$$\overline{U} = \nabla \Phi(x, y, t) = \lambda U = \frac{\partial x}{\partial x}, v = \frac{\partial y}{\partial y}$$

$$\nabla \cdot \overline{U} = \nabla^2 \Phi = 0$$
incompr. flow

$$y = \xi(x,t)$$

1. Kinem. cond.

$$\frac{\partial f}{\partial x} + n \frac{\partial f}{\partial x} = n$$

$$= n - \frac{\partial f}{\partial x} - n \frac{\partial f}{\partial x} = 0$$

$$\frac{\partial f}{\partial x} + n \frac{\partial f}{\partial x} = n \frac{\partial f}{\partial x} - (r \cdot r) = 0$$

2. Bernoulli law

$$\frac{3\nabla P}{3t} + \nabla (\frac{1}{3} + \frac{1}{2} u^2 + \chi) = 0$$

$$\frac{3P}{3t} + \frac{1}{3} u^2 + \chi = G(t)$$

B. law at the surf.:

Assumption: displacement is small w. respect to wavelength.

a-small param., a << L.

$$\begin{bmatrix} \frac{\partial \xi}{\partial t} + u \frac{\partial \xi}{\partial x} = V \\ + u \frac{\partial \xi}{\partial x} \end{bmatrix} (x, y, t) = \begin{bmatrix} \frac{\partial \xi}{\partial x} + u \frac{\partial \xi}{\partial x} \end{bmatrix} (x, 0, t) + O(2)$$

$$\approx \frac{\partial \xi}{\partial \xi}(x,0,t)$$

$$V(x,y,t) \approx V(x,0,t)$$

$$V(x,y,t) \approx V(x,0,t)$$
  
linearized. k. cond:  $\frac{\partial S}{\partial t} = V$  at  $y=0$ 

$$\frac{\partial P}{\partial t} + \frac{1}{2} \overline{U}^2 + 95 \rightarrow \frac{\partial P}{\partial t} + 95 = 0$$
 at  $y = 0$ 

Assumpt.: 
$$g(x,t) = A cos(kx - \omega t)$$

$$- \frac{\partial \varepsilon}{\partial t} = V = \frac{\partial \varphi}{\partial y} = \omega A \sin(kx - \omega t)$$

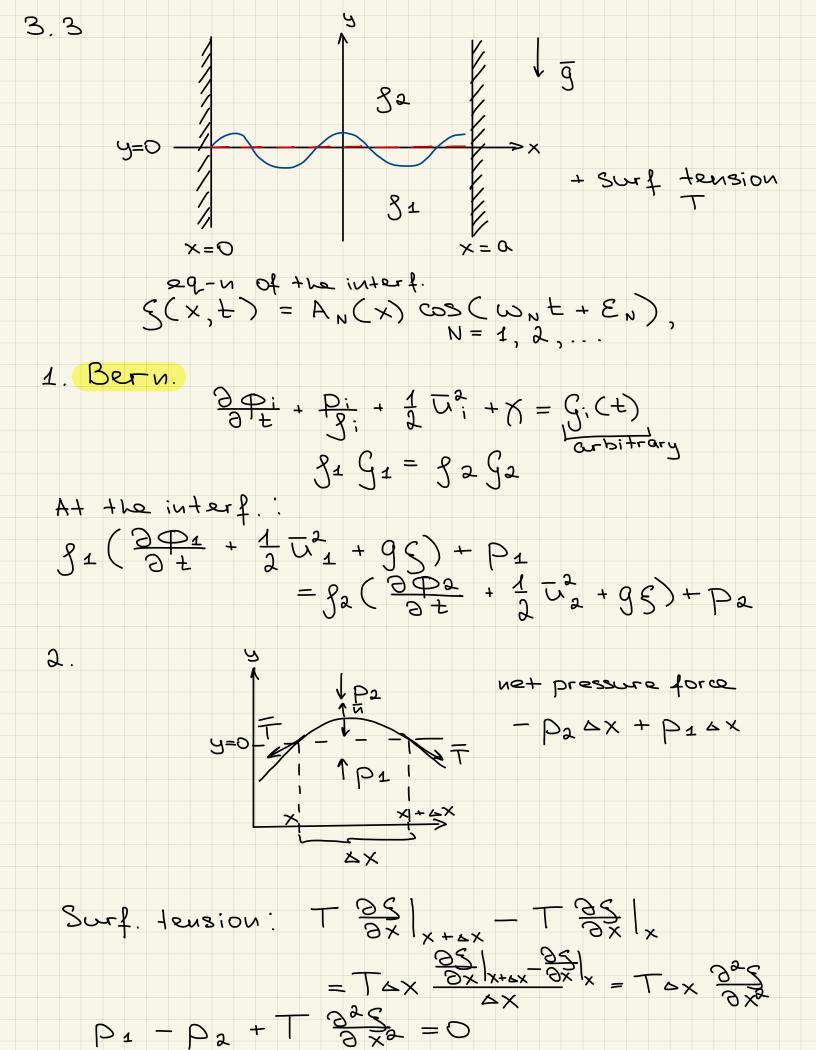
$$(4\omega - xs) \approx A = -95 = -95$$

$$= -\varphi(x,y,t) = f(y) \sin(kx - \omega t)$$

$$=> f'' - k^2 f = 0$$

y=-h no penetration of the wall

$$y = \frac{\partial Q}{\partial y} = 0 = -\frac{1}{2}(-h) = \frac{1}{2}kBe^{-kn} - \frac{1}{2}(e^{kn} = 0)$$
 $y = 0$ 
 $y = 0$ 



Solve for  $f_1(x)$ :  $f''_1 - Kf_1 = 0$ 1)  $K > 0 \rightarrow f_1(x) = Ae^{\sqrt{E}x} + Be^{-\sqrt{K}x}$ 

No penetration at x = 0 and x = a  $= 2 u = \frac{2Pi}{8x} = 0$  $\Rightarrow$  A = B = 0.

2) 
$$K = 0 = A = B = 0$$
.

3) 
$$K \geq 0$$
,  $K = -\lambda^2 \geq 0$ .  
 $f_1(x) = A \cos \lambda x + B \sin \lambda x$ 

$$f'_{1}(0) = \lambda B = 0$$

$$f'_{1}(a) = -\lambda A \sin \lambda a = 0$$

$$= \lambda = \frac{N\pi}{a}, \quad N = 1, 2, ...$$

$$f'_{1}(x) = A \cos \left(\frac{N\pi x}{a}\right)$$
For  $f_{1}(y) = C e^{\frac{N\pi}{a}y} + D e^{-\frac{N\pi}{a}y}$ 

$$At y = +\infty : \quad P_{2} = 0$$
For fluid 1:
$$P_{1}(x, y, t) = \tilde{A} \cos \left(\frac{N\pi x}{a}\right) e^{N\pi y} \sin \left(\omega_{1} x + \epsilon_{1}\right),$$

$$\tilde{A} = AC.$$
For fluid 2:
$$P_{2}(x, y, t) = \tilde{B} \cos \left(\frac{N\pi x}{a}\right) e^{N\pi y} \sin \left(\omega_{1} x + \epsilon_{1}\right),$$

$$\tilde{B} = AD.$$
Fin. cond. at  $y = 0$ :  $\frac{\partial y}{\partial t} = \frac{\partial p_{2}}{\partial y} = \frac{\partial p_{2}}{\partial y}$ 

$$\tilde{B} = AD.$$
Kin. cond. at  $y = 0$ :  $\frac{\partial y}{\partial t} = \frac{\partial p_{2}}{\partial y} = \frac{\partial p_{2}}{\partial y}$ 

$$\tilde{A} \cos \left(\frac{N\pi x}{a}\right) \sin \left(\omega_{1} x + \epsilon_{1}\right) = -\frac{N\pi}{a} \tilde{A} \cos \left(\frac{N\pi x}{a}\right) \sin \left(\omega_{1} x + \epsilon_{2}\right)$$

$$= \lambda \tilde{A} - \tilde{B}$$

$$- \omega_{1} A_{1}(x) \sin \left(\omega_{2} x + \epsilon_{2}\right) = \frac{N\pi}{a} \tilde{A} \cos \left(\frac{N\pi x}{a}\right) \sin \left(\frac{N\pi x}{a}\right)$$

$$\tilde{A}_{1} = -\frac{N\pi}{\omega_{1}} \tilde{A} \cos \left(\frac{N\pi x}{a}\right) = \tilde{A}_{1} \cos \left(\frac{N\pi x}{a}\right)$$

$$\tilde{A}_{2} = -\frac{N\pi}{\omega_{2}} \tilde{A} = \cos x + \frac{N\pi}{a} = \cos x + \frac{N\pi}{a}$$

$$g(x,t) = \tilde{A}_{N} \cos\left(\frac{NTX}{\alpha}\right) \cos(\omega_{N}t + \epsilon_{N})$$

$$g_{\perp}\left(\frac{\partial P_{\perp}}{\partial t} + g_{S}\right) = g_{2}\left(\frac{\partial P_{\perp}}{\partial t} + g_{S}\right) + T \frac{\partial^{2}Q}{\partial x^{2}}$$

$$g_{\perp}\left[\tilde{A}_{L} \omega_{N} \cos\left(\frac{NTX}{\alpha}\right) \cos\left(\frac{NTX}{\alpha}\right) + \frac{2}{NL} \tilde{A}_{L} \cos\left(\frac{NTX}{\alpha}\right) \cos\left(\frac{NTX}{\alpha$$

No surf. tension.

We assume 
$$\xi(x,t) = A\cos(kx - \omega t)$$

1. Balance of forces

2. Bernoulli

91 > 92

$$g_{1}(\frac{\partial P_{1}}{\partial t} + U_{1}^{2} + \chi) + P_{1} = g_{2}(\frac{\partial P_{2}}{\partial t} + U_{2}^{2} + \chi) + P_{2}$$

At the surf:
$$g_1(\frac{\partial p_1}{\partial t} + g_5) = g_2(\frac{\partial p_2}{\partial t} + g_5)$$
 at  $y=0$ 

Form of pi

$$\frac{\partial P_i}{\partial y} = A \omega \sin(kx - \omega t)$$

$$=> P_i = f_i(y) \sin(kx - \omega t)$$

$$\nabla^{2} P_{1} = 0 = - k^{2} f_{1} + f_{1}^{"} = 0$$

$$P_{1}(x,y,t) = (Be^{ky} + Ce^{-ky}) \sin(kx - \omega t)$$

$$P_{2}(x,y,t) = (De^{ky} + Ee^{-ky}) \sin(kx - \omega t)$$

$$P_{3}(x,y,t) = (De^{ky} + Ee^{-ky}) \sin(kx - \omega t)$$

$$P_{4} = Be^{ky} \sin(...), P_{2} = Ee^{-ky} \sin(...)$$

$$P_{4} = Be^{ky} \sin(...), P_{2} = Ee^{-ky} \sin(...)$$

$$P_{5} = A \omega \sin(kx - \omega t) = \frac{\partial P_{4}}{\partial y} = \frac{\partial P_{2}}{\partial y}$$

$$P_{5} = A \omega \sin(kx - \omega t) = \frac{\partial P_{4}}{\partial y} = \frac{\partial P_{4}}{\partial y}$$

$$P_{5} = Ee^{-ky} \sin(...) = \frac{\partial P_{4}}{\partial x} + \frac{\partial P_{4}}{\partial y}$$

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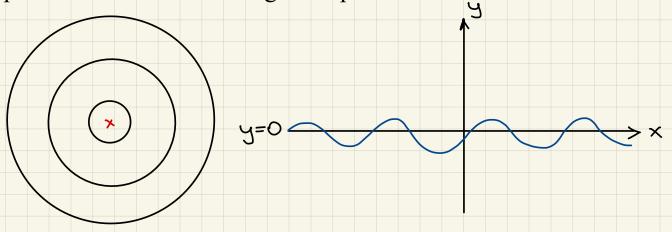
$$P_{5} = \frac{\partial P_{5}}{\partial y} + \frac{\partial P_{5$$

Note that in the book wanted -> if you assume an sin(-kx-wt) ( wave propagate in the opposite direction) -> the same disp. relation.

3.5 When a stone is dropped into a deep pond, waves are eventually observed only beyond a central region of calm water which expands in radius with time. Furthermore, the wavelength just beyond this calm region is constant, about 4.5 cm.

Use 2D plane wave theory, including both gravity and surface tension, to account broadly for these observations, and obtain an estimate for the

speed at which the calm region expands.



1) Balance forces
$$p_{\star} - p = T \frac{2^2 S}{8 \times 2} \text{ at } y = S(x,t)$$

At the surf: 
$$\frac{\partial Q}{\partial t} + gg = \frac{Pa}{g} - \frac{D}{g} = \frac{T}{g} \frac{\partial^2 G}{\partial x^2}$$
at  $y = 0$ .

Assumption: 
$$g(x,t) = A \cos(kx - \omega t)$$

-> use lin. kinem. cond.  $\frac{\partial g}{\partial t} = \frac{\partial \phi}{\partial y}$ 

->  $\phi(x,y,t) = f(y) \sin(kx - \omega t)$ 

$$\varphi = (Be^{ky} + Ce^{-ky}) \sin(kx - \omega +)$$
At  $y \Rightarrow -\infty \Rightarrow \varphi = 0 \Rightarrow C = 0$ .

$$\frac{\partial \xi}{\partial \xi} = A\omega \sin(...) = \frac{\partial \varphi}{\partial y} = k B \sin(...)$$
at  $y = 0$ 

$$= -B = \frac{\omega}{k} A$$
Bernoull;:  $-\frac{\omega^2}{k} A \cos(kx - \omega +) + g A \cos(kx - \omega +)$ 

$$= -\frac{1}{2} k^2 A \cos(kx - \omega +)$$

$$\omega^2 = gk + \frac{7k^3}{3}$$

$$\frac{disp. rel.}{for deep water}$$
Speed at which wavepacket propag. away from where you'se dropped a stoke.

$$C_g = \frac{d\omega}{dR}$$

$$\frac{d\omega^2}{dR} = 2\omega C_g = 9 + 2TR^2$$

$$= -C_g = \frac{9 + 2TR^2}{2\sqrt{gR} + TR^2} = C_g(R)$$

$$g = 9.81 \frac{m}{S^2}, S = 10^3 \frac{R_g}{M^3}, T = 0.074 \frac{N}{m}$$
Let's evaluate  $C_g$  for max. wavelength  $C_g$ 

$$\frac{dC_g}{dR} = 0$$

3.8 Surface waves generated by a mid-Atlantic storm arrive at the British coast with period 15 seconds. A day later the period of the waves arriving has dropped to 12.5 seconds. Roughly how far away did the storm occur?

$$y=0 \longrightarrow x$$

$$g(x,t) = A\cos(kx-\omega t)$$
Our model — surface waves on deep water.

At y=0:

1) Lin. kinem. cand.:  $\frac{2S}{2t} = V = \frac{BT}{3y}$ 
2) Lin. Bernoulli:  $\frac{C}{3t} + gS = 0$ 

see ex. 3.1

$$\frac{C}{3t} = V = \frac{BT}{3y}$$

$$\frac{C}{3t} = V = \frac{BT}{3t}$$

$$\frac{C}{3t} = V = \frac{T}{3t}$$

=> disp. relation is we = gk.

Speed at which the wave packet gen. by the storm moves:  $c_g = \frac{d\omega}{dk}, \quad 2\omega \frac{d\omega}{dk} = g$   $= > c_g = \frac{g}{2\omega} = \frac{1}{2\sqrt{gk}} = \frac{1}{2\sqrt{k}}$ It will propage at dist. x with veloce cg in ts:  $x = Cgt = = Cg = \frac{x}{t} = \frac{1}{2}\sqrt{\frac{g}{k}}$  $= -\frac{x^2}{t^2} = \frac{a}{4k}, k = \frac{at^2}{4x^2}$ local wavenumber at x and t Subs. into disp. relation:

 $\omega' = \frac{g + t}{g \times x}$ 

We know loc. freq. of this wave packet when they arrive at the Brit. coast after some time to the storm our dist. from the storm

And a day later to = t1 - 24h:  $\omega_2 = \frac{2\pi}{T_2} = \frac{g+2}{2x} \rightarrow t_2 = \frac{4\pi \times g}{gT_2}$  $t_1 + 24h = \frac{4\pi \times}{9T_1} + 86400s = \frac{4\pi \times}{9T_2}$  $-> \frac{4\pi x}{9} \left( \frac{1}{T_a} - \frac{1}{T_1} \right) = 86400s$  $x = 86400 \frac{T_1 I_2}{T_1 - T_2} \frac{9}{44} \approx 5059 \text{ Km}$