Isotope system	Maximum Moon[1]	Mars		Ordinary Chondrite (average)		CI-chondrite		References
		Value	> δf_T	Value	> δf_T	Value	> δf_T	
$\delta^{17}O$	0.016‰	0.32‰	-5.0%	~1‰	-1.6%	~0‰[2]	N/A	[7]
$\delta^{30}Si$	-0.34‰	-0.58‰	-59%	-0.59‰	-58%	-0.6‰	-57%	[10], [25]
$\varepsilon^{53}Cr$	~-0.1	~-0.22	-45%	~-0.48	-21%	~-0.43	-23%	[13]
$\varepsilon^{182}W$	~0.19	~0.1 to 3.2[3]	-100% to -5.9%	(~-2)[4]	-	(~-1.9)[4]	-	[14], [27]

[1] Value of the isotopic composition that the moon could have while still being compatible with observations (within 2σ).

[2] CI chondrites plot on the terrestrial fractionation line. Their $\delta^{17}O$ offset from this line is therefore ~0‰. See main text for a discussion of possible implications.

3 Mars shows a large variety of $\varepsilon^{182}W$ values. -100% is for $\varepsilon^{182}W = 0.19$, -5.9% for an impactor with $\varepsilon^{182}W = 3.2$.

[4] An $\varepsilon^{182}W$ value of ~-2 is expected only for undifferentiated chondrites, but not for a differentiated impactor as assumed here. If a lunar composition of $\varepsilon^{182}W = -0.01$ is assumed, a very constraining depletion of -0.5% would result for an undifferentiated chondritic impactor.