Do we need demographic data to forecast population responses to climate change?

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Abstract. Rapid climate change has generated growing interest in forecasts of future population trajectories. Traditional population models, typically built using detailed demographic observations from one study site, can address climate change impacts at one location, but are difficult to scale up to the landscape and regional scales relevant to management decisions. An alternative is to build models using population-level data that are much easier to collect over broad spatial scales than individual-level data. However, such models ignore the fact that climate drives population growth through its influence on individual performance. Here, we test the consequences of aggregating individual responses when forecasting climate change impacts on four perennial grass species in a semi-arid grassland in Montana, USA. We parameterized two population models for each species, one based on individual-level data (survival, growth and recruitment) and one on population-level data (percent cover), and compared their accuracy, precision, and sensitivity to climate variables.

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For both models we used Bayesian ridge regression to identify the optimal predictive model in terms of climate covariate strengths. The individual-level model was more accurate and precise than the aggregated model when predicting out-of-sample observations. When comparing climate effects from both models, the population-level model missed important climate effects from at least one vital rate for each species. Increasing the sample size at the population-level would not necessarily reduce forecast uncertainty; the way to reduce uncertainty is to capture unique climate dependence of individual vital rates. Our analysis indicates that there is no shortcut to forecasting climate change impacts on plant populations — detailed demographic data are essential. Despite the superiority of the individual-level model, the forecasts it generated still were too uncertain to be useful for decision-makers. We need new methods to collect demographic data efficiently across environmental gradients in space and time.

Key words: forecasting, climate change, grassland, integral projection model, population model

$_{40}$ Introduction

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- Perhaps the greatest challenge for ecology in the 21st century is to forecast the impacts
- 42 of environmental change (Clark et al. 2001, Petchey et al. 2015). Forecasts require so-
- 43 phisticated modeling approaches that fully account for uncertainty and variability in both
- ecological process and model parameters (Luo et al. 2011, but see Perretti et al. 2013
- for an argument against modeling the ecological process). The increasing statistical so-
- 46 phistication of population models (Rees and Ellner 2009) makes them promising tools for
- 47 predicting the impacts of environmental change on species persistence and abundance.
- But reconciling the scales at which population models are parameterized and the scales at
- 49 which environmental changes play out remains a challenge (Clark et al. 2010, 2012, Freck-
- be leton et al. 2011, Queenborough et al. 2011). The problem is that most population models

tracking the fates of individuals plants, is so difficult. The resulting models cannot be applied to the landscape and regional scales relevant to decision-making without information about how the fitted parameters respond to spatial variation in biotic and abiotic drivers (Sæther et al. 2007). The limited spatial extent of individual-level demographic datasets constrains our ability to use population models to address applied questions about the consequences of climate change. The inability of most population models to address landscape-scale problems may explain why land managers and conservation planners have embraced species distribution models (SDMs) (see Guisan and Thuiller 2005 for a review). SDMs typically rely on easy-to-collect presence/absence data (but see Clark et al. 2014 for new methods) and remotely-sensed environmental covariates that allow researchers to model large spatial extents (e.g., Maiorano et al. 2013). Thus, it is relatively straightforward to parameterize and project SDMs over landscapes and regions. However, the limitations of SDMs are well known (Pearson and Dawson 2003, Elith and Leathwick 2009, Araújo and Peterson 2012). Ideally, researchers would provide managers with landscape-scale population models, combining the extent of SDMs with information about dynamics and species abundances (Schurr et al. 2012, Merow et al. 2014). Aggregate measures of population status, rather than individual performance, offer an intriguing alternative for modeling populations (Clark and Bjørnstad 2004, Freckleton et al. 2011). Population-level data cannot provide inference about demographic mechanisms, 71 but might be sufficient for modeling future population states, especially since such data are feasible to collect across broad spatial extents (e.g., Queenborough et al. 2011). The choice between individual and population-level data involves a difficult trade-off: while individual-level data leads to more mechanistic models, population-level data leads to models that can be applied over greater spatial and temporal extents. An open question is how much forecasting skill is lost when we build models based on population rather than

are built using data from a single study site because collecting those data, which involves

⁷⁸ individual-level data.

attempts to capitalize on population-level measures. An important exception was an effort 80 by Taylor and Hastings (2004) to model the population growth rate of an invasive species to identify the best strategies for invasion control. They used a "density-structured" model 82 where the state variable is a discrete density state rather than a continuous density measure. Such models do not require individual-level demographic data and can adequately describe population dynamics. Building on Taylor and Hastings (2004), Freckleton et al. (2011) showed that density-structured models compare well to continuous models in theory, and Queenborough et al. (2011) provide empirical evidence that density-structured models are capable of reproducing population dynamics at landscape spatial scales, even if some precision is lost when compared to fully continuous models. The appeal of densitystructured approaches is clear. However, none of these models included environmental covariates. Addressing climate change questions with models fit to population-level data is potentially problematic. It is individuals, not populations, that respond to climate variables (Clark 93 et al. 2012). Ignoring this fact amounts to an "ecological fallacy", where inference about the individual relies on statistical inference on the group (Piantadosi et al. 1988). Population growth (or decline) is the outcome of demographic processes such as survival, growth, and recruitment that occur at the level of individual plants. Climate can affect each demographic process in unique, potentially opossing, ways (Dalgleish et al. 2011). These unique climate responses may be difficult to resolve in statistical models based on population-level data where demographic processes are not identifiable. If important climate effects are 100 missed because of the aggregation inherent in in population-level data, then population 101 models built with such data will make uninformative or unreliable forecasts. 102 Here, we compare the forecasting skill of statistical and population models based on aggre-

To date, most empirical population modelers have relied on individual-level data, with few

gated, population-level data with models based on individual-level data. We used a unique demographic dataset that tracks the fates of individual plants from four species over 14 105 years to build two kinds of single-species population models, traditional models using indi-106 vidual growth, survival, and recruitment data and alternative models based on basal cover. 107 In both models, interannual variation is explained, in part, by climate covariates. We first 108 quantify forecasting skill using cross-validation. We then performed simulations to quan-100 tify the sensitivities of species' cover to small perturbations in average precipitation and 110 temperature. Based on the cross-validation results, predictions of individual level models 111 were clearly better, but, unfortunately, still too uncertain to inform management decisions. 112

113 Materials and Methods

114 Study site and data

Our demographic data come from the Fort Keogh Livestock and Range Research Labo-115 ratory in eastern Montana's northern mixed prairie near Miles City, Montana, USA (46° 116 19' N, 105° 48' W). The dataset is freely available on Ecological Archives¹ (Anderson et al. 117 2011), and interested readers should refer to the metadata for a complete description. The 118 site is about 800 m above sea level and mean annual precipitation (1878-2009) is 334 mm, with most annual precipitation falling from April through September. The community 120 is grass-dominated, and we focused on the four most abundant grass species: Bouteloua 121 gracilis (BOGR), Hesperostipa comata (HECO), Pascopyrum smithii (PASM), and Poa 122 secunda (POSE) (Fig. 1). B. gracilis is a warm-season perennial grass, whereas H. comata, 123 P. smithii, and P. secunda are cool-season perennial grasses. All species typically begin 124 growth in the early spring, reach maximum growth and flower in early to mid summer 125 (May-June), and disperse seed in mid to late summer (July-September). 126

 $_{127}$ From 1932 to 1945 individual plants were identified and mapped annually in 44 1-m²

¹http://esapubs.org/archive/ecol/E092/143/

quadrats using a pantograph. The quadrats were distributed in six pastures, each assigned a grazing treatment of light (1.24 ha/animal unit month), moderate (0.92 ha/aum), 129 and heavy (0.76 ha/aum) stocking rates (two pastures per treatment). In this analysis 130 we account for potential differences among the grazing treatments, but do not focus on 131 grazing×climate interactions. The annual maps of the quadrats were digitized and the 132 fates of individual plants tracked and extracted using a computer program (Lauenroth and 133 Adler 2008, Chu et al. 2014). The permanent quadrats have not been relocated, but their 134 distribution in six different pastures means the data represent a broad spatial distribution 135 for the study area. Daily climate data are available for the duration of the data collection 136 period (1932 - 1945) from the Miles City airport, Wiley Field, 9 km from the study site. 137 We modeled each grass population based on two levels of data: individual and quadrat 138 (Fig. 2). The individual data is the "raw" data. For the quadrat-level we data we simply 139 sum individual basal cover for each quadrat by species. This is equivalent to a near-perfect 140 census of quadrat percent cover because previous analysis shows that measurement error 141 at the individual-level is small (Chu and Adler 2014). Based on these two datasets we can 142 compare population models built using individual-level data and aggregated, quadrat-level 143 data. At the individual level we explicitly model three vital rates: growth (13 year-to-year 144 transitions, 29 quadrats, 18,730 records), survival (13 years, 33 quadrats, 29,353 records), 145 and recruitment (13 years, 33 quadrats, 304 records). At the quadrat level we model popu-146 lation growth as change in percent cover of quadrats with non-zero cover in year t and in 147 year t-1 (13 year-to-year transitions, 29 quadrats, 866 records). For modeling population 148 growth at the quadrat level we ignore within-quadrat extirpation and colonization events 149 because they are very rare in our time series (N = 16 and N = 13, respectively, across all 150 species). Given the broad spatial distribution of the quadrats we are studying, it is safe to 151 assume that these events are in fact rare enough to be ignored for our purposes. 152

53 All R code and data necessary to reproduce our analysis is archived on GitHub as release

v1.0² (http://github.com/atredennick/MicroMesoForecast/releases). That stable release
will remain static as a record of this analysis, but subsequent versions may appear if we
update this work. We have also deposited the v1.0 release on Dryad (*link here after acceptance*).

At both levels of inference (individual and quadrat), the building blocks of our population

Stastical models of vital rates

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models are vital rate regressions. For individual-level data, we fit regressions for survival, 160 growth, and recruitment for each species. At the quadrat-level, we fit a single regression 161 model for population growth. We describe the statistical models separately since fitting 162 the models required different approaches. All models contain five climate covariates that 163 we chose a priori: "water year" precipitation at t-2 (lagppt); April through June precipi-164 tation at t-1 and t (ppt1 and ppt2, respectively) and April through June temperature at 165 t-1 and t (TmeanSpr1 and TmeanSpr2, respectively), where t-1 to t is the transition of 166 interest. We also include interactions among same-year climate covariates (e.g., ppt1 × 167 TmeansSpr1). 168 We fit all models using a hierarchical Bayesian approach. The models are fully descibed in Appendix A, so here we focus on the main process and the model likelihood. For the likelihood models, \mathbf{y}^{X} is always the relevant vector of observations for vital rate X (X = S, G, R, or P) for survival, growth, recruitment, or population growth). For example, \mathbf{y}^S is a vector of 0's and 1's indicating whether a genet survives from t to t+1, or not. All model 173 parameters are species-specific, but we omit subscripts for species in model descriptions 174 below to reduce visual clutter. 175

 $^{^2}Note\ to\ reviewers$: so that v1.0 will be associated with the published version of the manuscript, we have released v0.1 to be associated with this review version.

Vital rate models at the individual level We used logistic regression to model survival probability (s) of a genet:

$$\mathbf{y}^S \sim \text{Bernoulli}(\mathbf{s})$$
 (1)

$$logit[s(x, \mathbf{z}_t, w)] = \beta_{0,t} + \beta_{s,t}x + \beta_Q + \mathbf{z}_t' \boldsymbol{\beta}_c + \beta_{d,1}w + \beta_{d,2}(xw)$$
(2)

where x is the log of genet basal area, $\beta_{0,t}$ is a year specific intercept intercept, β_Q is the random effect of quadrat group location, $\beta_{s,t}$ is the year-specific slope parameter for size, \mathbf{z} is a vector of i climate covariates specific to year t, $\boldsymbol{\beta}_c$ is a vector of fixed climate effects of length i, $\beta_{d,1}$ is the effect of intraspecific crowding experienced by the focal genet (w), and $\beta_{d,2}$ is a size by crowding (xw) interaction effect.

We modeled growth as a Gaussian process describing genet size at time t+1 as a function of size at t and climate covariates:

$$\mathbf{y}^G \sim \text{Normal}(\boldsymbol{\mu}, \sigma_{xt}^2)$$
 (3)

$$\mu(x, \mathbf{z}_t, w) = \beta_{0,t} + \beta_{s,t} x + \beta_Q + \mathbf{z}_t' \boldsymbol{\beta}_c + \beta_{d,1} w + \beta_{d,2}(xw)$$

$$\tag{4}$$

where $\mu(x, \mathbf{z}_t, w)$ is log of predicted genet size at time t+1, and all other parameters are as described for the survival regression. We capture non-constant error variance in growth by modeling the variance around the growth regression $(\sigma_{x,t}^2)$ as a nonlinear function of predicted genet size:

$$\sigma_{x,t}^2 = a \exp[b \times \mu(x, \mathbf{z}_t, w)] \tag{5}$$

where $\mu(x, \mathbf{z}_t, w)$ is log of predicted genet size predicted from the growth regression (Eq. 4), and a and b are constants.

Our data allows us to track new recruits, but we cannot assign a specific parent to new genets. Therefore, we model recruitment at the quadrat level. We assume the number of individuals, y^R , recruiting at time t+1 in quadrat q follows a negative binomial distribution:

$$y_{q,t+1}^R \sim \text{NegBin}(\lambda_{q,t+1}, \phi)$$
 (6)

where λ is the mean intensity and ϕ is the size parameter. We define λ as a function of quadrat composition and climate in the previous year:

$$\lambda_{q,t+1} = C'_{q,t} \exp\left(\beta_{0,t} + \beta_Q + \mathbf{z}'_t \boldsymbol{\beta}_c + \beta_d \sqrt{C'_{q,t}}\right)$$
 (7)

where C' is effective cover (cm^2) of the focal species in quadrat q, and all other terms are as in the survival and growth regressions. Effective cover is a mixture of observed cover (C) in the focal quadrat (q) and the mean cover across the entire group (\bar{C}) of Q quadrats in which q is located:

$$C'_{q,t} = pC_{q,t} + (1-p)\bar{C}_{Q,t} \tag{8}$$

where p is a mixing fraction between 0 and 1 that is estimated within the model.

Population model at the quadrat level The statistical approach used to model ag-207 gregated data depends on the type of data collected. We have percent cover data, which 208 can easily be transformed to proportion data. An obvious choice for fitting a linear model 209 to proportion data is beta regression because the support of the beta distribution is [0,1], 210 not including true zeros or ones. However, when we used fitted model parameters from 211 a beta regression in a quadrat-based population model, the simulated population tended 212 toward 100% cover for all species. We therefore chose a modeling approach based on a 213 truncated log-normal likelihood. The model for quadrat cover change from time t to t+1 is 214

$$\mathbf{y}^P \sim \text{LogNormal}(\mu(x, \mathbf{z}_t), \sigma^2) \mathbf{T}[0, 1]$$
 (9)

$$\mu(x, \mathbf{z}_t) = \beta_{0,t} + \beta_{s,t} x + \beta_Q + \mathbf{z}_t' \boldsymbol{\beta}_c \tag{10}$$

where $\mu(x, \mathbf{z}_t)$ is the log of proportional cover in quadrat q at time t+1, and all other
parameters are as in the individual-level growth model (Eq. 4) except that x now represents log of proportional cover. The log normal likelihood includes a truncation (T[0,1]) to
ensure that predicted values do not exceed 100% cover.

220 Model fitting and stastical regularization

Model fitting Our Bayesian approach to fitting the vital rate models required choosing appropriate priors for unknown parameters and deciding which, if any, of those pri-222 ors should be hierarchical. We decided to fit models where all terms were fit by species. 223 Within a species, we fit yearly size effects and yearly intercepts hierarchically where year-224 specific coefficients were drawn from global distributions representing the mean size effect 225 and intercept. Quadrat random effects were also fit hierarchically, with quadrat offsets 226 being drawn from distributions with mean zero and a shared variance term (independent 227 Gaussian priors, Appendix A). Climate effects were not modeled hierarchically, and each 228 was given a diffuse prior distribution. We used standard diffuse priors for all unknown 220 parameters (Appendix A). 230 All of our analyses (model fitting and simulating) were conducted in R (R Core Team 231 2013). We used the 'No-U-Turn' MCMC sampler in Stan (Stan Development Team 2014a) to estimate the posterior distributions of model parameters using the package 'rstan' (Stan 233 Development Team 2014b). We obtained posterior distributions for all model parame-234 ters from three parallel MCMC chains run for 1,000 iterations after discarding an initial 235 1,000 iterations. Such short MCMC chains may surprise readers more familiar with other 236 MCMC samplers (i.e. JAGS or WinBUGS), but the Stan sampler is exceptionally efficient, 237 which reduces the number of iterations needed to achieve convergence. We assessed con-238 vergence visually and made sure scale reduction factors for all parameters were less than 239 1.01. For the purposes of including parameter uncertainty in our population models, we 240 saved the final 1,000 iterations from each of the three MCMC chains to be used as randomly drawn values during population simulation. This step alleviates the need to reduce model parameters by model selection since sampling from the full parameter space in the MCMC ensures that if a parameter broadly overlaps zero, on average the effect in the population models will also be near zero. We report the posterior mean, standard deviation, and 95% Bayesian Credible Intervals for every parameter of each model for each species in Appendix B.

Statistical regularization: Bayesian ridge regression Our objective is to model 248 the response of our focal grass species to interannual variation in climate, even if those 249 responses are weak. Therefore, we avoid selecting among models with all possible com-250 binations of climate covariates, and instead use Bayesian ridge regression to regulate, or 251 constrain, the posterior distributions of each climate covariate (Hooten and Hobbs 2015). 252 Ridge regression is a specific application of statistical regularization that seeks to opti-253 mize model generality by trading off bias and variance. As the name implies, statistical 254 regularization involves the use of a regulator that constrains an optimization. The natural 255 regulator in a Bayesian application is the prior on the coefficient of interest. Each of our 256 statistical models includes the effects of climate covariates via the term $\mathbf{z}_t'\boldsymbol{\beta}_c$ with prior 257 $\boldsymbol{\beta}_c \sim \text{Normal}(\boldsymbol{\mu}_{\beta_c}, \sigma_{\beta_c}^2)$. Since we standardized all climate covariates to have mean zero 258 and variance one, we can set $\mu_{\beta_c} = 0$ and let $\sigma_{\beta_c}^2$ serve as the regulator that can shrink 259 covariates toward zero – the smaller the prior variance, the more the posteriors of β_c are 260 shrunk toward zero, and the stronger the penalty (Hooten and Hobbs 2015). 261 To find the optimal penalty (i.e., optimal value of the hyperparameter $\sigma_{\beta_c}^2$), we fit each 262 statistical model with a range of values for $\sigma_{\beta_c}^2$ and compared predictive scores from leave-263 one-year-out cross-validation. We performed the grid search over 24 values of $\sigma_{\beta_c}^2$, ranging 264 from $\sigma_{\beta_c}^2 = 0.01$ to $\sigma_{\beta_c}^2 = 2.25$. For each statistical model and each species we fit 13×10^{-2} 265 24 = 312 models (13 years to leave out for cross-validation and 24 values of $\sigma_{\beta_c}^2$) – a total 266 of 4,992 models. We calculated the log pointwise predictive density (lppd) to score each 267

model's ability to predict the left-out data. Thus, for training data y_{train} and held-out data y_{hold} at a given value of σ_{θ}^2 across all MCMC samples s = 1, 2, ..., S and all hold outs of data from year t to year T, and letting θ represent all unknowns, lppd is

$$lppd_{CV} = \sum_{t=1}^{T} log_e \int [y_{t,hold}|\theta] [\theta|y_{train}] d\theta,$$
(11)

272 and computed as

$$\sum_{t=1}^{T} \log_e \left(\frac{1}{S} \sum_{s=1}^{S} (y_{t,hold} | \theta_{ts}) \right). \tag{12}$$

We chose the optimal prior variance for each species-statistical model combination as the one that produced the highest lppd and then fit each species-statistical model combination using the full data set for each species and the optimal prior variance.

277 Population models

With the posterior distribution of the vital rate statistical models in hand, it is straightforward to simulate the population models. We used an Integral Projection Model (IPM) to
model populations based on individual-level data (Ellner and Rees 2006) and a quadratbased version of an individually-based model (Quadrat-Based Model, QBM) to model
populations based on quadrat-level data. We describe each in turn.

Integral projection model We use a stochastic IPM (Rees and Ellner 2009) that includes the climate covariates from the vital rate statistical models. In all simulations we ignore the random year effects so that interannual variation is driven solely by climate. We fit the random year effects in the vital rate regressions to avoid over-attributing variation to climate covariates. Our IPM follows the specification of Chu and Adler (2015) where the population of species j is a density function $n(u_j, t)$ giving the density of sized-u genets at time t. Genet size is on the natural log scale, so that $n(u_j, t)du$ is the number of genets

whose area (on the arithmetic scale) is between e^{u_j} and e^{u_j+du} . The density function for any size v at time t+1 is 291

$$n(v_j, t+1) = \int_{L_j}^{U_j} k_j(v_j, u_j, \bar{w}_j(u_j)) n(u_j, t)$$
(13)

where $k_j(v_j, u_j, \bar{w}_j)$ is the population kernel that describes all possible transitions from size u to v and \bar{w}_i is a scalar representing the average intraspecific crowding experienced by 294 a genet of size u_j and species j. The integral is evaluated over all possible sizes between 295 predefined lower (L) and upper (U) size limits that extend beyond the range of observed 296 genet sizes. 297 Since the IPM is spatially-implicit, we cannot calculate neighborhood crowding for spe-298 cific genets (w_{ij}) . Instead, we use an approximation (\bar{w}_i) that captures the essential fea-299 tures of neighborhood interactions (Adler et al. 2010). This approximation relies on a 'no-overlap' rule for conspecific genets to approximate the overdispersion of large genets in space (Adler et al. 2010). The population kernel is defined as the joint contributions of survival (S), growth (G), and 303 recruitment (R):

$$k_j(v_j, u_j, \bar{w}_j) = S_j(u_j, \bar{w}_j(u_j))G_j(v_j, u_j, \bar{w}_j(u_j)) + R_j(v_j, u_j, \bar{w}_j),$$
(14)

which means we are calculating growth (G) for individuals that survive (S) from time t to t+1 and adding in newly recruited (R) individuals of an average sized one-year-old 307 genet for the focal species. Our stastical model for recruitment (R, described above) re-308 turns the number of new recruit produced per quadrat. Following previous work (Adler 309 et al. 2012, Chu and Adler 2015), we assume that fecundity increases linearly with size 310 $(R_j(v_j,u_j,\bar{w}_j)=e^{u_j}R_j(v_j,\bar{w}_j))$ to incorporate the recruitment function in the spatially-311 implicit IPM. 312

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We used random draws from the final 1,000 iterations from each of three MCMC chains

to introduce stochasticity into our population models. At each time step, we randomly selected climate covariates from one of the 14 observed years. Then, we drew the full pa-315 rameter set (climate effects and density-dependence fixed effects) from a randomly se-316 lected MCMC iteration. Using this approach, rather than simply using coefficient point 317 estimates, captures the effect of parameter uncertainty. Relatively unimportant climate 318 covariates (those that broadly overlap 0) will have little effect on the mean of the simula-319 tion results, but can contribute to their variation. Since our focus was on the contribution 320 of climate covariates to population states, we set the random year effects and the random 321 group effects to zero. 322

Quadrat-level statistical model (Eqs. 9-10). We use the same approach for drawing parameter values as described for the IPM. After drawing the appropriate parameter set, we calculate the mean response (log cover at $t+1 = \mu_{t+1}$) according to Eq. 10. We then make a random draw from a [0,1] truncated lognormal distribution with mean equal to μ_{t+1} from Eq. 10 and the variance estimate from the fitted model. We can then project the model forward by drawing a new parameter set (unique to climate year and MCMC iteration) at each timestep. As with the IPM, random year effects are ignored for all simulations.

331 Model validation

To test each model's ability to forecast population state, we made out-of-sample predictions using leave-one-year-out cross validation. For both levels of modeling, we fit the vital rate models using observations from all years except one, and then used those fitted parameters in the population models to perform a one-step-ahead forecast for the year whose observations were withheld from model fitting. Within each observation year, several quadrats were sampled. We made predictions for each observed quadrat in the focal year, initializing each simulation with cover in the quadrat the previous year. Since we

were making quadrat-specific predictions, we incorporated the group random effect on the intercept for both models. We repeated this procedure for all 13 observation years, making 100 one-step-ahead forecasts for each quadrat-year combination with parameter uncer-341 tainty included via random draw from the MCMC chain as described above. Random year 342 effects were set to zero since year effects cannot be assigned to unobserved years. 343 This cross-validation procedure allowed us to compare accuracy and precision of the two 344 modeling approaches (IPM versus QBM). We first calculated the median predicted cover 345 across the 100 simulations for each quadrat-year and then calculated the absolute error as 346 the absolute value of the difference between the observed cover for a given quadrat-year 347 and the median prediction. To arrive at mean absolute error (MAE), we then averaged the 348 absolute error within each species across the quadrat-year specific errors. We use MAE as our measure of accuracy. To measure precision we calculated the distance between the upper and lower 90th quantiles of the 100 predictions and averaged this value over 351 quadrat-years for each species. 352

353 Testing sensitivity to climate covariates

With our fitted and validated models in hand, we ran simulations for each model type

(IPM and QBM) under four climate perturbation scenarios: (1) observed climate, (2) pre
cipitation increased by 1%, (3) temperature increased by 1%, and (4) precipitation and

temperature increased by 1%. We ran the simulations for 2,500 time steps, enough to

estimate equilibrium cover after discarding an initial 500 time steps as burn-in. Each simu
lation was run under two parameter scenarios: (1) using mean parameter estimates and (2)

using randomly drawn parameters from the MCMC chain. We use (1) to detect the overall

sensitivity of equilibrium cover to climate, and we use (2) to show the impact of model

and parameter uncertainty on forecast precision.

ran simulations designed to quantify the sensitivities of individual and combined vital rates to climate for the IPM. Specifically, we ran simulations for the above climate scenar-365 ios, but applied the perturbed climate covariates to survival, growth, or recruitment vital 366 rates individually and in pairwise combinations. This allowed us to isolate the vital rate(s) 367 most sensitive to climate. For this analysis, we used mean parameter estimates to reduce 368 the sources of uncertainty in the sensitivity estimates. 360 We expected the IPM to produce more accurate and precise forecasts due to either (1) 370 the smaller sample size of the quadrat level data sets compared to the individual level 371 data sets, leading to larger parameter uncertainty for the QBM, or (2) the QBM climate 372 effects being weakly associated with one or more vital rate climate effects at the individual 373 level. To assess the impact of sample size on QBM parameter uncertainty we refit the 374 QBM statistical model (Eqs. 9-10) after removing sets of 2, 5, 10, and 15 quadrats. We 375 fit 10 models at each level of quadrat removal (2, 5, 10, 15 quadrats), removing a different randomly selected set of quadrats for each fit. We calculated the standard deviation of 377 climate main effects (pptLag, ppt1, ppt2, TmeanSpr1, and TmeanSpr2) for each model and averaged those over replicates within each set of quadrat removals. This allowed us to 379 regress parameter uncertainty against sample size. 380 To deterime if the QBM climate effects are correlated with climate effects for each vital 381 rate model in the IPM, we simply regressed the QBM climate coefficients against each 382 vital rate model's climate coefficients and calculated Pearson's ρ . Strong correlations indi-383 cate the QBM is capable of detecting climate effects associated with individual vital rates. 384 A weak correlation indicates the QBM "misses" the climate effect on a particular vital 385 rate. 386

Results

388 Comparison of forecast models

389 Sensitivity of models to climate

The response of a population to climate change is a result of the aggregate effects of cli-390 mate on individual vital rates. Since the IPM approach relies on vital rate regressions, we 391 were able to quantify the sensitivity of each vital rate in isolation and in pairwise combi-392 nations. Across all species, climate covariates can have opposing effects on different vital rates (Fig. 3). Growth was the most sensitive vital rate for all species, showing a negative response to increased precipitation, and stronger positive response to increased tempera-395 ture, and a mostly positive response when both climate factors are increased (Fig. 3). B. 396 qracilis survival rates were sensitive to temperature, resulting in an increase in plant cover 397 under increased temperature (Fig. 3a). In isolation, recruitment and survival were insen-398 sitive to climate factors for H. comata (Fig. 3b). Survival and recruitment of P. smithii 399 were both sensitive, negatively, to temperature and precipitation (Fig. 3c). P. secunda 400 equilibrium cover was sensitive to the climate effects on survival and recruitment, showing 401 a negative effect on both vital rates for increased precipition, but a strong positive effect 402 on survival with increased temperature (Fig. 3d). Equilibrium cover responded negatively 403 when increased precipitation and temperature affect recruitment (Fig. 3d). At least two of 404 three vital rates were sensitive to climate for each species (Fig. 3). 405

406 Sources of uncertainty in the QBM

Sample size had a relatively weak effect on QBM climate parameter uncertainty after the number of quadrats used in fitting exceeded about 10 (Fig. 5). Inverse-gaussian fits show that increasing sample size beyond the number of quadrats we used would result in diminishing returns in terms of parameter certainty (Fig. 5). Climate effects estimated from the QBM are most correlated with climate effects from
the growth regression at the individual level (Fig. 6). In no case does the QBM statistical
model have strong correlations across all three vital rates (Fig. 6). QBM climate effects
were most weakly correlated with those from individual-level recruitment models for B.

gracilis, H. comata, and P. secunda (Fig. 6a,b,d). For P. smithii, QBM climate effects
showed no correlation with the survival model effects (Fig. 6c).

417 Model forecasts

Forecasts based on 1% climate changes were extremely uncertain when we considered 418 model error and parameter uncertainty (Fig. 6; simulations with mean parameters are in 419 Appendix D for comparison). As expected based on model validation (Table 1), QBM pro-420 jections were more uncertain than IPM projections for all species except P. smithiii (Fig. 421 6). IPM forecasts for P. smithiii were very uncertain due to a very high instrinsic rate of 422 recruitment combined with uncertainty in climate coefficients which lead to high recruit-423 ment boom years and subsequent busts when young plants suffer high mortality (Appendx 424 C). When we included model error and parameter uncertainty, forecast changes in propor-425 tional cover always spanned a wide range of negative to positive values. In other words, 426 neither model could predict whether a climate perturbation would increase or decrease 427 equilibrium population size. 428

429 Discussion

Population models built using individual-level data allow inference on demographic processes, but they can only forecast future population states across the (typically limited)
spatial extent of the observations. Population-level data are much easier to collect across
broad spatial extents, so models built using such data offer an appealing alternative to
traditional population models (Queenborough et al. 2011). However, density-structured

models rely on the aggregation of individual-level data. This creates a potential problem if
such models are to be used in a climate change context because it is individuals, not populations, which respond to climate (Clark et al. 2012). Are models based on populationlevel metrics as sensitive to climate as models based on individual-level metrics? Do these
two types of models produce consistent forecasts? Do we need detailed demographic data
to forecast the impacts of climate change?

The importance of demographic data

Our comparison of a traditional, demographic population model (the IPM) with a model 442 inspired by density-structured models (the QBM) showed that the IPM outperformed 443 the QBM: the IPM was more accurate and precise than the QBM in out-of-sample cross 444 validation (Table 1). The superiority of the IPM could reflect either differences in sample 445 size or the effect of averaging over unique effects of climate on each individual-level vital 446 rate. Although increasing sample size of quadrat percent-cover observations would be easy 447 to do in the field, we found little evidence that it would lead to higher precision of climate 448 coefficient estimates (Fig. 4). 449 We did, however, find evidence that the QBM statistical model failed to identify climate dependence for some vital rates (Fig. 5). For no species were climate effects from the 451 QBM strongly correlated with all three vital rates (Fig. 5). Freckleton et al. (2011) ac-452 knowledge that averaging over complex stage dependence will lead to poorly specified models. This is analogous to our situation, but instead of averaging over complex life his-454 tories, we are averaging over complex climate dependence. Though our work here focused 455 on plant species, this finding is applicable to any species with vital rates that respond 456 uniquely to weather/climate. 457 Our interpretation is that the QBM is "missing" climate signals associated with at least 458 one vital rate for each species. This leads to inaccurate and imprecise forecasts because 459

the QBM statistical model struggles to explain variation due to climate variables that have positive and negative impacts on different vital rates. When this is the case, as it is for all 461 our species to varying degrees (Fig. 3), forecasts from models based on population-level 462 data will fail. Our result is consistent with related work on the importance of individual-463 level data to forecast population responses to exogenous drivers (Clark et al. 2011a, 2011b, 464 2012, Galván et al. 2014). 465 Detailed demographic data appears to be necessary to forecast climate change impacts on 466 plant populations when vital rates have unique climate responses. How then can we build 467 models to make forecasts for the landscape and regional scales beyond the scope of tradi-468 tional population models (Queenborough et al. 2011)? There are alternatives to density-469 structured models. For example, Clark et al. (2011a) use Forest Inventory and Analysis (FIA) data to parameterize a population model with multiple vital rates and climate de-471 pendence. Distributed efforts such as PlantPopNet (http://plantago.plantpopnet.com) will allow researchers to estimate variation around climate responses for widespread species by taking advantage of spatial variation in climate (e.g. Doak and Morris 2010). Finally, new approaches on the horizon that leverage photo/video of plots and advanced object 475 recognition algorithms (e.g. Liu et al. 2014) will increase the efficiency of plant mapping 476 and digitizing efforts. 477

The challenge of uncertainty

An important, but unexpected, result of our analysis was the great uncertainty in forecasts, even for our best model. The typical approach in ecology is to use point estimates
of model parameters to project populations forward according to the specified model, usually allowing for some variability around the determinstic process (e.g. Battin et al. 2007,
Jenouvrier et al. 2009, Adler et al. 2012). If we follow tradition and calculate the mean
response to climate perturbation with only model error and interannual variation included,

the IPM and the QBM produce opposing forecasts for three of four species (Fig. D1). It would be tempting to interpret this inconsistency as further evidence for the superiority of 486 the IPM. However, if we introduce parameter uncertainty, the forecasts are actually indis-487 tinguishable (Fig. 6), though the IPM projections are generally more precise (consistent 488 with our cross-validation results). The real story is that both models produce highly un-489 certain forecasts. For all species, the 90% quantiles of predicted changes in population size 490 overlapped zero; we cannot even predict whether a change in precipitation or temperature 491 will cause populations to increase or decrease. This result held when we tried perturbing 492 climate by 10% and 20% as well. 493 Our results highlight the state of affairs in ecology when it comes to forecasting the im-494 pacts of climate change. The analysis we conducted here could be considered at the fore-495 front of ecological forecasting with respect to the statistical approach employed (hierarchical Bayesian), the type of population model we used (density-dependent, stochastic IPM with parameter uncertainty), and the amount of high quality data we had at our disposal 498 (14 years of individual-level data). Yet, model predictions proved so uncertain that any forecast, when bounded with model and parameter uncertainty, would be uninformative. 500 How might we improve on this state of affairs? First, forecasts could be improved by 501 matching the spatial scale of predictor variables with the spatial scale of observations. One 502 of the major limitations of the models we fit here is that the climate data are collected 503 at a larger scale than the individual-level observations of plant size. Climate covariates 504 only vary by year, with no spatial variability within years. Thus, even if we fit models 505 to individual-level data, we are missing the key interaction point between weather and 506 individual plants (Clark et al. 2011b) because all observations share the same climate co-507 variates. Demographic studies should be designed with at least plot-level measurements of climate related variables (e.g., soil moisture). Second, accurately detecting climate signals will take even longer time series. Recent theoretical work on detecting climate signals in noisy data suggests that even advanced approaches to parameter fitting require 20-25 year

time series (Teller et al. 2016). Third, ecologists need a stronger commitment to reporting uncertainty. Although most modeling studies explicitly consider model uncertainty, param-513 eter uncertainty is often ignored. In some cases this is because the most convenient sta-514 tistical methods make it difficult to propagate parameter uncertainty. Yet even Bayesian 515 approaches that allow integration of model fitting and forecasting (Hobbs and Hooten 516 2015) are not simple when using modeling approaches like integral projection models that 517 separate the model fitting and simulation stages (Rees and Ellner 2009). However, as we 518 have done here, it is still possible to include parameter uncertainty by drawing parameter 519 values from MCMC iterations, taking care to draw all parameters from the same chain 520 and iteration to account for their correlations. Only by being honest about our forecasts 521 can we begin to produce better ones, and forecasts reported without parameter error are 522 disingenuous. Ignoring parameter error may be justifiable when the goal is investigating 523 basic processes, but it is indefensible when forecasting is the goal. 524

525 Conclusions

This work is not a critique of density-structured population models. We are confident that 526 density-structured models will prove to be a valuable tool for many applications. However, 527 our analysis represents the first comparison, to our knowledge, of population models based on individual and aggregated forms of the same data in a climate change context. Our results confirm theoretical arguments (Clark et al. 2011b) and empirical evidence (Clark et al. 2011a, 2012) that individual responses are critical for predicting species' responses 531 to climate change. It seems there is no short cut to producing accurate and precise pop-532 ulation forecasts: we need detailed demographic data to forecast the impacts of climate 533 change on populations. Given the importance of demographic data and its current collec-534 tion cost, we need modern methods to collect demographic data more efficiently across 535 environmental gradients in space and time.

Our results also offer a cautionary tale because forecast uncertainty was large for both model types. Even with 14 years of detailed demographic data and sophisticated modeling techniques, our projections contained too much uncertainty to be informative. Uncertainty in demographic responses to climate can be reduced by collecting (1) longer time series and (2) climate covariates that match the scale of inference (e.g., plot rather than landscape level climate/weather metrics).

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Tables Tables

Table 1: Accuracy (mean absolute error, MAE) and precision (90% Distance) of out of sample predictions. Forecasts were made without random year effects; only climate covariates could explain year-to-year variation. 90% Distance refers to the average distance between the upper and lower 90th percentiles of the 100 predicted values for each quadrat-year combination.

Species	Model	MAE	90% Distance	Mean Obs. Cover
BOGR	IPM	4.03	24.74	9.43
BOGR	QBM	2.80	29.46	8.35
HECO	IPM	0.34	1.87	1.15
HECO	QBM	0.29	5.53	1.18
PASM	IPM	0.14	0.50	0.42
PASM	QBM	0.17	1.57	0.42
POSE	IPM	0.45	1.63	1.25
POSE	QBM	0.41	3.51	1.27

554 Figure Legends

Figure 1. Time series of average percent cover over all quadrats for our four focal species: 555 Bouteloua gracilis (BOGR), Hesperostipa comata (HECO), Pascopyrum smithii (PASM), 556 and Poa secunda (POSE). Light grey lines show trajectories of individual quadrats. Note 557 the different y-axis scales across panels. 558 Figure 2. Work flow of the data aggregation, model fitting, and population simulating. Figure 3. Sensitivity of equilibrium cover simulated from the IPM to each climate scenario 560 applied to individual and combined vital rates. For example, the points associated with G 561 show the median cover from IPM simulations where a climate perturbation is applied only 562 to the growth regression climate covariates. These simulations use mean parameter values 563 for clarity. 564 Figure 4. Effect of quadrat sample size on the precision (standard deviation) of main cli-565 mate effect estimates in the QBM. Increasing the number of quadrats results in diminish-566 ing returns in terms of parameter certainty. Light dashed lines show individual climate 567 effects at five quadrat sample sizes. Thick dark lines are inverse gaussian fits showing the 568 mean effect of increasing quadrat sample size on parameter precision. 560 Figure 5. Correlations (r) between QBM and IPM estimates of climate effects. We ignore 570 sizeXclimate interactions since these are not directly comparable across model types. The 571 QBM does not have multiple vital rates, so its values are repeated across panels within 572 each species. Across top panels, 'growth' = growth regression, 'rec' = recruitment regres-573 sion, 'surv' = survival regression. 574 Figure 6. Mean (points) and 90% quantiles (errorbars) for the proportional difference 575 between baseline simulations (using observed climate) and the climate pertubation simula-576 tion on the x-axis. We calculated proportional difference as log(perturbed climate cover) 577 - log(observed climate cover), where 'perturbed' and 'observed' refer to the climate time

series used to drive interannual variation in the simulations. Model error and parameter

 $_{580}$ uncertainty were propagated through the simulation phase. Climate simulations are as in $_{581}$ Figure 3.

Figures 582

Figure 1: Time series of average percent cover over all quadrats for our four focal species: Bouteloua gracilis (BOGR), Hesperostipa comata (HECO), Pascopyrum smithii (PASM), and Poa secunda (POSE). Light grey lines show trajectories of individual quadrats. Note the different y-axis scales across panels.



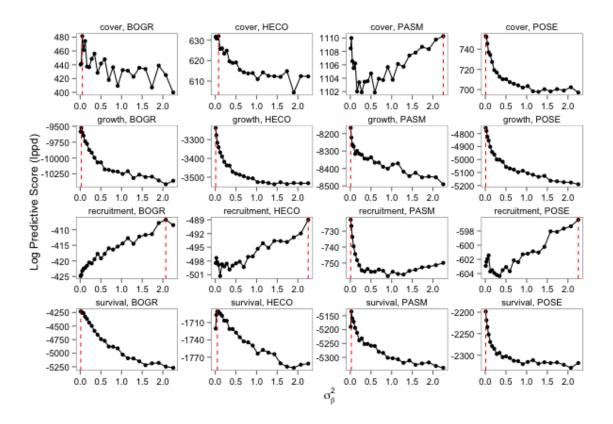
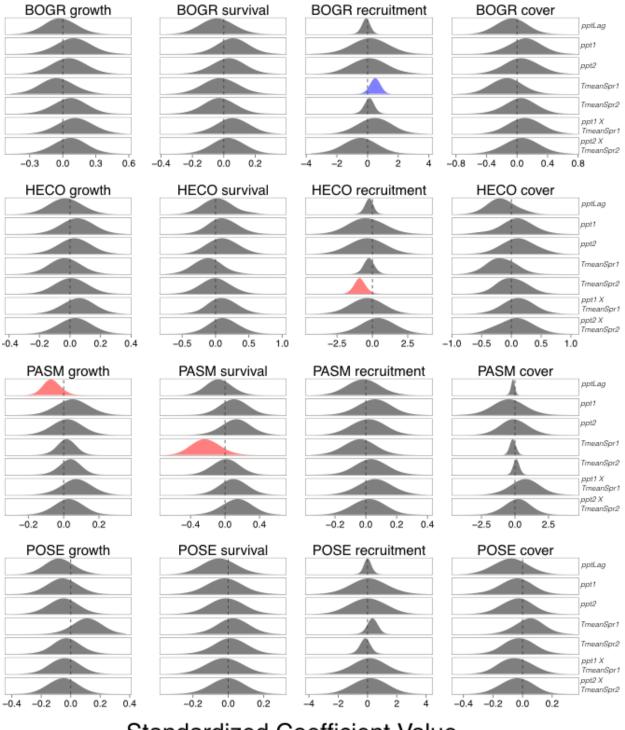


Figure 3: Cross validation scores (summed log pointwise predictive density) plotted against the prior variance for β_C . The optimal score for prediction for each species-vital rate combination is shown with a vertical red line.



Standardized Coefficient Value

Figure 4: Posterior distributions of climate effects (β_C) for each species and vital rate statistical model. Since our priors were constrained via ridge-regression, we highlight climate effects whose 80% credible intervals do not overlap zero (red for negative coefficients, blue for positive coefficients).

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