

Seismic Inversion in the Bayesian Framework for Infinite-Dimensional Inverse Problems

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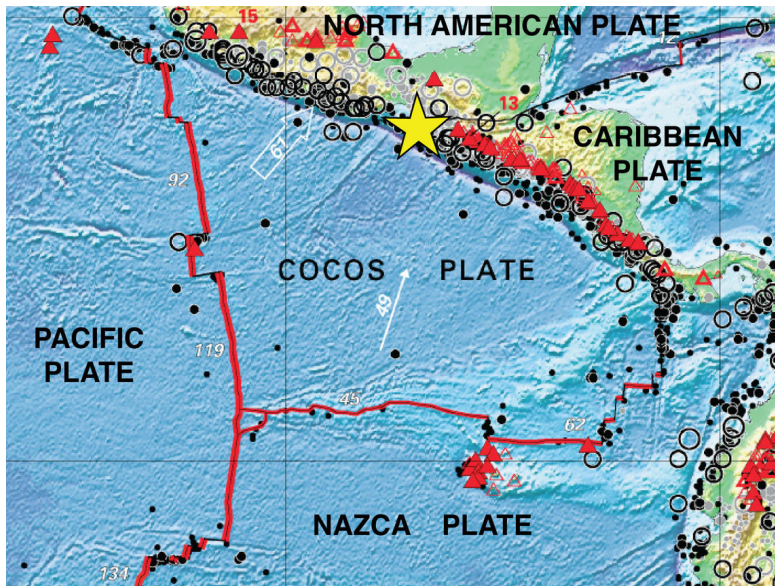


Figure: [Coc06]

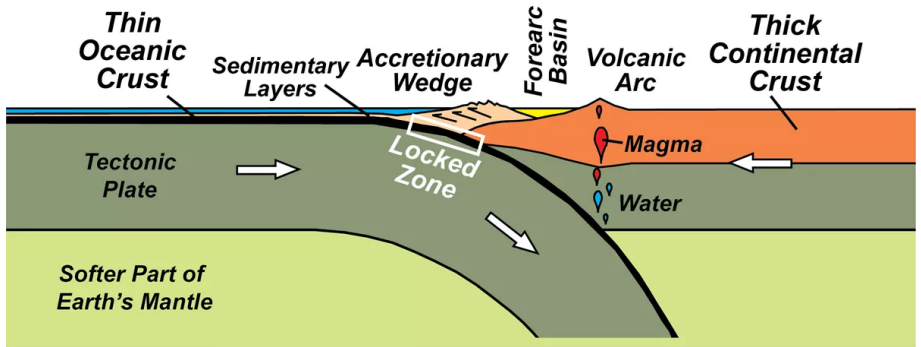


Figure: [Lil17]

Understand the subduction zone from collected observations.

Goals

Understand the subduction zone from collected observations.

Do so while quantifying measurement uncertainties.

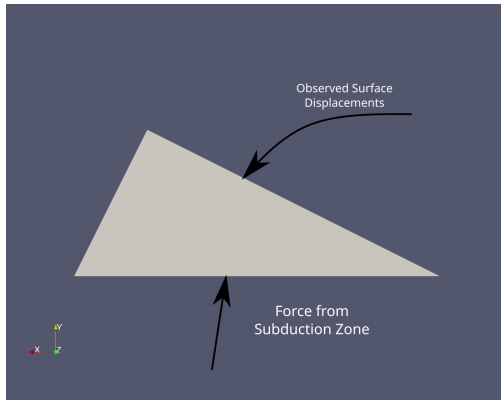
Table of Contents

- 1 Seismic Inversion Model Problem
- 2 Infinite-Dimensional Inverse Problem Setting
- 3 Numerical Discretization and Considerations
- 4 Results
- 5 Discussion and Future Work
- 6 References

Table of Contents

- 1 Seismic Inversion Model Problem
- 2 Infinite-Dimensional Inverse Problem Setting
- 3 Numerical Discretization and Considerations
- 4 Results
- 5 Discussion and Future Work
- 6 References

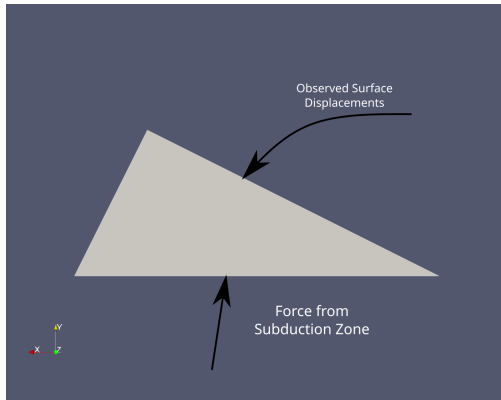
Model Assumptions



For the sake of modeling convenience, assume:

- **Governing PDE** (forward model): Linear elasticity

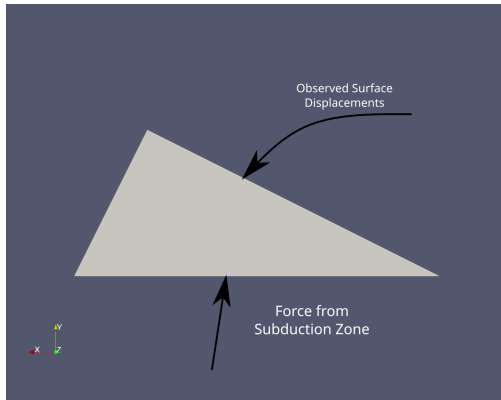
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Model Assumptions



For the sake of modeling convenience, assume:

- **Governing PDE** (forward model): Linear elasticity
- **Uncertain parameter**: Displacement on fault plane
- **Inverse Problem**: Given measurements of surface deformation \mathbf{u}^{obs} reconstruct fault plane displacement.

$$-\nabla \cdot \boldsymbol{\sigma}(\mathbf{u}) = \mathbf{0} \text{ in } \Omega, \quad (1)$$

where:

- $\boldsymbol{\sigma}(\mathbf{u}) = \mathbb{C}\boldsymbol{\varepsilon}(\mathbf{u})$ with
 - $\mathbb{C}[\boldsymbol{\varepsilon}] = 2\mu\boldsymbol{\varepsilon} + \lambda \text{tr}(\boldsymbol{\varepsilon})\mathbf{I}$ the fourth-order linear elasticity tensor:
 - $\boldsymbol{\varepsilon}(\mathbf{u}) = \frac{1}{2} [\nabla \mathbf{u} + (\nabla \mathbf{u})^T]$ the strain tensor.
- μ and λ are known as the Lamé constants.

$$-\nabla \left[\mu(\nabla \mathbf{u} + (\nabla \mathbf{u})^T) + \lambda \nabla \cdot \mathbf{u} \mathbf{I} \right] = \mathbf{0} \quad \text{in } \Omega, \quad (2a)$$

$$\boldsymbol{\sigma}(\mathbf{u})\mathbf{n} = \mathbf{0} \quad \text{on } \Gamma_t \quad (2b)$$

$$\mathbf{u} + \beta \boldsymbol{\sigma}(\mathbf{u})\mathbf{n} = \mathbf{h} \quad \text{on } \Gamma_s \quad (2c)$$

$$\mathbf{u} \cdot \mathbf{n} = 0 \quad \text{on } \Gamma_b \quad (2d)$$

$$\delta \mathbf{T}(\boldsymbol{\sigma}(\mathbf{u})\mathbf{n}) + \mathbf{T}\mathbf{u} = \mathbf{m} \quad \text{on } \Gamma_b \quad (2e)$$

- \mathbf{T} is the tangential operator $\mathbf{T}\mathbf{u} = (\mathbf{I} - \mathbf{n} \otimes \mathbf{n})\mathbf{u} = \mathbf{u} - (\mathbf{n}^T \mathbf{u})\mathbf{n}$.

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- The right hand side of (2e) is the displacement on the fault plane that is being inverted for.

Forward Model

$$-\nabla \left[\mu(\nabla \mathbf{u} + (\nabla \mathbf{u})^T) + \lambda \nabla \cdot \mathbf{u} \mathbf{I} \right] = \mathbf{0} \quad \text{in } \Omega, \quad (2a)$$

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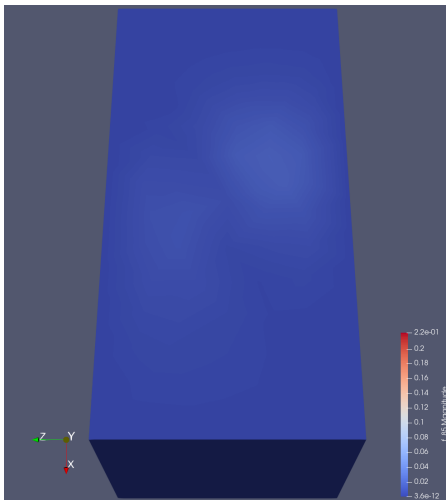
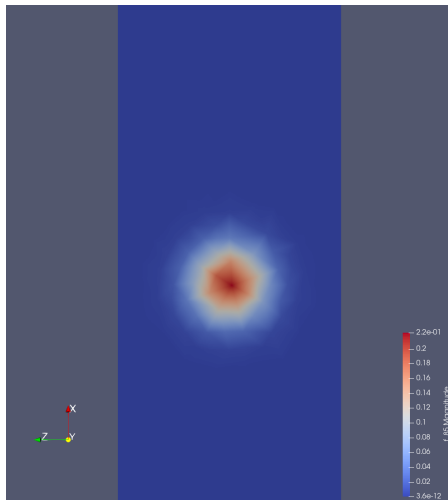
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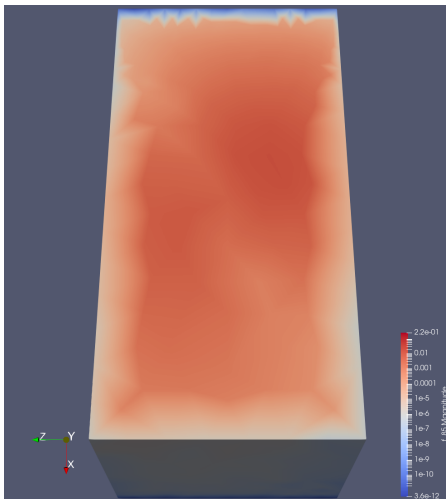
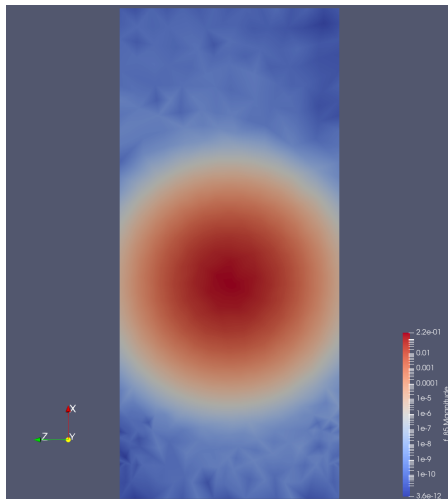
$$\delta \mathbf{T}(\boldsymbol{\sigma}(\mathbf{u})\mathbf{n}) + \mathbf{T}\mathbf{u} = \mathbf{m} \quad \text{on } \Gamma_b \quad (2e)$$

- \mathbf{T} is the tangential operator $\mathbf{T}\mathbf{u} = (\mathbf{I} - \mathbf{n} \otimes \mathbf{n})\mathbf{u} = \mathbf{u} - (\mathbf{n}^T \mathbf{u})\mathbf{n}$.
- The right hand side of (2e) is the displacement on the fault plane that is being inverted for.
- (2e) can be understood as a regularized Dirichlet condition.

Forward Solution



Forward Solution



Weak Formulation

Define:

$$\mathbf{V} := \{\mathbf{u} \in H^1(\Omega)^3 : \mathbf{u} \cdot \mathbf{n} = 0 \text{ on } \Gamma_b\}$$

Then the weak formulation of the forward model is given by:

$$\begin{aligned} \int_{\Gamma_s} \beta^{-1}(\mathbf{u} - \mathbf{h}) \cdot \mathbf{v} \, ds + \int_{\Gamma_b} \delta^{-1}(\mathbf{T}\mathbf{u} - \mathbf{m}) \cdot \mathbf{v} \, ds \\ + \int_{\Omega} \boldsymbol{\varepsilon}(\mathbf{u}) : \mathbb{C}[\boldsymbol{\varepsilon}(\mathbf{v})] \, d\mathbf{x} = 0, \quad \forall \mathbf{v} \in \mathbf{V} \quad (3) \end{aligned}$$

Table of Contents

- 1 Seismic Inversion Model Problem
- 2 Infinite-Dimensional Inverse Problem Setting**
- 3 Numerical Discretization and Considerations
- 4 Results
- 5 Discussion and Future Work
- 6 References

The Deterministic Inverse Problem

To reconstruct the fault displacement we construct the PDE-constrained optimization problem:

$$\mathcal{J}(\mathbf{m}) = \min_{\mathbf{m}} \frac{1}{2} \|\mathcal{B}\mathbf{u}(\mathbf{m}) - \mathbf{u}^{\text{obs}}\|^2 + \frac{1}{2} \|\mathcal{A}\mathbf{m}\|^2$$

where \mathbf{u} is given by the solution of the linear elasticity equation.

- $\mathbf{u}(\mathbf{m})$ is given by the forward model.
- $\mathcal{B} : (L^2(\Omega))^3 \rightarrow \mathbb{R}^N$ is an observation operator.
- $\mathbf{u}^{\text{obs}} \in \mathbb{R}^N$ where N is the number of data points.

¹In literature, called the *parameter-to-observable operator*.

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If we let \mathcal{S} be the forward model operator, i.e. $\mathbf{u} = \mathcal{S}\mathbf{m}$, and let¹ $\mathcal{F} = \mathcal{B}\mathcal{S}$, then:

$$\mathcal{J}(\mathbf{m}) = \min_{\mathbf{m}} \frac{1}{2} \|\mathcal{F}\mathbf{m} - \mathbf{u}^{\text{obs}}\|^2 + \frac{1}{2} \|\mathcal{A}\mathbf{m}\|^2$$

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- 6 References

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- 6 References

Table of Contents

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- 2 Infinite-Dimensional Inverse Problem Setting
- 3 Numerical Discretization and Considerations
- 4 Results
- 5 Discussion and Future Work**
- 6 References

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- 1 Seismic Inversion Model Problem
- 2 Infinite-Dimensional Inverse Problem Setting
- 3 Numerical Discretization and Considerations
- 4 Results
- 5 Discussion and Future Work
- 6 References**

References I



This dynamic planet: World map of volcanoes, earthquakes, impact craters and plate tectonics, Tech. report, 2006.



Robert J Lillie, *Oregon's Island In The Sky: Geology Road Guide to Marys Peak*, 2017 (English), OCLC: 979996650.