

Dynamic Landscape Evolution

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Abstract

This is a fine-scale, short term, process-based landscape evolution model using simulated erosion and deposition to generate a time-series of digital elevation models and compute the net change in elevation. This model uses a path sampling method to solve water and sediment flow continuity equations and model mass flows over complex topographies based on topographic, land cover, soil, and rainfall parameters. This either steady state or dynamic model can simulate landscape evolution for a range of hydrologic soil erosion regimes.

Keywords: landscape evolution, dynamic model

Contents	5	Implementation	6
1 Introduction	2	6 Case studies	7
1.1 Theory	2	6.1 Fort Bragg	7
1.2 Conceptual model	2	6.2 Patterson Branch Creek	7
		6.3 Benchmarks	7
2 Simulation of water erosion model	2	7 Tangible landscape evolution	8
2.1 Erosion regime	2	8 Discussion	8
2.2 Shallow water flow	2	8.1 Future work	8
2.3 Sediment flow	3	9 Conclusion	8
2.4 Erosion-deposition	3		
2.5 Landscape evolution	3	Appendix A Supporting information	8
3 Revised universal soil loss equation 3D model	4	Appendix A.1 Code	8
3.1 Event-based r-factor derivation	4	Appendix A.2 Data	8
3.2 Flow accumulation	4	Appendix A.3 3D models	8
3.3 3D topographic factor	4	Appendix A.4 Tangible Landscape	8
3.4 Sediment flow	4		
3.5 Landscape evolution	5		
4 Unit streampower erosion deposition model	5		
4.1 Topographic sediment transport factor	5		
4.2 Sediment flow at transport capacity	5		
4.3 Erosion-deposition at transport capacity	5		
4.4 Landscape evolution	5		

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1. Introduction

This process-based, spatially distributed, dynamic model uses a path sampling method to solve the water and sediment flow equations [8] and model mass flows over complex topographies based on topographic, land cover, soil, and rainfall parameters. The modeled flow of sediment – a function of the flow of water and soil detachment and transport parameters – is then used to estimate the net erosion and deposition rates and the associated short-term evolution of the topography.

This highly adaptable model can simulate landscape evolution for different soil erosion regimes across a range of spatiotemporal scales using either the simulated water erosion (SIMWE) model, the unit stream power erosion deposition (USPED) model, or the 3-dimensional revised universal soil loss equation (RUSLE 3D) model.

The dynamic landscape evolution has been implemented as the add-on module *r.evolution*¹ for a free, open-source geographic information system (GIS) – Geographic Resources Analysis Support System (GRASS) GIS. It supports multithreading and parallel processing for the efficient computation of physics-based simulations for large, high resolution topographic datasets.

1.1. Theory

1.2. Conceptual model

2. Simulation of water erosion model

SIMWE – the Simulation of Water Erosion model – is a physics-based simulation of shallow overland water and sediment flow that uses a path sampling method to solve the continuity and momentum equations with a 2D diffusive wave approximation [4, 7, 8]. It has been implemented in GRASS GIS as the modules *r.sim.water*² and *r.sim.sediment*³.

In SIMWE mode for each time step the dynamic landscape evolution model determines the soil erosion regime, simulates water and sediment flows, and then evolves the topography. In an erosion-deposition regime the model computes the partial derivatives of the topography, simulates shallow water flow and erosion-deposition, and then evolves the topography based on the erosion-deposition rate and gravitational diffusion.

The same process is used in a transport capacity limited regime except that the topography is evolved based on the transport limited erosion-deposition rate and gravitational diffusion. In a detachment capacity limited regime the model instead computes the partial derivatives of the topography, simulates shallow water flow and sediment flow, and then evolves the topography based on the sediment flow rate and gravitational diffusion.

The model simulates dynamic landscape evolution when the time step is less than the travel time for a drop of water or a particle of sediment to cross the landscape. With longer time steps the model simulates steady state dynamics.

2.1. Erosion regime

This model can switch erosion regimes at each time step based on the rainfall intensity i_r and the balance of the sediment detachment capacity D_c and the sediment transport capacity T_c represented by the first order reaction term σ which depends on soil and landcover properties. The detachment capacity is the maximum potential detachment rate by overland flow, while the sediment transport capacity is the maximum potential sediment flow rate.

$$\sigma = \frac{D_c}{T_c} \quad (1)$$

where:

σ is a first order reaction term (m^{-1})

D_c is the sediment detachment capacity ($kg\ m^{-1}\ s^{-1}$)

T_c is the sediment transport capacity ($kg\ m^{-1}\ s^{-1}$)

When rainfall intensity is very high ($i_r \geq 60\ mm\ hr^{-1}$) or σ is low ($\sigma \leq 0.01\ m^{-1}$), then the regime is detachment capacity limited. When rainfall intensity is not very high ($i_r < 60\ mm\ hr^{-1}$) and σ is high ($\sigma \geq 100\ m^{-1}$), then the regime is transport capacity limited. When rainfall intensity is not very high ($i_r < 60\ mm\ hr^{-1}$) and σ is neither high nor low ($0.01\ m^{-1} < \sigma < 100\ m^{-1}$), then there is an erosion-deposition regime.

2.2. Shallow water flow

The SIMWE model simulates shallow overland water flow controlled by spatially variable topographic, soil, landcover, and rainfall parameters by solving the continuity and momentum equations for steady state water flow with a path sampling method. Shallow water flow $q(x, y, t)$ can be approximated by the bivariate form of the St. Venant equation:

¹https://github.com/baharmon/landscape_evolution

²<https://grass.osgeo.org/grass75/manuals/r.sim.water.html>

³<https://grass.osgeo.org/grass75/manuals/r.sim.sediment.html>

Table 1: Select landscape evolution models

Model	Spatial scale	Temporal scale	Representation	Dynamics	GIS
SIBERIA	landscape	continuous	raster		
CHILD	landscape	continuous	mesh		
r.landscape.evol	landscape	continuous	raster		✓
r.evolution	watershed – landscape	event	raster	dynamic & steady state	✓

$$\frac{\partial h(x, y, t)}{\partial t} = i_e(x, y, t) - \nabla q(x, y, t) \quad (2)$$

where:

(x, y) is the position (m)

t is the time (s)

$h(x, y, t)$ is the depth of overland flow (m)

$i_e(x, y, t)$ is the rainfall excess ($m s^{-1}$)

(i.e. rainfall intensity – infiltration – vegetation intercept)

∇ is the divergence of the flow vector field

$q(x, y, t)$ is the water flow per unit width ($m^2 s^{-1}$).

Diffusive wave effects can be approximated so that water can flow through depressions by integrating a diffusion term $\propto \nabla^2[h^{5/3}(x, y)]$ into the solution of the continuity and momentum equations for steady state water flow. This equation is solved using a Green's function Monte Carlo path sampling method.

$$-\frac{\varepsilon(x, y)}{2} \nabla^2[h^{5/3}(x, y)] + \nabla[h(x, y) v(x, y)] = i_e(x, y) \quad (3)$$

where:

$\varepsilon(x, y)$ is a spatially variable diffusion coefficient.

2.3. Sediment flow

In SIMWE the sediment flow rate $q_s(x, y, t)$ is estimated as a function of water flow and sediment concentration:

$$q_s(x, y, t) = \rho_s(x, y, t) q(x, y, t) \quad (4)$$

where:

$q_s(x, y, t)$ is the sediment flow rate per unit width ($kg m^{-1} s^{-1}$)

$\rho_s(x, y, t)$ is sediment mass density ($kg m^{-3}$).

2.4. Erosion-deposition

In SIMWE the net erosion-deposition rate is estimated using the bivariate form of sediment continuity equation to model sediment storage and flow based

on effective sources and sinks. Net erosion-deposition $d_s(x, y, t)$ – the difference between sources and sinks – is approximated by the steady state sediment flow equation with diffusion:

$$d_s(x, y, t) = \frac{\partial[\rho_s c(x, y, t) h(x, y, t)]}{\partial t} + \nabla q_s(x, y, t) \quad (5)$$

where:

$d_s(x, y, t)$ is net erosion-deposition ($kg m^{-2} s^{-1}$).

2.5. Landscape evolution

The simulated change in elevation $\Delta z(x, y, t)$ due to water erosion and deposition is a function of the change in time, the net erosion-deposition rate, and the sediment mass density [5]:

$$\Delta z(x, y, t) = \Delta t d_s(x, y, t) \rho_s^{-1} \quad (6)$$

In a detachment limited erosion regime the simulated change in elevation $\Delta z(x, y, t)$ is a function of the change in time, the sediment flow rate, and the mass of water carried sediment per unit area [5]:

$$\Delta z(x, y, t) = \Delta t q_s(x, y, t) \varrho_s^{-1} \quad (7)$$

where:

ϱ_s is the mass of sediment per unit area ($kg m^{-2}$).

Gravitational diffusion is then applied to the evolved topography to simulate the settling of sediment particles. The simulated change in elevation $\Delta z(x, y, t)$ due to gravitational diffusion is a function of the change in time, the sediment mass density, the gravitational diffusion coefficient, and topographic divergence – i.e. the sum of the second order derivatives of elevation [11]:

$$\Delta z(x, y, t) = \Delta t \rho_s^{-1} \varepsilon_g \nabla(x, y, t) \quad (8)$$

where:

ε_g is the gravitational diffusion coefficient ($m^2 s^{-1}$)

$\nabla(x, y, t)$ is the topographic divergence (m^{-1}).

3. Revised universal soil loss equation 3D model

The Revised Universal Soil Loss Equation for Complex Terrain (RUSLE3D) is an empirical equation for computing erosion in a detachment-capacity limited soil erosion regime for watersheds with complex topography [6]. It is based on the Universal Soil Loss Equation (USLE), an empirical equation for estimating the average sheet and rill soil erosion from rainfall and runoff on agricultural fields and rangelands with simple topography [12]. It models erosion dominated regimes without deposition in which sediment transport capacity is uniformly greater than detachment capacity. As an empirical equation the predicted soil loss is spatially and temporally averaged. In USLE soil loss per unit area is determined by an erosivity factor R , a soil erodibility factor K , a slope length factor L , a slope steepness factor S , a cover management factor C , and a prevention measures factor P . These factors are empirical constants derived from an extensive collection of measurements on 22.1m standard plots with an average slope of 0.09°. RUSLE3D was designed to account for more complex, 3D topography with converging and diverging flows. In RUSLE3D the topographic potential for erosion at any given point is represented by a 3D topographic factor LS_{3D} , which is a function of the upslope contributing area and the angle of the slope.

In this spatially and temporally distributed model RUSLE3D is modified by the use of a event-based r -factor derived from the rainfall intensity at each time step. For each time step this model computes the parameters for RUSLE3D – an event-based erosivity factor, the slope of the topography, the flow accumulation, and the 3D topographic factor – and then solves the RUSLE3D equation for sediment flow. The sediment flow is used to simulate landscape evolution in a detachment capacity limited soil erosion regime.

3.1. Event-based r -factor derivation

In USLE and RUSLE the erosivity factor R is the combination of the total energy and peak intensity of a rainfall event, representing the interaction between the detachment of sediment particles and the transport capacity of the flow. It can be calculated as the product of the the kinetic energy of the rainfall event E and its maximum 30-minute intensity I_{30} [1, 10]. In this model, however, the erosivity factor is derived at each time step as a function of kinetic energy, rainfall volume, rainfall intensity, and time. First rain energy is derived from rainfall intensity [1]:

$$e_r = 0.29 (1 - 0.72 \exp(-0.05 i_r)) \quad (9)$$

where:

e_r is unit rain energy ($MJ ha^{-1} mm^{-1}$)

i_r is rainfall intensity ($mm h^{-1}$).

Then the event-based erosivity index R_e is calculated as:

$$R_e = e_r v_r i_r t_r \quad (10)$$

R_e is the event-based erosivity index ($MJ mm ha^{-1} hr^{-1}$)

v_r is rainfall volume (mm) derived from $v_r = i_r t_r$

t_r is time interval s

3.2. Flow accumulation

The upslope contributing area is determined by flow accumulation. Flow accumulation is calculated using a multiple flow direction algorithm [3] based on A^T least cost paths [2]. The multiple flow direction algorithm implemented in GRASS GIS as the module *r.watershed*⁴ is computationally efficient and can navigate nested depressions and other obstacles.

3.3. 3D topographic factor

The 3D topographic factor $LS_{3D}(x, y)$ is calculated as a function of the flow accumulation, representing the upslope contributing area, and the slope. The empirical coefficients m and n for the upslope contributing area and the slope can range from 0.2 to 0.6 and 1.0 to 1.3 respectively with low values representing dominant sheet flow and high values representing dominant rill flow.

$$LS_{3D}(x, y) = (m + 1.0) (a(x, y) a_0^{-1})^m (\sin(\beta)/\beta_0)^n \quad (11)$$

where:

LS_{3D} is the dimensionless topographic (length-slope) factor

a is flow accumulation (m)

a_0 is the length of the standard USLE plot (22.1m)

β is the slope angle (°)

m is an empirical coefficient

n is an empirical coefficient

β_0 is the slope of the standard USLE plot (0.09°)

3.4. Sediment flow

The sediment flow is a function of the event-based erosivity factor, the soil erodibility factor, the 3D topographic factor, cover factor, and the prevention measures factor:

$$E = R_e K LS_{3D} C P \quad (12)$$

⁴<https://grass.osgeo.org/grass72/manuals/r.watershed.html>

where:

E is soil loss ($kg\ m^{-2}\ min^{-1}$)

R_e is the event-based erosivity factor ($MJ\ mm\ ha^{-1}\ hr^{-1}$)

K is the soil erodibility factor ($ton\ ha\ hr\ ha^{-1}\ MJ^{-1}\ mm^{-1}$)

LS_{3D} is the dimensionless topographic (length-slope) factor

C is the dimensionless land cover factor

P is the dimensionless prevention measures factor

3.5. Landscape evolution

With RUSLE3D the simulated change in elevation $\Delta z(x, y, t)$ is derived from equation 7 for landscape evolution in an detachment limited soil erosion regime and then equation 8 for the settling of sediment particles due to gravitational diffusion.

4. Unit streampower erosion deposition model

The Unit Stream Power Erosion Deposition (USPED) model estimates net erosion-deposition as the divergence of sediment flow in transport capacity limited soil erosion regimes. At transport capacity shallow flows of water are carrying as much sediment possible – more sediment is being detached than can be transported. As a transport capacity limited model USPED predicts erosion where transport capacity increases and deposition where transport capacity decreases. In USPED the influence of topography on erosion and deposition is represented by a topographic sediment transport factor, while the influence of soil and landcover are represented by factors adopted from USLE and RUSLE [6].

With USPED net erosion-deposition is estimated by computing the event-based erosivity factor R_e using Eq. 10, the slope and aspect of the topography, the flow accumulation with a multiple flow direction algorithm, the topographic sediment transport factor, the sediment flow at transport capacity, and the divergence of the sediment flow.

4.1. Topographic sediment transport factor

For USPED the 3D topographic factor (Eq. 11 for RUSLE3D) is adapted to represent the topographic sediment transport factor LST – the topographic component of overland flow at sediment transport capacity:

$$LST = U^m (\sin \beta)^n \quad (13)$$

where:

LST is the topographic sediment transport factor

U is the flow accumulation (m)

β is the angle of the slope ($^\circ$)

m is an empirical coefficient

n is an empirical coefficient.

4.2. Sediment flow at transport capacity

The sediment flow at transport capacity is a function of the event-based rainfall factor, the soil erodibility factor, the topographic component of overland flow, the landcover factor, and the prevention measures factor:

$$T = R_e K C P LST \quad (14)$$

where:

T is sediment flow at transport capacity ($kg\ m^{-1}\ s^{-1}$)

R_e is the event-based rainfall factor ($MJ\ mm\ ha^{-1}\ hr^{-1}$)

K is the soil erodibility factor ($ton\ ha\ hr\ ha^{-1}\ MJ^{-1}\ mm^{-1}$)

C is the dimensionless land cover factor

P is the dimensionless prevention measures factor.

4.3. Erosion-deposition at transport capacity

Net erosion-deposition is estimated as the divergence of sediment flow:

$$d_s(x, y, t) = \frac{\partial(T \cos \alpha)}{\partial x} + \frac{\partial(T \sin \alpha)}{\partial y} \quad (15)$$

where:

$d_s(x, y, t)$ is net erosion-deposition ($kg\ m^{-2}\ s^{-1}$)

α is the aspect of the topography ($^\circ$).

4.4. Landscape evolution

With USPED the simulated change in elevation $\Delta z(x, y, t)$ is derived from equation 6 for landscape evolution and then equation 8 for the settling of sediment particles due to gravitational diffusion.

5. Implementation

The GRASS GIS add-on module written in Python is available on Github at https://github.com/baharmon/landscape_evolution released under the GNU General Public License version 2. GRASS GIS is an open source project released under the GNU General Public License version 2. GRASS GIS is available at <https://grass.osgeo.org/>.

6. Case studies

6.1. Fort Bragg

6.2. Patterson Branch Creek

6.3. Benchmarks

7. Tangible landscape evolution

Tangible Landscape – a tangible user interface tightly integrated with a geographic information system for intuitively sketching in 3D [9]. Conceptually, Tangible Landscape couples a physical model with a digital model in a real-time feedback cycle of 3D scanning, geospatial modeling and simulation, and projection in order to physically manifest digital data as tangible bits. With tangible bits users can directly, physically feel and manipulate data with their bodies – naturally, intuitively understanding space, form, and process. Tangible Landscape is available on Github at <https://github.com/ncsu-osgeorel/grass-tangible-landscape>.

We coupled Tangible Landscape with the landscape evolution model to test the model and experiment with strategies for restoration. We used Tangible Landscape to computationally steer the landscape evolution model and interactively explore the relationship between overland flow patterns and changes in topography. By manually changing the physical model of the landscape we change the topography used by the model.

8. Discussion

8.1. Future work

1. Test the model on historical data
2. Test the model with UAS SfM time-series
3. Implement as a Tangible Landscape analysis
4. Live, in-situ fabrication in polymer-enriched sand with a robotic arm

9. Conclusion

Appendix A. Supporting information

Appendix A.1. Code

Github repository

Appendix A.2. Data

GRASS GIS Mapset

Appendix A.3. 3D models

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Appendix A.4. Tangible Landscape

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Figure 1: **Rapid prototyping.** 3-axis CNC fabrication of the evolved landscape in polymer-enriched sand using a plunge cut.