

Surface Temperature Extremes and Distribution Shape Package

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The surface temperature extremes and distribution shape package computes statistics that relate to the shape of the two-meter temperature distribution and its influence on extreme temperature exceedances. These metrics evaluate model fidelity in capturing moments of the temperature distribution and distribution tail properties, as well as the large-scale meteorological patterns associated with extreme temperature exceedance days. Statistics include temperature mean, standard deviation, and skewness, along with PDFs of the underlying temperature distribution and the Gaussian fit to the distribution core of these values to identify non-Gaussianity at a location. Frequency of exceedances of a fixed extreme temperature threshold after a uniform warm shift across the distribution, the simplest prototype for future warming, provides a measure of Gaussianity as well as insight into the complexity of future changes in temperature extremes.

Version & Contact Information

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Functionality

This package consists of the following functionalities:

- (1) Moments of Surface Temperature Probability Distribution
(TempExtDistShape_Moments.py)
- (2) Shifted Underlying-to-Gaussian Distribution Tail Exceedances Ratio
(TempExtDistShape_ShiftRatio.py)
- (3) Frequency Distributions at Non-Gaussian Tail Locations
(TempExtDistShape_FreqDist.py)
- (4) Composite Circulation at Non-Gaussian Tail Locations
(TempExtDistShape_CircComps.py)

As a module of the MDTF code package, all scripts of this package can be found under
../diagnostics/temp_extremes_distshape/
and predigested observational data under
../inputdata/obs_data/temp_extremes_distshape/

Required Programming Language and Libraries

This package is written in Python 3, and requires the following Python packages:
os, json, numpy, scipy, matplotlib, mpl_toolkits, h5py, netCDF4, netcdftime, cftime,
cartopy

These Python packages are already included in the standard Anaconda installation

Required Model Output Variables

The following 3-D (lat-lon-time) model fields at a daily resolution are required:

- (1) Two-meter temperature (units: K or °C)
- (2) Sea level pressure (units: hPa or Pa)
- (3) Geopotential height (units: m)

Recommended timeseries length to facilitate statistical robustness for main features:

Preferred is 30 years matching the observational analysis; *Minimum* is 15 years

Observational Data Summary

Gridded two-meter temperature, sea level pressure (SLP), and 500 hPa geopotential height (Z500) data are from the Modern Era Retrospective-Analysis for Research and Applications reanalysis version 2 (MERRA-2; Gelaro et al. 2017, <https://doi.org/10.1175/JCLI-D-16-0758.1>). Daily means constructed from hourly data at the native horizontal resolution $0.5^\circ \times 0.625^\circ$ are provided as observational data. Analyses of MERRA-2 use the 30-year period 1980–2009 following Loikith and Neelin 2019.

This POD can be run directly on MERRA-2 data when set as the casename and input model in the default_tests.jsonc file with netcdf files supplied in the ../inputdata/model/ folder (within its own named folder “temp_extremes_distshape” per general usage instructions). To rerun MERRA-2 data with different default settings (see below) change variables under “pod_env_vars” in settings.jsonc.

Default Settings Options

Location: Frequency distributions and composite circulation maps can be output for user-specified locations. By default, the six selected locations for cold-tail frequency distributions are Yellowknife, Pendleton, Rennes St.-Jacques, Berlin, Adelaide, and North Platte (see Loikith and Neelin 2019), and the selected location for composite circulation maps is Yellowknife. Changes to locations for the frequency distributions can be made in the User Specified Section of the “...usp.py” file (TempExtDistShape_FreqDist_usp.py), and the selected location for composite maps can be altered in settings.jsonc (“city”). Note: plotting parameters for composite circulation maps may need to be altered to encompass the range of data at an alternate location, and are located in the “...usp.py” file.

Season: By default, the POD outputs cold-tail figures for the season spanning the months of December, January, and February (DJF). To output figures for the season spanning June, July, and August (JJA), change the “monthstr” and “monthsub” variables in settings.jsonc. The “monthstr” value is the string of letters corresponding to the names (‘JJA’ or ‘DJF’), whereas “monthsub” is the array of integers ([6,7,8] or [1, 2, 12]). Note: this changes the output figures for three of the functionalities only. The figure for cold-tail frequency distributions at non-Gaussian locations includes examples in both seasons and thus remains unchanged. Changes to locations and seasons in this functionality must be made within the User Specified Section of the source code (TempExtDistShape_FreqDist_usp.py).

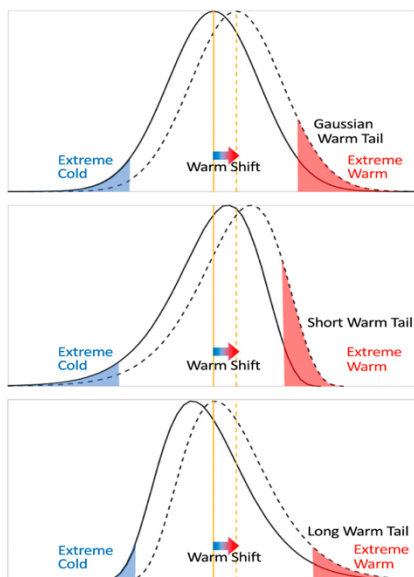
Tail: By default, the POD outputs figures for the cold-side tail of the surface temperature distribution, using the 5th percentile as the threshold from which to measure exceedances. To output figures for the warm-side tail (95th percentile), change the “ptile” percentile value in settings.jsonc. The only functionalities affected by this change are the

shifted underlying-to-Gaussian distribution tail exceedances ratio and the composite circulation figure. This is because 1) the moments of the temperature distribution are for the full distribution and not just the tail, and 2) non-Gaussian tail locations selected for the frequency distribution function are specifically chosen to represent shorter-than-Gaussian and longer-than-Gaussian cold tails following Loikith and Neelin 2019.

References

- Ruff, T. W., and J. D. Neelin, 2012: Long tails in regional surface temperature probability distributions with implications for extremes under global warming. *Geophys. Res. Lett.*, **39**, L04704, <https://doi.org/10.1029/2011GL050610>.
- Loikith, P. C., and J. D. Neelin, 2015: Short-tailed temperature distributions over North America and implications for future changes in extremes. *Geophys. Res. Lett.*, **42**, <https://doi.org/10.1002/2015GL065602>.
- Loikith, P. C., J. D. Neelin, J. Meyerson, and J. S. Hunter, 2018: Short warm-side temperature distribution tails drive hotspots of warm temperature extreme increases under near-future warming. *J. Climate*, **31**, 9469–9487, <https://doi.org/10.1175/JCLI-D-17-0878.1>.
- Loikith, P. C., and J. D. Neelin, 2019: Non-Gaussian cold-side temperature distribution tails and associated synoptic meteorology. *J. Climate*, **32**, 8399–8414, <https://doi.org/10.1175/JCLI-D-19-0344.1>.
- Catalano, A. J., P. C. Loikith, and J. D. Neelin, 2020: Evaluating CMIP6 model fidelity at simulating non-Gaussian temperature distribution tails. *Environ. Res. Lett.*, **15**, 074026, <https://doi.org/10.1088/1748-9326/ab8cd0>.

More About This Diagnostic



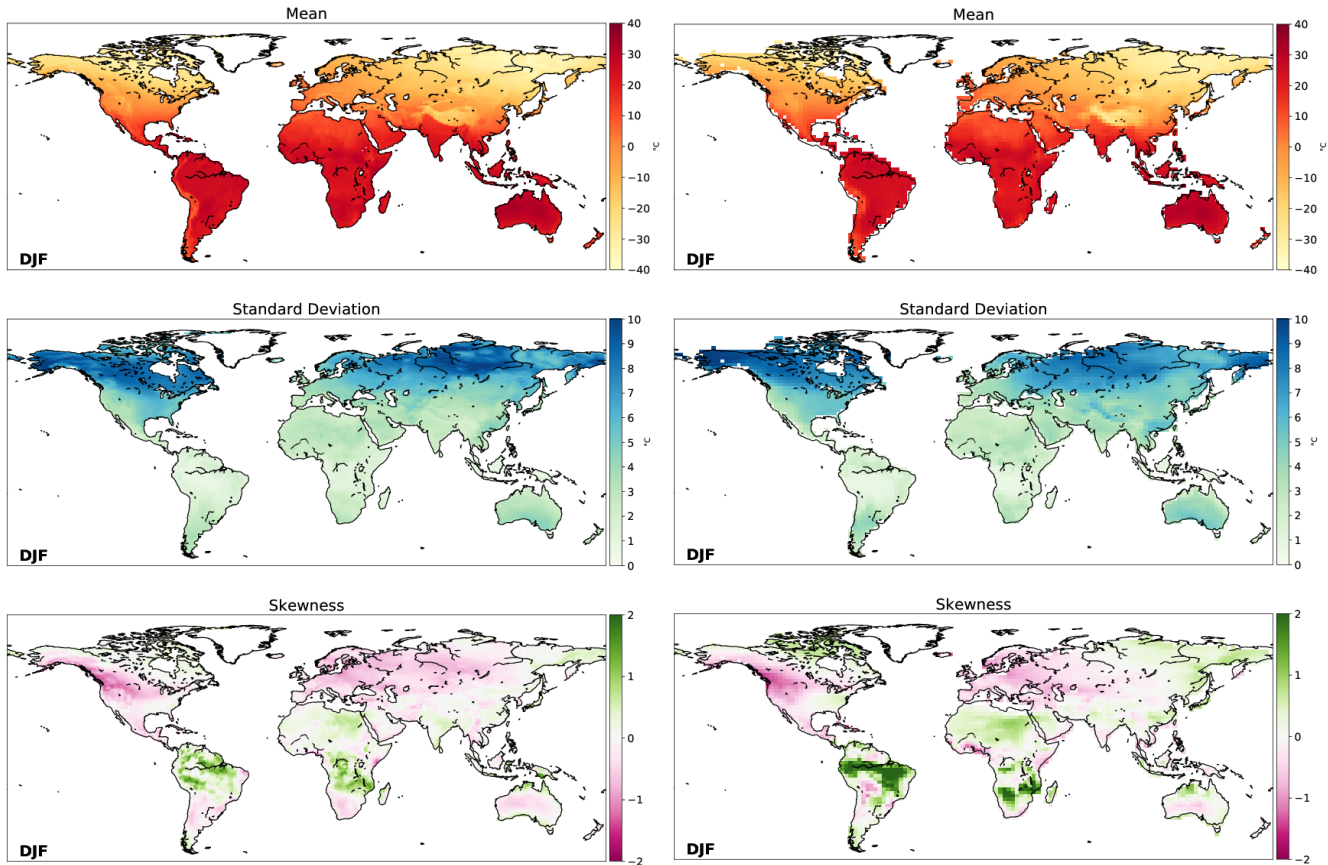
A schematic illustrating the effect of non-Gaussian tails on changes in the frequency of exceeding a fixed extreme temperature threshold under a uniform warm shift. Solid curves are the temperature frequency distributions for the pre-shifted climate while dotted curves are the temperature frequency distributions after the warm shift. Vertical yellow lines are pre-shift (solid) and post-shift (dashed) means. The red shaded area represents temperatures above the 95th percentile of the pre-shifted distribution while the blue shaded area represents the 5th percentile. Illustrations are for (top) a Gaussian, (middle) a distribution with a shorter-than-Gaussian warm tail and a longer-than-Gaussian cold tail, and (bottom) a distribution with a longer-than-Gaussian warm tail and a shorter-than-Gaussian cold tail.

The current version of the surface temperature extremes and distribution shape package produces four sets of figures for both pre-digested observations and model output, including (1) moments of the distribution (i.e. mean, variance, skewness), (2) shift ratio, (3) frequency distributions, and (4) composite large-scale meteorological patterns. The shift

ratio is computed by shifting the underlying daily temperature distribution uniformly by a fixed amount, tabulating the frequency of days that exceed a fixed pre-shifted extreme temperature threshold, and dividing that by the expected number of threshold exceedances from shifting a Gaussian by the same amount. For example, if the underlying distribution has a short warm tail, exceedances will be greater under a uniform warm shift than for a Gaussian distribution. Thus, the shift ratio identifies regions with a shorter- or longer-than-Gaussian tail. Locations

exhibiting a non-Gaussian tail can be selected for further analysis of frequency distributions and composite circulation maps. In the following, we will show an example set of the figures for a simulation of the GFDL-CM3 model (included in Phase 5 of the IPCC Coupled Model Intercomparison Project) that are produced by the package.

1) Moments of the Surface Temperature Probability Distribution

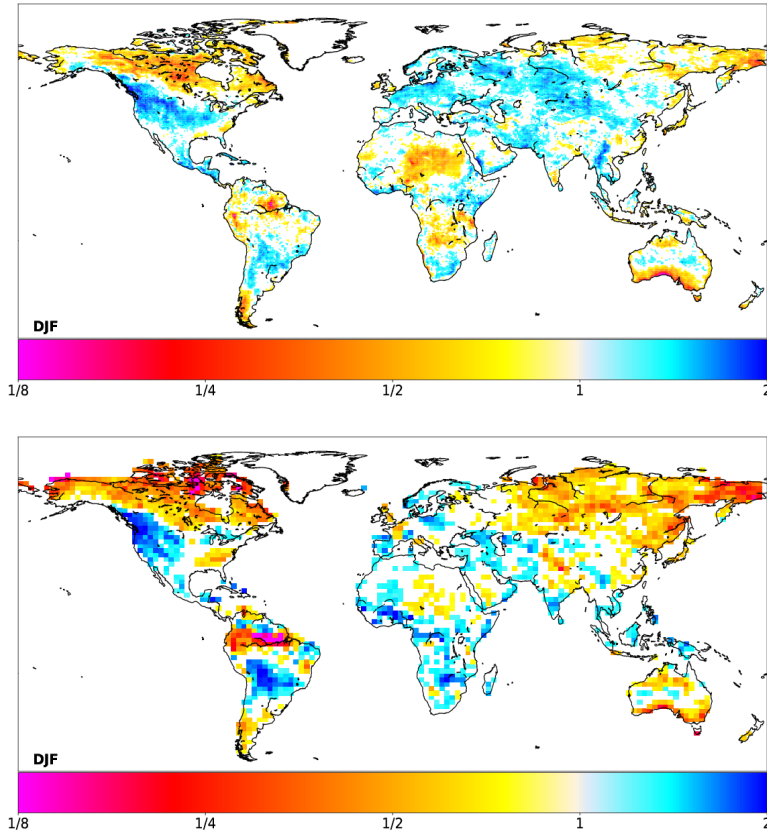


Moments for the DJF daily two-meter temperature distributions (°C) including the (*top*) mean, (*middle*) standard deviation, and (*bottom*) skewness. The left column are results from MERRA-2, and the right column are results from GFDL-CM3.

The three moments of the temperature distribution characterize its shape by measuring the central tendency, spread, and symmetry. During DJF, mean temperatures are coolest in the higher latitudes and altitudes of the Northern Hemisphere and warmest in the Southern Hemisphere. The standard deviation is largest in Northern North America and Eurasia owing to synoptic variability. Skewness measures the asymmetry of the full distribution of two-meter temperatures at each location, with positive values generally indicating a longer cold tail and negative indicating a shorter cold tail. Positively skewed regions include northern Africa, eastern Siberia, and the tropical regions of South America and Africa. Negatively skewed regions include South Africa, central Australia, Eurasia, and western North America. There is broad spatial agreement between GFDL-CM3 and MERRA-2 in the mean and standard deviation of two-meter temperature distributions over DJF, although the model does not capture some fine-scale features of standard deviation over Siberia and Yukon, which may be a result of the lower spatial resolution. GFDL-CM3 captures regions of negative skewness over western North America and Europe, and positive skewness in tropical South America and Africa, though the magnitude is higher in GFDL-CM3. Also, the model does not

adequately simulate skewness direction over some regions including Northern Russia and Southern India, which may affect how future changes in extreme temperature exceedances are manifested.

2) Shift Ratio

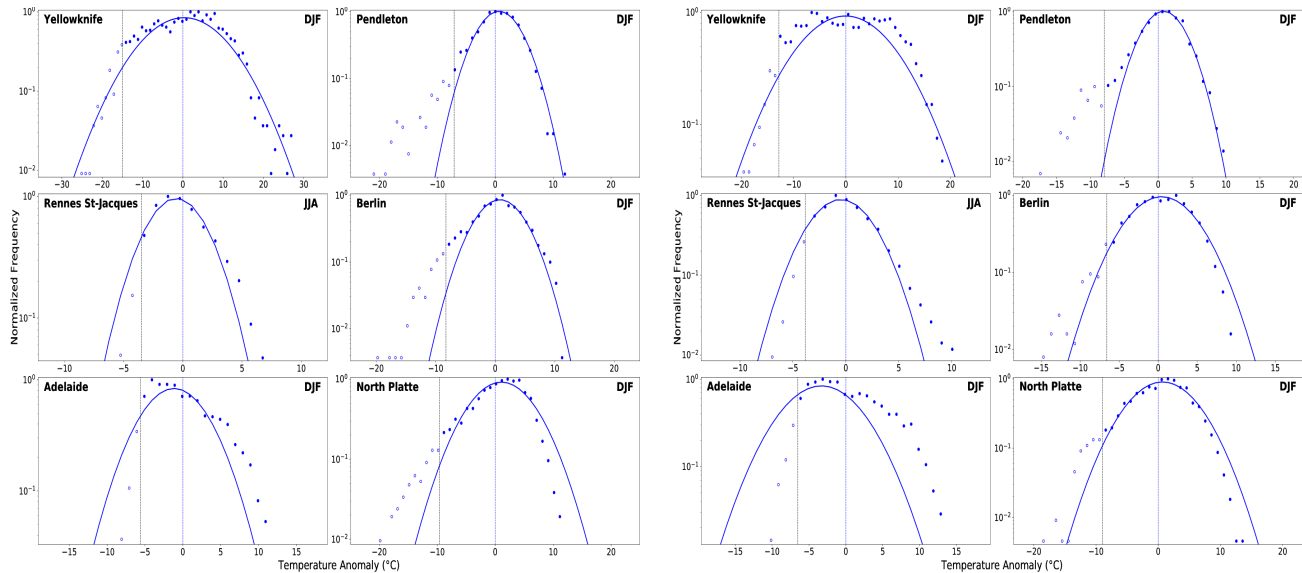


Shift ratio for cold tails during DJF, using the 5th-percentile as the fixed threshold to measure exceedances, in (top) MERRA-2 and (bottom) GFDL-CM3. Values greater than one indicate longer-than-Gaussian cold tails and values less than one indicate shorter-than-Gaussian cold tails. Only grid cells where the cold tail deviates from Gaussian with statistical significance at the 5% confidence level are shaded.

While skewness characterizes the asymmetry of the full distribution, the shift ratio provides an understanding of non-Gaussianity specific to the tail. The shift ratio identifies regions of non-Gaussianity and regions statistically indistinguishable from Gaussian based on a ratio of distribution tail exceedances following the application of a uniform shift (0.5σ). This procedure acts as both a pragmatic way of demonstrating the effects of non-Gaussian tails on extremes

exceedances and also provides a statistically robust measure of tail departure from Gaussian. Coherent regions with longer-than-Gaussian cold tails during DJF, such as northwestern United States, will experience a smaller decrease in 5th-percentile exceedances under future warming than regions with shorter-than-Gaussian cold tails (e.g., southern Australia). GFDL-CM3 captures spatially coherent short-tailed regions such as southern Australia, Canada, and northeastern Siberia, as well as long-tailed regions such as northwestern United States and central South America, although ratio amplitude is higher in GFDL-CM3. GFDL-CM3 tail shape differs from MERRA-2 over Europe and western Russia, as values are greater than one in MERRA-2 but below or insignificant in the model. This region also differs between datasets in skewness direction (see Moments of the Distribution section above), which is a related measure of asymmetry. These differences suggest that the model may simulate a smaller and potentially slower decline in DJF extreme cold temperatures over this region under future warming.

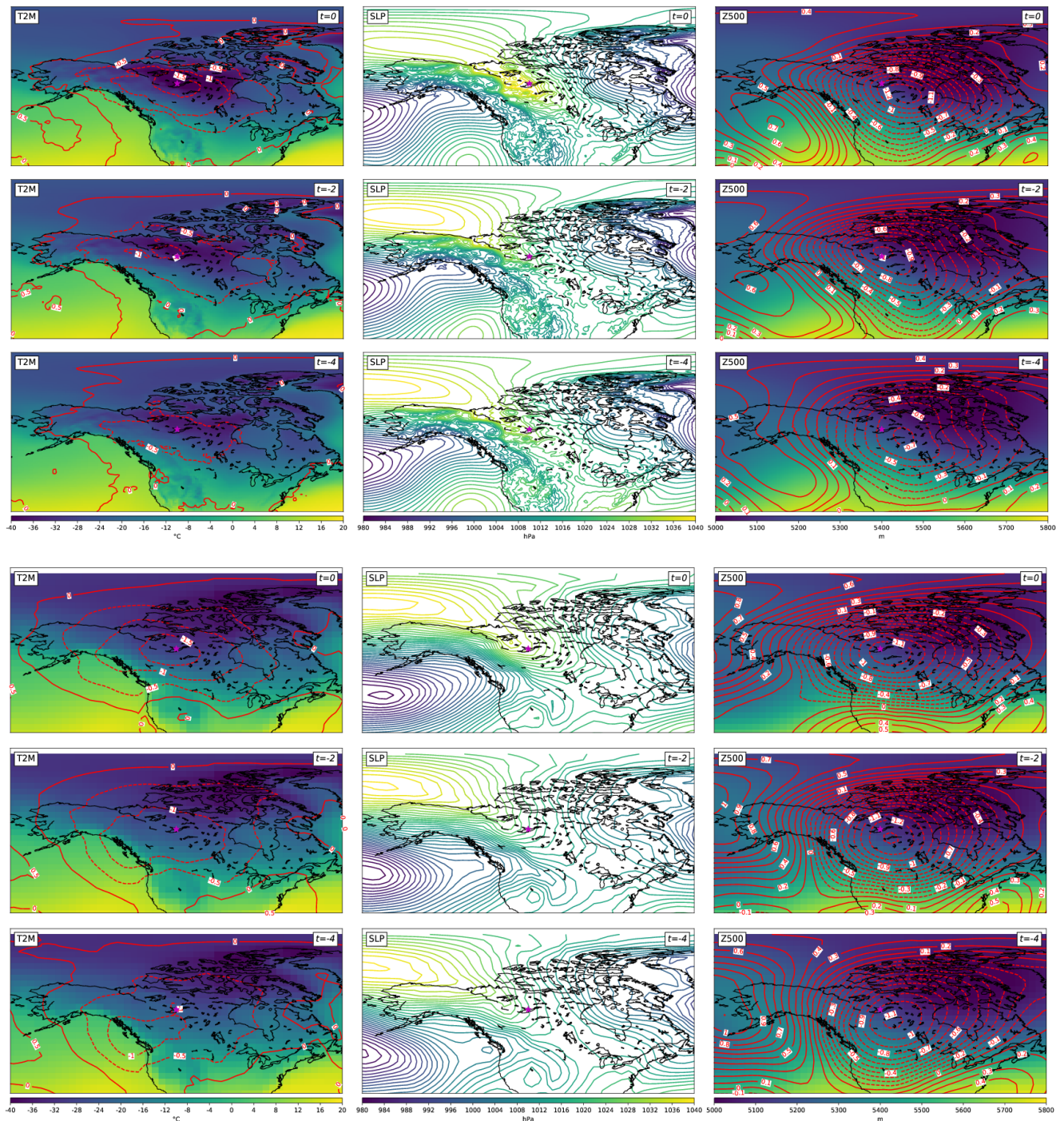
3) Frequency Distributions



Frequency distributions of two-meter temperature at grid cells nearest to select non-Gaussian locations in (*left*) MERRA-2 and (*right*) GFDL-CM3. Circles represent temperatures binned in 1°C intervals, with open circles identifying the cold tail of the distribution below the 5th percentile (black dashed line). The dashed blue line is the distribution mean, and the solid blue line is the Gaussian curve fit to the core of the distribution (values greater than 0.3 of the maximum). Frequency is provided on a log scale to highlight the tails. Short-tailed locations during respective seasons include: Yellowknife, Canada; Rennes St-Jacques, France; and Adelaide, Australia. Long-tailed locations include: Pendleton, United States; Berlin, Germany; and North Platte, United States. Note, the Gaussian at Yellowknife is fit to the entire distribution rather than the core as the Gaussian core fit was too wide, the result of which masks short-tail behavior (see Loikith et al. 2019 for more details).

Normalized frequency distributions of two-meter temperature anomalies provide a closer examination of distribution shape at selected locations exhibiting non-Gaussian behavior in the tail in the MERRA-2 distributions. A Gaussian curve is fit to the core of the underlying temperature distribution by polynomial regression, with the core being defined as values exceeding a threshold of 0.3 of the distribution maximum (Ruff and Neelin 2012). The fitted curve highlights the departures from Gaussianity in tails at select locations for the season specified. For example, the temperature distribution at Rennes St-Jacques, France exhibits a shorter-than-Gaussian cold-side tail during JJA, with bin counts (open circles) less than those estimated from the curve. Conversely, Pendleton, United States exhibits a longer-than-Gaussian cold-side tail during DJF, with binned anomalies in the cold tail nearly 10 degrees cooler than those estimated from the Gaussian fit. Overall, GFDL-CM3 simulates appropriate cold-tail non-Gaussianity, but there are some subtle differences in distributions. The Gaussian curve at Adelaide, Australia is narrower in MERRA-2 than in GFDL-CM3, indicating a flatter temperature distribution core simulated by the model (to which the curve is fit). Also, the simulated distribution at Yellowknife, Canada underestimates the extreme temperatures experienced, as values in MERRA-2 are around $\pm 30^\circ\text{F}$ but closer to $\pm 20^\circ\text{F}$ in GFDL-CM3. These differences influence the standard deviation of the distribution at each location, which is integral to the shift applied for computing the ratio of tail exceedances under this simplest prototype for future warming.

4) Composite Circulation



Composite mean of large-scale meteorological patterns in (top) MERRA-2 and (bottom) GFDL-CM3 at Yellowknife (magenta star) for days below the 5th percentile of the temperature distribution ($t=0$), two days prior to these exceedances ($t=-2$), and four days prior ($t=-4$) during DJF. Left column shading represents two-meter temperature ($^{\circ}\text{C}$), middle column contours represent sea level pressure (hPa), and right column shading represents 500-hPa geopotential heights (m). Red contours in left and right columns are temperature and geopotential height standardized anomalies, respectively.

After identifying areas exhibiting shorter-than-Gaussian or longer-than-Gaussian tails, composites of large-scale meteorological patterns lagged from tail exceedance days can be mapped to help diagnose the physical processes underlying the tail behavior. Large-scale meteorological patterns include surface temperature (T2M) and anomalies, SLP, and Z500 and anomalies for zero, two, and four days prior to tail exceedance days. The inclusion of these composites in the package is also process-oriented in that it enables the user to hypothesize possible reasons their model may or may not capture particular features. The user can specify the location at which to composite (see “Options...” below).

Here, the location example Yellowknife is used, where the DJF temperature distribution exhibits a shorter-than-Gaussian cold-side tail. Four days prior to an average cold-tail exceedance day, temperatures in the vicinity of Yellowknife are approximately -30°C , with anomalies around 0.5σ below the climatological average. An area of high pressure is present north of Alaska and extending southeastward towards Yellowknife, and Z500 anomalies indicate a broad trough is present centered east of Yellowknife. By $t=0$, temperature anomalies at the location exceed 1.5σ , SLP values have increased substantially, and the Z500 trough to the southeast has deepened. The progression of these patterns on days prior to the cold-tail exceedance day indicates that the coldest temperatures within this area of short tails are not primarily the product of transport of air from much colder places, but rather conditions that promote the development of anomalously cold air in place (Loikith and Neelin 2019). As Yellowknife is situated near the coldest temperatures in the hemisphere, there is a limit on advection of highly anomalous cold air, which explains the short tails in this region.

GFDL-CM3 generally captures the features of the circulation and surface temperatures associated with DJF cold days at Yellowknife at each lag time, which supports the notion that the model is producing a shorter-than-Gaussian tail for plausible physical reasons. Notable differences include lower SLP values in the model, particularly at $t=0$ when the high pressure is not centered on the location but rather located north of Alaska, as well as the positioning of the Z500 trough, which is centered over Yellowknife in the model. These discrepancies in general circulation features could lead to potential biases in distribution shape characteristics, which influence how temperature extremes are manifested under climate change.