

Ocean Dynamics Cheat Sheet

Primitive Equations

$$\begin{cases} \frac{Du}{Dt} - fv = -\frac{1}{\rho} \frac{\partial p}{\partial x} \\ \frac{Dv}{Dt} + fu = -\frac{1}{\rho} \frac{\partial p}{\partial y} \\ \frac{Dw}{Dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g \end{cases}$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{v}) = 0$$

$$\begin{cases} \frac{DT}{Dt} = F_T - Q_T \\ \frac{DS}{Dt} = F_S - Q_S \end{cases} \quad \text{or} \quad \frac{D\rho}{Dt} = \frac{D}{Dt} \left(\rho_0 - \frac{\rho p}{c_s^2} \right) = 0$$

Equation of state: $\rho = \rho_0 \left[1 - \beta_T \Delta T + \beta_S \Delta S + \beta_P \Delta P \right]$

Boussinesq Equations

Continuity: $\nabla \cdot \vec{v} = 0$

Momentum: $\frac{D\vec{u}}{Dt} + \vec{f} \times \vec{u} = -\nabla_h \phi$

$$\frac{\partial \phi}{\partial z} = b$$

$$\rho = \rho_0 + \delta \rho$$

$$\phi = \frac{\delta p}{\rho_0}$$

$$b = -g \frac{\rho - \rho_0}{\rho_0}$$

~~Thermodynamic~~

Equation of state: $b = g \left[\beta_T \Delta T - \beta_S \Delta S - \beta_P \Delta P \right]$
 $= g \left[\beta_T \Delta T - \beta_S \Delta S - \frac{gz}{c_s^2} \right]$

$$c_s^2 = \frac{\partial p}{\partial \rho}$$

Thermodynamic: $\frac{Db}{Dt} = \dot{b}$, or $\frac{\partial b}{\partial t} = 0$

- With $b = \tilde{b}(z) + b'$: $\frac{D\tilde{b}'}{Dt} + N^2 w = 0$ where $N^2 = \frac{\partial \tilde{b}}{\partial z}$

Dispersion relation: $\omega^2 = \frac{(k^2 + l^2) N^2}{k^2 + l^2 + m^2}$

with rotation: $\omega^2 = \frac{N^2(k^2 + l^2) + f^2 m^2}{k^2 + l^2 + m^2}$

Geostrophy

$$\begin{cases} f u_g = -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ f v_g = \frac{1}{\rho_0} \frac{\partial p}{\partial x} \end{cases}$$

Streamfunction: If $f = f_0$, set $\psi = \frac{p}{f_0 \rho_0}$

$$\psi_x = v_g$$

$$-\psi_y = u_g$$

Thermal wind balance:

$$\begin{cases} f \frac{\partial v_g}{\partial z} = \frac{\partial b}{\partial x} \\ f \frac{\partial u_g}{\partial z} = -\frac{\partial b}{\partial y} \end{cases}$$

$$\Rightarrow f \frac{\partial \vec{u}}{\partial z} = k \times \nabla b$$

Rossby Number

$$Ro = \frac{U}{fL}$$

Stability / Buoyancy Frequency

$$N^2 = \frac{-g}{\rho_0} \left(\frac{\partial \rho_0}{\partial z} \right)_E \approx \frac{-g}{\rho_0} \left(\frac{\partial \rho_\theta}{\partial z} \right)_E = \left(\frac{\partial b_\theta}{\partial z} \right)_E.$$

Stable if $N^2 > 0$

Unstable if $N^2 < 0$

Shallow Water Equations

• Barotropic:

Continuity: $\frac{\partial h}{\partial t} + \nabla \cdot (h \vec{u}) = 0$

$$h = H + \eta$$

Momentum: $\frac{D\vec{u}}{Dt} + \vec{f} \times \vec{u} = -g \nabla \eta$

• Waves:

$$\omega^2 = f_0^2 + gH(k^2 + l^2)$$

• Reduced gravity: $\vec{u}_2 = 0$, $h = \eta_0 - \eta$

Continuity: $\frac{Dh}{Dt} + h \nabla \cdot \vec{u} = 0$

Momentum: $\frac{D\vec{u}}{Dt} + \vec{f} \times \vec{u} = g' \nabla \eta$

• Shallow water PV:

$$\frac{D}{Dt} \left(\frac{f + \zeta}{h} \right) = 0$$

Planetary Geostrophy

• One layer: $\vec{f} \times \vec{u} = -g \nabla \eta$
 $\frac{\partial h}{\partial t} + \nabla \cdot (h \vec{u}) = 0$

• Boussinesq

$$\vec{f} \times \vec{u} = -\nabla \phi$$

$$\nabla \cdot \vec{u} = 0$$

$$\frac{\partial \phi}{\partial z} = b$$

$$\frac{Dh}{Dt} = 0$$

Quasigeostrophy

$$\frac{Dq}{Dt_g} = 0 \quad q = \beta \zeta_g + \beta y - \frac{f_0}{H} \eta \quad \frac{D}{Dt_g} := \frac{\partial}{\partial t} + \vec{u}_g \cdot \nabla$$

• Streamfunction $\psi = \frac{gH}{f_0} q$ $q = \nabla^2 \psi + \beta y - \frac{1}{L_D^2} \psi$ $L_D = \frac{\sqrt{gH}}{f_0}$

$$\frac{\partial q}{\partial t} + J[\psi, q] = 0$$

$$J[A, B] = A_x B_y - A_y B_x$$

• Dispersion relation $\omega = \frac{-\beta k}{k^2 + l^2 + L_D^{-2}}$

- Continuously stratified QG

$$\frac{Dq}{Dt} = 0 \quad q = \zeta_g + \beta y + \frac{2}{\sigma_2} \left(\frac{f_0 b'}{N^2} \right)$$

$$q = \nabla^2 \psi + \beta y + f_0^2 \frac{\partial}{\partial z} \left(\frac{1}{N^2} \frac{\partial \psi}{\partial z} \right)$$

- Dispersion relation $\omega = \frac{-\beta k}{k^2 + l^2 + m^2 \frac{f_0^2}{N^2}}$

Viscous Stress

• $\nabla \cdot \overleftrightarrow{\tau} = -\nabla p + \nabla \cdot \overleftrightarrow{\tau}$ surface stress tensor

\uparrow normal stresses \uparrow shear stresses

- For Newtonian Fluids: $\overleftrightarrow{\tau} = \mu \nabla^2 \vec{v}$
- Acceleration due to viscosity: $\frac{1}{\rho_0} \nabla \cdot \overleftrightarrow{\tau} = \nu \nabla^2 \vec{v}$
- Scaling: $\nu \nabla^2 \vec{v} \approx \nu \frac{\partial^2 \vec{v}}{\partial z^2} =: \frac{1}{\rho_0} \frac{\partial \overleftrightarrow{\tau}}{\partial z}$ $\overleftrightarrow{\tau} = (\overleftrightarrow{\tau}_{zx}, \overleftrightarrow{\tau}_{zy})$

Ekman

• $f \times \vec{u} = -\nabla_h \phi + \frac{1}{\rho_0} \frac{\partial \overleftrightarrow{\tau}}{\partial z}$ Frictional geostrophic Balance

• Ekman flow $\vec{u}_{Ek} = \vec{u} - \vec{u}_g$

$$f \times \vec{u}_{Ek} = \frac{1}{\rho_0} \frac{\partial \overleftrightarrow{\tau}}{\partial z}$$

- Ekman transport

$$\vec{M}_{Ek} = \int_D^0 \vec{u}_{Ek} dz$$

$$\boxed{\vec{M}_{Ek} = \frac{1}{\rho_0 f} \overleftrightarrow{\tau}_s \times \hat{k}}$$

Ekman Number

$$Ek = \frac{\nu}{f_0 H^2}$$

Ekman Depth

$$D = \sqrt{\frac{\nu}{f_0}}$$

- Ekman pumping

$$\omega_{Ek} = \frac{1}{\rho_0} \nabla \times \left(\frac{\overleftrightarrow{\tau}_s}{f} \right) \cdot \hat{k} = \frac{1}{\rho_0} \text{curl} \left(\frac{\overleftrightarrow{\tau}_s}{f} \right)$$

Barotropic Vorticity Equation

$$\frac{\partial \zeta_{BT}}{\partial t} + \text{curl} \left[\int_H^{\eta} \vec{v} \cdot \nabla \vec{u} \, dz \right] + \beta V_{BT}$$

$$= -\text{curl} \left[\phi_B \nabla H \right] + \text{curl} \left[\int_H^{\eta} v \nabla^2 \vec{u} \, dz \right] + \frac{1}{g_0} \text{curl} \vec{\tau}$$

• Sverdrup balance

$$\boxed{\beta V_{BT} = \frac{1}{g_0} \nabla \times \vec{\tau} \cdot \hat{k}}$$