (Freundt and Bursik, 1998; Bursik et al., 2005; Saucedo et al., 2005; Kelfoun and Druitt, 2005; Charbonnier and Gertisser, 2009; Kelfoun et al., 2009; Procter et al., 2010; Sulpizio et al., 2010; Kelfoun, 2011; Charbonnier et al., 2013). Volcanoes are great sources for a rich variety of geophysical flow types in addition to the collected field data from past flow events.

In this study, we use TITAN2D software that is a publicly available toolkit for simulation of granular flows over natural terrains by solving depth-averaged equations using Digital Elevation Model (DEM) of terrain (Pitman et al., 2003b; Patra et al., 2005, 2006; Yu et al., 2009; Aghakhani et al., 2016). In the 4th release of TITAN2D, the capability of having multiple rheology options (MC, PF and VS rheologies) for simulation of flow in the same computational platform have been implemented. Therefore, we can perfectly make use of this opportunity and only focus on modeling characteristics of these rheology options without concerns about effects of diverse numerical frameworks. So far, TITAN2D achieved many successful applications in the simulation of different geophysical mass flows with specific peculiarities (Sheridan et al., 2005; Rupp et al., 2006; Norini et al., 2009; Charbonnier and Gertisser, 2009; Procter et al., 2010; Sheridan et al.; Sulpizio et al., 2010; Capra et al., 2011). Several studies involving TITAN2D were recently directed towards a statistical study of geophysical flows, focusing on the uncertainty quantification and propagation (Dalbey et al., 2008; Dalbey, 2009; Stefanescu et al., 2012a,b), or on the more efficient production of hazard maps (Bayarri et al., 2009; Spiller et al., 2014; Bayarri et al., 2015; Ogburn et al., 2016).

This study for the first time applies a statistical approach to the detailed investigation of the material model *stress components* and their associated *powers*, as well as the *Froude Number Fr*, the *acceleration* and observable quantities as the flow *velocity*, *height*, *lateral extent*, and *area*. The stress components have a strong link with the dynamic equation terms, while *Fr* can give information on the flow regime; moreover, the observable quantities have a direct link to field data and hazard assessments. This can provide a complete description of the geophysical flow. Our main purpose is to obtain a quantitative and statistical understanding of the consequences of the different physical assumptions, both on the forces inside the flow and on the observable outputs in time and space.

In particular, what we present is a procedure for the improved exploration and quantitative comparison of physical models and their assumptions through the collection of full statistical data. Behind each physical model there are different physical assumptions, and therefore it is more appropriate to look at those instead than at the entire model results. Our statistical approach shall enable a first step towards a data driven selection of the best modeling assumptions to use.

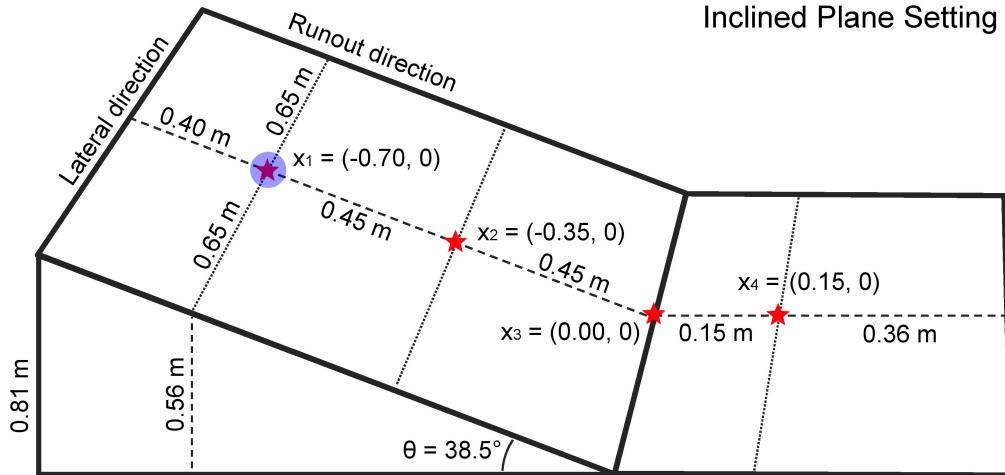


Figure 1: Inclined plane description, including local samples sites (red stars). Pile location is marked by a blue dot.

1.2 Overview of the case studies

The first case study assumes very simple boundary conditions, and corresponds to an experiment fully described in Webb (2004); Bursik et al. (2005); Webb and Bursik (2016). It is a classical flow down an inclined plane set-up, including a change in slope to an horizontal plane (Fig. 1). Four locations are selected among the center line of the flow to accomplish local testing. These are: the initial pile location $L_1 = (-0.7, 0)$ m, the middle of the inclined plane $L_2 = (-0.35, 0)$ m, the change in slope $L_3 = (0, 0)$ m, the middle of the flat plane $L_4 = (0.15, 0)$ m (see Section 4).

The second case study is a block and ash flow down the slope of Volcán de Colima (México) - an andesitic stratovolcano that rises to 3,860 m above sea level, situated in the western portion of the Trans-Mexican Volcanic Belt (Fig. 2). Historically, it has been the most active volcano in México (la Cruz-Reyna, 1993; Zobin et al., 2002; González et al., 2002). The modeling of pyroclastic flows generated by explosive eruptions and lava dome collapses of Volcán de Colima is a well studied problem (Martin Del Pozzo et al., 1995; Sheridan and Macías, 1995; Saucedo et al., 2002, 2004, 2005; Sarocchi et al., 2011; Capra et al., 2015). The volcano has been already used as a case study in several research involving the Titan2D code (Rupp, 2004; Rupp et al., 2006; Dalbey et al., 2008; Yu et al., 2009; Sulpizio et al., 2010; Capra et al., 2011; Aghakhani et al., 2016). During July 10th-11th, 2015, the volcano underwent its most intense eruptive phase since its Subplinian-Plinian 1913 AD eruption (Saucedo et al., 2010; Zobin et al., 2015; Reyes-Dávila et al., 2016; Capra et al., 2016).

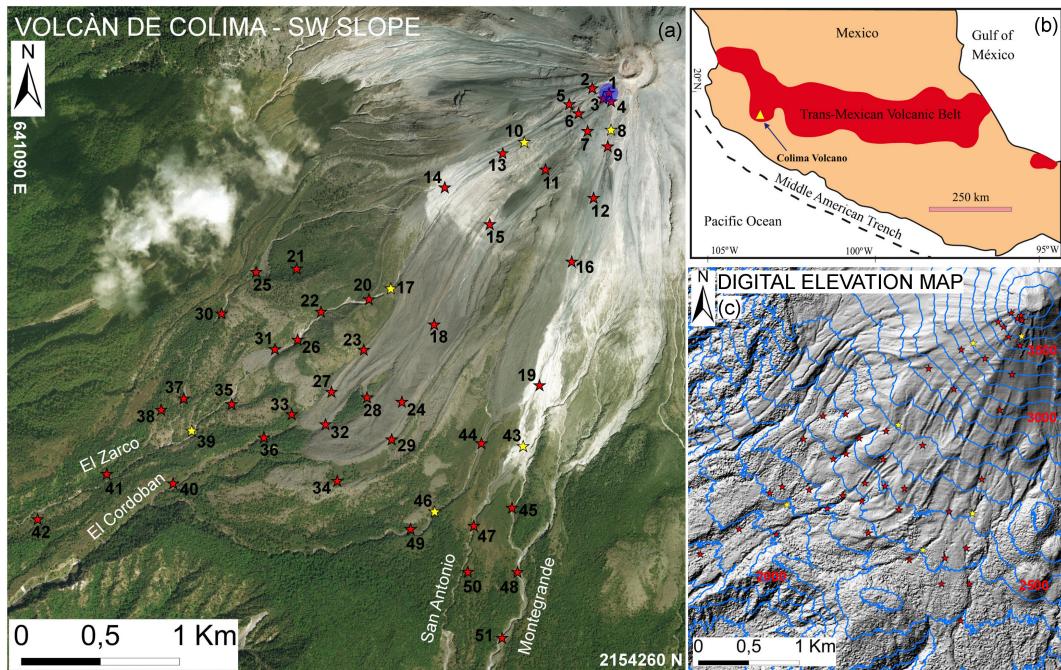


Figure 2: (a) Volcán de Colima (México) overview, including 51 numbered local sample sites (stars) and four labeled major ravines channeling the flow. Pile location is marked by a blue dot. Reported coordinates are in UTM zone 13N. Background is a satellite photo. (b) Regional geology map. (c) Digital elevation map. Six points that are adopted as preferred locations are highlighted in yellow. Elevation isolines are included in blue, elevation values in red.

We assume the flow to be generated by the gravitational collapse of a paraboloid dome placed close to the summit area. A dome collapse occurs when there is a significant amount of recently-extruded highly-viscous lava piled up in an unstable configuration around a vent. Further extrusion and/or external forces can cause the still hot dome of viscous lava to collapse, disintegrate, and avalanche downhill (Bursik et al., 2005; Wolpert et al., 2016). The hot, dense blocks in this “block and ash” flow (BAF) will typically range from centimeters to a few meters in size. The matrix is composed of fine ash from the

commuted blocks. Computations were performed on a DEM with 5m-pixel resolution, obtained from LiDAR data acquired in 2005 (Davila et al., 2007; Sulpizio et al., 2010). We select 51 locations along the flow inundated area to accomplish local testing. Six of them are then adopted as preferred locations, being representative of different flow regimes (see Section 5).

2 Overview of the models

2.1 Modeling Assumptions

The models of geophysical mass flows described in this study rely on several physical and mathematical assumptions. We will classify them in two groups - the general assumptions and the rheology assumptions. The main focus of this study is on the latter group, but, in principle, the same procedure could be applied to the former.

General Assumptions

- The *shallowness* approximation is at the base of the depth-averaging procedure. In this approximation, the flow depth is assumed to be at least an order of magnitude less than the characteristic length of the flowing material. Variation of velocity within the flow depth is neglected.
- The material is assumed to be *continuous*. The real flows consist of granular material and, often, interstitial fluid. Usually the “typical particle” diameter is small compared to the depth and length of the flowing mass, and the rheological properties are imposed to approximate the bulk behavior that is expected of an actual mass flow.
- The moving mass is assumed to be *volume preserving*. In contrast, phenomena of erosion and deposition of material may violate this assumption.
- The flowing mass is assumed to be consisted of an *incompressible* fluid. This means that we are ignoring variations of density due to the void ratio (i.e. the density is uniform in space and time).
- The body of mass is supposed to be in *isothermal* state or, if not, thermal effects can be ignored.
- The mass flow is assumed to be *free surface*. Air entrainment is instead frequently observed in the dilute real flows.
- The stress tensor is assumed *symmetric* with respect to the z axis.

The shallowness assumption is very important, and brings several implications which may be considered additional assumptions on their own.

Consequences of Shallowness

- The shallowness assumption implies a *hydrostatic* expression for the normal pressure in the direction perpendicular to the basal surface. Moreover, the downslope and cross-slope pressure components are assumed to be varying linearly with the normal pressure component through the flow depth.
- The *boundary layer* where the shearing deformation takes place is collapsed to zero thickness. The sliding and shearing velocity are combined to a single sliding law with a somewhat larger modelled sliding velocity.
- The *lateral shear stresses can be neglected*, compared to the basal shear stresses. Motion of the bulk of mass consists of “shearing within the deforming mass” and “sliding along the basal surface”.

The list rheology assumptions is detailed when the models are presented.

2.2 Governing Equations and Boundary Conditions

The motion of the mass flow is described by the basic conservation of mass and momentum for an incompressible medium:

$$\begin{aligned} \nabla \cdot \underline{\mathbf{u}} &= 0, \\ \frac{\partial}{\partial t} (\rho \underline{\mathbf{u}}) + \nabla \cdot (\rho \underline{\mathbf{u}} \otimes \underline{\mathbf{u}}) &= \nabla \cdot \underline{\underline{\mathbf{T}}} + \rho \underline{\mathbf{g}}, \end{aligned} \quad (1)$$

Where $\underline{\mathbf{u}} = [u, v, w]^T$ is the material velocity in cartesian coordinates, ρ is its constant density, $\underline{\mathbf{T}}$ is the *Cauchy* stress tensor, and $\underline{\mathbf{g}}$ is the gravity acceleration vector.

The Cauchy stress tensor, $\underline{\mathbf{T}}$, depends on the rheology assumptions. The kinetic boundary conditions are defined on the free surface $F^s = s(x, y, t) - z = 0$ and basal surface $F^b = b(x, y, t) - z = 0$ interfaces:

$$\frac{\partial F^s}{\partial t} + \underline{\mathbf{u}} \cdot \nabla F^s = 0, \quad \text{at } F^s(x, y, t) = 0 \quad (2)$$

$$\frac{\partial F^b}{\partial t} + \underline{\mathbf{u}} \cdot \nabla F^b = 0, \quad \text{at } F^b(x, y, t) = 0 \quad (3)$$

The depth-averaged Saint-Venant equations that TITAN2D solves are:

$$\begin{aligned} \frac{\partial h}{\partial t} + \frac{\partial}{\partial x}(h\bar{u}) + \frac{\partial}{\partial y}(h\bar{v}) &= 0, \\ \frac{\partial}{\partial t}(h\bar{u}) + \frac{\partial}{\partial x}\left(h\bar{u}^2 + \frac{1}{2}kg_zh^2\right) + \frac{\partial}{\partial y}(h\bar{u}\bar{v}) &= S_x, \\ \frac{\partial}{\partial t}(h\bar{v}) + \frac{\partial}{\partial x}(h\bar{u}\bar{v}) + \frac{\partial}{\partial y}\left(h\bar{v}^2 + \frac{1}{2}kg_zh^2\right) &= S_y \end{aligned} \quad (4)$$

Here the cartesian coordinate system is aligned such that z is normal to the surface; h is the flow height in the z direction; $h\bar{u}$ and $h\bar{v}$ are respectively the components of momentum in the x and y directions; and k is the coefficient which relates the lateral stress components, $\bar{\sigma}_{xx}$ and $\bar{\sigma}_{yy}$, to the normal stress component, $\bar{\sigma}_{zz}$. The definition of this coefficient depends on the constitutive model of the flowing material we choose. Note that $\frac{1}{2}kg_zh^2$ is the contribution of hydrostatic pressure to the momentum fluxes. S_x and S_y are the sum local stresses: they include the gravitational driving forces, the basal friction force resisting to the motion of the material, and additional forces specific of rheology assumptions.

2.2.1 Numerical solver

Let $\mathbf{U} = [h, h\bar{u}, h\bar{v}]^T$ be the vector of conservative variables and $\mathbf{F} = [h\bar{u}, h\bar{u}^2 + 0.5k_{ap}g_zh^2, h\bar{v}\bar{u}]^T$ and $\mathbf{G} = [h\bar{v}, h\bar{u}\bar{v}, h\bar{v}^2 + 0.5k_{ap}g_zh^2]^T$ be the physical flux vectors. In order to solve the system of hyperbolic conservation laws (4), TITAN2D employs a finite volume Godunov solver, and the time-stepping procedure is achieved by an explicit Euler scheme (Pitman et al., 2003b; Patra et al., 2005, 2006). Assuming $\mathbf{S} = [S_h, S_x, S_y]^T$ as the source terms vector containing the effect of the flow rheology, the flow in the next time step updates as:

$$\mathbf{U}_{i,j}^{n+1} = \mathbf{U}_{i,j}^n - \frac{\Delta t}{\Delta x}\{\mathbf{F}_{i+\frac{1}{2},j}^n - \mathbf{F}_{i-\frac{1}{2},j}^n\} - \frac{\Delta t}{\Delta y}\{\mathbf{G}_{i,j+\frac{1}{2}}^n - \mathbf{G}_{i,j-\frac{1}{2}}^n\} + \Delta t \mathbf{S}_{i,j} \quad (5)$$

Where $\mathbf{F}_{i\pm\frac{1}{2},j}^n$ and $\mathbf{G}_{i,j\pm\frac{1}{2}}^n$ are the numerical flux terms at the inter-cell boundaries which are computed regarding the Harten-Lax-Van Leer Riemann solver (Toro, 2013). In fact, the evolution of the flow to the next time step depends on the advection flux at the cell interface, which results from the wave interaction at the boundaries between cells. On the other hand, adaptive mesh refinement allows for the concentration of computing power on regions of special interest. It captures the complex flow features along the flow boundaries, as well as the locations where there are large mass or momentum fluxes (which may include places where the topography changes abruptly). Mesh coarsening is also applied where the solution values are relatively constant or small to further improve the computational efficiency (Patra et al., 2005; Aghakhani et al., 2016).

2.3 Rheology assumptions

In the three following sections, we briefly describe *Mohr-Coulomb* (MC), *Pouliquen-Forsterre* (PF) and *Voellmy-Salm* (VS) models.

2.3.1 Mohr-Coulomb model

Based on the long history of studies in soil mechanics (Rankine, 1857), the Mohr-Coulomb rheology model was developed and used to represent the behavior of geophysical mass flows Savage and Hutter (1989).

Shear and normal stress are assumed to obey Coulomb friction equation, both within the flow and at its boundaries. In other words,

$$\tau = \sigma \tan \phi, \quad (6)$$

where τ and σ are respectively the shear and normal stresses on failure surfaces, and ϕ is a friction angle. This relationship does not depend on the flow speed.

Under the assumption of symmetry of the stress tensor w.r.t. the z axis, the earth pressure coefficient $k = k_{ap}$ can take on only one of three values indicated by the square dots shown in Figure 3. The material yield criterion is represented by the two straight lines at angles $\pm\phi$ (the internal friction angle) relative to horizontal direction. Similarly, the normal and shear stress at the bed are represented by the line $\tau = -\sigma \tan(\delta)$ where δ is the bed friction angle. The three possible stress states in the xz -plane are indicated by the square dots. The circular dots show the additional possible stress states for σ_{yy} for a non-axisymmetric stress tensor.

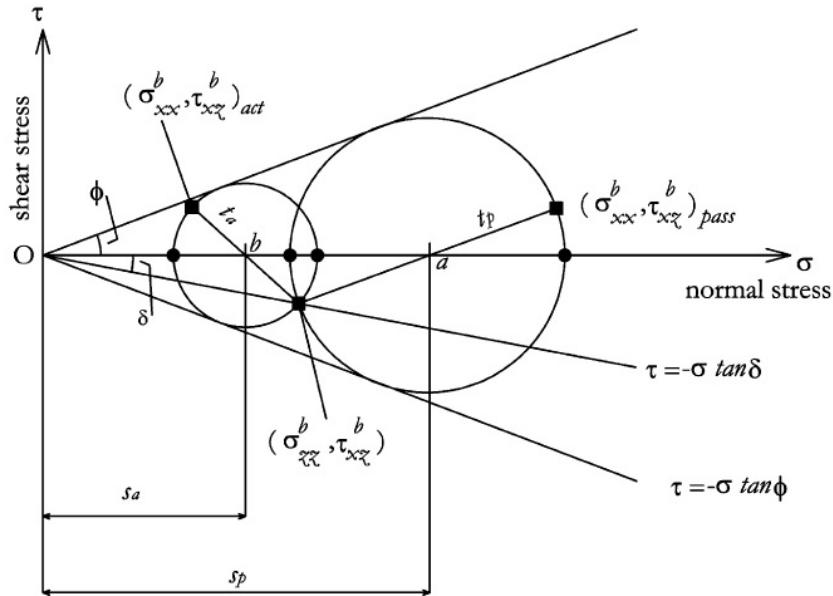


Figure 3: Mohr-circle-diagram representing the stress state within a granular medium (Pirulli et al., 2007).

Regarding the friction angle stated above, we have two material properties in this model:

- Internal friction angle, ϕ_{int} , which resists material shear.
- Bed friction angle, ϕ_{bed} , which resists motion of the material relative to the bed. This is a joint property of the material and the surface it flows over.

It is worth mentioning that the effective value of the friction angles can be strongly reduced by the presence of interstitial fluid, sometimes creating modeling issues. Specific models have been developed in case the effect of interstitial fluid is believed to be particularly relevant (Pitman and Le, 2005).

Earth pressure coefficient $k = k_{ap}$ is defined in MC as:

$$k_{ap} = \begin{cases} \frac{\sigma_{xx}^b \text{act}}{\sigma_{zz}^b} = 2 \frac{1 - \sqrt{1 - \cos^2(\phi_{int})(1 + \tan^2(\phi_{bed}))}}{\cos^2(\phi_{int})} - 1, & \nabla \cdot \tilde{\mathbf{u}} > 0, \text{ active} \\ \frac{\sigma_{xx}^b}{\sigma_{zz}^b} = 1, & \nabla \cdot \tilde{\mathbf{u}} = 0, \text{ neutral} \\ \frac{\sigma_{xx}^b \text{pass}}{\sigma_{zz}^b} = 2 \frac{1 + \sqrt{1 - \cos^2(\phi_{int})(1 + \tan^2(\phi_{bed}))}}{\cos^2(\phi_{int})} - 1, & \nabla \cdot \tilde{\mathbf{u}} < 0, \text{ passive} \end{cases} \quad (7)$$

MC equations As a result, we can write down the source terms of the Eqs. (4):

$$\begin{aligned} S_x &= g_x h - \frac{\bar{u}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\bar{u}^2}{r_x} \right) \tan(\phi_{bed}) \right] - h k_{ap} \operatorname{sgn} \left(\frac{\partial \bar{u}}{\partial y} \right) \frac{\partial(g_z h)}{\partial y} \sin(\phi_{int}), \\ S_y &= g_y h - \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\bar{v}^2}{r_y} \right) \tan(\phi_{bed}) \right] - h k_{ap} \operatorname{sgn} \left(\frac{\partial \bar{v}}{\partial x} \right) \frac{\partial(g_z h)}{\partial x} \sin(\phi_{int}) \end{aligned} \quad (8)$$

Where, $\tilde{\mathbf{u}} = (\bar{u}, \bar{v})$, is the depth-averaged velocity vector, r_x and r_y denote the radii of curvature of the local basal surface. The inverse of the radii of curvature is usually approximated with the partial derivatives of the basal slope, e.g., $1/r_x = \partial \theta_x / \partial x$, where θ_x is the local bed slope.

MC rheology assumptions In summary:

- *Basal Friction* is based on a constant friction angle.
- *Internal Friction* gives a not negligible contribution and it is based on a constant friction angle.
- *Earth pressure coefficient* formula depends on the Mohr-Coulomb circle.
- Velocity based curvature effects are included into the equations.

2.3.2 Pouliquen-Forterre model

The scaling properties for granular flows down rough inclined planes led to a new formulation of the basal friction stress as a function of the flow depth and velocity (Pouliquen, 1999).

Two critical slope inclination angles are defined as functions of the flow thickness, namely $\phi_{start}(h)$ and $\phi_{stop}(h)$. The function $\phi_{stop}(h)$ gives the slope angle at which a steady uniform flow leaves a deposit of thickness h , while $\phi_{start}(h)$ is the angle at which a layer of thickness h is mobilized. They define two different basal friction coefficients:

$$\mu_{start}(h) = \tan(\phi_{start}(h)) \quad (9)$$

$$\mu_{stop}(h) = \tan(\phi_{stop}(h)) \quad (10)$$

An empirical friction law $\mu_b(\|\tilde{\mathbf{u}}\|, h)$ is then defined in the whole range of velocity and thickness. The expression changes depending on two flow regimes, according to a parameter β and the Froude number $Fr = \|\tilde{\mathbf{u}}\| / \sqrt{hg_z}$.

Dynamic friction regime - $Fr \geq \beta$

$$\mu(h, Fr) = \mu_{stop}(h\beta/Fr) \quad (11)$$

Intermediate friction regime - $0 \leq Fr < \beta$

$$\mu(h, Fr) = \left(\frac{Fr}{\beta} \right)^\gamma [\mu_{stop}(h) - \mu_{start}(h)] + \mu_{start}(h), \quad (12)$$

where γ is the power of extrapolation, assumed equal to 10^{-3} in the sequel (Pouliquen and Forterre, 2002). In particular, if $Fr = \beta$, then $\mu(h, Fr) = \mu_{stop}(h)$, and if $Fr = 0$, then $\mu(h, Fr) = \mu_{start}(h)$.

The functions μ_{stop} and μ_{start} are defined by:

$$\mu_{stop}(h) = \tan \phi_1 + \frac{\tan \phi_2 - \tan \phi_1}{1 + h/\mathcal{L}} \quad (13)$$

and

$$\mu_{start}(h) = \tan \phi_3 + \frac{\tan \phi_2 - \tan \phi_1}{1 + h/\mathcal{L}} \quad (14)$$

The critical angles ϕ_1 , ϕ_2 and ϕ_3 and the parameters \mathcal{L}, β are the parameters of the model.

In particular, \mathcal{L} is the characteristic depth of the flow over which a transition between the angles ϕ_1 to ϕ_2 occurs, in the μ_{stop} formula. In practice, if $h \ll \mathcal{L}$, then $\mu_{stop}(h) \approx \tan \phi_2$, and if $h \gg \mathcal{L}$, then $\mu_{stop}(h) \approx \tan \phi_1$. The effect of the topographic local curvatures is also taken into account.

PF equations The depth-averaged Eqs. (4) source terms thus take the following form:

$$\begin{aligned} S_x &= g_x h - \frac{\bar{u}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\bar{u}^2}{r_x} \right) \mu_b(\|\tilde{\mathbf{u}}\|, h) \right] + g_z h \frac{\partial h}{\partial x} \\ S_y &= g_y h - \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\bar{v}^2}{r_y} \right) \mu_b(\|\tilde{\mathbf{u}}\|, h) \right] + g_z h \frac{\partial h}{\partial y} \end{aligned} \quad (15)$$

PF rheology assumptions In summary:

- *Basal Friction* is based on an interpolation of different friction angles, based on the flow regime and depth.
- *Internal Friction* is neglected.
- *Earth pressure coefficient* is equal to one.
- Normal stress is modified by a *hydrostatic pressure force* related to the flow height gradient.
- Velocity based *curvature effects* are included into the equations.

2.3.3 Voellmy-Salm model

The theoretical analysis of dense snow avalanches led to the VS rheology model (Voellmy, 1955; Salm, 1993). The following relation between shear and normal stresses holds:

$$\tau = \mu \sigma + \frac{\rho \|\mathbf{g}\|}{\xi} \|\tilde{\mathbf{u}}\|^2, \quad (16)$$

where, σ denotes the normal stress at the bottom of the fluid layer and $\mathbf{g} = (g_x, g_y, g_z)$ represents the gravity vector. The VS rheology adds a velocity dependent *turbulent* friction to the traditional velocity independent basal friction term which is proportional to the normal stress at the flow bottom. The two parameters of the model are the bed friction coefficient μ and the turbulent friction coefficient ξ . By dimension analysis, ξ is equivalent to an acceleration, while μ is dimensionless. The decomposition of the total basal friction into velocity independent and dependent parts allows the modeling of either the dynamic regime when the avalanche is flowing with a high velocity in the acceleration zone, and when it is close to stopping in the runout zone.

The effect of the topographic local curvatures is again taken into account by adding the terms containing the local radii of curvature r_x and r_y . In this case the formula is considering the modulus of velocity instead than the scalar component (Fischer et al., 2012).

VS equations Therefore, the final source terms take the following form:

$$\begin{aligned} S_x &= g_x h - \frac{\bar{u}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\|\tilde{\mathbf{u}}\|^2}{r_x} \right) \mu + \frac{\|\mathbf{g}\|}{\xi} \|\tilde{\mathbf{u}}\|^2 \right], \\ S_y &= g_y h - \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \left[h \left(g_z + \frac{\|\tilde{\mathbf{u}}\|^2}{r_y} \right) \mu + \frac{\|\mathbf{g}\|}{\xi} \|\tilde{\mathbf{u}}\|^2 \right]. \end{aligned} \quad (17)$$

VS rheology assumptions In summary:

- *Basal Friction* is based on a constant coefficient, similarly to the MC rheology.
- *Internal Friction* is neglected.
- *Earth pressure coefficient* is equal to one.
- Additional *turbulent friction* is based on the local velocity by a quadratic expression.
- Velocity based *curvature effects* are included into the equations, following an alternative formulation.

3 UQ Process

The key to a good forecasting capability in the context of mass flows requires the careful selection of the pair $(M(A), P_{M(A)})$, where A is a set of assumptions, $M(A)$ is the model which combines those assumptions, and P_M is a probability distribution in the parameter space of M . For the sake of simplicity we are always taking P_M uniformly distributed on selected parameter ranges.

Models and assumptions An assumption is a quite general concept - for example it can be a specific equation for the internal stress, the implementation of bed curvature effects, of active-passive material stretching, or the use of a specific correction on the pressure effects. Assumptions are what makes the models being different, and each model may be seen as the combined result of a set of assumptions. Sometimes a good model contains a useless assumption that may be removed, sometimes a good assumption should be implemented inside a different model - those are usually considered as subjective choices, not data driven. Moreover, the correct assumptions may change through time, making the analysis more difficult. We provided a list of assumptions behind SWE general implementation, and behind each rheology in section 2.

Parameters ranges It is worth mentioning that whereas the support of P_M can be restricted to a single point in case an inverse problem is solved for the optimal reconstruction of a particular flow, this is not possible if we are interested in the general predictive capabilities of the model, i.e. the target of probabilistic hazard assessment. In this study we always assume $P_M \sim \bigotimes_{i=1}^{N_M} Unif(a_{i,M}, b_{i,M})$, where N_M is the number of parameters of M . These parameter ranges will not be selected under the influence of a particular observation, but we will try to use the information gathered in literature about the physical meaning of those values, together with a preliminary testing, coupling the range of possible runouts. This coupling step is detailed in section 3.3, and required to focus the statistical comparison on a consistence range of flow regimes.

Simulated quantities The simulation algorithms can be schematized as:

(1) INPUT VARIABLES → (2) HIDDEN DYNAMIC QUANTITIES → (3) OBSERVABLE OUTPUTS

The *input variables* are the parameters of M , i.e. volume, rheology coefficients, but can include also the initiation site and geometry, and the digital elevation map. The *hidden quantities* include the stress terms in the Newton Equations that rule the simulation, and their powers. Those are hidden to the observation in a real flow, but they directly depend on the parameters, and represent a fundamental link between the parameters and the observable outputs. Moreover, the models share some of those terms while change others, and this enables a detailed comparison of the real physics below the curtain (see Sections 4.2, 5.3). Finally, the *observable outputs* include what can be measured in space and time: e.g. flow height, lateral extent, area, velocity, acceleration, and combined quantities as Fr (see Sections 4.1, 5.2). In the sequel (2) and (3) are also called quantities of interest (QoI).

Local samples and spatial averaging The QoI in the previous scheme which are evolving fields $f(\underline{x}, t)$ in space-time, can be either *locally sampled* or *spatially averaged* (see sections 4, 5). The spatial averaging is defined by $F(t) = \int_{\mathbb{R}^k} f(\underline{x}, t) d\underline{x}$. In the most of the cases $k = 2$, and $d\underline{x}$ is given by the area of the mesh elements. Sometimes $k = 3$, e.g. concerning speed, and $d\underline{x}$ is the element of volume corresponding to the mesh elements.

A fundamental part of this research was made possible by the new implementation of a local sampling option in the TITAN2D cyber-infrastructure. In particular, once a set of sample points $(x_i)_{i=1,\dots,N}$ is chosen, each field $f(\underline{x}, t)$ is calculated as a function of time on the elements of the numerical mesh which are found to contain the $(x_i)_{i=1,\dots,N}$. Exploring differences in time and location enable to constrain the changes in flow regime. Those are really important because if the flow behavior is radically different, then the significance and effects of the assumptions may also change.

Monte Carlo simulation In general, for each QoI, during a Monte Carlo simulation we sample the input variables and obtain a family of temporal graphs on time domains depending on the flow duration. These results are statistically summarized - plotting their expectation, and their 5th and 95th percentiles. In the following, we will detail the considered input variables and quantities of interest for each of our cases study.

Our sampling technique of the input variables is based on the Latin Hypercube Sampling (LHS) idea, and in particular, on the improved space-filling properties of the orthogonal array-based Latin Hypercubes (see Appendix 6). The LHS is performed over $[0, 1]^3$ for the MC and VS input parameters, and $[0, 1]^4$ for PF input parameters (see Section 3.3). Those adimensional samples are homothetically transformed to fill the required intervals.

3.1 Statistical analysis of force contributions

A statistical analysis capable of measuring the contributions of different force components in the rheology models is presented. That is helpful to explore and compare the effects of the assumptions behind some modeling choices, onto the time evolution of the flowing material. Let $(F_i(x, t))_{i=1,\dots,6}$ be the array of force components, where $x \in \mathbb{R}^2$ is a spatial location, and $t \in T$ is a time instant.

In our case study, $(F_1, F_2) := (\mathbf{LHS}_1, \mathbf{LHS}_2)$, i.e. the left-hand side terms in the momentum equation, while $(F_3, \dots, F_6) := (\mathbf{RHS}_1, \dots, \mathbf{RHS}_4)$ are the right-hand side terms. These forces can take different formulations in the different models, which are detailed in section 3.2. This type of procedure can be applied to any additive decomposition of the physical forces. The analysis of *RHS* can inform on local effects, while *LHS* concerns the of un-local dynamics in space and time.

In general we show the projection of the forces on the slope direction, or the product with the velocity field, i.e. the power related to them. Hence, in the following the forces are scalar and not vectorial terms. It is important to remark that all the forces are hidden dynamic quantities (see above), depending on the input variables, and they are thus considered as random variables. The degree of contribution of those force terms is significantly variable in space and time, and we define the *contribution coefficients* to model this. The definitions are not depending on the location, but all the results will significantly depend of that choice. Next notation will assume to be in a selected location $x = L_k$, where $k \geq 1$. In the following we will focus on **RHS** forces, although some testing including **LHS** forces has also been accomplished.

Definition 1 (Contribution coefficients) Let $(F_i)_{i \in I}$ be random variables on (Ω, \mathcal{F}, P) , representing the considered force components in location x at time t . Then, for each component i , the contribution coefficient is defined as:

$$C_i := \mathbb{E} \left[\frac{F_i}{\Phi} \right],$$

where Φ is a dominating function, i.e. $\Phi \geq |F_i|, \forall i \in I$.

The *total force*, i.e $\sum_i F_i$ excluding the inertial terms, is not a good candidate for a dominating function. Indeed, the terms often have opposite signs, and their sum can be really small. Another issue is given by the existence of a subset of times Θ characterized by the absence of flow in the selected location x . In Θ the dominant force is null, and cannot be the denominator of a fraction.

We have tested two alternative Φ choices: Φ_1 is based on the l^1 norm, and defined as:

$$\Phi_1 := \begin{cases} \sum_i \frac{|F_i|}{2}, & \text{if not null;} \\ 1, & \text{otherwise.} \end{cases}$$

In contrast Φ_2 is a dominant force based on the l^∞ norm:

$$\Phi_2 := \begin{cases} \max_i |F_i|, & \text{if not null;} \\ 1, & \text{otherwise.} \end{cases}$$

In particular, for a particular location x , time t , and parameter sample ω , we have $C_i = 0$ if there is no flow or all the forces are null. The expectation of C_i is reduced by the chance of F_i being small compared to the other terms, or by the chance of having no flow in (x, t) . Moreover, $\mathbb{E}[C_i] \in [-1, 1], \forall i$.

Proposition 2 *Let Φ_k , $k = 1, 2$ be the functions described above. Then they are well defined dominating function, i.e. $\Phi_k \geq |F_j|, \forall j, k$.*

Proof. If $k = 1$, the Newton equation states $\sum_{j \in J} F_j = \sum_{i \in I} F_i$, where $(F_j)_{j \in J}$ are inertial forces, and $(F_i)_{i \in I}$ the local forces. Then, $\forall h \in I$

$$F_h = \sum_{i \neq h, i \in I} F_i - \sum_{j \in J} F_j.$$

And hence

$$|F_h| \leq |F_h| + \sum_{i \neq k, i \in I} |F_i| - \sum_{j \in J} |F_j| = 2\Phi_1.$$

A similar equation is valid $\forall k \in J$. If $k = 2$, or if the inertial forces are not included, the proof is trivial. \square

Furthermore, assuming $\Phi = \Phi_2$, there is a useful result explaining the meaning of those coefficients through the conditional expectation.

Proposition 3 *Let $(F_i)_{i \in I}$ be random variables on (Ω, \mathcal{F}, P) , representing the considered force components in location x at time t . For each i , let C_i be the contribution coefficient of force F_i , assuming $\Phi = \Phi_2$. Then we have the following expression:*

$$C_i = \sum_j p_j \mathbb{E} \left[\frac{F_i}{|F_j|} \mid \Phi = |F_j| \right],$$

where $p_j := P\{\Phi = |F_j|\}$.

Proof. Let Z be a discrete random variable such that, for each $j \in \mathbb{N}$, $(Z = j) \iff (\Phi = |F_j|)$. Then, by the rule of chain expectation:

$$\begin{aligned} C_i &= \mathbb{E} \left[\frac{F_i}{\Phi} \right] = \mathbb{E} \left[\mathbb{E} \left[\frac{F_i}{\Phi} \mid Z = j \right] \right] = \\ &= \mathbb{E} \left[\mathbb{E} \left[\frac{F_i}{|F_j|} \mid Z = j \right] \right] = \sum_j P\{Z = j\} \mathbb{E} \left[\frac{F_i}{|F_j|} \mid Z = j \right]. \end{aligned}$$

Moreover, by definition, $p_j = P\{Z = j\}$. This completes the proof. \square

The last proposition brings to the definition of the *dominance factors* $(p_j)_{j=1, \dots, k}$, i.e. the probability of each F_j to be the dominant force in (x, t) .

Definition 4 (Dominance factors and conditional contributions) *Let $(F_i)_{i=1, \dots, k}$ be random variables on (Ω, \mathcal{F}, P) , representing arbitrary force components in location x at time t . Then, for each pair of components (i, j) , the conditional contribution $C_{i,j}$ is defined as:*

$$C_{i,j} := \mathbb{E} \left[\frac{F_i}{|F_j|} \mid \Phi = |F_j| \right],$$

where $\Phi = \Phi_2$ is the dominant force. In particular, for each component i , the dominance factor is defined as:

$$p_j := P\{\Phi = |F_j|\}.$$

In sections 4.2.4 and 5.3.3 we display the contribution coefficients as a function of time, assuming $\Phi = \Phi_1$. Moreover, we report the dominance factors plots, another very useful tool to read the dynamic significance of the forces as a function of time and location.

3.2 Forces terms definition

Forces analyzed have been classified according to the terms of the momentum equation in (4). Moreover, the forces terms in S_x have been additionally classified, accordingly to equations (8), (15), (17), describing the three rheology models considered in this study. What follows is a summary of the force terms, and this notation will be used in the following (see Sections 4.2, 5.3).

Left-hand side, LHS forces

$$\mathbf{LHS}_1 = \left[\frac{\partial}{\partial t}(h\bar{u}), \frac{\partial}{\partial t}(h\bar{v}) \right], \quad (18)$$

it is the temporal derivative of momentum.

$$\mathbf{LHS}_2 = \left[\frac{\partial}{\partial x} \left(h\bar{u}^2 + \frac{1}{2}kg_z h^2 \right) + \frac{\partial}{\partial y}(h\bar{u}\bar{v}), \frac{\partial}{\partial x}(h\bar{u}\bar{v}) + \frac{\partial}{\partial y} \left(h\bar{v}^2 + \frac{1}{2}kg_z h^2 \right) \right], \quad (19)$$

it is the material derivative of momentum, due to the advection.

Right-hand side, RHS forces

$$\mathbf{RHS}_1 = [g_x h, g_y h], \quad (20)$$

it is the gravitational force term, it has the same formulation in all models.

The formula of basal friction force \mathbf{RHS}_2 depends on the model:

$$\begin{aligned} \mathbf{RHS}_2 &= -hg_z \tan(\phi_{bed}) \left[\frac{\bar{u}}{\|\tilde{\mathbf{u}}\|}, \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \right], \text{ in MC model.} \\ \mathbf{RHS}_2 &= -hg_z \mu_b(\|\tilde{\mathbf{u}}\|, h) \left[\frac{\bar{u}}{\|\tilde{\mathbf{u}}\|}, \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \right], \text{ in PF model.} \\ \mathbf{RHS}_2 &= -hg_z \mu \left[\frac{\bar{u}}{\|\tilde{\mathbf{u}}\|}, \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \right], \text{ in VS model.} \end{aligned} \quad (21)$$

The formula of the force related to the topography curvature, \mathbf{RHS}_3 , also depends on the model:

$$\begin{aligned} \mathbf{RHS}_3 &= h \tan(\phi_{bed}) \left[-\frac{\bar{u}^3}{r_x \|\tilde{\mathbf{u}}\|}, -\frac{\bar{v}^3}{r_y \|\tilde{\mathbf{u}}\|} \right], \text{ in MC model.} \\ \mathbf{RHS}_3 &= h \mu_b(\|\tilde{\mathbf{u}}\|, h) \left[-\frac{\bar{u}^3}{r_x \|\tilde{\mathbf{u}}\|}, -\frac{\bar{v}^3}{r_y \|\tilde{\mathbf{u}}\|} \right], \text{ in PF model.} \\ \mathbf{RHS}_3 &= h \mu \left[-\frac{\bar{u} \|\tilde{\mathbf{u}}\|}{r_x} \mu, 0 \right], \text{ in VS model.} \end{aligned} \quad (22)$$

All the three models have an additional force term, having a different formula and meaning in the three models:

$$\begin{aligned} \mathbf{RHS}_4 &= -hk_{ap} \sin(\phi_{int}) \left[\operatorname{sgn} \left(\frac{\partial \bar{u}}{\partial y} \right) \frac{\partial(g_z h)}{\partial y}, \operatorname{sgn} \left(\frac{\partial \bar{v}}{\partial x} \right) \frac{\partial(g_z h)}{\partial x} \right], \text{ in MC model.} \\ \mathbf{RHS}_4 &= g_z h \left[\frac{\partial h}{\partial x}, \frac{\partial h}{\partial y} \right], \text{ in PF model.} \\ \mathbf{RHS}_4 &= -\frac{\|\mathbf{g}\|}{\xi} \|\tilde{\mathbf{u}}\|^2 \left[\frac{\bar{u}}{\|\tilde{\mathbf{u}}\|}, \frac{\bar{v}}{\|\tilde{\mathbf{u}}\|} \right], \text{ in VS model.} \end{aligned} \quad (23)$$

3.3 Definition and consistency of the input variables

In general, the three rheologies considered in this study, MC, PF, and VS, have totally different parameters. The statistical testing we perform requires to choose the parameter range P_M , and in principle it may be arbitrary. Nevertheless, if the total frictions of the models are not covering a similar span, the statistical comparison is dominated by trivial macroscopic differences, and cannot focus on the rheology details. Some degree of parameter consistency is required between the models.

We can find several instances of parameter choices in literature, and in general the choice depends on the case study. In all the models it is defined a basal friction stress (see RHS_2 in in 4). However, PF interpolates different effective friction angles, while VS includes a velocity based additional term. We tested that a direct correlation of effective friction angles is problematic to do. Hence, we did not follow such approach. In contrast, we make a preliminary testing on the extreme values of the parameter space, i.e. giving ***max volume – min resistance***, and ***min volume – max resistance***.

General definitions We assume what follows to simplify testing. Except when differently specified, parameters are sampled uniformly in linear scale.

- Material Volume V is an additional input parameter in all the models, on the same range.
- In MC, sampled input parameters are ϕ_{bed} , and $\Delta\phi := \phi_{int} - \phi_{bed}$. In particular, $\Delta\phi \in [2^\circ, 10^\circ]$ (Dalbey et al., 2008).
- In PF, sampled input parameters are ϕ_1 , $\Delta\phi_{12} := \phi_2 - \phi_1$, and β . In particular, $\Delta\phi_{12} \in [10^\circ, 15^\circ]$, and $\beta \in [0.1, 0.85]$. Moreover, $\phi_3 = \phi_1 + 1^\circ$, and \mathcal{L} is equal to 1 dm and 1 mm in the two case studies, respectively (Pouliquen and Forterre, 2002; Forterre and Pouliquen, 2003; Gray and Edwards, 2014; Barker et al., 2015).
- In VS, sampled input parameters are μ , and ξ . In detail, ξ uniform sampling is accomplished in log-scale. In fact, values of ξ between 250 and 4,000 m/s^2 have been described for snow avalanches (Salm, 1993; Bartelt et al., 1999; Gruber and Bartelt, 2007).

In summary, MC and VS have three-dimensional parameter spaces, while PF a four-dimensional parameter space.

Flow down an inclined plane In this case study, Dalbey et al. (2008) assumed $\phi_{bed} = [15^\circ, 30^\circ]$, while Webb and Bursik (2016) performed a series of laboratory experiments and found $\phi_{bed} = [18.2^\circ, 34.4^\circ]$. We relied on those published parameter choices to decide a comprehensive parameter range. Figure 4 displays the maps of max flow height and max velocities observed in the extreme cases tested. Simulation options are - `max_time = 2 s`, `height/radius = 1.34`, `length_scale = 1 m`, `number_of_cells_across_axis = 10`, `order = first`, `geoflow_tiny = 1e4` (Patra et al., 2005; Aghakhani et al., 2016).

- **Material Volume:** $[449.0, 607.0] \text{ cm}^3$, i.e. average of 528.0 cm^3 and uncertainty of $\pm 15\%$.
- **Rheology models' parameter space:**

The parameter ranges adopted in this case study are:

MC - $\phi_{bed} \in [18^\circ, 30^\circ]$.

PF - $\phi_1 \in [10^\circ, 22^\circ]$.

VS - $\mu \in [0.22, 0.45]$, $\log(\xi) \in [3, 4]$.

Even if maximum and minimum runout are both matching, the shape and lateral extent of the flow is remarkably different in the max. runout case, according to the three models. In particular, MC model can produce the largest lateral extent, and VS model displays an accentuated wedge-like shape, due to the increased friction in the lateral margins. These features will be detailed in Section 4.1.

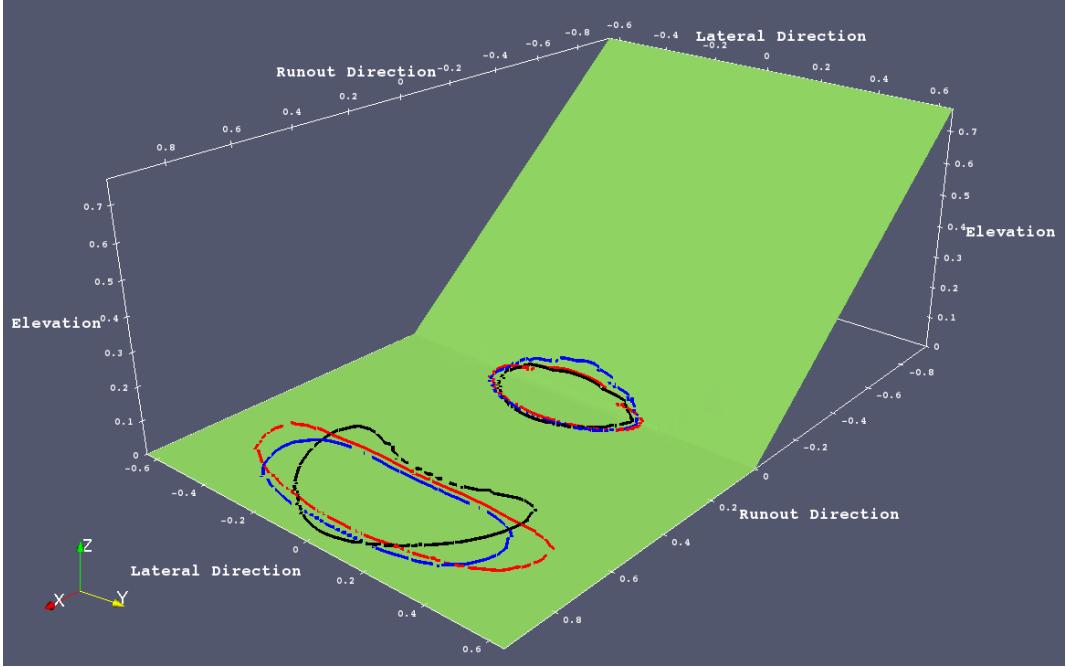


Figure 4: Inclined plane setup. Contours of $h = 1.0$ mm at last simulated snapshot ($t = 1.5$ sec) for simulated flows with *minimum runout* obtained from ***min volume – max resistance***, and *maximum runout* obtained from ***max volume – min resistance***. — : MC, — : PF, — : VS.

Volcán de Colima block and ash flow In this case study, Dallbey et al. (2008) assumed $\phi_{bed} = [15^\circ, 35^\circ]$, while (Capra et al., 2011) adopted $\phi_{bed} = 30^\circ$ in the simulation of a BAF in that same setting. Moreover, Spiller et al. (2014); Bayarri et al. (2015); Ogburn et al. (2016) found a statistical correlation between flow size and effective basal friction inferred from field observation of geophysical flows. The size of the BAF of this study would have $\phi_{bed} = [13^\circ, 18^\circ]$ according to their estimates. Figure 5 displays the maps of max flow height and max velocities observed in the extreme cases tested. Simulation options are - max_time = 7200 s (2 hours), height/radius = 0.55, length_scale = 4e3 m, number_of_cells_across_axis = 50, order = first, geoflow_tiny = 1e4 (Patra et al., 2005; Aghakhani et al., 2016).

- **Material Volume:** $[2.08, 3.12] \times 10^5 \text{ m}^3$, i.e. average of $2.6 \times 10^5 \text{ m}^3$ and uncertainty of $\pm 20\%$.
- **Rheology models' parameter space:**

The parameter ranges adopted in this case study are:

$$\text{MC} - \phi_{bed} \in [10^\circ, 25^\circ].$$

$$\text{PF} - \phi_1 \in [8^\circ, 18^\circ].$$

$$\text{VS} - \mu \in [0.15, 0.45], \quad \log(\xi) \in [1.7, 4].$$

The models possess different features, even if the maximum runout after channelization in the ravines is matching. In particular, VS lateral spreading is significant lower and material reaches higher thickness, whereas PF model seems to stop more gradually than MC with a more complex inundated area boundary lines. These features will be detailed in Section 5.2.



Figure 5: Volcán de Colima BAF. Comparison between *max flow height* maps of simulated flow, assuming Mohr-Coulomb (a),(b), Pouliquen-Forterre (c),(d), and Voellmy-Salm (e),(f) rheology. Extreme cases producing maximum and minimum runout, i.e. (a),(c),(e) **max volume – min resistance** and (b),(d),(f) **min volume – max resistance**.

4 QoIs and Data Collected - Flow down and inclined plane, and a change in slope to a flat plane

First case study is the small scale experimental setting described above. First we describe the observable outputs (Section 4.1), and then the hidden dynamic quantities (Section 4.2). In many plots we note the effect of cutting the flow when height is < 1 mm, avoiding the issues related to consider too thin flows (Aghakhani et al., 2016).

4.1 Observable outputs

Observable outputs include the flow height, speed, Fr , and acceleration as a function of time, measure on the four location L_1, \dots, L_4 displayed in Fig. 1. In addition, max lateral extension, flow area, and spatially averaged speed and Fr are displayed. UQ is also performed, accordingly to the parameter ranges described in Section 3.3.

4.1.1 Flow Height, local measurements

Figure 6 shows the flow height, $h(L, t)$, at the points $(L_i)_{i=1, \dots, 4}$, for the three rheology models.

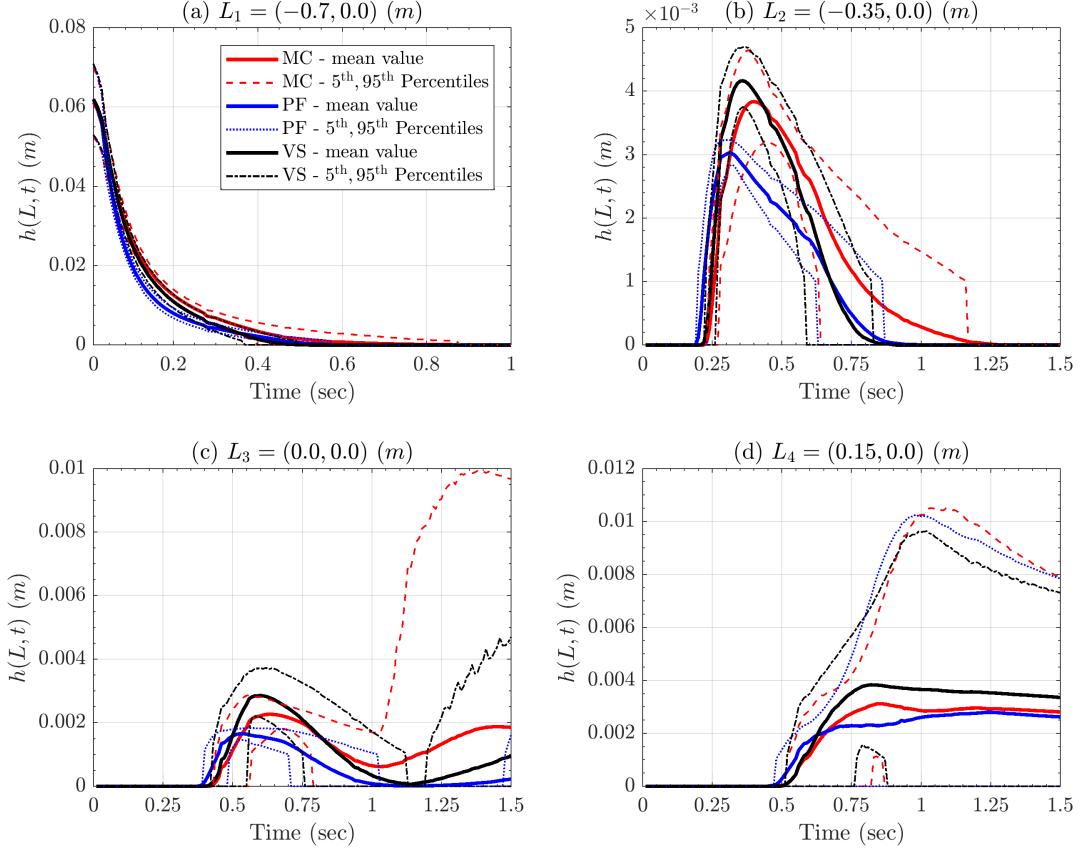


Figure 6: Records of flow height at four spatial locations of interest. Bold line is mean value, dashed/dotted lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. Plots are at different scale, for simplifying lecture.

In plot (a), related to point L_1 placed on the initial pile, the initial values of ~ 6 cm are the equal and only represent the pile height uncertainty, but then the flow height decreases slightly faster for the PF model, and slower for the MC, compared to the VS. Difference are more significant in plot (b), related to point L_2 , placed in the middle of the slope. Flow max height is greater for the VS, and smaller for the PF model, which decreases slower after the height peak. However, the value is ~ 15 times smaller than initial pile height, of ~ 4 mm. None of the models leaves a material deposit in L_1 or L_2 , i.e. the 95th percentile of the height is null at the end-time. In contrast, a deposit is left at points L_3 , placed at the change in slope, and L_4 , in the middle of the flat runout. MC's deposit is higher than the other models' at L_3 , while it is not significantly different at L_4 between the three models. It is of ~ 3 mm in average, but can range from 0 to 8 mm. In plots (b) and (c) we the elimination of material when height is < 1 mm, cuts the 95th percentile to zero when its value is below that threshold.

4.1.2 Flow Speed, local measurements

Figure 7 shows the flow speed, $\|\underline{\mathbf{u}}\|(L, t)$, at the points $(L_i)_{i=1,\dots,4}$, for the three rheology models.

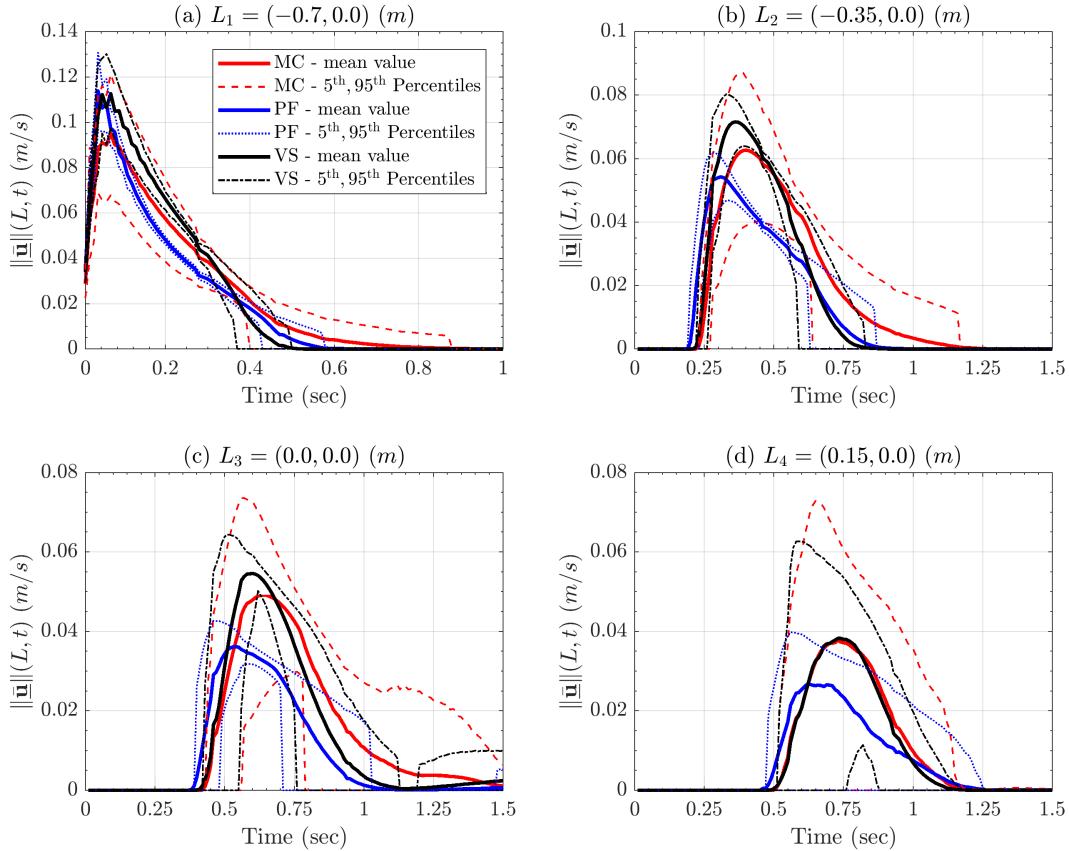


Figure 7: Records of flow speed at four spatial locations of interest. Bold line is mean value, dashed/dotted lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. Plots are at different scale.

In plot (a) and (b), related to points L_1 and L_2 , speed peaks are similar between the models, at ~ 10 cm/s and ~ 6 cm/s, respectively. In plot (a) VS peak is higher, but then its speed decreases more linearly, while MC and PF have a more concave profile, making MC the faster model after 0.5 s. In plot

(b) it is the PF model to decrease more linearly. UQ shows that MC model uncertainty is remarkably larger than in the other models, and produces higher values in the 95th percentile plots. Speed values below 1 cm/s are affected by the elimination of material below 1 mm flow height threshold. In plots (c) and (d), related to points L_3 and L_4 , the PF velocity profile is significantly lower than the other models, but it also decreases slower, and matches with the stopping times of them. Speed in the deposits is below 1 cm/s at L_3 , negligible at L_4 .

4.1.3 Froude Number, local measurements

Figure 8 shows the Froude Number, $\|\bar{\mathbf{u}}\|/\sqrt{gh}$, at the points $(L_i)_{i=1,\dots,4}$, for the three rheology models.

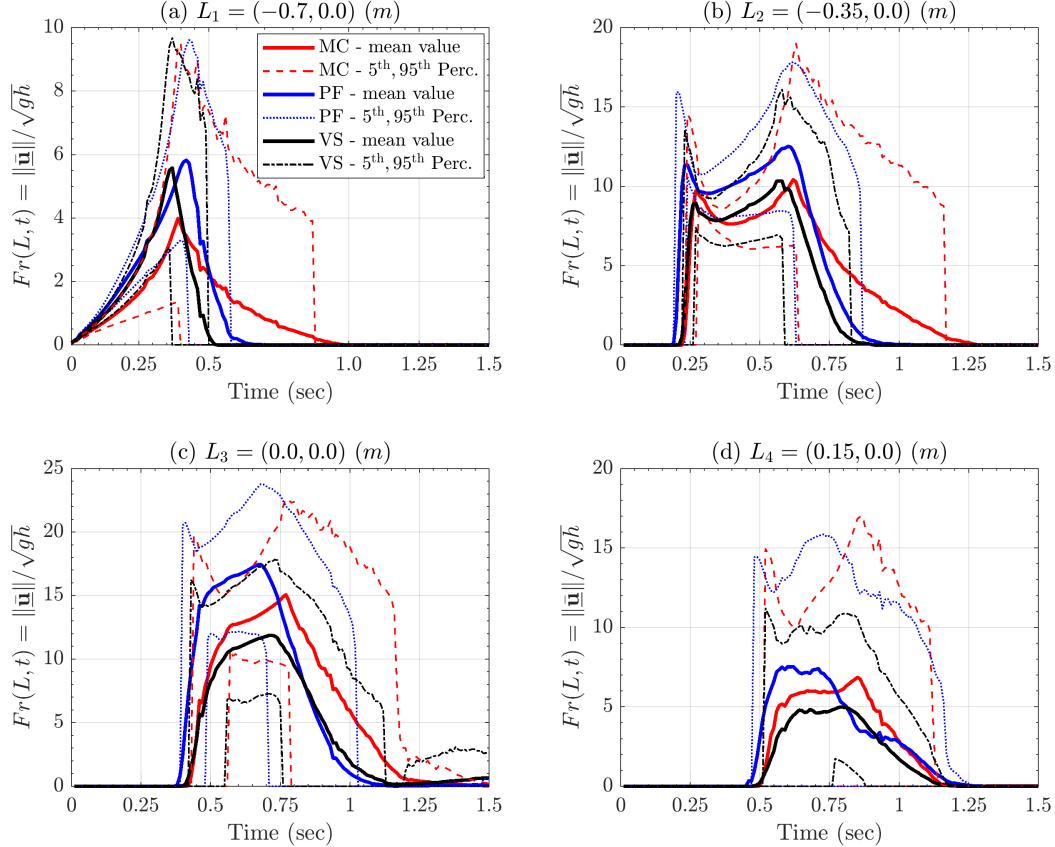


Figure 8: Records of Froude number at four spatial locations of interest. Bold line is mean value, dashed/dotted lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. Plots are at different scale.

Froude Number combines the estimates of flow height and speed, described above. In plot (a), related to point L_1 , Fr maximum average value is smaller for the MC model, ~ 4 , than for the others, ~ 6 . However, UQ tells us that 95th percentile almost reaches ~ 10 in all the three models. After the peak, the values decrease slower and more concavely in MC model than in the others. In plot (b), related to point L_2 , Fr shows a bimodal profile in time, with two separate peaks at ~ 10 on average, but reaching ~ 18 in the 95th percentile plot. In fact, first maximum, at ~ 0.25 s is due to the speed peak, and second, at ~ 0.6 s is related to \sqrt{h} decreasing while the speed is not significantly changing. In plot (c) and (d), related to points L_3 and L_4 , bimodality is less accentuated and becomes a plateau profile in $\sim [0.5, 0.75]$

s. PF model gives significantly larger Fr values, even > 20 at L_3 , due to a larger speed and a thinner flow. It is worth noting, for the sake of PF model, that $Fr > \beta$ and the flow is in the dynamic regime during the most of the time.

4.1.4 Flow Acceleration, local measurements

Figure 9 shows the flow speed, $\|\bar{a}\|(L, t)$, at the points $(L_i)_{i=1,\dots,4}$, for the three rheology models.

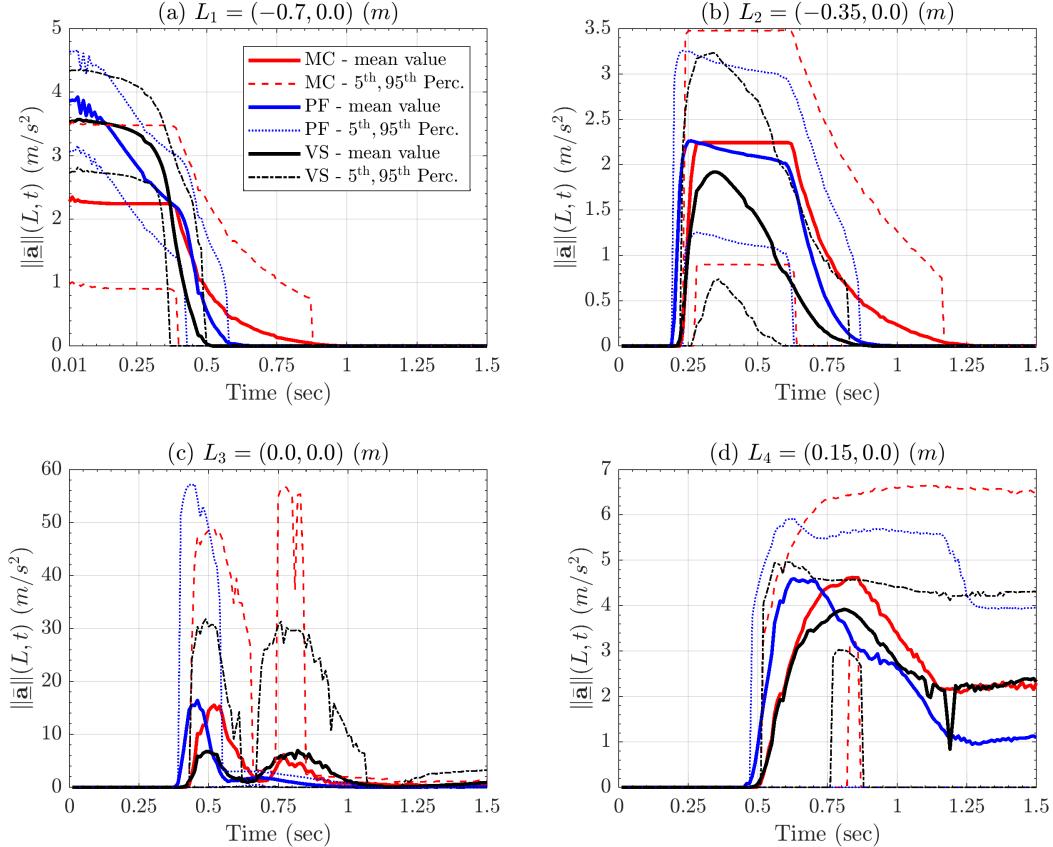


Figure 9: Records of flow acceleration magnitude (computed from LHS). Bold line is mean value, dashed lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. Plots are at different scale.

Acceleration is the link between force terms and observable motion. We calculated it from the LHS of the dynamical equation, but using the RHS terms produces very similar results, although not identical due to the numerical approximations in our solution. In plot (a), related to point L_1 , MC and VS show a plateau before ~ 0.4 s, at ~ 2.5 m/s 2 and ~ 3.5 m/s 2 , respectively, while PF linearly decreases between those same values. In plot (b), related to point L_2 , are MC and PF to show a plateau, at ~ 2.2 m/s 2 , while VS has a more bell-shaped profile. UQ tells us that PF has a smaller uncertainty than the other models. In plot (c), related to point L_3 , all the models show a bimodal profile, with peaks at ~ 0.5 s and 0.8 s. This is more accentuated for MC and VS, whereas the second peak is almost absent from PF profile. At the first peak, acceleration values are significant, with average peaks of MC and PF both at ~ 15 m/s 2 , and 95th percentile plot reaching ~ 50 m/s 2 and ~ 55 m/s 2 , respectively. VS shows about halved acceleration peak values. At the second peak, average acceleration values are similar for MC and

VS, at $\sim 5 \text{ m/s}^2$. In contrast, 95th percentile plot is $> 50 \text{ m/s}^2$ for MC, while $\sim 30 \text{ m/s}^2$ for VS. In plot (d), related to point L_4 , the acceleration has a first peak at $\sim 4 \text{ m/s}^2$, and a final asymptote at $\sim 2 \text{ m/s}^2$ for MC and VS, $\sim 1 \text{ m/s}^2$ for PF. These values generally mean flow deceleration, and uncertainty is more relevant for MC and PF than for VS.

4.1.5 Flow extents and averaged measurements (obsolete par. ranges, to update)

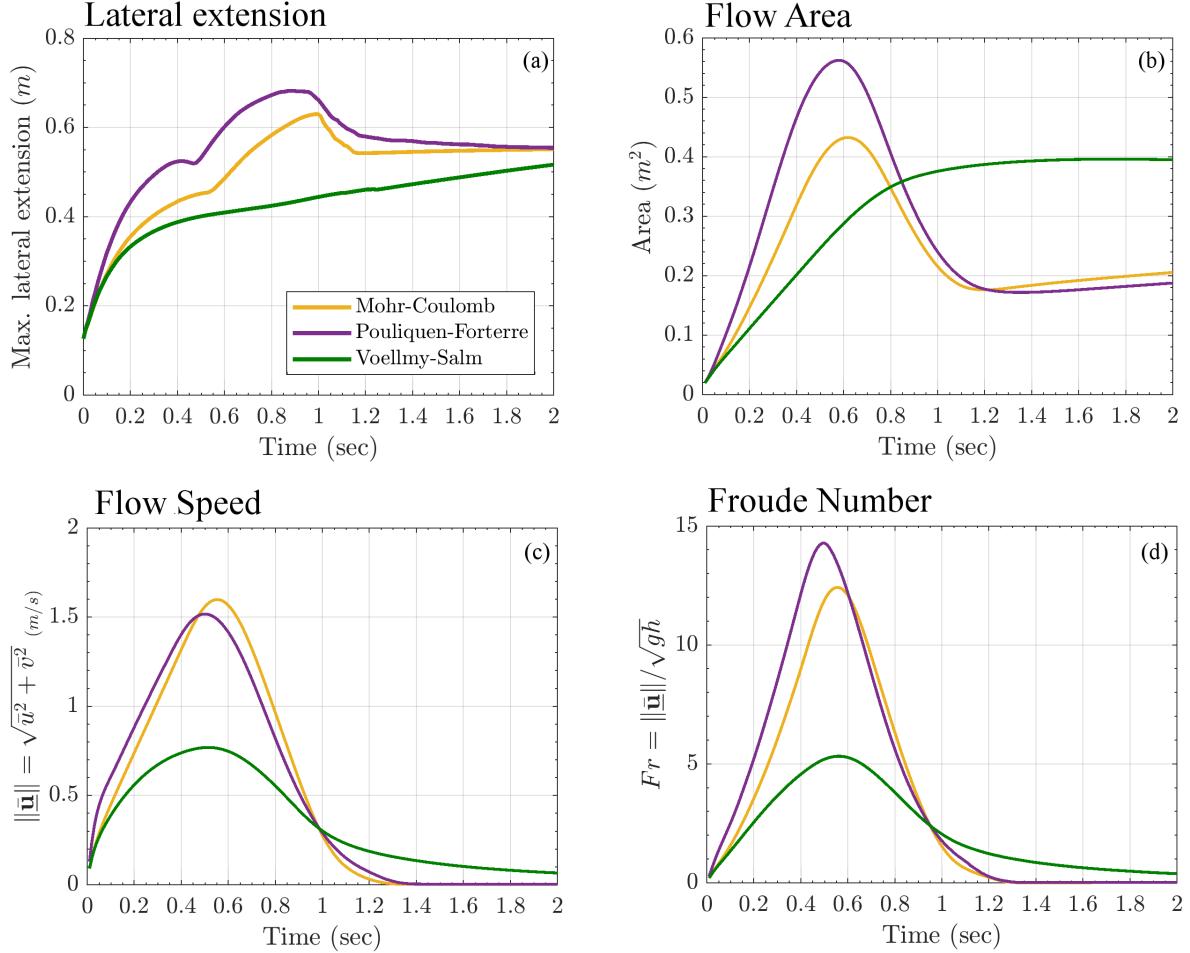


Figure 10: Comparison between mean values of flow (a) maximum lateral extent, and (b) areal extent, as a function of time. Spatial average of (c) flow speed, and (d) Froude Number.

4.2 Hidden dynamic quantities

These include force terms and related powers, as well as their statistical analysis.

4.2.1 Forces, spatial average (obsolete par. ranges, to update)

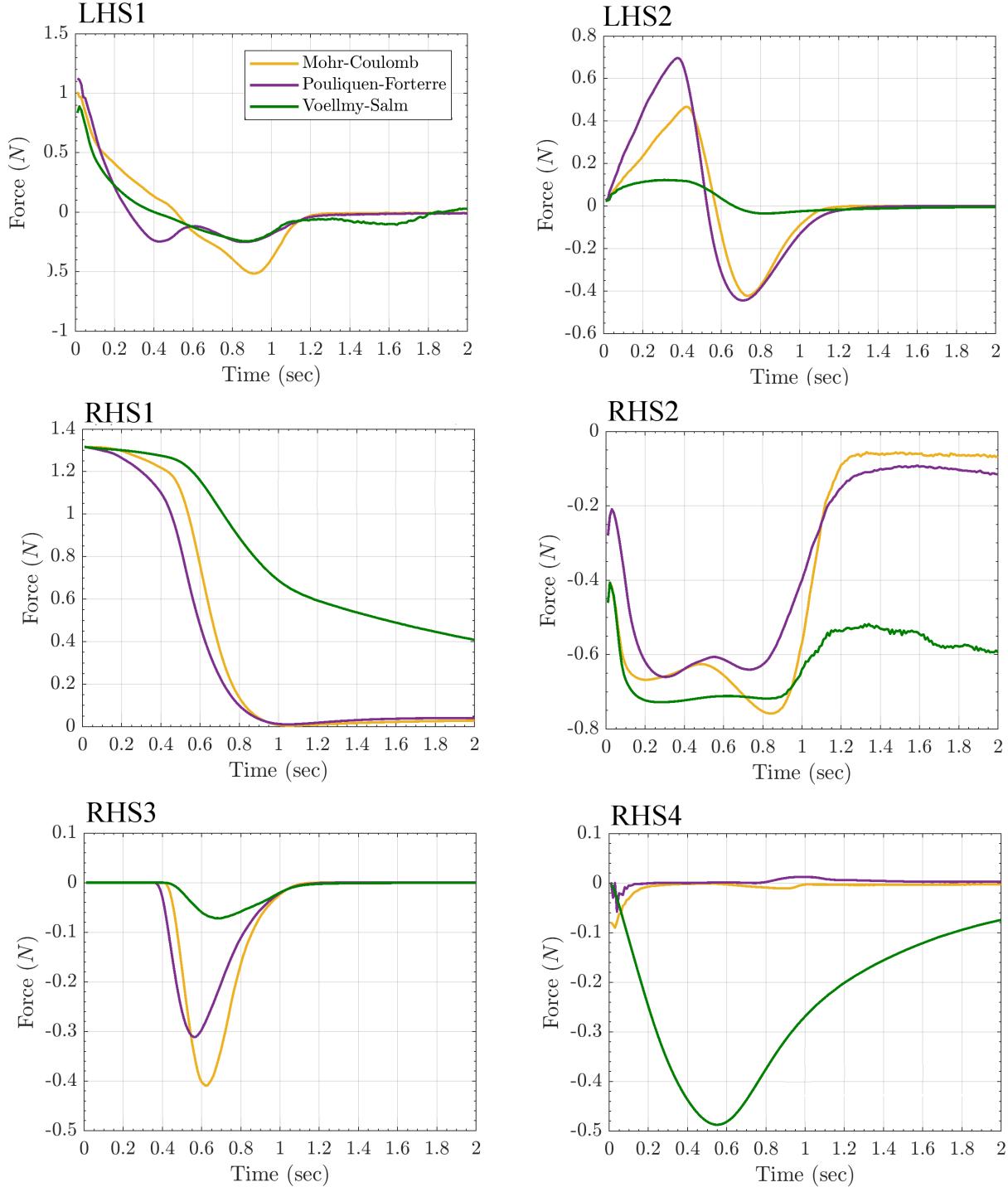


Figure 11: Spatial average of LHS and RHS forces in the slope direction. Bold line is mean value, dashed lines are 5th and 95th percentile bounds.

4.2.2 Powers, spatial average (obsolete par. ranges, to update)

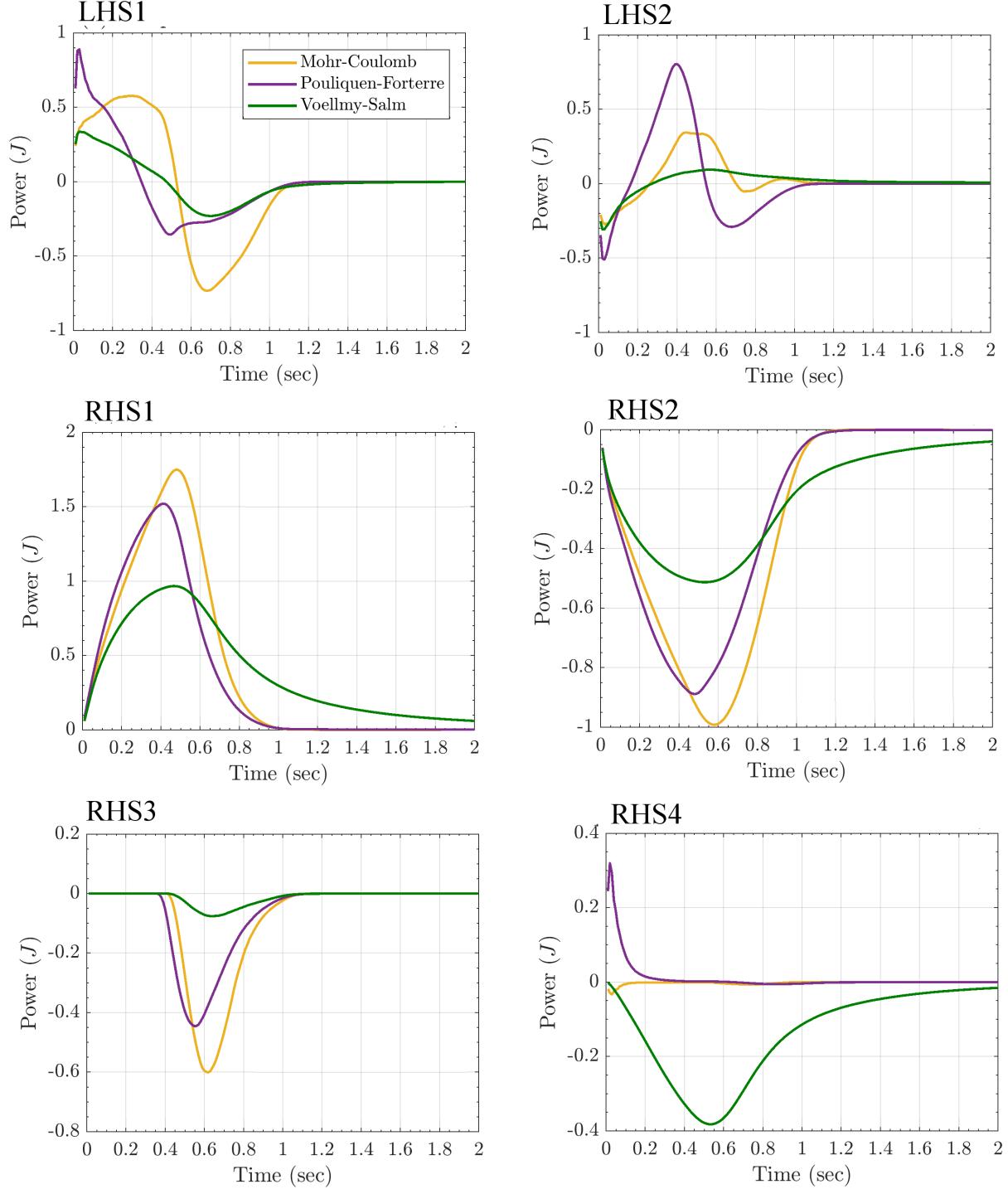


Figure 12: Spatial average of the LHS and RHS powers. Bold line is mean value, dashed lines are 5th and 95th percentile bounds. Model comparison on the mean value is also displayed.

4.2.3 Force contribution coefficients, RHS terms

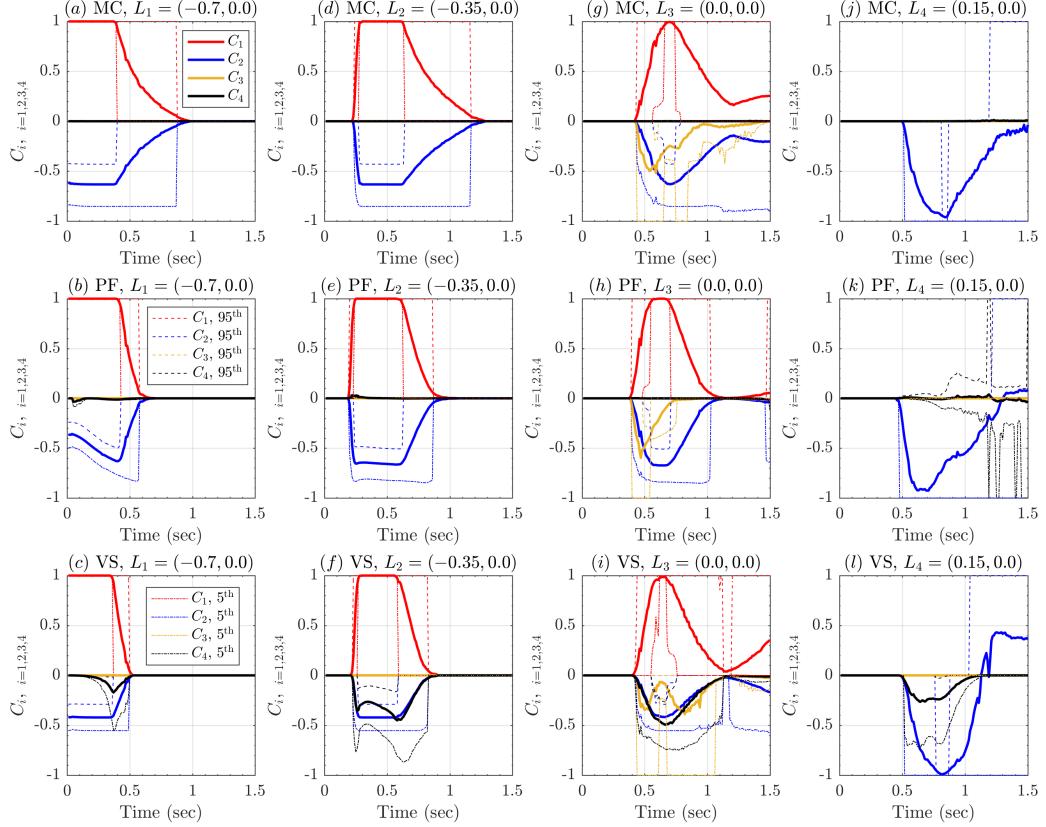


Figure 13: Records of contribution coefficients of **RHS** forces, in the slope direction, at four spatial locations of interest. Bold line is mean value, dashed/dotted lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. Dominant factor is calculated based on the l^1 norm (see Section 3.1)

4.2.4 Force dominance factors, RHS terms

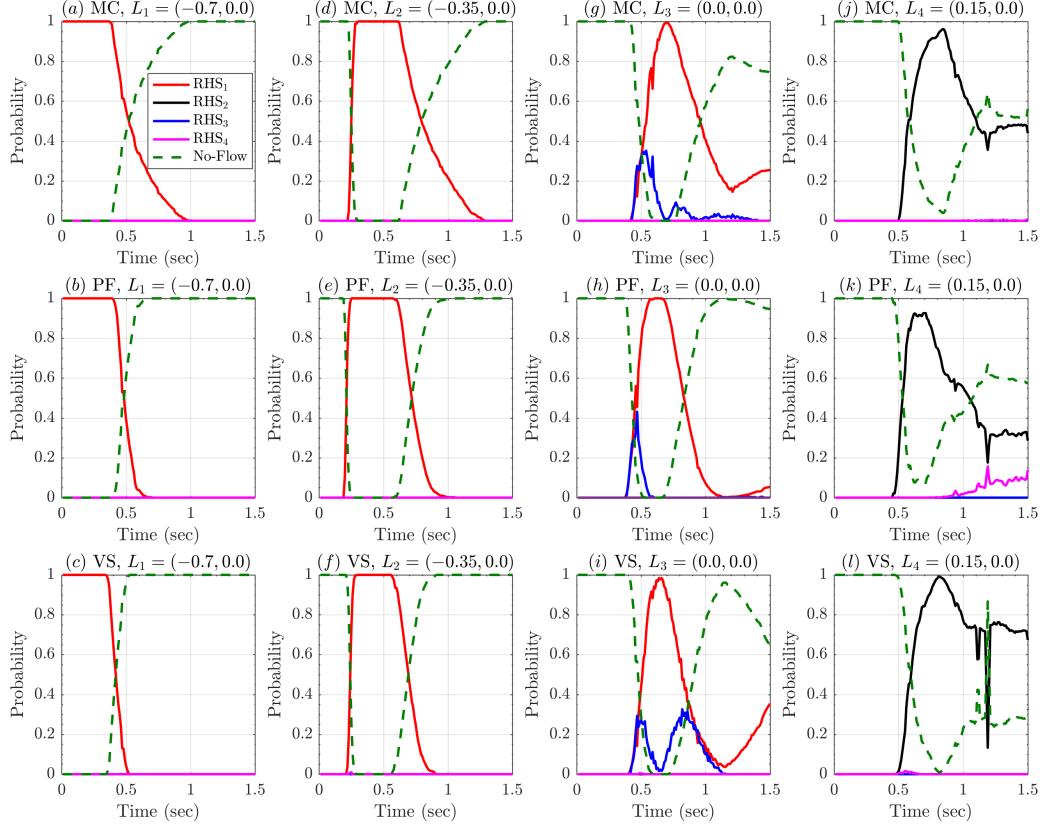


Figure 14: Records of dominance probabilities of **RHS** forces, in the slope direction, at four spatial locations of interest. Bold line is mean value, dashed/dotted lines are 5th and 95th percentile bounds. Different rheology models are displayed with different colors. No-flow probability is also displayed.

5 QoIs and Data Collected - Volcán de Colima BAF

5.1 Observable outputs overview - MC model

In our second case study, the number of spatial locations is significantly higher than previously. In fact, we placed 51 points to span the entire inundated area, in search of different flow regimes, displayed in Fig. 2. First we show the average flow height, speed, and Fr in all the locations, then, based on that, we select six points which we find representative of interesting flow regimes. Figures 15, 16, 17 assume MC model, while the results related to PF and VS models are included in the Supporting Material.

5.1.1 Average flow Height, local measurements.

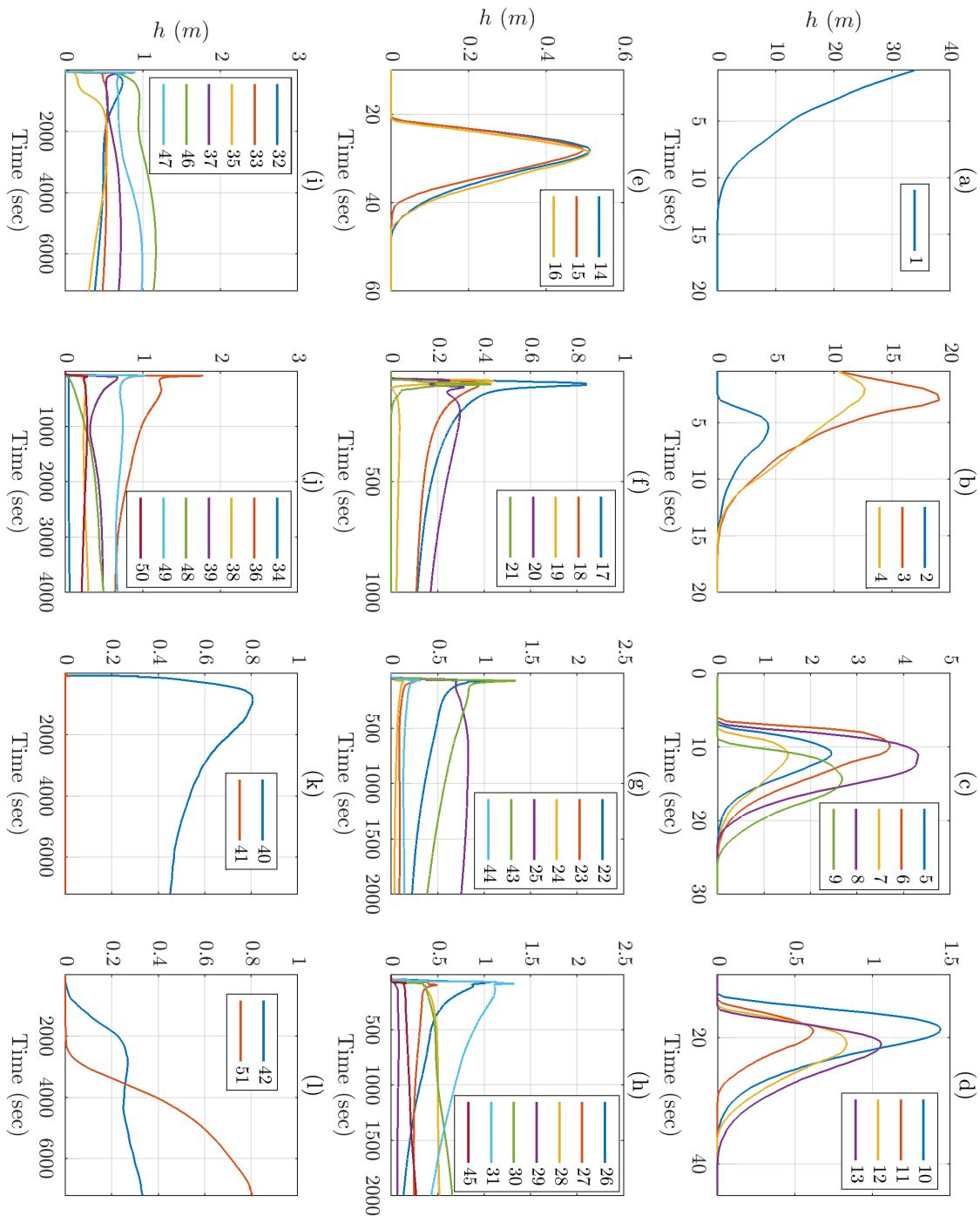


Figure 15: MC model, records of average flow height, $h(L, t)$, at 51 spatial locations of interest (Fig. 2).

5.1.2 Average flow Speed, local measurements.

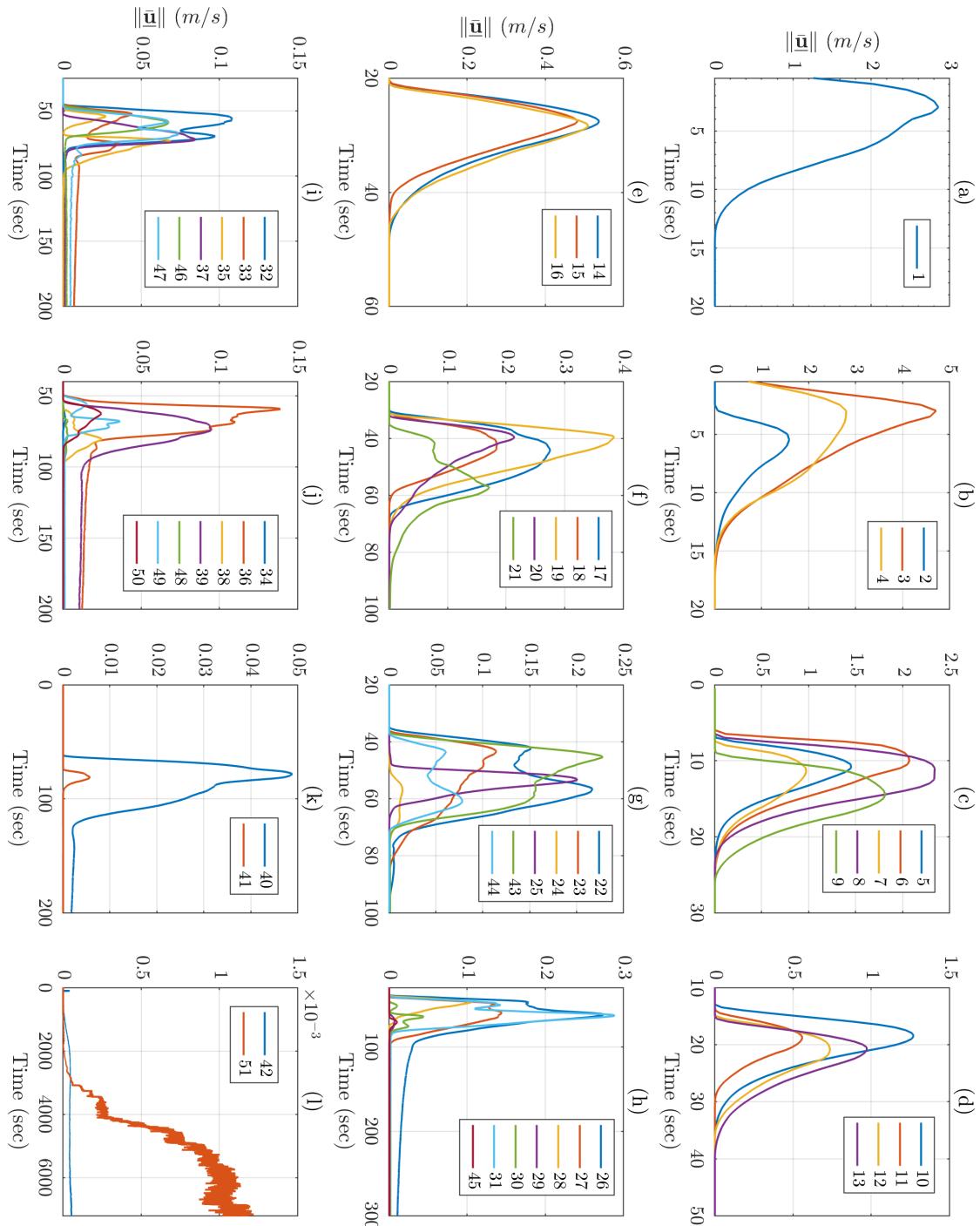


Figure 16: MC model, records of average flow speed, $|\underline{u}|(L, t)$, at 51 spatial locations of interest (Fig. 2).

5.1.3 Average Froude Number, local measurements.

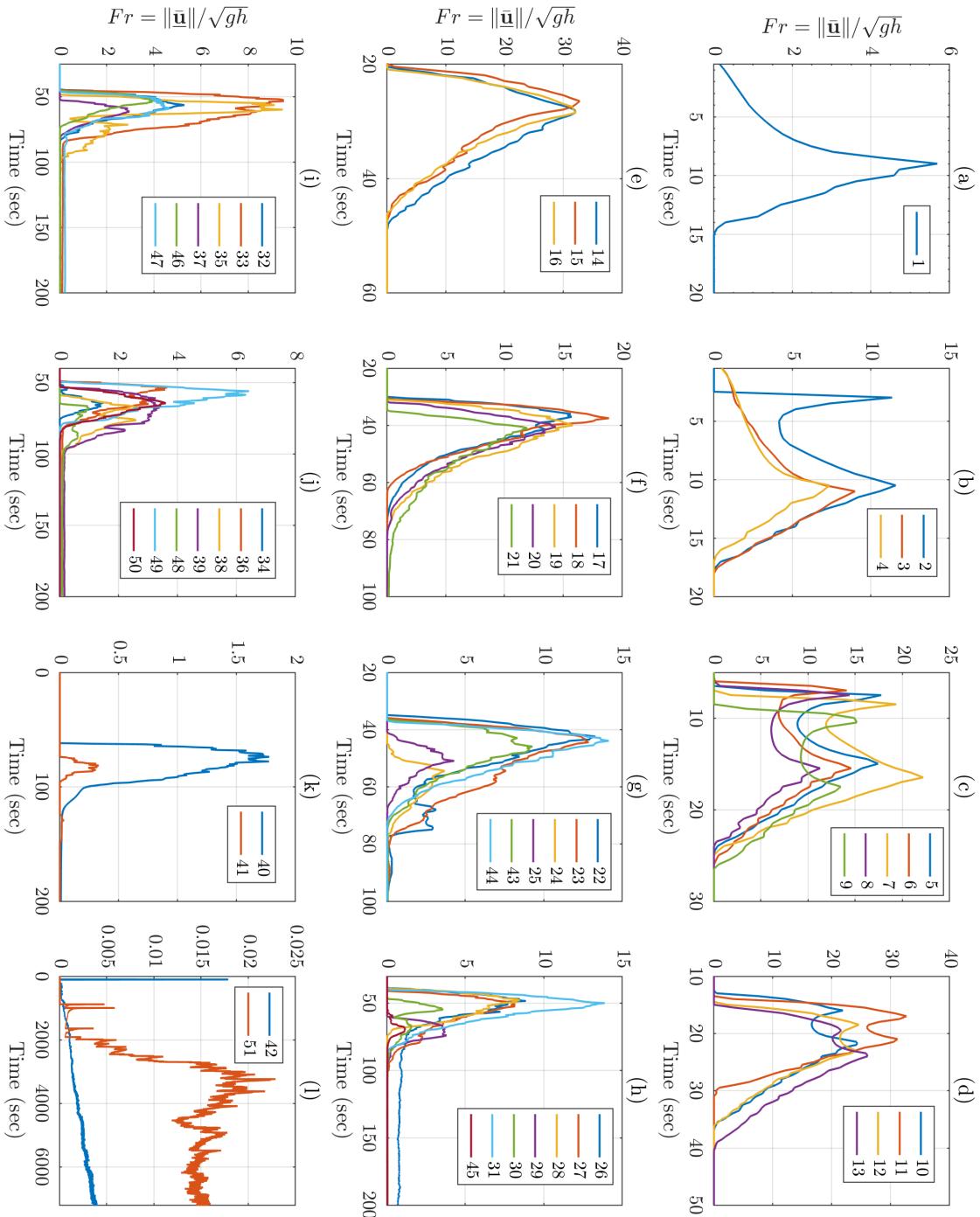


Figure 17: MC model, records of average Froude Number at 51 spatial locations of interest (Fig. 2).

5.2 Observable outputs - six selected locations

The six selected locations are $[L_8, L_{10}, L_{17}, L_{39}, L_{43}, L_{46}]$, (see Fig. 2). First two points, L_8 and L_{10} are both significantly close to the initiation pile, the points L_{17} and L_{43} are placed where the slope is reducing and the ravine channels start, and L_{39} and L_{46} are placed in the channels, further down-slope. Moreover, L_8 , L_{43} , and L_{46} are placed at the western side of the inundated area, whereas L_{10} , L_{17} , and L_{39} are placed at the eastern side.

5.2.1 Flow Height, selected measurements

5.2.2 Flow Speed, selected measurements

5.2.3 Flow Acceleration, selected measurements

5.2.4 Flow area and spatially averaged flow speed

5.3 Hidden dynamic quantities

The current case study does not show a preferential direction, equivalent to the slope direction of the inclined plane. Hence, we will directly focus on the powers related to the force terms detailed in Section 4.2.

5.3.1 Powers, spatial average

5.3.2 Power contribution coefficients, RHS terms

5.3.3 Power dominance factors, RHS terms

6 Discussion

Appendix A: Latin Hypercubes and orthogonal arrays

The Latin Hypercube Sampling (LHS) is a well established procedure for defining pseudo-random designs of samples in \mathbb{R}^d , with good properties with respect to the uniform probability distribution on an hypercube $[0, 1]^d$ (McKay et al., 1979; Owen, 1992a; Stein, 1987; Ranjan and Spencer, 2014; Ai et al., 2016). In particular, compared to a random sampling, a LHS: (i) enhances the capability to fill the d-dimensional space with a finite number of points, (ii) in case $d > 1$, avoids the overlapping of point locations in the one dimensional projections, (iii) reduces the dependence of the number of points necessary on the dimensionality d .

Definition 5 (Latin hypercube sampling) Let $\Xi = \{\xi_i : i = 1, \dots, N\}$ be a set of points inside the d -dimensional hypercube $C = [0, 1]^d$. Let $[0, 1] = \bigcup_{j=1}^N I_j$, where $I_j = [\frac{(j-1)}{N}, \frac{j}{N}]$. Let $\xi_i = (\xi_i^1, \dots, \xi_i^d)$, and for each $k \in \{1, \dots, d\}$, let $\Xi^k = \{\xi_i^k : i = 1, \dots, N\}$. Let λ^d be the uniform probability measure supported inside C , called Lebesgue measure. Then Ξ is a latin hypercube w.r.t. $\lambda^d \iff \forall j \in \{1, \dots, N\}, \forall k \in \{1, \dots, d\}, |I_j \cap \Xi^k| = 1$.

The procedure is simple: once the desired number of samples $N \in \mathbb{N}$ is selected, and $[0, 1]$ is divided in N equal bins, then each bin will contain one and only one projection of the samples over every coordinate. The LHS definition is trivially generalized over $C = \prod_i^d [a_i, b_i]$, i.e. the cartesian product of d arbitrary intervals. That will be applied in this study, defining LHS over the parameter domain of the flow models.

There are a large number of possible designs, corresponding the number of permutations of the bins in the d -projections, i.e. $d \cdot N!$. If the permutations are randomly sampled there is a high possibility that the design will have good properties. However, this is not assured, and clusters or points or regions of void space may be observed in C . For this reason, we base our design on the orthogonal arrays (OA) (Owen, 1992b; Tang, 1993).

Definition 6 (Orthogonal arrays) Let $S = \{1, \dots, s\}$, where $s \geq 2$. Let $Q \in S^{n \times m}$ be a matrix of such integer values. Then Q is called an $OA(n, m, s, r) \iff$ each $n \times r$ submatrix of Q contains all possible $1 \times r$ row vectors with the same frequency $\lambda = n/s^r$, which is called the index of the array. In particular, r is called the strength, n the size, ($m \geq r$) the constraints, and s the levels of the array.

Orthogonal arrays are very useful for defining latin hypercubes which are also forced to fill the space (or its r-dimensional subspaces) in a more robust way, at the cost of potentially requiring a larger number of points than a traditional LHS.

Proposition 7 *Let Q be an $OA(n, m, s, r)$. Then let $U \in \mathbb{R}^{n \times m}$ be defined as follows:*

$$\forall k \in \{1, \dots, s\}, \forall j \in \{1, \dots, m\}, \{Q[\cdot, j] : Q[i, j] = k\} = \Pi(\{(k-1)\lambda s^{r-1}, \dots, k\lambda s^{r-1}\}),$$

where Π is a random permutation of λs^{r-1} elements. Then $\Xi = \{\xi_i = U[i, \cdot] : i = 1, \dots, n\}$ is a LHS w.r.t to λ^m over $C = [0, 1]^m$. Moreover, let $[0, 1]^r = \bigcup_{(h_1, \dots, h_r)=1}^s I_{(h_i)}$, where $I_{(h_i)} = \prod_i [\frac{h_i-1}{s}, \frac{h_i}{s}]$. Then $\forall D = (d_1, \dots, d_r) \subseteq \{1, \dots, m\}$, let $\Xi^D = \{(\xi^{d_1}, \dots, \xi^{d_r}) : i = 1, \dots, n\}$. We have that

$$\forall k \in \{1, \dots, s\}, \forall (h_i : i = 1, \dots, r) \in \{1, \dots, d\}^r, |I_{(k_i)} \cap \Xi^D| = \lambda.$$

For each column of Q we are replacing the λs^{r-1} elements with entry k by a random permutation of $((k-1)\lambda s^{r-1} + h)_{h \in \{1, \dots, \lambda s^{r-1}\}}$. After the replacement procedure is done, the newly obtained matrix U is equivalent to a LHS which inherits from Q the property of fully covering s^r equal r-dimensional hypercubes in every r-dimensional projection. Each hypercube contains λ points. In other words, inside each r-dimensional projection, the design associated to U fills the space like a regular grid at the scale of those s^r hypercubes, but it is still an LHS at a finer scale, i.e. the λs^{r-1} one dimensional bins. A complete proof can be found in [Tang \(1993\)](#) and it is a straightforward verification of the required properties.

However, even in an LHS based on an $OA(n, m, s, r)$, if $r < m$ what happens in the projections with dimension $r' > r$ is not controlled, and randomizing procedures are made more difficult by the additional structure imposed by the OA . Moreover, the total number of points necessary to achieve a full design increases with r , and hence is affected by dimensionality issues.

Dealing with relatively small d , i.e. $d \in \{3, 4\}$, we adopt a LHS U created by a $OA(s^d, d, s, d)$. The strength is equal to the dimension d , hence the design fills the entire space like a d dimensional grid, but it is a LHS as well. In this case there is one point in each hypercube, and $\lambda = 1$. We take $s = 8$ for the 3-dimensional designs over the parameter space of Mohr-Coulomb and Voellmy-Salm models, i.e. 512 points; we took $s = 6$ for the 4-dimensional designs over the more complex parameter space of the Pouliquen-Forterre model, i.e. 1296 points.

References

- Aghakhani, H., Dalbey, K., Salac, D., and Patra, A. K. (2016). Heuristic and Eulerian interface capturing approaches for shallow water type flow and application to granular flows. *Computer Methods in Applied Mechanics and Engineering*, 304:243–264.
- Ai, M., Kong, X., and Li, K. (2016). A general theory for orthogonal array based latin hypercube sampling. *Statistica Sinica*, (26):761–777.
- Balmforth, N. J. and Kerswell, R. R. (2005). Granular collapse in two dimensions. *Journal of Fluid Mechanics*, 538:399–428.
- Barker, T., Schaeffer, D. G., Bohorquez, P., and Gray, J. M. N. T. (2015). Well-posed and ill-posed behaviour of the $\mu(i)$ -rheology for granular flow. *Journal of Fluid Mechanics*, 779:794–818.
- Bartelt, P. and McArdell, B. (2009). Granulometric investigations of snow avalanches. *Journal of Glaciology*, 55(193):829–833(5).
- Bartelt, P., Salm, B., and Gruber, U. (1999). Calculating dense-snow avalanche runout using a voellmy-fluid model with active/passive longitudinal straining. *Journal of Glaciology*, 45(150):242–254.
- Batchelor, G. (2000). *An Introduction to Fluid Dynamics*. Cambridge Mathematical Libraries. Cambridge University Press.
- Bayarri, M. J., Berger, J. O., Calder, E. S., Dalbey, K., Lunagomez, S., Patra, A. K., Pitman, E. B., Spiller, E. T., and Wolpert, R. L. (2009). Using Statistical and Computer Models to Quantify Volcanic Hazards. *Technometrics*, 51(4):402–413.
- Bayarri, M. J., Berger, J. O., Calder, E. S., P. A. K. P. E. B., Spiller, E. T., and Wolpert, R. L. (2015). Probabilistic Quantification of Hazards: A Methodology Using Small Ensembles of Physics-Based Simulations and Statistical Surrogates. *International Journal for Uncertainty Quantification*, 5(4):297–325.
- Bursik, M., Patra, A., Pitman, E. B., Nichita, C., Macias, J. L., Saucedo, R., and Girina, O. (2005). Advances in studies of dense volcanic granular flows. *Reports on Progress in Physics*, 68(2):271.
- Capra, L., Gavilanes-Ruiz, J. C., Bonasia, R., Saucedo-Giron, R., and Sulpizio, R. (2015). Re-assessing volcanic hazard zonation of volcán de colima, méxico. *Natural Hazards*, 76(1):41–61.
- Capra, L., Macís, J., Cortés, A., Dávila, N., Saucedo, R., Osorio-Ocampo, S., Arce, J., Gavilanes-Ruiz, J., Corona-Chávez, P., García-Sánchez, L., Sosa-Ceballos, G., and Vázquez, R. (2016). Preliminary report on the july 10–11, 2015 eruption at volcán de colima: Pyroclastic density currents with exceptional runouts and volume. *Journal of Volcanology and Geothermal Research*, 310(Supplement C):39 – 49.
- Capra, L., Manea, V. C., Manea, M., and Norini, G. (2011). The importance of digital elevation model resolution on granular flow simulations: a test case for colima volcano using titan2d computational routine. *Natural Hazards*, 59(2):665–680.
- Charbonnier, S., Germa, A., Connor, C., Gertisser, R., Preece, K., Komorowski, J.-C., Lavigne, F., Dixon, T., and Connor, L. (2013). Evaluation of the impact of the 2010 pyroclastic density currents at merapi volcano from high-resolution satellite imagery, field investigations and numerical simulations. *Journal of Volcanology and Geothermal Research*, 261(Supplement C):295 – 315. Merapi eruption.
- Charbonnier, S. J. and Gertisser, R. (2009). Numerical simulations of block-and-ash flows using the titan2d flow model: examples from the 2006 eruption of merapi volcano, java, indonesia. *Bulletin of Volcanology*, 71(8):953–959.
- Christen, M., Kowalski, J., and Bartelt, P. (2010). RAMMS: Numerical simulation of dense snow avalanches in three-dimensional terrain. *Cold Regions Science and Technology*, 63:1–14.
- DaCruz, F., Emam, S., Prochnow, M., Roux, J.-N., and Chevoir, F. (2005). Rheophysics of dense granular materials: Discrete simulation of plane shear flows. *Physical Review E*, 72:021309.

- Dade, W. B. and Huppert, H. E. (1998). Long-runout rockfalls. *Geology*, 26(9):803–806.
- Dalbey, K., Patra, A. K., Pitman, E. B., Bursik, M. I., and Sheridan, M. F. (2008). Input uncertainty propagation methods and hazard mapping of geophysical mass flows. *Journal of Geophysical Research: Solid Earth*, 113:1–16.
- Dalbey, K. R. (2009). *Predictive Simulation and Model Based Hazard Maps*. PhD thesis, University at Buffalo.
- Davila, N., Capra, L., Gavilanes-Ruiz, J., Varley, N., Norini, G., and Vazquez, A. G. (2007). Recent lahars at volcán de colima (mexico): Drainage variation and spectral classification. *Journal of Volcanology and Geothermal Research*, 165(3):127 – 141.
- Denlinger, R. P. and Iverson, R. M. (2001). Flow of variably fluidized granular masses across three-dimensional terrain: 2. Numerical predictions and experimental tests. *Journal of Geophysical Research*, 106(B1):553–566.
- Denlinger, R. P. and Iverson, R. M. (2004). Granular avalanches across irregular three-dimensional terrain: 1. Theory and computation. *Journal of Geophysical Research (Earth Surface)*, 109(F1):F01014.
- Edwards, A. N. and Gray, J. M. N. T. (2015). Erosion-deposition waves in shallow granular free-surface flows. *Journal of Fluid Mechanics*, 762:35–67.
- Fischer, J., Kowalski, J., and Pudasaini, S. P. (2012). Topographic curvature effects in applied avalanche modeling. *Cold Regions Science and Technology*, 74–75:21–30.
- Forterre, Y. (2006). Kapiza waves as a test for three-dimensional granular flow rheology. 563(*Journal of Fluid Mechanics*):123–132.
- Forterre, Y. and Pouliquen, O. (2002). Stability analysis of rapid granular chute flows: formation of longitudinal vortices. *Journal of Fluid Mechanics*, 467:361–387.
- Forterre, Y. and Pouliquen, O. (2003). Long-surface-wave instability in dense granular flows. *Journal of Fluid Mechanics*, 486:21–50.
- Forterre, Y. and Pouliquen, O. (2008). Flows of dense granular media. *Annual Review of Fluid Mechanics*, 40(1):1–24.
- Fraccarollo, L. and Toro, E. F. (1995). Experimental and numerical assessment of the shallow water model for two dimensional dam-break type problems. *Journal of Hydraulic Research*, 33:843–864.
- Freundt, A. and Bursik, M. I. (1998). Pyroclastic flow transport mechanisms. In Freundt, A. and Rosi, M., editors, *From Magma to Tephra: Modelling Physical Processes of Explosive Volcanism*, volume 4 of *Advances in Volcanology*, pages 151–222. Elsevier Science, Amsterdam.
- González, M. B., Ramírez, J. J., and Navarro, C. (2002). Summary of the historical eruptive activity of volcán de colima, mexico 1519–2000. *Journal of Volcanology and Geothermal Research*, 117(1):21 – 46.
- Gray, J. M. N. T. and Edwards, A. N. (2014). A depth-averaged $\mu(I)$ -rheology for shallow granular free-surface flows. *Journal of Fluid Mechanics*, 755:503–534.
- Gray, J. M. N. T., Tai, Y. C., and Noelle, S. (2003). Shock waves, dead zones and particle-free regions in rapid granular free-surface flows. *Journal of Fluid Mechanics*, 491:161–181.
- Gray, J. M. N. T., Wieland, M., and Hutter, K. (1999). Gravity-driven free surface flow of granular avalanches over complex basal topography. *Proceedings of the Royal Society A: Mathematical, Physical and Engineering Sciences*, 455(1985):1841–1874.
- Greve, R., Koch, T., and Hutter, K. (1994). Unconfined flow of granular avalanches along a partly curved surface: experiments and theoretical predictions. *Proc. R. Soc. London Ser. A*, 342:573–600.

- Gruber, U. and Bartelt, P. (2007). Snow avalanche hazard modelling of large areas using shallow water numerical models and GIS. *Environmrntal Modelling & Software*, 22:1472–1481.
- Hokanardottir, K. M. and Hogg, A. J. (2005). Oblique shocks in rapid granular flows. *Phys. Fluids*, 17:077101.
- Hutter, K., Siegel, M., Savage, S. B., and Nohguchi, Y. (1993). Two-dimensional spreading of a granular avalanche down an inclined plane part I. Theory. *Acta Mechanica*, 100:37–68.
- Iordanoff, I. P. and Khonsari, M. M. (2004). Granular Lubrication: Toward an Understanding of the Transition Between Kinetic and Quasi-Fluid Regime. *Journal of Tribology*, 126(1):137–145.
- Iverson, R. M. (1997). The physics of debris flows. *Reviews of Geophysics*, 35:245–296.
- Iverson, R. M. and Denlinger, R. P. (2001). Flow of variably fluidized granular masses across three-dimensional terrain: 1. Coulomb mixture theory. *Journal of Geophysical Research*, 106(B1):537.
- Iverson, R. M. and George, D. L. (2014). A depth-averaged debris-flow model that includes the effects of evolving dilatancy. i. physical basis. *Proceedings of the Royal Society of London A: Mathematical, Physical and Engineering Sciences*, 470(2170).
- Iverson, R. M., Logan, M., and Denlinger, R. P. (2004). Granular avalanches across irregular three-dimensional terrain: 2. experimental tests. *Journal of Geophysical Research: Earth Surface*, 109(F1):n/a–n/a. F01015.
- Jaeger, H. M., Liu, C., and Nagel, S. R. (1989). Relaxation at the Angle of Repose. *Physical Review Letters*, 62(1):40–43.
- Jop, P., Forterre, Y., and Pouliquen, O. (2006). A constitutive law for dense granular flows. *Nature*, 441:727–730.
- Kelfoun, K. (2011). Suitability of simple rheological laws for the numerical simulation of dense pyroclastic flows and long-runout volcanic avalanches. *Journal of Geophysical Research*, 116:B08209.
- Kelfoun, K. and Druitt, T. H. (2005). Numerical modeling of the emplacement of socompa rock avalanche, chile. *Journal of Geophysical Research: Solid Earth*, 110(B12).
- Kelfoun, K., Samaniego, P., Palacios, P., and Barba, D. (2009). Testing the suitability of frictional behaviour for pyroclastic flow simulation by comparison with a well-constrained eruption at Tungurahua volcano (Ecuador). *Bulletin of Volcanology*, 71:1057–1075.
- Kern, M., Bartelt, P., Sovilla, B., and Buser, O. (2009). Measured shear rates in large dry and wet snow avalanches. *Journal of Glaciology*, 55(190):327–338.
- la Cruz-Reyna, S. D. (1993). Random patterns of occurrence of explosive eruptions at colima volcano, mexico. *Journal of Volcanology and Geothermal Research*, 55(1):51 – 68.
- Lajeunesse, E., Monnier, J. B., and Homsy, G. M. (2005). Granular slumping on a horizontal surface. *Phys. Fluids*, 17:103302.
- Luca, I., Tai, Y., and Kuo, C. (2016). *Shallow Geophysical Mass Flows down Arbitrary Topography: Model Equations in Topography-fitted Coordinates, Numerical Simulation and Back-calculations of Disastrous Events*. Advances in Geophysical and Environmental Mechanics and Mathematics. Springer International Publishing.
- Mangeney-Castelnau, A., Bouchut, F., Thomas, N., Vilotte, J. P., and Bristeau, M.-O. (2007). Numerical modeling of self-channeling granular flows and of their levee/channel deposits. *J. Geophys. Res.*, 112:F02017.
- Mangeney-Castelnau, A., Bouchut, F., Vilotte, J. P., Lajeunesse, E., Aubertin, A., and Pirulli, M. (2005). On the use of Saint-Venant equations to simulate the spreading of a granular mass. *J. Geophys. Res.*, 110:B09103.

- Martin Del Pozzo, A. M., Sheridan, M. F., Barrera, M., Hubp, J. L., and Selem, L. V. (1995). Potential hazards from Colima Volcano, Mexico. *Geofisica Internacional*, 34:363–376.
- McKay, M. D., Beckman, R. J., and Conover, W. J. (1979). A comparison of three methods for selecting values of input variables in the analysis of output from a computer code. *Technometrics*, 21(2):239–245.
- Norini, G., De Beni, E., Andronico, D., Polacci, M., Burton, M., and Zucca, F. (2009). The 16 november 2006 flank collapse of the south-east crater at mount etna, italy: Study of the deposit and hazard assessment. *Journal of Geophysical Research: Solid Earth*, 114(B2):n/a–n/a. B02204.
- Ogburn, S. E., Berger, J., Calder, E. S., Lopes, D., Patra, A., Pitman, E. B., Rutarindwa, R., Spiller, E., and Wolpert, R. L. (2016). Pooling strength amongst limited datasets using hierarchical bayesian analysis, with application to pyroclastic density current mobility metrics. *Statistics in Volcanology*, 2:1–26.
- Owen, A. B. (1992a). A Central Limit Theorem for Latin Hypercube Sampling. *Journal of the Royal Statistical Society.*, 54(2):541–551.
- Owen, A. B. (1992b). Orthogonal arrays for computer experiments, integration and visualization. *Statistica Sinica*, 2(2):439–452.
- Patra, A., Nichita, C., Bauer, A., Pitman, E., Bursik, M., and Sheridan, M. (2006). Parallel adaptive discontinuous galerkin approximation for thin layer avalanche modeling. *Computers & Geosciences*, 32(7):912 – 926. Computer Simulation of natural phenomena for Hazard Assessment.
- Patra, A. K., Bauer, A. C., Nichita, C. C., Pitman, E. B., Sheridan, M. F., Bursik, M., Rupp, B., Webber, A., Stinton, A. J., Namikawa, L. M., and Renschler, C. S. (2005). Parallel adaptive numerical simulation of dry avalanches over natural terrain. *Journal of Volcanology and Geothermal Research*, 139(1-2):1–21.
- Pirulli, M., Bristeau, M. O., Mangeney, A., and Scavia, C. (2007). The effect of the earth pressure coefficients on the runout of granular material. *Environmental Modelling & Software*, 22(10):1437–1454.
- Pirulli, M. and Mangeney, A. (2008). Results of Back–Analysis of the Propagation of Rock Avalanches as a Function of the Assumed Rheology. *Rock Mechanics and Rock Engineering*, 41(1):59–84.
- Pitman, E., Nichita, C., Patra, A., Bauer, A., Bursik, M., and Webb, A. (2003a). A model of granular flows over an erodible surface. *Discrete and Continuous Dynamical Systems - Series B*, 3(4):589–599.
- Pitman, E. B. and Le, L. (2005). A two-fluid model for avalanche and debris flows. *Philosophical transactions. Series A, Mathematical, physical, and engineering sciences*, 363(1832):1573–601.
- Pitman, E. B., Nichita, C. C., Patra, A., Bauer, A., Sheridan, M., and Bursik, M. (2003b). Computing granular avalanches and landslides. *Physics of Fluids*, 15(12):3638.
- Pouliquen, O. (1999). Scaling laws in granular flows down rough inclined planes. *Physics of Fluids*, 11(3):542–548.
- Pouliquen, O. and Forterre, Y. (2002). Friction law for dense granular flows: application to the motion of a mass down a rough inclined plane. *Journal of Fluid Mechanics*, 453:133–151.
- Procter, J. N., Cronin, S. J., Platz, T., Patra, A., Dalbey, K., Sheridan, M., and Neall, V. (2010). Mapping block-and-ash flow hazards based on titan 2d simulations: a case study from mt. taranaki, nz. *Natural Hazards*, 53(3):483–501.
- Pudasaini, S. and Hutter, K. (2007). *Avalanche Dynamics: Dynamics of Rapid Flows of Dense Granular Avalanches*. Springer.
- Pudasaini, S. P. and Hutter, K. (2003). Rapid shear flows of dry granular masses down curved and twisted channels. *Journal of Fluid Mechanics*, 495:193–208.

- Ranjan, P. and Spencer, N. (2014). Space-filling latin hypercube designs based on randomization restrictions in factorial experiments. *Statistics and Probability Letters*, 94(Supplement C):239 – 247.
- Rankine, W. J. M. (1857). On the Stability of Loose Earth. *Phil. Trans. R. Soc. Lond.*, 147(2):9–27.
- Reyes-Dávila, G. A., Arámbula-Mendoza, R., Espinasa-Perena, R., Pankhurst, M. J., Navarro-Ochoa, C., Savov, I., Vargas-Bracamontes, D. M., Cortés-Cortés, A., Gutiérrez-Martínez, C., Valdés-González, C., Domínguez-Reyes, T., González-Amezcuia, M., Martínez-Fierros, A., Ramírez-Vázquez, C. A., Cárdenas-González, L., Castaneda-Bastida, E., de los Monteros, D. M. V. E., Nieto-Torres, A., Champion, R., Courtous, L., and Lee, P. D. (2016). Volcan de colima dome collapse of july, 2015 and associated pyroclastic density currents. *Journal of Volcanology and Geothermal Research*, 320(Supplement C):100 – 106.
- Rupp, B. (2004). An analysis of granular flows over natural terrain. Master's thesis, University at Buffalo.
- Rupp, B., Bursik, M., Namikawa, L., Webb, A., Patra, A. K., Saucedo, R., Macías, J. L., and Renschler, C. (2006). Computational modeling of the 1991 block and ash flows at Colima Volcano, México. *Geological Society of America Special Papers*, 402(11):223–237.
- Buyer-Quil, C. and Manneville, P. (2000). Improved modeling of flows down inclined planes. *Eur. Phys. J. B*, 15:357–369.
- Salm, B. (1993). Flow, flow transition and runout distances of flowing avalanches. *Annals of Glaciology*, 18:221–226.
- Salm, B., Burkard, A., and Gubler, H. (1990). Berechnung von Fliesslawinen: eine Anleitung für Praktiker mit Beispielen. *Mitteilungen des Eidgenössische Institutes für Schnee- und Lawinenforschung*, 47.
- Sarocchi, D., Sulpizio, R., Macías, J., and Saucedo, R. (2011). The 17 july 1999 block-and-ash flow (baf) at colima volcano: New insights on volcanic granular flows from textural analysis. *Journal of Volcanology and Geothermal Research*, 204(1):40 – 56.
- Saucedo, R., Macías, J., Bursik, M., Mora, J., Gavilanes, J., and Cortes, A. (2002). Emplacement of pyroclastic flows during the 1998–1999 eruption of volcán de colima, méxico. *Journal of Volcanology and Geothermal Research*, 117(1):129 – 153.
- Saucedo, R., Macías, J., Gavilanes, J., Arce, J., Komorowski, J., Gardner, J., and Valdez-Moreno, G. (2010). Eyewitness, stratigraphy, chemistry, and eruptive dynamics of the 1913 plinian eruption of volcán de colima, méxico. *Journal of Volcanology and Geothermal Research*, 191(3–4):149 – 166.
- Saucedo, R., Macías, J., Sheridan, M., Bursik, M., and Komorowski, J. (2005). Modeling of pyroclastic flows of colima volcano, mexico: implications for hazard assessment. *Journal of Volcanology and Geothermal Research*, 139(1):103 – 115. Modeling and Simulation of Geophysical Mass Flows.
- Saucedo, R., Macías, J. L., and Bursik, M. (2004). Pyroclastic flow deposits of the 1991 eruption of Volcán de Colima, Mexico. *Bulletin of Volcanology*, 66(4):291–306.
- Savage, S. B. and Hutter, K. (1989). The motion of a finite mass of granular material down a rough incline. *Journal of Fluid Mechanics*, 199:177.
- Savage, S. B. and Iverson, R. M. (2003). Surge dynamics coupled to pore-pressure evolution in debris flows. *Debris-Flow Hazards Mitigation: Mechanics, Prediction and Assessment*, edited by: Rickenmann, D. and Chen, C. L., Millpress, 1:503–514.
- Sheridan, M., Stinton, A., Patra, A., Pitman, E., Bauer, A., and Nichita, C. (2005). Evaluating titan2d mass-flow model using the 1963 little tahoma peak avalanches, mount rainier, washington. *Journal of Volcanology and Geothermal Research*, 139(1):89 – 102. Modeling and Simulation of Geophysical Mass Flows.
- Sheridan, M. F. and Macías, J. L. (1995). Estimation of risk probability for gravity-driven pyroclastic flows at volcan colima, mexico. *Journal of Volcanology and Geothermal Research*, 66(1):251 – 256. Models of Magnetic Processes and Volcanic Eruptions.

- Sheridan, M. F., Patra, A. K., Dalbey, K., and Hubbard, B.
- Silbert, L. E., Ertas, D., Grest, G. S., Halsey, T. C., Levine, D., and Plimpton, S. J. (2001). Granular flow down an inclined plane: Bagnold scaling and rheology. *Physical Review E*, 64(5):051302.
- Spiller, E. T., Bayarri, M. J., Berger, J. O., Calder, E. S., Patra, A. K., Pitman, E. B., and Wolpert, R. L. (2014). Automating emulator construction for geophysical hazard maps. *SIAM/ASA Journal on Uncertainty Quantification*, 2(1):126–152.
- Stefanescu, E. R., Bursik, M., Cordoba, G., Dalbey, K., Jones, M. D., Patra, A. K., Pieri, D. C., Pitman, E. B., and Sheridan, M. F. (2012a). Digital elevation model uncertainty and hazard analysis using a geophysical flow model. *Proceedings of the Royal Society of London A: Mathematical, Physical and Engineering Sciences*, 468(2142):1543–1563.
- Stefanescu, E. R., Bursik, M., and Patra, A. K. (2012b). Effect of digital elevation model on mohr-coulomb geophysical flow model output. *Natural Hazards*, 62(2):635–656.
- Stein, M. (1987). Large sample properties of simulations using latin hypercube sampling. *Technometrics*, 29(2):143–151.
- Sulpizio, R., Capra, L., Sarocchi, D., Saucedo, R., Gavilanes-Ruiz, J., and Varley, N. (2010). Predicting the block-and-ash flow inundation areas at volcán de colima (colima, mexico) based on the present day (february 2010) status. *Journal of Volcanology and Geothermal Research*, 193(1):49 – 66.
- Tang, B. (1993). Orthogonal array-based latin hypercubes. *Journal of the American Statistical Association*, 88(424):1392–1397.
- Toro, E. (2013). *Riemann Solvers and Numerical Methods for Fluid Dynamics: A Practical Introduction*. Springer Berlin Heidelberg.
- Voellmy, A. (1955). Über die Zerstörungskraft von Lawinen. *Schweiz Bauzeitung*, 73:159–165, 212–217, 246–249, 280–285.
- Webb, A. (2004). Granular flow experiments to validate numerical flow model, TITAN2D. Master's thesis, University at Buffalo.
- Webb, A. and Bursik, M. (2016). Granular flow experiments for validation of numerical flow models. Technical report, University at Buffalo.
- Wieland, J. M., Gray, J. M. N. T., and Hutter, K. (1999). Channelized free-surface flow of cohesionless granular avalanches in a chute with shallow lateral curvature. *Journal of Fluid Mechanics*, 392:73–100.
- Wolpert, R. L., Ogburn, S. E., and Calder, E. S. (2016). The longevity of lava dome eruptions. *Journal of Geophysical Research: Solid Earth*, 121(2):676–686. 2015JB012435.
- Yu, B., Dalbey, K., Webb, A., Bursik, M., Patra, A. K., Pitman, E. B., and Nichita, C. (2009). Numerical issues in computing inundation areas over natural terrains using Savage-Hutter theory. *Natural Hazards*, 50(2):249–267.
- Zobin, V., Luhr, J., Taran, Y., Bretón, M., Cortés, A., Cruz-Reyna, S. D. L., Domínguez, T., Galindo, I., Gavilanes, J., Muniz, J., Navarro, C., Ramírez, J., Reyes, G., Ursúa, M., Velasco, J., Alatorre, E., and Santiago, H. (2002). Overview of the 1997–2000 activity of volcán de colima, méxico. *Journal of Volcanology and Geothermal Research*, 117(1):1 – 19.
- Zobin, V. M., Arámbula, R., Bretón, M., Reyes, G., Plascencia, I., Navarro, C., Téllez, A., Campos, A., González, M., León, Z., Martínez, A., and Ramírez, C. (2015). Dynamics of the january 2013–june 2014 explosive-effusive episode in the eruption of volcán de colima, méxico: insights from seismic and video monitoring. *Bulletin of Volcanology*, 77(4):31.