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#### Abstract

We describe a high-resolution approach for identifying microbial contributions to soil C-cycling dynamics using nucleic acid stable isotope probing coupled with next generation sequencing (HR-SIP). We amended series of soil microcosms with a complex mixture of model carbon (C) substrates and inorganic nutrients similar to plant biomass. A single C constituent in the C substrate mixture was substituted for its  $^{13}$ C-labeled equivalent in each microcosm series. Specifically, in separate microcosms we substituted <sup>13</sup>C-xylose or <sup>13</sup>C-cellulose for their unlabeled equivalents. Xylose and cellulose were chosen to represent labile soluble C and polymeric insoluble C, respectively. Microcosm DNA was interrogated for <sup>13</sup>C incorporation at days 1, 3, 7, 14 and 30. Incorporation of <sup>13</sup>C from xylose into microcosm DNA was observed at days 1, 3, and 7, while incorporation of <sup>13</sup>C from cellulose was peaked at day 14 and was maintained through day 40. Of over 6,000 OTUs detected, a total of XX and XX unique OTUs assimilated  $^{13}\mathrm{C}$  from xylose and cellulose into DNA, respectively. Xylose assimilating OTUs were more abundant in the microcosm community than cellulose assimilating OTUs, while cellulose OTUs demonstrated a greater substrate specificity than xylose OTUs. 13C-xylose incorporating OTUs included members of the Firmicutes, Bacteroidetes and *Actinobacteria*. <sup>13</sup>C-cellulose incorporating OTUs included members of the Verrucomicrobia and Chloroflexi.

monolayer | structure | x-ray reflectivity | molecular electronics

Abbreviations: SAM, self-assembled monolayer; OTS, octadecyltrichlorosilane

#### Introduction

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Whe have only a rudimentary understanding of carbon flow through soil microbial communities. This deficiency is driven by the staggering complexity of soil microbial food webs and the opacity of these biological systems to current methods for describing microbial metabolism in the environment. Relating community composition to overall soil processes, such as nitrification and denitrification, which are mediated by defined functional groups has been a useful approach. However, carbon-cycling processes have proven more recalcitrant to study due to the wide range of organisms participating in these reactions and our inability to discern diagnostic functional genetic markers.

Excluding plant biomass, there are 2,300 Pg of carbon (C) stored in soils worldwide which accounts for  $\sim$ 80% of the global terrestrial C pool (BATJES, 1996; Amundson, 2001). When organic C from plants reaches soil it is degraded by fungi, archaea, and bacteria. This C is rapidly returned to the atmosphere as CO<sub>2</sub> or remains in the soil as humic substances that can persist up to 2000 years (Yanagita, 1990). The majority of plant biomass C in soil is respired and produces 10 times more CO2 than anthropogenic emissions on an annual basis (Chapin, 2002). Global changes in atmospheric CO<sub>2</sub>, temperature, and ecosystem nitrogen inputs, are expected to impact primary production and C inputs to soils (Groenigen et al., 2006) but it remains difficult to predict the response of soil processes to anthropogenic change (DAVIDSON et al., 2006). Current climate change models concur on atmospheric and ocean C predictions but not terrestrial (Friedlingstein et al., 2006). These contrasting terrestrial ecosystem model predictions reflect how little is known about soil C cycling dynamics and it has been suggested that incosistencies in terrestial modeling could be improved by elucidating the relationship between dissolved organic carbon and microbial communities in soils (Neff and Asner, 2001).

An estimated 80-90% of C cycling in soil is mediated by microorganisms (Nannipieri et al., 2003a; n.d.). Understanding microbial processing of nutrients in soils presents a special challenge due to the hetergeneous nature of soil ecosystems and methods limitations. Soils are biologically, chemically, and physically complex which affects microbial community composition, diversity, and structure (Nannipieri et al., 2003a). Confounding factors such as physical protection/aggregation, moisture content, pH, temperature, frequency and type of land disturbance, soil history, mineralogy, N quality and availability, and litter quality have all been shown to affect the ability of the soil microbial community to access and metabolize C substrates (Sollins et al., 1996; Kalbitz et al., 2000). Further, rates of metabolism are often measured without knowing the identity of the microbial species involved (Nannipieri et al., 2003b) leaving the importance of community membership towards maintaining ecosystem functions unknown (Nannipieri et al., 2003b; Allison and Martiny, 2008; Schimel and Schaeffer, 2012). Litter bag experiments have shown that the community composition of soils can have quantitative and qualitative impacts on the breakdown of plant materials (Schimel, 1995). Reciprocal exchange of litter type and microbial inocula under controlled environmental conditions reveals that differences in community composition can account for 85% of the variation in litter carbon mineralization (Strickland et al., 2009). In addition, assembled communities of cellulose degraders reveal that the composition of the community has significant impacts on the rate of cellulose degradation (Wohl et al., 2004).

An important step in understanding soil C cycling dynamics is to identify individual contributions of discrete microorganisms and to investigate the relationship between genetic diversity, community structure, and function (O'Donnell *et al.*, 2002). The vast majority of microorganisms continue to resist cultivation in the laboratory, and even when cultivation is achieved, the traits expressed by a microorganism in culture may not be representative of those expressed when in its natural habitat. Stable-isotope probing (SIP) provides a unique opportunity to link microbial identity to activity and has been utilized to expand our knowledge of a myriad of important biogeochemical processes (Chen and Murrell, 2010). The most successful applications of this technique have identified organisms which mediate processes performed by a narrow set of

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PNAS | Issue Date | Volume | Issue Number | 1–12





functional guilds such as methanogens (Lu, 2005). The technique has been less applicable to the study of soil C cycling because of limitations in resolving power as a result of simultaneous labeling of many different organisms in the community. Additionally, molecular applications such as TRFLP, DGGE, and cloning that are frequently used in conjunction with SIP provide insufficient resolution of taxon identity and depth of coverage. We have developed an approach that employs a complex mixture of substrates added to soil at a low concentration relative to soil organic matter pools along with massively parallel DNA sequencing. This greatly expands the ability of nucleic acid SIP to explore complex patterns of C-cycling in microbial communities with increased resolution.

A temporal cascade occurs in natural microbial communities during the plant biomass degradation in which labile C degradation preceeds polymeric C (Hu and Bruggen, 1997; Rui et al., 2009). The aim of this study is to track the temporal dynamics of C assimilation through discrete individuals of the soil microbial community to provide greater insight into soil C-cycling. Our experimental approach employs the addition of a soil organic matter (SOM) simulant (a complex mixture of model carbon sources and inorganic nutrients common to plant biomass), where a single C constituent is substituted for its  $^{13}$ C-labeled equivalent, to soil. Parallel incubations of soils amended with this complex C mixture allows us to test how different C substrates cascade through discrete taxa within the soil microbial community. In this study we use <sup>13</sup>C-xylose and <sup>13</sup>C-cellulose as a proxy for labile and polymeric C, respectively. Using a novel approach we couple nucleic acid stable isotope probing with next generation sequencing (SIP-NGS) to elucidating soil microbial community members responsible for specific C transformations. Amplicon sequencing of 16S rRNA gene fragments from many gradient fractions and multiple gradients make it possible to track C assimilation by hundreds of different taxa. Ultimately we identify discrete microorganisms responsible for the cycling of specific C substrates.

### Results

To observe C use dynamics by the soil microbial community, we conducted a nucleic acid SIP experiment wherein xylose or cellulose carried the isotopic label, and, we assayed SSU rRNA gene content of CsCl gradient fractions using high-throughput DNA sequencing technology. We set up three soil microcosm series. Microcosms in each series were were amended with a C substrate mixture that included cellulose and xylose. The C substrate mixture approximated freshly degrading plant biomass. The same substrate mixture was added to microcosms in each series, however, for each series except the control, one substrate was substituted for its  $^{13}$ C counterpart. In one series cellulose was  $^{13}$ Clabeled in another xylose was <sup>13</sup>C-labeled and in the control series no substrates were <sup>13</sup>C labeled. Microcosm amendments are shorthand identified in the following figures by the following code: "13CXPS" refers to the amendment with 13C-xylose (that is <sup>13</sup>C Xlose Plant Simulant), "13CCPS" refers to the <sup>13</sup>C-cellulose amendment and "12CCPS" refers to the amendment that only contained <sup>12</sup>C substrates. Xylose or cellulose were chosen to carry the isotopic label to contrast C assimilation for labile, soluble C (xylose) versus insoluble, polymeric C (cellulose). 5.3 mg of C substrate mixture per gram soil was added to each microcosm representing 18% of the total soil C. The mixture included 0.42 mg xylose-C and 0.88 mg cellulose-C g soil<sup>-1</sup>. Microcosms were harvested at days 1, 3, 7, 14 and 30 during a 30 day incubation.  $^{13}$ C-xylose assimilation peaked immediately and tapered over the 30 day incubation whereas  $^{13}$ C-cellulose assimilation peaked at two weeks of (Figure 3).

We sequenced SSU rRNA gene amplicons from a total of 277 CsCl gradient fractions from 14 CsCl gradients and 12 bulk mi-

crocosm DNA samples. The SSU rRNA gene data set contained 1,376,008 total sequences. The average number of sequences per sample was 3,816 (sd 3,629) and 265 samples had over 1,000 sequences. We sequenced SSU rRNA gene amplicons from an average of 19.8 fractions per CsCl gradient (sd 0.57). The average density between fractions was 0.0040 g mL<sup>-1</sup> The sequencing effort recovered a total of 5,940 OTUs. 2,943 of the total 5,940 OTUs were observed in bulk samples. We observed 33 unique phylum and 340 unique genus annotations.

Soil microcosm microbial community changes with time. Changes in the soil microcosm microbial community structure and membership correlated with incubation time (Figure 9B, pvalue 0.23, R<sup>2</sup> 0.63, Adonis test Anderson (2001)). The <sup>1</sup> composition of the C-substrate addition did not significantly correlate with soil microcosm community structure and membership (p-value 0.35). Additionally, bulk sample beta diversity was significantly less than gradient fraction beta diversity (p-value 0.003, Anderson et al. (2006)). Twenty-nine OTUs significantly changed in relative abundance with time ("BH" adjusted p-value < 0.10, Y Benjamini (1995)). OTUs that significantly increased in relative abundance with time included OTUs in the Verrucomicrobia, Proteobacteria, Planctomycetes, Cyanobacteria, Chloroflexi and Acidobacteria. OTUs that significantly decreased in relative abundance with time included OTUs in the Proteobacteria, Firmicutes, Bacteroidetes and Actinobacteria (Figure XX). Proteobacteria was the only phylum that had OTUs that significantly increased and OTUs that significantly decreased in abundance with time. If sequences were grouped by taxonomic annotations at the class level, only four classes significantly changed in abundance, Bacilli (decreased), Flavobacteria (decreased), Gammaproteobacteria (decreased) and Herpetosiphonales (increased) (Figure XX). Of the 29 OTUs that changed significantly in relative abundance with time, 14 are labeled substrate responders (Figure XX).

OTUs that assimilated  $^{13}$ C from xylose. Isotope incorporation by an OTU is revealed by enrichment of the OTU in heavy CsCl gradient fractions containing  $^{13}$ C labeled DNA relative to heavy fractions from control gradients containing no  $^{13}$ C labeled DNA. We refer to OTUs that putatively incorporated  $^{13}$ C into DNA from an isotopically labeled substrate as a substrate "responder". Within the first 7 days of incubation 63% on average of  $^{13}$ C-xylose was respired and only an additional 6% more was respired from day 7 to 30. At the end of the 30 day incubation 30% of the  $^{13}$ C from added xylose remained in the soils. The  $^{13}$ C remaining in the soil from  $^{13}$ C-xylose addition was likely stabilized by assimilation into microbial biomass and/or microbial conversion into other forms of organic matter. It is also possible that some  $^{13}$ C-xylose remains unavailable to microbes due to abiotic interactions in soil (Kalbitz *et al.*, 2000).

At day 1, 84% of <sup>13</sup>C-xylose responsive OTUs belong to *Firmicutes*, 11% to *Proteobacteria* and 5% to *Bacteroidetes*. *Firmicutes* responders decreased to from 16 OTUs at day 1 to 1 OTU at day 3 while *Bacteroidetes* responders increased from 1 OTU at day 1 to 12 OTUs at day 3. The remaining day 3 responders are members of the *Proteobacteria* (26%) and the *Verrucomicrobia* (5%). Day 7 responders were 53% *Actinobacteria*, 40% *Proteobacteria*, and 7% *Firmicutes*. The identities of <sup>13</sup>C-xylose responders change with time. The numerically dominant <sup>13</sup>C-xylose responder phylum shifts from *Firmicutes* to *Bacteroidetes* and then to *Actinobacteria* across days 1, 3 and 7 (Figure 4, Figure 5).

All of the <sup>13</sup>C-xylose responders in the *Firmicutes* phylum are closely related (at least 99% sequence identity) to cultured isolates from genera that are known to form endospores (Table 2). Each <sup>13</sup>C-xylose responder is closely related to isolates annotated as members of *Bacillus*, *Paenibacillus* or *Lysinibacillus*. *Bacteroidetes* <sup>13</sup>C-xylose responders are predominantly





closely related to Flavobacterium species (5 of 8 total responders) (Table 2. Only one Bacteroidetes 13 C-xylose responder is not closely related to a cultured isolate, "OTU.183" (closest LTP BLAST hit, Chitinophaca sp., 89.5% sequence identity, Table 2). OTU.183 shares high sequence identity with environmental clones derived from rhizosphere samples (accession AM158371, unpublished) and the skin microbiome (accession JF219881, Kong et al. (2012)). Other Bacteroidetes responders share high sequence identities with canonical soil genera including Dyadobacer, Solibius and Terrimonas. Six of the 8 Actinobacteria  $^{13}$ C-xylose responders are in the Micrococcales order. One  $^{13}$ C-xylose responding Actinobacteria OTU shares 100% sequence identity with Agromyces ramosus (Table 2). A. ramosus is a known predatory bacterium but is not dependent on a host for growth in culture (Casida, n.d.). It is not possible to determine the specific origin of assimilated <sup>13</sup>C in a DNA-SIP experiment. <sup>13</sup>C can be passed down through trophic levels although heavy isotope representation in C pools targeted by cross-feeders and predators would be diluted with depth into the trophic cascade. It's possible, however, that the <sup>13</sup>C labeled Agromyces OTU was assimilating <sup>13</sup>C primarily by predation if the Agromyces OTU was selective enough with respect to its prey that it primarily attacked <sup>13</sup>C-xylose assimilating

Cellulose OTUs. Only 2 and 5 OTUs were found to have incorporated <sup>13</sup>C from <sup>13</sup>C-cellulose at days 3 and 7, respectively. At days 14 and 30, however, 42 and 39 OTUs were found to incorporate <sup>13</sup>C from <sup>13</sup>C-cellulose into biomass. An average 16% of the <sup>13</sup>Ccellulose added was respired within the first 7 days, 38% by day 14, and 60% by day 30. A Cellvibrio and Sandaracinaceae OTU assimilated <sup>13</sup>C from <sup>13</sup>C-cellulose at day 3. Day 7 <sup>13</sup>C-cellulose responders included the same Cellvibrio responder as day 3, a Verrucomicrobia OTU and three Chloroflexi OTUs. 50% of Day 14 responders belong to Proteobacteria (66% Alpha-, 19% Gamma-, and 14% Beta-) followed by 17% Planctomycetes, 14% Verrucomicrobia, 10% Chloroflexi, 7% Actinobacteria and 2% cyanobacteria. *Bacteroidetes* OTUs begin to incorporate <sup>13</sup>C from cellulose at day 30 (13% of day 30 responders). Other day 30 responding phyla include Proteobacteria (30% of day 30 responders; 42% Alpha-, 42% Delta, 8% Gamma-, and 8% Beta-), Planctomycetes (20%), Verrucomicrobia (20%), Chloroflexi (13%) and cyanobacteria (3%). Proteobacteria, Verrucomicrobia, and Chloroflexi had relatively high numbers of responders with strong response across multiple time points (Figure 4).

Proteobacteria represent 46% of all <sup>13</sup>C-cellulose responding OTUs identified. Cellvibrio accounted for 3% of all proteobacterial  $^{13}\mathrm{C}\text{-cellulose}$  responding OTUs detected.  $\mathit{Cellvibrio}$  was one of the first identified cellulose degrading bacteria and was originally described by Winogradsky in 1929 who named it for its cellulose degrading abilities (Boone, 2001). All 13 C-cellulose responding Proteobacteria share high sequence identity with 16S rRNA genes from sequenced cultured isolates (Table 1) except for "OTU.442" (best cultured isolate match 92% sequence identity in the Chrondomyces genus, Table 1) and "OTU.663" (best cultured isolate match outside Proteobacteria entirely, Clostridium genus, 89% sequence identity, Table 1). Some Proteobacteria responders share high sequence identity with type strains for genera known to possess cellulose degraders including Rhizobium. Devosia, Stenotrophomonas and Cellvibrio. One Proteobacteria OTU shares high sequence identity with a Brevundimonas cultured isolate. Brevundimonas has not previously been identified as a cellulose degrader, but has been shown to degrade cellouronic acid, an oxidized form of cellulose (Tavernier et al., 2008)

Verrucomicrobia, a cosmopolitan soil phylum often found in high abundance (Fierer et al., 2013), are hypothesized to degrade polysaccharides in many environments (Fierer et al., 2013;

Herlemann et al., 2013; Chin et al., n.d.). Verrucomicrobia comprise 16% of the total 13C-cellulose responder OTUs detected. 40% of Verrucomicrobia 13 C-cellulose responders belong to the uncultured "FukuN18" family originally identified in freshwater lakes (Parveen et al., 2013). The Verrucomicrobia OTU with the strongest Verrucomicrobial response to 13 C-cellulose shared high sequence identity (97%) with an isolate from Norway tundra soil (Jiang et al., 2011) although growth on cellulose was not assessed for this isolate. Only one other <sup>13</sup>C-cellulose responding verrucomicrobium shared high DNA sequence identity with a sequenced type strain, "OTU.638" (Table 1) with Roseimicrobium gellanilyticum (100% sequence identity). Roseimicrobium gellanilyticum grows on soluble cellulose (Otsuka et al., 2012). The remaining <sup>13</sup>C-cellulose *Verrucomicrobia* responders did not share high sequence identity with any cultured isolates (maximum sequence identity with any cultured isolate 93%).

Chloroflexi are traditionally known for metabolically dynamic lifestyles ranging from anoxygenic phototrophy to organohalide respiration (Hug et al., 2013). Recent studies have focused on Chloroflexi roles in C cycling (Goldfarb et al., 2011; Cole et al., 2013; Hug et al., 2013) and several Chloroflexi utilize cellulose (Goldfarb et al., 2011; Cole et al., 2013; Hug et al., 2013). Four closely related OTUs in an undescribed Chloroflexi lineage (closest matching cultured isolate for all four OTUs: Herpetosiphon geysericola, 89% sequence identity, Table 1) responded to <sup>13</sup>C-cellulose (Figure 6). One additional OTU also from a poorly characterized Chloroflexi lineage (closest cultured isolate match a proteobacterium at 78% sequence identity) responded to <sup>13</sup>C-cellulose (Figure 6).

Other notable 13C-cellulose responders include a Bacteroidetes OTU that shares high sequence identity (99%) to Sporocytophaga myxococcoides a known cellulose degrader (Vance et al., 1980), and three Actinobacteria OTUs that share high sequence identity (100%) with sequenced cultured isolates. One of the three Actinobacteria <sup>13</sup>C-cellulose responders is in the Streptomyces, a genus known to possess cellulose degraders, while the other two share high sequence identity to cultured isolates Allokutzneriz albata (Tomita et al., 1993; Labeda and Kroppenstedt, 2008) and Lentzea waywayandensis (LABEDA and LYONS, 1989; Labeda et al., 2001); neither isolate decomposes cellulose in culture. Nine *Planctomycetes* OTUs responded to <sup>13</sup>C-cellulose but none are within described genera (closest cultured isolate match 91% sequence identity, Table 1) (Figure 6). Interestingly, one <sup>13</sup>C-cellulose responder is annotated as "cyanobacteria". The cyanobacteria phylum annotation is misleading as the OTU is not closely related to any oxygenic phototrophs (closest cultured isolate match Vampirovibrio chlorellavorus, 95% sequence identity, Table 1). A sister clade to the oxygenic phototrophs classically annotated as "cyanobacteria" in SSU rRNA gene reference databases but does not possess any known phototrophs has recently been proposed to constitute its own phylum, "Melainabacteria" Rienzi et al. (2013). Although the phylogenetic position of "Melainabacteria" is debated (Soo et al., 2014). The catalog of metabolic capabilities associated with cyanobacteria (or candidate phyla previously annotated as cyanobacteria) are quickly expanding (Rienzi et al., 2013; Soo et al., 2014). Our findings provide evidence for cellulose degradation within a lineage closely related to but apart from oxygenic phototrophs. Notably, polysaccharide degradation is suggested by an analysis of a "Melainabacteria" genome (Rienzi et al., 2013). Although we highlight <sup>13</sup>C-cellulose responders that share high sequence identity with described genera, most <sup>13</sup>C-cellulose responders uncovered in this experiment are not closely related to cultured isolates (Table 1).

Xylose responders are more abundant members of the soil community than cellulose responders. <sup>13</sup>C-xylose responders are generally more abundant members based on relative abun-

Footline Author

PNAS | Issue Date | Volume | Issue Number |  $\bf 3$ 





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dance in bulk DNA SSU rRNA gene content than  $^{13}$ C-cellulose responders (Figure 7, p-value 0.00028). However, both abundant and rare OTUs responded to  $^{13}$ C-xylose and  $^{13}$ C-cellulose (Figure 7). For instance, a *Delftia*  $^{13}$ C-cellulose responder is fairly abundant in the bulk samples ("OTU.5", Table 1). OTU.5 was on average the 13th most abundant OTU in bulk samples. A  $^{13}$ C-xylose responder ("OTU.1040", Table 2) has a mean relative abundance in bulk samples of  $3.57x10^{-05}$ . Two  $^{13}$ C-cellulose responders were not found in any bulk samples ("OTU.862" and "OTU.1312", Table 1). Of the 10 most abundant responders 8 are  $^{13}$ C-xylose responders and 6 of these 8 are consistently among the 10 most abundant OTUs in bulk samples.

Responder abundances summed at phylum level generally increased for <sup>13</sup>C-cellulose (Figure XX) whereas <sup>13</sup>C-xylose responder abundances summed at the phylum level decreased over time for *Firmicutes*, *Bacteroidetes*, *Actinobacteria* and *Proteobacteria* although *Proteobacteria* spiked at day 14 (Figure 10). Bulk abundance trends are roughly consistent with <sup>13</sup>C assimilation

Cellulose degraders exhibit higher substrate specificity than xylose utilizers. Cellulose responders exhibited a greater shift in BD than xylose responders in response to isotope incorporation (Figure 7, p-value 1.8610x<sup>-06</sup>). <sup>13</sup>C-cellulose responders shifted on average 0.0163 g/mL (sd 0.0094) whereas xylose responders shifted on average 0.0097 (sd 0.0094). For reference, 100% <sup>13</sup>C DNA shifts X.XX g/mL relative to the BD of its  $^{12}$ C counterpart. DNA BD increases as its ratio of  $^{13}$ C to  $^{12}$ C increases. An organism that only assimilates C into DNA from a  $^{13}$ C isotopically labeled source, will have a greater  $^{13}$ C: $^{12}$ C ratio in its DNA than an organism utilizing a mixture of isotopically labeled and unlabeled C sources. Upon labeling, DNA from an organism that incorporates exclusively <sup>13</sup>C will increase in buoyant density more than DNA from an organism that does not exclusively utilize isotopically labeled C. Therefore the magnitude DNA buoyant density shifts indicate substrate specificity given our experimental design as only one substrate was labeled in each amendment. We measured density shift as the change in an OTU's density profile center of mass between corresponding control and labeled gradients. Density shifts, however, should not be evaluated on an individual OTU basis as a small number of density shifts are observed for each OTU and the variance of the density shift metric at the level of individual OTUs is unknown. It is therefore more informative to compare density shifts among substrate responder groups. Further, density shifts are based on relative abundance profiles and would be theoretically muted in comparison to density shifts based on absolute abundance profiles and should be interpreted with this transformation in mind. It should also be noted that there was overlap in observed density shifts between <sup>13</sup>C-cellulose and <sup>13</sup>C-xylose responder groups suggesting that although cellulose degraders are generally more substrate specific than xylose utilizers, some cellulose degraders show less substrate specificity for cellulose than some xylose utilizers for xylose (Figure 7), and, each responder group exhibits a range of substrate specificities (Figure 7).

Estimated rrn gene copy number in substrate responder groups. <sup>13</sup>C-xylose responder estimated rrn gene copy number is inversely related to time of first response (p-value  $2.02 \times 10^{-15}$ , Figure 8). OTUs that first respond at later time points have fewer estimated rrn copy number than OTUs that first respond earlier (Figure 8). rrn copy number estimation is a recent advance in microbiome science (Kembel et al., 2012) while the relationship of rrn copy number per genome with ecological strategy is well established (Klappenbach et al., 2000). Microorganisms with a high rrn copy number tend to be fast growers specialized to take advantage of boom-bust environments whereas microorganisms with low rrn copy number favor slower growth under lower and more consis-

tent nutrient input (Klappenbach *et al.*, 2000). At the beginning of our incubation, OTUs with estimated high *rrn* copy number or "fast-growers" assimilate xylose into biomass and with time slower growers (lower *rrn* copy number) begin to incorporate <sup>13</sup>C from xylose. Further, <sup>13</sup>C-xylose responders have more estimated rRNA operon copy numbers per genome than <sup>13</sup>C-cellulose responders (p-value 1.878x10<sup>-09</sup>) suggesting xylose respiring microbes are generally faster growers than cellulose degraders.

### Discussion

Nucleic-acid SIP coupled to microbiome fingerprinting techniques has progressed from simple proof-of-concept experiments (Radajewski et al., 2000) to studies utilizing non-DNA-sequencing microbial community profiling methods such as DGGE, ARISA and/or tRFLP (El Zahar Haichar et al., 2007), and currently to large experiments employing multiple labeled substrates and high-throughput amplicon and/or shotgun DNA sequencing (Verastegui et al., 2014). We present a high-resolution nucleic acid SIP (HR-SIP) approach that expands upon classical nucleic acid SIP methods in three dimensions: 1) temporally, we sample isotopically labeled substrate amended microcosms at multiple time points; 2) spatially, we assay more fractions along the CsCl gradients; and 3), bioinformatically, we interrogate taxa at the level of OTU (97% sequence identity) for isotope incorporation employing cutting edge statistics for assessing differential abundance in microbiome datasets.

Ordination of CsCl gradient fraction OTU profiles can be used to observe fraction-level  $^{13}\mathrm{C}$  assimilation dynamics and membership differences. Many techniques are utilized to identify  $^{13}$ C-DNA in DNA-SIP studies (reviewed by CITE) including the use of <sup>13</sup>C carrier DNA, visualization of CsCl gradient DNA with DNA intercalating dyes CITE, and screening of CsCl gradient fractions with microbial community fingerprinting techniques or qPCR CITE. Indeed the greatest challenge in DNA-SIP studies is teasing apart  $^{13}$ C DNA from high G+C  $^{12}$ C-DNA (Buckley *et al.*, 2007). High throughput DNA sequencing allows microbial ecologists to survey the microbial composition of thousands of samples CITE EMP. We leveraged the high throughput nature of next generation DNA sequencing technology to survey the microbial composition of entire CsCl gradients. Understanding the composition of entire gradients is essential for robustly identifying and disentangling microbes that have incorporated <sup>13</sup>C into DNA from those microbes with only <sup>2</sup>C-DNA.

Each CsCl gradient fraction possesses a unique composition of SSU rRNA gene phylogenetic types. DNA buoyant density (BD) drives differences in CsCl gradient fraction SSU rRNA gene composition (see Figure 3). For instance, lighter DNA is more abundant in fractions at lighter densities so DNA with lower G+C will be found in greater abundance at the light end of the CsCl gradient and vice versa. Duplicate gradients receiving entirely 12C DNA with the same bulk or non-fractionated SSU rRNA gene phylogenetic composition would have the same overall profile of SSU rRNA gene phylogenetic types across the density gradient. We fed microcosms identical C substrate mixtures save for the identity of a <sup>13</sup>C labeled substrate, and by design, DNA from all microcosms harvested at a time point will be similar in bulk phylogenetic composition. Therefore, SSU rRNA gene profile differences between gradients harvested at the same time are due to  $^{13}\mathrm{C}$  incorporation into bulk community DNA.  $^{13}\mathrm{C}\text{-DNA}$  shifts from its  $^{12}\mathrm{C}$  BD position towards the heavy end of the density gradient. This causes heavy fractions in gradients that received  $^{13}\mathrm{C}\text{-DNA}$  to be different in phylogenetic content than corresponding heavy fractions from gradients that received 12C-DNA of the same bulk phylogenetic composition.





Ordination of CsCl gradient fraction phylogenetic profiles reveals differences and similarities between gradients. It's clear that microcosms incorporated  $^{13}{\rm C}$  from both  $^{13}{\rm C}$ -xylose and  $^{13}{\rm C}$ cellulose as gradients from both 13C-xylose and 13C-cellulose microcosms differ from corresponding control gradients (Figure 3). These differences from control gradients are focused in the heavy fractions (Figure 3). Analysis of SSU rRNA gene surveys has greatly benefited from utilizing conventional methods for data exploration in ecology such as ordination (Lozupone and Knight, 2008). SSU rRNA gene phylogenetic profiles in CsCl gradient fractions have only recently been surveyed with high-throughput DNA sequencing technology and subsequently explored via ordination (Angel and Conrad, 2013; Verastegui et al., 2014). Ordination of CsCl gradient fraction phylogenetic profiles has reveled the relative influence of buoyant density and soil type on gradient phylogenetic profile variance, however, ordination has not demonstrated isotope incorporation. Demonstrating isotope incorporation requires careful comparisons between control and labeled gradients over the same buoyant density range. By sequencing CsCl gradient fractions from both control and labeled gradients across the full density gradient with DNA harvested from microcosms at multiple time points, we can observe where in the density gradient  $^{13}\mathrm{C}$  isotope incorporation signal is strongest and when  $^{13}\mathrm{C}$  isotope incorporation begins (Figure 3). <sup>13</sup>C incorporation from xylose and cellulose is most apparent at days 1/3/7 and days 14/30, respectively (Figure 3). Moreover, labeled gradient fraction phylogenetic profiles diverge from controls most dramatically at relatively heavy buoyant densities (Figure 3). Also, <sup>13</sup>C-DNA from <sup>13</sup>C-xylose microcosms is different in phylogenetic composition from <sup>13</sup>C-cellulose microcosm  $^{13}\mathrm{C\text{-}DNA}$  indicating that xylose and cellulose were assimilated by different microbial community members (Figure 3). Lastly, ordination indicates organisms that assimilated <sup>13</sup>C from <sup>13</sup>C-xylose changed in phylogenetic type over incubation days 1, 3 and 7 (Figure 3).

Cellulose degraders identified from undescribed lineages and cosmopolitan soil taxa for which functional attributes are not established. Verrucomicrobia are ubiquitous in soil worldwide (Bergmann et al., 2011). Verrucomicrobia can constitute 23% of 16S rRNA gene sequences in high-throughput DNA sequencing surveys of SSU rRNA genes in soil (Bergmann et al., 2011) and have been shown to represent as high as 9.8% of soil 16S rRNA (Buckley and Schmidt, 2001). Many Verrucomicrobia cultivars have been established in the last decade Wertz et al., 2011 but only one of the 15 most abundant verrucomicrobial phylotypes in a global soil sample collection shared greater than 93% sequence identity with an isolate (Bergmann et al., 2011). Genomic analyses and physiological profiling of Verrucomicrobia isolates have revealed methanotrophy and diazotrophy (Wertz et al., 2011) within Verrucomicria (CITE and reviewed by Wertz et al. (2011)). Notably, the genetic capacity to degrade cellulose and cellulose degradation in culture have been demonstrated in Verrucomicrobia (Wertz et al., 2011; Otsuka et al., 2012). Although, we have learned many functional roles of Verrucomicrobia in the environment, the function and or global significance of soil Verrucomicrobia in global C-cycling is unknown. For example, only one of the putative verrucomicrobial cellulose degraders identified in this experiment are closely related to named cultivars (OTU.XX, Table 1) and only XX% of all verrucomicrobial OTUs found in this study share at lease 97% sequence identity with isolates. Seven of 10  $^{13}$ C-cellulose responding verrucomicrobial OTUs were classified belonging to the Spartobacteria order. Spartobacteria order was overwhelmingly the numerically dominant order of Verrucomicrobia in SSU rRNA gene surveys of 181 globally distributed soil samples (Bergmann et al., 2011). HR-SIP identifies key players in soil C-cycling and Verrucomicrobia lineages particularly Spartobacteria, given their ubiquity and abundance in soil as well as their demonstrated incorporation of  $^{13}{\rm C}$  from  $^{13}{\rm C}\text{-cellulose},$  may be significant players in global soil cellulose respiration.

Soil *Chloroflexi* have been found to assimilate cellulose before in DNA-SIP studies with <sup>13</sup>C-cellulose (Schellenberger *et al.*, 2010). Previously identified *Chloroflexi* <sup>13</sup>C-cellulose-responders in were similarly underrepresented in culture collections as the *Chloroflexi* identified as cellulose degraders in this study (Schellenberger *et al.*, 2010). *Chloroflexi* is on average XX% abundant in soil samples screened by the Earth Microbiome Project (EMP, CITE, Figure XX) and is found in XX of XX EMP soil samples (XX%) and XX of all XX EMP samples (Figure XX). Although not as abundant as *Spartobacteria* in EMP samples (Figure XX), *Chloroflexi* are still a cosmopolitan soil phylum. Similarly to *Spartobacteria*, *Chloroflexi* are largely uncharacterized functionally in soil. Here we present a lineage in the *Chloroflexi* that incorporated <sup>13</sup>C from cellulose in our microcosms. *Chloroflexi* Banfield paper

Importance of identifying lineages participating in SOM decomposition. SOM is a major global C reservoir and the storage and flux of SOM is largely mediated by microorganisms. Moreover, SOM decomposition is more sensitive to temperature changes than primary productivity CITE. Climate change can affect terrestrial microbial communities on a global scale CITE Garcia-Pichel. Therefore, climate change has the potential to significantly influence SOM flux and storage. Predicting climate how climate change will affect SOM flux and storage will require experiments that investigate how temperature affects decomposition of various quality CITE in addition to understanding how climate change alters the abundance and activity of key microbial players in decomposition. Before we can assess how climate change will impact SOM decomposition by observing temperature influences on soil microbial composition, we must first identify which microbial phylogenetic types participate in the decomposition of different SOM C components. Here we demonstrate heretofore undefined functional roles of cosmopolitan soil microbial phylotypes.

The succession hypothesis of decomposition hypothesizes a succession from microbial types that utilize labile C to those that utilize recalcitrant polymeric C over time CITE. We observed a *sub*-succession during the degradation of labile C. Soluble C more robust to temperature changes because or redundancy? But we found similar numbers of xylose and cellulose degrader

Ecological strategies of soil microorganisms participating in the decomposition of organic matter. Ecological strategies of soil microorganisms have mostly been elucidated by cultivationdependent studies CITE and include characterization of growth rates, ... We observed several ecological strategies among substrate responders although as boundaries between groups were not well defined the observed ecological strategy space appeared more like a continuum than discrete points. In general, 13Ccellulose responders were slow-growers with low rrn gene copy number. 13C-cellulose responders were also typically low abundance members of the microbial community and exhibited higher substrate specificity than <sup>13</sup>C-xylose responders (Figure 7). Within the <sup>13</sup>C-cellulose responders there was faint but apparent subecological strategy groups. For instance, Verrucomicrobia, Chloroflexi and Planctomycetes showed had distinct patterns of substrate specificity, rrn gene copy number and temporal dynamics in the bulk community (Figure XX).

We found that control gradients yield amplifiable DNA well into the heavy fractions. In fact, we observe XX OTUs in the heavy fractions of control gradients. Therefore, presence in heavy fractions of  $^{13}\mathrm{C}$  gradients alone is not evidence of  $^{13}\mathrm{C}$  incorporation into DNA. Only enrichment in  $^{13}\mathrm{C}$  heavy fractions relative to control gradient fractions demonstrates incorporation even if heavy fractions are shown to be different from controls by fingerprinting techniques CITE Neufeld 2010.

PNAS | Issue Date | Volume | Issue Number | 5



Footline Author



Xylose responders change over time. Implications for DNA-SIP. Succession within succession.

Response not consistent across phyla.

Ecological strategies of xylose and cellulose degrading OTUs in soil. HR-SIP allows us to assess C substrate specificity and temporal dynamics of <sup>13</sup>C-incorporation. C substrate specificity is assessed by measuring the BD shift of OTU DNA upon <sup>13</sup>C incorporation. OTUs that incorporate more <sup>13</sup>C per unit DNA have greater specify for the labeled substrate than OTUs that incorporate less  $^{13}\mathrm{C}$  per unit DNA.  $^{13}\mathrm{C}$ -cellulose incorporating OTUs as a group displayed greater substrate specificity than <sup>13</sup>C-xylose incorporating OTUs. This suggests that polymeric C-degraders tend to be specialists tuned to particular C-substrates such as cellulose or lignin whereas labile C-degraders are generalists able to assimilate C from many different labile sources. Although we observed a succession of <sup>13</sup>C-xylose responders (Figure 4 and 5), there was no discernible difference in substrate specificity between <sup>13</sup>C-xylose responders that first responded at days 1, 3 or 7. There was, however, a strong positive relationship between  $^{13}\mbox{C-xylose}$  responders that first responded at days 1, 3 or 7 in rrn copy number. <sup>13</sup>C-xylose responders that first responded at day 1 had higher estimated rrn copy number than responders that first responded at day 3 which had higher rrn copy number than responders that first responded at day 7. Therefore, OTUs that grow faster assimilate C from xylose faster intuitively. However, fast growers are replaced with respect to xylose C assimilation with slower growers as xylose diminishes. There was a succession of xylose degraders with time from fast growing spore-formers to Bacteroidetes types and finally Actinobacteria in our microcosms. The succession hypothesis of decomposition groups ecological units by substrate CITE, however, our results suggest there is a succession of microbial activity for even a single substrate. Hence, in soil, there is an ecological hierarchy coarsely defined by the ability to assimilate C from labile or polymeric sources but within the labile substrate degraders there are ecological subunits tuned to specific substrate concentrations.

#### Methods

Additional information on sample collection and analytical methods is provided in SI Materials and Methods.

Twelve soil cores (5 cm diameter x 10 cm depth) were collected from six random sampling locations within an organically managed agricultural field in Penn Yan, New York. Soils were pretreated by sieving (2 mm), homogenizing sieved soil, and preincubating 10 g of dry soil weight in flasks for 2 weeks. Soils were amended with a 5.3 mg g soil-1 carbon mixture; representative of natural concentrations Schneckenberger et al., 2008. Mixture contained 38% cellulose, 23% lignin, 20% xylose, 3% arabinose, 1% galactose, 1% glucose, and 0.5% mannose by mass, with the remaining 13.5% mass composed of an amino acid (in-house made replica of Teknova C0705) and basal salt mixture (Murashige and Skoog, Sigma Aldrich M5524). Three parallel treatments were performed; (1) unlabeled control,(2)<sup>13</sup>C-cellulose, (3)<sup>13</sup>C-xylose (98 atom% <sup>13</sup>C, Sigma Aldrich). Each treatment had 2 replicates per time point (n = 4) except day 30 which had 4 replicates; total microcosms per treatment n = 12, except  $^{13}$ C-cellulose which was not sampled at day 1, n = 10. Other details relating to substrate addition can be found in SI. Microcosms were sampled destructively (stored at -80°C until nucleic acid processing) at days 1 (control and xylose only), 3, 7, 14, and 30.

Nucleic acids were extracted using a modified Griffiths procotol Griffiths *et al.*, 2000. To prepare nucleic acid extracts for isopycnic centrifugation as previously described Buckley *et al.*, 2007, DNA was size selected (>4kb) using 1% low melt agarose gel and

 $\beta$ -agarase I enzyme extraction per manufacturers protocol (New England Biolab, M0392S). For each time point in the series isopycnic gradients were setup using a modified protocol Neufeld *et al.*, 2007 for a total of five <sup>12</sup>C-control, five <sup>13</sup>C-xylose, and four <sup>13</sup>Ccellulose microcosms. A density gradient (average density 1.69 g mL<sup>-1</sup>) solution of 1.762 g cesium chloride (CsCl) ml<sup>-1</sup> in gradient buffer solution (pH 8.0 15 mM Tris-HCl, 15 mM EDTA, 15 mM KCl) was used to separate <sup>13</sup>C-enriched and <sup>12</sup>C-nonenriched DNA. Each gradient was loaded with approximately 5  $\mu g$  of DNA and ultracentrifuged for 66 h at 55,000 rpm and room temperature (RT). Fractions of  $\sim$ 100  $\mu$ L were collected from below by displacing the DNA-CsCl-gradient buffer solution in the centrifugation tube with water using a syringe pump at a flow rate of 3.3  $\mu$ L s<sup>-1</sup> Manefield et al., 2002 into Acroprep™ 96 filter plate (Pall Life Sciences 5035). The refractive index of each fraction was measured using a Reichart AR200 digital refractometer modified as previously described Buckley et al., 2007 to measure a volume of 5  $\mu$ L. Then buoyant density was calculated from the refractive index as previously described Buckley et al., 2007 (see also SI). The collected DNA fractions were purified by repetitive washing of Acroprep filter wells with TE. Finally, 50  $\mu$ L TE was added to each fraction then resuspended DNA was pipetted off the filter into a new microfuge

For every gradient, 20 fractions were chosen for sequencing between the density range 1.67-1.75 g mL<sup>-1</sup>. Barcoded 454 primers were designed using 454-specific adapter B, 10 bp barcodes Hamady et al., 2008, a 2 bp linker (5'-CA-3'), and 806R primer for reverse primer (BA806R); and 454-specific adapter A, a 2 bp linker (5'-TC-3'), and 515F primer for forward primer (BA515F). Each fraction was PCR amplified using 0.25  $\mu$ L 5 U  $\mu$ I AmpliTaq Gold (Life Technologies, Grand Island, NY; N8080243), 2.5  $\mu$ L 10X Buffer II (100 mM Tris-HCl, pH 8.3, 500 mM KCl), 2.5  $\mu$ L 25 mM MgCl<sub>2</sub>, 4  $\mu$ L 5 mM dNTP, 1.25  $\mu$ L 10 mg mL<sup>-1</sup> BSA, 0.5  $\mu$ L 10  $\mu$ M BA515F, 1  $\mu$ L 5  $\mu$ M BA806R, 3  $\mu$ L H<sub>2</sub>O, 10  $\mu$ L 1:30 DNA template) in triplicate. Samples were normalized either using Pico green quantification and manual calculation or by SequalPrep<sup>™</sup> normalization plates (Invitrogen, Carlsbad, CA; A10510), then pooled in equimolar concentrations. Pooled DNA was gel extracted from a 1% agarose gel using Wizard SV gel and PCR clean-up system (Promega, Madison, WI; A9281) per manufacturer's protocol. Amplicons were sequenced on Roche 454 FLX system using titanium chemistry at Selah Genomics (formerly EnGenCore, Columbia, SC)

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PNAS | Issue Date | Volume | Issue Number | 7





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## **Figures**

Table 1:  $^{13}$ C-cellulose responders BLAST against Living Tree Project

OTU ID	Fold change	Top BLAST hits	BLAST %ID	Phylum;Class;Order
OTU.862	5.87	Allokutzneria albata	100.0	Actinobacteria Pseudonocardiales Pseudonocardiaceae
OTU.257	2.94	Lentzea waywayandensis, Lentzea flaviverrucosa	100.0	Actinobacteria Pseudonocardiales Pseudonocardiaceae
OTU.132	2.81	Streptomyces spp.	100.0	Actinobacteria Streptomycetales Streptomycetaceae
OTU.465	3.79	Ohtaekwangia kribbensis	92.73	Bacteroidetes Cytophagia Cytophagales
OTU.1094	3.69	Sporocytophaga myxococcoides	99.55	Bacteroidetes Cytophagia Cytophagales
OTU.669	3.34	Ohtaekwangia koreensis	92.69	Bacteroidetes Cytophagia Cytophagales
OTU.573	3.03	Adhaeribacter aerophilus	92.76	Bacteroidetes Cytophagia Cytophagales
OTU.670	2.87	Adhaeribacter aerophilus	91.78	Bacteroidetes Cytophagia Cytophagales
OTU.971	3.68	No hits of at least 90% identity	78.57	Chloroflexi Anaerolineae Anaerolineales
OTU.64	4.31	No hits of at least 90% identity	89.5	Chloroflexi Herpetosiphonales Herpetosiphonaceae
OTU.4322	4.19	No hits of at least 90% identity	89.14	Chloroflexi Herpetosiphonales Herpetosiphonaceae
OTU.98	3.68	No hits of at least 90% identity	88.18	Chloroflexi Herpetosiphonales Herpetosiphonaceae
OTU.5190	3.6	No hits of at least 90% identity	88.13	Chloroflexi Herpetosiphonales Herpetosiphonaceae
OTU.120	4.76	Vampirovibrio chlorellavorus	94.52	Cyanobacteria SM1D11 uncultured-bacterium
OTU.1065	5.31	No hits of at least 90% identity	84.55	Planctomycetes Planctomycetacia Planctomycetales
OTU.484	4.92	No hits of at least 90% identity	89.09	Planctomycetes Planctomycetacia Planctomycetales
OTU.1204	4.32	Planctomyces limnophilus	91.78	Planctomycetes Planctomycetacia Planctomycetales
OTU.150	4.06	No hits of at least 90% identity	86.76	Planctomycetes Planctomycetacia Planctomycetales
OTU.663	3.63	Pirellula staleyi DSM 6068	90.87	Planctomycetes Planctomycetacia Planctomycetales
OTU.473	3.58	Pirellula staleyi DSM 6068	90.91	Planctomycetes Planctomycetacia Planctomycetales
OTU.285	3.55	Blastopirellula marina	90.87	Planctomycetes Planctomycetacia Planctomycetales
OTU.351	3.54	Pirellula staleyi DSM 6068	91.86	Planctomycetes Planctomycetacia Planctomycetales
OTU.600	3.48	No hits of at least 90% identity	80.37	Planctomycetes Planctomycetacia Planctomycetales
OTU.900	4.87	Brevundimonas vesicularis, Brevundimonas nasdae	100.0	Proteobacteria Alphaproteobacteria Caulobacterales
OTU.1754	4.48	Asticcacaulis biprosthecium, Asticcacaulis benevestitus	96.8	Proteobacteria Alphaproteobacteria Caulobacterales
OTU.119	3.31	Brevundimonas alba	100.0	Proteobacteria Alphaproteobacteria Caulobacterales
OTU.327	2.99	Asticcacaulis biprosthecium, Asticcacaulis benevestitus	98.63	Proteobacteria Alphaproteobacteria Caulobacterales
OTU.982	4.47	Devosia neptuniae	100.0	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.1087	4.32	Devosia soli, Devosia crocina, Devosia riboflavina	99.09	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.5539	4.01	Devosia subaequoris	98.17	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.3775	3.88	Devosia glacialis, Devosia chinhatensis, Devosia geojensis, Devosia yakushimensis	98.63	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.429	3.7	Devosia limi, Devosia psychrophila	97.72	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.766	3.21	Devosia insulae	99.54	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.165	3.1	Rhizobium spp.	100.0	Proteobacteria Alphaproteobacteria Rhizobiales







Table 1 – continued from previous page

OTU ID	Fold change	Top BLAST hits	BLAST %ID	Phylum;Class;Order
OTU.28	2.59	Rhizobium giardinii, Rhizobium tubonense, Rhizobium tibeticum, Rhizobium mesoamericanum CCGE 501 Rhizobium herbae, Rhizobium endophyticum	99.54	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.19	2.44	Rhizobium spp., Arthrobacter spp.	99.54	Proteobacteria Alphaproteobacteria Rhizobiales
OTU.90	2.94	Sphingopyxis panaciterrae, Sphingopyxis chilensis, Sphingopyxis sp. BZ30, Sphingomonas sp.	100.0	Proteobacteria Alphaproteobacteria Sphingomonadales
OTU.518	4.8	Hydrogenophaga intermedia	100.0	Proteobacteria Betaproteobacteria Burkholderiales
OTU.1312	4.07	Paucimonas lemoignei	99.54	Proteobacteria Betaproteobacteria Burkholderiales
OTU.5	3.69	Delftia tsuruhatensis, Delftia lacustris	100.0	Proteobacteria Betaproteobacteria Burkholderiales
OTU.114	2.78	Herbaspirillum sp. SUEMI03, Herbaspirillum sp. SUEMI10, Oxalicibacterium solurbis, Herminiimonas fonticola, Oxalicibacterium horti	100.0	Proteobacteria Betaproteobacteria Burkholderiales
OTU.633	3.84	No hits of at least 90% identity	89.5	Proteobacteria Deltaproteobacteria Myxococcales
OTU.3594	3.83	Chondromyces robustus	90.41	Proteobacteria Deltaproteobacteria Myxococcales
OTU.442	3.05	Chondromyces robustus	92.24	Proteobacteria Deltaproteobacteria Myxococcales
OTU.32	3.0	Sandaracinus amylolyticus	94.98	Proteobacteria Deltaproteobacteria Myxococcales
OTU.228	2.54	Sorangium cellulosum	98.17	Proteobacteria Deltaproteobacteria Myxococcales
OTU.899	2.28	Enhygromyxa salina	97.72	Proteobacteria Deltaproteobacteria Myxococcales
OTU.6	3.62	Cellvibrio fulvus	100.0	Proteobacteria Gammaproteobacteria Pseudomonadales
OTU.11	5.25	Stenotrophomonas pavanii, Stenotrophomonas maltophilia, Pseudomonas geniculata	99.54	Proteobacteria Gammaproteobacteria Xanthomonadales
OTU.6062	4.83	Dokdonella sp. DC-3, Luteibacter rhizovicinus	97.26	Proteobacteria Gammaproteobacteria Xanthomonadales
OTU.154	3.24	Pseudoxanthomonas mexicana, Pseudoxanthomonas japonensis	100.0	Proteobacteria Gammaproteobacteria Xanthomonadales
OTU.100	2.66	Pseudoxanthomonas sacheonensis, Pseudoxanthomonas dokdonensis	100.0	Proteobacteria Gammaproteobacteria Xanthomonadales
OTU.1023	4.61	No hits of at least 90% identity	80.54	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.266	4.54	No hits of at least 90% identity	83.64	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.541	4.49	No hits of at least 90% identity	84.23	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.185	4.37	No hits of at least 90% identity	85.14	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.2192	3.49	No hits of at least 90% identity	83.56	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.1533	3.43	No hits of at least 90% identity	82.27	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.241	3.38	No hits of at least 90% identity	87.73	Verrucomicrobia Spartobacteria Chthoniobacterales
OTU.83	5.61	Luteolibacter sp. CCTCC AB 2010415	97.72	Verrucomicrobia Verrucomicrobiae Verrucomicrobiales
OTU.627	4.43	Verrucomicrobiaceae bacterium DC2a-G	7 100.0	Verrucomicrobia Verrucomicrobiae Verrucomicrobiales
OTU.638	4.0	Luteolibacter sp. CCTCC AB 2010415, Luteolibacter algae	93.61	Verrucomicrobia Verrucomicrobiae Verrucomicrobiales









Table 2:  $^{13}\mathrm{C}\text{-xylose}$  responders BLAST against Living Tree Project

OTU ID	Fold change	Top BLAST hits	<b>BLAST %ID</b>	Phylum;Class;Order
OTU.4446	3.49	Catenuloplanes niger, Catenuloplanes castaneus, Catenuloplanes atrovinosus, Catenuloplanes crispus, Catenuloplanes nepalensis, Catenuloplanes japonicus	97.72	Actinobacteria Frankiales Nakamurellaceae
OTU.62	2.57	Nakamurella flavida	100.0	Actinobacteria Frankiales Nakamurellaceae
OTU.24	2.81	Cellulomonas aerilata, Cellulomonas humilata, Cellulomonas terrae, Cellulomonas soli, Cellulomonas xylanilytica	100.0	Actinobacteria Micrococcales Cellulomonadaceae
OTU.4	2.84	Agromyces ramosus	100.0	Actinobacteria Micrococcales Microbacteriaceae
OTU.37	2.68	Phycicola gilvus, Microterricola viridarii, Frigoribacterium faeni, Frondihabitans sp. RS-15, Frondihabitans australicus	100.0	Actinobacteria Micrococcales Microbacteriaceae
OTU.5284	3.56	Isoptericola nanjingensis, Isoptericola hypogeus, Isoptericola variabilis	98.63	Actinobacteria Micrococcales Promicromonosporaceae
OTU.252	3.34	Promicromonospora thailandica	100.0	Actinobacteria Micrococcales Promicromonosporaceae
OTU.244	3.08	Cellulosimicrobium funkei, Cellulosimicrobium terreum	100.0	Actinobacteria Micrococcales Promicromonosporaceae
OTU.760	2.89	Dyadobacter hamtensis	98.63	Bacteroidetes Cytophagia Cytophagales
OTU.14	3.92	Flavobacterium oncorhynchi, Flavobacterium glycines, Flavobacterium succinicans	99.09	Bacteroidetes Flavobacteria Flavobacteriales
OTU.6203	3.32	Flavobacterium granuli, Flavobacterium glaciei	100.0	Bacteroidetes Flavobacteria Flavobacteriales
OTU.159	3.16	Flavobacterium hibernum	98.17	Bacteroidetes Flavobacteria Flavobacteriales
OTU.2379	3.1	Flavobacterium pectinovorum, Flavobacterium sp. CS100	97.72	Bacteroidetes Flavobacteria Flavobacteriales
OTU.131	3.07	Flavobacterium fluvii, Flavobacteria bacterium HMD1033, Flavobacterium sp. HMD1001	100.0	Bacteroidetes Flavobacteria Flavobacteriales
OTU.3540	2.52	Flavobacterium terrigena	99.54	Bacteroidetes Flavobacteria Flavobacteriales
OTU.107	2.25	Flavobacterium sp. 15C3, Flavobacterium banpakuense	99.54	Bacteroidetes Flavobacteria Flavobacteriales
OTU.277	3.52	Solibius ginsengiterrae	95.43	Bacteroidetes Sphingobacteriia Sphingobacteriales
OTU.183	3.31	No hits of at least 90% identity	89.5	Bacteroidetes Sphingobacteriia Sphingobacteriales
OTU.5906	3.16	Terrimonas sp. M-8	96.8	Bacteroidetes Sphingobacteriia Sphingobacteriales
OTU.360	2.98	Flavisolibacter ginsengisoli	95.0	Bacteroidetes Sphingobacteriia Sphingobacteriales
OTU.369	5.05	Paenibacillus sp. D75, Paenibacillus glycanilyticus	100.0	Firmicutes Bacilli Bacillales
OTU.267	4.97	Paenibacillus pabuli, Paenibacillus tundrae, Paenibacillus taichungensis, Paenibacillus xylanexedens, Paenibacillus xylanilyticus	100.0	Firmicutes Bacilli Bacillales
OTU.1040	4.78	Paenibacillus daejeonensis	100.0	Firmicutes Bacilli Bacillales

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PNAS | Issue Date | Volume | Issue Number | 11







Table 2 – continued from previous page

OTU ID	Fold change	Top BLAST hits	BLAST %ID	Phylum;Class;Order
OTU.57	4.39	Paenibacillus castaneae	98.62	Firmicutes Bacilli Bacillales
OTU.394	4.06	Paenibacillus pocheonensis	100.0	Firmicutes Bacilli Bacillales
OTU.319	3.98	Paenibacillus xinjiangensis	97.25	Firmicutes Bacilli Bacillales
OTU.5603	3.96	Paenibacillus uliginis	100.0	Firmicutes Bacilli Bacillales
OTU.1069	3.85	Paenibacillus terrigena	100.0	Firmicutes Bacilli Bacillales
OTU.843	3.62	Paenibacillus agarexedens	100.0	Firmicutes Bacilli Bacillales
OTU.2040	2.91	Paenibacillus pectinilyticus	100.0	Firmicutes Bacilli Bacillales
OTU.3	2.61	[Brevibacterium] frigoritolerans, Bacillus sp. LMG 20238, Bacillus coahuilensis m4-4, Bacillus simplex	100.0	Firmicutes Bacilli Bacillales
OTU.335	2.53	Paenibacillus thailandensis	98.17	Firmicutes Bacilli Bacillales
OTU.3507	2.36	Bacillus spp.	98.63	Firmicutes Bacilli Bacillales
OTU.8	2.26	Bacillus niacini	100.0	Firmicutes Bacilli Bacillales
OTU.4743	2.24	Lysinibacillus fusiformis, Lysinibacillus sphaericus	99.09	Firmicutes Bacilli Bacillales
OTU.9	2.04	Bacillus megaterium, Bacillus flexus	100.0	Firmicutes Bacilli Bacillales
OTU.22	2.8	Paracoccus sp. NB88	99.09	Proteobacteria Alphaproteobacteria Rhodobacterales
OTU.346	3.44	Pseudoduganella violaceinigra	99.54	Proteobacteria Betaproteobacteria Burkholderiales
OTU.68	3.74	Shigella flexneri, Escherichia fergusonii, Escherichia coli, Shigella sonnei	100.0	Proteobacteria Gammaproteobacteria Enterobacteriales
OTU.290	3.59	Pantoea spp., Kluyvera spp., Klebsiella spp., Erwinia spp., Enterobacter spp., Buttiauxella spp.	100.0	Proteobacteria Gammaproteobacteria Enterobacteriales
OTU.48	2.99	Aeromonas spp.	100.0	Proteobacteria Gammaproteobacteria aaa34a10









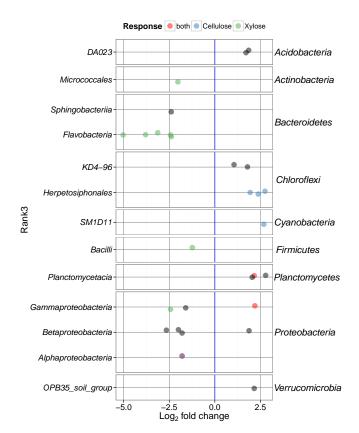


Fig. 1. Fold change time  $^{-1}$  for OTUs that changed significantly in abundance over time. One panel per phylum (phyla indicated on the right). Taxonomic class indicated on the left.

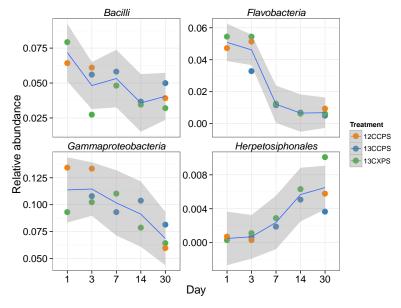


Fig. 2. Relative abundance versus day for classes that changed significantly in relative abundance with time.





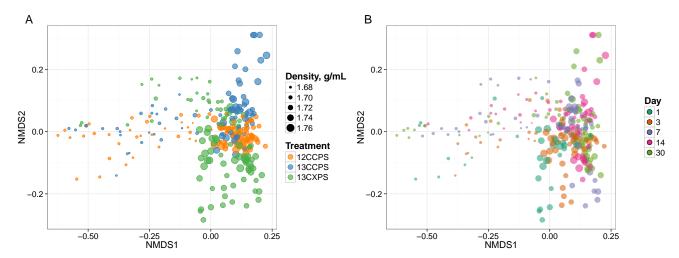


Fig. 3. NMDS analysis from weighted unifrac distances of 454 sequence data from SIP fractions of each treatment over time. Twenty fractions from a CsCl gradient fractionation for each treatment at each time point were sequenced (Fig. S1). Each point on the NMDS represents the bacterial composition based on 16S sequencing for a single fraction where the size of the point is representative of the density of that fraction and the colors represent the treatments (A) or days (B).

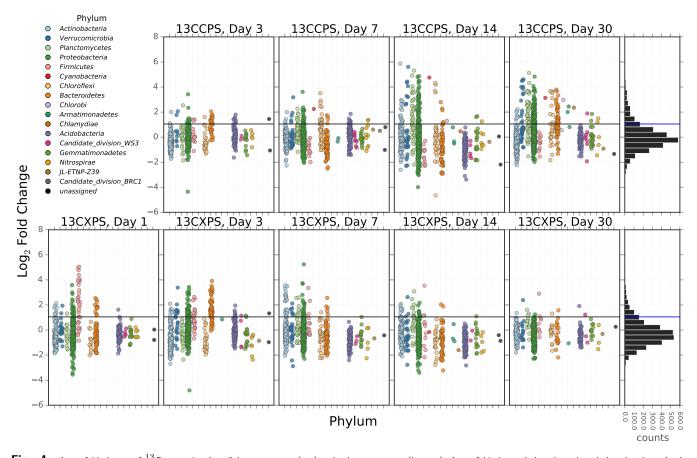


Fig. 4.  $\log_2$  fold change of  $^{13}$ C-responders in cellulose treatment (top) and xylose treatment (bottom).  $\log_2$  fold change is based on the relative abundance in the experimental treatment compared to the control within the density range 1.7125-1.755 g ml $^{-1}$ . Taxa are colored by phylum. 'Counts' is a histogram of  $\log_2$  fold change values.



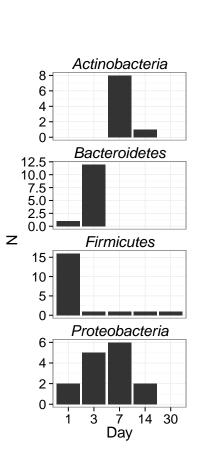


Fig. 5. Counts of <sup>13</sup>C-xylose responders in the *Actinobacteria*, *Bacteroidetes*, *Firmicutes* and *Proteobacteria* at days 1, 3, 7 and 30.







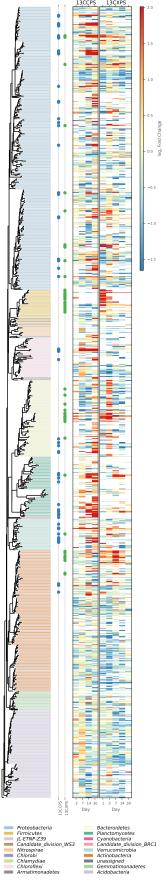


Fig. 6. Phylogenetic tree of sequences passing a user defined sparcity threshold (0.6) for at least one day of the time series. Branches are colored by phylum. <sup>13</sup>C-responders for cellulose (blue) and xylose (green) are indicated by a point beside the respective branch. Heatmap demonstrates log<sub>2</sub> fold change of each taxa through the full time series for both treatments (cellulose, left; xylose, right).









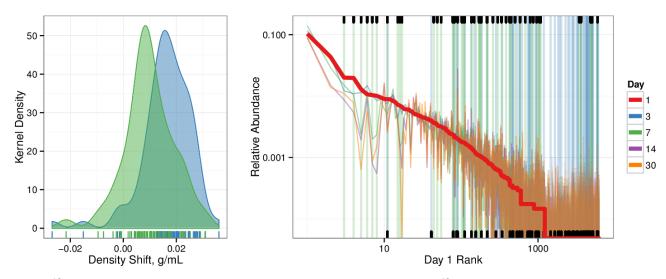


Fig. 7.  $^{13}$ C-responder characteristics based on density shift (A) and rank (B). Kernel density estimation of  $^{13}$ C-responder's density shift in cellulose treatment (blue) and xylose treatment (green) demonstrates degree of labeling for responders for each respective substrate.  $^{13}$ C-responders in rank abundance are labeled by substrate (cellulose, blue; xylose, green). Ticks at top indicate location of  $^{13}$ C-cyclose responders in bulk community. Ticks at bottom indicate location of  $^{13}$ C-cellulose responders in bulk community. OTU rank was assessed from day 1 bulk samples.

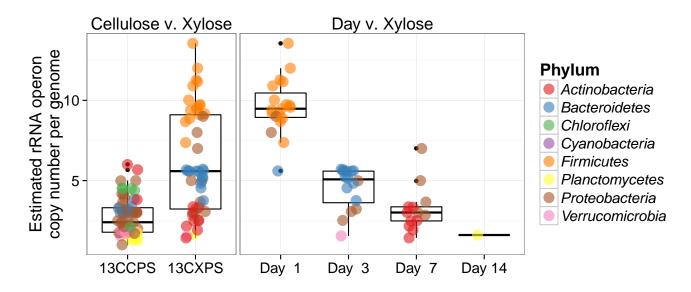


Fig. 8. Estimated rRNA operon copy number per genome for  $^{13}$ C responding OTUS. Panel titles indicate which labeled substrate(s) are depicted.

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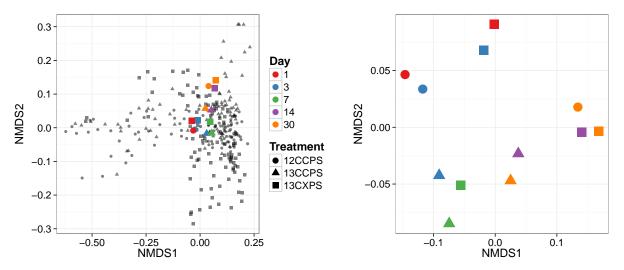
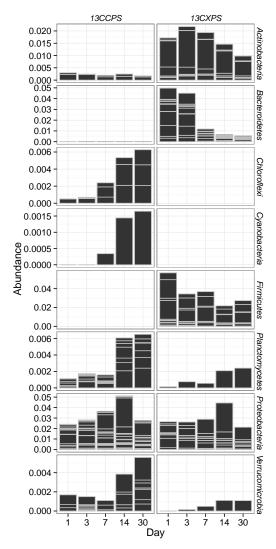


Fig. 9. Ordination of bulk gradient fraction phylogenetic profiles.



 $\textbf{Fig. 10.} \quad \text{Sum of bulk abundances with each phylum for responder OTUs.}$ 

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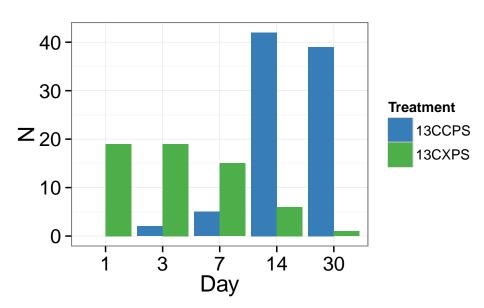


Fig. 11. Counts of responders to each isotopically labeled substrate (cellulose and xylose) over time.