

Spatio temporal analysis of extreme wind velocities for infrastructure desing. Case  
study Colombia

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A Thesis  
Presented to  
The Division of Instituto for Geoinformatics - IFGI  
University of Münster

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In Partial Fulfillment  
of the Requirements for the Degree  
Master of Science in Geospatial Technologies

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Approved for the Division  
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# Acknowledgements

I want to thank a few people.



# Preface

This is an example of a thesis setup to use the reed thesis document class (for LaTeX) and the R bookdown package, in general.





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# Abstract

The preface pretty much says it all.

Second paragraph of abstract starts here.





# Dedication

You can have a dedication here if you wish.



# Introduction

Placeholder



# Chapter 1

## Data

Placeholder

### 1.1 IDEAM

### 1.2 ISD

### 1.3 ERA5

### 1.4 Data Download and Organization

### 1.5 Data Standardization



# Chapter 2

## Theoretical Framework

### 2.1 Probability Concepts

Poisson process is an stochastic method that relies in the concepts of probability distributions. The main functions related to probability for extreme value analysis will be described below.

#### 2.1.1 Probability Density Function - *pdf*

Pdf defines the probability that a continuous variable falls between two points, this is, in *pdf* the probability is related to the area below the curve (integral) between two points, as for continuous probability distributions the probability at a single point is zero. The term density is directly related to the probability of a portion of the curve, if the density function has high values the probability will be greater in comparison with the same portion of curve for low values.

$$\int_a^b f(x)dx = Pr[a \leq X \leq b]$$

Equation (2.1) is the Gumbel *pdf*.

$$f(x) = \frac{1}{\beta} \exp \left\{ -\frac{x - \mu}{\beta} \right\} \exp \left\{ -\exp \left\{ -\left( \frac{x - \mu}{\beta} \right) \right\} \right\}, \quad -\infty < x < \infty \quad (2.1)$$

where  $\exp \{.\} \mapsto e^{\{.\}}$ ,  $\beta$  is the scale parameter, and  $\mu$  is the location parameter. Location ( $\mu$ ) has the effect to shift the *pdf* to left or right along 'x' axis, thus, if location value is changed the effect is a movement of *pdf* to the left (small value for location), or to the right (big value for location). Scale has the effect to stretch ( $\beta > 1$ ) or compress ( $0 < \beta < 1$ ) the *pdf*, if scale parameter is close to zero the pdf approaches a spike.

Figure 2.1 shows *pdf* with location ( $\mu$ ) = 100 and scale ( $\beta$ ) = 40, using equation (2.1).

```

location = 100
scale = 40
.x <- seq(0, 300, length.out=1000)
pdfG <- function(x) {
  1/location * exp(-(x-location)/scale) * exp(-exp(-(x-location)/scale))
}
.y = pdfG(.x)
plot(.x, .y, col="green", lty=4,
     xlab="Velocities Km/h", ylab="Density Function - Gumbel Distribution",
     main=paste("Gumbel - Density Function Gumbel Distribution\n", "Location=",
               round(location,2), " Scale=", round(scale,2)), type="l",
     cex.axis = 0.5, cex.lab= 0.6, cex.main=0.7, cex.sub=0.6)

```

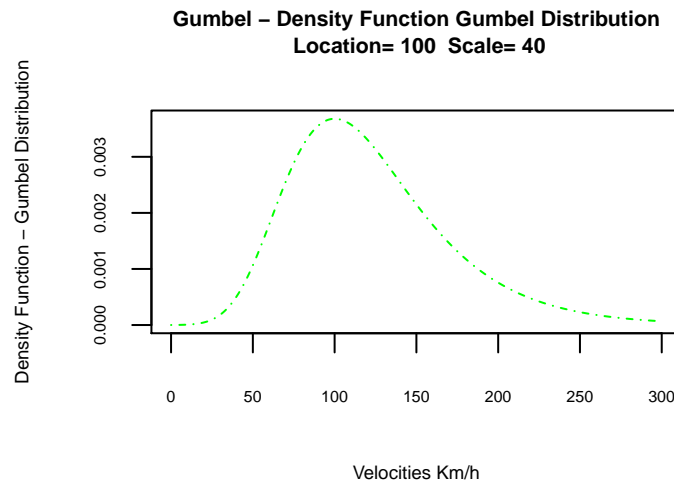


Figure 2.1: Gumbel pdf

Figure 2.2 shows *pdf* with location ( $\mu$ ) = 100 and scale ( $\beta$ ) = 40, using function *dgumbel* of the package *RcmdrMisc*

```

location = 100
scale = 40
.x <- seq(0, 300, length.out=1000)
dfG = dgumbel(.x, location=location, scale=scale)
plot(.x, dfG, col="red", lty=4,
     xlab="Velocities Km/h", ylab="Density Function - Gumbel Distribution",
     main=paste("Gumbel - Density Function Gumbel Distribution\n", "Location=",
               round(location,2), " Scale=", round(scale,2)), type="l",
     cex.axis = 0.5, cex.lab= 0.6, cex.main=0.7, cex.sub=0.6)

```



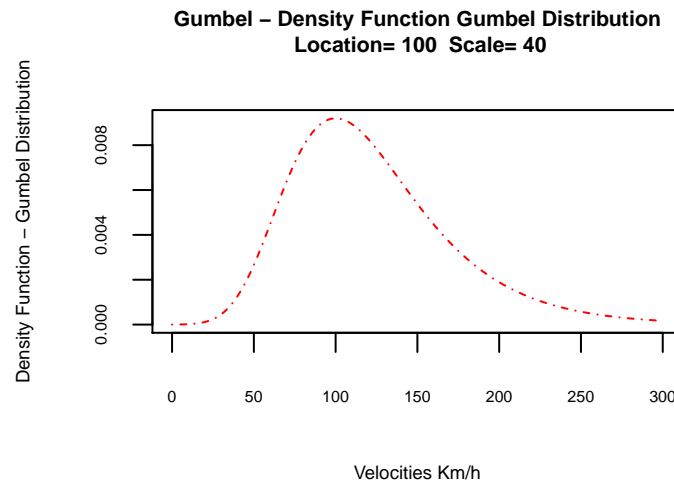


Figure 2.2: Gumbel pdf - dgumbel function

### 2.1.2 Cumulative Distribution Function - *cdf*

*Cdf* is the probability of taking a value less than or equal to  $x$ . That is

$$F(x) = Pr[X \leq x] = \alpha$$

For a continuous variable, *cdf* can be expressed as the integral of its *pdf*.

$$F(x) = \int_{-\infty}^x f(x)dx$$

Equation (2.2) is the Gumbel *cdf*.

$$F(x) = \exp \left\{ -\exp \left[ -\left( \frac{x - \mu}{\beta} \right) \right] \right\}, \quad -\infty < x < \infty \quad (2.2)$$

Figure 2.3 shows Gumbel *cdf* with location ( $\mu$ ) = 100 and scale ( $\beta$ ) = 40, using equation (2.2). As previously done with *pdf*, similar result can be achieved using function `pgumbel` of package `RcmdrMisc`.

```
location = 100
scale = 40
.x <- seq(0, 300, length.out=1000)
cdfG <- function(x) {
  exp(-exp(-(x-location)/scale))
}
.y = cdfG(.x)
plot(.x, .y, col="green", lty=4,
     xlab="Velocities Km/h", ylab="Probability",
     main=paste("Gumbel - Cumulative Distribution Function\n", "Location=",
               round(location,2), " Scale=", round(scale,2)), type="l",
     cex.axis = 0.5, cex.lab= 0.6, cex.main=0.7, cex.sub=0.6)
```

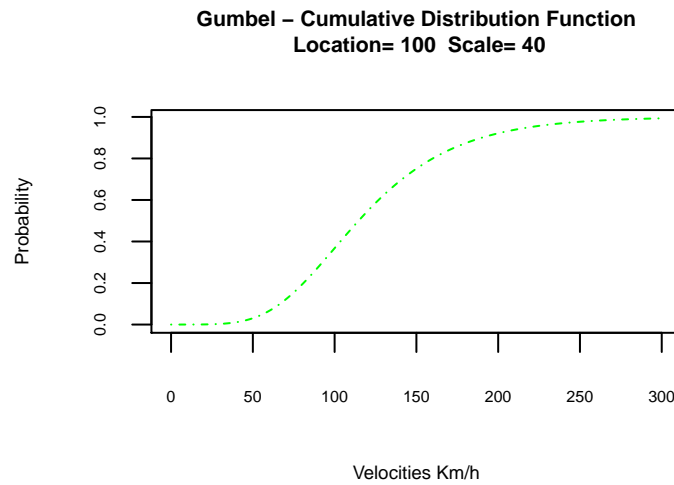


Figure 2.3: Gumbel cdf

### 2.1.3 Percent Point Function - *ppf*

*Ppf* is the inverse of *cdf*, also called the *quantile* function. This is, from a specific probability get the corresponding value  $x$  of the variable.

$$x = G(\alpha) = G(F(x))$$

Equation (2.3) is the Gumbel *ppf*.

$$G(\alpha) = \mu - \beta \ln(-\ln(\alpha)) \quad 0 < \alpha < 1 \quad (2.3)$$

Figure 2.4 shows Gumbel *ppf*, using equation (2.3). Similar result can be achieved using function `qgumbel` of package `RcmdrMisc`.

```
location = 100
scale = 40
.x <- seq(0, 1, length.out=1000)
ppfG <- function(x) {
  location - (scale*log(-log(x)))
}
.y = ppfG(.x)
plot(.x, .y, col="green", lty=4,
     ylab="Velocities Km/h", xlab="Probability",
     main=paste("Gumbel - Percent Point Function\n", "Location=",
               round(location,2), " Scale=", round(scale,2)), type="l",
     cex.axis = 0.5, cex.lab= 0.6, cex.main=0.7, cex.sub=0.6)
```

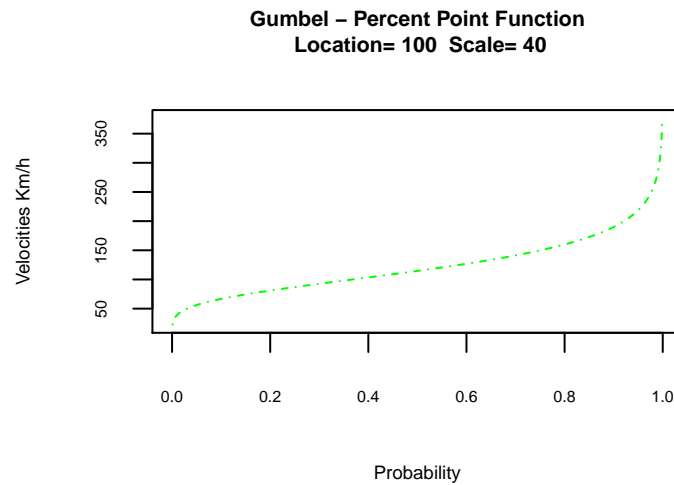


Figure 2.4: Gumbel cdf

### 2.1.4 Hazard Function - $hf$

Using  $S(x) = 1 - F(x)$  as survival function - $sf$ , the probability that a variable takes a value greater than  $x$   $S(x) = Pr[X > x] = 1 - F(x)$ , the  $hf$  is the ratio between  $pdf$  and  $sf$ .

$$h(x) = \frac{f(x)}{S(x)} = \frac{f(x)}{1 - F(x)}$$

Equation (2.4) is the Gumbel  $ppf$ .

$$h(x) = \frac{1}{\beta} \frac{\exp(-(x - \mu)/\beta)}{\exp(\exp(-(x - \mu)/\beta)) - 1} \quad (2.4)$$

Figure 2.5 shows Gumbel  $hf$ , using equation (2.4).

```
location = 100
scale = 40
.x <- seq(0, 3000, length.out=1000)
hfG <- function(x) {
  (1/scale)*(exp(-(x-location)/scale))/(exp(exp(-(x-location)/scale))-1)
}
.y = hfG(.x)
plot(.x, .y, col="green", lty=4,
     xlab="Velocities Km/h", ylab="Hazard",
     main=paste("Gumbel - Hazard Function\n", "Location=",
               round(location,2), " Scale=", round(scale,2)), type="l",
     cex.axis = 0.5, cex.lab= 0.6, cex.main=0.7, cex.sub=0.6)
```

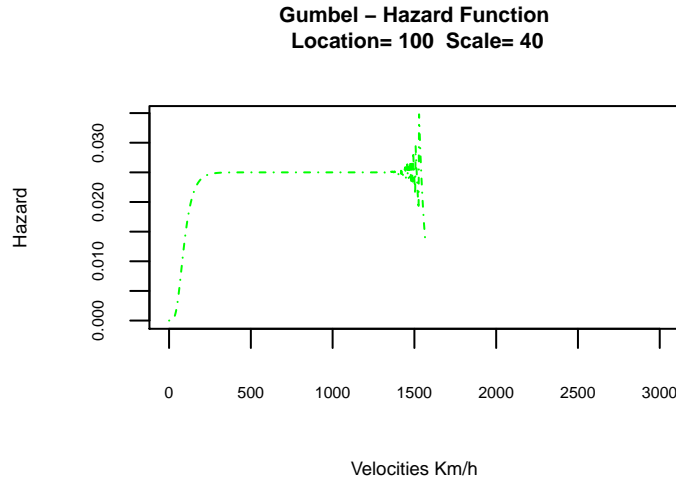


Figure 2.5: Gumbel cdf

```
#library(reliaR)
#plot(.x, hgumbel(.x, mu=location, sigma=scale))
#plot(.x, hra.gumbel(.x, mu=location, sigma=scale))
```

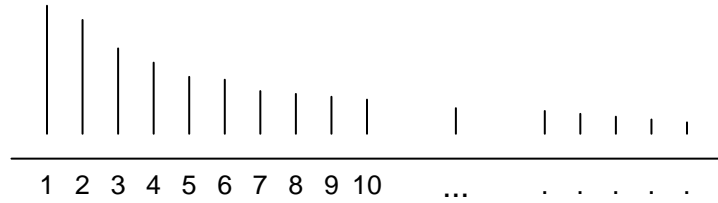
## 2.2 Introductory concepts for statistical analysis of extreme values

In order to approach the extreme value analysis, some statistical concepts are needed to understand the theoretical framework behind this knowledge area. In this section will be introduced the concepts annual exceedance probability, mean recurrence interval - MRI, exposure time, and compound probability for any given exposure time and MRI.

As an hypothetical example, a simulated database of extreme wind speed will be used. This database is supposed to have 10.000 years of simulated wind speeds.

### 2.2.1 Annual Excedance Probability - $P_e$

Using the previously described database, a question arises to calculate the probability to exceed the highest probable loss due to the simulated winds. It is possible to conclude that there is only one event grather or equal (in this case equal) to the highest probable causing loss in 10.000 years, and it is the *highest wind*. If we sort the database by wind magnitude in descending order (small winds last), the question is solved calculating the annual exceedance probability  $P_e$  with next formula



Event Index – Ordered Winds by Magnitud

Figure 2.6: Gumbel cdf

$$P_e = \frac{\text{Event index after descending sorting}}{\text{Years of simulations}} = \frac{1}{10.000} = 0.001 = 0.01\%$$

because the highest wind will be the first in the sorted list. Same exercise can be done with all winds to construct the annual exceedance probability curve, that in this case will represent the probability to equal or exceed different probable losses due to wind.

### 2.2.2 Return Period - Mean Recurrence Interval

Continuing with the previous section, if the inverse of the exceedance probability is taken, the return period is obtained. The return period or Mean Recurrence Interval - MRI.

### 2.2.3 Compound Exceedance Probability - Pn

## 2.3 Extreme Value Analysis Overview

Analysis of extreme values is related with statistical inference to calculate probabilities of extreme events. Main methods to analyze extreme data are epochal, Peaks Over Threshold - POT, and extreme index. The epochal method, also known as block maxima, uses the most extreme value for a specific frame of time, typically, one year. POT is based in the selection of a single threshold value to do the analysis only with values above the threshold. But there are different POT approaches, the most common one is Generalized Pareto Distribution - POT-GPD, but also it is possible to use the Poisson process approach.

In both methods (Epochal and POT), the first step is to fit the data to an appropriate probability distribution model, among them the most used are, - Extreme Value Type I (Gumbel), Extreme Value Type II (Frechet), Weibull, Generalized Pareto - GPD, and Generalized Extreme Value - GEV.

Distribution models are fitted based in the estimation of its parameters, commonly called location, scale and shape, nonetheless each model has its own parameters names. There are different methods to estimate parameters, among them, - method of moments (modified moments - see Kubler (1994), and L moments - see Hosking & Wallis (1997)), - method of maximum likelihood MLE, see Harris & Stocker (1998), which is problematic for GPD and GEV, - probability plot correlation coefficient, and - elemental percentiles (for GPD and GEV)

Once candidate parameters are available, it is necessary to assess the goodness of fit of the selected model, using one of the next methods, - Kolmogorov-Smirnov (KS) goodness of fit test, and - Anderson-Darling goodness of fit test. Here a visual assessment is also useful using a probability plot or a kernel density plot with the fitted *pdf* overlaid.

The main use of the fitted model is the estimation of mean return intervals - MRI, and extreme wind speeds (return levels),

$$MRI = \frac{1}{1 - F(y)}$$

with  $F(y)$  as the *cdf*. If  $1 - F(y)$  is the annual exceedance probability, MRI is its inverse, see Simiu & Scanlan (1996) for more details about MRI. If  $y$  is solved from previous equation using a given MRI of N-years, its value represents the  $Y_N$  wind speed return level,

$$Y_N = G\left(1 - \frac{1}{\lambda N}\right)$$

where  $G$  is the *ppf* (quantile function) and  $\lambda$  is the number of wind speeds over the threshold per year.

The CRAN Task View “Extreme Value Analysis” <https://cran.r-project.org/web/views/ExtremeValue.html> shows available **R** for block maxima, POT by GPD, and external indexes estimation approaches. Most important to consider are `evd`, `extremes`, `evir`, `POT`, `extremeStat`, `ismev`, and `Renext`.

## 2.4 Peaks Over Threshold - Poisson Process

According to Pintar, Simiu, Lombardo, & Levitan (2015) the stochastic poisson process is mainly defined by its intensity function. As the intensity function is not uniform over the domain, the poisson process considered here is non-homogeneous, and due to the intensity function dependence of magnitude and time, it is also bi-dimensional. Poisson Process was described for the first time in Pickands (1971), then extended in

Smith (1989).

$$\lambda(y, t) \begin{cases} \lambda_t(y), & \text{for } t \text{ in thunderstorm period} \\ \lambda_{nt}(y), & \text{for } t \text{ in non - thunderstorm period} \end{cases} \quad (2.5)$$

Generic equation (??) shows the intensity function, which is defined in the domain  $D = D_t \cup D_{nt}$ , and allow to fit the poisson process at each station to the observed data  $\{t_i, y_i\}_{i=1}^I$  for al the times ( $t_i$ ) of threshold crossing observations and its corresponding wind speeds magnitudes ( $y_i$ ). Thus, only data above the threshold is used.

Intensity function of the Poisson Process is defined in Smith (2004),

$$\frac{1}{\psi_t} \left( 1 + \zeta_t \frac{y - \omega_t}{\psi_t} \right)_+^{-\frac{1}{\zeta_t} - 1}$$

Where  $\zeta_t$  controls the tail lengh of the intensity function at a given time  $t$ , but to facilitate the estimation of the parameters then  $\zeta_t$  is taken to be zero, then doing the limit, the resulting intensity function is the same as the the GEV type I or Gumbel distribution,

$$\frac{1}{\psi_t} \exp \left\{ \frac{-(y - \omega_t)}{\psi_t} \right\}$$

In this study, the used intensity functions are shown in ecuation (2.6).

$$\lambda(y, t) \begin{cases} \frac{1}{\psi_s} \exp \left\{ \frac{-(y - \omega_s)}{\psi_s} \right\}, & \text{for } t \text{ in thunderstorm period} \\ \frac{1}{\psi_{nt}} \exp \left\{ \frac{-(y - \omega_{nt})}{\psi_{nt}} \right\}, & \text{for } t \text{ in non - thunderstorm period} \end{cases} \quad (2.6)$$

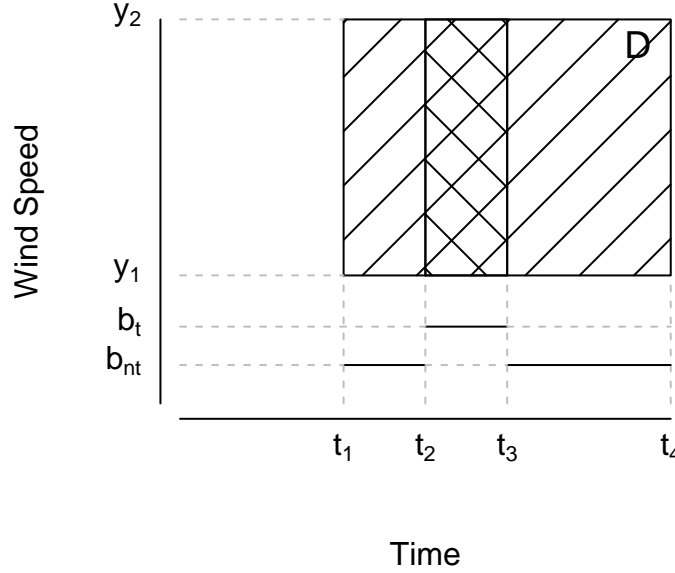


Figure 2.7: Domain of the Poisson Process

Figure 2.7 represent the domain  $D$  of the Poisson process. In time, the domain represents the station service period from first sample  $t_1$  to last sample  $t_4$ .  $D$  is the union of all thunderstorm periods  $D_t$  (from  $t_2$  to  $t_3$ ), and all non-thunderstorm periods  $D_{nt}$  (periods  $t_1$  to  $t_2$  and  $t_3$  to  $t_4$ ). In magnitude, only thunderstorm data above its threshold  $b_t$ , and only non-thunderstorm data above its threshold  $b_{nt}$  are used.

Thunderstorms and non-thunderstorms are modeled independently:

1. Observations in domain  $D$  follow a Poisson distribution with mean  $\int_D \lambda(t, y) dt dy$
2. For each disjoint subdomain  $D_1$  or  $D_2$  inside  $D$ , the observations in  $D_1$  or  $D_2$  are independent random variables.

Visual representation of the intensity function for the Poisson Process can be seen in figure 2.8. In vertical axis, two surfaces were drawn representing independent intensity functions for thunderstorm  $\lambda_t(y)$  and for non-thunderstorm  $\lambda_{nt}(y)$ . The volume under each surface for its corresponding time periods and peak (over threshold) velocities, is the mean of the Poisson Process.



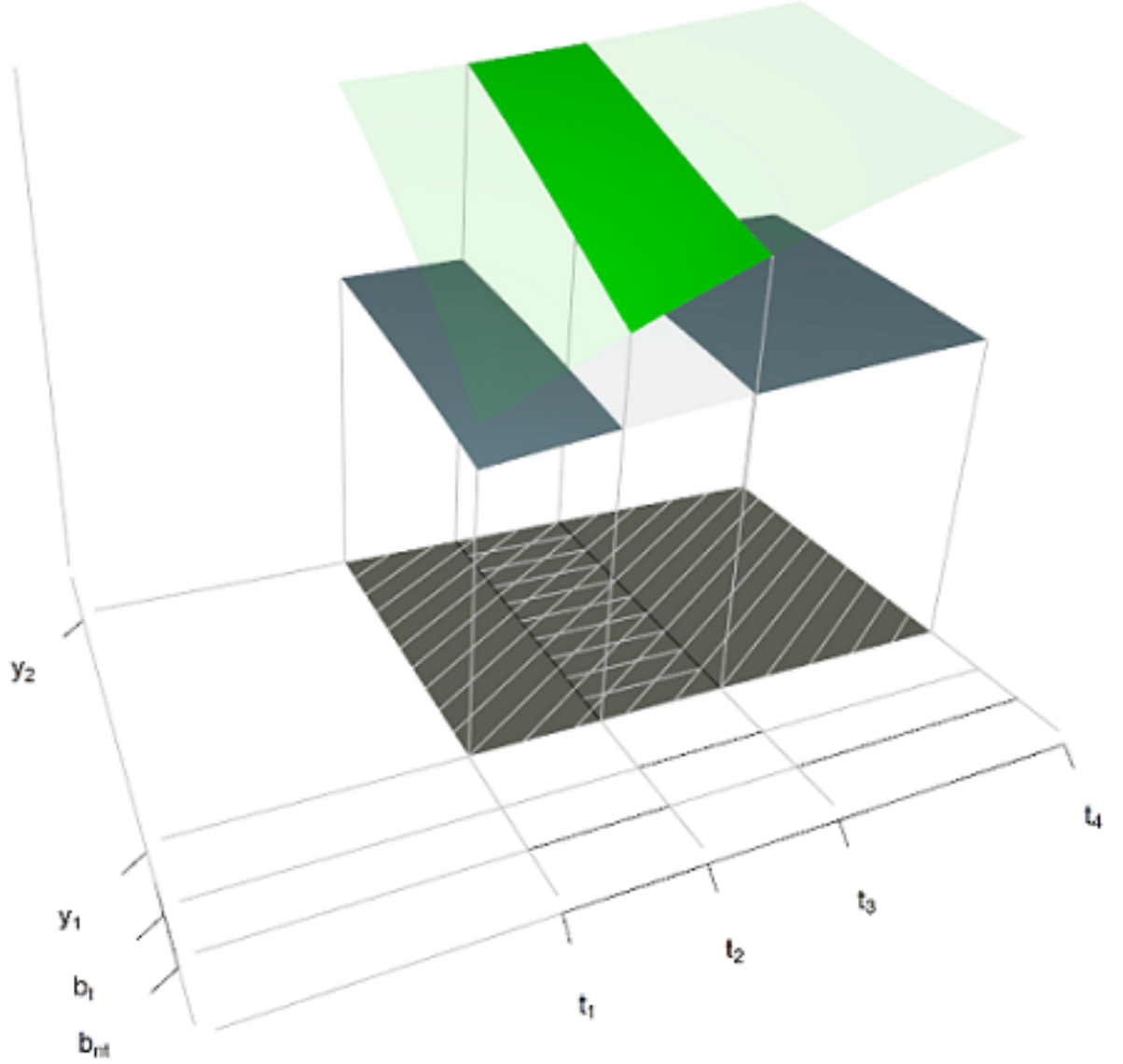


Figure 2.8: Volume under surfaces represents the mean of the Poisson process

The method of maximum likelihood is used to estimate the parameters of the Poisson process, the selected vector of parameters  $\eta$  are the  $\hat{\eta}$  values that maximize the function

$$L(\eta) = \left( \prod_{i=1}^I \lambda(y_i, t_i) \right) \exp \left\{ - \int_D \lambda(y, t) dy dt \right\} \quad (2.7)$$

$\hat{\eta}$  values need to be calculated using a numerical approach because there is not an analytical solution available.

Once the Poisson process is fitted to the data, the model will provide extreme wind velocities (return levels), for different return periods (mean recurrence intervals).

A  $Y_N$  extreme wind velocity, called the return level (RL) belonging to the  $N$ -years return period, has an expected frequency to occur or to be exceeded (annual exceedance

probability)  $P_e = \frac{1}{N}$ , and also has a probability that the event does not occur (annual non-excedance probability)  $P_{ne} = 1 - \frac{1}{N}$ .  $Y_N$  will be the resulting value of the  $G$  (ppf or quantile) function using a probability equal to  $P_{ne}$ .  $Y_N = \text{quantile}(y, p = P_{ne}) = G(x, p = P_{ne}) = \text{ppf}(x, p = P_{ne})$ . As for this study  $\zeta = 0$ , the  $G$  function to use is the Gumbel quantile function.  $Y_N$  can be understood as the wind extreme value expected to be exceeded on average once every  $N$  years.

For different POT approaches, as POT-GPD described –, the value of the probability passed to the  $G$  function, has to be modified with the  $\lambda$  parameter, as is described in next equation.  $\lambda$  is the number of wind speed over the threshold per year.

$$Y_N = G\left(y, 1 - \frac{1}{\lambda N}\right)$$

For the Poisson process  $Y_N$  is also the solution to the next equation, which is defined in terms of the intensity function,

$$\int_{Y_N}^{\infty} \int_0^1 \lambda(y, t) dy dt = A_t \int_{Y_N}^{\infty} \lambda_t(y) dy + A_{nt} \int_{Y_N}^{\infty} \lambda_{nt}(y) dy = \frac{1}{N} \quad (2.8)$$

where  $A_t$ , is the multiplication of the average number of thunderstorm per year and the average length of a thunderstorm (taken to be 1 hour as defined in Pintar et al. (2015)), and  $A_{nt} = 1 - A_t$ . The average length of a non-thunderstorm event is variable, and it is adjusted in each station to guarantee that  $A_{nt} + A_t = 1$

The same thunderstorm event is considered to occur if the time lag distance between sucesive thunderstorm samples is small than six hours, and for non-thunderstorm this time is 4 days. For the Poisson process, all the measurements belonging to the same event (thunderstorm or non tunderstorm), need to be declustered to leave only one maximun value. In other words, the number of thunderstorm in the time serie is the number of time lag distances grather than 6 hours, and for non-thunderstorm grather than 4 days.

###Threshold Selection

$$U = F(Y)$$

$$W = -\log(1 - U)$$

# Chapter 3

## Methodology

Placeholder

## **3.1 Input Data Selection and Standarization**

### **3.1.1 Data Selection**

### **3.1.2 Data Standarization**

Anemometer height - 10 m

Surface Roughness - 0.03 m

Averaging Time - 3-s gust

### **3.1.3 Data Filterng**

## **3.2 Fit data to a POT - Poisson Process**

### **3.2.1 Data Requirements**

### **3.2.2 Exploratory Data Analysis and Data Preparation**

Declustering of observations

Exclude no-data periods

Threshold selection

### **3.2.3 Parameters Estimation**

Intensity function

Density function

Distribution function

Maximun likelihood estimation

### **3.2.4 Velocities at Return Periods**

## **3.3 spatial Interpolation**

# Conclusion

If we don't want Conclusion to have a chapter number next to it, we can add the `{-}` attribute.

## **More info**

And here's some other random info: the first paragraph after a chapter title or section head *shouldn't be* indented, because indents are to tell the reader that you're starting a new paragraph. Since that's obvious after a chapter or section title, proper typesetting doesn't add an indent there.



# Appendix A

## The First Appendix

This first appendix includes all of the R chunks of code that were hidden throughout the document (using the `include = FALSE` chunk tag) to help with readability and/or setup.

**In the main Rmd file**

**In Chapter 3:**





## Appendix B

The Second Appendix, for Fun



# References

Placeholder

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