











The influence of sea surface temperature on tropical-cyclone intensity

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Introduction

For the North Atlantic (N. Atl.) basin, it has been shown [1] that the probability distribution of the so-called power-dissipation index (*PDI*, a rough estimation of released energy) is indeed affected by the annual and basin-wide averaged sea surface temperature (SST), displacing towards more extreme values on warm years (high-SST). As the *PDI* integrates (cubic) wind speed over tropical-cyclone (TC) lifetime, it is an open question where the *PDI* increase comes from (higher speed, longer lifetime, or both).

Data

To characterise a TC one needs to define a physically relevant measure of released energy. The released energy of each TC is summarised as

$$PDI = \sum_{t} v_t^3 \Delta t. \tag{1}$$

The raw hurricane best track data (HURDAT2) is provided by the National Hurricane Center. We intentionally limit this study to the satellite era (1966–2016), as it is the most reliable.

Then, the hurricane observational data is classified into occurrences in low-SST and high-SST years depending on whether they are lower or greater than

$$\langle SST \rangle = \sum_{y \in Y} \frac{SST(y)}{Y},$$
 (2)

where SST(y) is the mean SST of the year y, and Y is the total number of years studied. The SST data (HadISST1) is provided by the Met Office Hadley Centre.

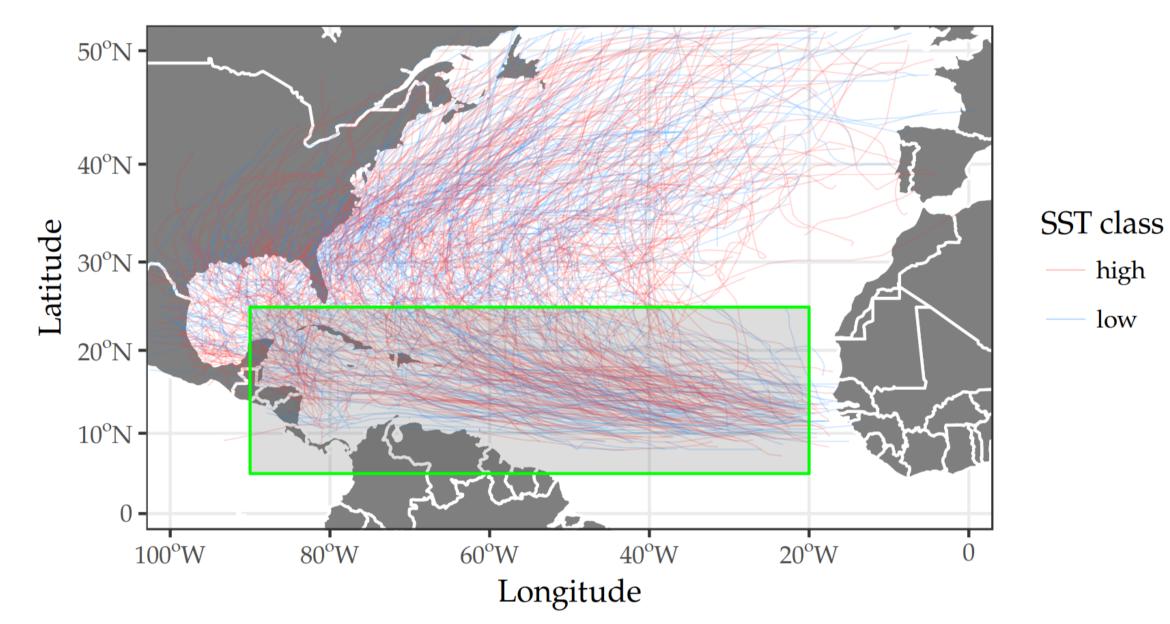


Figure: Tropical-cyclones best tracks for the North Atlantic basin

Hypothesis

The hypothesis is that the SST does not directly affect the maximum wind speed of a TC: storms of equal duration should, in theory, have the same wind speed and *PDI*, and have the same joint distribution:

$$f(Y \mid X = x)_{\text{low}} = f(Y \mid X = x)_{\text{high}}.$$
 (3)

The physical reasoning behind this is that once the cyclone is activated, the wind speed should not depend on its underlying SST.

Instead of working with the exact joint distributions f, we study the expected value of the distributions:

$$E(Y \mid X = x)_{\text{low}} = E(Y \mid X = x)_{\text{high}}, \tag{4}$$

where $E(Y \mid X = x)$ is estimated by performing a ordinary least squares (OLS) regression analysis on the data sets.

Data shows

Our empirical results show a remarkable correlation in the joint distribution of lifetime and wind speed, as seen in Figure 2. The marginal distributions can be seen in Figure 3.

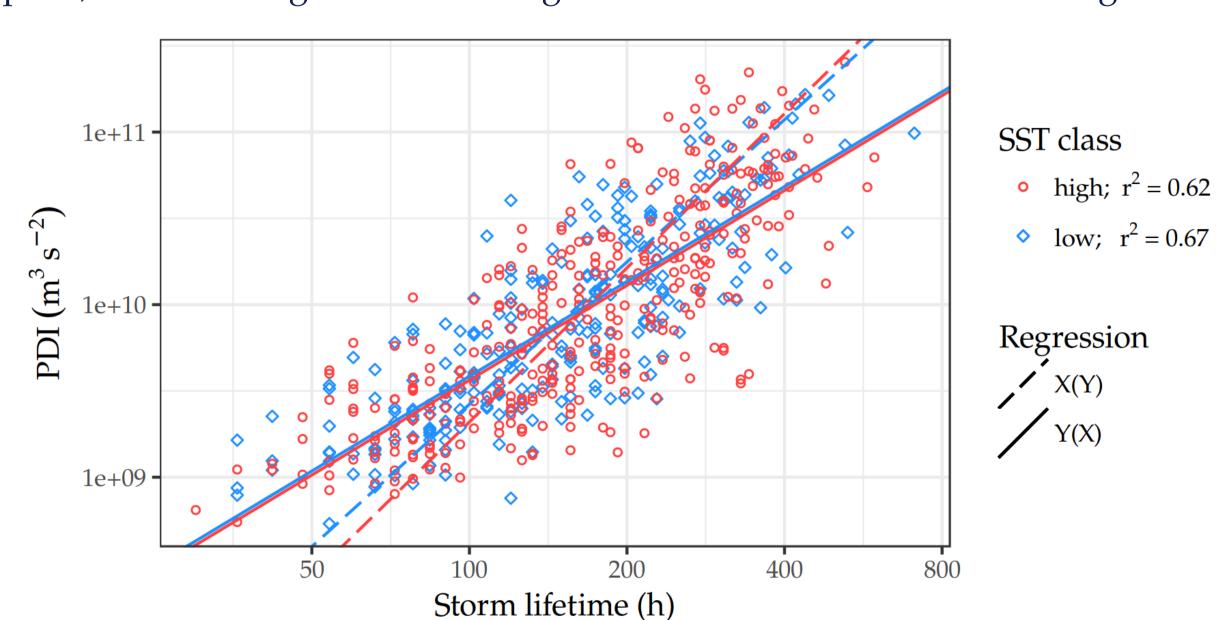


Figure: *PDI* vs lifetime regression analysis for tropical-cyclones (excluding tropical depressions) on the North Atlantic basin

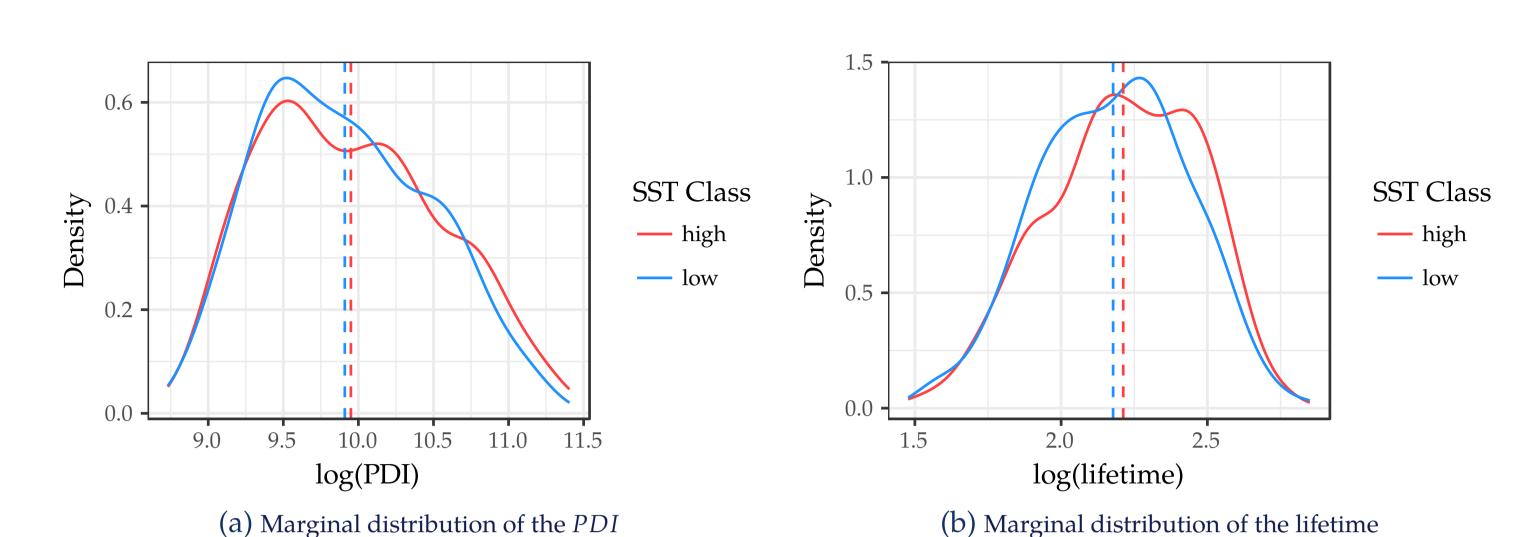


Figure: Marginal analysis for the variables of the bivariate lognormal distribution for the North Atlantic basin data

It is important to notice that the relationships between the TC variables are of non-linear nature. This naturally means that our regressions need to follow a so-called log-log model:

$$\log \Psi = \alpha + \beta \log \Phi + \epsilon, \tag{5}$$

where $\log \Psi \equiv Y$ and $\log \Phi \equiv X$.

Methodology

In essence, our methodology will consist in comparing the low-SST and high-SST regression models using non-parametric permutation tests.

For the permutation tests, we use bootstrap in order to get a more robust estimation of the value of the different studied statistical coefficients (such as $\hat{\beta}_0$, $\hat{\beta}_1$, R^2) and their standard errors than the one obtained using the value obtained using the OLS method. Withal, we compare the results obtained using both methods.

The following null hypothesis is proposed to test the data:

$$H_0: \hat{\beta}_{0,h} = \hat{\beta}_{0,l} \wedge \hat{\beta}_{1,h} = \hat{\beta}_{1,l}. \tag{6}$$

We use the following statistics:

$$T^{(1)} = |\hat{\beta}_{0,h} - \hat{\beta}_{0,l}|, \quad T^{(2)} = |\hat{\beta}_{1,h} - \hat{\beta}_{1,l}|, \quad T^{(3)} = |R_h^2 - R_l^2|. \tag{7}$$

Polko-Zając [3] proposes alternative statistics that not only consider the nominal value of the coefficient estimates, but take into account their standard errors as well:

$$T^{(4)} = \frac{|\hat{\beta}_{0,h} - \hat{\beta}_{0,l}|}{\widehat{\operatorname{se}}(\hat{\beta}_{0,h} - \hat{\beta}_{0,l})}, \quad T^{(5)} = \frac{|\hat{\beta}_{1,h} - \hat{\beta}_{1,l}|}{\widehat{\operatorname{se}}(\hat{\beta}_{1,h} - \hat{\beta}_{1,l})}, \quad T^{(6)} = T^{(4)} + T^{(5)}.$$
(8)

Results

The value of the studied statistics $T^{(i)}$ obtained from the N. Atl. data set can be seen in Table 1.

| X | Υ | $T^{(1)}$ | $T^{(2)}$ | $T^{(3)}$ | $T^{(4)}$ | $T^{(5)}$ | $T^{(6)}$ |
|----------|----------|-----------|-----------|-----------|-----------|-----------|-----------|
| lifetime | PDI | 0.289 | 0.027 | 0.048 | 1.404 | 1.324 | 2.727 |
| PDI | lifetime | 0.007 | 0.007 | 0.054 | 0.031 | 0.065 | 0.095 |

Table: Value of the studied statistics for North Atlantic basin data set

The results of the permutation tests performed on the N. Atl. data can be seen in Table 2.

| X | Υ | $T^{(1)}$ | $T^{(2)}$ | $T^{(3)}$ | $T^{(4)}$ | $T^{(5)}$ | $T^{(6)}$ |
|----------|----------|-----------|-----------|-----------|-----------|-----------|-----------|
| lifetime | PDI | 0.146 | 0.121 | 0.711 | 0.137 | 0.117 | 0.12 |
| PDI | lifetime | 0.870 | 0.806 | 0.705 | 0.864 | 0.795 | 0.83 |

Table: List of *p*-values of the bootstrap-powered permutation test for the North Atlantic basin data

The results show that none of the performed permutation tests is able to reject the null hypothesis H_0 .

Conclusion

Our conclusions are compatible with the view of tropical cyclones as an activation process, in which, once the event has started, its intensity is kept in critical balance between attenuation and intensification (and so, higher SST does not trigger more intensification).

References

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