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The repository principally contains the compiled products rather than the source for size reasons.

- Associated Python code (as Jupyter notebooks mostly) will be held on the same repository. The source data however might be big, so I am going to be naughty and possibly just refer you to where you might get the data if that is the case (e.g. JRA-55 data). I know I should make properly reproducible binders etc., but I didn't...
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OCES 2003 : Descriptive Physical Oceanography

(a.k.a. physical oceanography by drawing pictures)

Lecture 15: Dynamics 1 (intro to waves)

Outline

- Recap: circulation and dependence on small-scale dynamics
- waves: fundamental concepts
 - → periodicity, crest/trough/node
 - → wavelength + period
 - → frequency + wavenumber
 - → restoring force + dispersion relation
 - → propagation, phase/group velocity

Key terms: waves, wavenumber, frequency, period, dispersion relation, phase/group velocity

Reacp: MOC

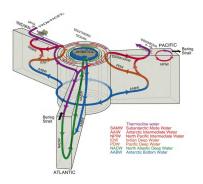


Figure: Schematic of the 3d MOC with watermass distributions. From Talley *et al.* (2011), *Descriptive Physical Oceanography*; see more in their Fig. 14.11. Format after Arnold Gordon (1991).

- MOC important for climate, carbon storage, ecology, etc.
 - → e.g. warming of Watern Europe by AMOC
 - \rightarrow e.g. carbon storage by deep water formation
- mostly along-isopycnal flow
- isolated places for watermass transformation + deep/abyssal water formation (deep convection)

Recap: what goes down must come up

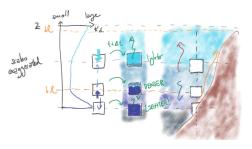


Figure: Schematic of the diffusive upwelling.

▶ diapycnal mixing contribute upwelling, strongest in boundary layers
 → broad diffusive boundary intensified upwelling

what causes the bounary intensification of κ_d ?

Recap: what goes down must come up

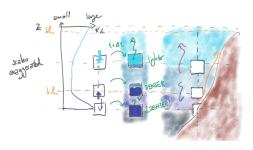


Figure: Schematic of the diffusive upwelling.

▶ diapycnal mixing contribute upwelling, strongest in boundary layers
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what causes the bounary intensification of κ_d ? dynamics!

- at the surface, lots of things... (convection, waves, Langmuir turbulence etc.)
- ▶ at the bottom, probably tidal conversion (Lec. 18) \rightarrow internal gravity waves (Lec. 16) \rightarrow shear instabilities (Lec. 17)



Recap: form stress and SO overturning

Role also of baroclinic instability (Lec. 13, see also Lec. 17), important for

- vertical momentum transfer by interfacial form stress
- scale transfer of energy
 - → mesoscale eddies, conduit between large-scales and submesoscales
- along-isopycnal mixing and also MOC

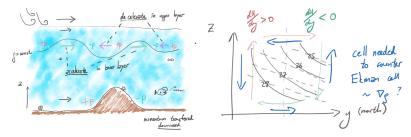


Figure: Schematic of form stress and eddy induced overturning cell in Southern Ocean (see Lec. 14)

Outlook of the next few lectures

Dynamics important, next four lectures on

- waves (this Lec. + 16, 18) and instabilities (Lec. 17)
 - → because waves are easier to talk about without maths...

Highlight gross features (i.e. those that can be drawn...)

- how to describe waves (Lec. 15)
- types of waves (Lec. 16)
 - → consequence + leading to instabilities
- ▶ instabilities (Lec. 17)
 - → parcel-type (mechanistic) arguments for instability
- ▶ tides (particularly as internal gravity waves) (Lec. 18)



Figure: Gravity waves with signal at the sea surface (as darker and lighter bands). Taken at HKUST.

Surface gravity waves¹

- restoring force is buoyancy
 - \rightarrow treat air-sea as one fluid with a giant jump in density

According to Richard Feynman, while water waves are "...easily seen by everyone... are the worse possible example [of waves], because... they have all the complications that waves can have (From Feynman, Leighton & Sands, 1971, Lectures of Physics).

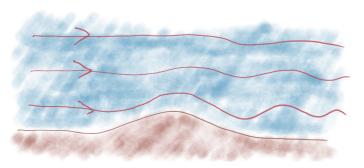


Figure: Flow over topography (e.g. tidal motion) leading to wave generation.

Tides and/or internal gravity waves forced by tidal motion

- restoring force is still buoyancy
 - → wave breaking contributing to mixing



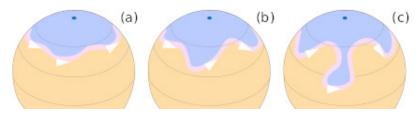


Figure: Example of jet stream meanders as Rossby waves. Figure from Wikipedia.

Rossby waves (cf. same Rossby as Rossby number)

- restoring force is Coriolis (or background gradient in vorticity)
 - \rightarrow planetary-scale waves, generic in rotating systems (e.g. rotating tanks, Jupiter, galactic disks)



Figure: Example of some other waves in systems that support waves: Alfven waves (magnetic field + Lorentz force), gravitational waves (spacetime + gravity), electromagnetic waves (but also wave-partical duality), sound waves (mechanical forcing + any medium). All figures from Wikipedia except the cello one.

Some observations:

- waves have some oscillation/periodicity
 - \rightarrow want a measure of **period**
- waves to propagate
 - \rightarrow **speed/velocity** associated with waves
- waves need a medium to travel through
 - → subtlety with Electro-Magnetic waves (not touched on here)
- characterised by a restoring force (follows from medium)
- waves can increase in amplitude and steepen
 - \rightarrow wave breaking and mixing
- can disperse, refract, interfere etc. (used in Lec. 17, 20)

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physics \Rightarrow dispersion relation, identifies the type of waves





Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

ightharpoonup described by (could also be sine)

$$\eta \sim \cos(x)$$

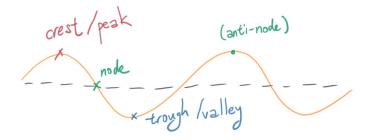


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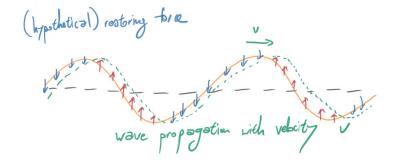


Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

• displacement η described by (could also be sine)

$$\eta \sim \cos(x - vt)$$



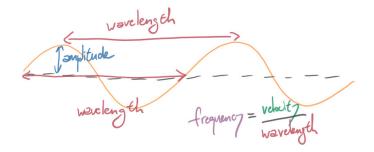


Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

ightharpoonup described by (could also be sine)

$$\eta \sim A \cos(x - vt), \qquad \gamma = v/\lambda$$



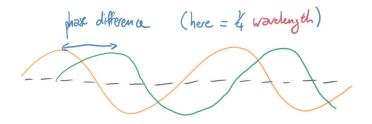


Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

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$$\eta \sim A\cos(x), \qquad \eta \sim A\cos(x - \lambda/4)$$



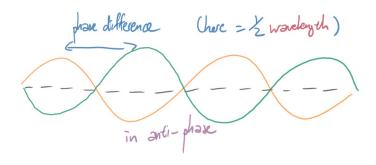


Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

displacement η described by (could also be sine)

$$\eta \sim A \cos(x), \qquad \eta \sim A \cos(x - \frac{\lambda}{2}) \sim -A \sin(x)$$



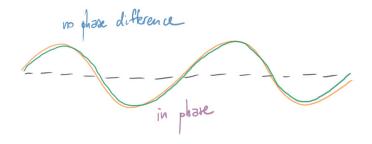


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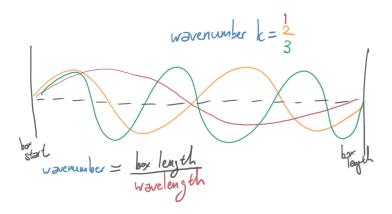


Figure: Schematic of wave features. Box length $L=2\pi$ for simplicity.

• displacement η described by (could also be sine)

$$\eta \sim A\cos(2x), \qquad \eta \sim A\cos(1x), \qquad \eta \sim A\cos(3x)$$

$$\gamma = \frac{v}{\lambda}, \qquad k = \frac{2\pi}{\lambda}$$

- γ the frequency (units: $s^{-1} = Hz$)
 - \rightarrow how quickly the wave oscillates
- $ightharpoonup v = c_p$ the **phase** velocity
 - \rightarrow how fast the wave itself moves around
- \triangleright λ the wavelength
 - \rightarrow how long the wave is
- ▶ *k* the wavenumber
 - \rightarrow intuitively how many waves can you fit in a box (so $k \sim \lambda^{-1}$)
 - \rightarrow does not necessarily have to be an integer



Features of waves: dispersion relation

Usually describe waves in terms of wavenumber k and the angular frequency $\omega = 2\pi\gamma$, i.e.

$$\eta = A\cos(\mathbf{k}x - \omega t)$$

• generally, for x = (x, y, z), we would have

$$\eta = A\cos(\mathbf{k} \cdot \mathbf{x} - \omega t)$$

 \rightarrow *k* is the wavevector

Features of waves: dispersion relation

Usually describe waves in terms of wavenumber k and the angular frequency $\omega = 2\pi\gamma$, i.e.

$$\eta = A\cos(kx - \omega t)$$

ightharpoonup generally, for x = (x, y, z), we would have

$$\eta = A\cos(\mathbf{k} \cdot \mathbf{x} - \omega t)$$

 $\rightarrow k$ is the wavevector

Aside: If you know your complex numbers, the above is neatly encapsualted as

$$\eta = \text{Real}\left[Ae^{i(\mathbf{k}\cdot\mathbf{x} - \omega t - \theta_0)}\right],$$

where e is Euler's number, $i=\sqrt{-1}$, and θ_0 denotes a phase shift if any (could be sucked into the amplitude). For calculating the dispersion relation this form is substantially nicer to deal with (don't have to keep track of sines and cosines when taking derivatives).



Features of waves: dispersion relation

Note that

$$\omega = 2\pi \gamma = 2\pi \frac{v}{\lambda} = vk$$

- the physics tells you how v = v(k)
- the dispersion relation is given by

$$\omega = \mathcal{F}(k; \ldots)$$

for some function \mathcal{F}

 \rightarrow the dispersion identifies the types of wave (see Lec. 16), e.g.

$$\omega = \sqrt{gk}, \qquad \omega = -\frac{\beta}{k}, \qquad \omega = B_0 k, \qquad \omega = \frac{\hbar k^2}{2m},$$

Superposition

(Linear) waves can be superimposed, leading to interference

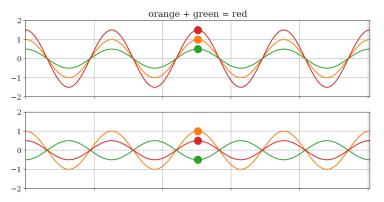
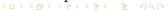


Figure: Interference of waves with Red = Orange + Green. For waves in phase (constructive inteference) and waves in anti-phase (destructive inteference).

Q. but what about waves not quite in phase or anti-phase?



Superposition

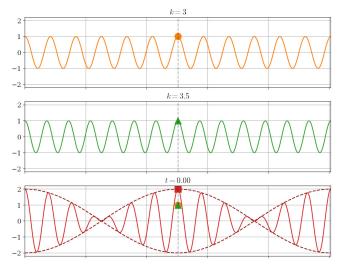


Figure: Superposition of two waves slightly out of phase, again with Red = Orange + Green. The crests have a marker marked on to track its progress later.

Recall that v = v(k) is the wave velocity. More precisely,

• the phase speed in a direction (= v) is defined as

$$c_{p,x} = \frac{\omega}{k}$$

 \rightarrow how the wave by itself travels

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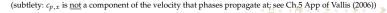
- \rightarrow how the wave by itself travels
- ▶ the group velocity c_g is defined as

$$c_{g,x} = \frac{\partial \omega}{\partial k}$$

in higher space dimensions,

$$c_{p,x} = \frac{\omega}{k_x}, \qquad c_{p,y} = \frac{\omega}{k_y}, \qquad c_{p,z} = \frac{\omega}{k_z}, \qquad c_g = \nabla_k \omega$$

 \rightarrow <u>NOTE!</u> Phase propagates in the direction of *k*





Example: 1d Rossby waves (animation)

- group velocity describes
 - \rightarrow how a collection of waves travel as a **group** or wavepacket
 - \rightarrow velocity that "stuff" propagates at
- a type of wave is non-dispersive if

$$c_g = c_p$$

- \rightarrow e.g. $\omega = B_0 k$ and $\omega = k \sqrt{gH}$ are non-dispersive
- \rightarrow if non-dispersive, wavepacket and phase travel together
- example just now is dispersive
 - ightarrow 1d Rossby waves, $\omega = -\beta/k$ (exercise: show $c_g = -c_p$ for this case)



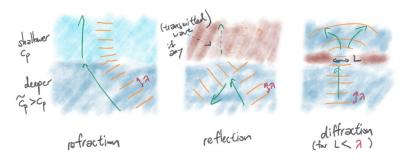


Figure: Schematic of refraction, reflection (and transmission), and diffraction, nominally using **monochromatic** (i.e. one choice of *k*) surface gravity wave as an example. The orange lines are phase lines (e.g. think wave crests).

refraction, reflection, and diffraction (used in Lec. 17, 20)

 resulting interference of waves can lead to wave steepening and wave breaking





Figure: Picture of (presumably non-monochromatic) waves over the Arabian sea. Image taken from https://www.earthqlance.com/post/133835790223/wave-diffraction-on-the-arabian-sea.

Summary

smaller-scale dynamics affects large-scale circulation

- waves are ubiquitous physical features
 - \rightarrow depends on physics
- wave described by the dispersion relation $\omega = \mathcal{F}(k)$
 - \rightarrow physics of the system dictates what $\mathcal{F}(k)$ is
 - \rightarrow usually use angular frequency ω and wavenumber k (absorbs factors of 2π floating around)
 - $\rightarrow k \sim \lambda^{-1}$ sometimes used to characterise scale of motion (more on this in Lec. 18)
- ightharpoonup difference in c_p and c_g
 - → individual (former) and collective (latter) velocity

wave breaking contributes to diapycnal mixing (see Lec. 16 + 17)

